



# Polarimetric Radar Verification of GPM Satellite-Based Retrievals of the Raindrop Size Distribution



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## Outline

- **Motivation and Requirements**
- **Approach, Methods, Data**
- **Verification of basic mission requirement and GPM DSD "drill down"**
- **Summary**

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# GPM “Core” Satellite Science Requirements

## (Termed “Level -1” or “L1”)

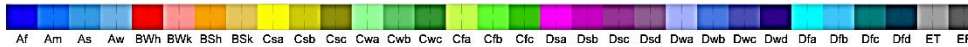
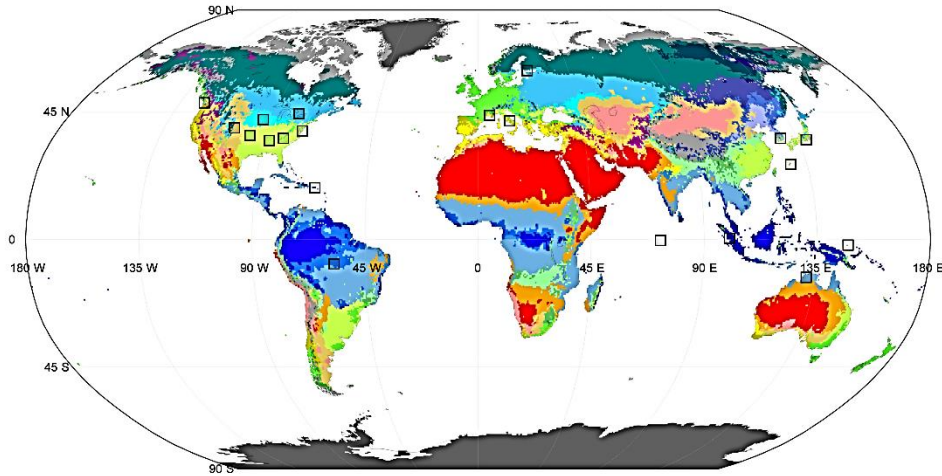
- DPR: *quantify rain rates between 0.22 and 110 mm hr<sup>-1</sup> and demonstrate the detection of snowfall at an effective resolution of 5 km.*
- GMI: *quantify rain rates between 0.22 and 60 mm hr<sup>-1</sup> and demonstrate the detection of snowfall at an effective resolution of 15 km.*
- **Core observatory radar estimation of the Drop Size Distribution (DSD)- specifically,  $D_m$  to within +/- 0.5 mm. [note- no  $N_w$  requirement]**
- Core observatory *instantaneous* rain rate estimates at a resolution of 50 km with *bias and random error < 50% at 1 mm hr<sup>-1</sup> and < 25% at 10 mm hr<sup>-1</sup>, relative to GV*



# Overarching Philosophy



World map of Köppen climate classification for 1901–2010

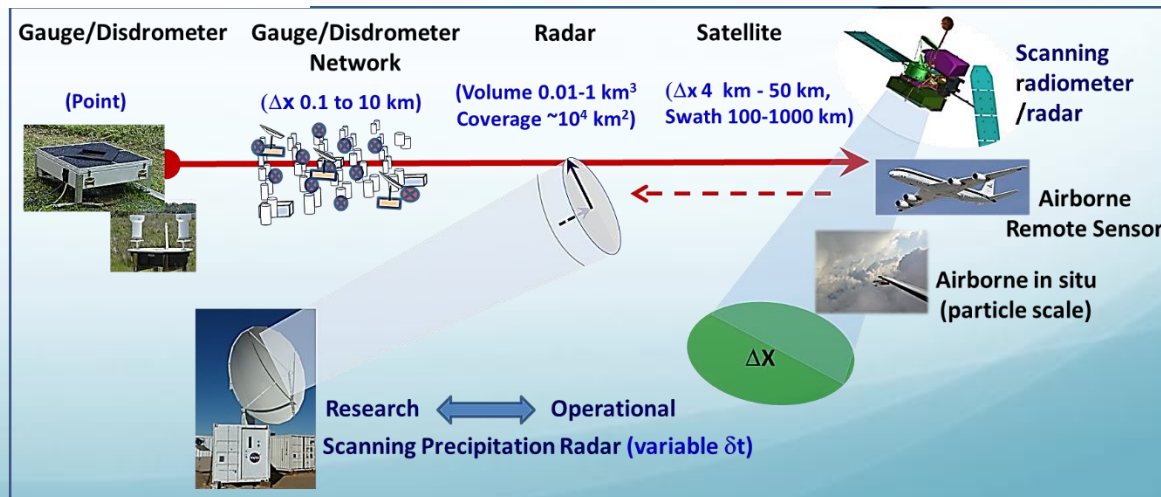


First letter	Second letter	Third letter
A: Tropical	f: Fully humid	T: Tundra
B: Dry	m: Monsoon	F: Frost
C: Mild temperate	s: Dry summer	
D: Snow	w: Dry winter	
E: Polar	W: Desert	
	S: Steppe	

Data source: Terrestrial Air Temperature/Precipitation: 1900-2010 Gridded Monthly Time Series (V 3.01)  
 Resolution: 0.5 degree latitude/longitude  
 Website: <http://hanschen.org/koppen>  
 Ref: Chen, D. and H. W. Chen, 2013: Using the Köppen classification to quantify climate variation and change: An example for 1901–2010. Environmental Development, 6, 69-79, 10.1016/j.envdev.2013.03.007.

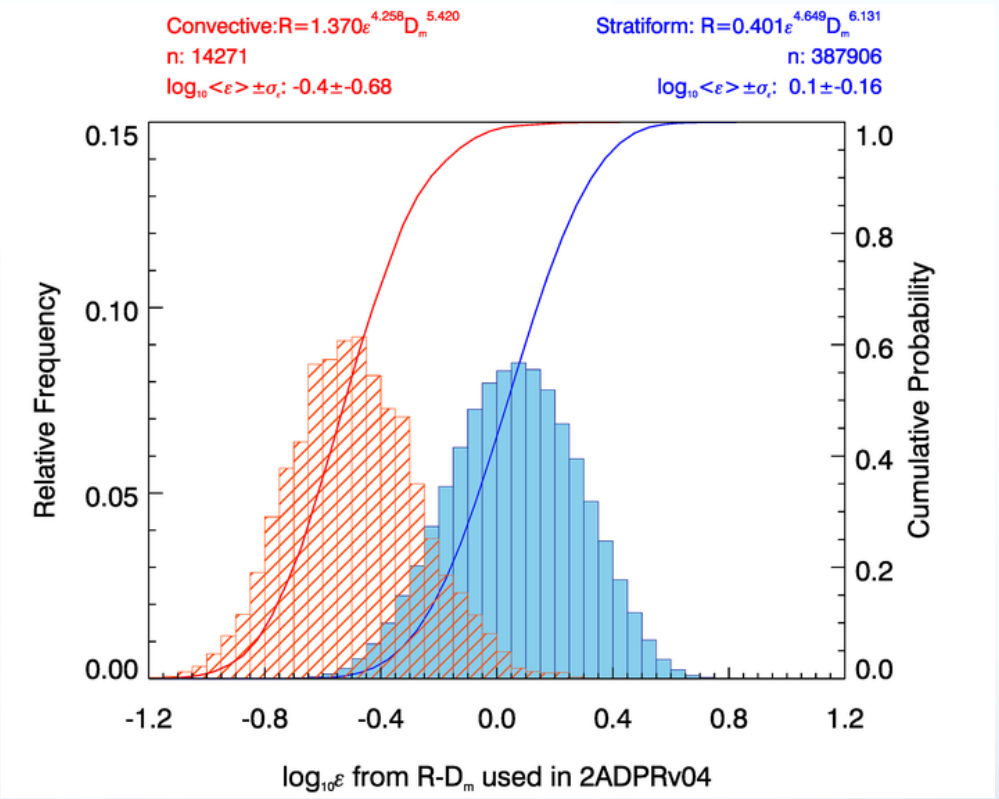
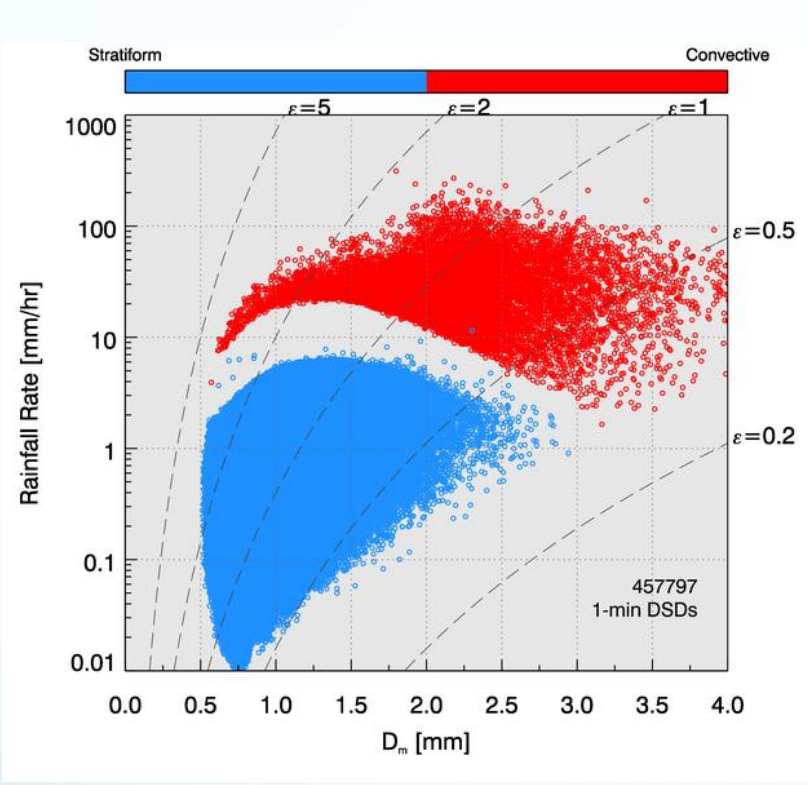
2D Video disdrometer data collected at numerous locations, regimes, and point scales.....

.....references dual-pol radar that functions as a "translator" to GPM footprint and swath scales





# Ensemble Point Data Useful for Verification of DPR DSD-related Algorithm Assumptions



Algorithm assumes  $R = C \epsilon^a D_m^b$   
 coefficients  $a$   $f(\text{rain type})$  and  $\epsilon$   
 range  $[5, 0.2]$ ;

DPR Algorithm- assumes  
 $\log$  normal  $\epsilon$ , with  $\langle \epsilon \rangle \sim 1$ ;

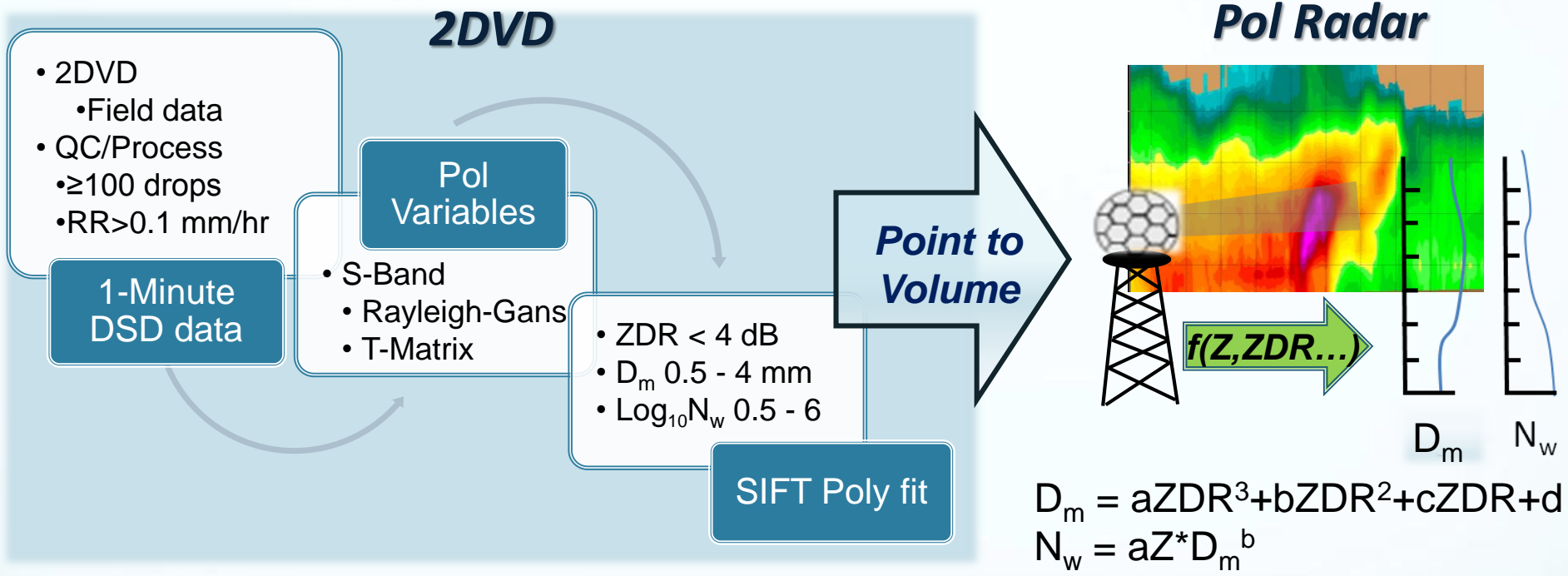
Disdrometer data suggests  $\epsilon$  is smaller  
 for convective vs. stratiform - consistent  
 (with analysis of DPR retrievals);



# Need Footprint Comparisons for L1 Requirements



## 2DVD to Radar: Methodology

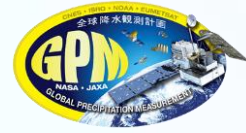


- Empirical models developed for NASA field campaign "regimes" (Oklahoma, Iowa, Alabama, Mid-Atlantic Coastal, Washington Coast, Appalachians/Piedmont....)
- Aggregated to make "**ALL-regimes**" relationship developed for U.S. continental-scale statistical verification (> 200,000 minutes used)

**"ALL" DSD model-fit relative errors: BIAS < 10%, MAE < 15%**



# Individual Field Campaign and Aggregate Retrievals

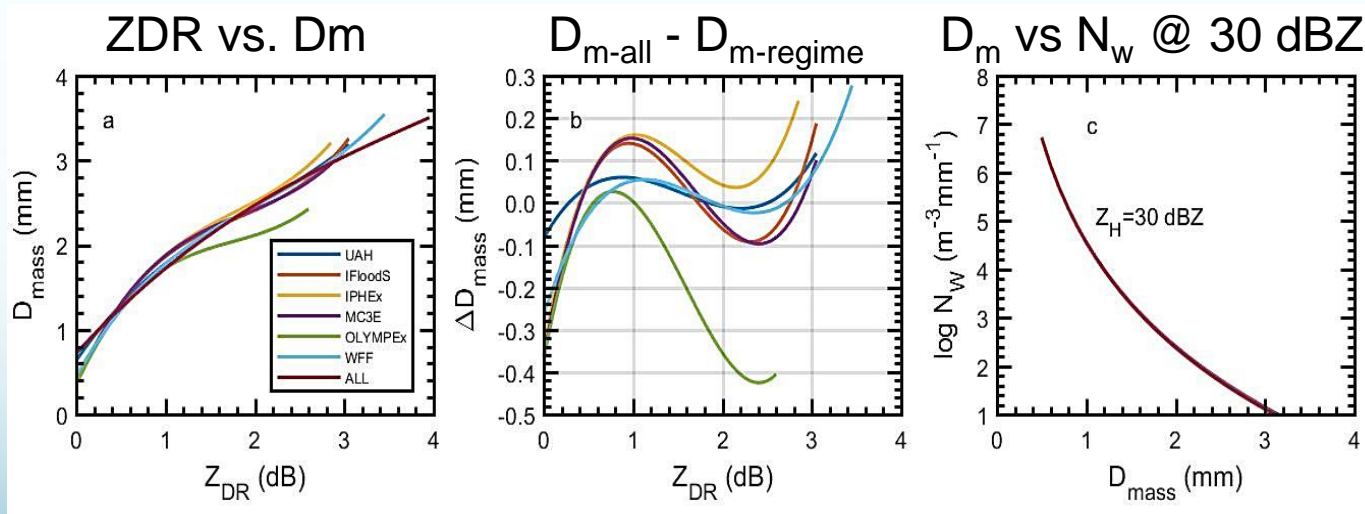


## Regime Sub-sample comparisons to NPOL

Field Campaign	Bias	Absolute Bias	Samples
<b><math>D_m</math></b>			
IFloods	0.00	0.42	6,610
IPHEX	0.07	0.34	1,058
OLYMPEX	0.03	0.34	1,008
<b>LOG10[<math>N_w</math>]</b>			
IFloods	0.04	0.90	6,610
IPHEX	-0.12	0.89	1,058
OLYMPEX	0.21	0.89	1,008

- **Sanity check:** Examine regime  $D_m$ ,  $N_w$  fits against NPOL observations;
- Examine departure of regime fits from the "ALL" relationship

Tokay et al. 2017 (in preparation)



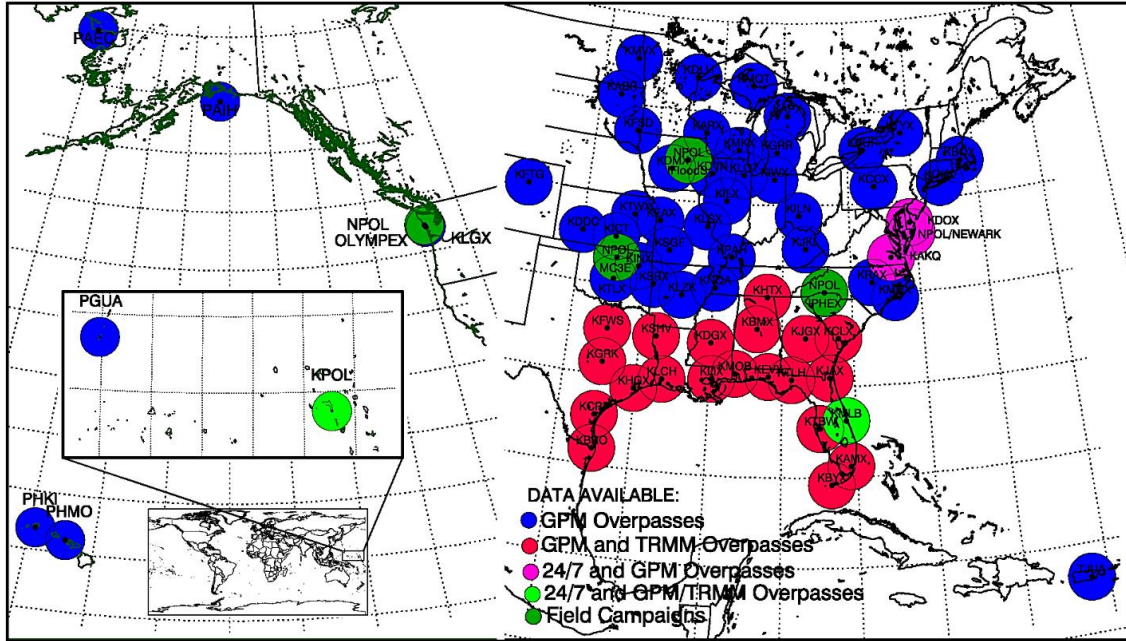
Application of the "ALL" relationship to certain regimes (e.g., OLYMPEX) and/or the less-frequently sampled large ZDR introduces more uncertainty in  $D_m$ ;  $N_w$  more stable.



# Radar to GPM: Validation Network (VN) Radar Processing



GPM-GV Radar Sites



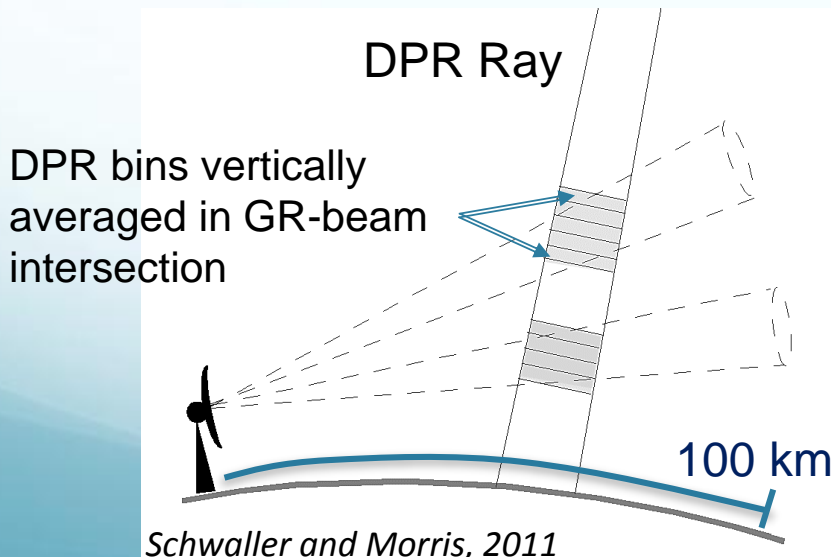
## 88Ds, NPOL, KWAJ

Network radar datasets used for "statistical" science requirements verification of the DSD

## VN Matching

For each GV radar beam, range gates within 100 km of a given radar are geometrically volume-matched to intersecting DPR rays

Products (e.g., select DPR variables, Polarimetric moments, **DSD**, HID, RR...) are stored in the VN-database.

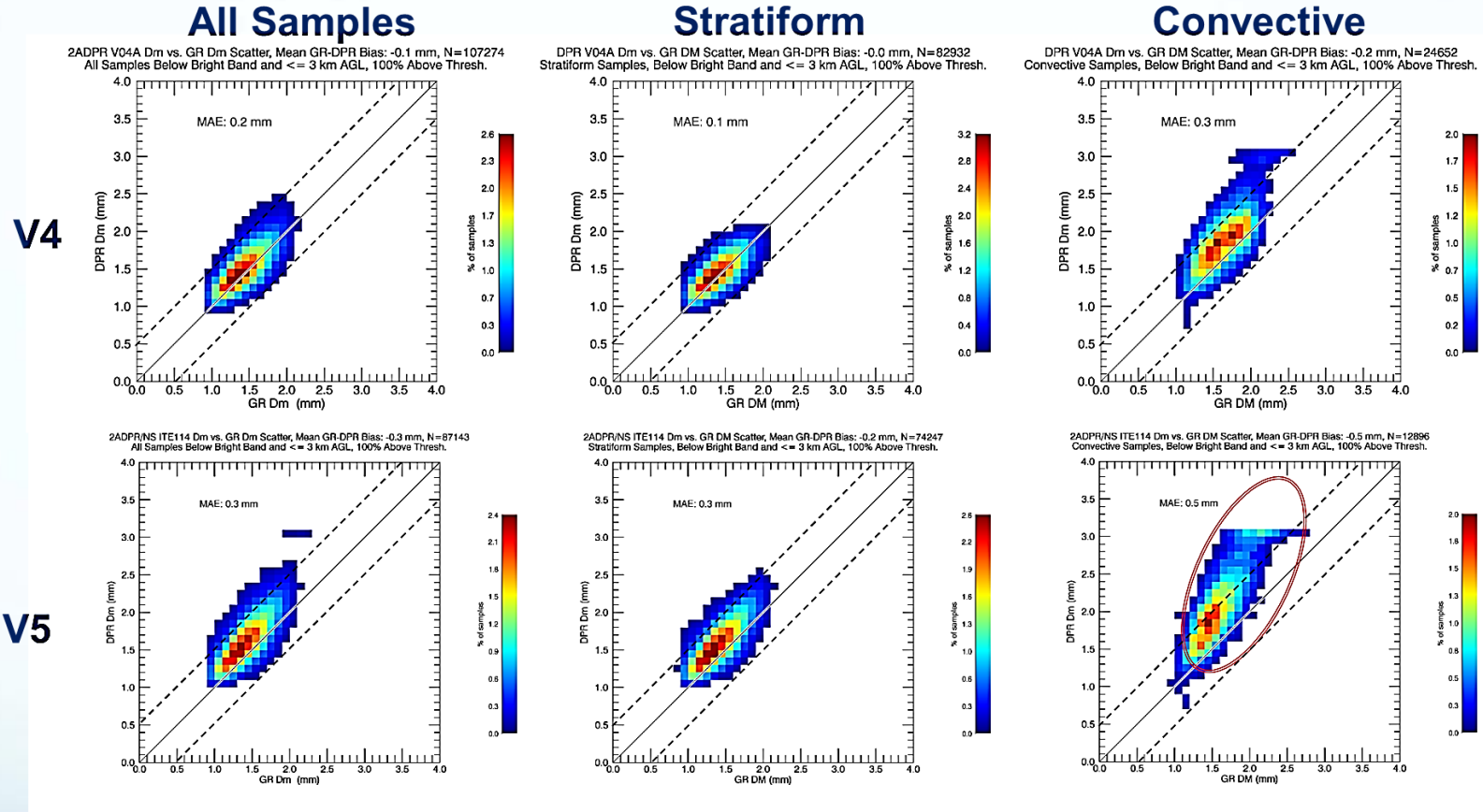




# L1 Requirement: Continental Scale VN-GPM Comparisons



## L1 DSD: DPR MS Version 4, Version 5 vs. GV Radar $D_m$



L1 science requirement: Satisfied as a whole. However, stratiform samples dominate and V5 inner swath of NS (MS) possesses an increasingly positive bias in  $D_m$  relative to GV;

2ADPR Convective  $D_m$  in V5 deviates more from GV and secondary mode in convective  $D_m$  more pronounced at large values of DPR  $D_m$  (?)





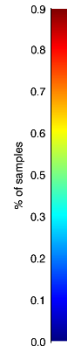
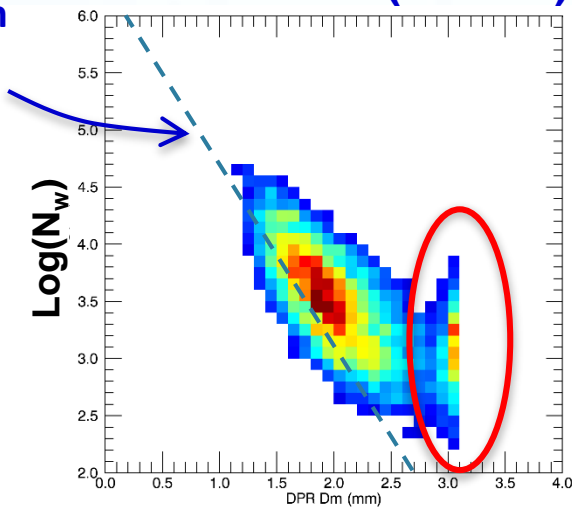
# 2ADPR Convective $N_w$ vs. $D_m$ against GV Radar



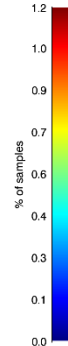
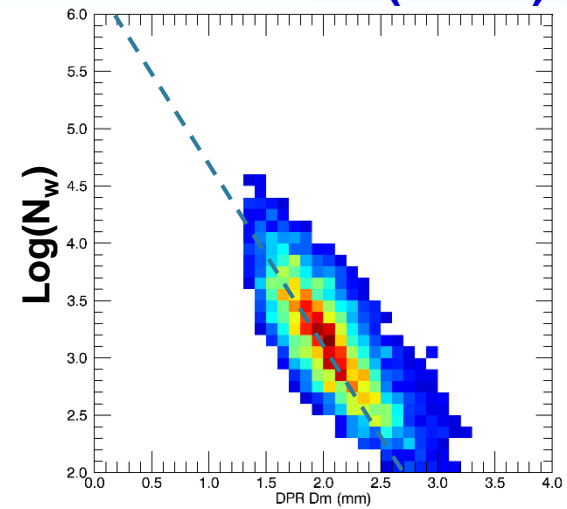
Ref. C/S Separation line (e.g., Bringi et al., 2009; Thurai et al. 2015)

DPR

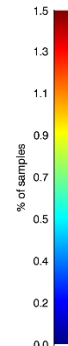
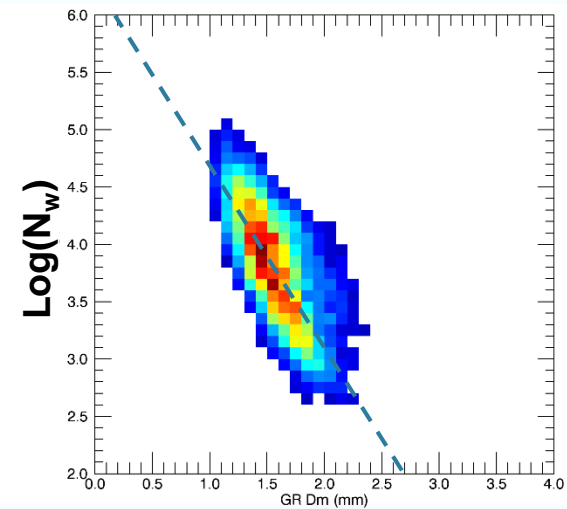
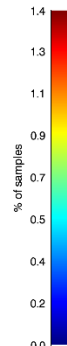
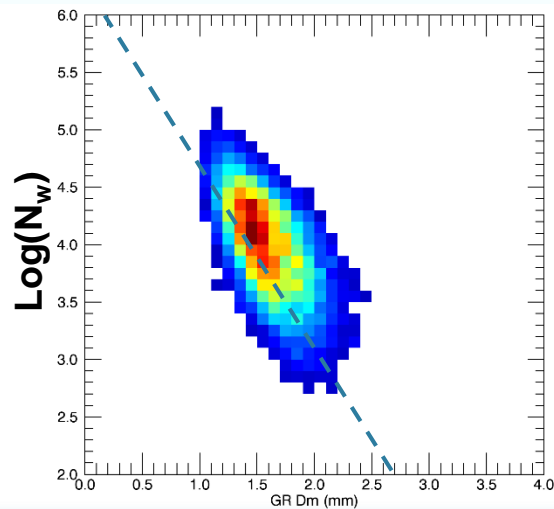
Inner Swath (Ka+Ku)



Outer Swath (KuPR)

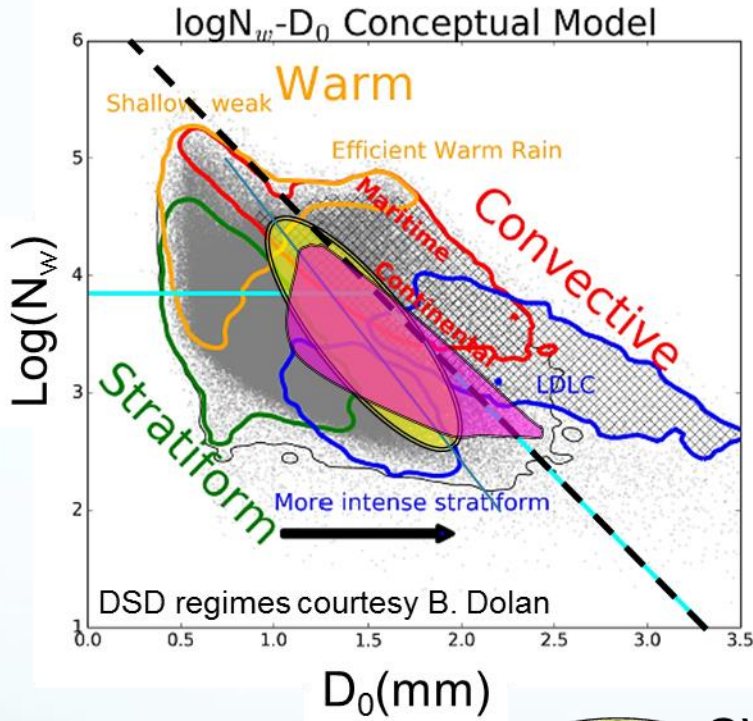


GV

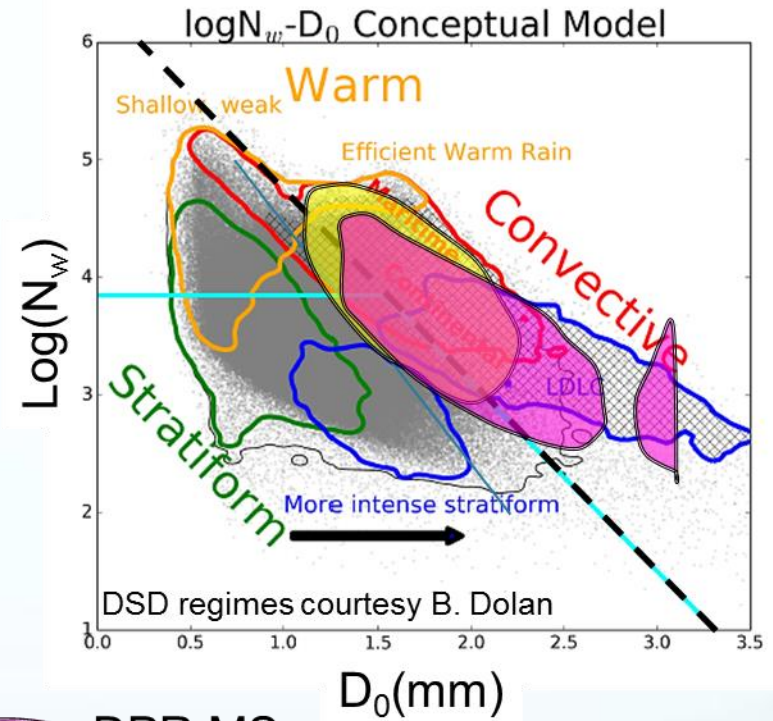


- $D_m$  offset results in lower  $N_w$  in DPR retrievals and mode in *inner* dual-freq. swath
- Differences marked between *inner* (DPR) and *outer* (2AKu) retrieval swaths
- Slope of  $N_w$  vs.  $D_m$  is reasonably similar between retrievals and GV pol relationship

## V5 Stratiform



## V5 Convective

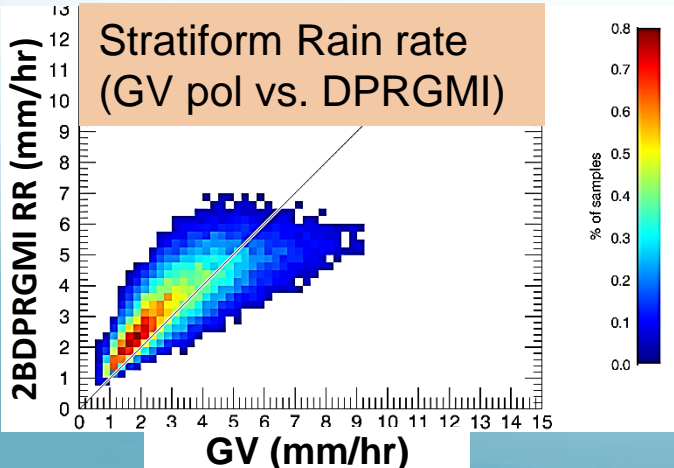
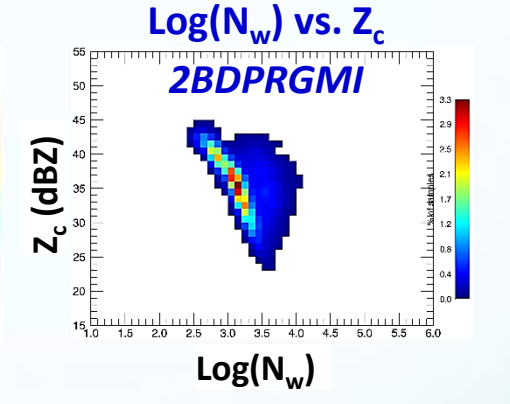
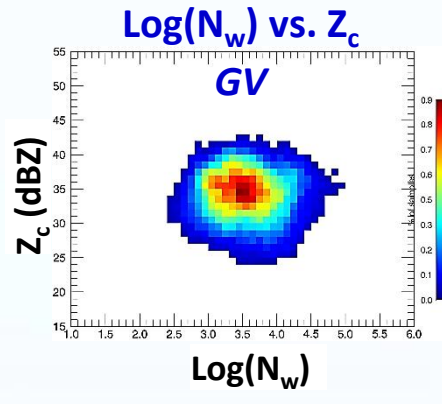
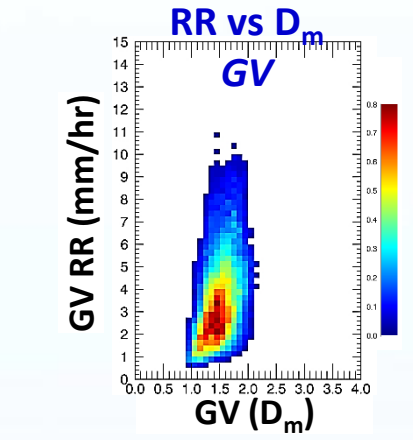
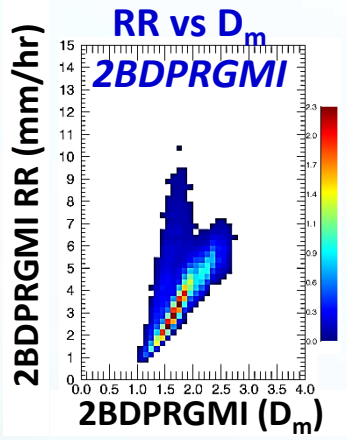
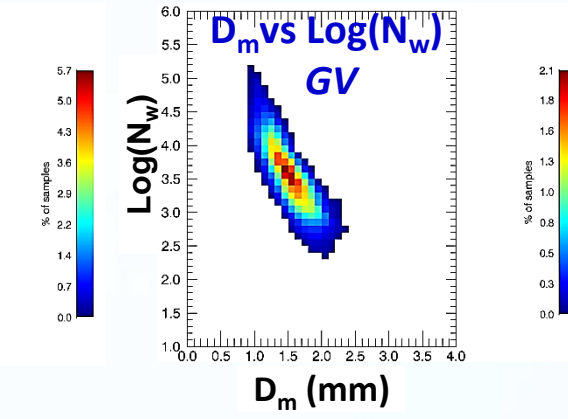
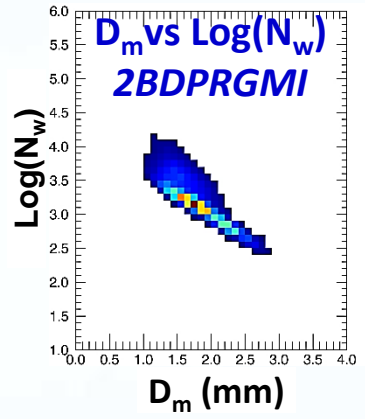
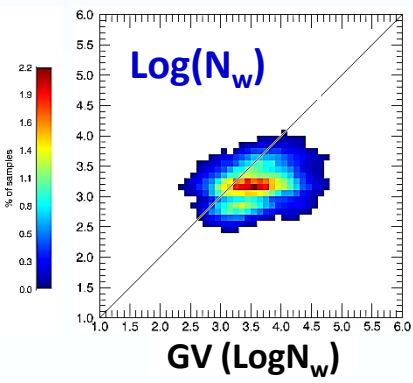
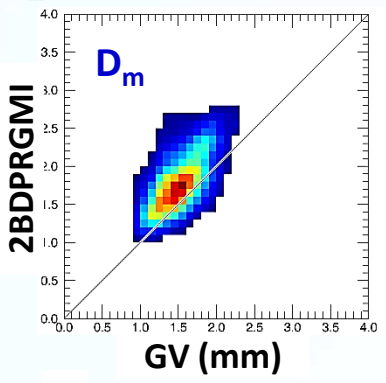


GV    DPR MS

- V5 fits GV sample space (Assuming  $D_m \approx D_0$ ); behavior qualitatively similar to GPM GV Radar
- Shift to larger  $D_m$  and smaller  $N_w$  relative to GV; secondary mode at large  $D_m$



# 2BDPRGMI: MS Swath with GV (DSD, Rain, Z...)



- Modes in the 2BDPRGMI DSD ( $N_w$ ?).
- In aggregate 2BDPRGMI produces a footprint rain rate similar to GV (GV-pol, and the MRMS!)

# Summary

## Approach:

- Polarimetric radar-based DSD retrievals ( $D_m$ ,  $N_w$ ) developed using 2DVD data for multiple rainfall regimes; scale translation to GPM satellite footprints/swaths.
- VN architecture for comparing GPM Core satellite DPR to GV on CONUS scale

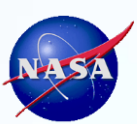
## Result:

- GPM Level 1 Requirements on  $D_m$  (+/- 0.5 mm of GV) are satisfied relative to GV measurement;
- $D_m$  positive bias- accentuated in convective precip;  $N_w$  in DPR somewhat similar to GV but responds to  $D_m$  bias; Combined-Algorithm  $N_w$ - odder behavior compared to GV.
- Sensitivity of comparisons to rain type (Convective vs. Stratiform) and swath (e.g., inner Ka/Ku vs. outer KuPR, Combined MS)- algorithms/sampling vs. physics?

## Moving ahead:

- Further analysis work to parse/isolate DSD behavior as a function of 3-D GPM and ancillary observables to guide algorithm approaches (e.g.,  $\mu$ , PIA,  $N_w$  selection in Combined algorithm,  $\varepsilon$  and associated parameter behavior in DPR/KuPR algorithms....)
- Further GV work on defining the DSD for light rain/small  $D_m$ - Generalized Gamma approach?

**EXTRAS**

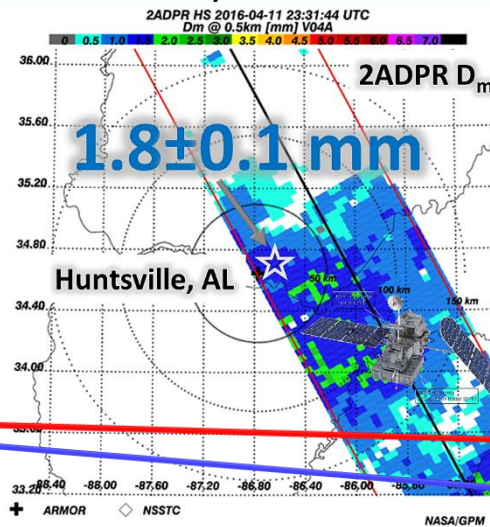


# Next: How do we handle light rain DSD?



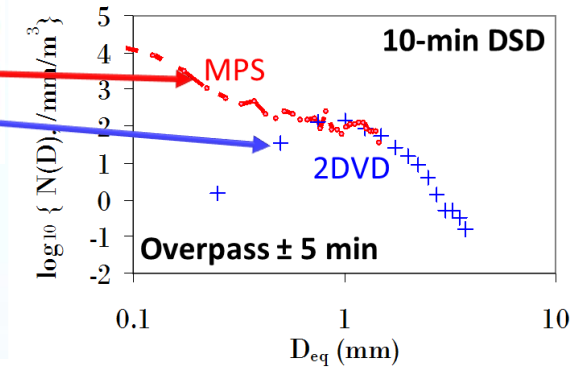
Do current DSD assumptions for GPM adequately represent the small rain drop sizes?

GPM dual-frequency precipitation radar (DPR) swath as it samples rain over Huntsville disdrometers



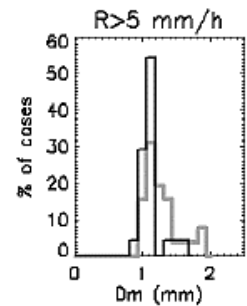
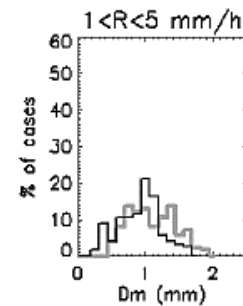
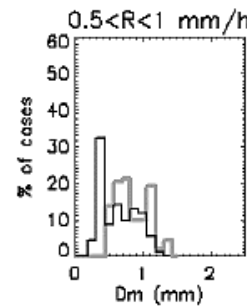
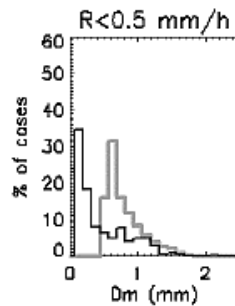
Right answer ...right reason?

DSD measured by GV  
MPS+2DVD (2DVD)  $D_m = 1.61$  (1.73) mm



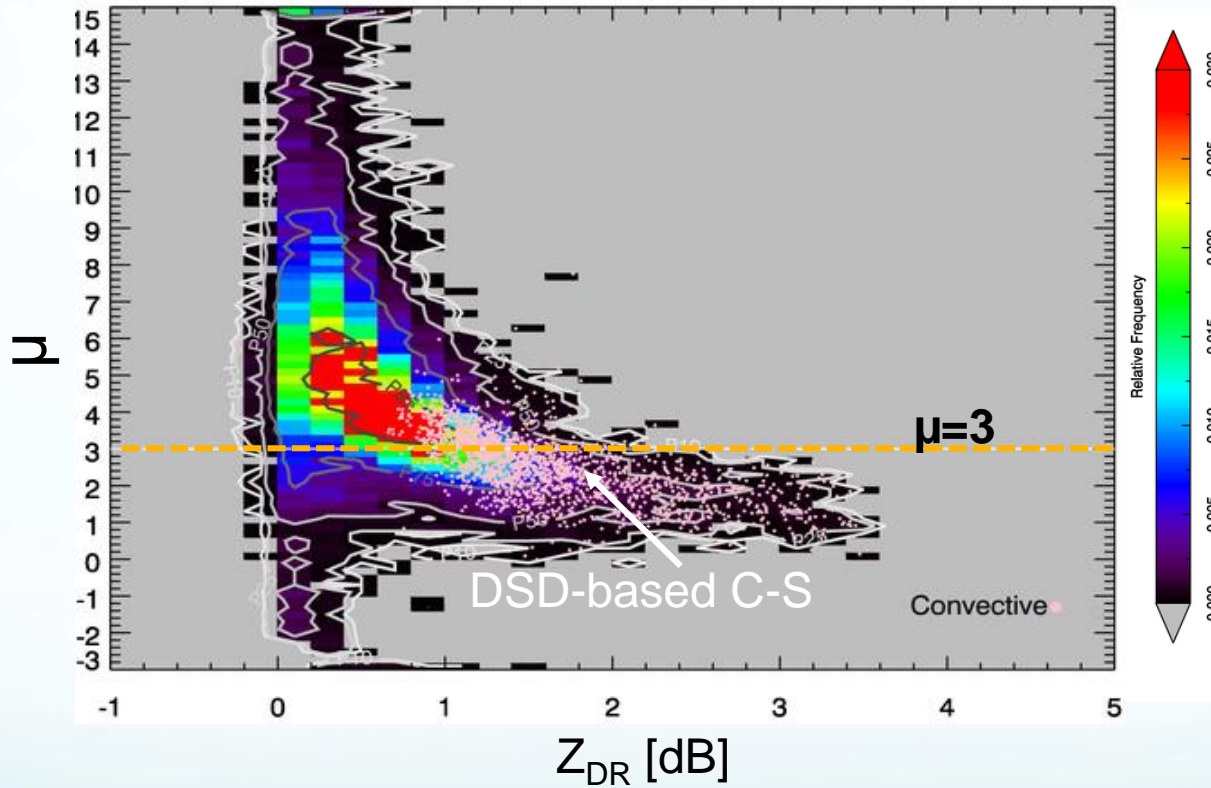
We do not properly represent the small-drop end (< 0.7 mm) of the drop size distribution-

Likely important for light rain estimation.



Reference: Thurai et al. 2017, JAMC

# $\mu = 3 ?$



Convective (stratiform)  $\mu$  almost always  $< (>) 3$  in MC3E and Alabama 2DVD data [DPR  $\mu=3$ , Combined  $\mu=2$ ]