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UNIVERSITY OF ILLINOIS  
URBANA

# AERONOMY REPORT NO. 38

## D-REGION ION CHEMISTRY

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by

S. M. Radicella

D. W. Stowe

July 15, 1970

Supported by

National Aeronautics and Space Administration

Grant NGR-013 *NGR-14-005-013*

Aeronomy Laboratory

Department of Electrical Engineering

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Urbana, Illinois

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A E R O N O M Y R E P O R T

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## ABSTRACT

The neutral composition of the mesosphere is reviewed with emphasis on nitrogen oxides and oxygen allotropes. Then ionization sources in the D region are studied including the effects of recent experimental results. Finally, the negative ion chemistry is discussed including the calculation of an "effective attachment factor,  $\beta_{\text{eff}}$ ." On the basis of the experimental  $\beta_{\text{eff}}$ , a new fast binary attachment process is considered.

In the appendix, relevant experimental data is tabulated, and the diurnal variation of atomic oxygen and ozone concentrations is presented, based on recent values for the relevant rate coefficients.



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## 1. D-REGION NEGATIVE-ION CHEMISTRY

### 1.1 Introduction

Electron-density distribution is, at this time, the best known parameter in the D region. From recent work, a general picture of this parameter can be deduced (Mechtly and Smith 1968a,b; Mechtly and Shirke 1968). Unlike the upper ionized regions, D-region structure is characterized by the attachment of electrons to neutral molecules to form, through complex chemistry, a large number of negative ions of variable distribution and mass. Electron density is therefore strongly controlled by the negative-ion chemistry. This, in turn, is determined by the rates of reactions that produce or eliminate different species of negative ions. Reaction rates depend upon the rate coefficients now quite well established by laboratory measurements (Fehsenfeld et al, 1967; Fehsenfeld and Ferguson, 1968) and upon the concentration of neutral species such as O, O<sub>3</sub> and NO. The experimental knowledge about the odd allotropes of oxygen is very limited, and theoretical computations depend upon several factors still doubtful. Recent direct estimates of NO concentration (Barth 1966; Pearce 1968) show much larger values than those obtained with indirect ionospheric methods based upon electron-density data and ion-reaction schemes widely accepted at this time. To accept the experimental data for the NO concentration will force a re-evaluation of the negative-ion chemistry, in a search for missing reactions, in order to reconcile those data and direct measurements of electron density.

The first chapter of this report will review the neutral composition of the mesosphere, where D-region ionized species are very minor constituents. Nitrogen oxides and oxygen allotropes distributions will be discussed in more detail.

The second chapter will be devoted to a review of the sources of ionization in the undisturbed D region for both daytime and nighttime. The impact of high values of NO concentrations (Pearce, 1968) in relation to the ionization of nitric oxide by Lyman-alpha radiation will be described. The importance of precipitation of electrons, as a source of ionization during nighttime will be considered.

In the third chapter, the negative-ion chemistry will be discussed. An "effective attachment factor"  $\beta_{\text{eff}}$  is defined, this being the only term in the continuity equation for electrons and ions that depends upon the negative-ion schemes. Assuming steady state conditions,  $\beta_{\text{eff}}$  is computed from experimental electron-density profiles, calculated rate of ionization and laboratory deionization data. The problem introduced by different values of NO concentration will become evident during the discussion. It will be shown that in order to obtain theoretically the experimental  $\beta_{\text{eff}}$  a new fast binary attachment process should be postulated when NO concentration measurements are used.

## 2. NEUTRAL COMPOSITION OF THE MESOSPHERE

### 2.1 Introduction

The atmospheric region between roughly 50 and 90 km, called the mesosphere, is chemically very active. Photodissociation produces atomic species that, in a relatively dense medium, can react rapidly with other constituents forming a large number of more or less stable minor species. The absorption of solar ultraviolet radiation by ozone at wavelengths of 2100-3000 Å is a strong source of heating. The effect of this is the large temperature observed around 50 km. A decrease in temperature is observed above that height up to 80-90 km. This altitude gradient of temperature is important to the chemistry of the mesosphere because many reaction rate coefficients are strongly temperature dependent.

In the region of 50 to 90 km there are a number of three-body reactions among both neutral and ionic species. The passive third body is represented in most cases by the general atmospheric molecule, i.e., the concentration of the third body involved is the atmospheric number density itself.

Both the COSPAR International Reference Atmosphere (1965) and the U.S. Standard Atmosphere Supplement (1966) give experimental models of the neutral gas densities, temperature, and other parameters as a function of height, season and latitude. Below approximately 80 km a constant mean molecular mass or, what is equivalent, a constant mixing ratio can be assumed for the major constituents. The neutral atmospheric region between 80 and 120 km is the least known experimentally, and constituent distribution is still uncertain. Variations of O<sub>2</sub> distribution are very important because dissociation of molecular oxygen controls the chemistry of the odd allotropes of oxygen and

also some aspects of the ion chemistry. Figure 1 shows the concentration of  $O_2$  from winter and summer at  $45^\circ$  latitude and low solar activity from the U.S. Standard Atmosphere Supplement (1966).

As a product of the chemical activity of the mesosphere, numerous neutral species are present in that region. Most of them have concentrations smaller by several orders of magnitude of  $O_2$  or  $N_2$ . Such species are called minor constituents. Photodissociation of  $O_2$ ,  $N_2$ ,  $H_2O$  and  $CH_4$  is the first step in the formation of:  $O$ ,  $O_3$ ,  $NO$ ,  $NO_2$ ,  $N_2O$ ,  $H$ ,  $OH$ ,  $HO_2$ ,  $H_2O$ , etc. Other minor constituents in the mesosphere could be clusters of neutral molecules like  $(H_2O) \cdot NO$  could be produced in the region of low temperature around 80 km. Also chemically important is  $CO_2$ .

At this time, all the species mentioned above, the odd allotropes of oxygen and the oxides of nitrogen are considered the most important ones to the negative-ion chemistry. Hydrogenous compounds appear to be of indirect importance because they represent only perturbations introduced in the chemistry of oxygen. The role played by neutral clusters is unknown.

## 2.2 Odd Allotropes of Oxygen

The experimental knowledge about  $O$  and  $O_3$  distribution at D-region altitudes is very limited. Only one nighttime preliminary profile of mesospheric oxygen concentration has been published (Fedynskii et al, 1967). Daytime ozone measurements have been carried out by Johnson et al, (1954) and by Rawcliffe et al, (1963). Nighttime profiles have been obtained by Carver et al. (1966), Mikirov (1965) and Reed (1968). Figure 2 summarizes these experimental results.

Theoretical computations of  $O$  and  $O_3$  distribution are based upon the knowledge about: fluxes of the solar radiation capable of photodissociating

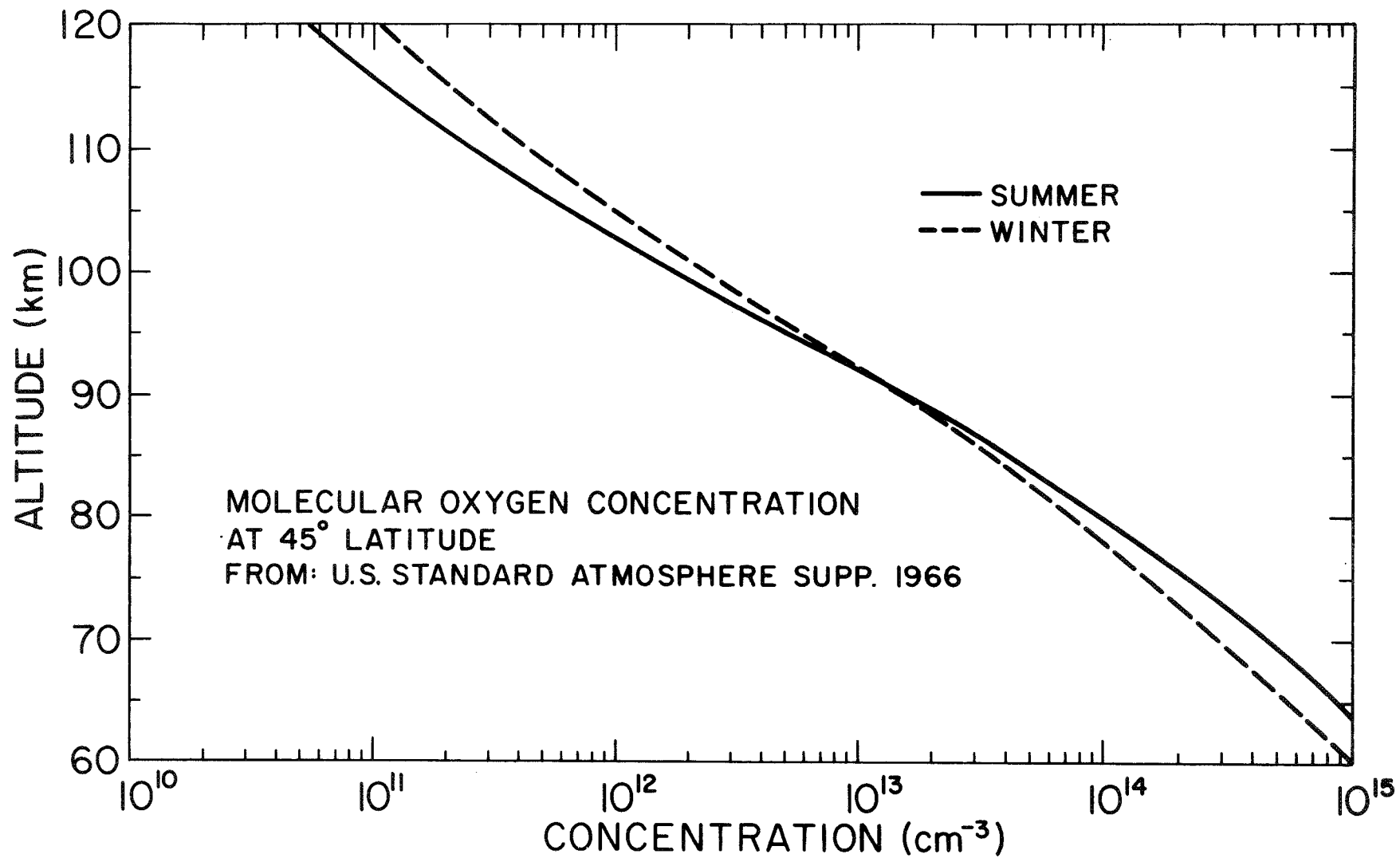


Figure 1. Molecular oxygen concentration at 45° latitude derived from U.S. Standard Atmosphere Supplement 1966. Full line indicates summer conditions; broken line indicates winter conditions.

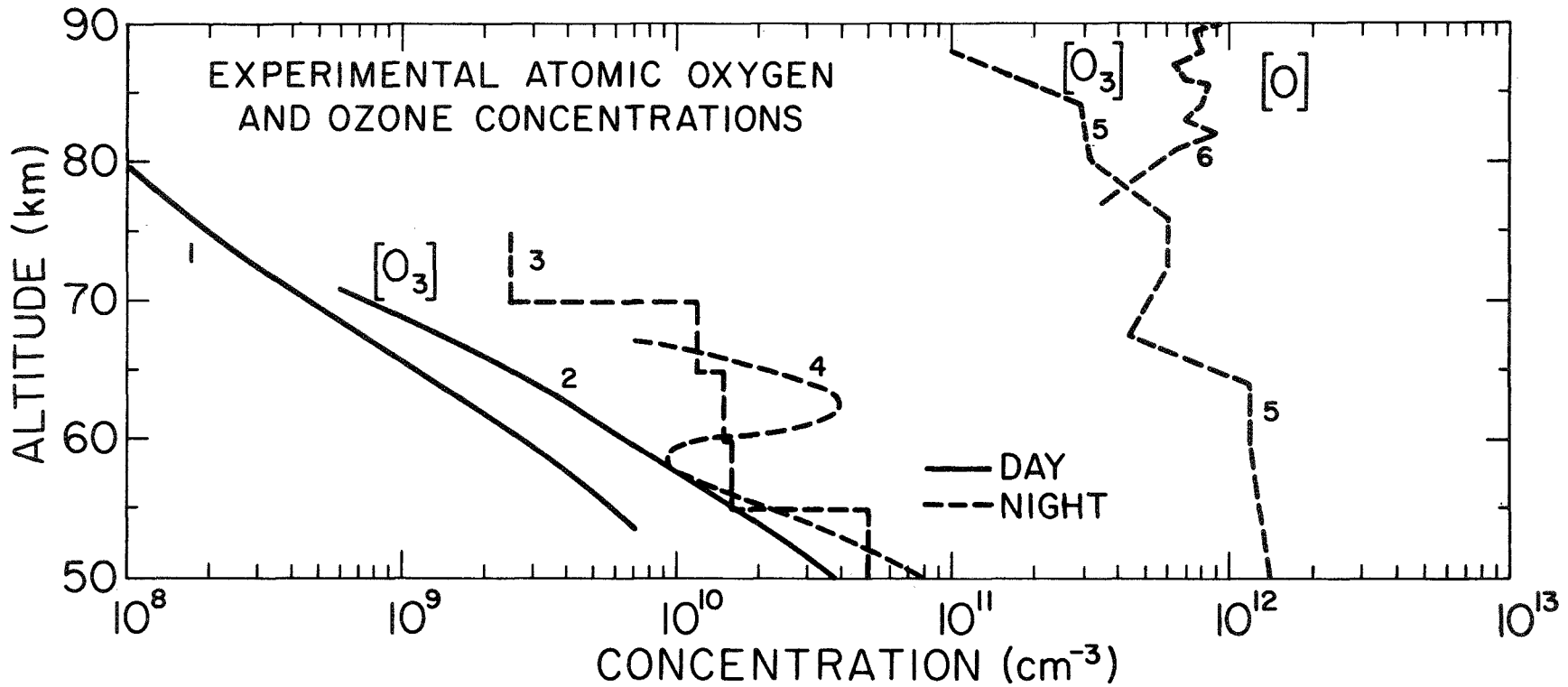
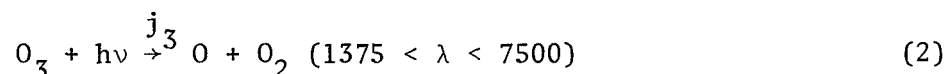
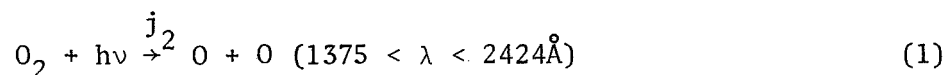


Figure 2. Experimental profiles of atomic oxygen and ozone concentrations. 1. Daytime ozone concentration by Johnson, et al. (1954); 2. Daytime ozone concentration by Rawcliffe, et al. (1963); 3. Nighttime ozone concentration by Carver, et al. (1966); 4. Nighttime ozone concentration by Reed (1968); 5. Nighttime ozone concentration by Mikirov (1965) and 6. Atomic oxygen nighttime concentration by Fedynskii, et al. (1967).

$O_2$  and  $O_3$ , dissociation cross sections as a function of the radiation wavelength, and rates of the chemical reactions involved. The latter represents the less known factor.

The basic photochemical reactions of atmospheric oxygen are:



Cross sections and fluxes of radiation involved in reactions (1) and (2) are shown in Figures 3 and 4.

The diversity of published laboratory values for the rate coefficients of (3), (4) and (5) is very wide. It reflects the complications introduced in the experimental determinations by the effect of hydrogenous and nitrogenous trace impurities and the formation of metastable species (Schofield, 1967).

A recent discussion of experimental studies of reactions (3), (4) and (5) has been given by Kaufman (1967). This author adopts for the range between 190 and 400°K, the value of  $k_3 = 3.0 \times 10^{-33} \left(\frac{T}{300}\right)^{-(2.6 \pm 0.4)}$  obtained by combining several data when  $M \equiv N_2$ . Also  $k_5 = (1.4 \pm 0.3) \times 10^{-12} \exp \left[ -\frac{(3.0 \pm 0.4)}{RT} \right]$  is adopted by the same author. It must be noted that some discrepancies among different experimental studies of these reactions still exists. Table 1 shows the individual results of some of these studies



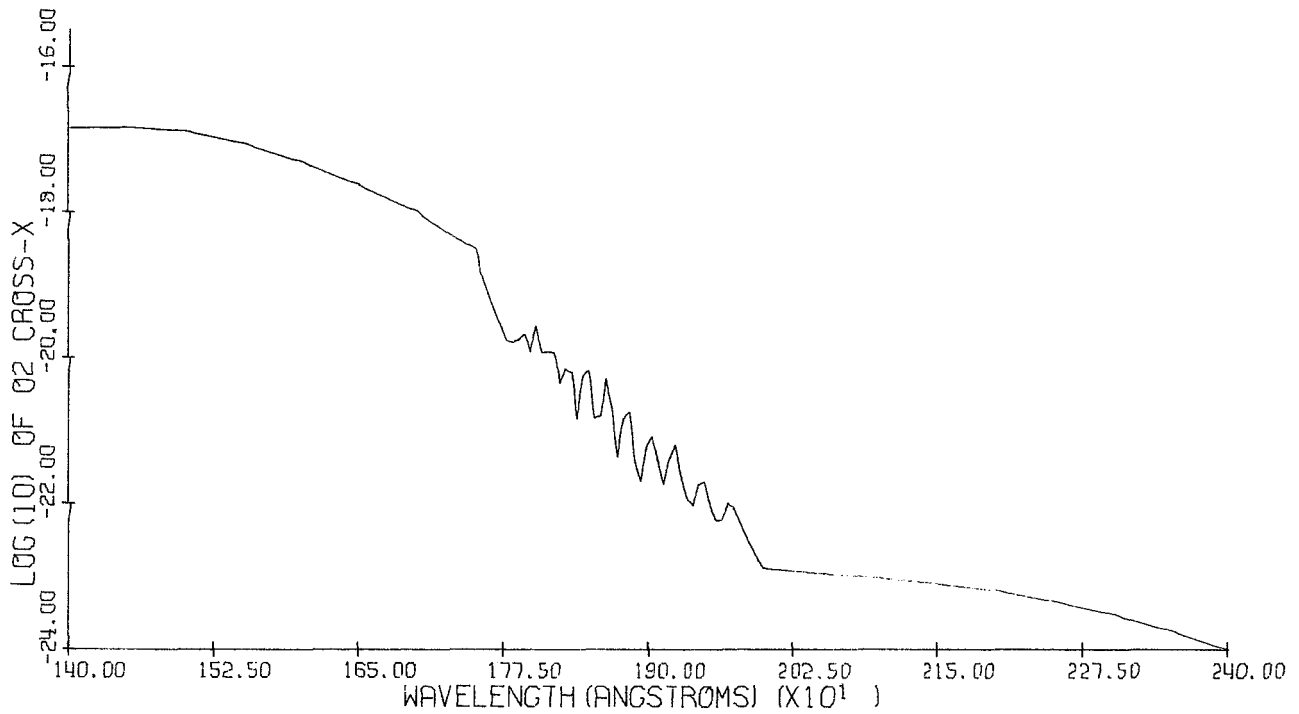


Figure 3a. O<sub>2</sub> absorption cross sections as a function of radiation wavelength (see Appendix for references).

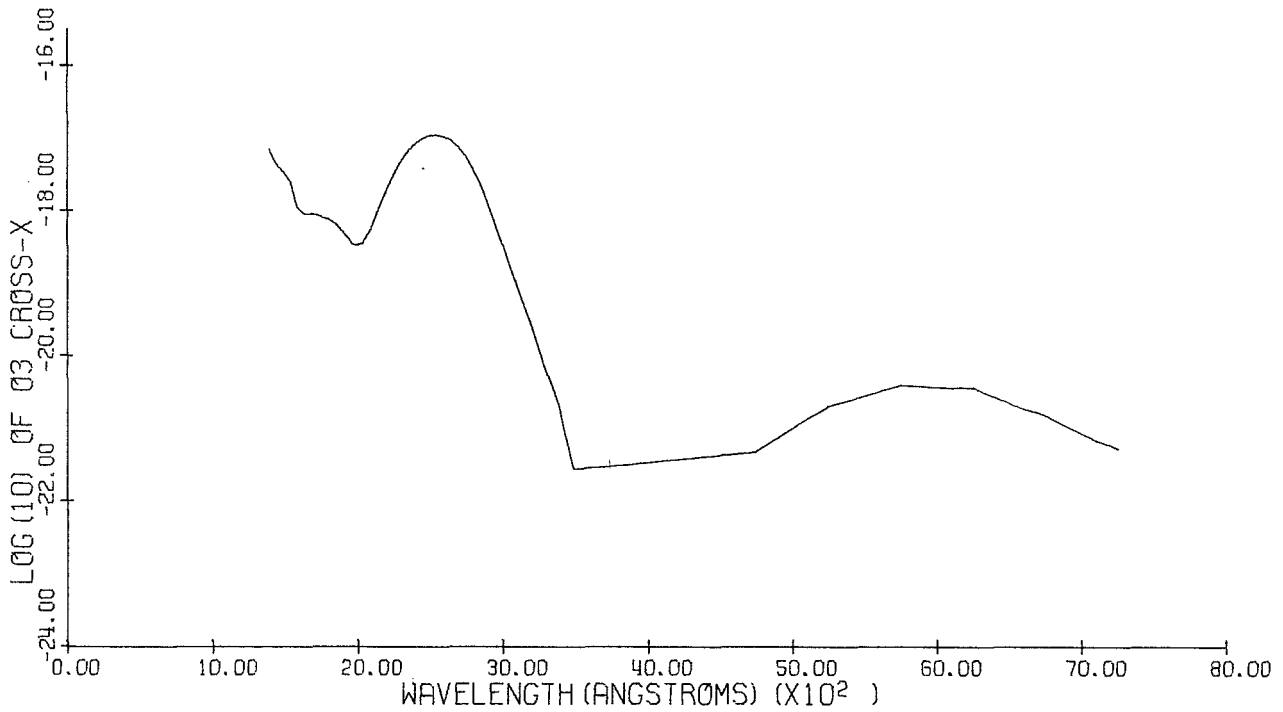


Figure 3b. Ozone absorption cross sections as a function of radiation wavelength (see Appendix for references).

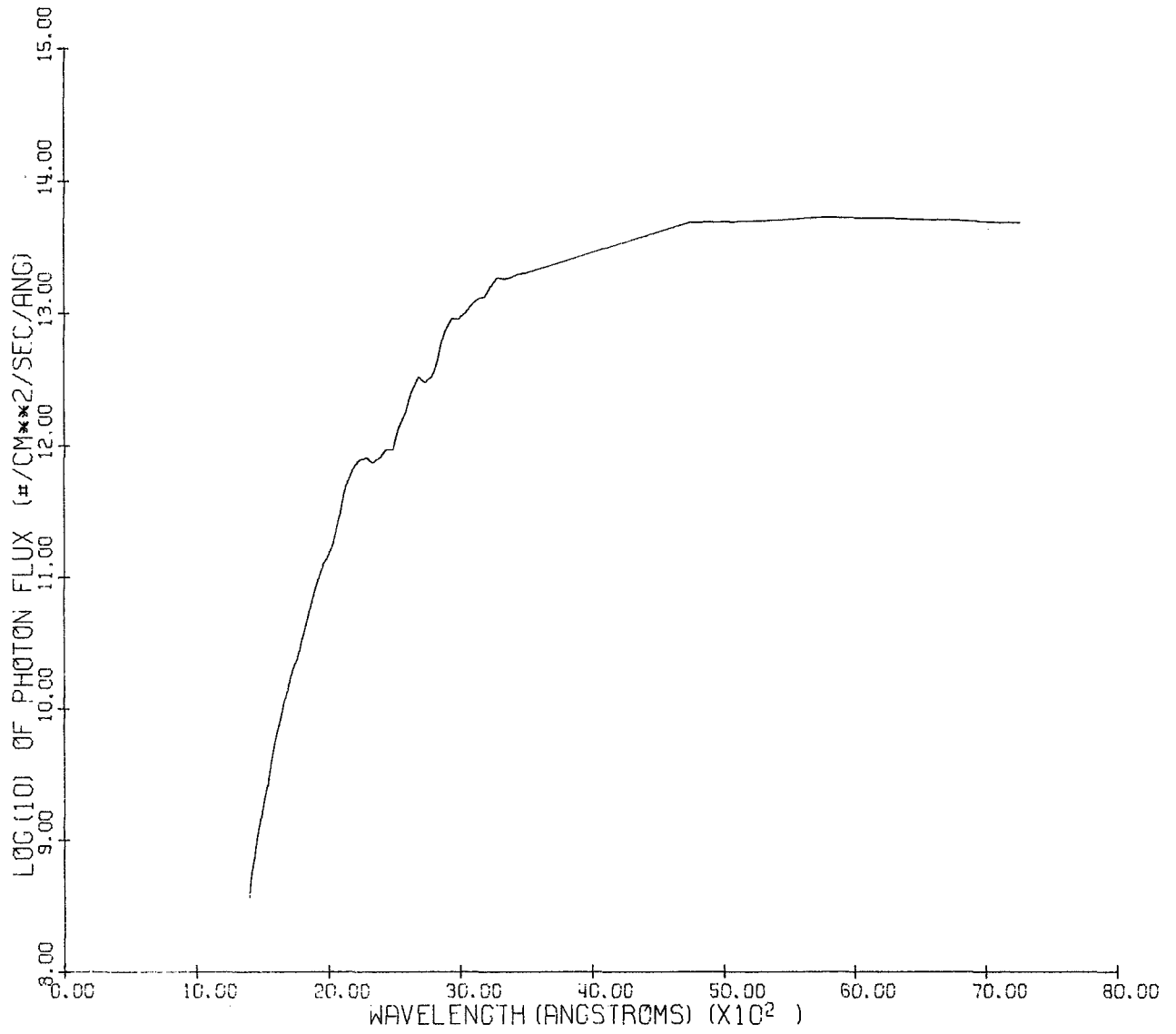


Figure 4. Solar flux radiation as a function of wavelength (see Appendix for references).

TABLE 1

Reaction rate coefficients for oxygen including Kaufman's adopted values

10

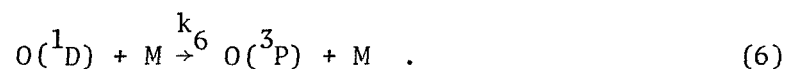
<u>Reaction</u>	<u>Rate Coefficient</u>	<u>Author</u>	<u>Comments</u>
$O + O + M \xrightarrow{k_5} O_2 + M$	$2.8 \times 10^{-33} \text{ (cm}^6 \text{ sec}^{-1}\text{)}$	Morgan and Schiff (1963)	$T = 294^\circ\text{K}$ , $M \equiv N_2$
	$1.5 \times 10^{-34}$	Young and Black (1966)	$T = 300^\circ\text{K}$ , $M \equiv O$
	$3.0 \times 10^{-33} \left(\frac{T}{300}\right)^{-2.9 \pm 0.4}$	Campbell and Thrush (1968)	$190 < T < 350^\circ\text{K}$ , $M \equiv N_2$
	$3.0 \times 10^{-33} \left(\frac{T}{300}\right)^{-2.9 \pm 0.4}$	Kaufman (1967)	adopted by $M \equiv N_2$
$O + O_2 + M \xrightarrow{k_4} O_3 + M$	$8.2 \times 10^{-35} \exp\left(\frac{890}{RT}\right) \text{ (cm}^6 \text{ sec}^{-1}\text{)}$	Benson and Axworthy (1965)	near $T = 360^\circ\text{K}$ , $M \equiv O_2$
	$(2.6 \pm 0.3) \times 10^{-35}$	Jones and Davidson (1962)	$M \equiv N_2$
	$\exp\left(\frac{1700 \pm 300}{RT}\right)$		
	$6.5 \times 10^{-34}$	Kaufman and Kelso (1964)	$T = 300^\circ\text{K}$ , $M \equiv O_2$
	$5.5 \times 10^{-34}$	Kaufman and Kelso (1964)	$T = 300^\circ\text{K}$ , $M \equiv N_2$
	$2.5 \times 10^{-35} \exp\left(\frac{1800 \pm 400}{RT}\right)$	Clyne <u>et al.</u> (1965)	$188 < T < 373^\circ\text{K}$ , $M \equiv \text{Ar}$
	$5.5 \times 10^{-34} \left(\frac{T}{300}\right)^{-2.6 \pm .4}$	Kaufman (1967)	adopted by, $M \equiv N_2$
$O + O_3 \xrightarrow{k_5} O_2 + O_2$	$5.6 \times 10^{-11} \exp\left(-\frac{5700}{RT}\right) \text{ (cm}^2 \text{ sec}^{-1}\text{)}$	Benson and Axworthy (1965)	$340 < T < 390^\circ\text{K}$
	$3.9 \times 10^{-14}$	Phillips and Schiff (1962)	Considered questionable value by Schofield (1967)
	$8.0 \times 10^{-15}$	Mathias and Schiff (1964)	$T = 300^\circ\text{K}$
	$(1.4 \pm 0.3) \times 10^{-12}$	Kaufman (1967)	adopted by
	$\exp\left(-\frac{3000 \pm 400}{RT}\right)$		

 $R = 1.987 \text{ cal}/(^{\circ}\text{K mole})$

for the range of atmospheric temperatures, together with the data adopted by Kaufman (1967).

Theoretical models of O and O<sub>3</sub> have been calculated by several authors. Some of them considered a photochemical reaction scheme based upon an oxygen alone atmosphere as represented by reactions (1) and (5), (Craig, 1950; Hunt, 1966a; London, 1967; Maeda and Aikin, 1968). However, some of these authors are mainly concerned with stratospheric ozone. It must be noted that the rate coefficients used in the above-mentioned studies are notably different from the values adopted by Kaufman (1967) using all experimental data available at this time.

Models of O and O<sub>3</sub> that take into account reactions of hydrogenous compounds in the chemical scheme have been published by Hampson (1965) and Hunt (1966b). However, they have used reaction rate coefficients that differ sensibly from the ones that are considered correct at this time. Of particular importance, besides the basic oxygen reactions rates already discussed, is the rate coefficient of reaction:



Hunt (1966a) assumed a value of  $k_6 = 10^{-12}$  (cm<sup>3</sup> sec<sup>-1</sup>). A more realistic one appears to be:  $k_6 = 5 \times 10^{-11}$  as deduced from McGrath and McGarvey (1967).

A more sophisticated approach is to consider an oxygen-hydrogen atmosphere including also the dynamics terms. Konashenok (1967) and Hesstvedt (1968) introduced vertical eddy transport in their calculations. From the work of these authors it is clear that the dynamical terms can change the vertical and geographic distribution of atomic oxygen and ozone. It must be

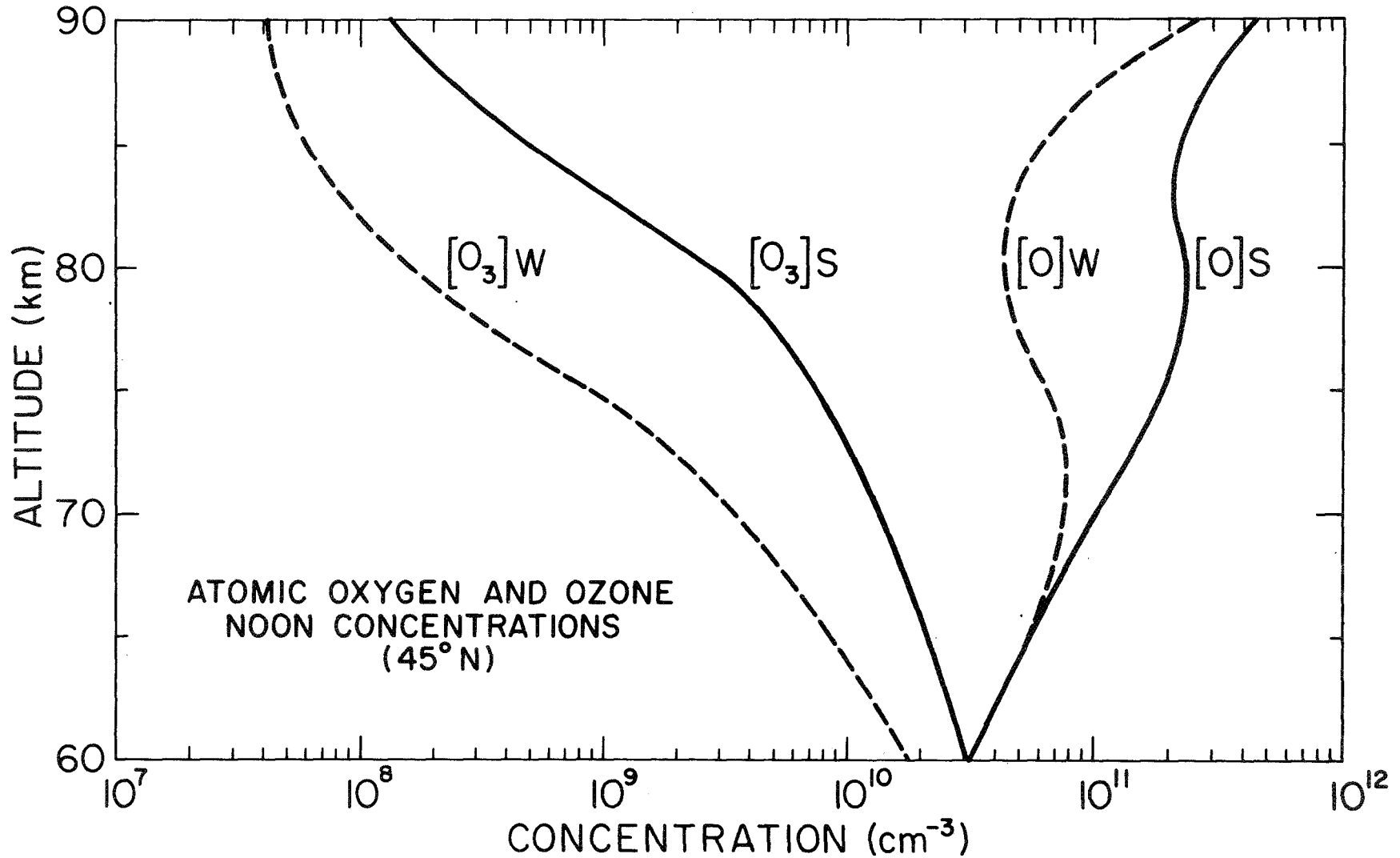


Figure 5. Theoretical profiles of noon atomic oxygen and ozone concentration at 45° N. Full line indicates summer profiles; broken line indicates winter profiles. These values were computed for oxygen alone atmosphere.

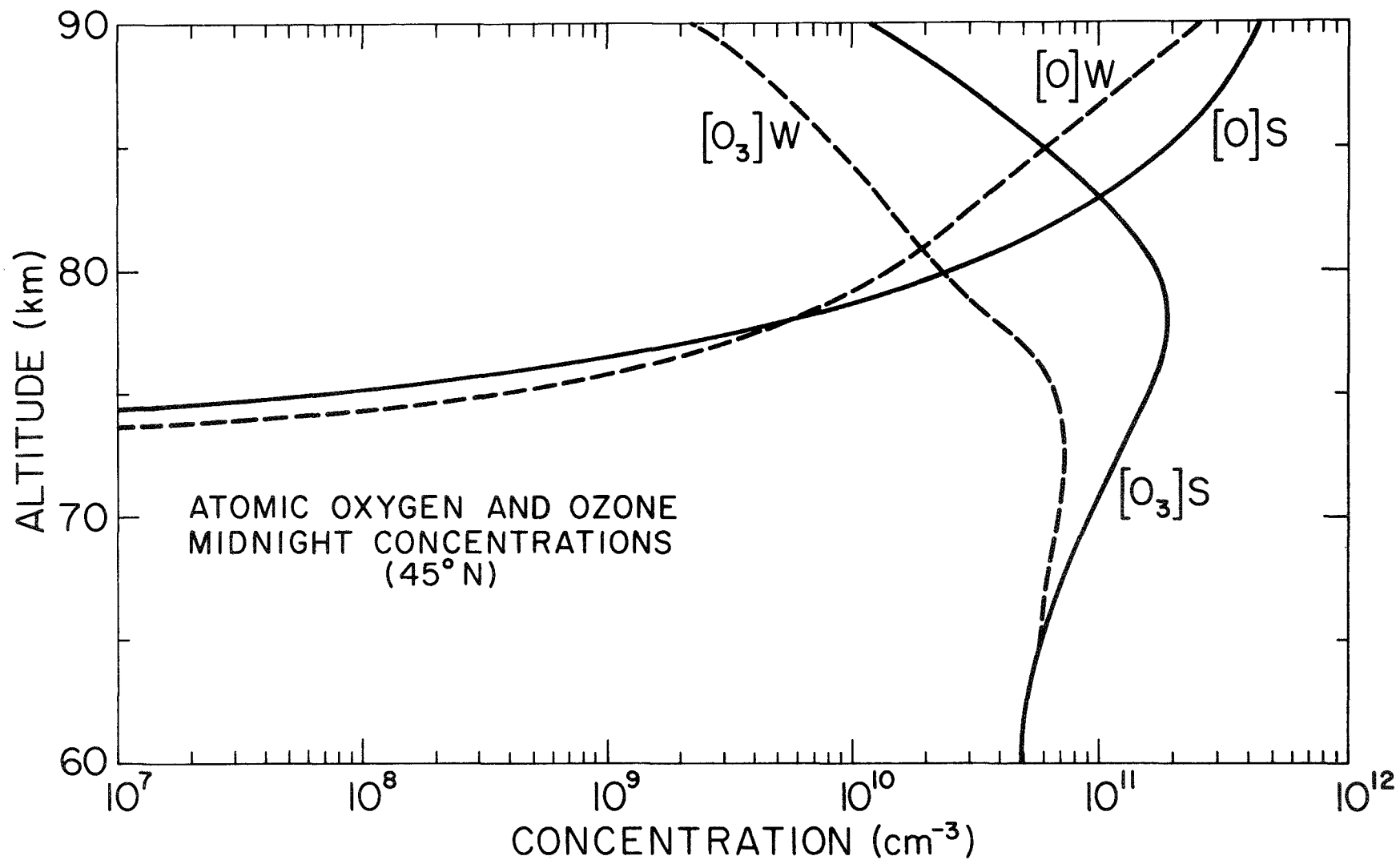


Figure 6. Theoretical profiles of midnight atomic oxygen and ozone concentration at 45° N. Full line indicates summer profiles; broken line indicates winter profiles. These values were computed for oxygen alone atmosphere.

noted however that the reaction rate coefficients of reactions (3) and (5) adopted in all the studies mentioned above, differ notably from the Kaufman (1967) values which, at this time, should be considered more representative.

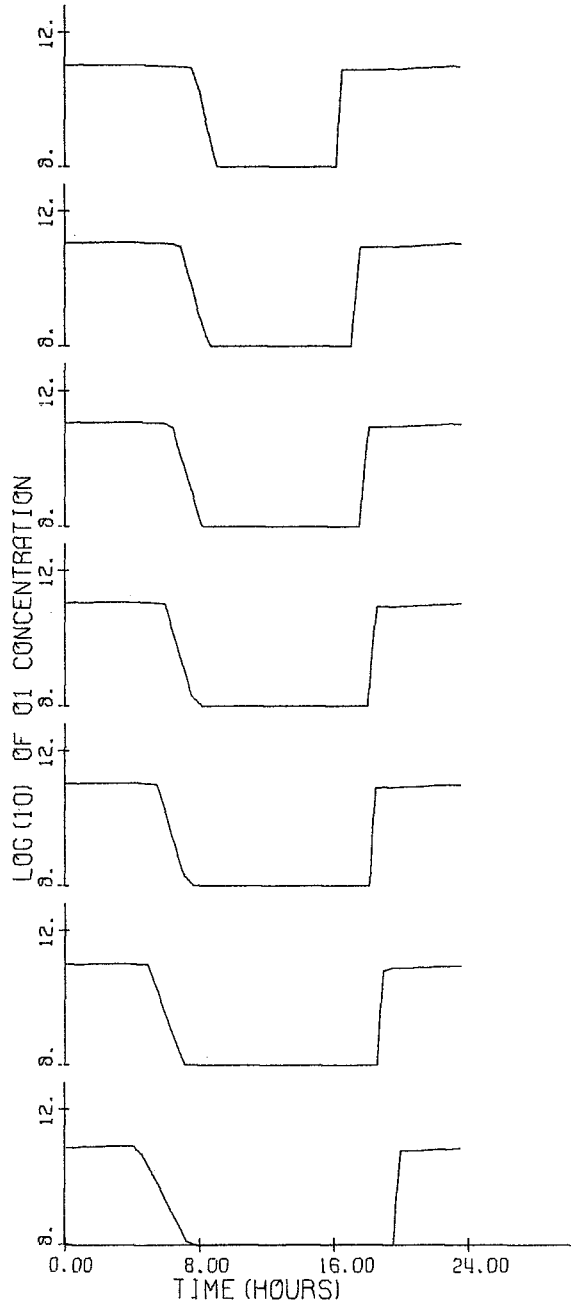
It is convenient to compute the temporal and spatial variations of O and O<sub>3</sub> based on the Kaufman rate coefficients for the reactions (3) and (5) and also on recent values of solar fluxes. In this way a basic photochemical model and a function of time and latitude is obtained. It can be changed as a second step, by introducing hydrogenous compound reactions and dynamical terms. Time variations as a function of height, latitude and season are computed by solving the continuity equations in the differential forms:

$$\frac{d[O]}{dt} = 2j_2 [O_2] + j_3 [O_3] - 2k_3 [O]^2 [M] - k_4 [O][O_2] [M] - k_5 [O][O_3] \quad (7)$$

$$\frac{d[O_3]}{dt} = k_4 [O][O_2] [M] - k_5 [O][O_3] - j_3 [O_3] \quad (8)$$

The computer programs used and the input values of cross sections, solar radiation fluxes, atmospheric temperature, and densities adopted in the calculations are given in the Appendix. Detailed tabulation of the output data (atomic oxygen and ozone concentrations) are also given there. Figures 5 and 6 show the computed O and O<sub>3</sub> profiles for 45° of latitude north, summer and winter, and noon and midnight. Figures 7, 8 and 9 show the latitudinal variation of atomic oxygen and ozone concentrations at 70, 80 and 90 km. A full discussion of these data will be given elsewhere.

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LAT=45.0



ALT=70.0  
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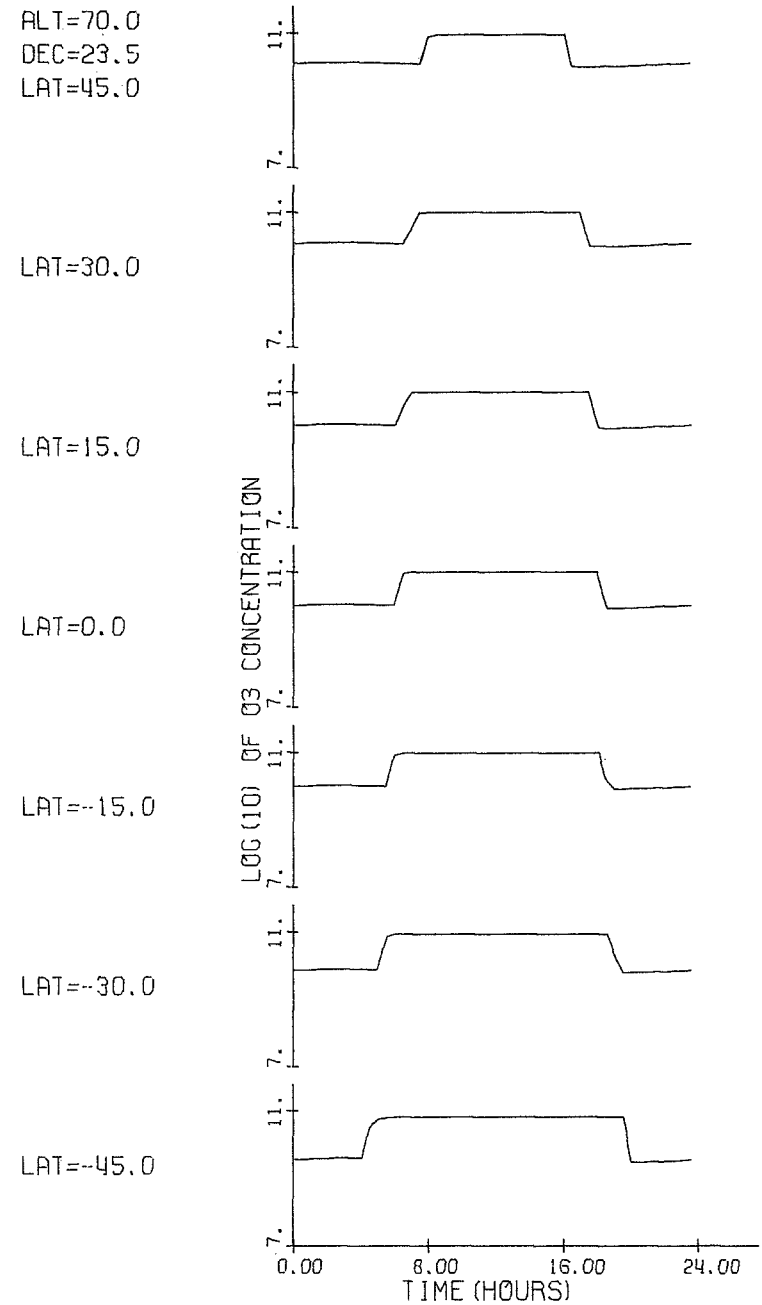


Figure 7. Theoretical values of diurnal latitudinal and seasonal variation of atomic oxygen and ozone for an oxygen alone atmosphere (70 km).



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LAT=30.0

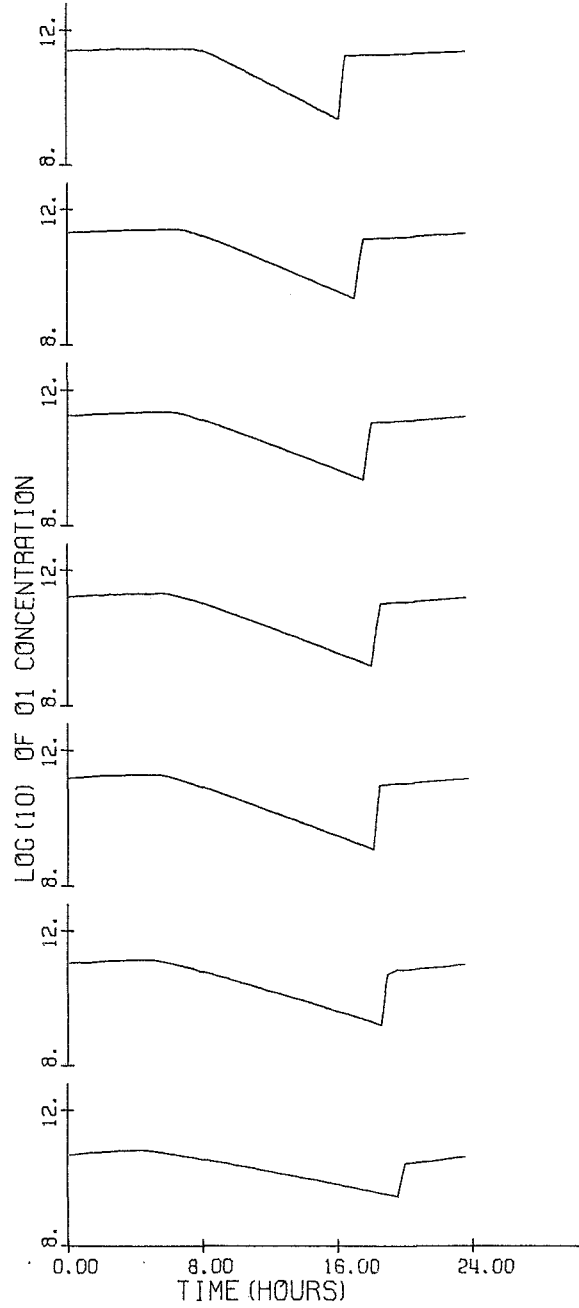
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LAT=-45.0



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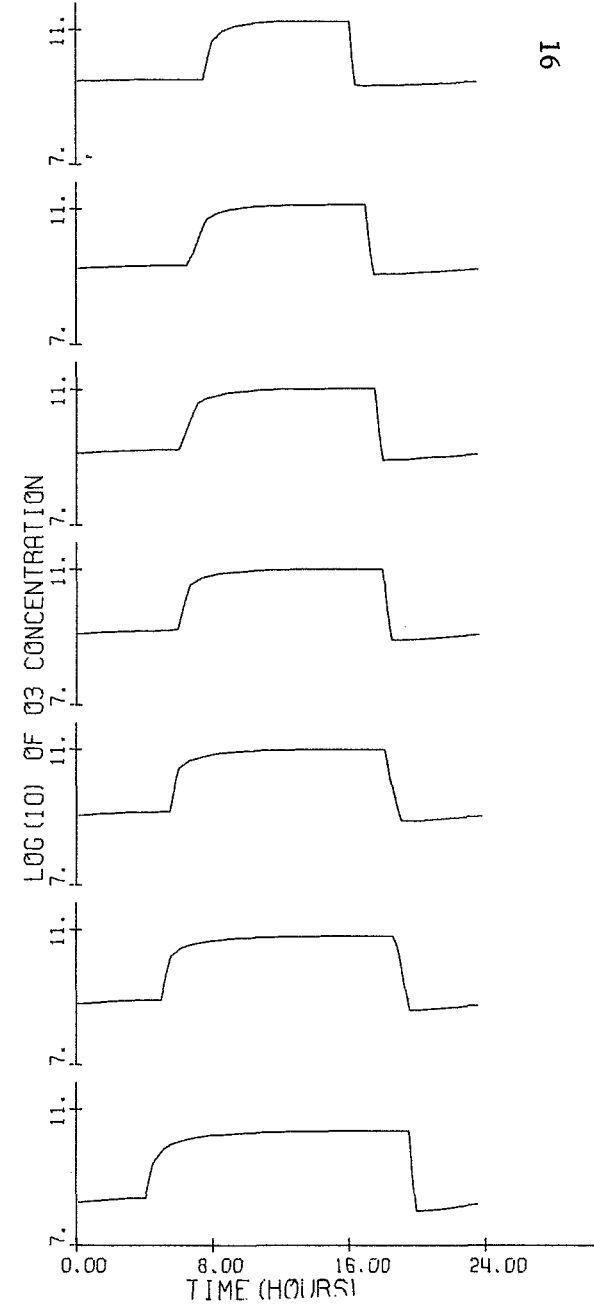


Figure 8. Theoretical values of diurnal latitudinal and seasonal variation of atomic oxygen and ozone for an oxygen alone atmosphere (80 km).

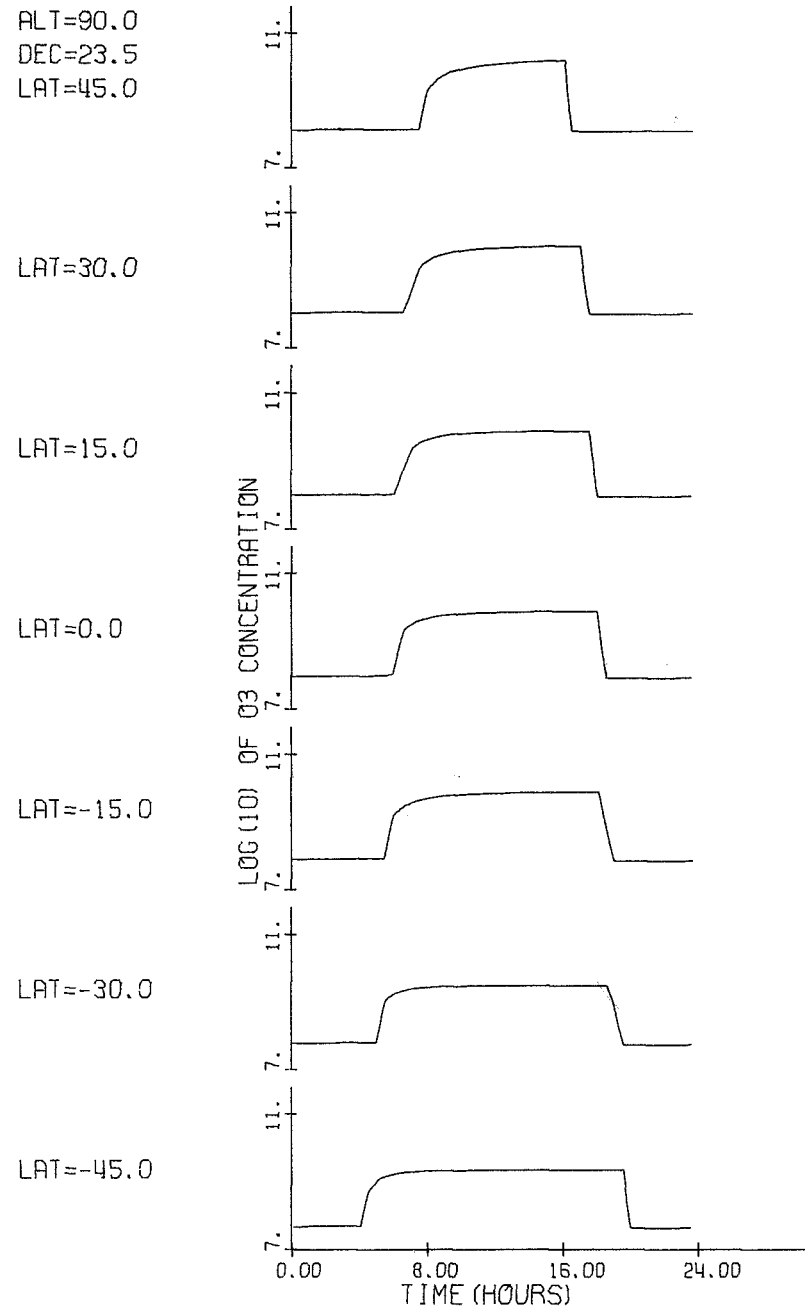
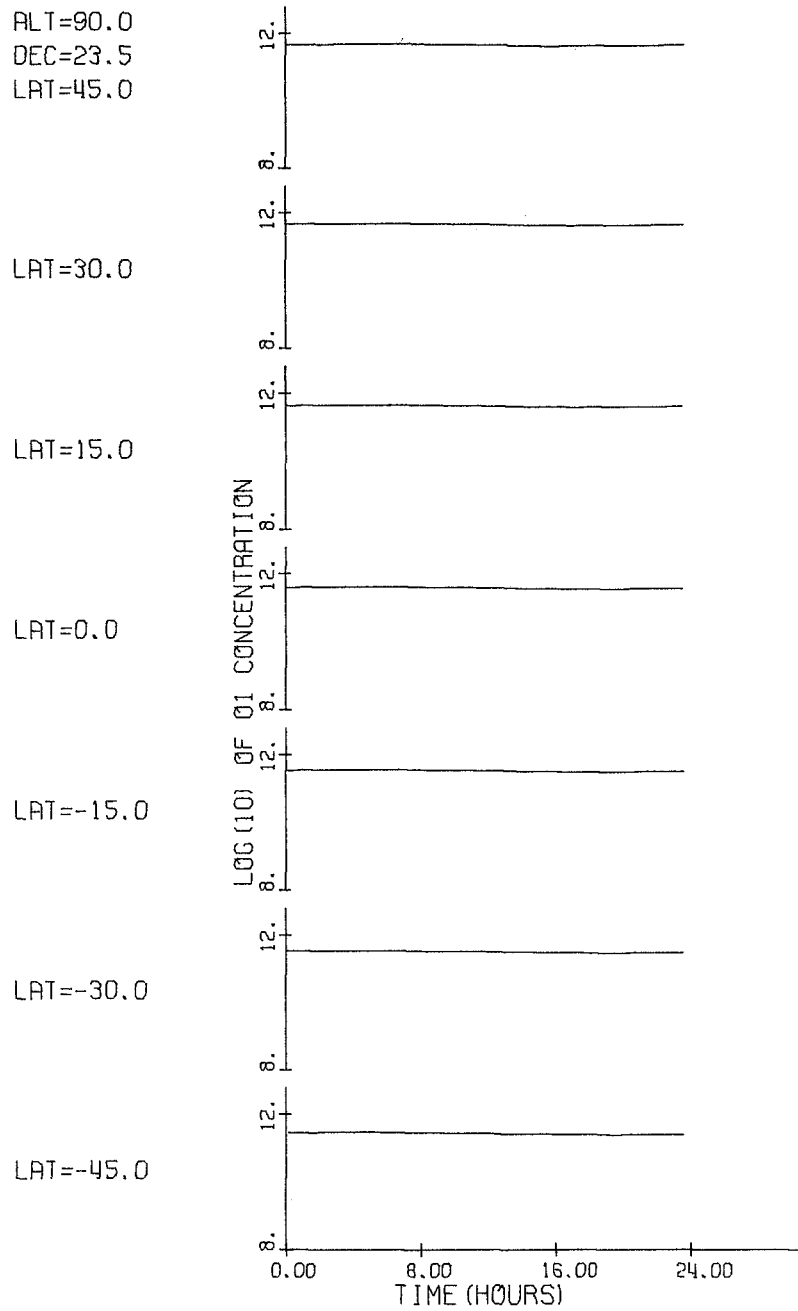


Figure 9. Theoretical values of diurnal latitudinal and seasonal variation of atomic oxygen and ozone for an oxygen alone atmosphere (90 km).

### 2.3 Nitrogen Oxides

Minor constituents of great importance to the chemical aeronomy and, in particular, to the D-region ionization are nitric oxide and, in a minor scale, nitrogen dioxide. NO is the relevant specie to the chemistry of the D region as considered by Nicolet and Aikin (1960). Their paper represents one of the first attempts to explain theoretically what was known at the end of the fifties about the D region. These authors introduced the hypothesis that the basic process of formation during low solar activity was the ionization of NO by Lyman-alpha solar radiation with a production peak between 75 and 80 km. Until 1964 no direct experimental data on the nitric oxide distribution were published. The only methods used to derive NO concentration were the indirect ionospheric estimates as the ones obtained by Mitra (1966) and Wagner (1966) for D-region heights. A final model deduced using the same methods has been published recently by Mitra (1968). His NO distribution is expressed by:

$$[\text{NO}] = 4 \times 10^{-1} \exp\left(-\frac{3700}{T}\right) [\text{O}_2] + 5 \times 10^{-7} [\text{O}] \quad (9)$$

Figure 10 shows Mitra's (1968) distribution of NO.

Direct photometric measurements of the NO airglow were made by Barth (1964, 1966) from Wallops Island during low solar activity. The data obtained from a rocket flight in November 1963 shows an NO concentration of  $6.2 \times 10^7 \text{ cm}^{-3}$  at 75 km and  $6.0 \times 10^7$  at 95 km. These values are much higher than the ones deduced by indirect ionospheric methods. For some time the Barth (1966) values were considered doubtful, but a recent rocket photometric experiment by Pearce (1968) again shows very high values of nitric oxide concentration at all heights in the mesosphere. His values are also plotted in Figure 10.

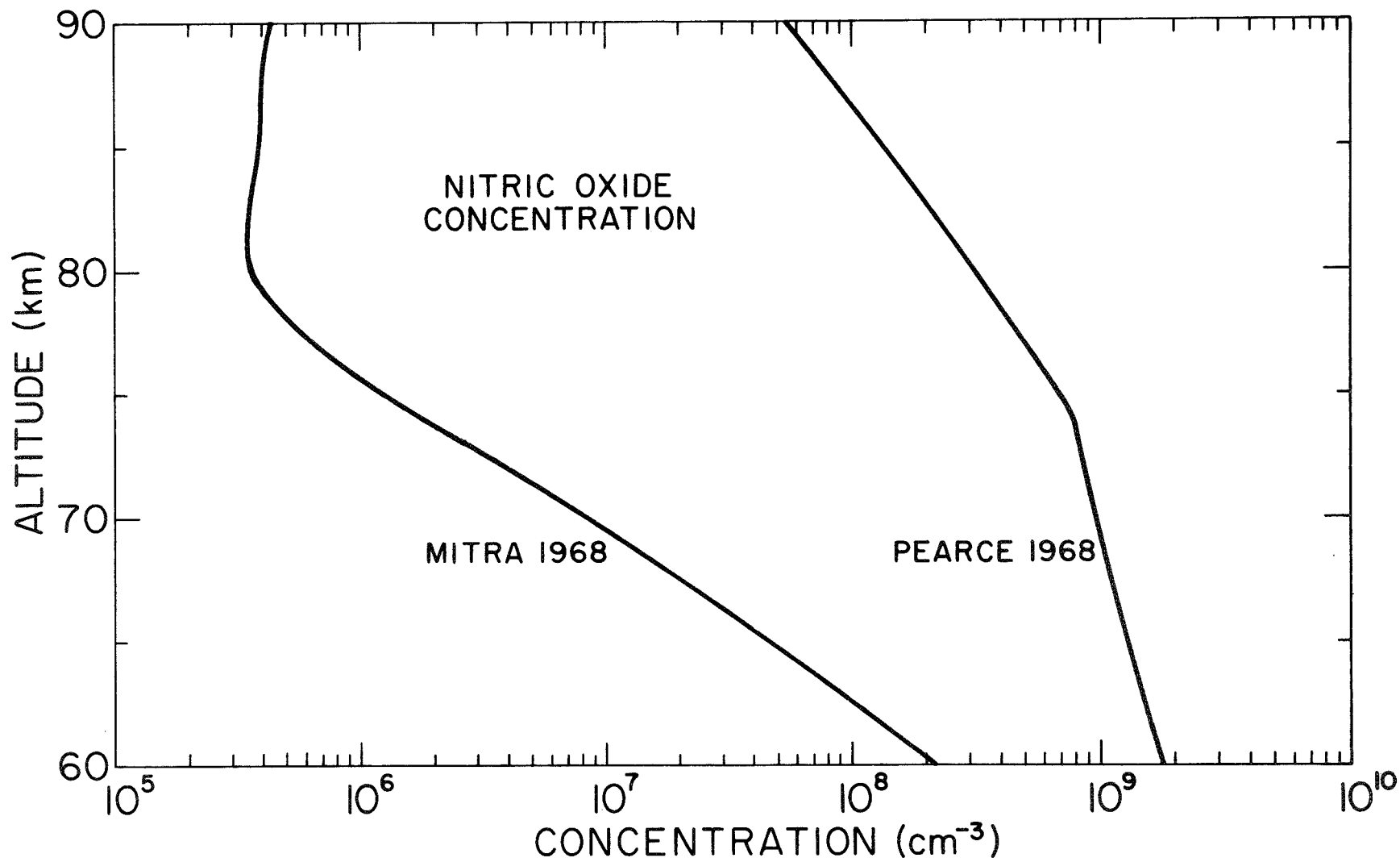
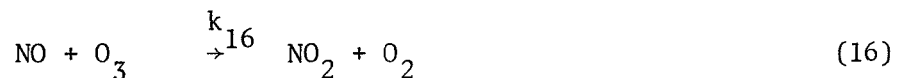
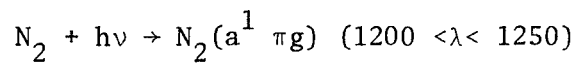


Figure 10. Nitric oxide concentration as a function of height by Mitra (1968) and Pearce (1968).

The more comprehensive theoretical study of the NO distribution has been published by Nicolet (1965a,b). This author shows the importance of atomic oxygen and ozone as reactants in the chemistry of the oxides of nitrogen in the atmosphere. He shows also that in order to conciliate the values given by Barth (1964) it was necessary to introduce ionic reactions in the chemical scheme.

Nitrogen dioxide distribution in the atmosphere have not been determined experimentally. However its chemistry is strongly related to that of nitric oxide.

The basic reaction scheme of the oxides of nitrogen appears to be (Nicolet, 1965a,b) as follows:



Rate coefficients for these reactions are given in Table 2.

TABLE 2  
Reaction rate coefficients for the oxides of nitrogen

<u>Reaction</u>	<u>Reaction Rate Coefficient</u>	<u>Source</u>
$N + O_3 \xrightarrow{k_{11}} NO + O_2$	$5.7 \times 10^{-13}$	Phillips and Schiff (1962)
$N + O + M \xrightarrow{k_{12}} NO + M$	$1.1 \times 10^{-33} \left(\frac{T}{300}\right)^{-0.5 \pm 0.2}$	Kaufman (1968)
$N + O_2 \xrightarrow{k_{13}} NO + O$	$(1.4 \pm 0.2) \times 10^{-11}$ $\exp - \left(\frac{7100 \pm 400}{RT}\right)$	Clyne and Thrush (1961)
$N + NO \xrightarrow{k_{15}} O + N_2$	$(2.2 \pm 0.6) \times 10^{-11}$	Phillips and Schiff (1962)
$NO + O_3 \xrightarrow{k_{16}} NO_2 + O_2$	$9.5 \times 10^{-13} \exp - \left(\frac{2450 \pm 150}{RT}\right)$	Clyne <u>et al.</u> (1964)

R = 1.987 K cal (°K mole)

Mesospheric distribution of NO obtained by using the above-mentioned radiation scheme give values much lower than the experimental data obtained by Barth (1966) and Pearce (1968). A possible way of explaining high values of NO concentration around 80 km is by taking into account vertical downward motions of NO from the lower thermosphere as discussed recently by Geisler and Dickinson (1968). Large amounts of nitric oxide can be produced below 120 km by the ion-molecule reaction:



as shown by Ryan (1968).

#### 2.4 Hydrogenous Compounds

The basic hydrogenous compound in the atmosphere below 90 km is water vapor. The only experimental data available on the concentration of water vapor are the data by Fedynskii et al. (1967) which show very large values in the mesosphere up to an altitude of approximately 85 km. Water vapor is dissociated by solar radiation of wavelength less than 2400Å:



Both atomic hydrogen and hydroxyl are chemically very active and are capable of undergoing a very complex series of reactions with atomic oxygen and ozone. The problem of the presence and distribution of hydrogenous compounds in the mesosphere and stratosphere has been studied in detail by Hesstvedt (1964, 1965, 1968). The chemistry of water vapor is closely related to the chemistry of oxygen. For example, the reaction:



is capable of reducing the concentration of ozone in the mesosphere as indicated originally by Nicolet (1964) and shown by Hunt (1966b).

Other basic reactions of hydrogen and hydrogenous compounds are:



Table 3 gives the rate coefficients of these reactions. Hydrogen peroxide is an additional compound that could be present in the mesosphere but in very small concentration. Its presence gives rise to a large number of secondary reactions. A very detailed chemical scheme has been given recently by Hesstvedt (1968). In the same paper, theoretical models of hydrogen and hydrogenous compounds, distributions are shown. As a comment to these models, it should be noted that when dynamical vertical transport is taken into account, the dominant hydrogenous constituent is water vapor up to a height of about 85 km. This fact indicates that H and OH (both very active and important for the oxygen chemistry) are less abundant than expected, reducing their effect on the distribution of other species.



TABLE 3

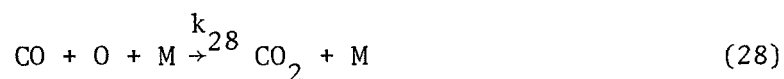
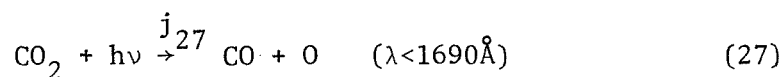
Reaction rate coefficients for hydrogen and hydrogenous compounds

<u>Reaction</u>	<u>Reaction Rate Coefficient</u>	<u>Source</u>
$H + O_3 \xrightarrow{k_{20}} OH + O_2$	$(2.6 \pm 0.5) \times 10^{-11}$	Phillips and Schiff (1962)
$OH + O \xrightarrow{k_{21}} H + O_2$	$(5 \pm 2) \times 10^{-11}$	Kaufman (1968)
$H + O_2 + M \xrightarrow{k_{22}} HO_2 + M$	$3 \times 10^{-32} \left(\frac{273}{T}\right)^{1.3}$	Schofield (1967)
$HO_2 + O \xrightarrow{k_{23}} OH + O_2$	$\geq 10^{-11}$	Kaufman (1964)
$HO_2 + O_3 \xrightarrow{k_{25}} OH + 2O_2$	$\sim 10^{-14}$ (estimate)	Hunt (1963)

## 2.5 Carbon Dioxide

The importance of carbon dioxide in the heat balance of the region below 90 km has been recognized for a long time (Gowan 1947a,b). In their study of the photochemistry of the atmospheric CO<sub>2</sub>, Bates and Witherspoon (1952) deduced that carbon dioxide should be in constant mixing ratio up to 100 km, approximately. A recent discussion of the role of CO<sub>2</sub> in the mesosphere has been given by Craig (1965). However, no detailed calculations have been made of the distribution variations that can be expected. Also, no direct measurements of this constituent above the troposphere are available at this time.

The photochemistry of CO<sub>2</sub> can be expressed basically by the reactions:



Carbon dioxide absorption cross sections are smaller by an order of magnitude with respect to the equivalent molecular oxygen cross sections. The rate coefficient  $k_{28}$  is still doubtful, (Kaufman, 1968) but is considered very slow. At 293°K Clyne and Thrush (1962) found  $K_{28} < 8 \times 10^{-35}$ .

In summary, it can be said that the chemistry of atmospheric carbon oxides is at this time practically unknown.

### 3. RATES OF IONIZATION

#### 3.1 Introduction

Several sources of ionization can act in the atmosphere below 90 km. The most important one is considered to be, at this time, solar Lyman-alpha radiation ionizing nitric oxide. However, the uncertainties about the mesospheric concentration of this constituent which were mentioned in the last chapter, make difficult a correct estimate of the relative importance of Lyman-alpha and other radiation in the ionization of the D region. Solar X-rays and ultraviolet radiation of 1027-1118Å wavelengths between 85 and 90 km, and cosmic rays below 70 km, are capable of ionizing the atmospheric gas. At night, scattered Lyman-alpha radiation and precipitated electrons could be effective sources of ionization between 70 and 90 km.

#### 3.2 Ionization by Solar Lyman-Alpha

A very intense radiation is emitted from the sun at a wavelength of 1215.7Å. Its photon flux is equal, for average low solar activity, to 2.7 photons  $\text{cm}^{-2} \text{sec}^{-1}$  equivalent to 6 erg  $\text{cm}^{-2} \text{sec}^{-1}$ . The hydrogen Lyman-alpha line originates in the absorbing region between the photosphere and the chromosphere. The kinetic energy of solar protons in that region is lost through the radiative recombination with electrons. Lyman-alpha radiation is mostly absorbed by  $\text{O}_2$  and an interesting fact is that one of the absorption "windows" of molecular oxygen atmosphere corresponds to the same wavelength as shown in Figure 11. The corresponding absorption cross section is  $\sigma_{\text{O}_2} = 1.04 \times 10^{20} \text{ cm}^2$  (Watanabe, 1958), Lyman-alpha radiation can penetrate as low as 65-70 km. The only atmospheric constituent that can be ionized by this radiation, below 90 km, is nitric oxide, and it becomes a strong

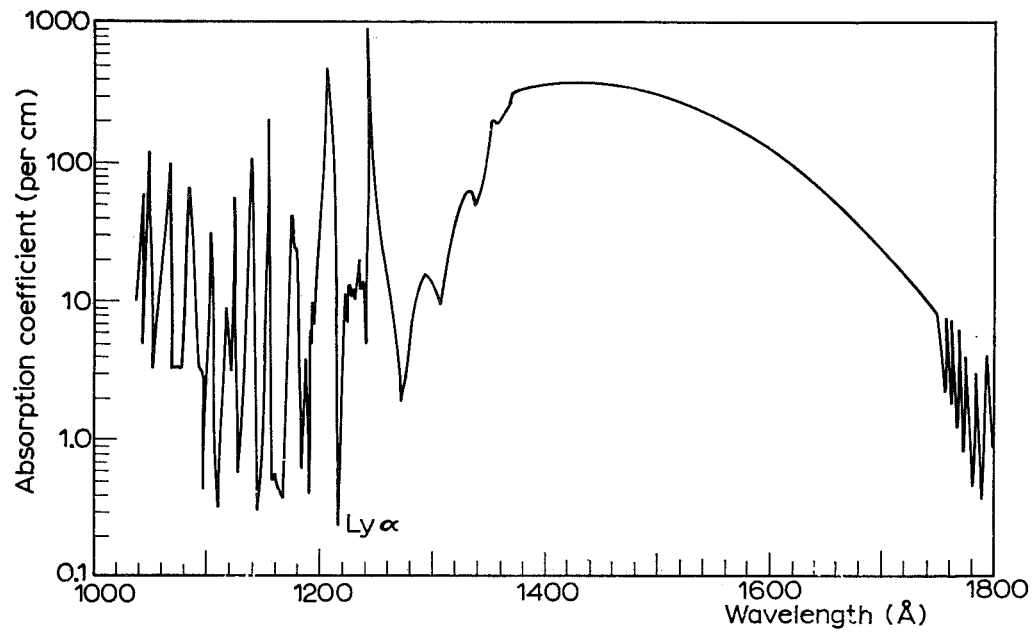


Figure 11. Absorption coefficient of molecular oxygen in the far ultraviolet (Watanabe, et al., 1953).

source of ionization in the D region. NO ionization cross section is  $\sigma'_{NO} = 2 \times 10^{-18} \text{ cm}^2$  (Watanabe, 1954). With cross section and flux data it is possible to compute the rate of NO ionization by Lyman-alpha. This is done by using the expression:

$$q_{L\alpha}(h) = \sigma'_{NO} [NO] \exp(- \sec \chi \sigma_{O_2} \int_h^{\infty} d[O_2] d h) \quad (29)$$

where  $q_{L\alpha}$  indicates the rate of ionization,  $h$  the altitude, brackets indicate concentration, and  $\chi$  is the solar zenith angle.  $\sec \chi$  is substituted by the Chapman function for angles greater than  $80^\circ$ .

By using Mitra (1968) and Pearce (1968) nitric oxide distributions shown in Figure 3, different rates are computed and are drawn in Figure 12. The Pearce data results in a very large rate of ionization which overdraw the effects of X-rays and in part, of cosmic rays as effective sources of ionization as will be seen below.

### 3.3 Ionization by X-Rays

Quasi-thermal X-ray emissions of wavelength  $< 100\text{\AA}$  originates in the corona of the sun. This radiation is highly variable as is shown in Figure 13, reproduced from de Jager (1967).

Solar X-rays of wavelength between 1 and  $40\text{\AA}$  can penetrate in the D region down to an altitude of 50 km, ionizing all the constituents of the mesosphere. However, the low photon fluxes available during low solar activity are not capable of producing important ionization below 85 km. Under such conditions the X-ray contribution to the formation of the D region is limited to its upper part. During solar flares when the X-ray fluxes increase by several orders of magnitude, this radiation increases greatly the ionization of all the D region.

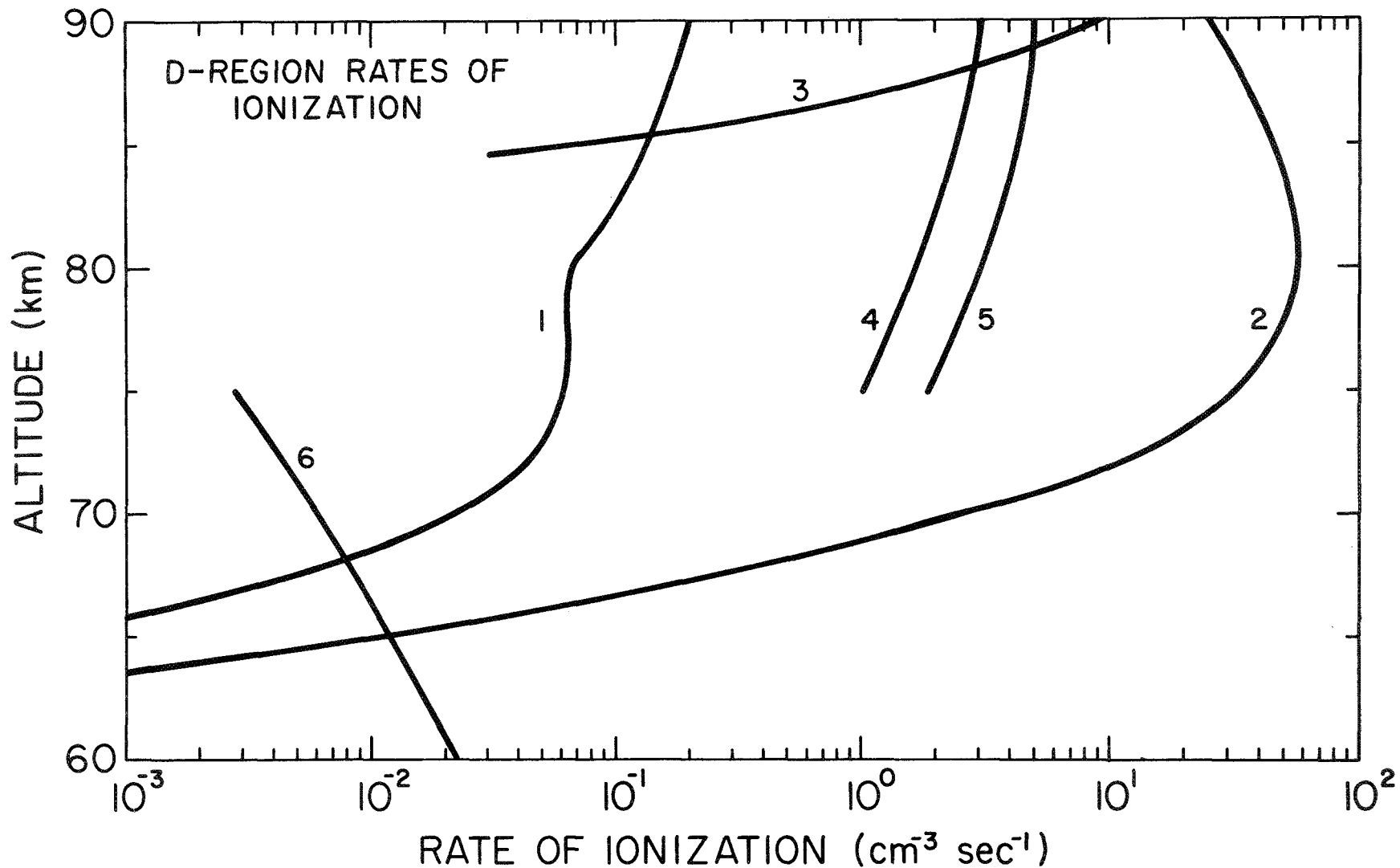


Figure 12. Daytime D-region rates of ionization due to different sources. 1 is the ion-pair production by  $L\alpha$  radiation on NO concentration given by Mitra (1968). 2 is ionization of NO with Pearce (1968) distribution. 3 is x-ray ion-pair production from Ohshio, et al. (1966) and Hinterreger, et al. (1965). 4 and 5 are rates of ionization of  $\text{O}_2(^1\Delta_g)$  given by Hunten and McElroy (1968).

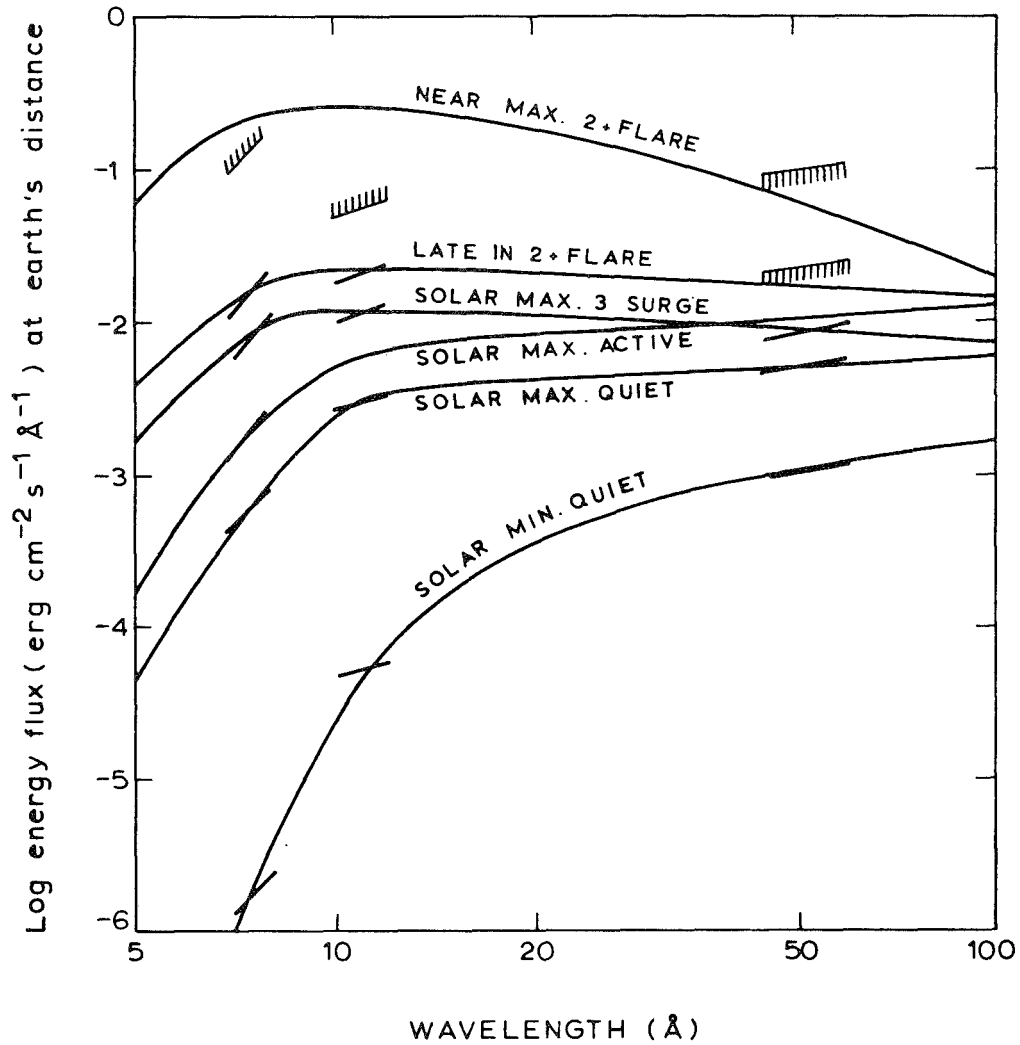


Figure 13. Solar x-ray emission for various solar conditions. The curves drawn are based on measurements in three wavelength bands, as indicated by the heavy bar segments. The slopes of the bar segments are the slopes of the assumed x-ray emission functions used to reduce the photometer responses to the plotted energy fluxes. The energy fluxes refer to values observed at the earth's distance above the atmosphere (from H. Friedman, 1963).

A summary of the absorption and ionization cross sections for X-rays and  $O_2$  and  $N_2$  has been given by Hinterreger et al. (1965) and are reproduced in Table 4, together with their estimates of photon fluxes for quiet conditions of low solar activity. From these data the rate of ionization due to X-rays of wavelength from 1 to  $40\text{\AA}$ , can be computed by using the expression:

$$q_{XR}(h, \chi) = \sum_i \sum_j \sigma_{ij} [X]_j \exp - \sec \chi_j \sum_h \sigma_{ij} \int_h^\infty [X]_j dh \quad (31)$$

where  $i$  represents wavelength,  $j$  constituent, and the other symbols have been defined above. Ohshio et al. (1966) using Hinterreger et al. (1965) cross sections, computed the rate of ionization produced by a unit photon flux for all solar radiation. They included the band from 1 to  $40\text{\AA}$  X-rays. With their data and the fluxes of Table 4, the rate of ionization for undisturbed low solar activity are also shown in Figure 7.

### 3.4 Ionization of Metastable Molecular Oxygen

Recently (Hunten and McElroy, 1968) it has been suggested that a major source of ionization in the upper D region could be the ionization of the metastable species  $O_2(^1\Delta_g)$  by solar ultraviolet radiation on the wavelength band  $1027\text{-}1184\text{\AA}$ . Ionization of the metastable molecular oxygen requires less energy than the non-excited  $O_2$ . Experimental measurements of  $O_2(^1\Delta_g)$  shows that its concentration is large in the mesosphere as is shown in Figure 14, taken from Hunten and McElroy, (1968). The rate of ionization obtained by the same authors is also shown in Figure 12.

### 3.5 Ionization by Precipitated Electrons

It is known that the auroral excitation and ionization in the lower ionosphere is produced mostly by the precipitation of energetic electrons in the atmosphere (Hultqvist, 1965 and references therein). Electrons



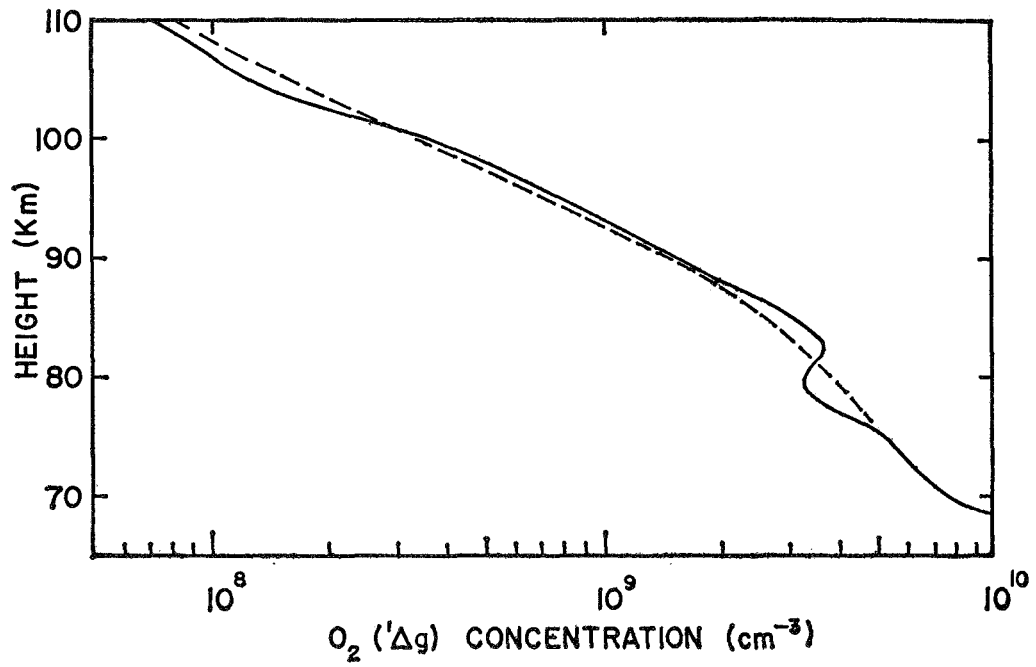


Figure 14. The concentrations of  $O_2(^1\Delta_g)$  above 70 km, from the dayglow measurement of Evans, et al. (1968). The smooth version shown dashed was used for the calculations in this paper.

TABLE 4

Absorption and ionization coefficients for O<sub>2</sub> and N<sub>2</sub> in the X-ray region

Wavelength Region or line (Å)	Photon Energy (eV)	Flux		Effective Cross Sections ( $10^{-18} \text{ cm}^2$ )			
		( $10^9 \frac{\text{ph}}{\text{s}} \text{ cm}^{-2}$ )	( $\text{erg} \text{ cm}^{-2} \text{ s}^{-1}$ )	Total O <sub>2</sub>	N <sub>2</sub>	Ionization O <sub>2</sub>	N <sub>2</sub>
41 - 31	300-400	0.15	0.083	0.2	0.07	0.2	0.07
31 - 22.8	400-540	$5 \times 10^{-3}$	0.005	0.09	1.0	0.09	1.0
22.8 - 15	540-830	$3 \times 10^{-3}$	0.003	0.7	0.36	0.7	0.36
15 - 10	830-1240	$5 \times 10^{-4}$	0.001	0.27	0.15	0.27	0.15
10 - 5	1240-2480	$1.5 \times 10^{-4}$	0.001	0.075	0.045	0.075	0.045
5 - 3	2480-4140	$2 \times 10^{-6}$	$<10^{-5}$	0.012	0.0065	0.012	0.0065
3 - 1	4140-12400	$10^{-8}$	$<10^{-7}$	0.002	0.0015	0.002	0.0015

precipitating in the region of the South Atlantic anomaly (Ghielmetti et al., 1965; Paulikas et al., 1966 and references therein) can be a substantial source of ionization in that area, both at night and in the day (Zmuda, 1966).

In July 1964, O'Brien et al. (1965) measured for the first time simultaneously, both the flux of precipitating electrons ( $E \geq 40$  keV) and nightglow emissions including the  $3914\text{\AA}$  line using rocket techniques. On that occasion, they recorded a flux of  $40 \text{ electron cm}^{-2} \text{ sec}^{-1}$ . Tulinov (1967) and Tulinov et al. (1969) have measured directly the flux of daytime and nighttime energetic particles with a system of Geiger counters aboard rockets. Values of particles fluxes as a function of latitude obtained are shown in Figure 15. The same author computed the rate of ionization that is drawn in Figure 16. The figure also shows the rate of ionization computed by using O'Brien et al. (1965) particle flux and assumed spectrum. This computation has been carried out by using the formulation of Lazarev (1967) which is essentially the same as developed by Rees (1963) but more adequate for a complete computation.

It has been shown (Radicella, 1968) that nighttime ionization in the upper D region can be explained assuming ionization by electron precipitation. His computations were based upon a chemical scheme that will be discussed in the next chapter and on the rate of nighttime ionization given by Ivanov-Kholodny (1965). By using the same formulation given by Radicella (1968) and the rate of ionization obtained from O'Brien et al. (1965) data, the electron density shown in Figure 17 is computed. The figure also shows an experimental profile measured by Mechtly and Smith (1968a). It can be seen that measured electron precipitation can produce at night the necessary ionization in the upper D region.

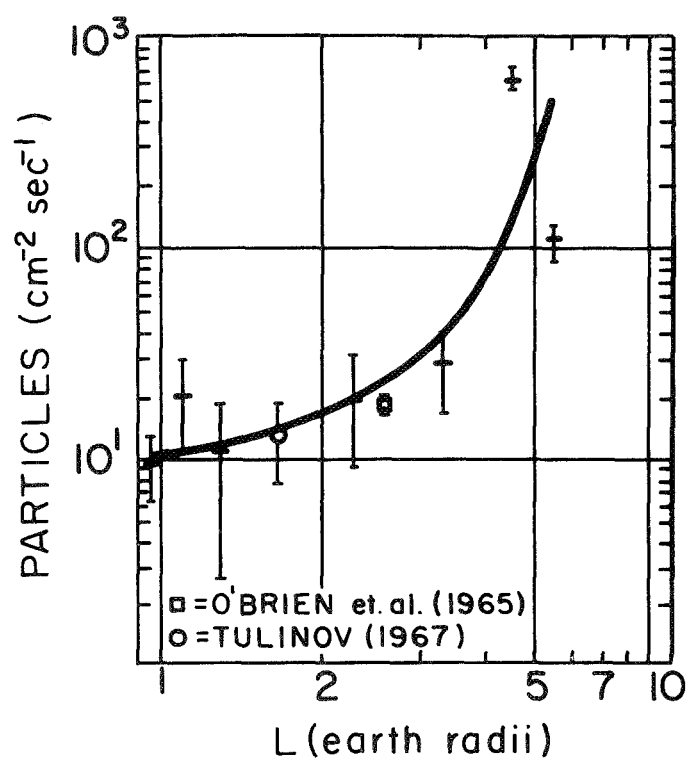


Figure 15. Flux measurements of precipitating energetic particles.

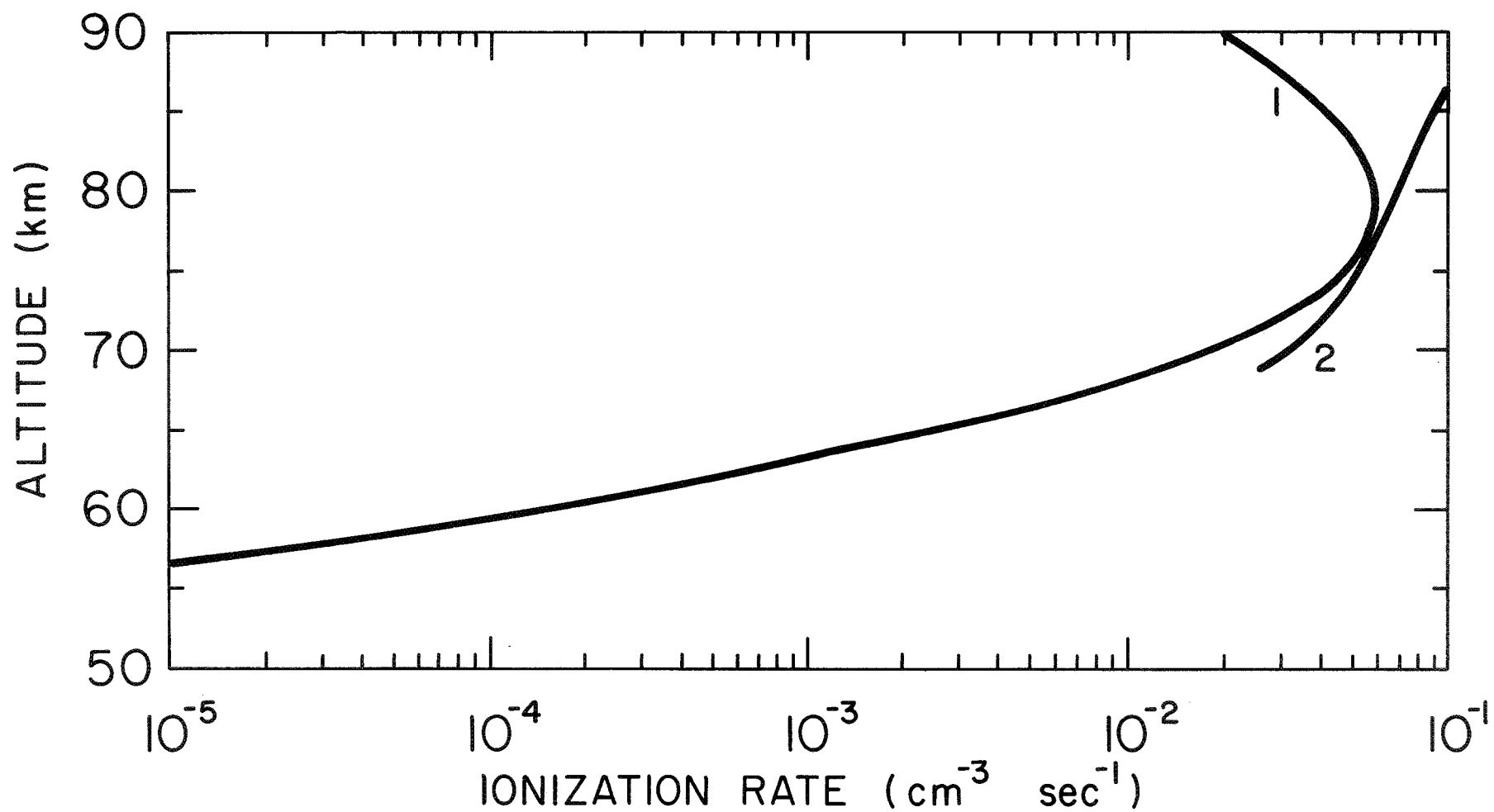


Figure 16. 1 Rate of ionization computed from O'Brien, et al. (1965) measured electron fluxes and estimated spectrum. 2 Tulinov, et al. (1968) rate of ionization from precipitated electrons.

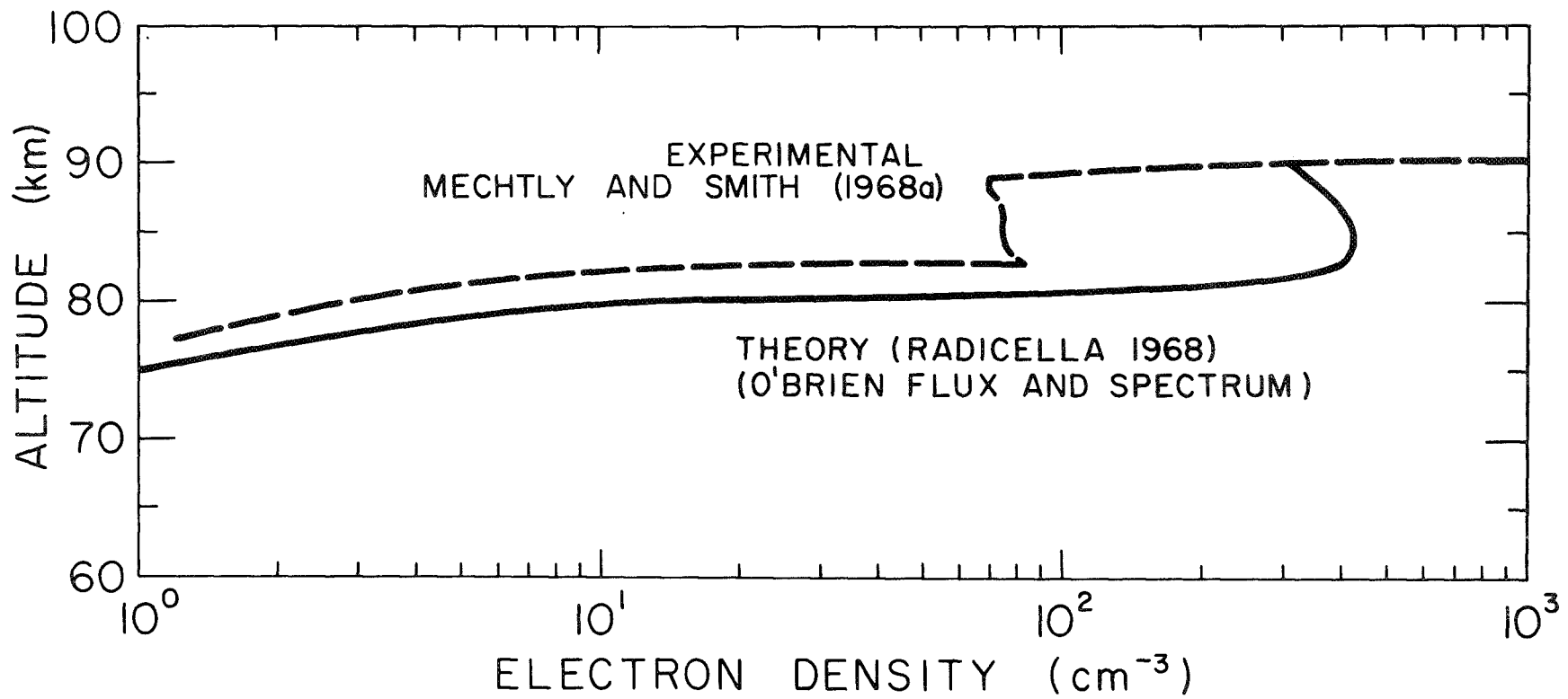


Figure 17. Comparison of experimental and theoretical electron densities in the nighttime D region.

### 3.6 Ionization by Cosmic Rays

Almost forty years ago Hulburt (1930) used Millikan's measurements to compute ionization by cosmic radiation in the region below 60 km. Nicolet and Aikin (1960), in their early theory of the D region, attributed the ionization of the lower part of that region to galactic cosmic rays. Values of the rate of ionization produced by galactic cosmic rays is shown also in Figure 7. The computed values were obtained by using the empirical relation given by Arnold and Pierce (1963):

$$q = 1.5 \times 10^{-18} [M] \cos^{-4} \theta \left[1 - \frac{3}{22} (y - 1964)\right] \quad (32)$$

where

$q$  = ion pair production ( $\text{cm}^{-3} \text{sec}^{-1}$ )

$\theta$  = geomagnetic latitude

$y$  = year

$[M]$  = neutral gas density.

#### 4. NEGATIVE-ION CHEMISTRY IN THE D REGION

##### 4.1 Laboratory Measurements of Negative-Ion Reaction Rates

On the basis of the known laboratory measurements of attachment reactions and mesospheric composition it appears that the basic process of negative-ion formation in the D region is:



with a rate coefficient (Phelps, 1967):

$$k_{33} = 1.4 \times 10^{-29} \left(\frac{300}{T}\right) e^{-\frac{600}{T}} .$$

Other three-body attachment reactions involving minor constituents are unimportant. Reaction:



is very slow (Fehsenfeld and Ferguson, 1968) and as a result,  $O^-$  is not considered as an initial negative ion of importance. It should be considered also that binary attachment reactions involving heavy neutrals like  $NO_3$  could have rate coefficients as high as  $10^{-7} \text{ cm}^{-3} \text{ sec}^{-1}$ .

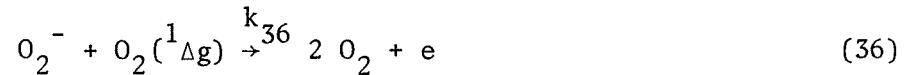
Molecular oxygen negative ions are short lived because electrons can be detached from them by collision with atomic and excited species, and because they can charge exchange with neutral molecular species.

The only abundant atomic species in the D region is atomic oxygen which reacts rapidly with  $O_2^-$ :





where  $k_{34} = 3 \times 10^{-10}$  (Fehsenfeld et al., 1967). Also the excited species  $O_2(^1\Delta g)$  detaches electrons through:



where  $k_{36} = 2 \times 10^{-10}$  (E. E. Ferguson, 1968 private communication).

Recently, several charge exchange reactions have been measured in the laboratory. In particular:



where  $k_{37} = 3 \times 10^{-10}$  (Fehsenfeld et al., 1967) and  $k_{38} = 1 \times 10^{-9}$  (Fehsenfeld and Ferguson, 1968; Rutherford and Turner, 1967).

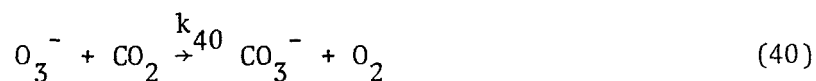
The reaction:



suggested by Nicolet (1965b) has been found to be slow in laboratory measurements, Ferguson (1967). At this time, reaction (35) appears as the basic charge exchange for  $O_2^-$ . Reaction 36 could be important only at night below 70 km where it will eventually provide an effective process in formation of  $NO_2^-$ .

Possible three-body reactions leading to the formation of heavy and stable negative ions should be investigated in the laboratory because of their potential importance at D-region heights. Experimental observations of negative ions spectra Narcisi (1968), for example, suggest the presence of heavy negative ions below 100 km.

Negative ion  $O_3^-$  can exchange its charge with molecular species:



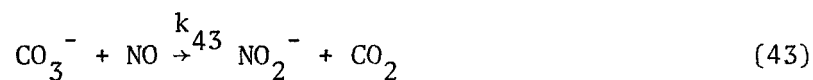
where  $k_{40} = 4 \times 10^{-10}$  and  $k_{41} = 1 \times 10^{-11}$  at room temperature, as found by Fehsenfeld et al., (1967). Reaction (40) is always more effective than reaction (41) because the ratio  $[CO_2]/[NO]$  is much larger than one below 90 km, also if the large values estimated by Barth (1966) and Pearce (1968) are accepted.

$NO_3^-$  is a stable negative ion. This is not the case with  $CO_3^-$  which can react with atomic oxygen regenerating  $O_2^-$  as shown by Fehsenfeld et al., (1967)



where  $k_{42} = 8 \times 10^{-11}$ .

Another possible reaction is:



where  $k_{43} = 9 \times 10^{-12}$  (Fehsenfeld et al., 1967). Reaction (42) is more effective than (43) if:

$$[O] \gg \frac{k_{43}}{k_{42}} [NO] \approx 10^{-1} [NO] .$$

It can be said safely that above 55 km during daytime and above 70 km at night, reaction (42) is the main charge exchange process of  $CO_3^-$ . However,

the formation of  $O_2^-$  by (42) and the rapid detachment that takes place through reactions (35) and (36), tend to produce an excess of electrons below 90 km during both daytime and nighttime, contrary to the experimental evidence. To avoid this difficulty in studying nighttime D region, Radicella (1968) postulated another fast charge exchange reaction dealing with a stable negative ion, suggested originally by Whitten and Poppoff (1964):



assuming  $k_{44} = 10^{-13} \pm 1$ . However, F. C. Fehsenfeld (private communication 1968) studied this reaction in the laboratory and found it to be very slow ( $k_{44} < 10^{-15} \text{ cm}^{-3} \text{ sec}^{-1}$ ) at room temperature.

It must be noticed, however, that the final implicit equation used by Radicella (1968) in order to compute nighttime electron-density profiles for steady-state conditions (see his Appendix 2, Equation A12) is independent of the particular reaction that produces the stable negative ion. That equation is formally identical with equation below.

Final processes that involve electrons and stable negative ions in the D region are recombination and mutual neutralization. The rate coefficient for dissociative recombination:



has been analyzed recently by Biondi (1967).

A value of  $k_{45} = 7 \times 10^{-7}$  at 200°K (D-region temperature) with a temperature dependence of the type  $T^{-1}$  appears to be reasonable, when  $XY^+ \equiv NO^+$ .

The present experimental knowledge about the rate of mutual neutralization:



is not very good. However, a rate coefficient  $k_{46} = 4 \times 10^{-8}$  for atmospheric ions has been measured by Hirsch et al. (1967) and this value agrees with the theoretical estimate of  $k_{46} = 5 \times 10^{-8}$  by Baulknight and Bortner (1964).

A summary of the rate coefficients of negative-ion reactions is given in Table 5. In the above discussion, we confined our attention on a rather reduced number of reactions involving the major constituents and the minor constituents of larger concentration. A simple inspection of the rate coefficients for the many negative-ion reactions involving the known minor atmospheric constituents including hydrogenous compounds. (DASA Reaction Rate Handbook: 1967) shows that they are of very little importance in the negative-ion chemistry of the normal D region.

#### 4.2 General Schemes of D-Region Negative-Ion Chemistry

Let reaction (33) be the initial step in the formation of negative ions. Time variations of the ionic species at a given altitude can be written:

$$\frac{d[e]}{dt} = \sum q + \delta [O_2^-] [P_1] - \beta [e] [M]^2 - \alpha_e [e] [XY^+] \quad (47)$$

$$\frac{d[O_2^-]}{dt} = \beta [e] [M]^2 - \delta [O_2^-] [P_1] - \gamma [O_2^-] [P_2] \quad (48)$$

$$\frac{d[M^-]}{dt} = \gamma [O_2^-] [P_2] - \alpha_i [XY^+] [M^-] \quad (49)$$

where brackets indicate concentrations,  $e$  indicates electron,  $P_1, P_2, \dots$  are neutral molecules or atoms,  $M$  is the atmospheric gas as a whole,  $N^-$  is the final molecular stable negative ion,  $XY^+$  is the general molecular positive ion, and  $\sum q$  is the ion pair production rate due to the sum of all possible sources of ionization.  $\alpha_e, \alpha_i, \beta, \gamma, \delta$  are, respectively, rate coefficients of recombination, mutual neutralization, attachment, charge exchange and detachment. It can be assumed also that:

$$[XY^+] \approx [e] + [N^-] \quad (50)$$

because the stable negative ion will be always more abundant than other intermediate negative ions like  $O_2^-$ . When:

$$\frac{d[e]}{dt} = \frac{d[O_2^-]}{dt} = \frac{d[M^-]}{dt} = 0 \quad (51)$$

is assumed, i.e. when steady-state conditions are considered, solutions of Equations (47), (48) and (49) will not be changed if intermediate charge exchange reactions take place. The steady-state solution of Equation (48) gives:

$$[O_2^-] = \frac{\beta[M]^2[e]}{\gamma[P_1] + \delta[P_2]} \quad (52)$$

By substitution in 47 the following equation is obtained:

$$\sum q - \frac{\gamma[P_2]}{\delta[P_1] + \gamma[P_2]} \beta[M]^2[e] - \alpha_e[XY^+][e] = 0 \quad (53)$$

or:

$$\sum q - \beta_{\text{eff}}[e] - \alpha_e[XY^+][e] = 0 \quad (54)$$

where:

$$\beta_{\text{eff}} = \frac{\gamma[P_2]}{\delta[P_1] + \gamma[P_2]} \beta[M]^2 \quad (55)$$

can be called an "effective attachment factor". Taking into account Equation (50), it can be written:

$$\sum q - (\beta_{\text{eff}} + \alpha_e [N^-])[e] - \alpha_e [e]^2 = 0 \quad (56)$$

$$\beta_{\text{eff}}[e] - \alpha_i [e][N^-] - \alpha_i [N^-]^2 = 0 \quad (57)$$

This system of equations can be solved in terms of  $e$  and  $N^-$  if the ion-pair function and coefficients are known. An implicit equation in  $\beta_{\text{eff}}$  is written as follows:

$$[e] = \frac{\sum q}{\alpha_e \left\{ \frac{[e]}{2} + \sqrt{\frac{[e]^2}{4} + \beta_{\text{eff}} \frac{[e]}{\alpha_i}} \right\} + \beta_{\text{eff}}} \quad (58)$$

which is formally equal to Equation A12 of Radicella (1968). If the electron density is known an implicit equation in  $\beta_{\text{eff}}$  is:

$$\beta_{\text{eff}} = \frac{\sum q}{[e]} - \alpha_e \left\{ \frac{[e]}{2} + \sqrt{\frac{[e]^2}{4} + \beta_{\text{eff}} \frac{[e]}{\alpha_i}} \right\} \quad (59)$$

Equations (47) to (49) have been written considering only the detachment process with a rate  $\delta[P_1][O_2^-]$ . If an additional process with a rate  $\delta'[P_1'] [O_2^1]$  is taken into account, reactions (56) and (57) are formally maintained, but now is:

$$\beta_{\text{ef}} = \frac{\gamma[P_2]}{\delta[P_1] + \delta'[P_1'] + \gamma[P_2]} \beta[M]^2 \quad (60)$$

If the charge exchange with rate  $\gamma[P_2][O_2^-]$  is such that:

$$\gamma [P_2] \gg \delta [P_1]$$

Equation (23) reduces to:

$$\beta_{\text{eff}} = \beta[M]^2 \quad (61)$$

Another possibility appears if the charge exchange reaction or chain of reactions is not capable of producing a large concentration of stable negative ions. The effect will be:

$$\beta_{\text{eff}} \ll \beta [M]^2 \quad (62)$$

that is equivalent, in terms of electron density, to reducing Equation (26) to:

$$[e] = \sqrt{\frac{\Sigma q}{\alpha_e}} \quad (63)$$

A final case results if an additional attachment process is postulated with a rate  $\beta^1[X]$ , where X is an unknown constituent, and such that Equation (23) becomes:

$$\beta_{\text{eff}} = \beta^1[X] + \frac{\gamma[P_2]}{\delta[P_1] + \gamma[P_2]} \beta[M]^2 \approx \beta^1[X] \quad (64)$$

From the above discussion it is clear that each of the different forms of  $\beta_{\text{eff}}$  reflects a different negative-ion scheme. It is possible now to analyze

the problem of the apparent contradictions between daytime-measured electron densities and NO concentration. If measured NO concentration (Pearce, 1968) is accepted together with reactions (33), (34), (35), (36), (37), (40), (41), (42), (43), (45), and (46), which represent the most accepted set, and which will refer as the "basic schemes", the effective attachment factor will be written as:

$$\beta_{\text{eff}} = 1 \frac{\frac{k_{43}[\text{NO}]}{k_{43}[\text{NO}] + k_{42}[\text{O}]}}{[\text{O}] + \frac{2}{3} [\text{O}_2(^1\Delta\text{g})] + \frac{k_{43}[\text{NO}]}{k_{43}[\text{NO}] + k_{42}[\text{O}]} [\text{O}_3]} k_{33}[\text{O}_2]^2 \quad (65)$$

As it is mentioned before, during daytime, above 55-60 km, and at night above 70-75 km:

$$k_{42}[\text{O}] \gg k_{43}[\text{NO}] \quad ,$$

Then:

$$\beta_{\text{eff}} \approx \frac{\frac{k_{43}[\text{NO}]}{k_{42}[\text{O}]} [\text{O}_3]}{[\text{O}] + \frac{2}{3} [\text{O}_2(^1\Delta\text{g})] + \frac{k_{43}[\text{NO}]}{k_{42}[\text{O}]} [\text{O}_3]} k_{33}[\text{O}_2]^2 \quad (66)$$

where:

$$\frac{\frac{k_{43}[\text{NO}]}{k_{42}[\text{O}]} [\text{O}_3]}{[\text{O}] + \frac{2}{3} [\text{O}_2(^1\Delta\text{g})] + \frac{k_{43}[\text{NO}]}{k_{42}[\text{O}]} [\text{O}_3]} \ll 1 \quad . \quad (67)$$

This case corresponds then to the condition given by Equation (63). The electron distribution can be computed by using Equations (58) and (66) and the ion-pair production rate given in Chapter 2 when Pearce (1968) NO



concentration is adopted. The result is shown in Figure 18 together with experimental data and it can be seen that below 85 - 90 km the electron density is much larger than expected on the basis of experimental evidence (Mechtly and Smith, 1968b).

By using the basic sequence plus an additional fast charge exchange of the type given by Radicella (1968):



where:

$$k_{68}[X] \gg k_{46}[CO_2] \quad .$$

The effective attachment factor will be:

$$\beta_{\text{eff}} = \frac{[O_3]}{[O] + [O_3] + \frac{2}{3} [O_2(^1\Delta g)]} k_{33} [O_2]^2 \quad (69)$$

which is equivalent to Equation (60). Again by using (58) and (69) the electron density is computed and it is shown in Figure 19. Once more, this theoretical computation shows a larger electron density than what expected experimentally. This is also the case if a fast charge exchange reaction of the type:



such that:

$$k_{70}[X] \gg k_{35} \left( [O] + \frac{2}{3} [O_2(^1\Delta g)] \right)$$

is accepted. The effective attachment factor reduces then to:

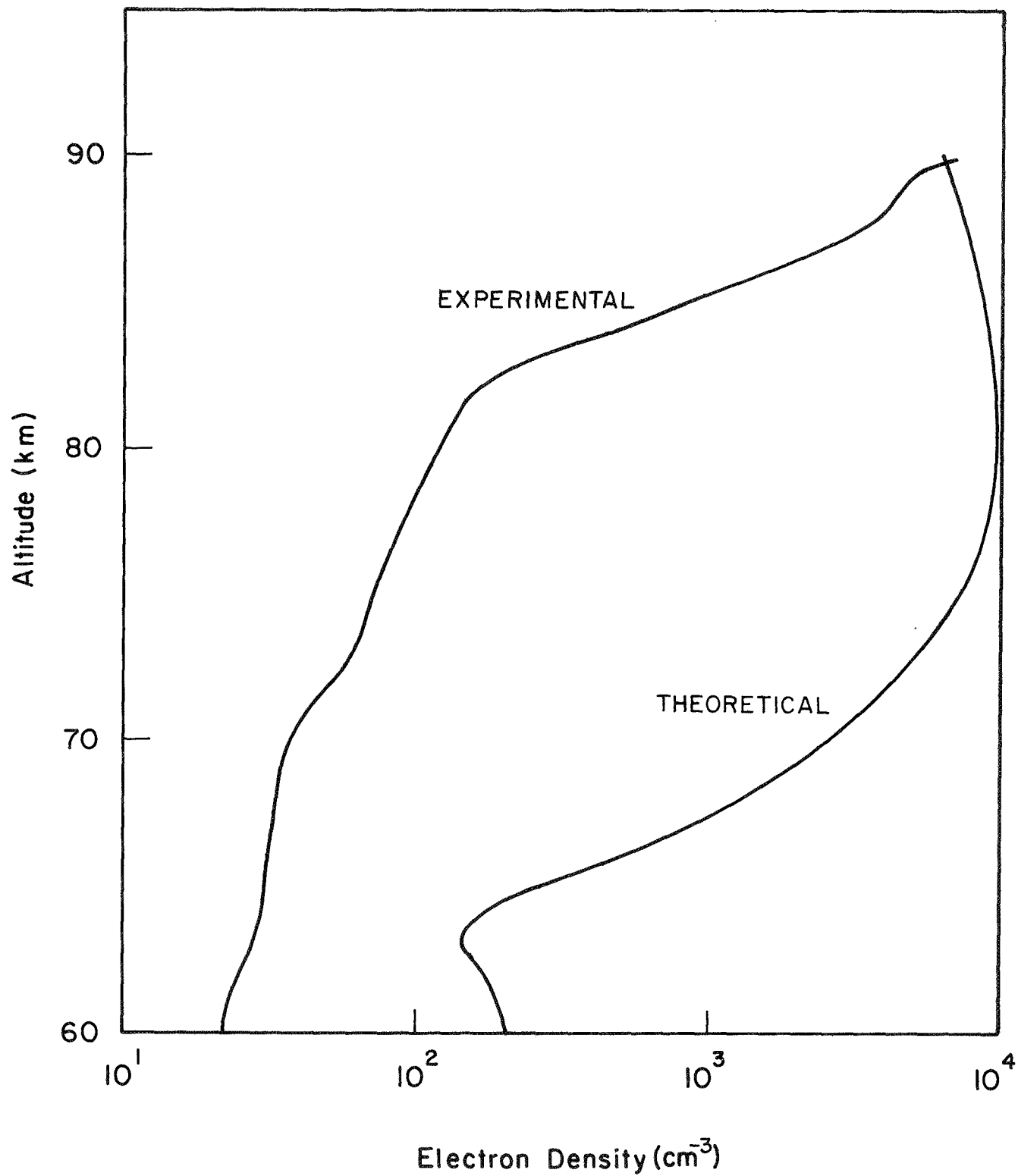


Figure 18. Comparison of daytime electron densities. Experimental by Mechtly and Smith (1968b); theoretical from Equations (58) and (66) and NO concentration by Pearce (1968).

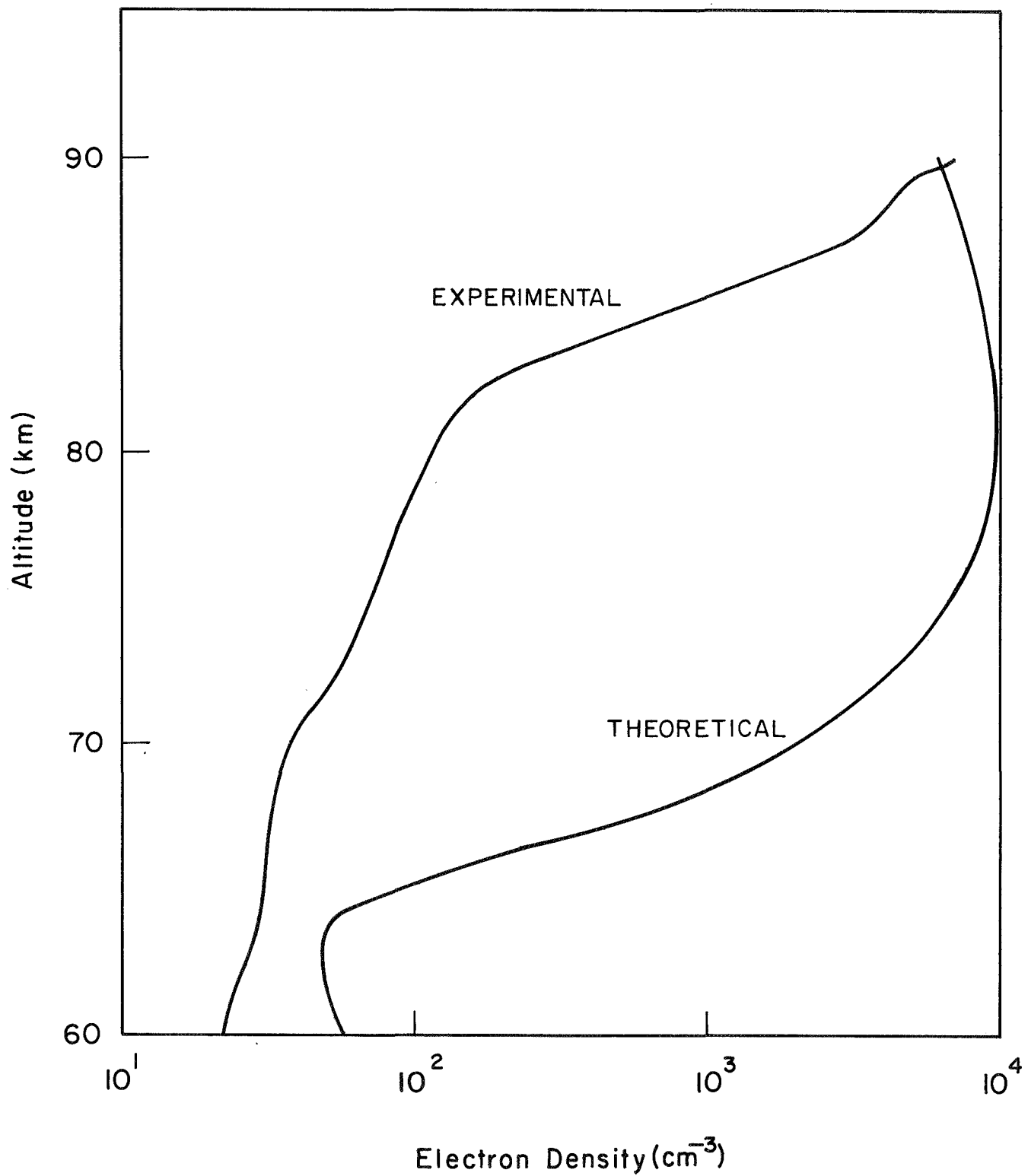


Figure 19. Comparison of daytime electron densities. Experimental by Mechtly and Smith (1968b); theoretical from Equations (58) and (69) and NO concentration by Pearce (1968).

$$\beta_{\text{eff}} = k_{33}[\text{O}_2]^2 \quad (71)$$

which is formally equal to Equation (61). The computed electron density for this case is shown in Figure 20 and again no agreement is reached with the experimental data.

The only possibility left in order to obtain an agreement between the experimental electron density and the measured NO concentration is to search for new attachment reactions faster than the attachment to  $\text{O}_2$  between 70 and 90 km. This is done by means of  $\beta_{\text{eff}}$ . It can be computed by using Equation (59), daytime experimental values of  $[\text{NO}]$  and  $[\text{e}]$ , and accepted deionization coefficients. The effective attachment factor can be also derived theoretically from (69) and (71). Figure 21 shows these computed values of  $\beta_{\text{eff}}$ . It can be seen that the profile based upon experimental data, is quite different from the two theoretical profiles. The only way to conciliate experimental values of  $[\text{NO}]$  and  $[\text{e}]$  and theoretical computation is to postulate a binary attachment reaction of the type:



where  $\text{Z}^-$  is a stable negative ion and such that:

$$k_{72}[\text{ZW}] \gg k_{33}[\text{O}_2]^2$$

or:

$$\beta_{\text{eff}} = k_{72}[\text{ZW}]$$

which corresponds to Equation (64).

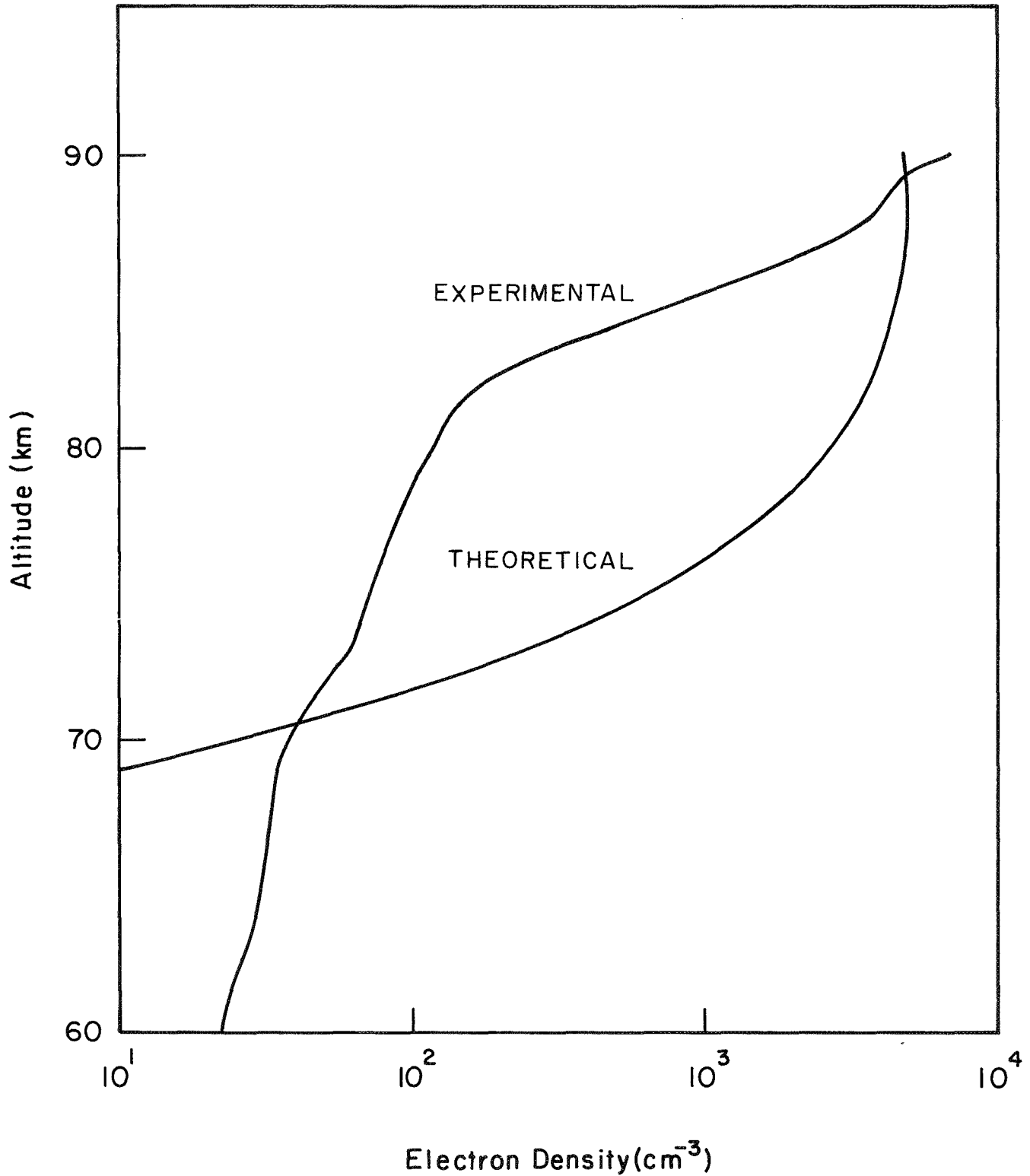


Figure 20. Comparison of daytime electron densities. Experimental by Mechtly and Smith (1968b); theoretical from Equations (58) and (71) and NO concentration by Pearce (1968).

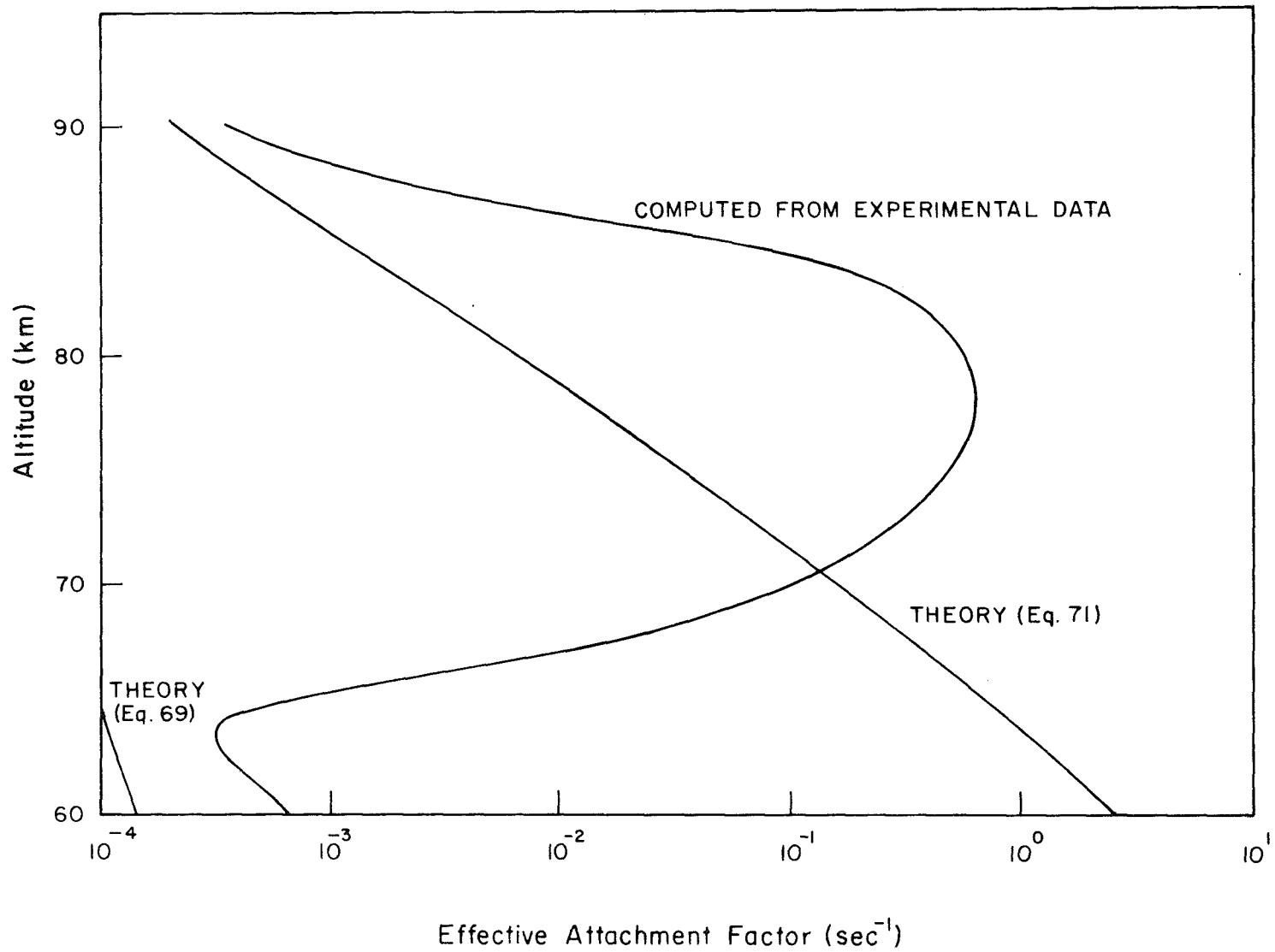


Figure 21. Comparison of effective attachment values. Computed from experimental data of Mechtly and Smith (1968b) and Pearce (1968) and from Equations (69) and (71).

Accepting reaction (72) where [ZW] could be a heavy neutral, and a tentative rate coefficient  $k_{72} \sim 10^{-7} \text{ cm}^{-3} \text{ sec}^{-1}$ ,  $\beta_{\text{eff}}$  experimental in Figure 21 will indicate that the specie [ZW] would be a minor constituent with a peak concentration of  $\sim 5 \times 10^6$  at around 80 km. Such a minor contamination at that height could be accepted, but its eventual nature is at this time unpredictable. Possible conglomerates of neutral particle in the region of low temperature could form heavy neutrals.

In summary it can be said, that the negative-ion chemistry can be represented by an effective attachment coefficient  $\beta_{\text{eff}}$  as shown by Equation (54) which takes different forms depending upon the negative-ion processes involved. To conciliate the high values of [NO] given by Pearce (1968) with experimental values of electron density and recombination, and mutual neutralization rates, a new attachment process must be postulated. This should be able to form a stable negative ion and could be a fast binary reaction involving heavy neutral molecular compounds or conglomerates formed under the low temperature conditions of the D region. However, further complications appear in the lower mesosphere where the experimental values of  $\beta_{\text{eff}}$  suggest the existence of fast detachment reactions.

## APPENDIX I

## 1. SOURCES OF DATA

1.1 Introduction

The computations for the oxygen atmosphere were based on Equations (7) and (8) of the text. The values of  $k_3$ ,  $k_4$  and  $k_5$  were taken from the Kaufman adopted data in Table 1.

The  $O_2$  and  $O_3$  absorption cross sections and the solar photon flux are tabulated in Table A1. The same data is shown graphically in Figures 3 and 4 of the text. A list of the references from which the cross sections and fluxes were taken is given in Table A2. The solar flux is the unattenuated photon flux outside the earth's atmosphere. The data numbers in Table A2 refer to those in Table A1. Below 2425Å the  $O_2$  absorption cross sections are very small according to Hunt (1963) SAD 123, so corrections below 2425 Å are assumed to be zero for these calculations. Between 1775 Å and 1975 Å the solar flux was read from a graph in the Handbook of Physics at 5 Å intervals. These values were then averaged over 50 Å intervals to give an average flux for each 50 Å interval.

1.2 Neutral Atmosphere Parameters  $[O_2]$ ,  $[M]$ 

The attenuation of solar radiation as it passes through the atmosphere depends on the number of absorbing molecules between the altitude of interest and the outside of the earth's atmosphere. For these calculations  $O_2$  was assumed to be the only species which significantly attenuated the solar flux. From 300 to 120 km, the values of  $[O_2]$  were obtained from the U.S. Standard Atmosphere Supplement 1966 for the 3 cases: Winter Model, Exospheric Temperature = 1000°K, Summer Model, Exospheric Temperature = 1000°K, and Spring/Fall Model, Exospheric Temperature = 1000°K.

Below 120 km it was necessary to convert the absolute atmospheric density



Table A1

DATUM NUMBER	CENTRAL WAVELENGTH (Å)	WAVELENGTH INTERVAL (Å)	PHOTON FLUX		O2 ABSORB. CROSS-X (cm <sup>-2</sup> )	O3 ABSORB. CROSS-X (cm <sup>-2</sup> )
			PER INTERVAL (phot. cm <sup>-2</sup> sec <sup>-1</sup> Å <sup>-1</sup> )	PER ANGSTROM (phot. cm <sup>-2</sup> sec <sup>-1</sup> Å <sup>-1</sup> )		
1	1215.	1.0	2.700D 11	2.700E 11	1.100D-20	3.500D-17
2	1400.	50.0	1.830D 10	3.660E 08	1.450D-17	7.000D-18
3	1450.	50.0	3.700D 10	7.400E 08	1.480D-17	4.500D-18
4	1500.	50.0	7.300D 10	1.460E 09	1.280D-17	3.500D-18
5	1550.	50.0	1.330D 11	2.660E 09	8.800D-18	2.500D-18
6	1600.	50.0	2.600D 11	5.200E 09	4.870D-18	1.100D-18
7	1650.	50.0	4.200D 11	8.400E 09	2.370D-18	9.000D-19
8	1700.	50.0	7.000D 11	1.400E 10	1.020D-18	9.000D-19
9	1750.	50.0	1.060D 12	2.120E 10	3.080D-19	8.500D-19
10	1775.	2.5	6.088D 10	2.435E 10	2.530D-20	8.000D-19
11	1780.	5.0	1.275D 11	2.550E 10	1.700D-20	7.900D-19
12	1785.	5.0	1.328D 11	2.656E 10	1.630D-20	7.900D-19
13	1790.	5.0	1.394D 11	2.788E 10	1.770D-20	7.800D-19
14	1795.	5.0	1.407D 11	2.814E 10	2.080D-20	7.800D-19
15	1800.	5.0	1.515D 11	3.030E 10	1.200D-20	7.700D-19
16	1805.	5.0	1.587D 11	3.174E 10	2.670D-20	7.700D-19
17	1810.	5.0	1.660D 11	3.320E 10	1.170D-20	7.600D-19
18	1815.	5.0	1.733D 11	3.466E 10	1.190D-20	7.500D-19
19	1820.	5.0	1.806D 11	3.612E 10	1.180D-20	7.400D-19
20	1825.	5.0	1.880D 11	3.760E 10	4.500D-21	7.300D-19
21	1830.	5.0	1.977D 11	3.954E 10	7.040D-21	7.200D-19
22	1835.	5.0	2.061D 11	4.122E 10	6.340D-21	7.100D-19
23	1840.	5.0	2.150D 11	4.300E 10	1.430D-21	7.000D-19
24	1845.	5.0	2.248D 11	4.496E 10	5.630D-21	6.900D-19
25	1850.	5.0	2.347D 11	4.694E 10	6.640D-21	6.800D-19
26	1855.	5.0	2.447D 11	4.894E 10	1.480D-21	6.700D-19
27	1860.	5.0	2.570D 11	5.140E 10	1.570D-21	6.550D-19
28	1865.	5.0	2.671D 11	5.342E 10	5.070D-21	6.400D-19
29	1870.	5.0	2.819D 11	5.638E 10	2.030D-21	6.250D-19
30	1875.	5.0	2.971D 11	5.842E 10	4.400D-22	6.100D-19
31	1880.	5.0	3.070D 11	6.140E 10	1.450D-21	5.950D-19
32	1885.	5.0	3.197D 11	6.394E 10	1.770D-21	5.800D-19
33	1890.	5.0	3.348D 11	6.696E 10	3.560D-22	5.650D-19
34	1895.	5.0	3.500D 11	7.000E 10	2.060D-22	5.500D-19
35	1900.	5.0	3.628D 11	7.256E 10	6.370D-22	5.300D-19
36	1905.	5.0	3.781D 11	7.562E 10	8.280D-22	5.200D-19
37	1910.	5.0	3.935D 11	7.870E 10	3.980D-22	5.100D-19
38	1915.	5.0	4.138D 11	8.276E 10	1.880D-22	5.000D-19
39	1920.	5.0	4.293D 11	8.586E 10	4.300D-22	4.800D-19
40	1925.	5.0	4.450D 11	8.903E 10	6.450D-22	4.700D-19
41	1930.	5.0	4.655D 11	9.310E 10	2.280D-22	4.600D-19
42	1935.	5.0	4.862D 11	9.724E 10	1.200D-22	4.500D-19
43	1940.	5.0	4.972D 11	9.944E 10	9.440D-23	4.400D-19
44	1945.	5.0	5.278D 11	1.056E 11	1.800D-22	4.300D-19
45	1950.	5.0	5.438D 11	1.088E 11	1.980D-22	4.200D-19
46	1955.	5.0	5.649D 11	1.130E 11	8.870D-23	4.100D-19
47	1960.	5.0	5.860D 11	1.172E 11	5.850D-23	3.900D-19
48	1965.	5.0	5.974D 11	1.195E 11	5.970D-23	3.750D-19
49	1970.	5.0	6.286D 11	1.257E 11	1.030D-22	3.600D-19
50	1975.	2.5	3.250D 11	1.300E 11	8.850D-23	3.450D-19
51	2000.	50.0	7.035D 12	1.407E 11	1.300D-23	3.300D-19
52	2050.	50.0	9.271D 12	1.854E 11	1.100D-23	3.520D-19
53	2100.	50.0	1.530D 13	3.060E 11	9.600D-24	5.490D-19
54	2150.	50.0	2.593D 13	5.186E 11	7.900D-24	9.860D-19
55	2200.	50.0	3.427D 13	6.854E 11	6.400D-24	1.750D-18
56	2250.	50.0	3.957D 13	7.914E 11	4.600D-24	2.890D-18
57	2300.	50.0	4.161D 13	8.322E 11	3.100D-24	4.450D-18
58	2350.	50.0	3.779D 13	7.558E 11	1.800D-24	6.250D-18
59	2400.	50.0	4.100D 13	8.200E 11	1.000D-24	8.050D-18
60	2450.	50.0	4.802D 13	9.604E 11	0.0	9.800D-18
61	2500.	50.0	4.774D 13	9.548E 11	0.0	1.090D-17
62	2550.	50.0	7.176D 13	1.435E 12	0.0	1.130D-17
63	2600.	50.0	9.146D 13	1.829E 12	0.0	1.080D-17
64	2650.	50.0	1.332D 14	2.664E 12	0.0	9.560D-18
65	2700.	50.0	1.696D 14	3.392E 12	0.0	7.770D-18
66	2750.	50.0	1.520D 14	3.040E 12	0.0	5.700D-18
67	2800.	50.0	1.688D 14	3.376E 12	0.0	3.840D-18
68	2850.	50.0	2.435D 14	4.870E 12	0.0	2.440D-18
69	2900.	50.0	3.789D 14	7.578E 12	0.0	1.300D-18
70	2950.	50.0	4.670D 14	9.340E 12	0.0	7.440D-18
71	3000.	50.0	4.598D 14	9.196E 12	0.0	3.840D-19
72	3050.	50.0	5.134D 14	1.027E 13	0.0	1.960D-18
73	3100.	50.0	5.920D 14	1.184E 13	0.0	1.020D-18
74	3150.	50.0	6.490D 14	1.298E 13	0.0	5.360D-20
75	3200.	50.0	6.834D 14	1.367E 13	0.0	7.940D-20
76	3250.	50.0	8.320D 14	1.666E 13	0.0	1.460D-20
77	3300.	50.0	9.535D 14	1.907E 13	0.0	7.040D-21
78	3350.	50.0	9.343D 14	1.869E 13	0.0	3.820D-21
79	3400.	50.0	9.482D 14	1.896E 13	0.0	1.940D-21
80	3450.	50.0	1.014D 15	2.028E 13	0.0	7.040D-22
81	3500.	50.0	1.038D 15	2.076E 13	0.0	2.740D-22
82	4750.	500.0	2.506D 16	5.012E 13	0.0	4.800D-22
83	5250.	500.0	2.575D 16	5.150E 13	0.0	2.060D-21
84	5750.	500.0	2.757D 16	5.514E 13	0.0	3.980D-21
85	6250.	500.0	2.698D 16	5.396E 13	0.0	3.660D-21
86	6750.	500.0	2.602D 16	5.204E 13	0.0	1.530D-21
87	7250.	500.0	2.474D 16	4.948E 13	0.0	5.230D-22

TABLE A2

## Ø2 Absorption Cross Sections

Data Numbers	Wavelength	Source
1-50	1215Å, 1375Å-1975Å	A
51-87	1975Å-3525Å, 4500Å-7500Å	B

## Ø3 Absorption Cross Sections

Data Numbers	Wavelength	Source
1-10	1215Å, 1375Å-1775Å	C
15,25,35,45	1800Å, 1850Å, 1900Å, 1950Å	D
51-87	1975Å-3525Å, 4500Å-7500Å	D

Between 1775Å and 1975Å, all cross sections not listed in this table are smooth curve interpolations of Hunt's data.

## Solar Flux

Data Numbers	Wavelength	Source
1-50	1215Å, 1375Å-1975Å	E
51-81	1975Å-3525Å	F
82-87	4500Å-7500Å	G

## TABLE A2 SUPPLEMENT

- A. Huffman, Robert E., DASA Reaction Rate Handbook, October 1967, Chapter 10, PHOTOCHEMICAL PROCESSES, page 10-7, Values read from graph, General Electrical Company Missile and Space Division, Space Sciences Lab., Philadelphia, Penn.
- B. Hunt, B. G., Technical Note SAD 123, Weapons Research Establishment, June 1963, Adelaide, South Australia, page 22.
- C. Craig, Richard A., The Upper Atmosphere, Meteorology and Physics, 1965, Academy Press, N.Y., Ed. J. Van Mieghem, page 168 read from graph, Chapter 4, The Sun's Radiation and the Upper Atmosphere.
- D. Hunt, SAD 123, page 24.
- E. Handbook of Geophysics, Data Numbers 1-9 from page 16-16, Data numbers 9-50 read from graph page 16-11.
- F. Tousey, Richard, Space Science Reviews, Vol. 2, 1963, The Extreme Ultraviolet Spectrum of the Sun, page 22.
- G. Handbook Geophysics, page 165.

given in the U. S. Standard Atmosphere Supplement 1966 to  $[O_2]$  and  $[M]$  by using the mean molecular weight and the percent composition of  $O_2$ . Below 90 km the densities were converted to  $[O_2]$  and  $[M]$  by assuming a constant mean molecular weight of 28.96 (CIRA 1965) and constant ratio of  $[O_2]$  to  $[M]$  of .2091 for all seasons. This ratio was calculated from the 90 km values of  $[N_2]$ ,  $[O_2]$ ,  $[O]$ , and  $[A]$  given in CIRA 1965 page 12. At 120 km the mean molecular weight and percent composition of  $O_2$  to  $M$  for each season was calculated from data in the U. S. Standard Atmosphere Supplement 1966. Between 90 and 120 km the mean molecular weight and the percent composition were assumed to vary linearly with altitude so as to match the boundary values of 90 and 120 km. The boundary values of composition and mean molecular mass at 90 km and 120 km are given in Table A3. The resultant values of  $[O_2]$  were then used to calculate the integral  $\int_z^{300 \text{ km}} [O_2] dh$  which is used in the calculation of the photodissociation rates. Above 300 km  $[O_2]$  contributed negligibly to the integral.

The photodissociation rate  $J_k$  of the  $k$ th constituent was calculated according to the formula:

$$J_k = \sum_{\lambda} \sigma_{k\lambda} R_{\lambda} \exp(-\sec\chi \sigma_{\lambda}(O_2) \int_z^{300 \text{ km}} [O_2] dh)$$

where:  $\sigma_{k\lambda}$  = the average value of the photodissociation cross section of the  $k$ th constituent in the  $\lambda$ th wavelength interval

$R_{\lambda}$  = the total solar photon flux with wavelength in the  $\lambda$ th interval

$\sec\chi$  = the secant of the solar zenith angle  $\chi$

$\sigma_{\lambda}(O_2)$  = the average absorption cross section of  $O_2$  in the  $\lambda$ th interval

$\int_z^{300 \text{ km}} [O_2] dh$  = the integral of the  $O_2$  concentration from 300 km to  $Z$ , the altitude of interest.

TABLE A3

	90 km	Mean Molecular Mass	Composition
Summer		28.96	20.91%
Winter		28.96	20.91%
Spring/Fall		28.96	20.91%

	120 km	Mean Molecular Mass	Composition
Summer		26.76	13.17%
Winter		27.12	14.40%
Spring/Fall		26.90	13.61%

In the calculation of the photodissociation rates, it was assumed that only  $O_2$  significantly attenuated solar radiation. It was also assumed that the quantum efficiency for  $O_2$  and  $O_3$  was unity; i.e. every photon absorbed by the molecule resulted in the dissociation of that molecule. Thus, the absorption cross sections are the same as the photodissociation cross sections.

### 1.3 Initial Values for the Oxygen Atmosphere

Initial values of  $O$  and  $O_3$  were calculated for latitudes ranging from  $45^\circ N$  to  $45^\circ S$  in  $15^\circ$  intervals with a constant solar declination of  $23.5^\circ N$ . The appropriate temperature profiles for the rate coefficients, as well as  $[O_2]$ , and the  $[O_2]$  integrals, were obtained from the data in U.S. Standard Atmosphere Supplement 1966. The solar zenith angle used in computing the photodissociation rates was calculated from the formula:

$$\cos \chi = \sin(\text{latitude}) \sin(\text{declination}) + \cos(\text{latitude}) \cos(\text{declination})$$

The steady state concentrations of  $O$  and  $O_3$  were approximated by:

$$[O] = \left[ \frac{J_2 [O_2]}{[M] (k_3 + k_4 k_5 [O_2])} \right]^{1/2}$$

$$[O_3] = \frac{k_4 [O] [O_2] [M]}{J_3}$$

$k_3, k_4, k_5$  are the Kaufman reaction rates given in Table 1 of the text.

$J_2, J_3$  = photodissociation rates of  $O_2$  and  $O_3$  respectively.

$[O], [O_2], [M], [O_3]$  are the concentrations of  $O, O_2, M,$  and  $O_3$ , respectively.

### 1.4 Time-Dependent Oxygen Atmosphere

The time-dependent oxygen atmosphere used the values of  $[O]$  and  $[O_2]$  from the steady-state approximation as initial values. The basic data for reaction rates, temperatures, cross-sections, solar fluxes, and  $[O_2]$  profiles were identical in both models. The time dependence was incorporated into the

solar zenith angle  $\chi$  in the form  $\cos\chi = \sin(\text{latitude}) \sin(\text{declination}) + \cos(\text{latitude}) \cos(\text{declination}) \cos(\text{hour angle})$  where the hour angle uniformly increased at the rate of 24 hours per day. An approximation to the Chapman function was used for the solar zenith angle  $\chi$  between  $80^\circ$  and  $90^\circ$  to lessen errors due to the curvature of the atmosphere. The values of the Chapman function  $Ch(\chi, x)$  were taken from M.V. Wiles. (In this notation  $\chi$  = solar zenith angle, and  $x = \frac{h+R_e}{H}$  where  $h$  is the altitude of observation above the earth's surface,  $R_e$  is the radius of the earth, and  $H$  is the scale height of  $O_2$  at that altitude.) At the altitudes and densities involved, the resultant values of  $x$  were quite large, of the order of 1000. To approximate the Chapman function, the region between  $80^\circ$  and  $90^\circ$  was broken into  $1^\circ$  intervals. Within each interval the Chapman function was approximated quite well as a linear function of  $x$  about the value at  $x = 1000$ . When  $\chi \geq 90^\circ$  all dissociation rates were set to zero because the solar radiation must pass through very dense layers of  $O_2$  for larger values of  $\chi$ . Near sunrise and sunset the equations are only rough approximations because only absorption by  $O_2$  is considered. In reality at sunset the sun's rays must pass through dense layers of the atmosphere so that absorption by constituents such as  $O_3$  also becomes important. The results of the time-dependent oxygen atmosphere are depicted graphically in Figures A1 through A24. Numerical tabulations of the  $O$  and  $O_3$  variation during a 24 hour period are given in Table A4 to A37. Profiles were computed and plotted for 90 km, 80 km, 70 km, and 60 km for each of the latitudes  $45^\circ$ ,  $30^\circ$ ,  $15^\circ$ ,  $0^\circ$ ,  $-15^\circ$ ,  $-30^\circ$ ,  $-45^\circ$  for a constant solar declination of  $+23.5^\circ$ , corresponding to a summer solstice in the + latitudes. In addition, profiles were computed at 85 km, 75 km, and 65 km for  $\pm 45^\circ$  latitude and declination of  $23.5^\circ$ . The solutions for each latitude and altitude were generated until the concentration

of  $O_3$  at successive noons was within 5% of each other or until a maximum time of 6 days was reached. At all altitudes, except 90 km, the solutions stabilized to within 5% at successive noons before six days passed, and at 90 km the profiles indicate that stabilization would have been reached within another 24 hour period.



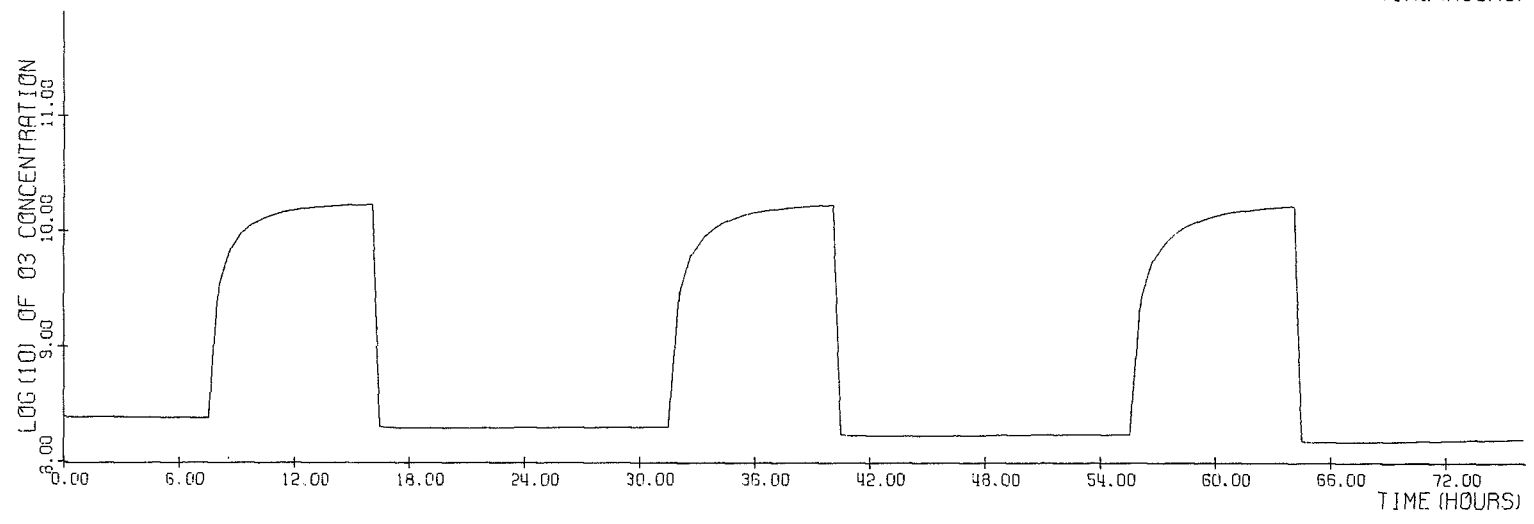
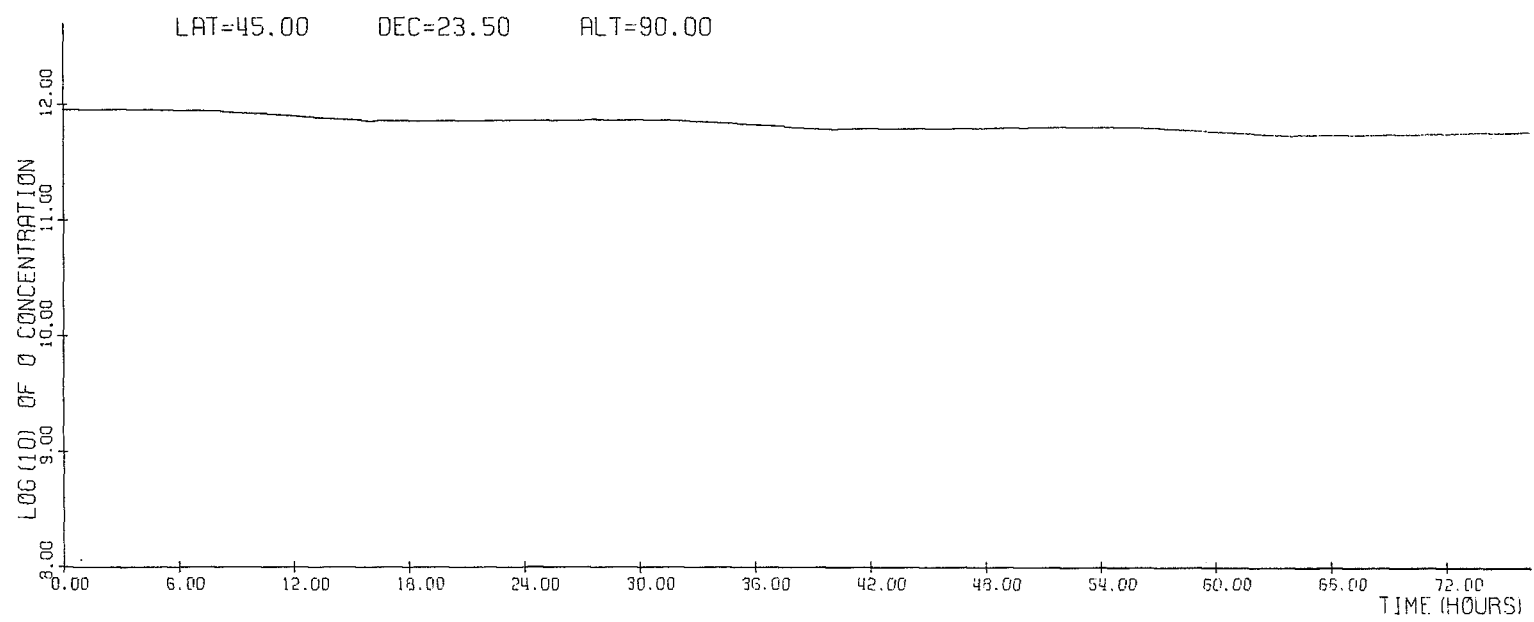


Figure A1

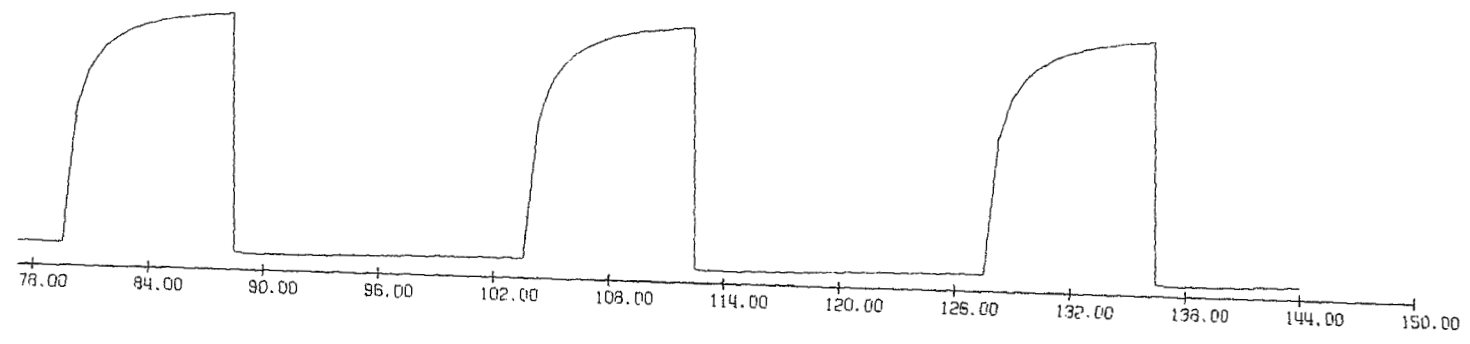
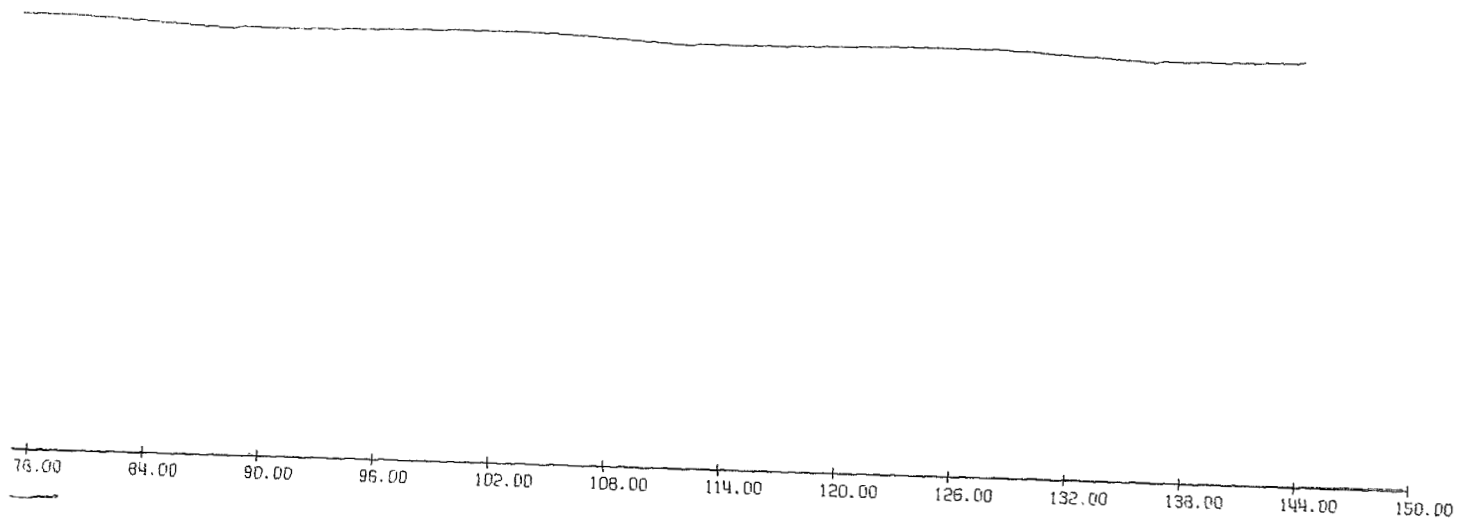


Figure A1 (continued)

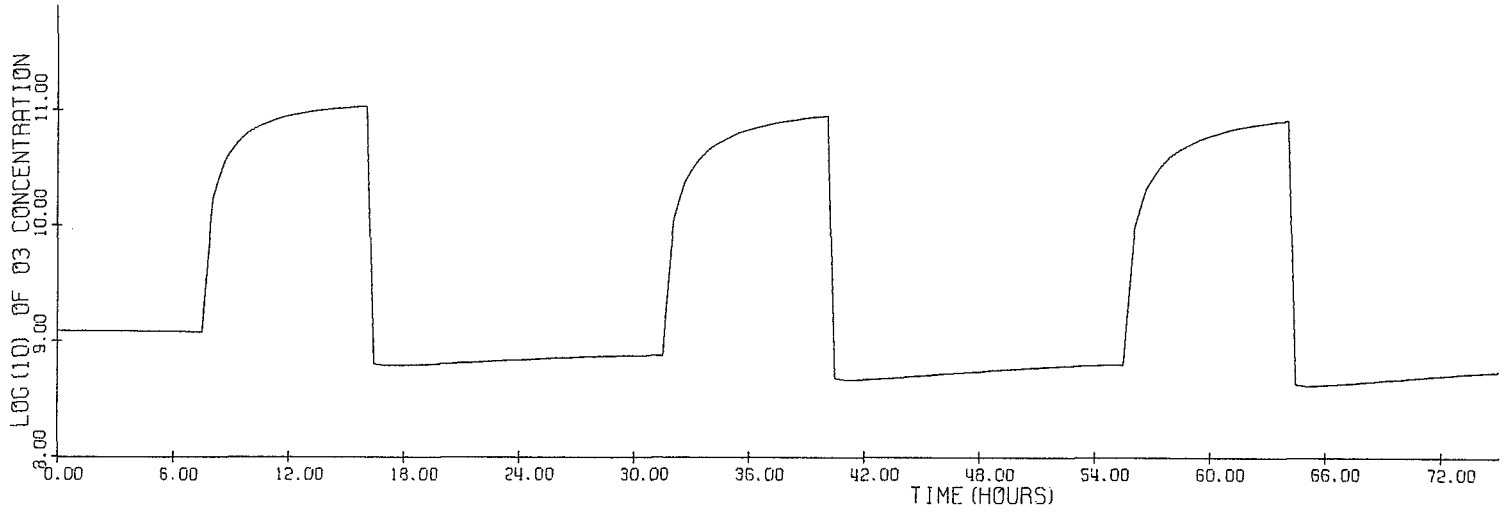
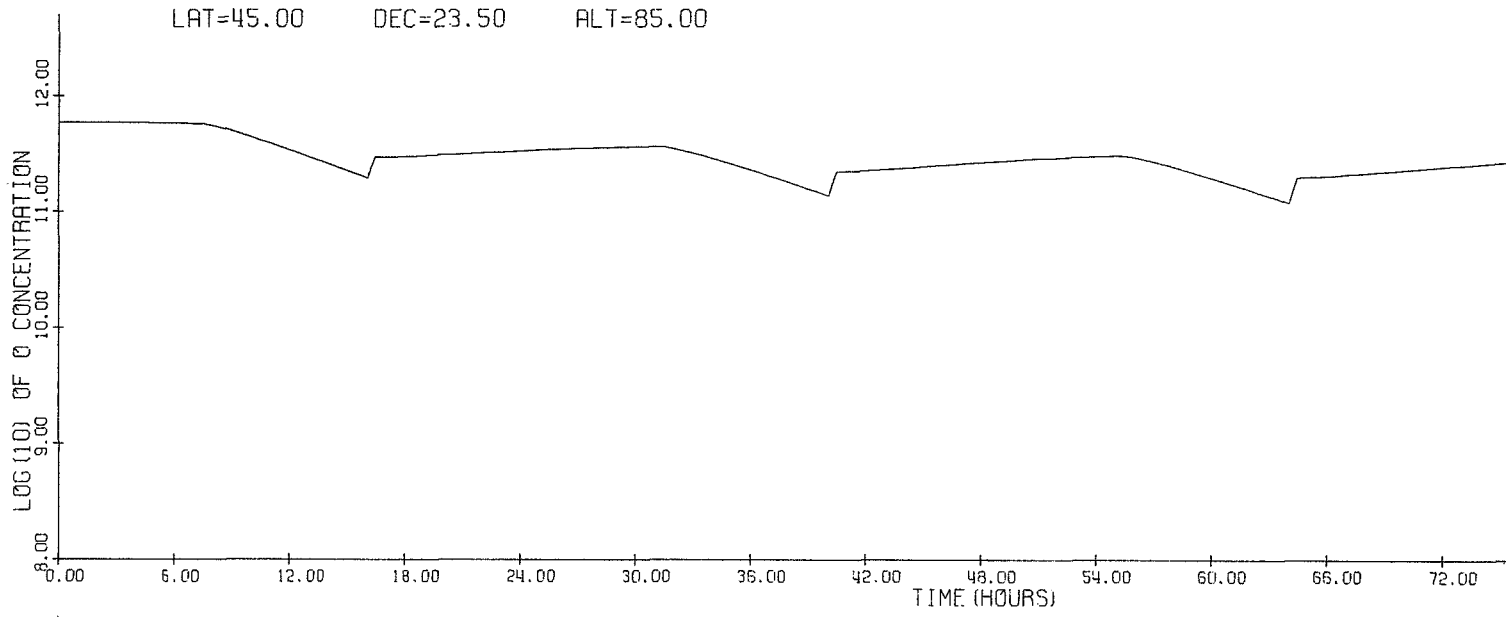


Figure A2

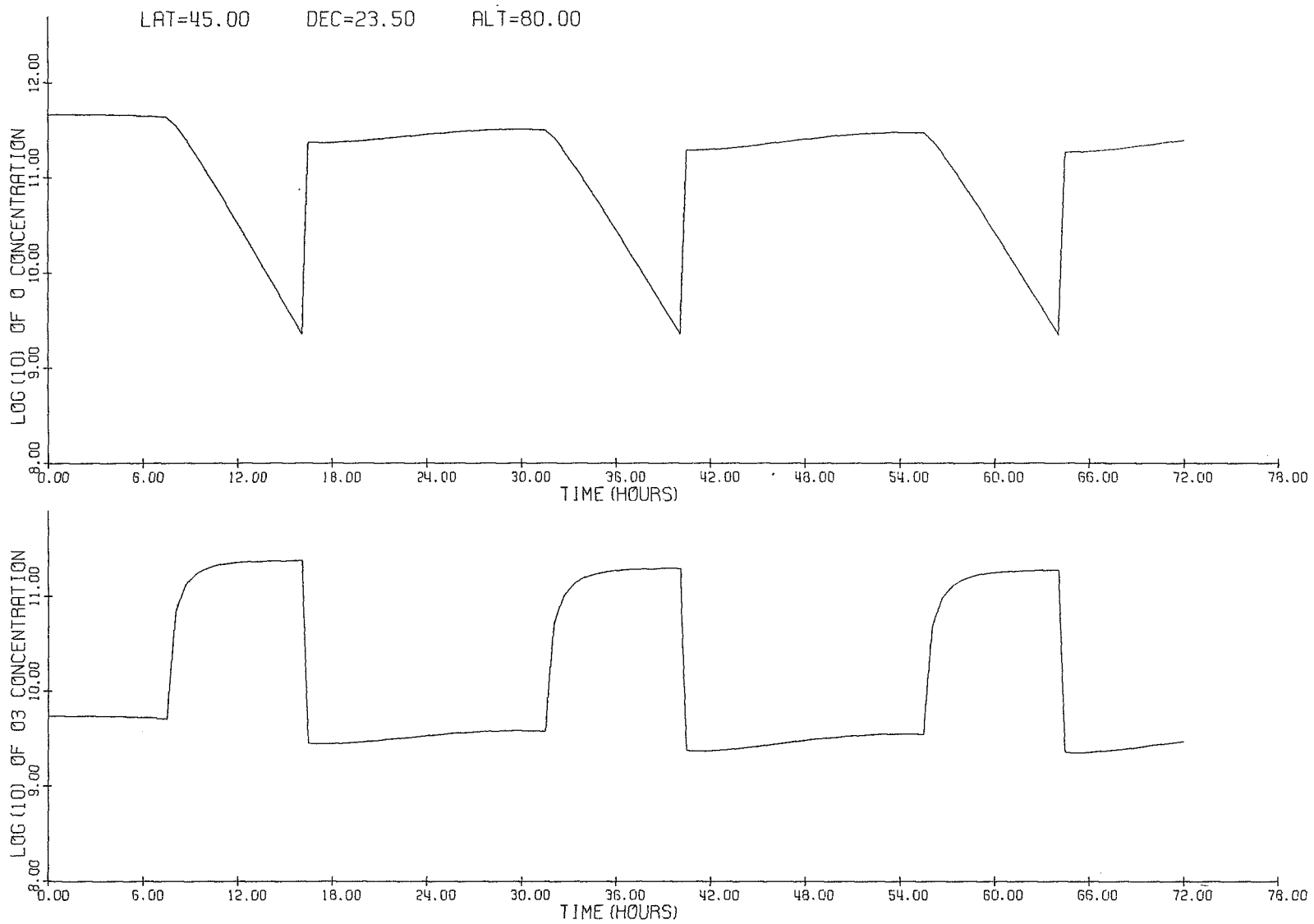


Figure A3

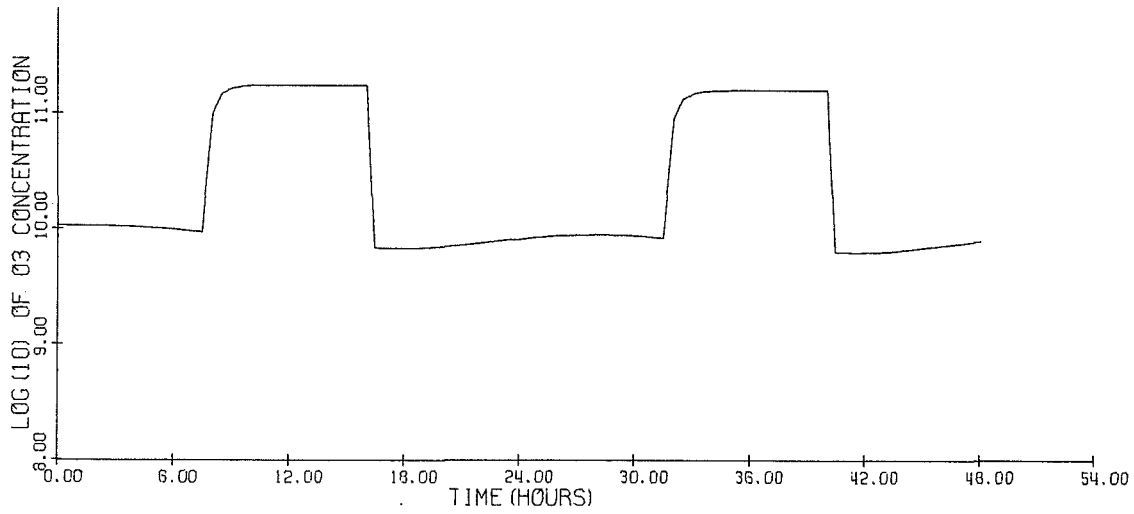
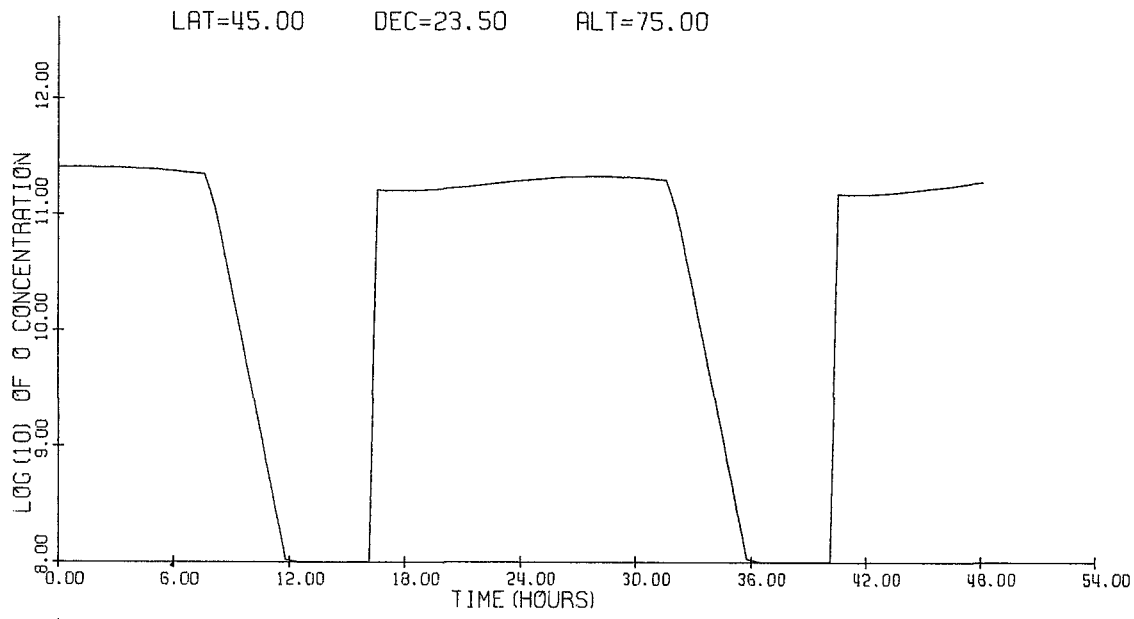


Figure A3 (continued)

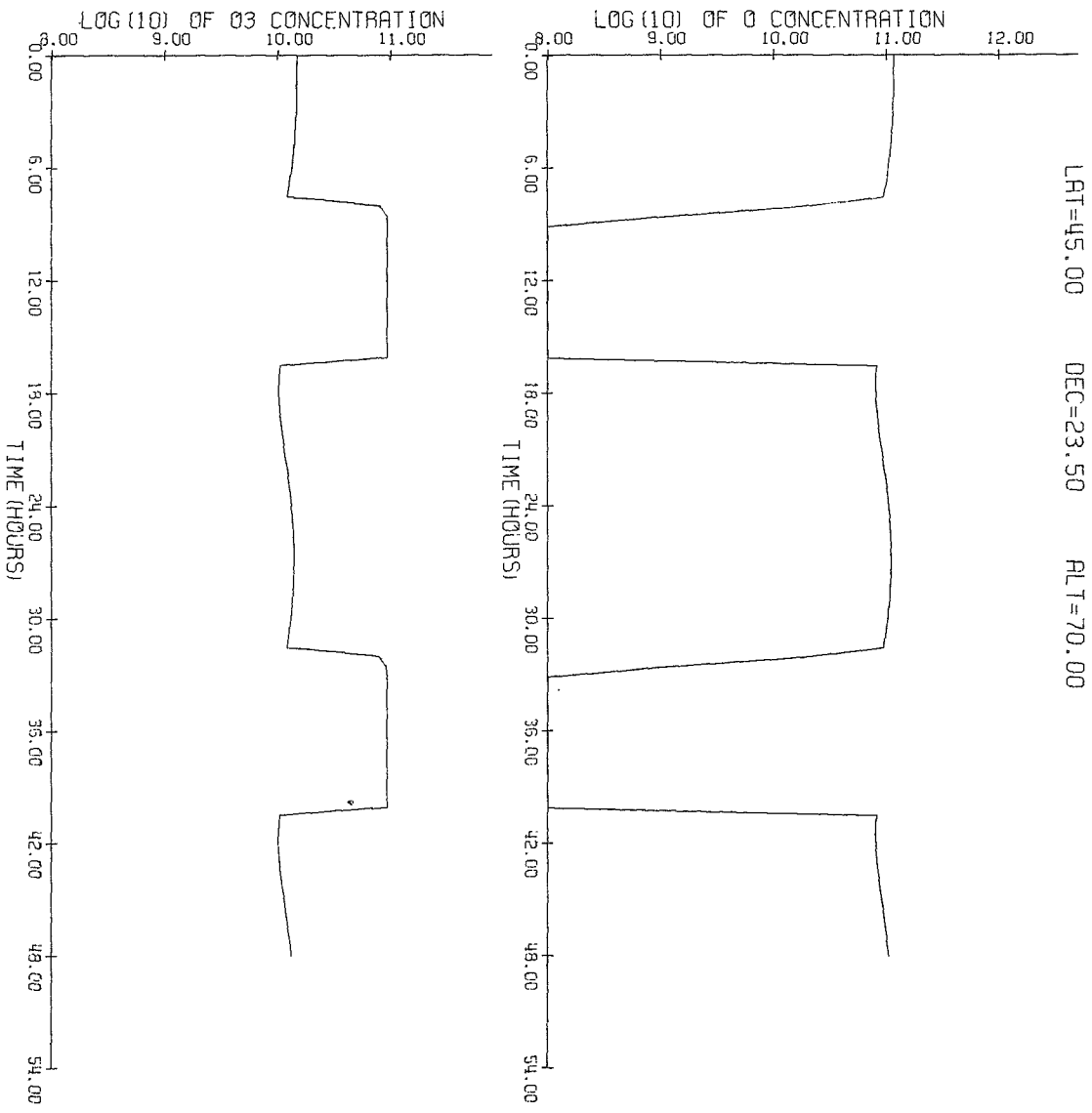


Figure A4

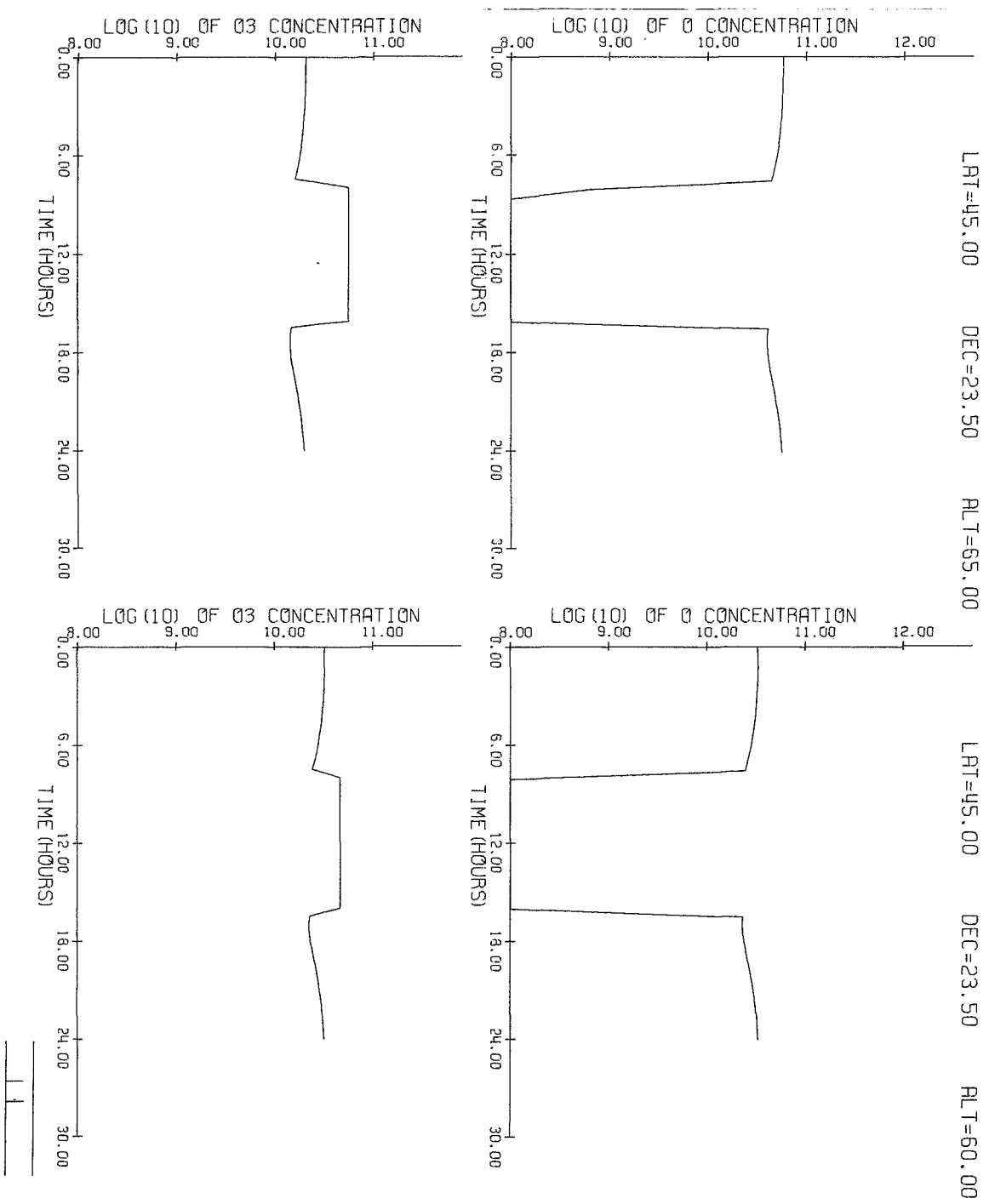


Figure A4 (continued)

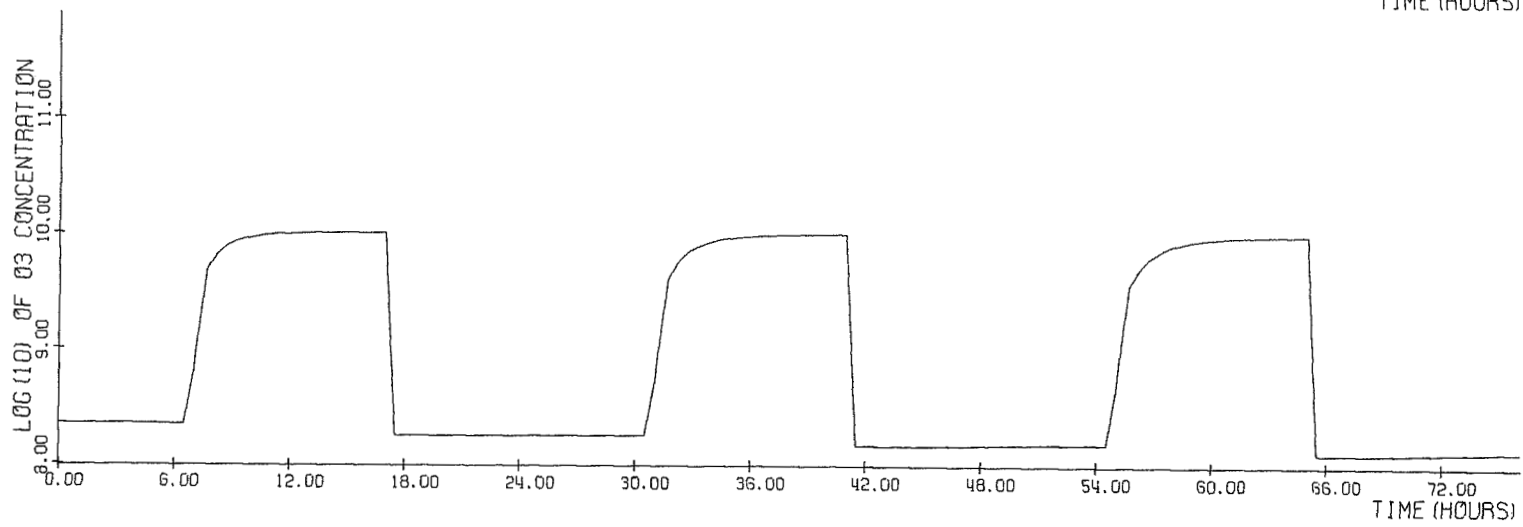
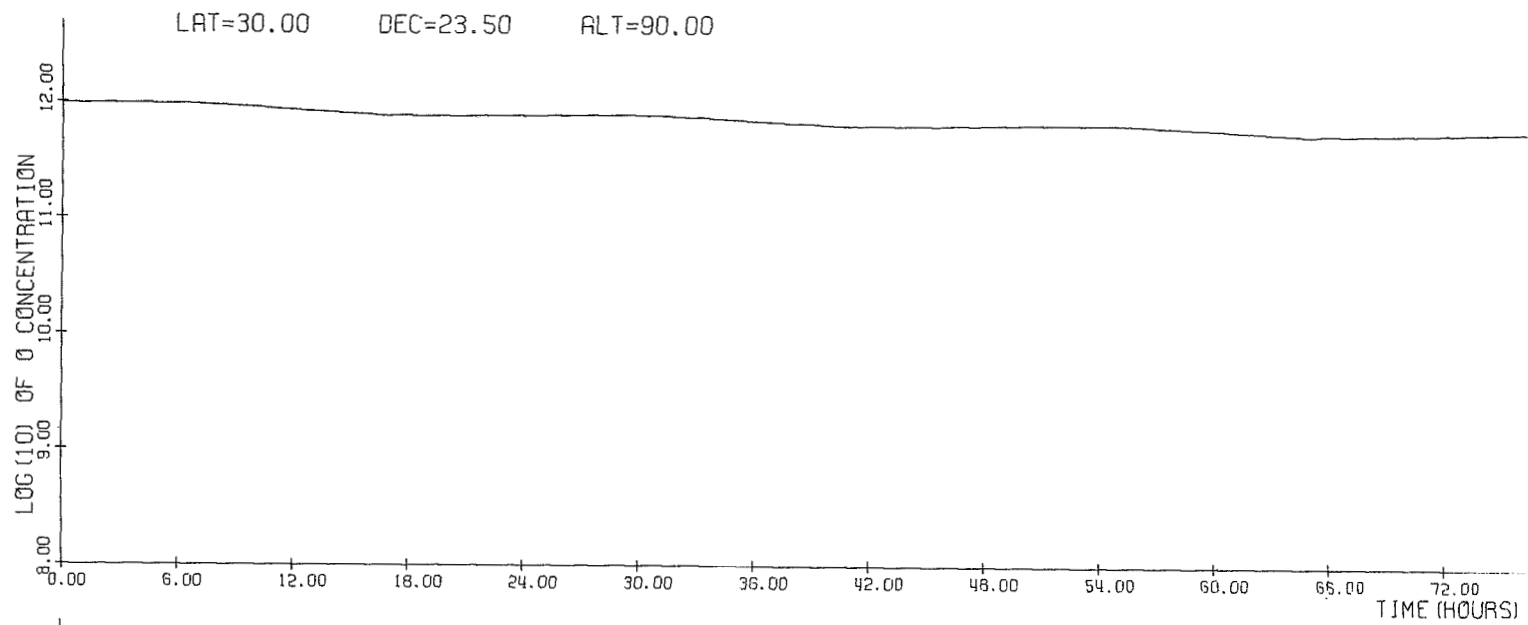
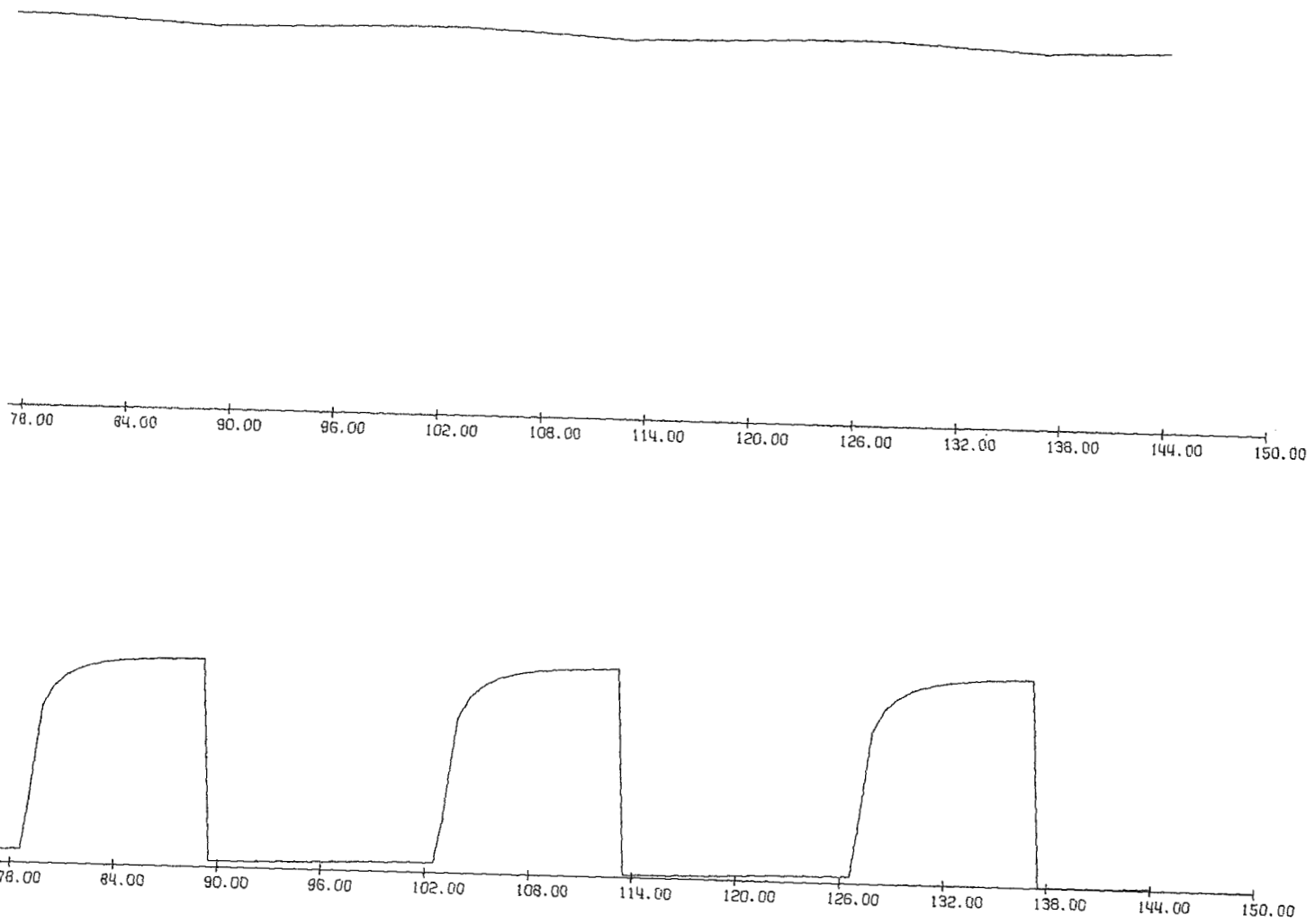
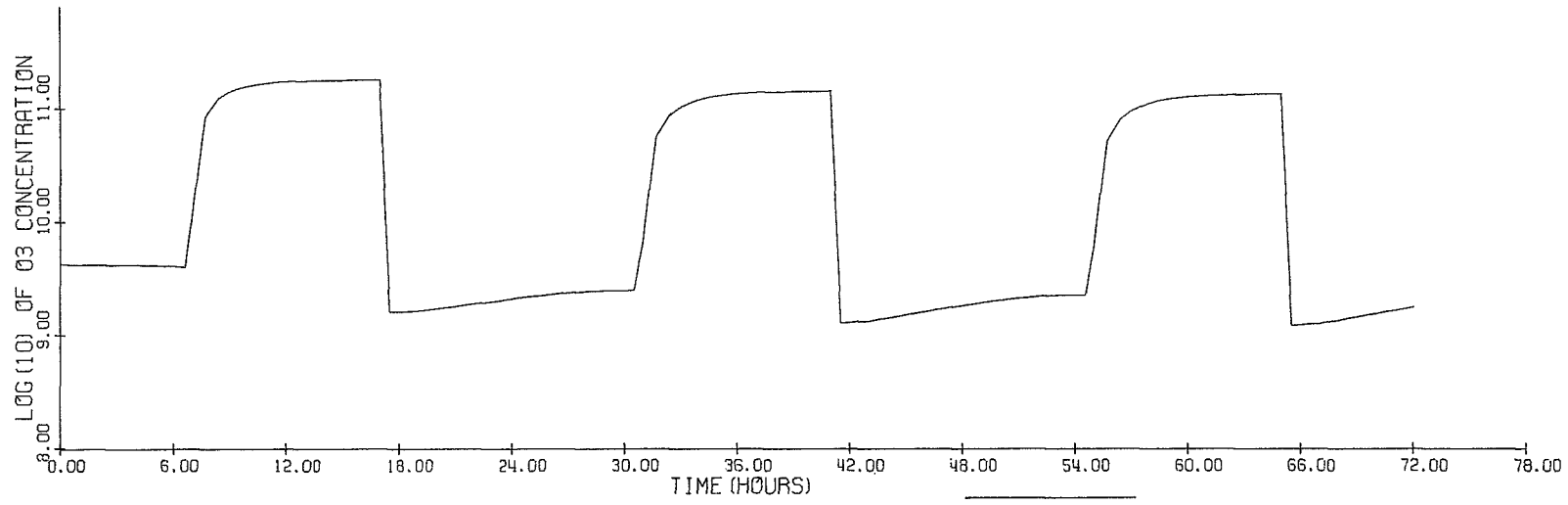
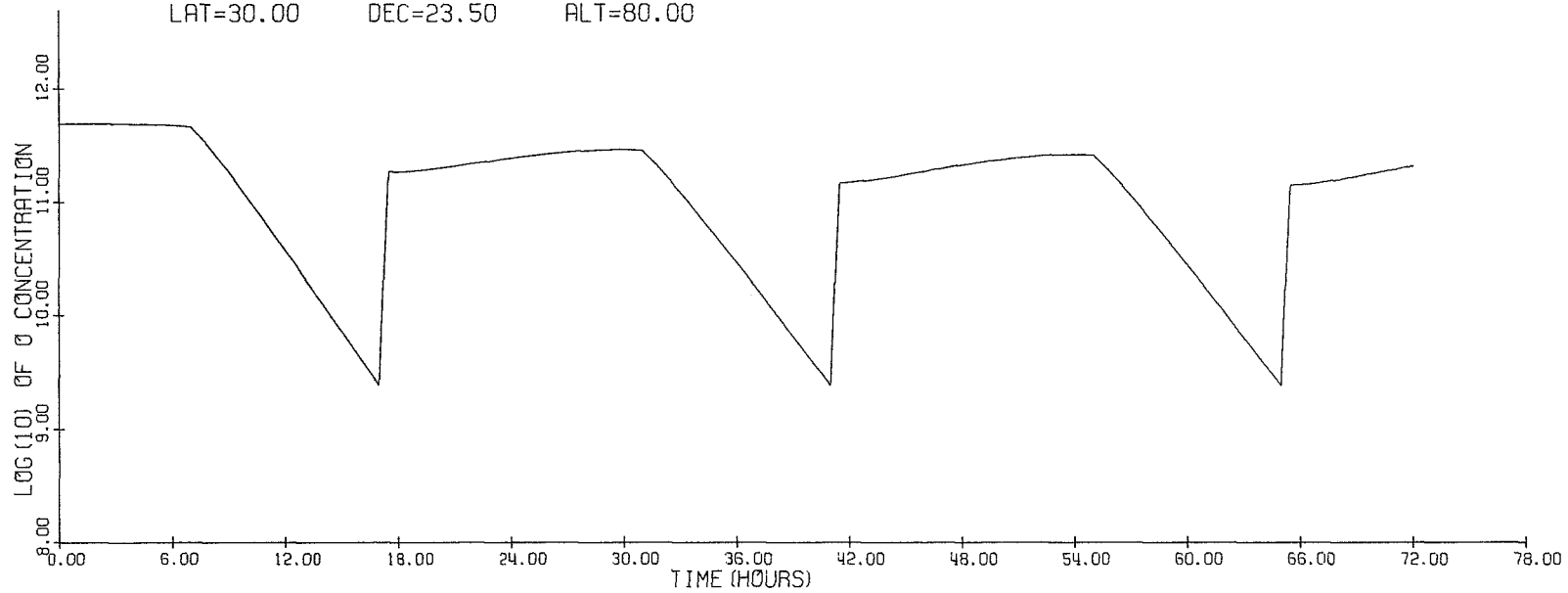


Figure A5





LAT=30.00 DEC=23.50 ALT=80.00



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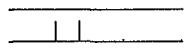


Figure A6

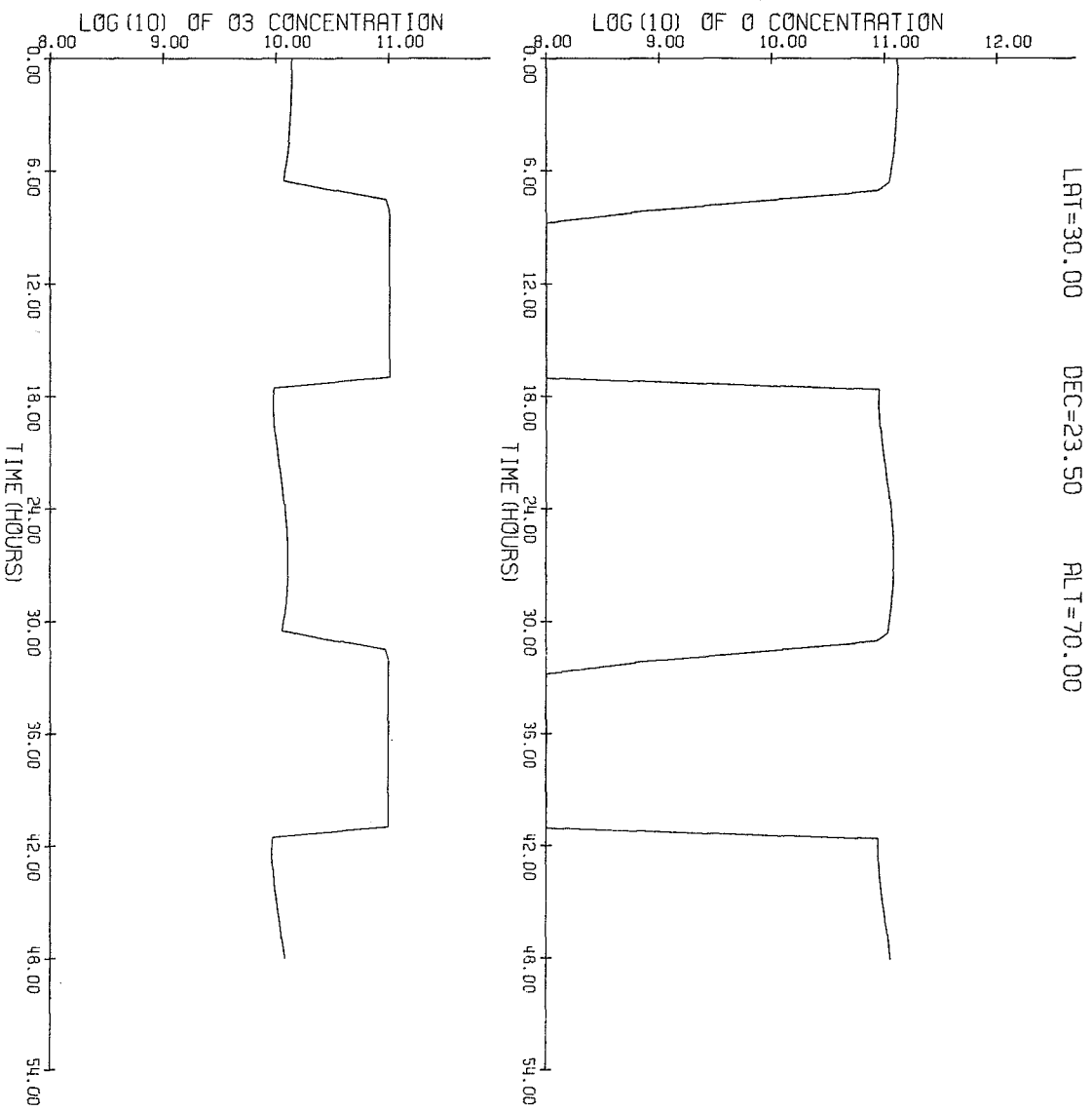


Figure A7

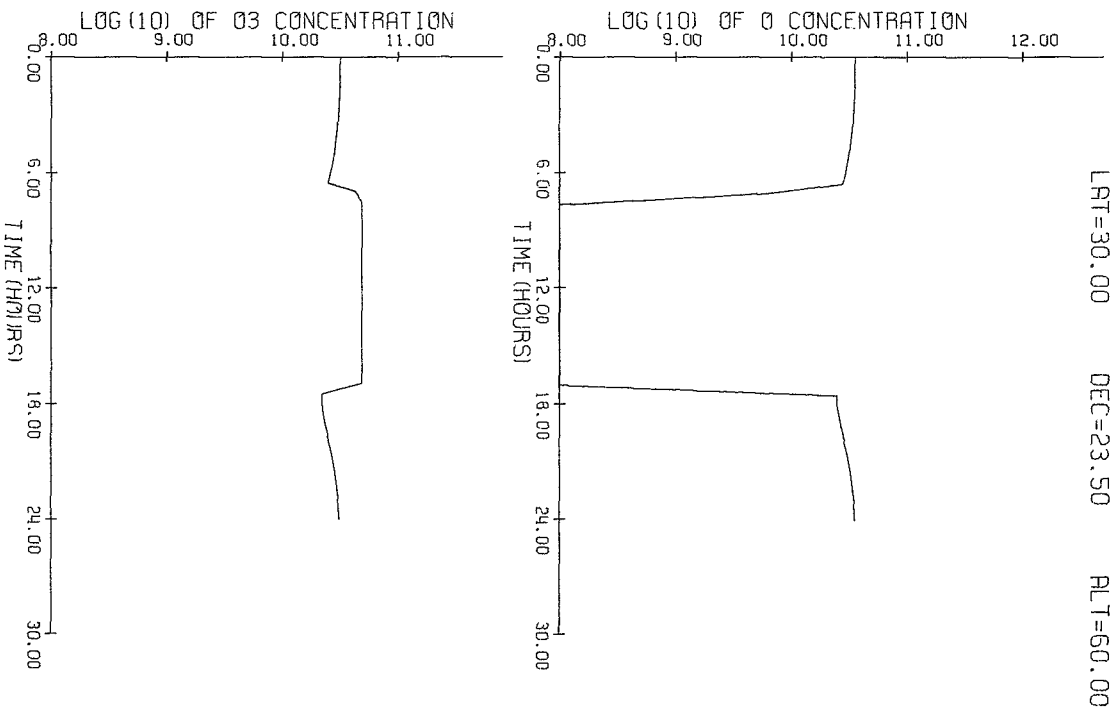


Figure A7 (continued)

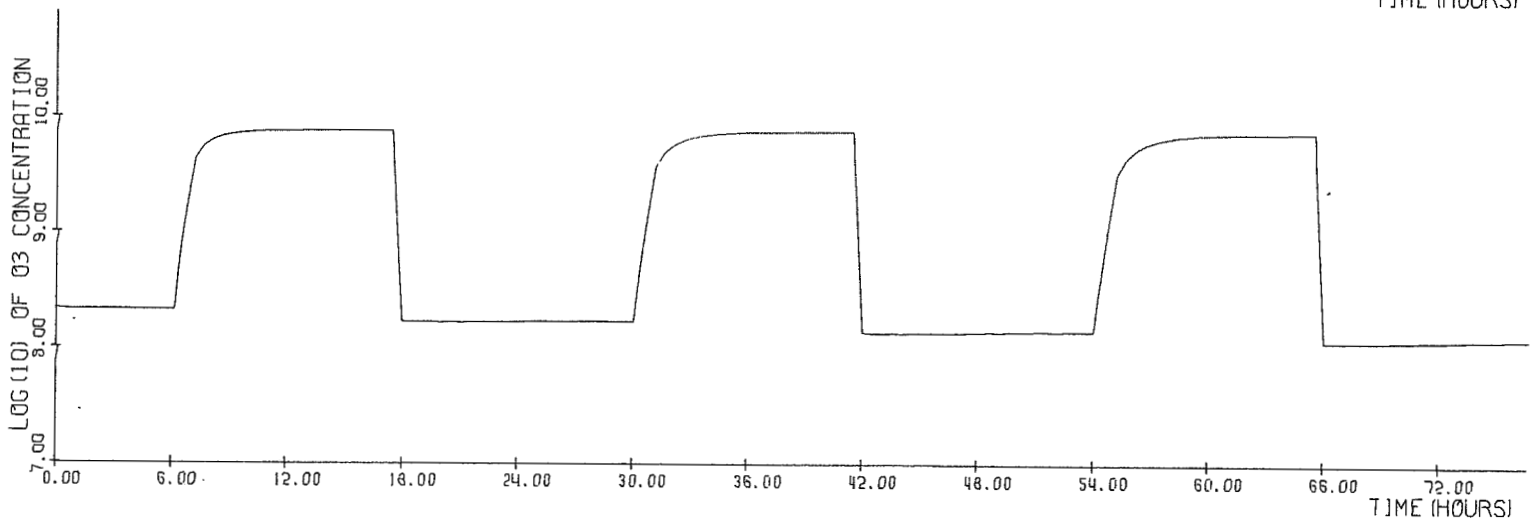
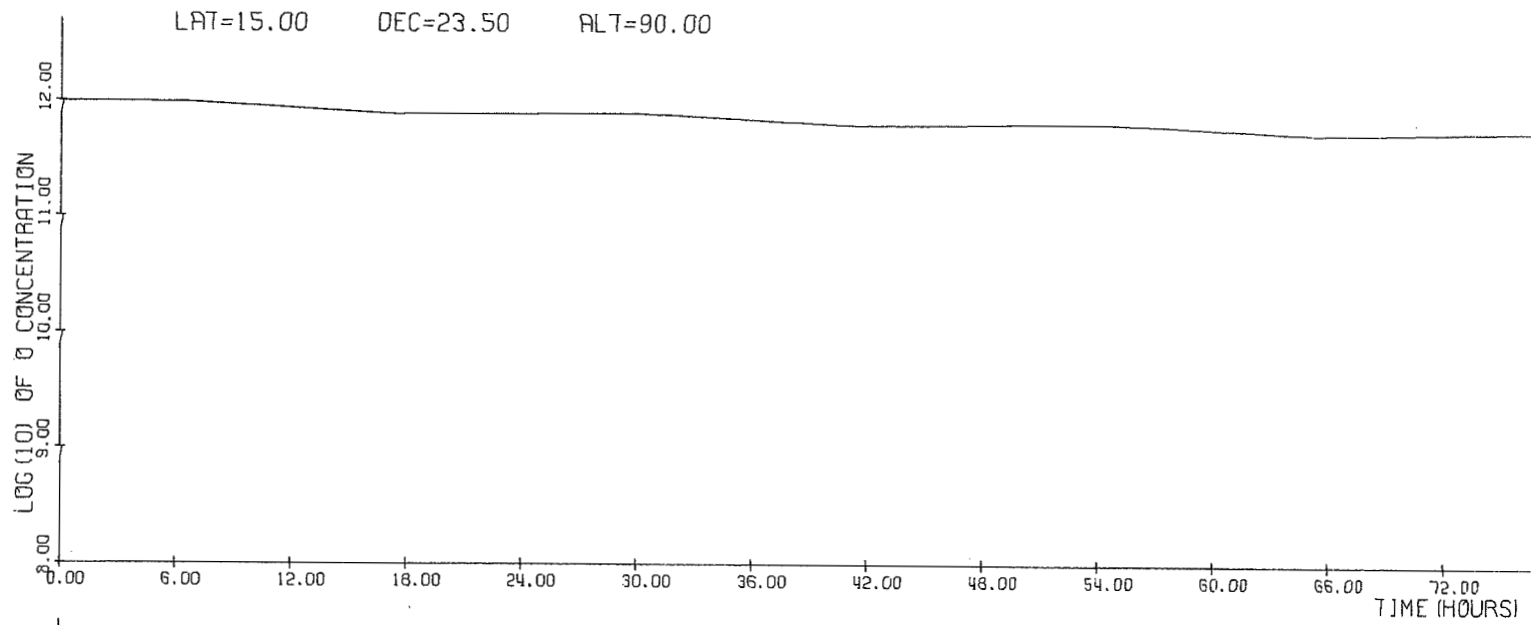


Figure A8

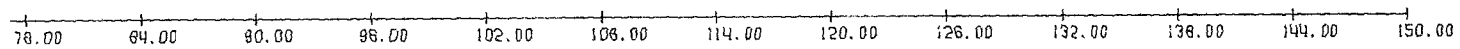
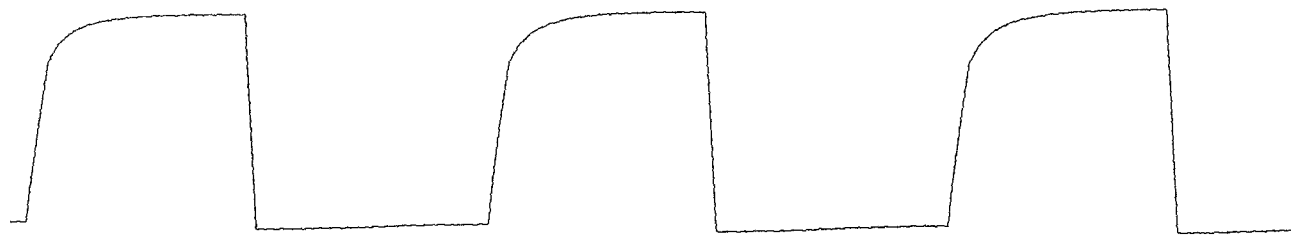
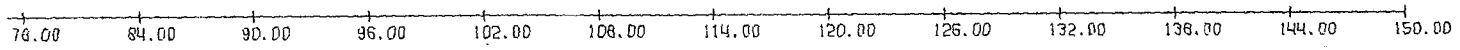


Figure A8 (continued)

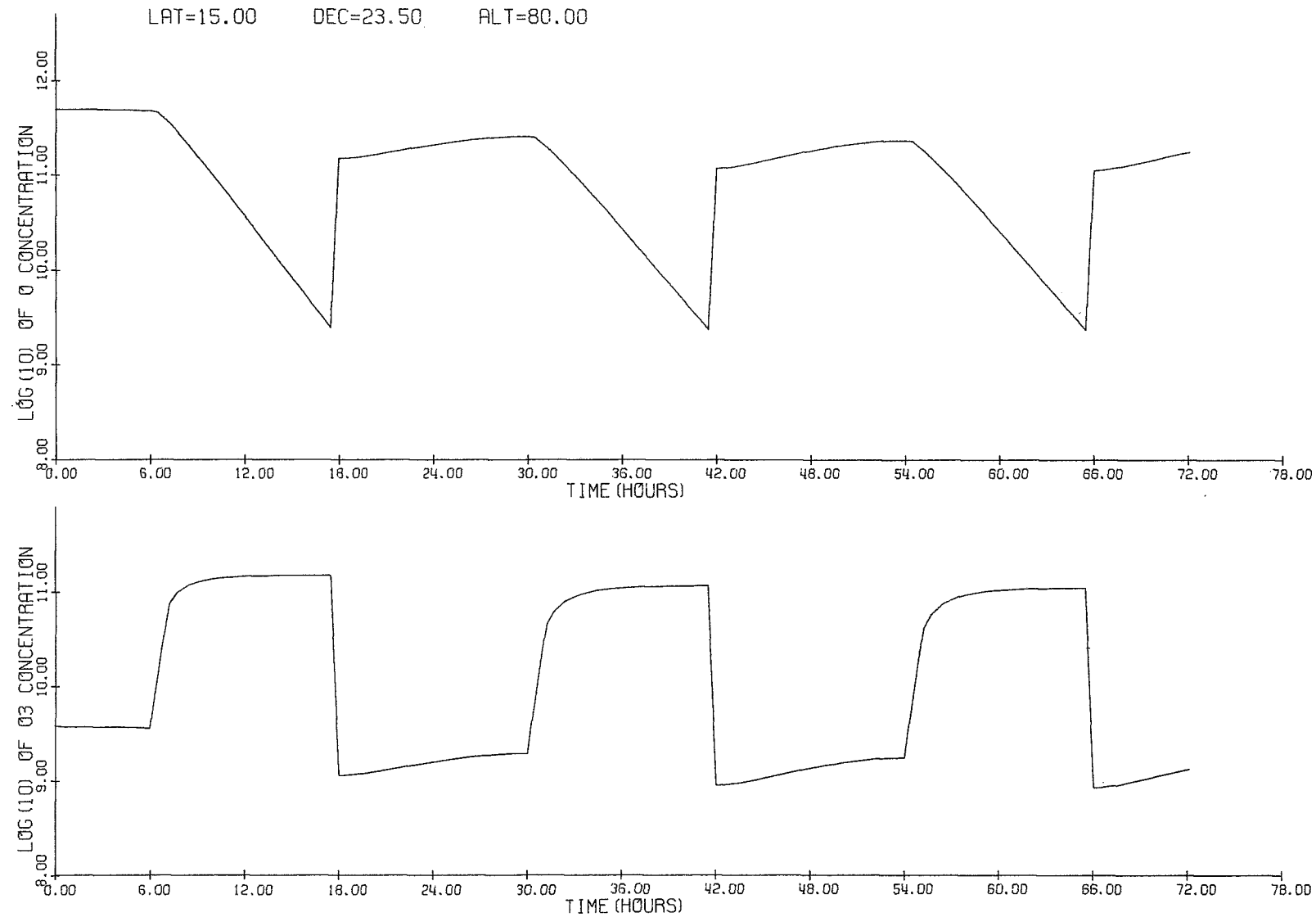


Figure A9

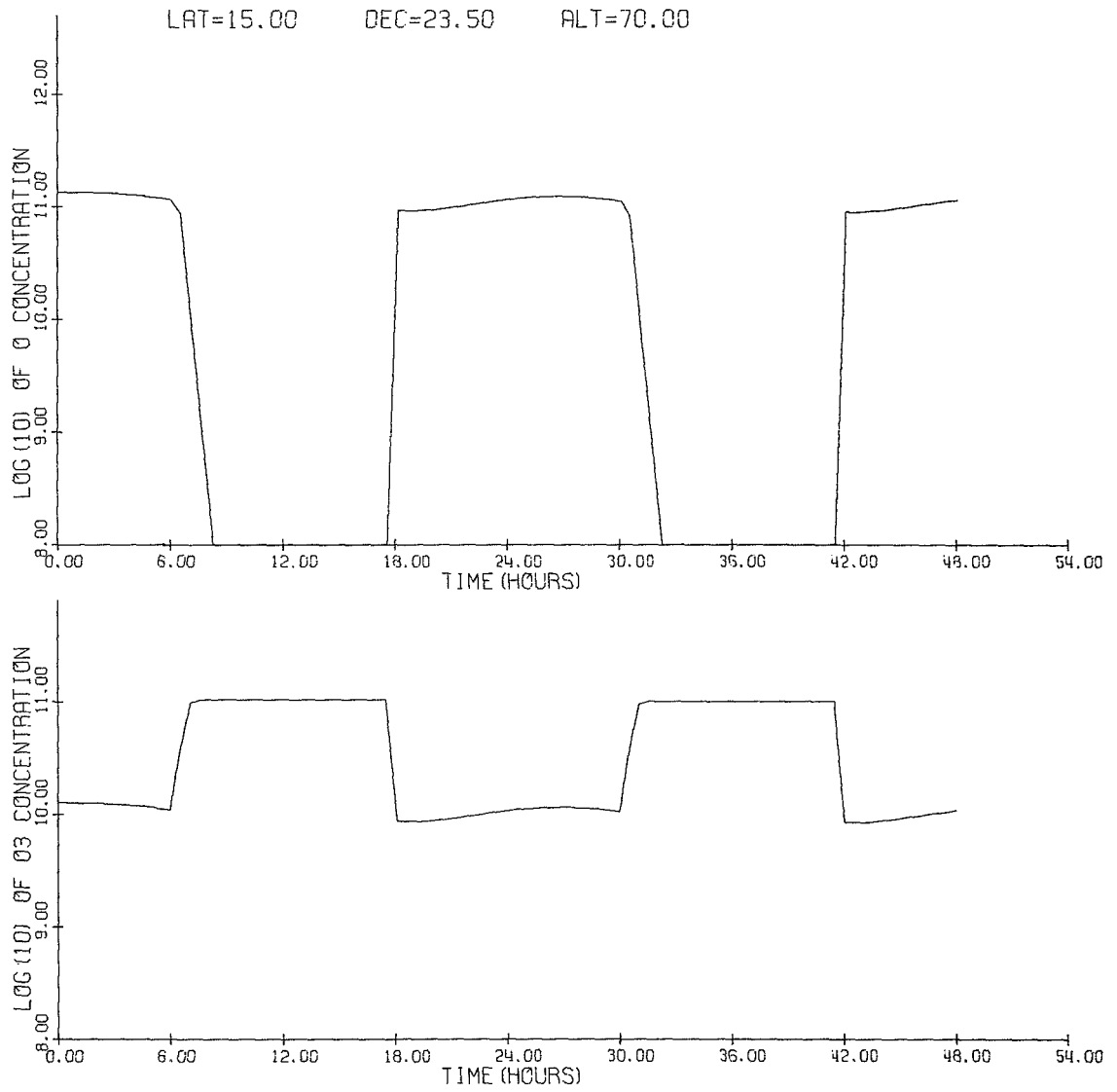


Figure A10



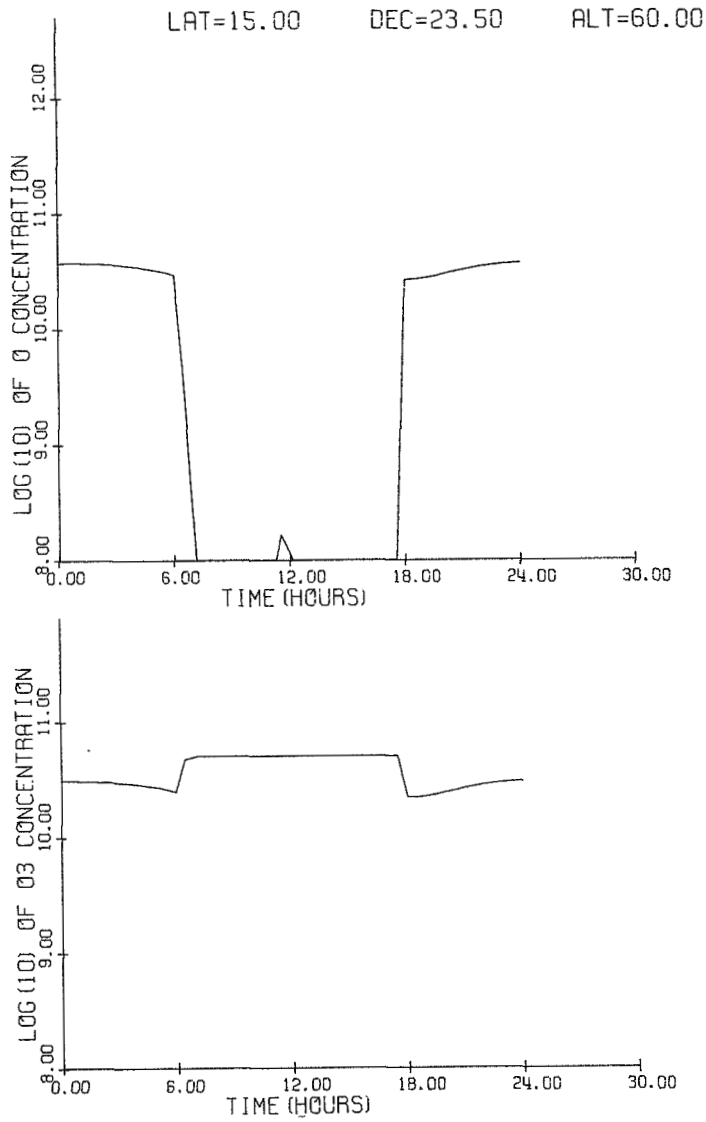


Figure A10 (continued)

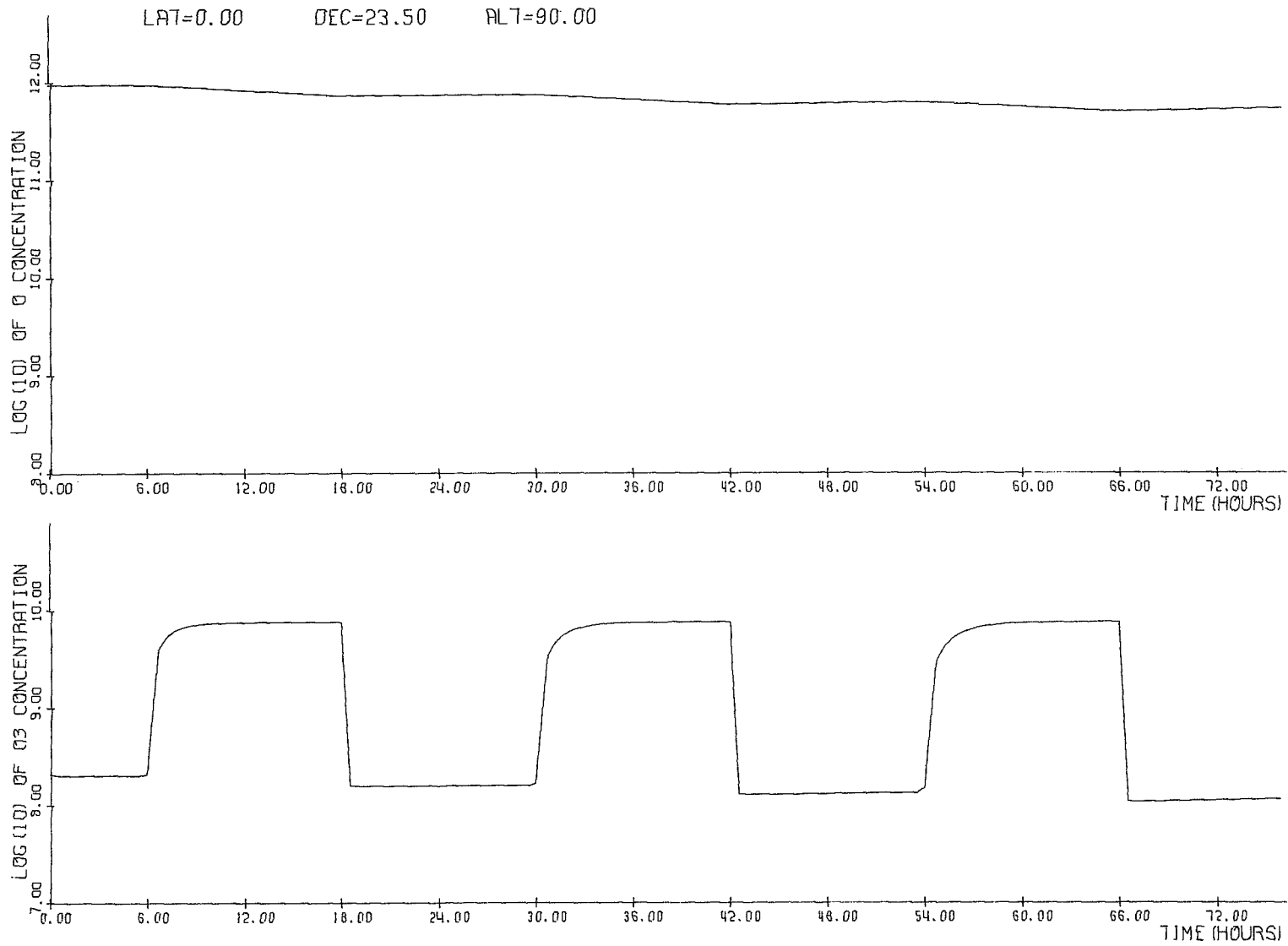


Figure A11

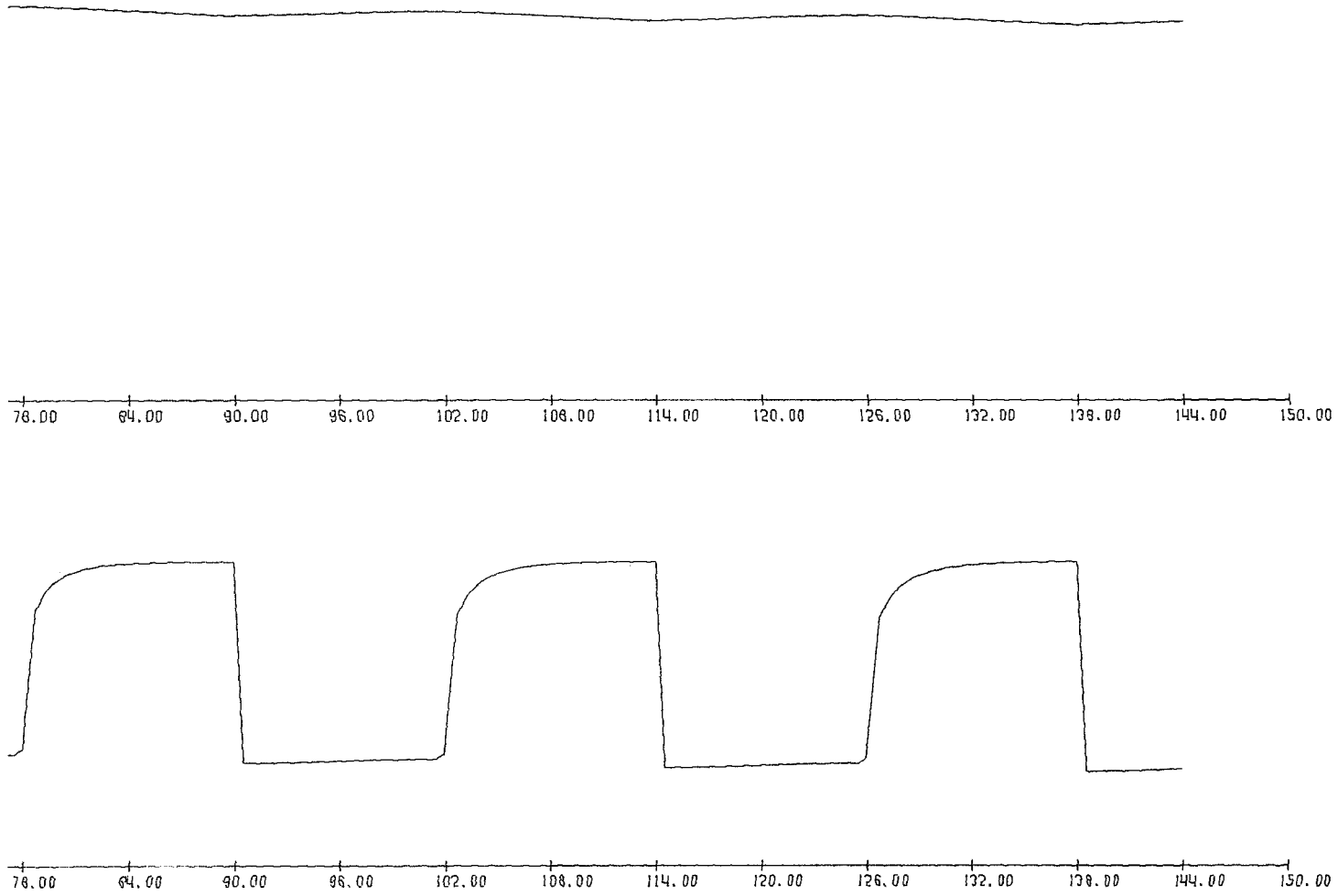


Figure A11 (continued)

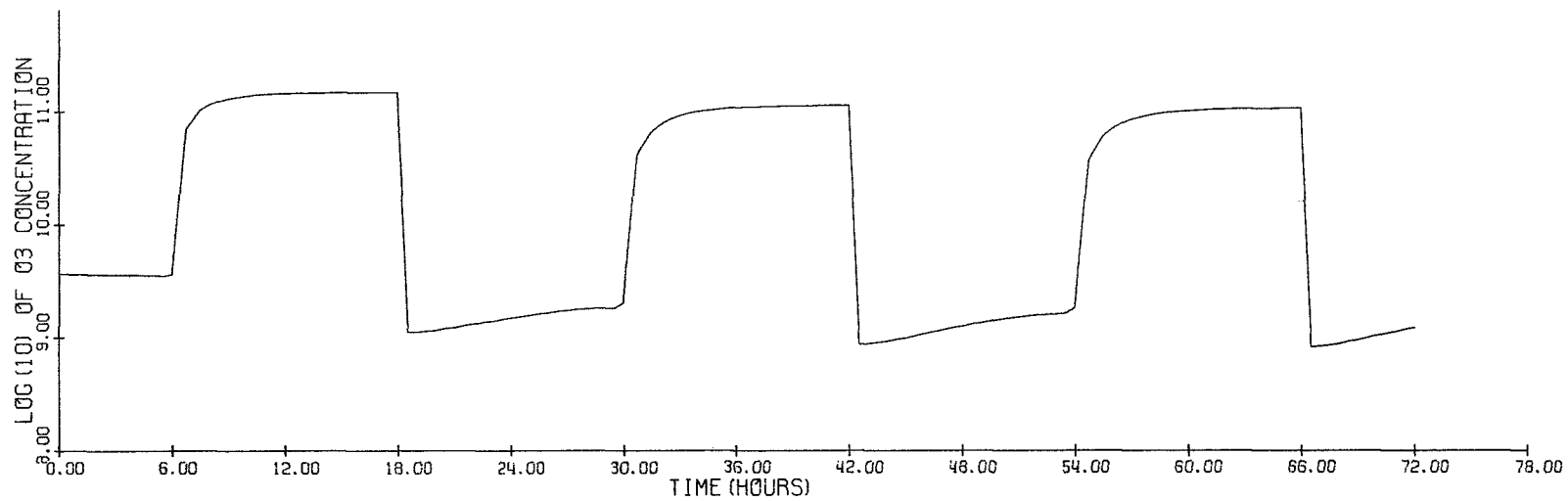
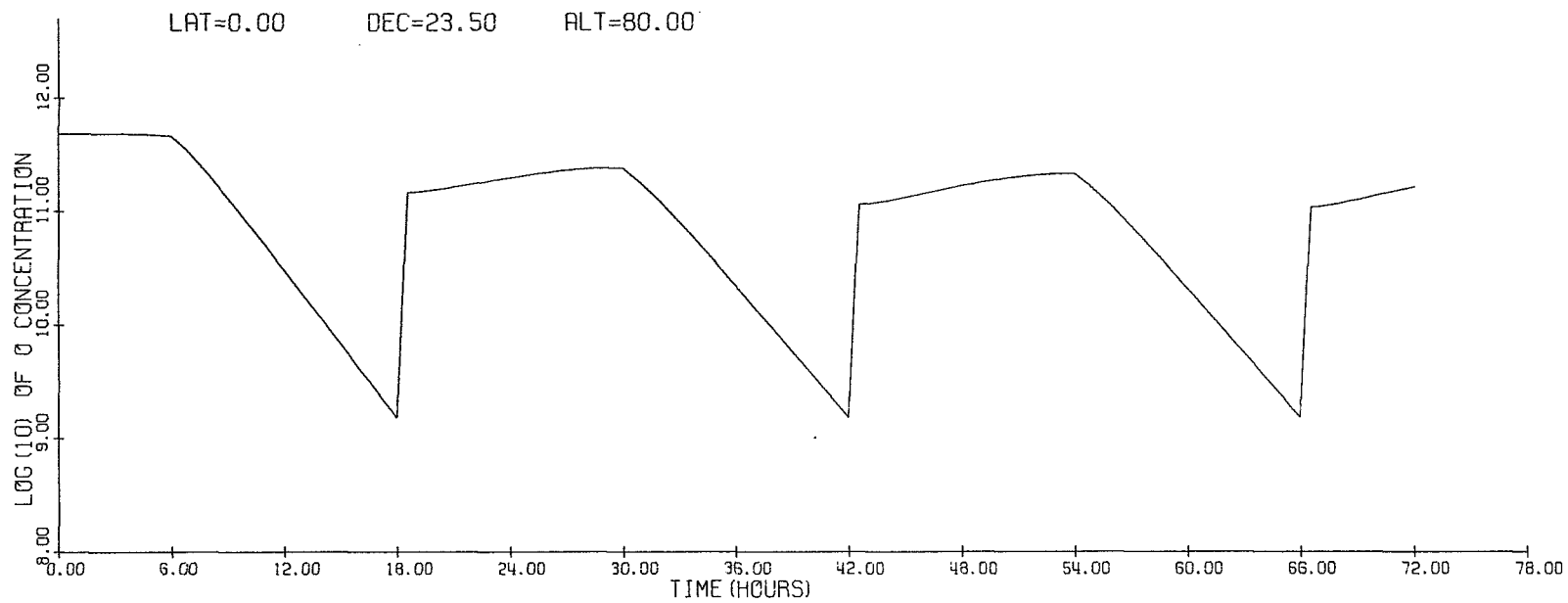


Figure A12

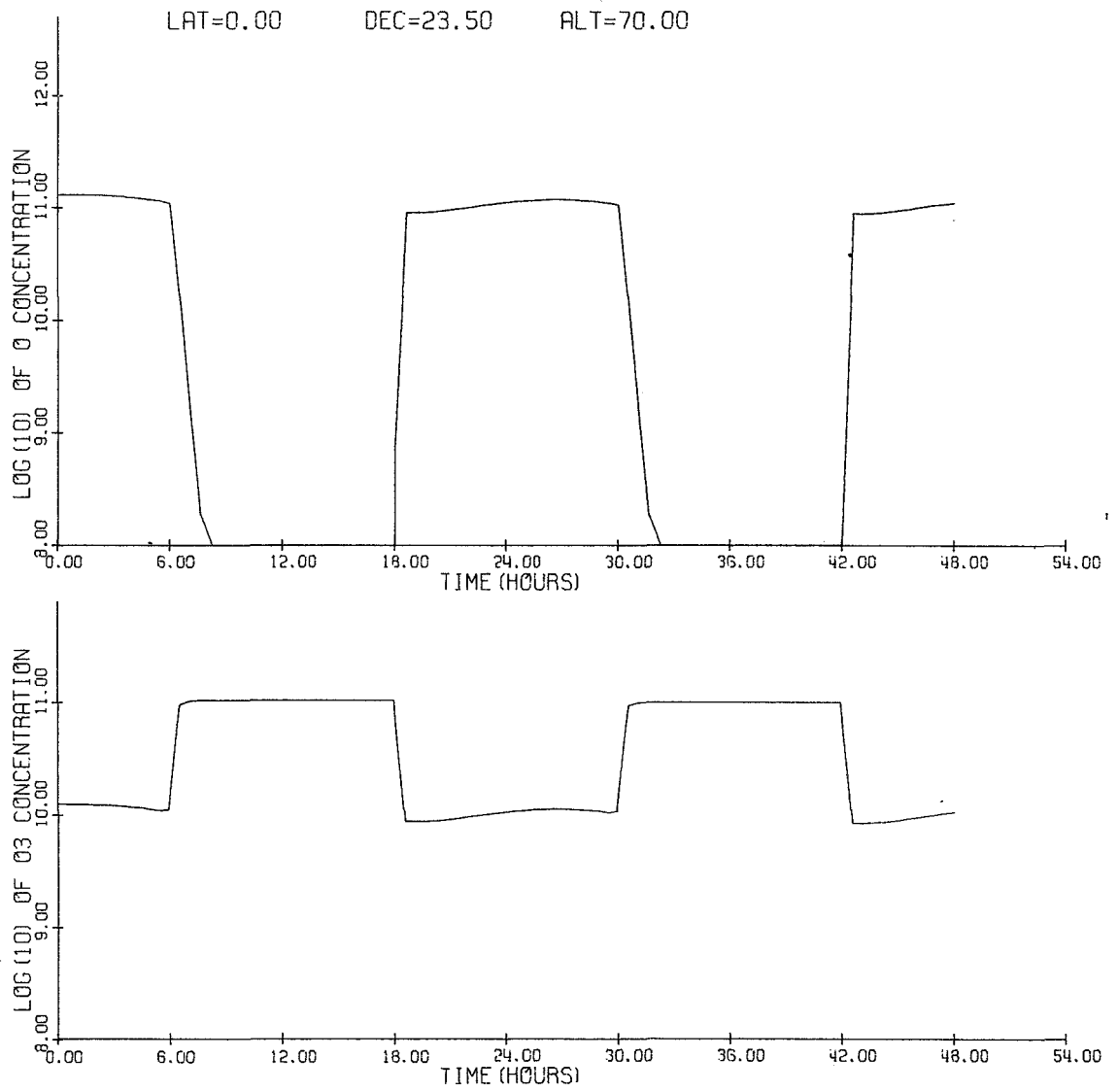


Figure A13

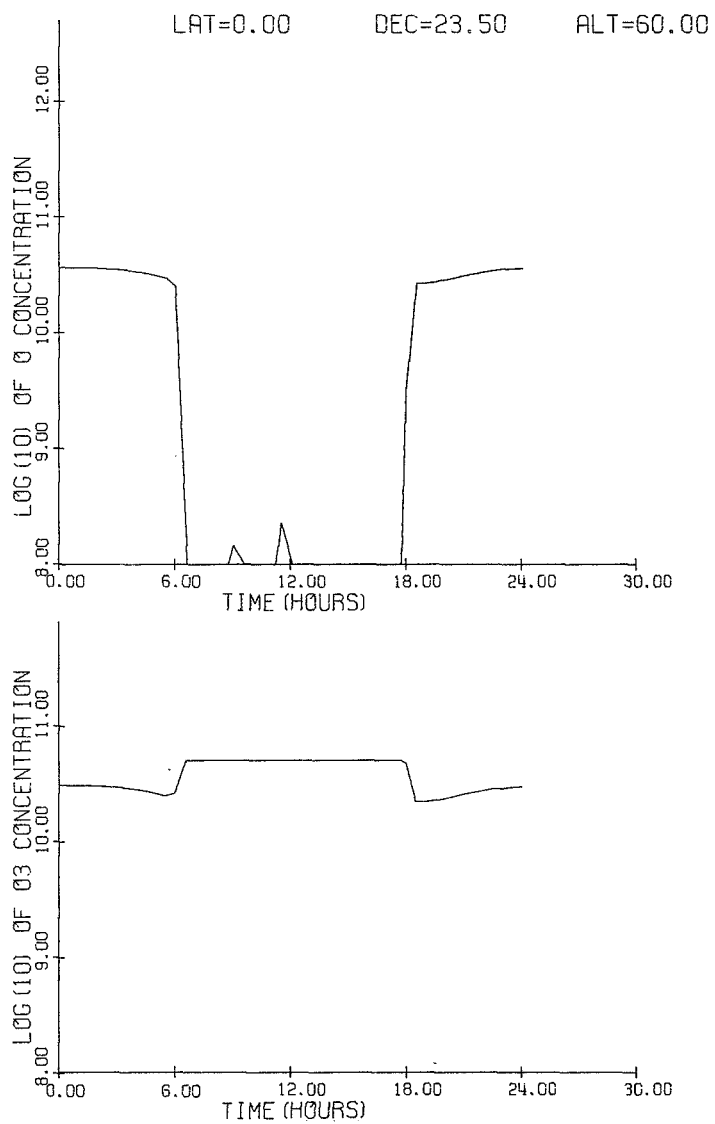


Figure A13 (continued)

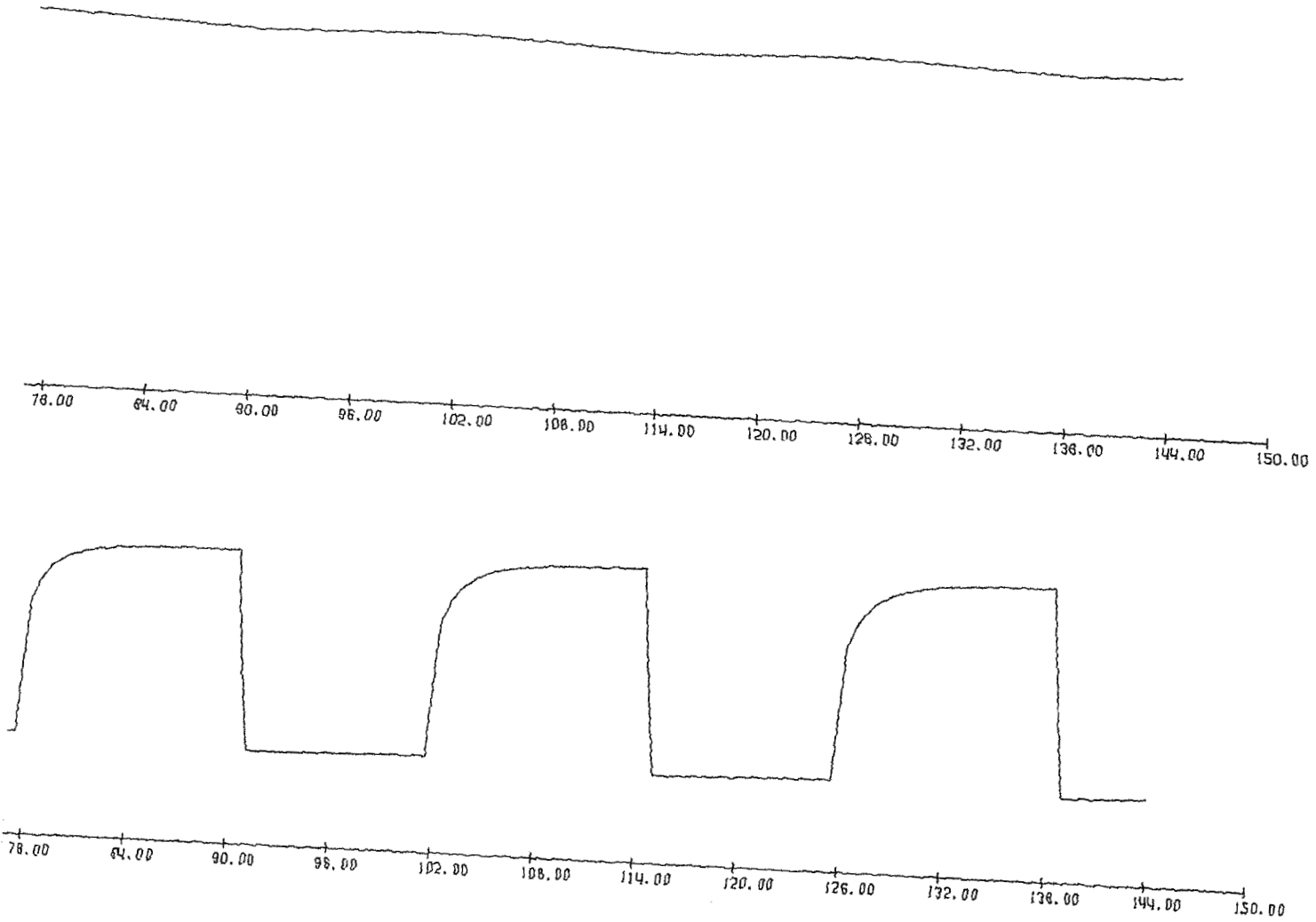


Figure A14

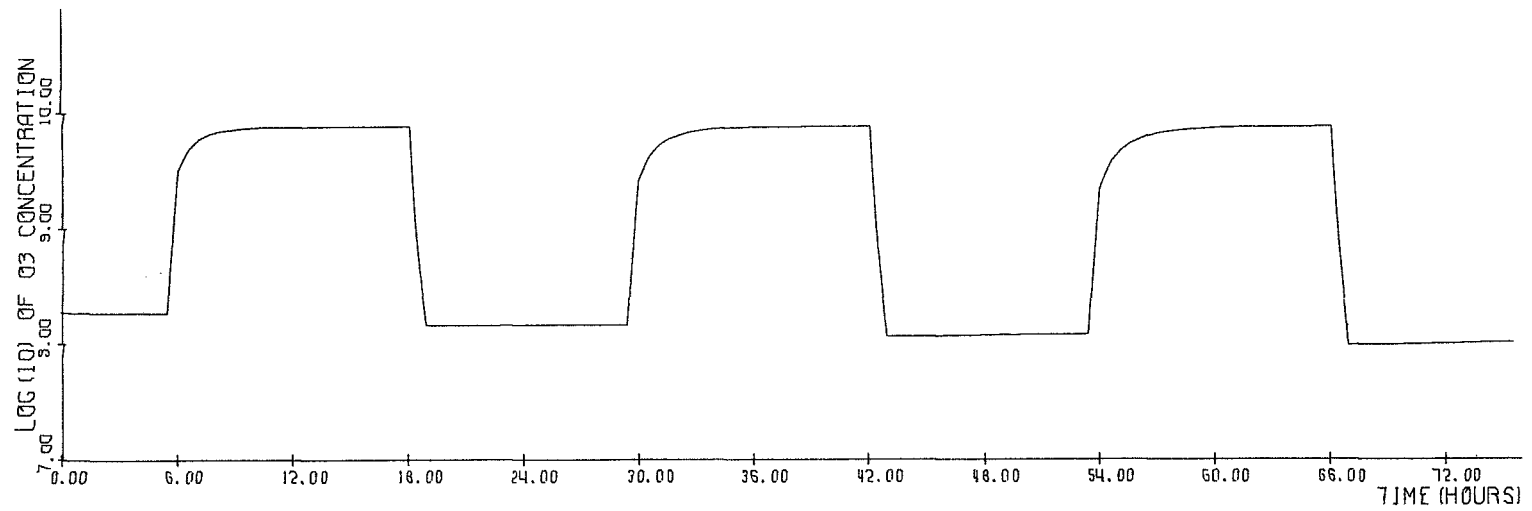
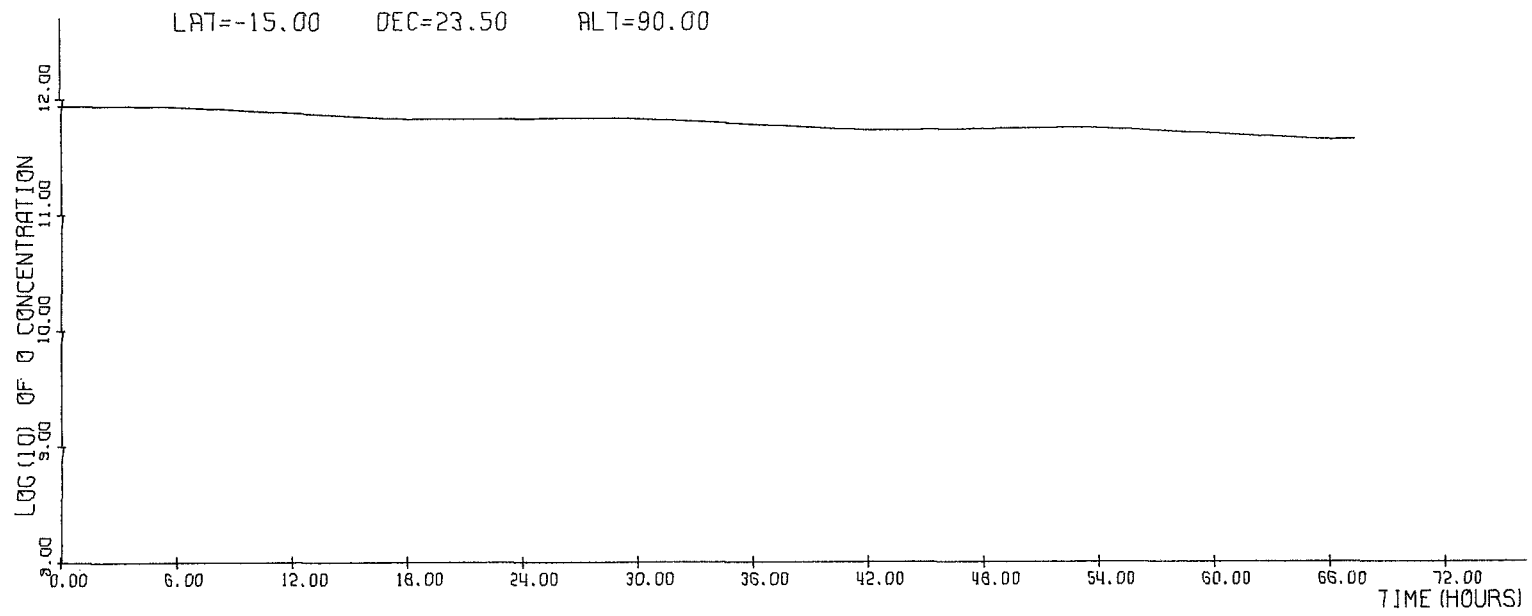


Figure A14 (continued)



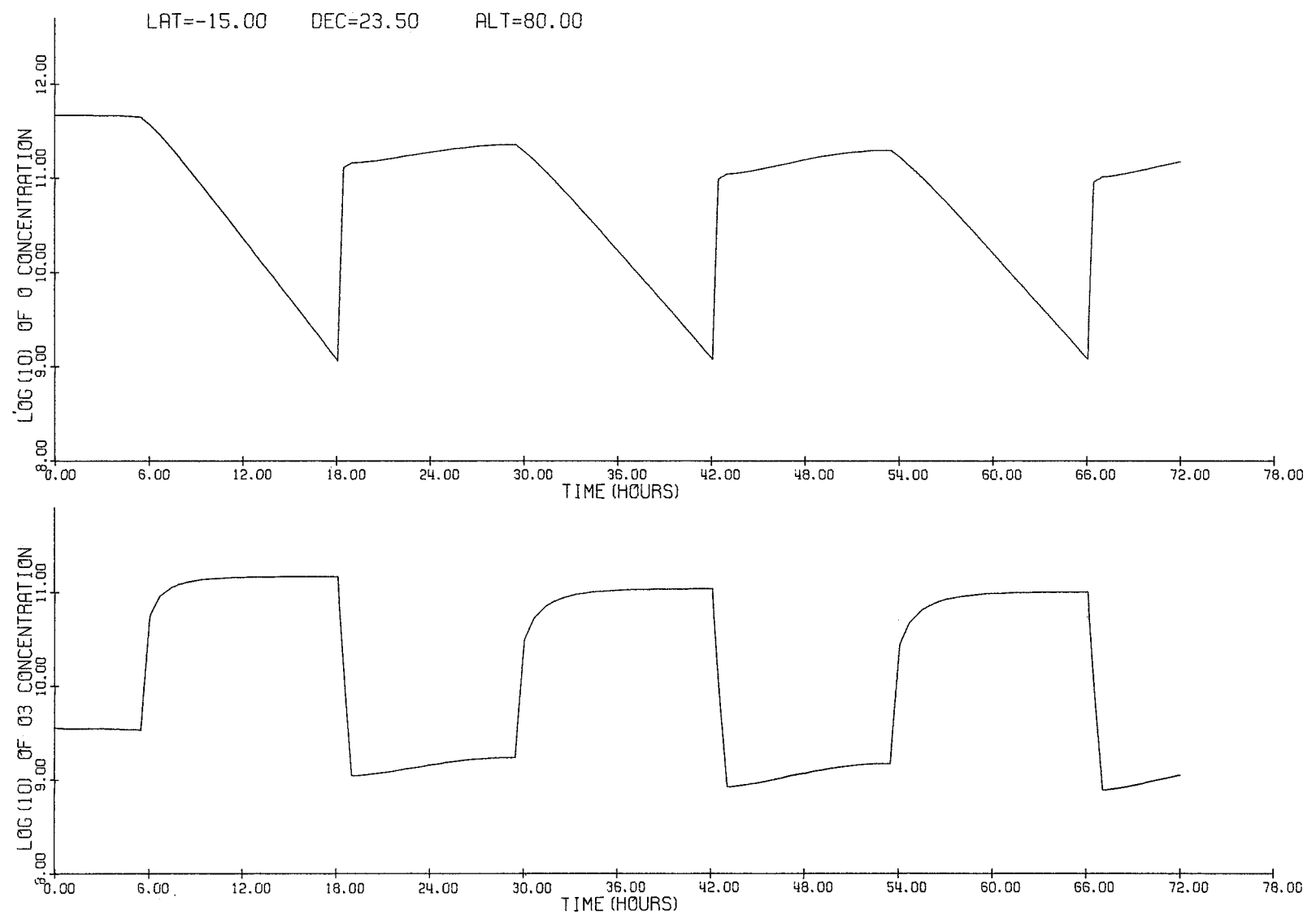


Figure A15

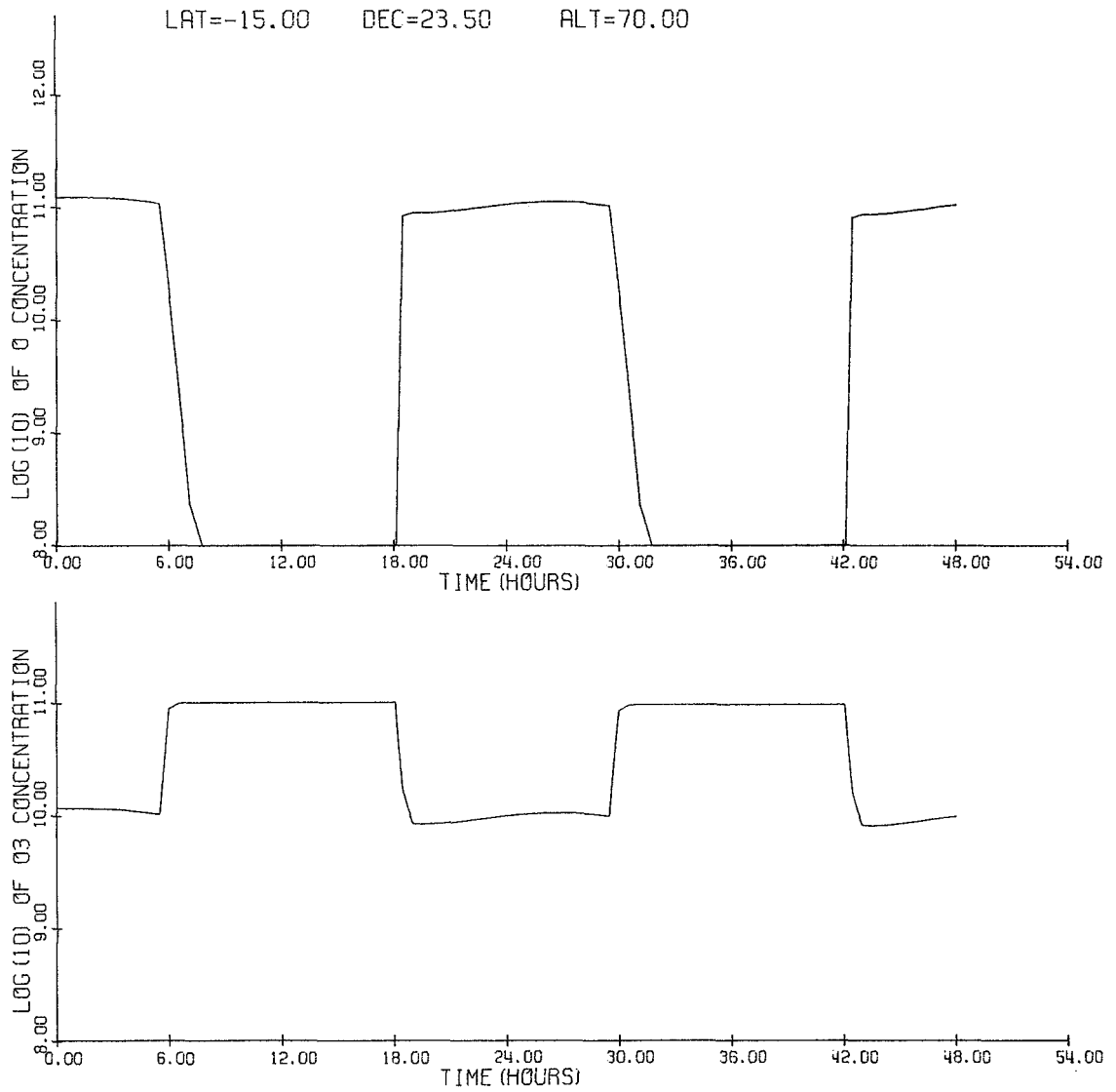


Figure A16

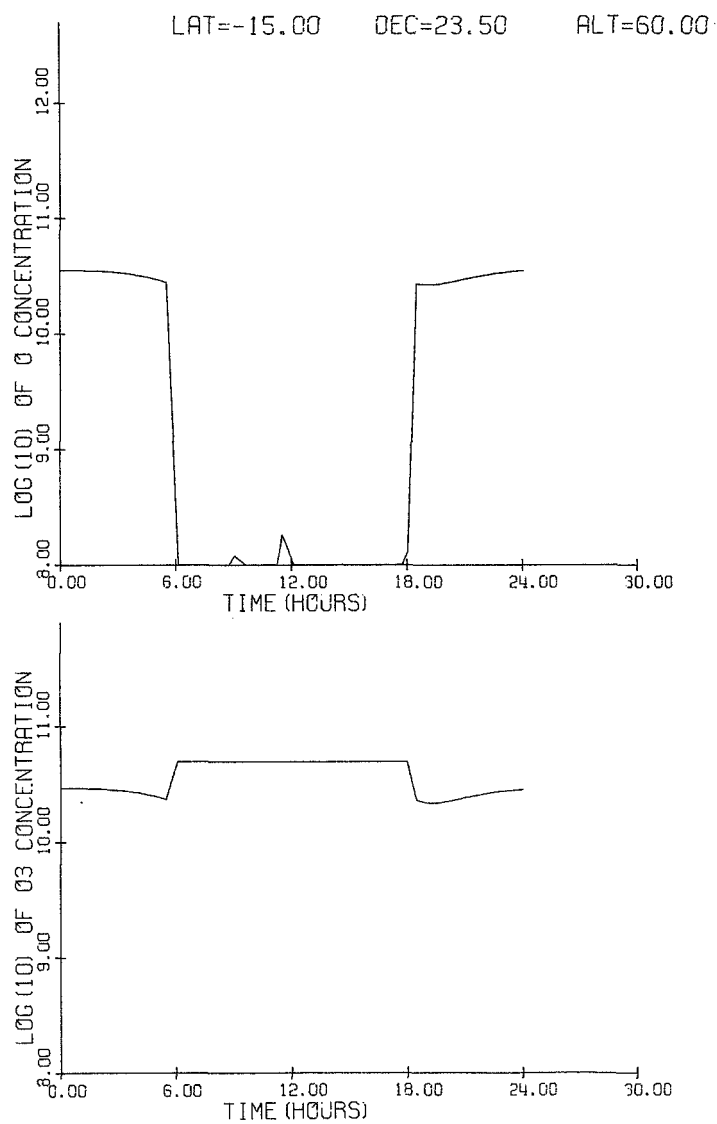


Figure A16 (continued)

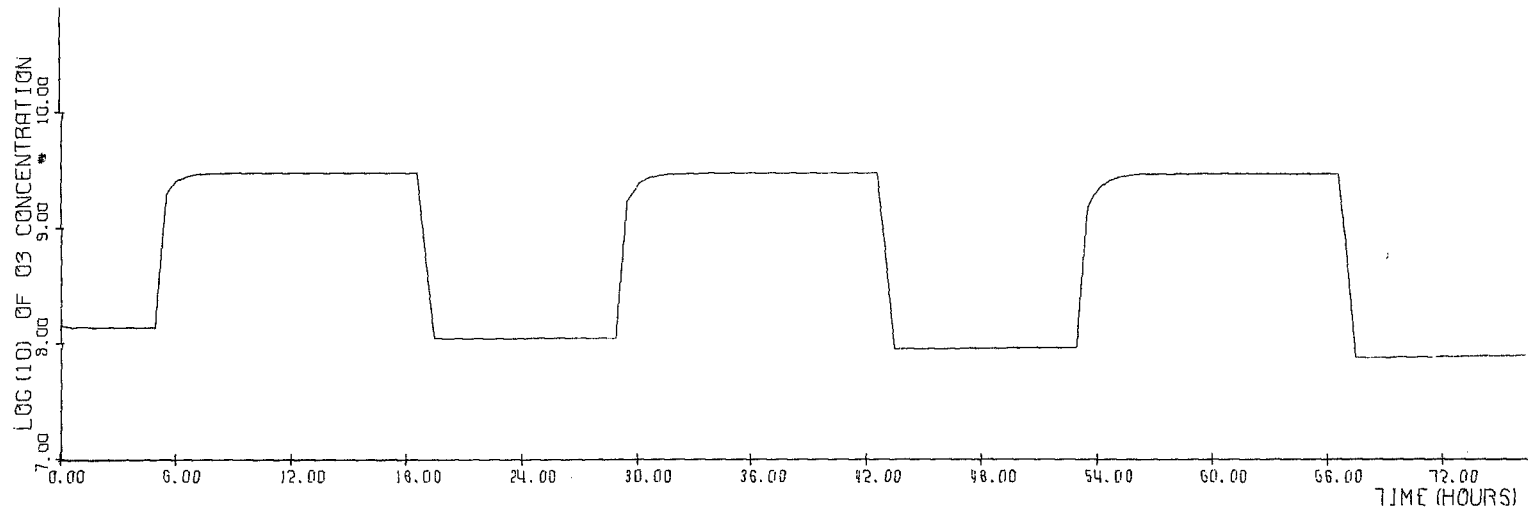
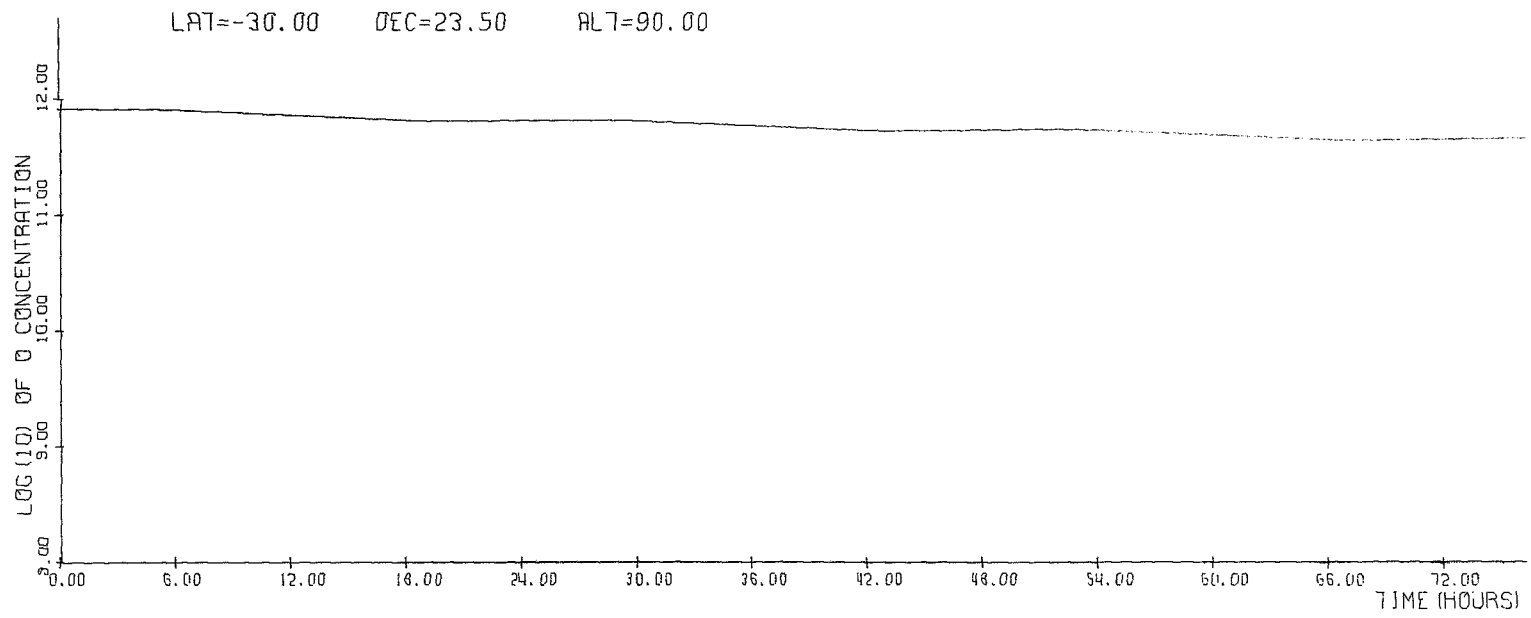


Figure A17

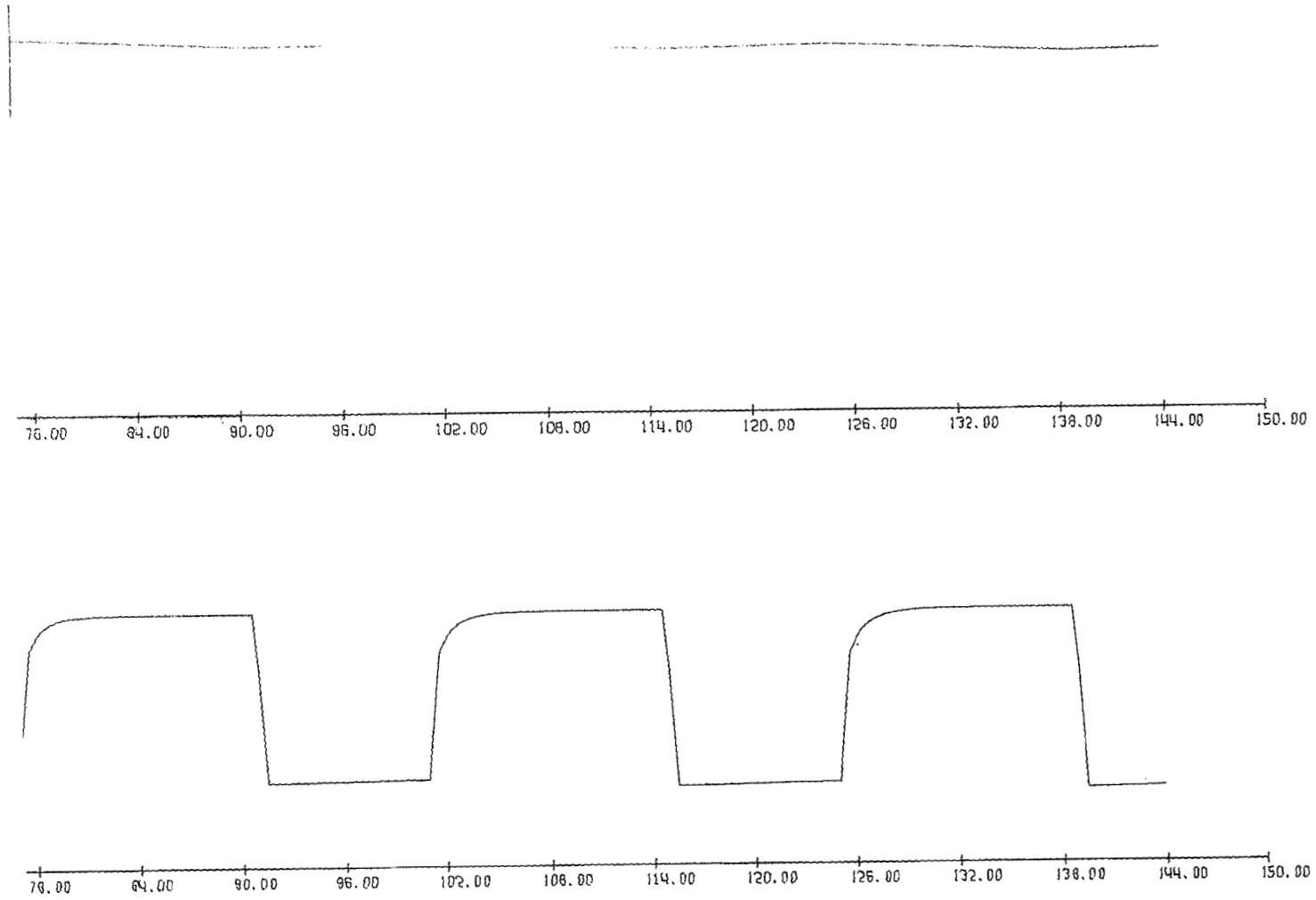


Figure A17 (continued)

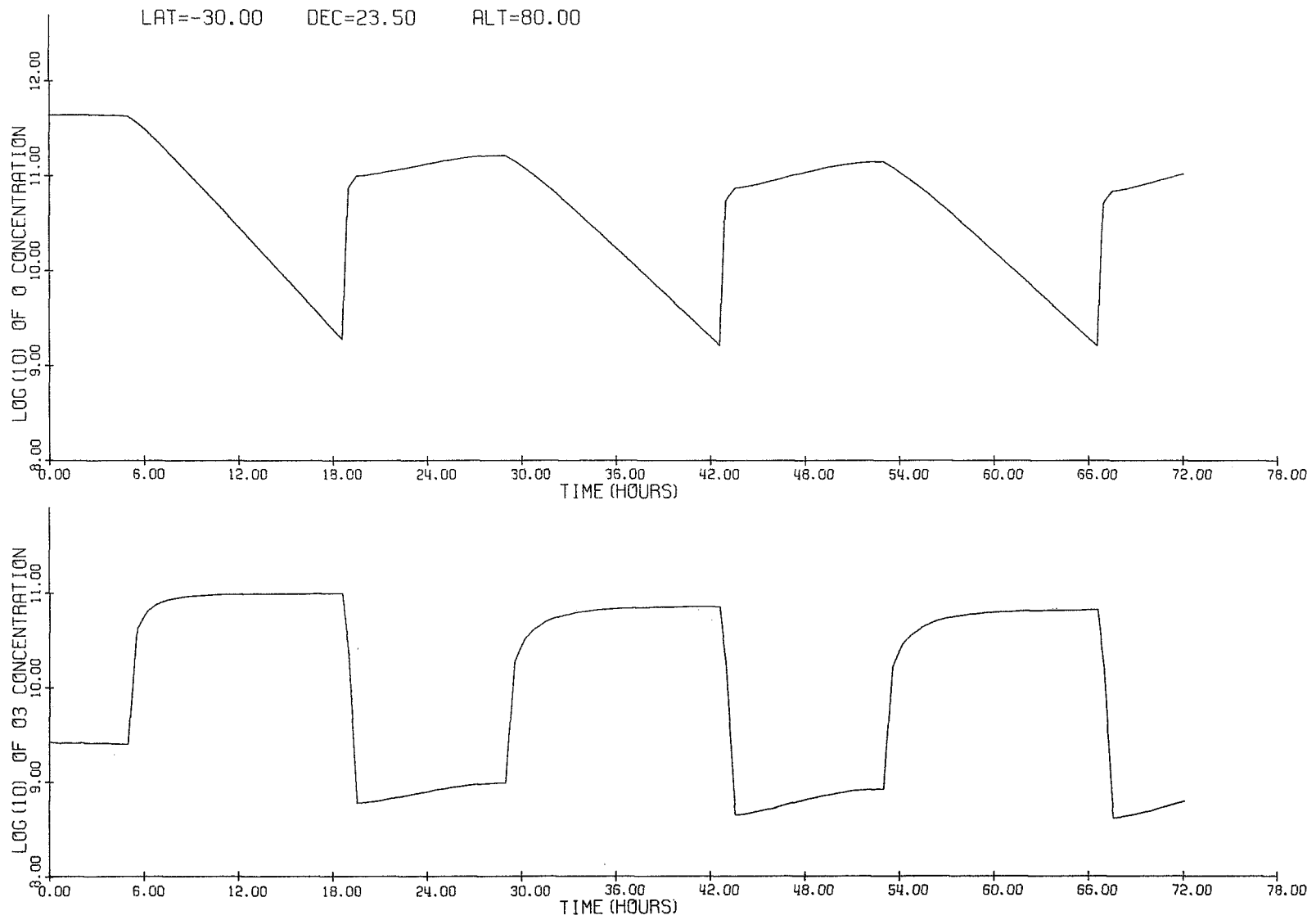


Figure A18

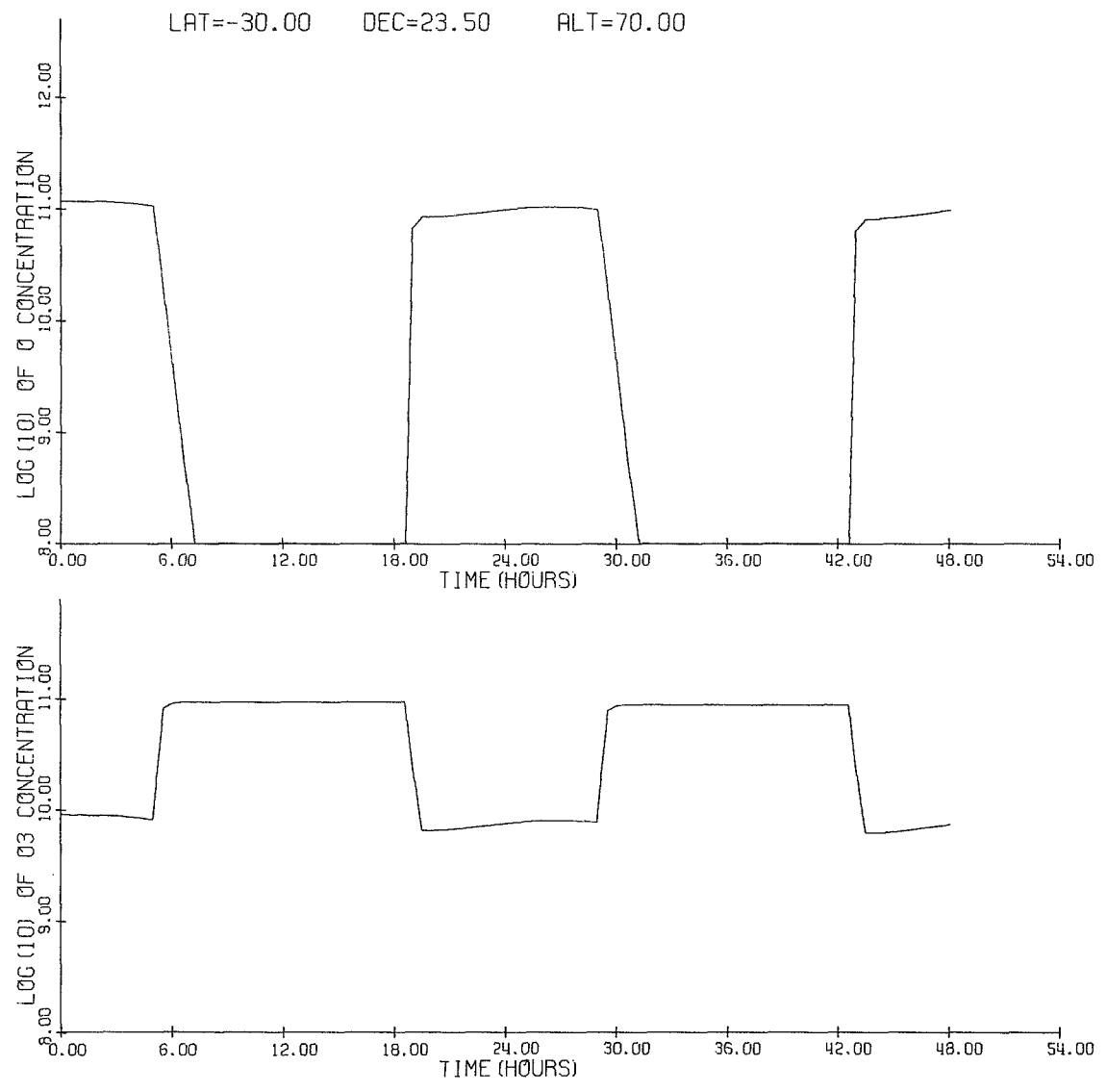


Figure A19

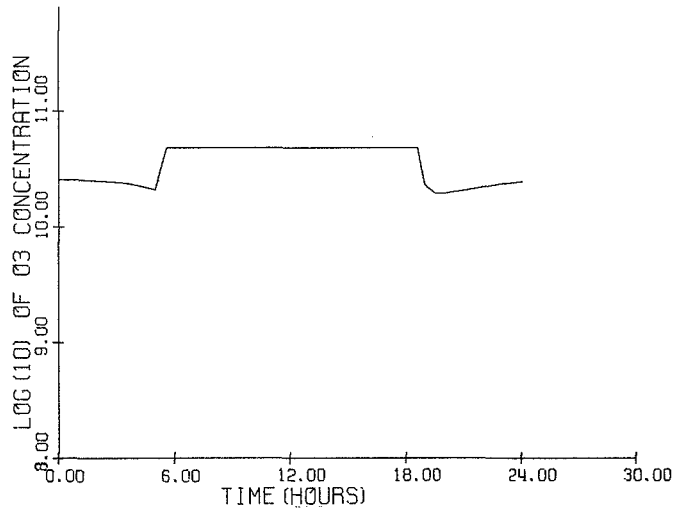
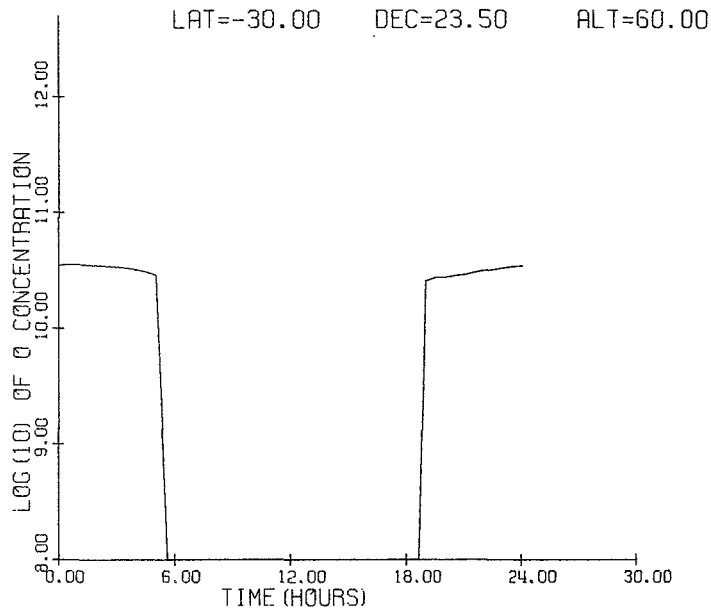


Figure A19 (continued)



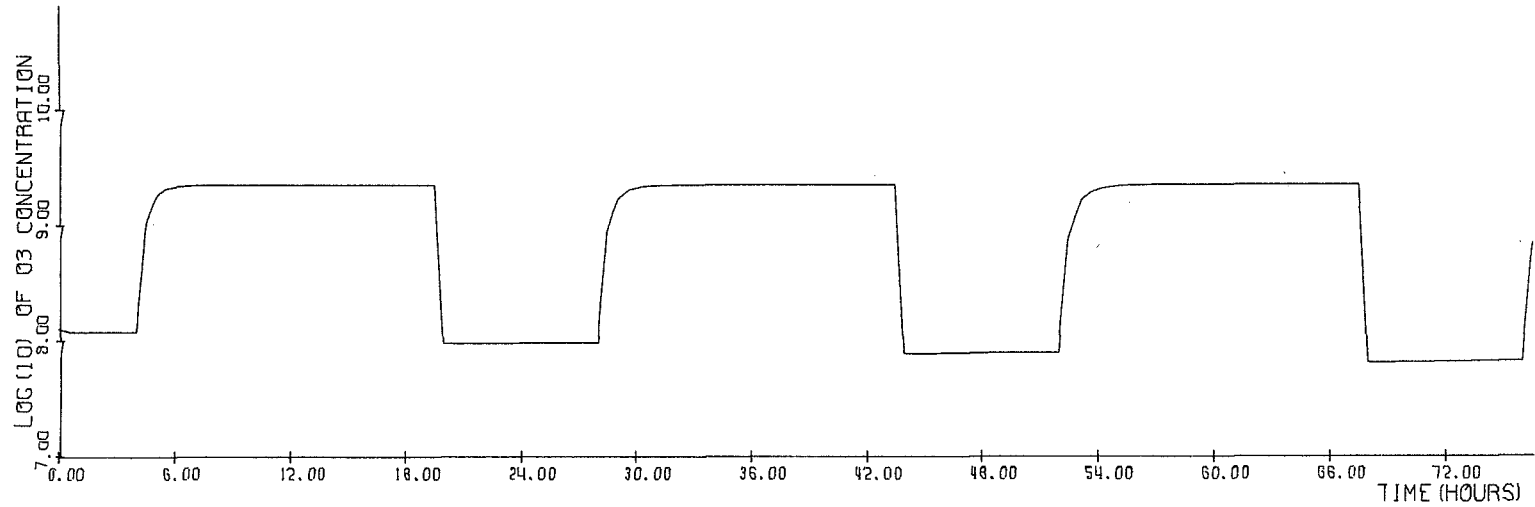
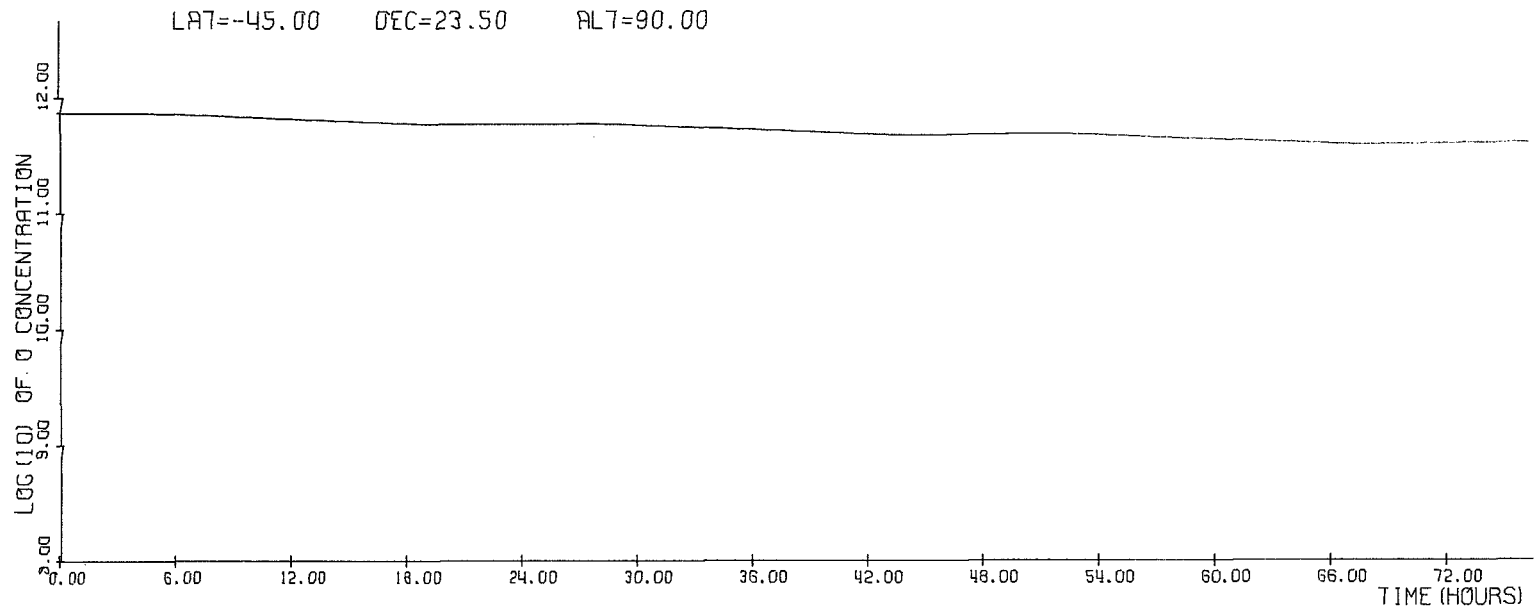


Figure A20

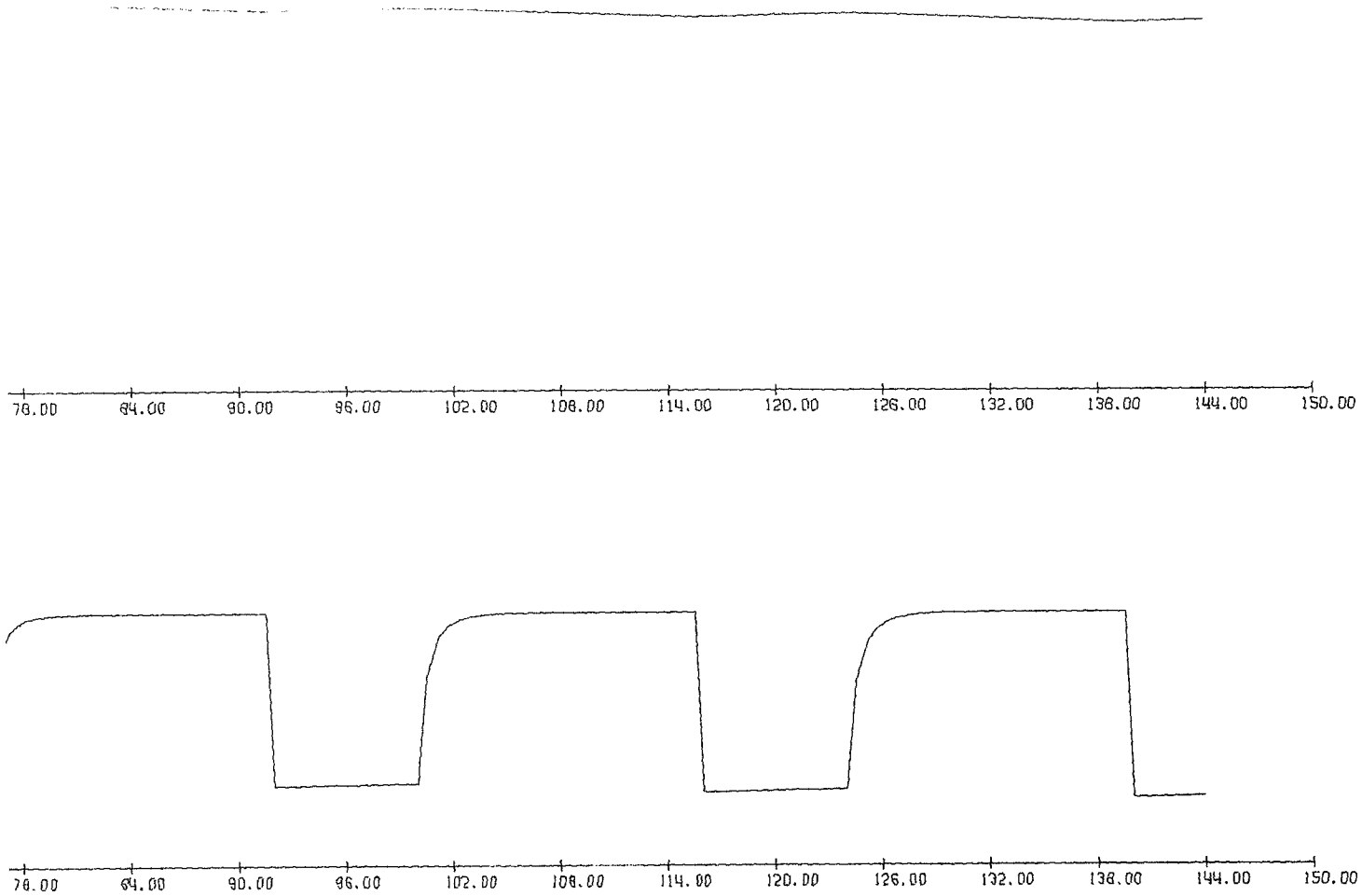


Figure A20 (continued)

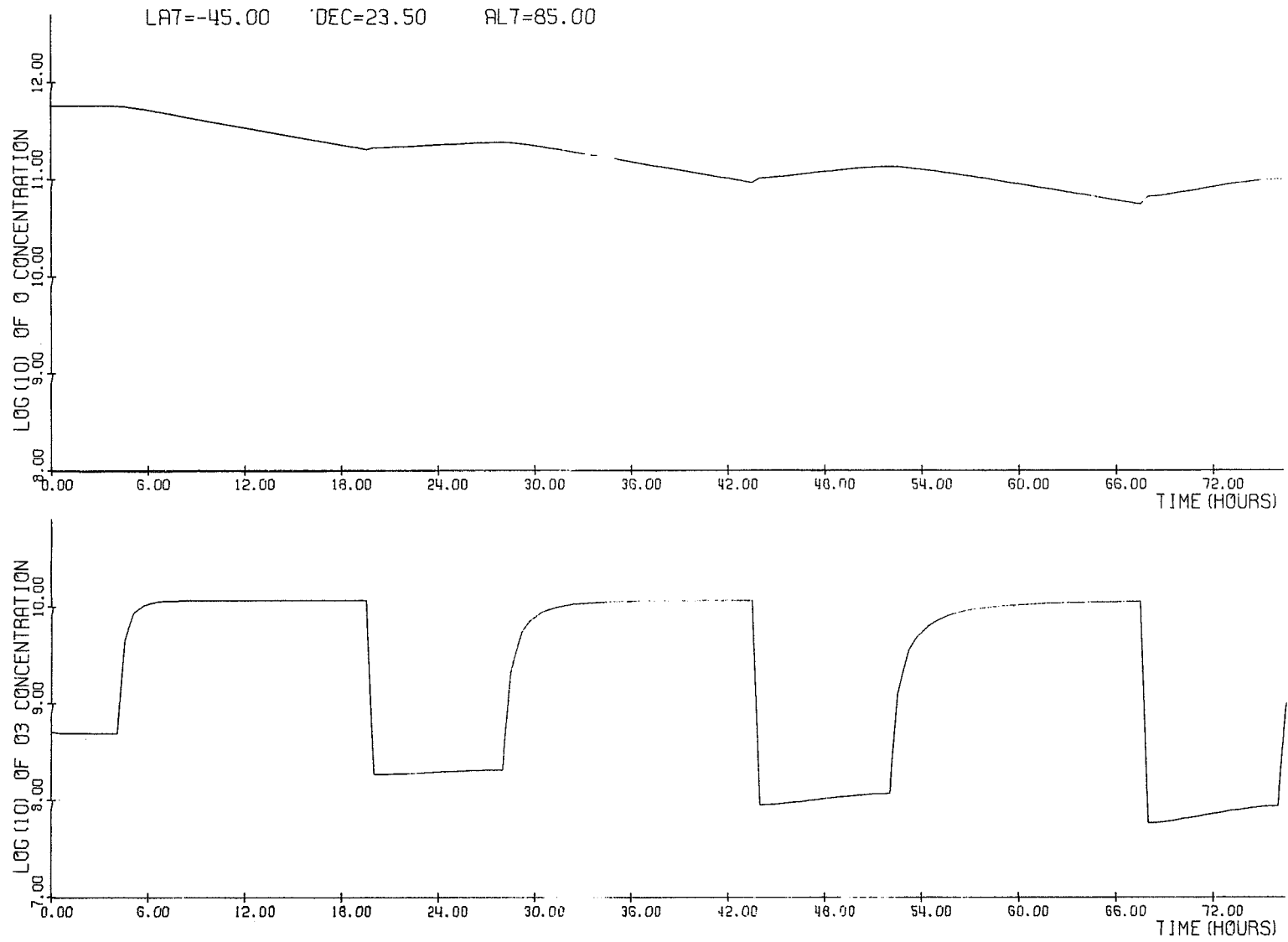


Figure A21

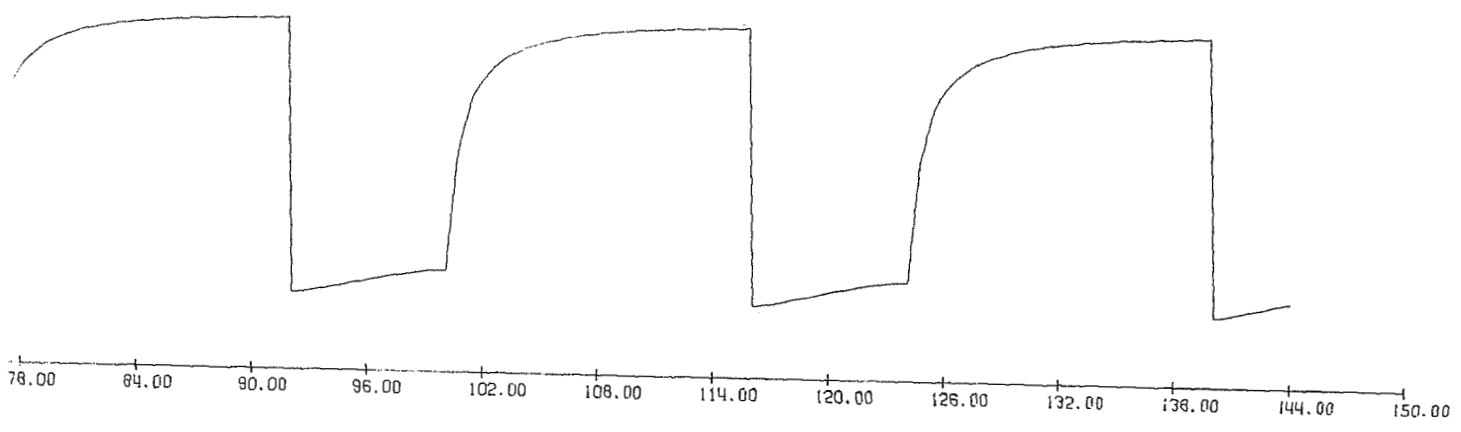
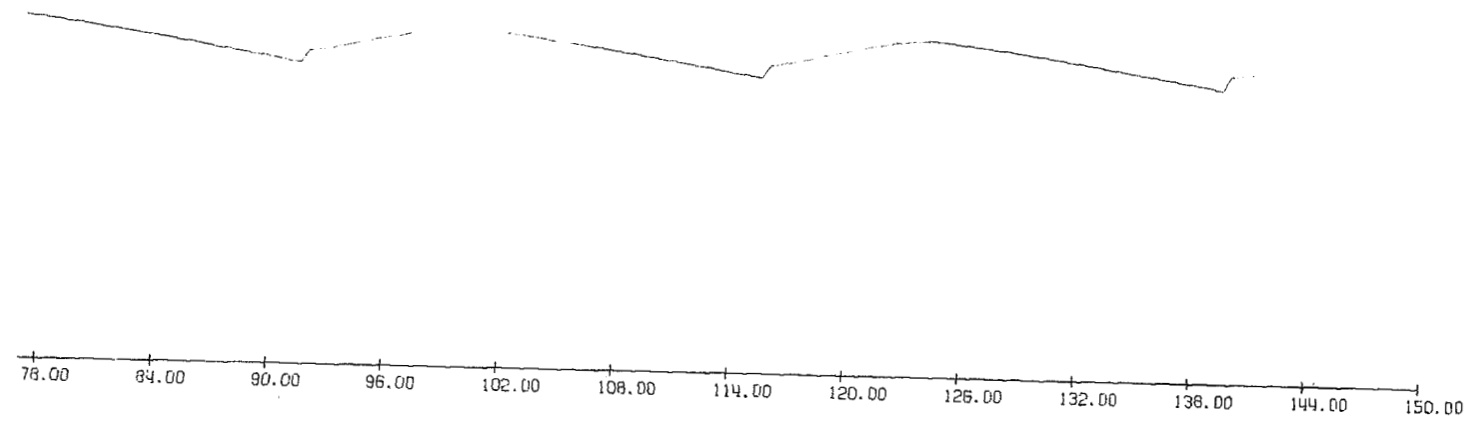


Figure A21 (continued)

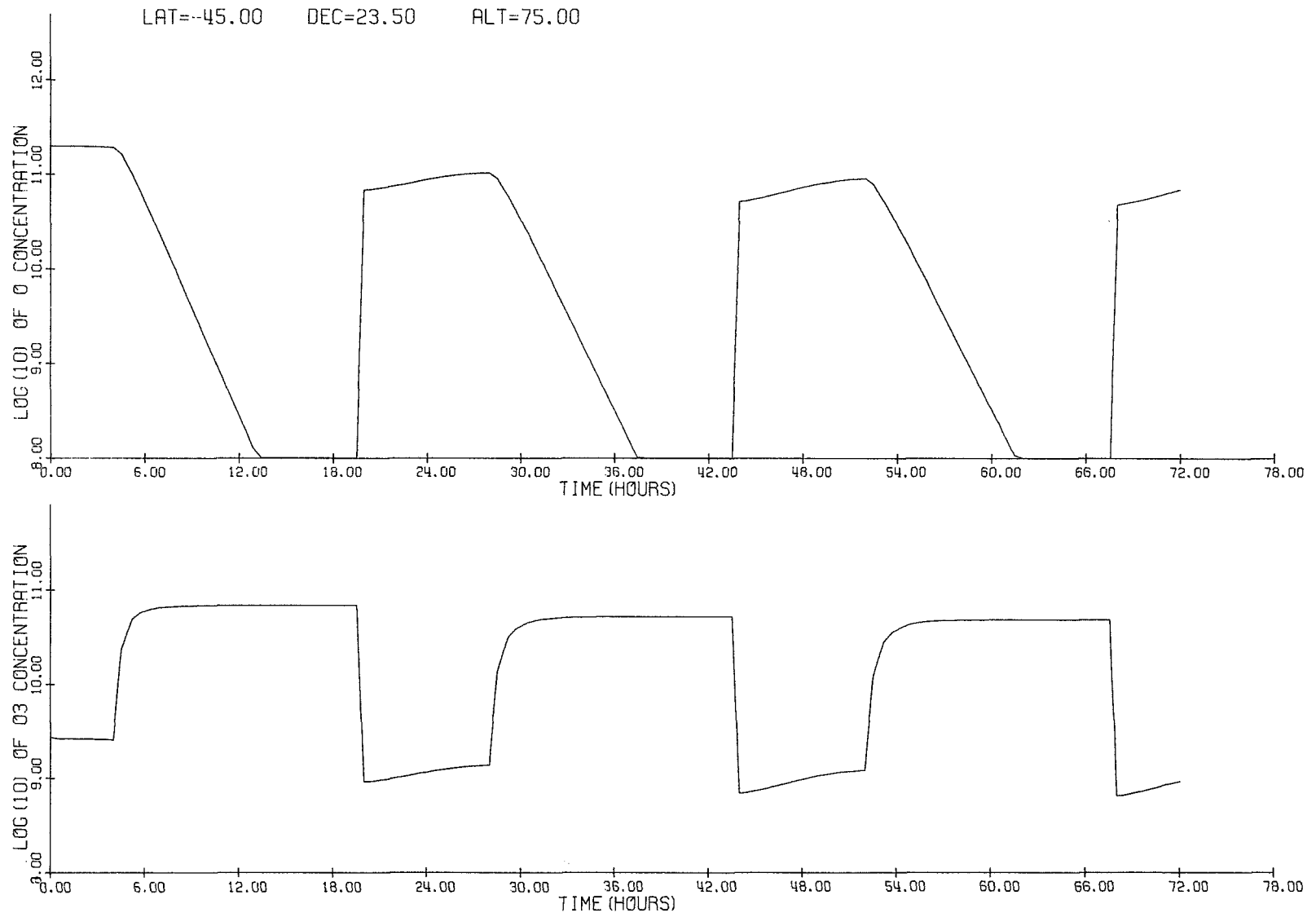


Figure A22

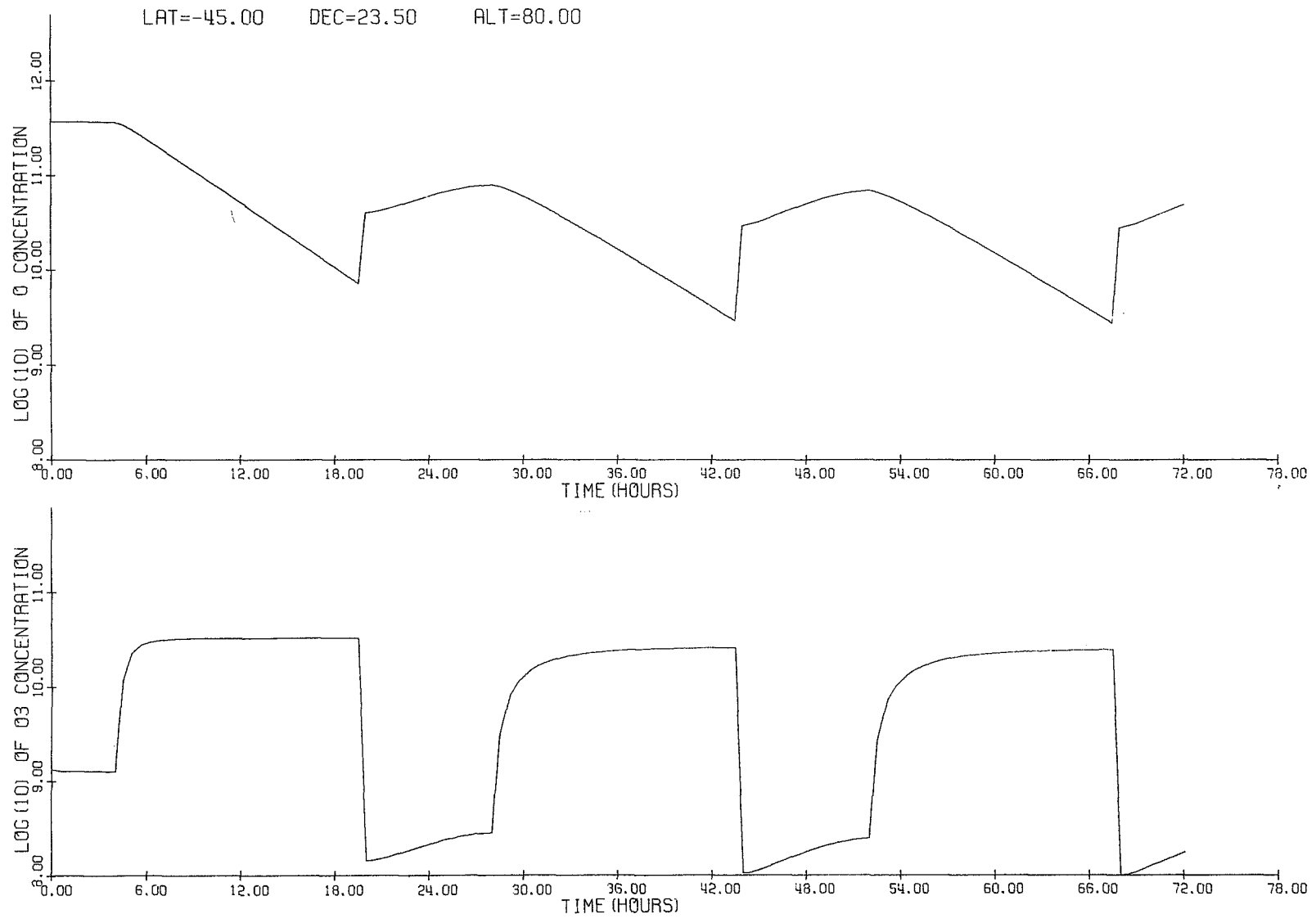


Figure A22 (continued)

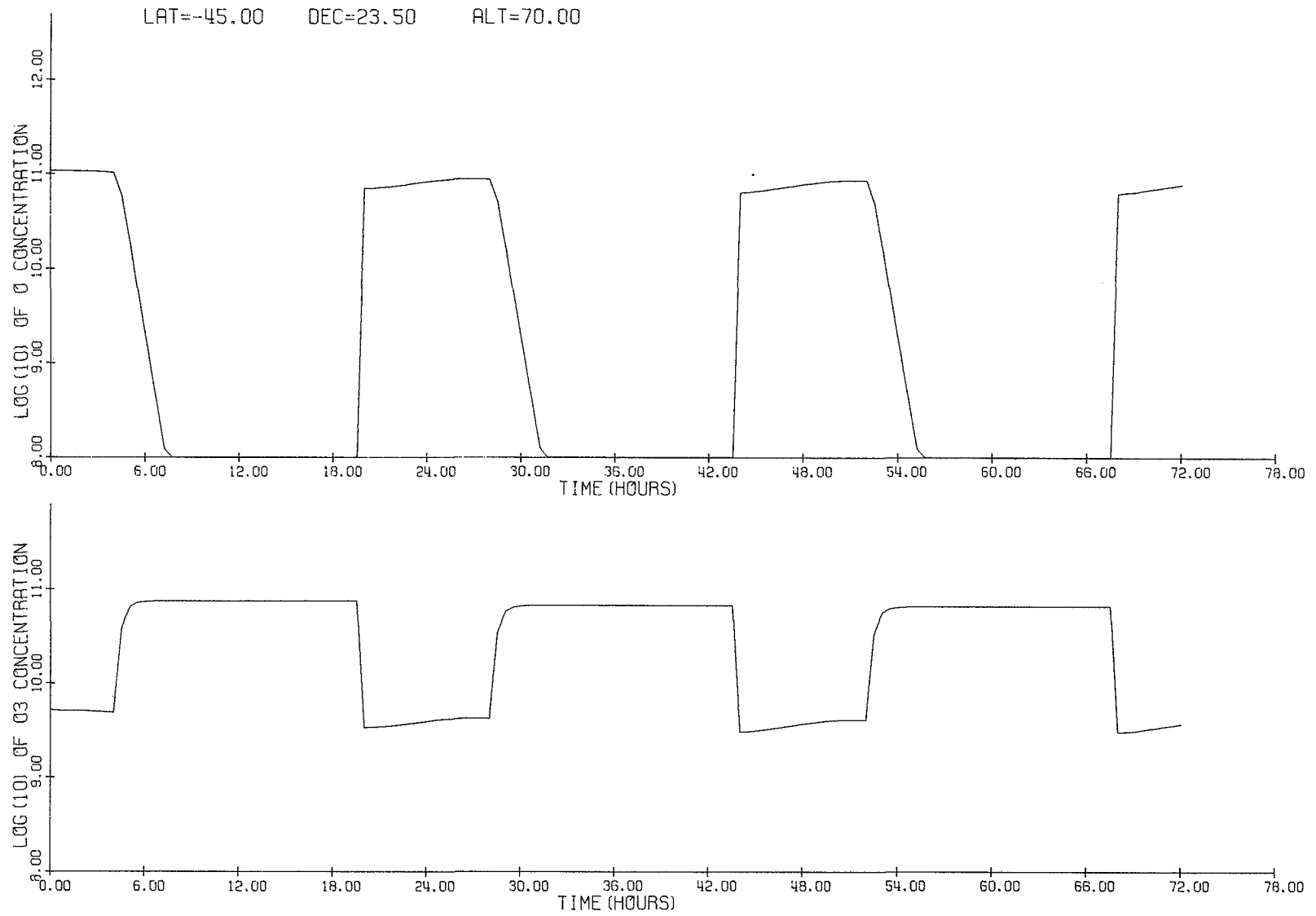


Figure A23

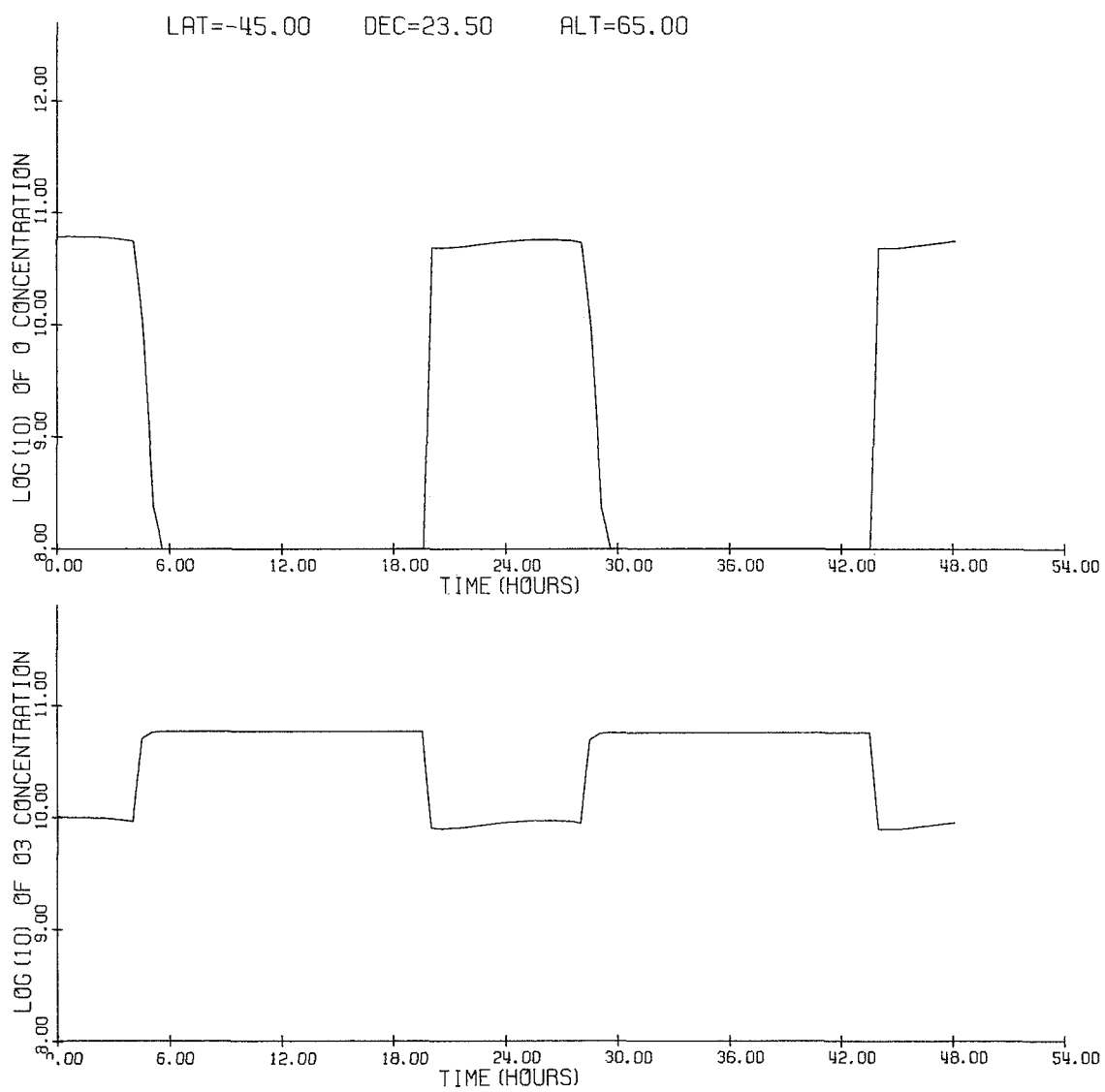


Figure A24



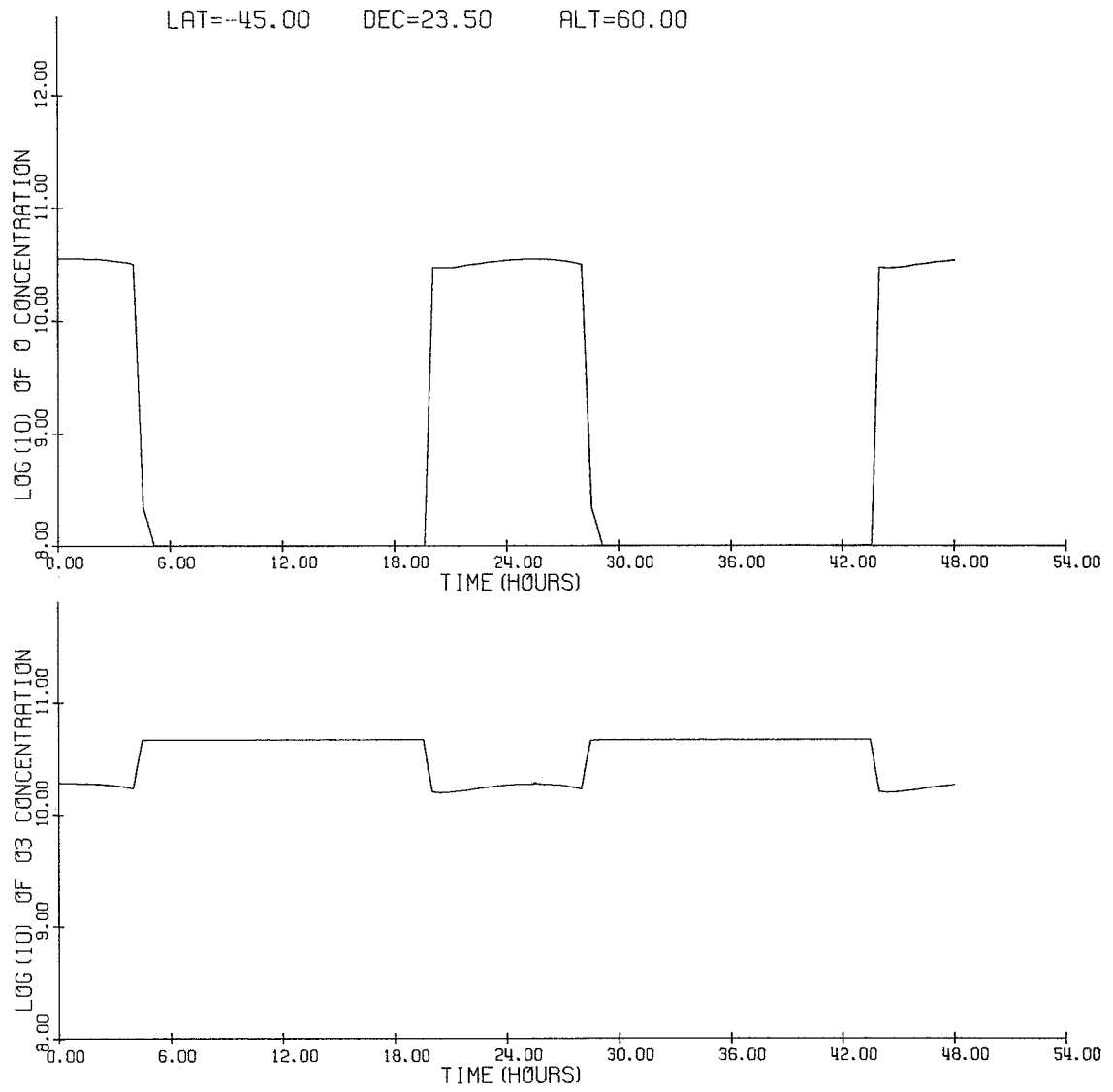


Figure A24 (continued)

Table A4

Table A5

LATITUDE= 45.0 DEG

ALTITUDE= 90.0 KM

LATITUDE= 45.0 DEG

ALTITUDE= 95.0 KM

DECLINATION=23.5 DEG

TIME(HR-MIN)	N(C)	N(C3)	TIME(HR-MIN)	N(C)	N(C3)
C- 7	4.786E 11	1.334E 08	C- 7	2.455E 11	5.129E 08
0-31	4.898E 11	1.343E 08	0-31	2.512E 11	5.200E 08
1- 7	4.898E 11	1.349E 08	1- 7	2.570E 11	5.284E 08
1-31	4.898E 11	1.358E 08	1-31	2.570E 11	5.358E 08
2- 7	4.898E 11	1.361E 08	2- 7	2.630E 11	5.432E 08
2-31	4.898E 11	1.368E 08	2-31	2.630E 11	5.508E 08
3- 7	5.012E 11	1.371E 08	3- 7	2.692E 11	5.572E 08
3-31	5.012E 11	1.380E 08	3-31	2.692E 11	5.649E 08
4- 7	5.012E 11	1.380E 08	4- 7	2.754E 11	5.715E 08
4-31	5.012E 11	1.387E 08	4-31	2.754E 11	5.781E 08
5- 7	5.012E 11	1.384E 08	5- 7	2.818E 11	5.821E 08
5-31	5.012E 11	1.390E 08	5-31	2.818E 11	5.888E 08
6- 7	5.012E 11	1.390E 08	6- 7	2.884E 11	5.916E 08
6-31	5.012E 11	1.396E 08	6-31	2.884E 11	5.970E 08
7- 7	5.012E 11	1.390E 08	7- 7	2.884E 11	5.957E 08
7-33	5.012E 11	1.390E 08	7-33	2.884E 11	5.957E 08
8- 7	5.012E 11	2.065E 09	8- 7	2.754E 11	8.933E 09
8-44	4.898E 11	4.603E 09	8-45	2.630E 11	2.042E 10
9-29	4.898E 11	7.096E 09	9-30	2.455E 11	3.162E 10
9-59	4.786E 11	8.472E 09	10- C	2.291E 11	3.802E 10
10-30	4.786E 11	9.660E 09	10-30	2.188E 11	4.467E 10
11- C	4.786E 11	1.072E 10	11- C	2.089E 11	4.898E 10
11-30	4.677E 11	1.148E 10	11-30	1.950E 11	5.370E 10
12- C	4.677E 11	1.230E 10	12- C	1.862E 11	5.888E 10
12-30	4.571E 11	1.288E 10	12-30	1.778E 11	6.166E 10
13- C	4.571E 11	1.349E 10	13- C	1.660E 11	6.607E 10
13-30	4.467E 11	1.413E 10	13-30	1.549E 11	6.918E 10
14- C	4.467E 11	1.445E 10	14- C	1.479E 11	7.079E 10
14-30	4.365E 11	1.479E 10	14-30	1.380E 11	7.413E 10
15- C	4.365E 11	1.514E 10	15- 0	1.318E 11	7.586E 10
15-30	4.266E 11	1.549E 10	15-30	1.230E 11	7.762E 10
16- C	4.266E 11	1.585E 10	16- C	1.148E 11	7.943E 10
16- 7	4.266E 11	1.585E 10	16- 7	1.148E 11	7.943E 10
16-30	4.365E 11	1.242E 08	16-31	1.905E 11	4.140E 08
17- 7	4.365E 11	1.211E 08	17- 7	1.905E 11	3.999E 08
17-37	4.365E 11	1.211E 08	17-37	1.950E 11	4.027E 08
18- 3	4.365E 11	1.213E 08	18- 3	1.950E 11	4.064E 08
18-37	4.365E 11	1.216E 08	18-37	1.995E 11	4.140E 08
19- 3	4.365E 11	1.216E 08	19- 3	1.995E 11	4.178E 08
19-37	4.467E 11	1.222E 08	19-37	2.042E 11	4.256E 08
20- 3	4.467E 11	1.225E 08	20- 3	2.089E 11	4.315E 08
20-33	4.467E 11	1.227E 08	20-33	2.138E 11	4.385E 08
21- 3	4.467E 11	1.233E 08	21- 7	2.138E 11	4.477E 08
21-37	4.467E 11	1.242E 08	21-31	2.188E 11	4.560E 08
22- 1	4.467E 11	1.253E 08	22- 7	2.239E 11	4.634E 08
22-37	4.571E 11	1.253E 08	22-31	2.291E 11	4.710E 08
23- 3	4.571E 11	1.262E 08	23- 7	2.291E 11	4.797E 08
23-37	4.571E 11	1.268E 08	23-31	2.344E 11	4.875E 08

Table A6

Table A7

LATITUDE= 45.C DEG

ALTITUDE= 80.0 KM

LATITUDE= 45.C DEG

ALTITUDE= 75.0 KM

DECLINATION=23.5 DEG

TIME(FR-MIN)	N(O)	N(C3)	TIME(FR-MIN)	N(O)	N(C3)
C- 7	2.570E 11	3.048E 09	C- 1	1.995E 11	8.166E 09
C-31	2.630E 11	3.119E 09	C-37	2.C42E 11	8.414E 09
1- 7	2.692E 11	3.177E 09	1- 3	2.C42E 11	8.531E 09
1-33	2.692E 11	3.228E 09	1-37	2.C89E 11	8.670E 09
2- 7	2.754E 11	3.289E 09	2- 3	2.C89E 11	8.770E 09
2-33	2.818E 11	3.342E 09	2-33	2.138E 11	8.851E 09
3- 3	2.884E 11	3.388E 09	3- 7	2.138E 11	8.933E 09
3-37	2.884E 11	3.436E 09	3-33	2.138E 11	8.954E 09
4- 1	2.951E 11	3.475E 09	4- 3	2.138E 11	8.974E 09
4-37	2.951E 11	3.508E 09	4-37	2.138E 11	8.974E 09
5- 1	2.951E 11	3.540E 09	5- 3	2.138E 11	8.933E 09
5-37	2.951E 11	3.540E 09	5-33	2.138E 11	8.872E 09
6- 1	2.951E 11	3.548E 09	6- 7	2.C89E 11	8.770E 09
6-33	2.951E 11	3.548E 09	6-33	2.C89E 11	8.690E 09
7- 7	2.951E 11	3.532E 09	7- 7	2.C42E 11	8.551E 09
7-33	2.951E 11	3.508E 09	7-33	2.C42E 11	8.414E 09
8- 7	2.455E 11	4.786E 10	8- 7	1.C23E 11	9.120E 10
8-45	1.778E 11	9.550E 10	8-37	4.C74E 10	1.349E 11
9-30	1.175E 11	1.318E 11	9-15	1.259E 10	1.514E 11
10- 0	8.710E 10	1.479E 11	9-45	4.909E 09	1.585E 11
10-30	6.607E 10	1.585E 11	10-15	1.892E 09	1.585E 11
11- 0	4.898E 10	1.660E 11	10-45	7.278E 08	1.585E 11
11-30	3.631E 10	1.738E 11	11-15	2.799E 08	1.585E 11
12- 0	2.692E 10	1.778E 11	11-45	1.C76E 08	1.585E 11
12-30	1.995E 10	1.820E 11	12-30	1.C00E 08	1.585E 11
13- 0	1.479E 10	1.820E 11	13- 0	1.C00E 08	1.585E 11
13-30	1.C96E 10	1.862E 11	13-30	1.C00E 08	1.585E 11
14- 0	8.147E 09	1.862E 11	14- 0	1.C00E 08	1.585E 11
14-30	6.C12E 09	1.862E 11	14-30	1.C00E 08	1.585E 11
15- 0	4.446E 09	1.862E 11	15- 0	1.C00E 08	1.585E 11
15-30	3.281E 09	1.862E 11	15-30	1.C00E 08	1.585E 11
16- 0	2.421E 09	1.862E 11	16- 0	1.C00E 08	1.585E 11
16- 7	2.244E 09	1.862E 11	16- 7	1.C00E 08	1.585E 11
16-31	1.862E 11	2.265E 09	16-31	1.514E 11	6.339E 09
17- 7	1.862E 11	2.234E 09	17- 1	1.514E 11	6.281E 09
17-37	1.905E 11	2.249E 09	17-37	1.514E 11	6.266E 09
18- 3	1.905E 11	2.265E 09	18- 3	1.514E 11	6.281E 09
18-37	1.950E 11	2.307E 09	18-33	1.514E 11	6.324E 09
19- 3	1.950E 11	2.333E 09	19- 7	1.549E 11	6.397E 09
19-37	1.995E 11	2.388E 09	19-33	1.549E 11	6.486E 09
20- 7	2.C42E 11	2.438E 09	20- 3	1.585E 11	6.592E 09
20-33	2.C89E 11	2.489E 09	20-37	1.622E 11	6.761E 09
21- 3	2.138E 11	2.547E 09	21- 3	1.660E 11	6.887E 09
21-37	2.188E 11	2.624E 09	21-33	1.698E 11	7.063E 09
22- 1	2.239E 11	2.685E 09	22- 7	1.738E 11	7.261E 09
22-37	2.344E 11	2.754E 09	22-33	1.778E 11	7.413E 09
23- 1	2.344E 11	2.825E 09	23- 7	1.820E 11	7.603E 09
23-37	2.455E 11	2.897E 09	23-33	1.862E 11	7.780E 09

Table A8

Table A9

LATITUDE= 45.0 DEG

ALTITUDE= 70.0 KM

LATITUDE= 45.0 DEG

ALTITUDE= 65.0 KM

DECLINATION=23.5 DEG

TIME(FR-MIN)	N(O)	N(O3)	TIME(FR-MIN)	N(O)	N(O3)
C- 1	1.047E 11	1.349E 10	C- 0	5.888E 10	2.089E 10
C-31	1.072E 11	1.349E 10	C-37	5.888E 10	2.089E 10
1- 7	1.096E 11	1.380E 10	1- 3	5.888E 10	2.089E 10
1-33	1.096E 11	1.413E 10	1-31	5.888E 10	2.089E 10
2- 3	1.096E 11	1.413E 10	2- 7	5.888E 10	2.042E 10
2-31	1.096E 11	1.413E 10	2-33	5.754E 10	2.042E 10
3- 1	1.096E 11	1.413E 10	3- 1	5.754E 10	2.042E 10
3-31	1.096E 11	1.413E 10	3-37	5.623E 10	1.995E 10
4- 1	1.096E 11	1.413E 10	4- 3	5.623E 10	1.995E 10
4-31	1.096E 11	1.380E 10	4-31	5.495E 10	1.950E 10
5- 1	1.072E 11	1.380E 10	5- 7	5.370E 10	1.905E 10
5-31	1.047E 11	1.349E 10	5-33	5.248E 10	1.862E 10
6- 1	1.023E 11	1.318E 10	6- 1	5.129E 10	1.820E 10
6-31	1.000E 11	1.288E 10	6-37	4.898E 10	1.738E 10
7- 1	9.772E 10	1.259E 10	7- 3	4.677E 10	1.660E 10
7-33	9.333E 10	1.202E 10	7-31	4.467E 10	1.585E 10
8- 3	1.660E 10	8.128E 10	8- 1	6.397E 08	5.623E 10
8-37	9.057E 08	9.333E 10	8-37	1.000E 08	5.754E 10
9- 7	1.000E 08	9.333E 10	9-15	1.000E 08	5.754E 10
9-45	1.000E 08	9.333E 10	9-45	1.000E 08	5.754E 10
10-15	1.000E 08	9.333E 10	10-15	1.000E 08	5.754E 10
10-45	1.000E 08	9.333E 10	10-45	1.000E 08	5.754E 10
11-30	1.000E 08	9.333E 10	11- 7	1.000E 08	5.754E 10
12- 0	1.000E 08	9.333E 10	11-45	1.000E 08	5.754E 10
12-30	1.000E 08	9.333E 10	12-15	1.000E 08	5.754E 10
13- 0	1.000E 08	9.333E 10	12-45	1.000E 08	5.754E 10
13-30	1.000E 08	9.333E 10	13- 7	1.000E 08	5.754E 10
14- 0	1.000E 08	9.333E 10	13-45	1.000E 08	5.754E 10
14-30	1.000E 08	9.333E 10	14-15	1.000E 08	5.754E 10
15- 0	1.000E 08	9.333E 10	14-45	1.000E 08	5.754E 10
15-30	1.000E 08	9.333E 10	15- 7	1.000E 08	5.754E 10
16- 0	1.000E 08	9.333E 10	15-45	1.000E 08	5.754E 10
16- 7	1.000E 08	9.333E 10	16- 7	1.000E 08	5.754E 10
16-31	8.128E 10	1.047E 10	16-31	4.074E 10	1.479E 10
17- 3	8.128E 10	1.023E 10	17- 3	3.981E 10	1.413E 10
17-37	7.943E 10	1.023E 10	17-33	3.981E 10	1.413E 10
18- 7	7.943E 10	1.023E 10	18- 3	4.074E 10	1.445E 10
18-33	8.128E 10	1.047E 10	18-31	4.169E 10	1.445E 10
19- 3	8.128E 10	1.047E 10	19- 7	4.365E 10	1.549E 10
19-31	8.318E 10	1.072E 10	19-31	4.467E 10	1.585E 10
20- 1	8.511E 10	1.096E 10	20- 7	4.677E 10	1.660E 10
20-31	8.710E 10	1.122E 10	20-37	4.786E 10	1.698E 10
21- 1	9.120E 10	1.148E 10	21- 3	5.012E 10	1.778E 10
21-31	9.333E 10	1.175E 10	21-31	5.129E 10	1.820E 10
22- 1	9.550E 10	1.230E 10	22- 7	5.248E 10	1.862E 10
22-31	9.772E 10	1.259E 10	22-33	5.370E 10	1.905E 10
23- 1	1.000E 11	1.288E 10	23- 1	5.495E 10	1.950E 10
23-31	1.023E 11	1.318E 10	23-37	5.623E 10	1.995E 10

Table A10

Table A11

LATITUDE= 45.0 DEG

ALTITUDE= 60.0 KM

LATITUDE= 30.0 DEG

ALTITUDE= 90.0 KM

DECLINATION=23.5 DEG

TIME(FR-MIN)	N(C)	N(C3)	TIME(FR-MIN)	N(C)	N(C3)
C- 0	3.311E 10	3.311E 1C	C- 1	4.677E 11	1.104E 08
C-3C	3.311E 10	3.311E 1C	C-37	4.786E 11	1.114E 08
1- 0	3.311E 10	3.311E 1C	1- 1	4.786E 11	1.117E 08
1-3C	3.311E 1C	3.236E 1C	1-37	4.786E 11	1.127E 08
2- 0	3.236E 10	3.236E 1C	2- 1	4.786E 11	1.130E 08
2-3C	3.236E 10	3.236E 1C	2-37	4.898E 11	1.138E 08
3- 0	3.236E 10	3.162E 1C	3- 1	4.898E 11	1.138E 08
3-3C	3.162E 10	3.162E 1C	3-37	4.898E 11	1.148E 08
4- 0	3.090E 10	3.090E 1C	4- 1	4.898E 11	1.146E 08
4-3C	3.020E 10	3.020E 1C	4-37	4.898E 11	1.153E 08
5- 0	2.951E 10	2.951E 1C	5- 1	4.898E 11	1.151E 08
5-3C	2.884E 10	2.884E 1C	5-37	4.898E 11	1.156E 08
6- 0	2.818E 10	2.818E 1C	6- 1	4.898E 11	1.153E 08
6-3C	2.692E 10	2.692E 1C	6-37	4.898E 11	1.159E 08
7- 0	2.570E 10	2.570E 1C	7- 0	4.898E 11	3.083E 08
7-21	2.455E 10	2.455E 10	7-44	4.898E 11	2.911E 09
8- 3	1.000E 08	4.677E 10	8-29	4.786E 11	4.842E 09
8-33	1.000E 08	4.677E 10	8-59	4.786E 11	5.834E 09
9- 7	1.000E 08	4.677E 1C	9-29	4.786E 11	6.637E 09
9-31	1.000E 08	4.677E 1C	9-59	4.677E 11	7.295E 09
10- 7	1.000E 08	4.677E 1C	10-30	4.677E 11	7.816E 09
10-31	1.000E 08	4.677E 1C	11- 0	4.571E 11	8.260E 09
11- 7	1.000E 08	4.677E 10	11-30	4.571E 11	8.610E 09
11-33	1.000E 08	4.677E 10	12- 0	4.571E 11	8.892E 09
12- 7	1.000E 08	4.677E 1C	12-30	4.467E 11	9.141E 09
12-33	1.000E 08	4.677E 10	13- 0	4.467E 11	9.333E 09
13- 7	1.000E 08	4.677E 1C	13-30	4.365E 11	9.484E 09
13-33	1.000E 08	4.677E 1C	14- 0	4.365E 11	9.616E 09
14- 7	1.000E 08	4.677E 10	14-30	4.266E 11	9.727E 09
14-33	1.000E 08	4.677E 10	15- 0	4.266E 11	9.817E 09
15- 3	1.000E 08	4.677E 10	15-30	4.266E 11	9.886E 09
15-37	1.000E 08	4.677E 10	16- 0	4.169E 11	9.954E 09
16- 1	1.000E 08	4.677E 10	16-30	4.169E 11	1.000E 10
16-31	2.291E 10	2.291E 10	17- 0	4.074E 11	1.000E 10
17- 1	2.291E 10	2.291E 10	17- 0	4.074E 11	1.000E 10
17-31	2.291E 10	2.291E 10	17-33	4.169E 11	9.852E 07
18- 1	2.344E 10	2.344E 1C	18- 3	4.169E 11	9.860E 07
18-31	2.455E 10	2.455E 1C	18-37	4.169E 11	9.884E 07
19- 1	2.570E 10	2.570E 1C	19- 1	4.169E 11	9.877E 07
19-31	2.630E 10	2.630E 10	19-37	4.266E 11	9.935E 07
20- 1	2.754E 10	2.754E 1C	20- 1	4.266E 11	9.936E 07
20-31	2.884E 10	2.818E 1C	20-37	4.266E 11	1.000E 08
21- 1	2.951E 10	2.951E 1C	21- 1	4.266E 11	1.002E 08
21-31	3.020E 10	3.020E 1C	21-37	4.266E 11	1.012E 08
22- 1	3.090E 10	3.090E 1C	22- 1	4.365E 11	1.014E 08
22-31	3.162E 10	3.162E 1C	22-37	4.365E 11	1.023E 08
23- 1	3.162E 10	3.162E 1C	23- 1	4.365E 11	1.026E 08
23-31	3.236E 10	3.236E 10	23-37	4.467E 11	1.038E 08

Table A12

Table A13

LATTITUDE= 30.0 DEG			ALTITUDE= 80.0 KM			LATTITUDE= 30.0 DEG			ALTITUDE= 70.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)						
0- 7	2.089E 11	1.858E 09	0- 3	1.148E 11	1.202E 10						
0-31	2.188E 11	1.910E 09	0-31	1.175E 11	1.230E 10						
1- 7	2.239E 11	1.963E 09	1- 1	1.175E 11	1.259E 10						
1-31	2.291E 11	2.018E 09	1-31	1.202E 11	1.259E 10						
2- 7	2.344E 11	2.065E 09	2- 1	1.202E 11	1.288E 10						
2-33	2.399E 11	2.109E 09	2-31	1.202E 11	1.288E 10						
3- 7	2.455E 11	2.153E 09	3- 1	1.202E 11	1.288E 10						
3-33	2.512E 11	2.188E 09	3-33	1.202E 11	1.288E 10						
4- 7	2.512E 11	2.234E 09	4- 3	1.202E 11	1.259E 10						
4-33	2.570E 11	2.254E 09	4-37	1.175E 11	1.259E 10						
5- 7	2.570E 11	2.280E 09	5- 3	1.175E 11	1.230E 10						
5-33	2.630E 11	2.291E 09	5-37	1.148E 11	1.202E 10						
6- 7	2.630E 11	2.301E 09	6- 7	1.096E 11	1.175E 10						
6-33	2.630E 11	2.296E 09	6-33	1.072E 11	1.148E 10						
7- 0	2.570E 11	6.067E 09	7- 0	8.710E 10	2.818E 10						
7-45	1.950E 11	5.248E 10	7-33	8.260E 09	9.550E 10						
8-30	1.413E 11	8.318E 10	8- 7	6.887E 08	1.000E 11						
9- 0	1.148E 11	9.550E 10	8-45	1.000E 08	1.023E 11						
9-30	9.120E 10	1.072E 11	9-15	1.000E 08	1.023E 11						
10- 0	7.244E 10	1.148E 11	9-45	1.000E 08	1.023E 11						
10-30	5.754E 10	1.202E 11	10-30	1.000E 08	1.023E 11						
11- 0	4.571E 10	1.230E 11	11- 0	1.000E 08	1.023E 11						
11-30	3.548E 10	1.288E 11	11-30	1.000E 08	1.023E 11						
12- 0	2.818E 10	1.288E 11	12- 0	1.000E 08	1.023E 11						
12-30	2.188E 10	1.318E 11	12-30	1.000E 08	1.023E 11						
13- 0	1.738E 10	1.349E 11	13- 0	1.000E 08	1.023E 11						
13-30	1.349E 10	1.349E 11	13-30	1.000E 08	1.023E 11						
14- 0	1.047E 10	1.349E 11	14- 0	1.000E 08	1.023E 11						
14-30	8.241E 09	1.349E 11	14-30	1.000E 08	1.023E 11						
15- 0	6.442E 09	1.380E 11	15- 0	1.000E 08	1.023E 11						
15-30	5.047E 09	1.380E 11	15-30	1.000E 08	1.023E 11						
16- 0	3.945E 09	1.380E 11	16- 0	1.000E 08	1.023E 11						
16-30	3.083E 09	1.380E 11	16-30	1.000E 08	1.023E 11						
17- 0	2.404E 09	1.380E 11	17- 0	1.000E 08	1.023E 11						
17- 0	2.399E 09	1.380E 11	17- 0	1.000E 08	1.023E 11						
17-33	1.380E 11	1.230E 09	17-37	8.913E 10	9.462E 09						
18- 3	1.413E 11	1.242E 09	18- 3	8.710E 10	9.376E 09						
18-37	1.445E 11	1.268E 09	18-33	8.710E 10	9.376E 09						
19- 1	1.479E 11	1.279E 09	19- 3	8.913E 10	9.462E 09						
19-37	1.514E 11	1.334E 09	19-33	8.913E 10	9.616E 09						
20- 3	1.549E 11	1.371E 09	20- 3	9.120E 10	9.795E 09						
20-37	1.622E 11	1.422E 09	20-31	9.333E 10	1.000E 10						
21- 3	1.660E 11	1.466E 09	21- 1	9.772E 10	1.023E 10						
21-37	1.738E 11	1.528E 09	21-31	1.000E 11	1.072E 10						
22- 3	1.778E 11	1.570E 09	22- 1	1.023E 11	1.096E 10						
22-37	1.862E 11	1.637E 09	22-31	1.047E 11	1.122E 10						
23- 3	1.905E 11	1.683E 09	23- 1	1.096E 11	1.148E 10						
23-37	1.995E 11	1.746E 09	23-33	1.122E 11	1.175E 10						

Table A14

Table A15

LATITUDE= 30.0 DEG			ALTIUDE= 60.0 KM			LATITUDE= 15.0 DEG			ALTIUDE= 90.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)						
0-0	3.548E 10	3.162E 10	0-7	4.467E 11	9.822E 07						
0-31	3.631E 10	3.162E 10	0-31	4.467E 11	9.846E 07						
1-1	3.548E 10	3.162E 10	1-7	4.571E 11	9.945E 07						
1-30	3.548E 10	3.162E 10	1-31	4.571E 11	9.962E 07						
2-1	3.548E 10	3.090E 10	2-7	4.571E 11	1.005E 08						
2-31	3.548E 10	3.090E 10	2-31	4.571E 11	1.007E 08						
3-1	3.467E 10	3.090E 10	3-7	4.677E 11	1.014E 08						
3-31	3.388E 10	3.020E 10	3-31	4.677E 11	1.014E 08						
4-0	3.388E 10	2.951E 10	4-7	4.677E 11	1.021E 08						
4-31	3.236E 10	2.884E 10	4-31	4.677E 11	1.019E 08						
5-1	3.162E 10	2.818E 10	5-7	4.677E 11	1.026E 08						
5-31	3.090E 10	2.692E 10	5-31	4.677E 11	1.023E 08						
6-0	2.951E 10	2.630E 10	6-7	4.677E 11	1.028E 08						
6-31	2.818E 10	2.455E 10	6-30	4.677E 11	3.581E 08						
7-0	6.123E 09	4.266E 10	7-14	4.677E 11	2.564E 09						
7-37	1.000E 08	4.898E 10	7-44	4.571E 11	3.664E 09						
8-3	1.000E 08	4.898E 10	8-29	4.571E 11	4.875E 09						
8-37	1.000E 08	4.898E 10	8-59	4.571E 11	5.470E 09						
9-3	1.000E 08	4.898E 10	9-29	4.467E 11	5.929E 09						
9-37	1.000E 08	4.898E 10	9-59	4.467E 11	6.281E 09						
10-7	1.000E 08	4.898E 10	10-30	4.365E 11	6.561E 09						
10-33	1.000E 08	4.898E 10	11-0	4.365E 11	6.792E 09						
11-7	1.000E 08	4.898E 10	11-30	4.365E 11	6.966E 09						
11-37	1.000E 08	4.898E 10	12-0	4.266E 11	7.096E 09						
12-1	1.000E 08	4.898E 10	12-30	4.266E 11	7.211E 09						
12-33	1.000E 08	4.898E 10	13-0	4.169E 11	7.295E 09						
13-7	1.000E 08	4.898E 10	13-30	4.169E 11	7.362E 09						
13-33	1.000E 08	4.898E 10	14-0	4.169E 11	7.430E 09						
14-7	1.000E 08	4.898E 10	14-30	4.074E 11	7.464E 09						
14-31	1.000E 08	4.898E 10	15-0	4.074E 11	7.499E 09						
15-7	1.000E 08	4.898E 10	15-30	3.981E 11	7.534E 09						
15-33	1.000E 08	4.898E 10	16-0	3.981E 11	7.551E 09						
16-7	1.000E 08	4.898E 10	16-30	3.981E 11	7.568E 09						
16-33	1.000E 08	4.898E 10	17-0	3.890E 11	7.586E 09						
17-0	1.000E 08	4.898E 10	17-15	3.890E 11	7.586E 09						
17-33	2.512E 10	2.188E 10	17-30	3.890E 11	7.603E 09						
18-1	2.512E 10	2.188E 10	18-3	3.981E 11	8.631E 07						
18-31	2.570E 10	2.239E 10	18-33	3.981E 11	8.641E 07						
19-3	2.692E 10	2.344E 10	19-7	3.981E 11	8.667E 07						
19-31	2.754E 10	2.455E 10	19-31	3.981E 11	8.652E 07						
20-1	2.884E 10	2.570E 10	20-7	3.981E 11	8.723E 07						
20-31	3.020E 10	2.630E 10	20-31	3.981E 11	8.712E 07						
21-1	3.162E 10	2.754E 10	21-7	3.981E 11	8.800E 07						
21-33	3.236E 10	2.818E 10	21-31	4.074E 11	8.796E 07						
22-3	3.311E 10	2.951E 10	22-7	4.074E 11	8.902E 07						
22-33	3.388E 10	3.020E 10	22-31	4.074E 11	8.902E 07						
23-3	3.467E 10	3.020E 10	23-7	4.169E 11	9.022E 07						
23-33	3.467E 10	3.090E 10	23-31	4.169E 11	9.022E 07						

Table A16

Table A17

LATITUDE= 15.0 DEG

ALTITUDE= 80.0 KM

LATITUDE= 15.0 DEG

ALTITUDE= 70.0 KM

DECLINATION=23.5 DEG

TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)
C- 3	1.778E 11	1.343E 09	C- 3	1.175E 11	1.096E 10
C-37	1.862E 11	1.396E 09	C-31	1.175E 11	1.096E 10
1- 3	1.905E 11	1.439E 09	1- 1	1.202E 11	1.122E 10
1-37	1.995E 11	1.486E 09	1-33	1.230E 11	1.148E 10
2- 3	2.042E 11	1.528E 09	2- 3	1.230E 11	1.148E 10
2-37	2.089E 11	1.570E 09	2-31	1.230E 11	1.148E 10
3- 3	2.138E 11	1.607E 09	3- 1	1.230E 11	1.148E 10
3-37	2.188E 11	1.644E 09	3-33	1.230E 11	1.148E 10
4- 3	2.239E 11	1.671E 09	4- 3	1.230E 11	1.148E 10
4-37	2.291E 11	1.702E 09	4-31	1.202E 11	1.122E 10
5- 7	2.291E 11	1.718E 09	5- 3	1.175E 11	1.096E 10
5-33	2.291E 11	1.726E 09	5-33	1.148E 11	1.072E 10
6- 3	2.291E 11	1.730E 09	6- 3	1.122E 11	1.047E 10
6-30	2.239E 11	5.998E 09	6-30	8.511E 10	3.311E 10
7-15	1.778E 11	4.169E 10	7- 7	8.110E 09	9.550E 10
7-45	1.479E 11	5.754E 10	7-37	1.112E 09	1.000E 11
8-30	1.122E 11	7.413E 10	8-15	1.000E 08	1.023E 11
9- 0	9.120E 10	8.318E 10	8-45	1.000E 08	1.023E 11
9-30	7.413E 10	8.913E 10	9-15	1.000E 08	1.023E 11
10- 0	6.026E 10	9.333E 10	9-45	1.000E 08	1.023E 11
10-30	4.858E 10	9.772E 10	10-30	1.000E 08	1.023E 11
11- 0	3.981E 10	1.000E 11	11- 0	1.000E 08	1.023E 11
11-30	3.236E 10	1.023E 11	11-30	1.000E 08	1.023E 11
12- 0	2.570E 10	1.047E 11	12- 0	1.000E 08	1.023E 11
12-30	2.089E 10	1.047E 11	12-30	1.000E 08	1.023E 11
13- 0	1.698E 10	1.072E 11	13- 0	1.000E 08	1.023E 11
13-30	1.349E 10	1.072E 11	13-30	1.000E 08	1.023E 11
14- 0	1.096E 10	1.072E 11	14- 0	1.000E 08	1.023E 11
14-30	8.710E 09	1.072E 11	14-30	1.000E 08	1.023E 11
15- 0	6.998E 09	1.096E 11	15- 0	1.000E 08	1.023E 11
15-30	5.623E 09	1.096E 11	15-30	1.000E 08	1.023E 11
16- 0	4.508E 09	1.096E 11	16- 0	1.000E 08	1.023E 11
16-30	3.622E 09	1.096E 11	16-30	1.000E 08	1.023E 11
17- 0	2.904E 09	1.096E 11	17- 0	1.000E 08	1.023E 11
17-15	2.600E 09	1.096E 11	17-15	1.000E 08	1.023E 11
17-31	2.312E 09	1.096E 11	17-31	1.000E 08	1.023E 11
18- 3	1.122E 11	8.433E 08	18- 7	8.913E 10	8.453E 09
18-33	1.148E 11	8.551E 08	18-33	8.913E 10	8.414E 09
19- 7	1.175E 11	8.810E 08	19- 3	8.913E 10	8.433E 09
19-31	1.202E 11	8.892E 08	19-33	9.120E 10	8.531E 09
20- 7	1.259E 11	9.441E 08	20- 1	9.333E 10	8.670E 09
20-33	1.288E 11	9.795E 08	20-33	9.550E 10	8.892E 09
21- 7	1.349E 11	1.026E 09	21- 3	9.772E 10	9.120E 09
21-33	1.413E 11	1.064E 09	21-33	1.000E 11	9.397E 09
22- 7	1.479E 11	1.119E 09	22- 3	1.023E 11	9.660E 09
22-33	1.549E 11	1.159E 09	22-31	1.072E 11	9.954E 09
23- 7	1.622E 11	1.216E 09	23- 1	1.096E 11	1.023E 10
23-33	1.660E 11	1.256E 09	23-33	1.122E 11	1.047E 10



Table A18

Table A19

LATITUDE= 15.0 DEG

ALTITUDE= 60.0 KM

LATITUDE= 0.0 DEG

ALTITUDE= 90.0 KM

DECLINATION=23.5 DEG

TIME(HR-MIN)	N(C)	N(C3)	TIME(HR-MIN)	N(C)	N(C3)
0-0	3.715E 10	3.090E 10	0-7	4.074E 11	8.935E 07
0-31	3.715E 10	3.090E 10	0-33	4.074E 11	8.973E 07
1-1	3.715E 10	3.090E 10	1-7	4.169E 11	9.054E 07
1-33	3.715E 10	3.090E 10	1-33	4.169E 11	9.084E 07
2-3	3.715E 10	3.020E 10	2-7	4.169E 11	9.142E 07
2-33	3.631E 10	3.020E 10	2-33	4.169E 11	9.180E 07
3-3	3.631E 10	2.951E 10	3-3	4.169E 11	9.220E 07
3-33	3.548E 10	2.884E 10	3-37	4.266E 11	9.256E 07
4-3	3.467E 10	2.818E 10	4-1	4.266E 11	9.311E 07
4-33	3.388E 10	2.754E 10	4-37	4.266E 11	9.307E 07
5-3	3.236E 10	2.692E 10	5-1	4.266E 11	9.371E 07
5-33	3.090E 10	2.570E 10	5-37	4.266E 11	9.345E 07
6-1	2.951E 10	2.455E 10	6-0	4.266E 11	1.045E 08
6-30	3.296E 09	4.677E 10	6-44	4.266E 11	2.244E 09
7-7	1.000E 08	5.012E 10	7-28	4.169E 11	3.767E 09
7-33	1.000E 08	5.012E 10	7-59	4.169E 11	4.529E 09
8-7	1.000E 08	5.012E 10	8-29	4.074E 11	5.140E 09
8-37	1.000E 08	5.012E 10	8-59	4.074E 11	5.623E 09
9-3	1.000E 08	5.012E 10	9-29	4.074E 11	6.012E 09
9-37	1.000E 08	5.012E 10	9-59	3.981E 11	6.324E 09
10-7	1.000E 08	5.012E 10	10-30	3.681E 11	6.577E 09
10-33	1.000E 08	5.012E 10	11-0	3.981E 11	6.776E 09
11-15	1.000E 08	5.012E 10	11-30	3.890E 11	6.934E 09
11-31	1.656E 08	5.012E 10	12-0	3.890E 11	7.079E 09
12-7	1.000E 08	5.012E 10	12-30	3.802E 11	7.178E 09
12-37	1.000E 08	5.012E 10	13-0	3.802E 11	7.261E 09
13-7	1.000E 08	5.012E 10	13-30	3.802E 11	7.328E 09
13-33	1.000E 08	5.012E 10	14-0	3.715E 11	7.396E 09
14-7	1.000E 08	5.012E 10	14-30	3.715E 11	7.430E 09
14-37	1.000E 08	5.012E 10	15-0	3.631E 11	7.464E 09
15-7	1.000E 08	5.012E 10	15-30	3.631E 11	7.499E 09
15-33	1.000E 08	5.012E 10	16-0	3.631E 11	7.534E 09
16-7	1.000E 08	5.012E 10	16-30	3.548E 11	7.551E 09
16-37	1.000E 08	5.012E 10	17-0	3.548E 11	7.568E 09
17-7	1.000E 08	5.012E 10	17-30	3.467E 11	7.586E 09
17-31	1.000E 08	5.012E 10	18-0	3.467E 11	7.586E 09
18-1	2.630E 10	2.188E 10	18-0	3.467E 11	7.586E 09
18-33	2.630E 10	2.188E 10	18-33	3.548E 11	7.775E 07
19-3	2.692E 10	2.239E 10	19-7	3.548E 11	7.771E 07
19-33	2.818E 10	2.291E 10	19-33	3.548E 11	7.815E 07
20-3	2.951E 10	2.399E 10	20-7	3.548E 11	7.847E 07
20-33	3.090E 10	2.512E 10	20-33	3.631E 11	7.876E 07
21-3	3.236E 10	2.630E 10	21-7	3.631E 11	7.926E 07
21-31	3.311E 10	2.692E 10	21-33	3.631E 11	7.958E 07
22-1	3.388E 10	2.818E 10	22-7	3.631E 11	8.022E 07
22-31	3.548E 10	2.884E 10	22-33	3.715E 11	8.061E 07
23-3	3.548E 10	2.951E 10	23-7	3.715E 11	8.136E 07
23-31	3.631E 10	3.020E 10	23-33	3.715E 11	8.178E 07

Table A20

Table A21

LATTITUDE= C.C DEG			ALTIUDE= 80.0 KM			LATTITUDE= 0.0 DEG			ALTIUDE= 70.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)						
C- 1	1.698E 11	1.265E 09	0- 3	1.122E 11	1.047E 10						
0-37	1.778E 11	1.327E 09	0-31	1.148E 11	1.072E 10						
1- 1	1.820E 11	1.355E 09	1- 1	1.175E 11	1.098E 10						
1-37	1.905E 11	1.416E 09	1-33	1.175E 11	1.096E 10						
2- 7	1.950E 11	1.462E 09	2- 3	1.202E 11	1.122E 10						
2-33	1.995E 11	1.493E 09	2-31	1.202E 11	1.122E 10						
3- 7	2.042E 11	1.535E 09	3- 1	1.202E 11	1.122E 10						
3-37	2.089E 11	1.567E 09	3-33	1.202E 11	1.122E 10						
4- 3	2.138E 11	1.589E 09	4- 3	1.175E 11	1.098E 10						
4-33	2.138E 11	1.611E 09	4-31	1.148E 11	1.096E 10						
5- 7	2.138E 11	1.626E 09	5- 1	1.148E 11	1.072E 10						
5-31	2.188E 11	1.637E 09	5-31	1.096E 11	1.047E 10						
6- 0	2.138E 11	1.820E 09	6- 0	1.072E 11	1.072E 10						
6-45	1.698E 11	3.715E 10	6-37	9.931E 09	9.333E 10						
7-30	1.318E 11	6.026E 10	7- 7	1.371E 09	1.000E 11						
8- 0	1.096E 11	7.079E 10	7-37	1.875E 08	1.000E 11						
8-30	8.913E 10	7.943E 10	8-15	1.000E 08	1.000E 11						
9- 0	7.244E 10	8.511E 10	8-45	1.000E 08	1.000E 11						
9-30	5.888E 10	8.913E 10	9-30	1.000E 08	1.000E 11						
10- 0	4.786E 10	9.333E 10	10- 0	1.000E 08	1.000E 11						
10-30	3.890E 10	9.550E 10	10-30	1.000E 08	1.000E 11						
11- 0	3.162E 10	9.772E 10	11- 0	1.000E 08	1.000E 11						
11-30	2.570E 10	1.000E 11	11-30	1.000E 08	1.000E 11						
12- 0	2.085E 10	1.023E 11	12- 0	1.000E 08	1.000E 11						
12-30	1.660E 10	1.023E 11	12-30	1.000E 08	1.000E 11						
13- 0	1.349E 10	1.047E 11	13- 0	1.000E 08	1.000E 11						
13-30	1.096E 10	1.047E 11	13-30	1.000E 08	1.000E 11						
14- 0	8.750E 09	1.047E 11	14- 0	1.000E 08	1.000E 11						
14-30	7.031E 09	1.047E 11	14-30	1.000E 08	1.000E 11						
15- 0	5.662E 09	1.047E 11	15- 0	1.000E 08	1.000E 11						
15-30	4.560E 09	1.072E 11	15-30	1.000E 08	1.000E 11						
16- 0	3.664E 09	1.072E 11	16- 0	1.000E 08	1.000E 11						
16-30	2.951E 09	1.072E 11	16-30	1.000E 08	1.000E 11						
17- 0	2.371E 09	1.072E 11	17- 0	1.000E 08	1.000E 11						
17-30	1.910E 09	1.072E 11	17-30	1.000E 08	1.000E 11						
18- 0	1.535E 09	1.072E 11	18- 0	1.000E 08	1.000E 11						
18- 0	1.730E 09	1.072E 11	18- 0	1.000E 08	1.000E 11						
18-33	1.072E 11	8.166E 08	18-37	8.913E 10	8.356E 09						
19- 3	1.096E 11	8.299E 08	19- 3	8.913E 10	8.318E 09						
19-37	1.148E 11	8.551E 08	19-33	8.913E 10	8.375E 09						
20- 1	1.175E 11	8.730E 08	20- 3	9.120E 10	8.472E 09						
20-37	1.230E 11	9.204E 08	20-33	9.120E 10	8.650E 09						
21- 1	1.259E 11	9.462E 08	21- 3	9.550E 10	8.851E 09						
21-37	1.318E 11	1.002E 09	21-31	9.772E 10	9.078E 09						
22- 1	1.380E 11	1.033E 09	22- 3	1.000E 11	9.354E 09						
22-37	1.445E 11	1.094E 09	22-33	1.023E 11	9.638E 09						
23- 1	1.514E 11	1.127E 09	23- 1	1.047E 11	9.886E 09						
23-37	1.585E 11	1.189E 09	23-33	1.096E 11	1.023E 10						

Table A22

Table A23

LATTITUDE= C.C DEG			LATTITUDE=-15.0 DEG		
ALTITUDE= 60.0 KM			ALTITUDE= 90.0 KM		
DECLINATION=23.5 DEG					
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)
0- 0	3.631E 10	3.020E 10	0- 7	3.548E 11	7.831E 07
0-31	3.715E 10	3.020E 10	0-33	3.548E 11	7.865E 07
1- 1	3.715E 10	3.020E 10	1- 7	3.631E 11	7.933E 07
1-31	3.631E 10	3.020E 10	1-33	3.631E 11	7.961E 07
2- 3	3.631E 10	2.951E 10	2- 7	3.631E 11	8.024E 07
2-33	3.631E 10	2.951E 10	2-33	3.631E 11	8.043E 07
3- 3	3.548E 10	2.884E 10	3- 7	3.715E 11	8.098E 07
3-33	3.467E 10	2.818E 10	3-33	3.715E 11	8.108E 07
4- 3	3.388E 10	2.754E 10	4- 7	3.715E 11	8.140E 07
4-33	3.236E 10	2.692E 10	4-33	3.715E 11	8.159E 07
5- 3	3.162E 10	2.570E 10	5- 7	3.715E 11	8.187E 07
5-31	3.020E 10	2.455E 10	5-30	3.715E 11	8.187E 07
6- 0	2.512E 10	2.570E 10	6- 7	3.715E 11	1.598E 09
6-37	1.000E 08	5.012E 10	6-44	3.715E 11	2.904E 09
7- 3	1.000E 08	5.012E 10	7-29	3.631E 11	4.093E 09
7-37	1.000E 08	5.012E 10	7-59	3.631E 11	4.710E 09
8- 3	1.000E 08	5.012E 10	8-29	3.548E 11	5.212E 09
8-45	1.000E 08	5.012E 10	8-59	3.548E 11	5.636E 09
9- 1	1.469E 08	5.012E 10	9-29	3.548E 11	5.970E 09
9-37	1.000E 08	5.012E 10	9-59	3.467E 11	6.252E 09
10- 7	1.000E 08	5.012E 10	10-30	3.467E 11	6.488E 09
10-33	1.000E 08	5.012E 10	11- 0	3.467E 11	6.683E 09
11-15	1.000E 08	5.012E 10	11-30	3.388E 11	6.839E 09
11-31	2.259E 08	5.012E 10	12- 0	3.388E 11	6.966E 09
12- 7	1.000E 08	5.012E 10	12-30	3.311E 11	7.079E 09
12-37	1.000E 08	5.012E 10	13- 0	3.311E 11	7.178E 09
13- 7	1.000E 08	5.012E 10	13-30	3.311E 11	7.244E 09
13-33	1.000E 08	5.012E 10	14- 0	3.236E 11	7.311E 09
14- 7	1.000E 08	5.012E 10	14-30	3.236E 11	7.362E 09
14-37	1.000E 08	5.012E 10	15- 0	3.162E 11	7.413E 09
15- 7	1.000E 08	5.012E 10	15-30	3.162E 11	7.447E 09
15-33	1.000E 08	5.012E 10	16- 0	3.162E 11	7.482E 09
16- 7	1.000E 08	5.012E 10	16-30	3.090E 11	7.499E 09
16-37	1.000E 08	5.012E 10	17- 0	3.090E 11	7.516E 09
17- 3	1.000E 08	5.012E 10	17-30	3.090E 11	7.534E 09
17-45	1.000E 08	5.012E 10	18- 0	3.020E 11	7.551E 09
18- 0	3.192E 09	4.677E 10	18- 7	3.020E 11	7.588E 09
18-31	2.630E 10	2.188E 10	18-30	3.090E 11	8.750E 08
19- 1	2.630E 10	2.188E 10	19- 3	3.090E 11	6.771E 07
19-31	2.754E 10	2.239E 10	19-33	3.090E 11	6.796E 07
20- 1	2.818E 10	2.344E 10	20- 7	3.090E 11	6.821E 07
20-31	2.951E 10	2.455E 10	20-37	3.090E 11	6.852E 07
21- 1	3.090E 10	2.512E 10	21- 3	3.162E 11	6.881E 07
21-31	3.236E 10	2.630E 10	21-33	3.162E 11	6.920E 07
22- 3	3.311E 10	2.754E 10	22- 3	3.162E 11	6.967E 07
22-31	3.388E 10	2.818E 10	22-37	3.162E 11	7.013E 07
23- 1	3.467E 10	2.884E 10	23- 3	3.236E 11	7.057E 07
23-31	3.548E 10	2.951E 10	23-37	3.236E 11	7.122E 07

Table A24

Table A25

LATITUDE=-15.0 DEG			LATITUDE=-15.0 DEG		
ALTITUDE= 80.0 KM			ALTITUDE= 70.0 KM		
DECLINATION=23.5 DEG					
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)
0- 1	1.549E 11	1.164E 09	0- 3	1.072E 11	9.931E 09
0-37	1.622E 11	1.227E 09	0-33	1.072E 11	1.000E 10
1- 7	1.698E 11	1.265E 09	1- 1	1.096E 11	1.023E 10
1-33	1.738E 11	1.306E 09	1-31	1.122E 11	1.047E 10
2- 7	1.778E 11	1.349E 09	2- 3	1.122E 11	1.047E 10
2-33	1.820E 11	1.380E 09	2-33	1.122E 11	1.047E 10
3- 7	1.862E 11	1.419E 09	3- 1	1.122E 11	1.047E 10
3-33	1.905E 11	1.439E 09	3-33	1.122E 11	1.047E 10
4- 7	1.950E 11	1.466E 09	4- 3	1.096E 11	1.023E 10
4-33	1.950E 11	1.476E 09	4-31	1.096E 11	1.023E 10
5- 3	1.950E 11	1.483E 09	5- 1	1.072E 11	9.977E 09
5-30	1.950E 11	1.486E 09	5-30	1.023E 11	9.727E 09
6- 7	1.660E 11	2.754E 10	6- 3	1.514E 10	8.511E 10
6-45	1.349E 11	4.786E 10	6-37	1.656E 09	9.550E 10
7-30	1.023E 11	6.457E 10	7- 7	2.291E 08	9.550E 10
8- 0	8.511E 10	7.244E 10	7-45	1.000E 08	9.772E 10
8-30	6.918E 10	7.943E 10	8-15	1.000E 08	9.772E 10
9- 0	5.623E 10	8.318E 10	8-45	1.000E 08	9.772E 10
9-30	4.677E 10	8.710E 10	9-30	1.000E 08	9.772E 10
10- 0	3.802E 10	9.120E 10	10- 0	1.000E 08	9.772E 10
10-30	3.090E 10	9.333E 10	10-30	1.000E 08	9.772E 10
11- 0	2.512E 10	9.550E 10	11- 0	1.000E 08	9.772E 10
11-30	1.995E 10	9.550E 10	11-30	1.000E 08	9.772E 10
12- 0	1.622E 10	9.772E 10	12- 0	1.000E 08	9.772E 10
12-30	1.318E 10	9.772E 10	12-30	1.000E 08	9.772E 10
13- 0	1.072E 10	1.000E 11	13- 0	1.000E 08	9.772E 10
13-30	8.650E 09	1.000E 11	13-30	1.000E 08	9.772E 10
14- 0	6.982E 09	1.000E 11	14- 0	1.000E 08	9.772E 10
14-30	5.649E 09	1.000E 11	14-30	1.000E 08	9.772E 10
15- 0	4.560E 09	1.000E 11	15- 0	1.000E 08	9.772E 10
15-30	3.681E 09	1.000E 11	15-30	1.000E 08	9.772E 10
16- 0	2.972E 09	1.023E 11	16- 0	1.000E 08	9.772E 10
16-30	2.404E 09	1.023E 11	16-30	1.000E 08	9.772E 10
17- 0	1.941E 09	1.023E 11	17- 0	1.000E 08	9.772E 10
17-30	1.567E 09	1.023E 11	17-30	1.000E 08	9.772E 10
18- 0	1.265E 09	1.023E 11	18- 0	1.000E 08	9.772E 10
18- 7	1.197E 09	1.023E 11	18- 7	1.000E 08	9.772E 10
18-30	9.120E 10	1.148E 10	18-30	7.943E 10	1.585E 10
19- 7	1.023E 11	7.780E 08	19- 3	8.511E 10	8.091E 09
19-33	1.047E 11	7.907E 08	19-33	8.511E 10	8.072E 09
20- 3	1.072E 11	8.128E 08	20- 3	8.710E 10	8.147E 09
20-37	1.122E 11	8.453E 08	20-33	8.710E 10	8.260E 09
21- 3	1.175E 11	8.730E 08	21- 3	8.913E 10	8.433E 09
21-37	1.230E 11	9.162E 08	21-33	9.120E 10	8.630E 09
22- 3	1.259E 11	9.528E 08	22- 3	9.550E 10	8.851E 09
22-37	1.318E 11	9.977E 08	22-33	9.772E 10	9.099E 09
23- 3	1.380E 11	1.038E 09	23- 1	1.000E 11	9.333E 09
23-45	1.445E 11	1.099E 09	23-31	1.023E 11	9.572E 09

Table A26

Table A27

LATTITUDE=-15.0 DEG			LATTITUDE=-30.0 DEG		
ALTITUDE= 60.0 KM			ALTITUDE= 90.0 KM		
DECLINATION=23.5 DEG					
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(C3)
0- 0	3.548E 10	2.951E 10	0- 7	3.467E 11	6.019E 07
0-31	3.548E 10	2.884E 10	0-33	3.467E 11	6.050E 07
1- 1	3.548E 10	2.884E 10	1- 7	3.467E 11	6.076E 07
1-31	3.548E 10	2.884E 10	1-33	3.467E 11	6.100E 07
2- 3	3.467E 10	2.884E 10	2- 7	3.548E 11	6.144E 07
2-33	3.467E 10	2.818E 10	2-33	3.548E 11	6.150E 07
3- 3	3.388E 10	2.754E 10	3- 7	3.548E 11	6.177E 07
3-33	3.311E 10	2.754E 10	3-33	3.548E 11	6.194E 07
4- 3	3.236E 10	2.630E 10	4- 3	3.548E 11	6.216E 07
4-33	3.090E 10	2.570E 10	4-37	3.548E 11	6.230E 07
5- 1	2.951E 10	2.455E 10	5- 0	3.548E 11	6.234E 07
5-30	2.818E 10	2.291E 10	5-37	3.548E 11	1.138E 09
6- 7	1.000E 08	5.012E 10	6-14	3.548E 11	1.845E 09
6-33	1.000E 08	5.012E 10	6-44	3.548E 11	2.213E 09
7- 7	1.000E 08	5.012E 10	7-14	3.467E 11	2.460E 09
7-37	1.000E 08	5.012E 10	7-44	3.467E 11	2.636E 09
8- 3	1.000E 08	5.012E 10	8-29	3.388E 11	2.793E 09
8-45	1.000E 08	5.012E 10	8-59	3.388E 11	2.864E 09
9- 1	1.200E 08	5.012E 10	9-29	3.388E 11	2.911E 09
9-37	1.000E 08	5.012E 10	9-59	3.388E 11	2.938E 09
10- 7	1.000E 08	5.012E 10	10-30	3.311E 11	2.965E 09
10-33	1.000E 08	5.012E 10	11- 0	3.311E 11	2.979E 09
11-15	1.000E 08	5.012E 10	11-30	3.311E 11	2.985E 09
11-31	1.820E 08	5.012E 10	12- 0	3.236E 11	2.992E 09
12- 7	1.000E 08	5.012E 10	12-30	3.236E 11	2.999E 09
12-37	1.000E 08	5.012E 10	13- 0	3.236E 11	3.006E 09
13- 7	1.000E 08	5.012E 10	13-30	3.162E 11	3.006E 09
13-33	1.000E 08	5.012E 10	14- 0	3.162E 11	3.006E 09
14- 7	1.000E 08	5.012E 10	14-30	3.162E 11	3.013E 09
14-37	1.000E 08	5.012E 10	15- 0	3.090E 11	3.013E 09
15- 7	1.000E 08	5.012E 10	15-30	3.090E 11	3.013E 09
15-33	1.000E 08	5.012E 10	16- 0	3.090E 11	3.013E 09
16- 7	1.000E 08	5.012E 10	16-30	3.020E 11	3.013E 09
16-37	1.000E 08	5.012E 10	17- 0	3.020E 11	3.013E 09
17- 3	1.000E 08	5.012E 10	17-30	3.020E 11	3.013E 09
17-45	1.000E 08	5.012E 10	18- 0	2.951E 11	3.013E 09
18- 1	1.276E 08	5.012E 10	18-15	2.951E 11	3.013E 09
18-30	2.630E 10	2.291E 10	18-37	2.951E 11	3.013E 09
19- 1	2.630E 10	2.138E 10	19- 0	2.951E 11	7.907E 08
19-33	2.630E 10	2.188E 10	19-33	2.951E 11	5.196E 07
20- 3	2.692E 10	2.239E 10	20- 3	2.951E 11	5.219E 07
20-33	2.818E 10	2.344E 10	20-37	3.020E 11	5.234E 07
21- 3	2.951E 10	2.399E 10	21- 3	3.020E 11	5.254E 07
21-33	3.090E 10	2.512E 10	21-37	3.020E 11	5.297E 07
22- 3	3.162E 10	2.570E 10	22- 3	3.020E 11	5.304E 07
22-33	3.236E 10	2.692E 10	22-37	3.090E 11	5.337E 07
23- 3	3.311E 10	2.754E 10	23- 3	3.090E 11	5.361E 07
23-33	3.388E 10	2.818E 10	23-33	3.090E 11	5.393E 07

Table A28

Table A29

LATITUDE=-30.0 DEG			LATITUDE=-30.0 DEG		
ALTITUDE= 80.0 KM			ALTITUDE= 70.0 KM		
DECLINATION=23.5 DEG					
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)
C- 7	1.096E 11	6.501E 08	C- 7	1.000E 11	7.516E 09
C-33	1.148E 11	6.761E 08	C-37	1.000E 11	7.656E 09
1- 3	1.175E 11	7.047E 08	1- 3	1.023E 11	7.762E 09
1-37	1.230E 11	7.345E 08	1-37	1.047E 11	7.852E 09
2- 1	1.288E 11	7.621E 08	2- 3	1.047E 11	7.925E 09
2-37	1.318E 11	7.834E 08	2-37	1.047E 11	7.962E 09
3- 1	1.349E 11	8.091E 08	3- 3	1.047E 11	7.962E 09
3-37	1.380E 11	8.204E 08	3-37	1.047E 11	7.907E 09
4- 1	1.413E 11	8.318E 08	4- 3	1.023E 11	7.852E 09
4-37	1.413E 11	8.395E 08	4-33	1.023E 11	7.827E 09
5- 0	1.413E 11	8.414E 08	5- 0	1.000E 11	7.568E 09
5-37	1.230E 11	1.660E 10	5-37	1.549E 10	7.762E 10
6-15	1.047E 11	2.884E 10	6- 7	3.020E 09	8.710E 10
6-45	9.120E 10	3.631E 10	6-37	5.781E 08	8.710E 10
7-30	7.244E 10	4.467E 10	7-15	1.000E 08	8.913E 10
8- 0	6.166E 10	4.898E 10	7-45	1.000E 08	8.913E 10
8-30	5.248E 10	5.129E 10	8-30	1.000E 08	8.913E 10
9- 0	4.467E 10	5.495E 10	9- 0	1.000E 08	8.913E 10
9-30	3.802E 10	5.623E 10	9-30	1.000E 08	8.913E 10
10- 0	3.236E 10	5.888E 10	10- 0	1.000E 08	8.913E 10
10-30	2.692E 10	6.026E 10	10-30	1.000E 08	8.913E 10
11- 0	2.291E 10	6.166E 10	11- 0	1.000E 08	8.913E 10
11-30	1.950E 10	6.166E 10	11-30	1.000E 08	8.913E 10
12- 0	1.622E 10	6.310E 10	12- 0	1.000E 08	8.913E 10
12-30	1.380E 10	6.310E 10	12-30	1.000E 08	8.913E 10
13- 0	1.148E 10	6.457E 10	13- 0	1.000E 08	8.913E 10
13-30	9.683E 09	6.457E 10	13-30	1.000E 08	8.913E 10
14- 0	8.147E 09	6.457E 10	14- 0	1.000E 08	8.913E 10
14-30	6.839E 09	6.607E 10	14-30	1.000E 08	8.913E 10
15- 0	5.754E 09	6.607E 10	15- 0	1.000E 08	8.913E 10
15-30	4.831E 09	6.607E 10	15-30	1.000E 08	8.913E 10
16- 0	4.055E 09	6.607E 10	16- 0	1.000E 08	8.913E 10
16-30	3.404E 09	6.607E 10	16-30	1.000E 08	8.913E 10
17- 0	2.858E 09	6.607E 10	17- 0	1.000E 08	8.913E 10
17-30	2.404E 09	6.607E 10	17-30	1.000E 08	8.913E 10
18- 0	2.018E 09	6.607E 10	18- 0	1.000E 08	8.913E 10
18-15	1.845E 09	6.607E 10	18-15	1.000E 08	8.913E 10
18-37	1.618E 09	6.607E 10	18-37	1.000E 08	8.913E 10
19- 0	5.129E 10	1.698E 10	19- 0	6.310E 10	2.455E 10
19-37	6.918E 10	4.102E 08	19-33	8.128E 10	6.138E 09
20- 3	7.079E 10	4.198E 08	20- 1	8.128E 10	6.180E 09
20-33	7.244E 10	4.345E 08	20-37	8.128E 10	6.237E 09
21- 3	7.586E 10	4.539E 08	21- 3	8.318E 10	6.339E 09
21-37	8.128E 10	4.808E 08	21-33	8.511E 10	6.471E 09
22- 1	8.511E 10	4.920E 08	22- 3	8.710E 10	6.637E 09
22-37	8.913E 10	5.333E 08	22-33	8.913E 10	6.808E 09
23- 7	9.550E 10	5.649E 08	23- 3	9.120E 10	6.982E 09
23-33	1.000E 11	5.888E 08	23-33	9.333E 10	7.161E 09

Table A30

Table A31

LATITUDE=-30.0 DEG			ALTITUDE= 60.0 KM			LATITUDE=-45.0 DEG			ALTITUDE= 90.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(C)	N(C3)	TIME(HR-MIN)	N(C)	N(C3)						
0- 0	3.548E 10	2.570E 10	0- 7	2.951E 11	4.889E 07						
0-33	3.548E 10	2.512E 10	0-31	2.951E 11	4.872E 07						
1- 3	3.548E 10	2.512E 10	1- 7	2.951E 11	4.934E 07						
1-31	3.548E 10	2.512E 10	1-33	2.951E 11	4.953E 07						
2- 3	3.467E 10	2.512E 10	2- 7	3.020E 11	4.977E 07						
2-33	3.467E 10	2.455E 10	2-31	3.020E 11	4.972E 07						
3- 1	3.388E 10	2.399E 10	3- 7	3.020E 11	5.015E 07						
3-31	3.311E 10	2.344E 10	3-33	3.020E 11	5.028E 07						
4- 1	3.162E 10	2.291E 10	4- 3	3.020E 11	5.038E 07						
4-31	3.020E 10	2.188E 10	4-33	3.020E 11	5.082E 08						
5- 0	2.884E 10	2.089E 10	5-14	3.020E 11	1.265E 09						
5-37	1.000E 08	4.898E 10	5-44	2.951E 11	1.600E 09						
6- 7	1.000E 08	4.898E 10	6-14	2.951E 11	1.820E 09						
6-37	1.000E 08	4.898E 10	6-44	2.951E 11	1.968E 09						
7- 7	1.000E 08	4.898E 10	7-29	2.884E 11	2.099E 09						
7-37	1.000E 08	4.898E 10	7-59	2.884E 11	2.153E 09						
8- 7	1.000E 08	4.898E 10	8-29	2.884E 11	2.188E 09						
8-37	1.000E 08	4.898E 10	8-59	2.884E 11	2.208E 09						
9- 7	1.000E 08	4.898E 10	9-29	2.818E 11	2.228E 09						
9-37	1.000E 08	4.898E 10	9-59	2.818E 11	2.239E 09						
10- 7	1.000E 08	4.898E 10	10-30	2.818E 11	2.244E 09						
10-37	1.000E 08	4.898E 10	11- 0	2.754E 11	2.249E 09						
11- 7	1.000E 08	4.898E 10	11-30	2.754E 11	2.254E 09						
11-33	1.000E 08	4.898E 10	12- 0	2.754E 11	2.254E 09						
12- 7	1.000E 08	4.898E 10	12-30	2.692E 11	2.254E 09						
12-33	1.000E 08	4.898E 10	13- 0	2.692E 11	2.259E 09						
13- 7	1.000E 08	4.898E 10	13-30	2.692E 11	2.259E 09						
13-33	1.000E 08	4.898E 10	14- 0	2.692E 11	2.259E 09						
14- 7	1.000E 08	4.898E 10	14-30	2.630E 11	2.259E 09						
14-33	1.000E 08	4.898E 10	15- 0	2.630E 11	2.259E 09						
15- 7	1.000E 08	4.898E 10	15-30	2.630E 11	2.259E 09						
15-33	1.000E 08	4.898E 10	16- 0	2.570E 11	2.259E 09						
16- 7	1.000E 08	4.898E 10	16-30	2.570E 11	2.259E 09						
16-33	1.000E 08	4.898E 10	17- 0	2.570E 11	2.259E 09						
17- 7	1.000E 08	4.898E 10	17-30	2.512E 11	2.259E 09						
17-37	1.000E 08	4.898E 10	18- 0	2.512E 11	2.259E 09						
18- 7	1.000E 08	4.898E 10	18-30	2.512E 11	2.259E 09						
18-37	1.000E 08	4.898E 10	19- 0	2.512E 11	2.259E 09						
19- 0	2.570E 10	2.291E 10	19-15	2.455E 11	2.259E 09						
19-31	2.754E 10	1.950E 10	19-33	2.455E 11	2.259E 09						
20- 1	2.754E 10	1.950E 10	20- 1	2.512E 11	4.163E 07						
20-33	2.818E 10	1.995E 10	20-33	2.512E 11	4.160E 07						
21- 3	2.884E 10	2.089E 10	21- 7	2.512E 11	4.188E 07						
21-33	3.020E 10	2.138E 10	21-31	2.512E 11	4.241E 07						
22- 3	3.090E 10	2.239E 10	22- 7	2.512E 11	4.229E 07						
22-33	3.236E 10	2.291E 10	22-33	2.570E 11	4.258E 07						
23- 3	3.311E 10	2.344E 10	23- 7	2.570E 11	4.275E 07						
23-33	3.388E 10	2.399E 10	23-33	2.570E 11	4.296E 07						

Table A32

Table A33

LATITUDE=-45.0 DEG			ALTIUDE= 85.0 KM			LATITUDE=-45.0 DEG			ALTIUDE= 80.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(C)	N(O3)						
0- 7	6.457E 10	5.795E 07	0- 7	5.012E 10	1.786E 08						
0-33	6.761E 10	5.972E 07	0-33	5.248E 10	1.879E 08						
1- 7	7.079E 10	6.265E 07	1- 7	5.623E 10	2.009E 08						
1-33	7.244E 10	6.431E 07	1-33	5.888E 10	2.104E 08						
2- 7	7.586E 10	6.700E 07	2- 3	6.166E 10	2.203E 08						
2-33	7.762E 10	6.834E 07	2-37	6.457E 10	2.301E 08						
3- 7	7.943E 10	7.053E 07	3- 1	6.607E 10	2.382E 08						
3-33	7.943E 10	7.125E 07	3-37	6.761E 10	2.421E 08						
4- 3	8.128E 10	7.170E 07	4- 3	6.918E 10	2.455E 08						
4-33	8.128E 10	7.762E 08	4-33	6.607E 10	2.612E 09						
5-14	7.762E 10	2.312E 09	5-15	6.026E 10	7.396E 09						
5-44	7.762E 10	3.273E 09	5-45	5.495E 10	1.023E 10						
6-29	7.413E 10	4.498E 09	6-30	4.786E 10	1.349E 10						
6-59	7.244E 10	5.200E 09	7- 0	4.365E 10	1.514E 10						
7-29	7.079E 10	5.808E 09	7-30	3.981E 10	1.660E 10						
7-59	6.918E 10	6.353E 09	8- 0	3.548E 10	1.778E 10						
8-29	6.761E 10	6.839E 09	8-30	3.236E 10	1.862E 10						
8-59	6.607E 10	7.278E 09	9- 0	2.884E 10	1.950E 10						
9-29	6.310E 10	7.656E 09	9-30	2.630E 10	1.995E 10						
9-59	6.166E 10	7.998E 09	10- 0	2.344E 10	2.089E 10						
10-30	6.026E 10	8.299E 09	10-30	2.138E 10	2.138E 10						
11- 0	5.888E 10	8.570E 09	11- 0	1.905E 10	2.188E 10						
11-30	5.754E 10	8.810E 09	11-30	1.698E 10	2.239E 10						
12- 0	5.495E 10	9.037E 09	12- 0	1.514E 10	2.239E 10						
12-30	5.370E 10	9.226E 09	12-30	1.349E 10	2.291E 10						
13- 0	5.248E 10	9.397E 09	13- 0	1.230E 10	2.291E 10						
13-30	5.129E 10	9.550E 09	13-30	1.096E 10	2.344E 10						
14- 0	4.898E 10	9.705E 09	14- 0	9.750E 09	2.344E 10						
14-30	4.786E 10	9.840E 09	14-30	8.710E 09	2.344E 10						
15- 0	4.677E 10	9.954E 09	15- 0	7.780E 09	2.399E 10						
15-30	4.467E 10	1.000E 10	15-30	6.950E 09	2.399E 10						
16- 0	4.365E 10	1.023E 10	16- 0	6.209E 09	2.399E 10						
16-30	4.266E 10	1.023E 10	16-30	5.546E 09	2.399E 10						
17- 0	4.074E 10	1.023E 10	17- 0	4.943E 09	2.399E 10						
17-30	3.981E 10	1.047E 10	17-30	4.416E 09	2.399E 10						
18- 0	3.890E 10	1.047E 10	18- 0	3.936E 09	2.455E 10						
18-30	3.715E 10	1.047E 10	18-30	3.508E 09	2.455E 10						
19- 0	3.631E 10	1.072E 10	19- 0	3.133E 09	2.455E 10						
19-15	3.548E 10	1.072E 10	19-15	2.958E 09	2.455E 10						
19-33	3.467E 10	1.072E 10	19-33	2.754E 09	2.455E 10						
20- 1	4.571E 10	4.076E 07	20- 1	2.754E 10	1.000E 08						
20-33	4.677E 10	4.138E 07	20-33	2.884E 10	1.026E 08						
21- 7	4.786E 10	4.308E 07	21- 7	3.090E 10	1.102E 08						
21-33	5.012E 10	4.456E 07	21-33	3.236E 10	1.169E 08						
22- 7	5.248E 10	4.713E 07	22- 7	3.548E 10	1.282E 08						
22-33	5.495E 10	4.869E 07	22-33	3.802E 10	1.361E 08						
23- 7	5.754E 10	5.059E 07	23- 7	4.169E 10	1.500E 08						
23-33	6.026E 10	5.333E 07	23-33	4.467E 10	1.592E 08						



Table A34

Table A35

LATTITUDE=-45.0 DEG			ALTITUDE= 75.0 KM			LATTITUDE=-45.0 DEG			ALTITUDE= 70.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(C)	N(O3)	TIME(HR-MIN)	N(C)	N(O3)						
C- 3	7.244E 10	9.705E 08	0- 3	7.762E 10	3.750E 09						
C-37	7.586E 10	1.021E 09	0-37	8.128E 10	3.855E 09						
1- 3	7.762E 10	1.054E 09	1- 3	8.318E 10	3.936E 09						
1-37	8.128E 10	1.099E 09	1-37	8.318E 10	4.009E 09						
2- 7	8.511E 10	1.132E 09	2- 7	8.511E 10	4.074E 09						
2-33	8.511E 10	1.156E 09	2-33	8.710E 10	4.112E 09						
3- 3	8.710E 10	1.180E 09	3- 7	8.710E 10	4.130E 09						
3-37	8.913E 10	1.200E 09	3-33	8.710E 10	4.130E 09						
4- 3	8.913E 10	1.211E 09	4- 3	8.511E 10	4.112E 09						
4-33	7.762E 10	1.202E 10	4-33	5.012E 10	3.388E 10						
5-15	5.129E 10	2.884E 10	5- 7	1.514E 10	5.754E 10						
5-45	3.631E 10	3.548E 10	5-37	4.932E 09	6.310E 10						
6-30	2.138E 10	4.169E 10	6-15	1.208E 09	6.607E 10						
7- 0	1.479E 10	4.365E 10	6-45	3.908E 08	6.607E 10						
7-30	1.023E 10	4.467E 10	7-15	1.262E 08	6.607E 10						
8- 0	6.582E 09	4.571E 10	7-45	1.000E 08	6.607E 10						
8-30	4.786E 09	4.677E 10	8-30	1.000E 08	6.607E 10						
9- 0	3.281E 09	4.677E 10	9- 0	1.000E 08	6.607E 10						
9-30	2.244E 09	4.786E 10	9-30	1.000E 08	6.607E 10						
10- 0	1.535E 09	4.786E 10	10- 0	1.000E 08	6.607E 10						
10-30	1.047E 09	4.786E 10	10-30	1.000E 08	6.607E 10						
11- 0	7.161E 08	4.786E 10	11- 0	1.000E 08	6.607E 10						
11-30	4.887E 08	4.786E 10	11-30	1.000E 08	6.607E 10						
12- 0	3.334E 08	4.786E 10	12- 0	1.000E 08	6.607E 10						
12-30	2.280E 08	4.786E 10	12-30	1.000E 08	6.607E 10						
13- 0	1.556E 08	4.786E 10	13- 0	1.000E 08	6.607E 10						
13-30	1.062E 08	4.786E 10	13-30	1.000E 08	6.607E 10						
14- 0	1.000E 08	4.786E 10	14- 0	1.000E 08	6.607E 10						
14-30	1.000E 08	4.786E 10	14-30	1.000E 08	6.607E 10						
15- 0	1.000E 08	4.786E 10	15- 0	1.000E 08	6.607E 10						
15-30	1.000E 08	4.786E 10	15-30	1.000E 08	6.607E 10						
16- 0	1.000E 08	4.786E 10	16- 0	1.000E 08	6.607E 10						
16-30	1.000E 08	4.786E 10	16-30	1.000E 08	6.607E 10						
17- 0	1.000E 08	4.786E 10	17- 0	1.000E 08	6.607E 10						
17-30	1.000E 08	4.786E 10	17-30	1.000E 08	6.607E 10						
18- 0	1.000E 08	4.786E 10	18- 0	1.000E 08	6.607E 10						
18-30	1.000E 08	4.786E 10	18-30	1.000E 08	6.607E 10						
19- 0	1.000E 08	4.786E 10	19- 0	1.000E 08	6.607E 10						
19-15	1.000E 08	4.786E 10	19-15	1.000E 08	6.607E 10						
19-33	1.000E 08	4.786E 10	19-33	1.000E 08	6.607E 10						
20- 1	4.786E 10	6.486E 08	20- 1	6.310E 10	3.006E 09						
20-21	4.898E 10	6.637E 08	20-33	6.310E 10	3.041E 09						
21- 7	5.129E 10	6.966E 08	21- 3	6.457E 10	3.105E 09						
21-33	5.370E 10	7.244E 08	21-33	6.607E 10	3.192E 09						
22- 3	5.623E 10	7.603E 08	22- 7	6.918E 10	3.296E 09						
22-37	6.026E 10	8.035E 08	22-33	7.079E 10	3.388E 09						
23- 1	6.166E 10	8.492E 08	23- 3	7.244E 10	3.499E 09						
23-37	6.607E 10	8.892E 08	23-33	7.586E 10	3.606E 09						

Table A36

Table A37

LATITUDE=-45.0 DEG			ALTIITUDE= 65.0 KM			LATITUDE=-45.0 DEG			ALTIITUDE= 60.0 KM		
DECLINATION=23.5 DEG											
TIME(HR-MIN)	N(O)	N(O3)	TIME(HR-MIN)	N(O)	N(O3)						
0- 7	5.623E 10	8.995E 09	0- 0	3.388E 10	1.820E 10						
0-37	5.623E 10	9.141E 09	0-33	3.467E 10	1.820E 10						
1- 3	5.754E 10	9.247E 09	1- 3	3.467E 10	1.862E 10						
1-31	5.754E 10	9.311E 09	1-33	3.467E 10	1.862E 10						
2- 1	5.754E 10	9.333E 09	2- 1	3.467E 10	1.862E 10						
2-37	5.754E 10	9.311E 09	2-21	3.467E 10	1.820E 10						
3- 7	5.623E 10	9.204E 09	3- 3	3.388E 10	1.778E 10						
3-33	5.623E 10	9.078E 09	3-33	3.311E 10	1.738E 10						
4- 3	5.495E 10	8.851E 09	4- 1	3.162E 10	1.698E 10						
4-33	9.099E 09	4.898E 10	4-31	2.173E 08	4.571E 10						
5- 7	2.350E 08	5.623E 10	5- 7	1.000E 08	4.571E 10						
5-37	1.000E 08	5.623E 10	5-45	1.000E 08	4.571E 10						
6-15	1.000E 08	5.623E 10	6-15	1.000E 08	4.571E 10						
6-45	1.000E 08	5.623E 10	6-33	1.000E 08	4.571E 10						
7-30	1.000E 08	5.623E 10	7- 7	1.000E 08	4.571E 10						
8- 0	1.000E 08	5.623E 10	7-45	1.000E 08	4.571E 10						
8-30	1.000E 08	5.623E 10	8-15	1.000E 08	4.571E 10						
9- 0	1.000E 08	5.623E 10	8-33	1.000E 08	4.571E 10						
9-30	1.000E 08	5.623E 10	9- 7	1.000E 08	4.571E 10						
10- 0	1.000E 08	5.623E 10	9-45	1.000E 08	4.571E 10						
10-30	1.000E 08	5.623E 10	10- 7	1.000E 08	4.571E 10						
11- 0	1.000E 08	5.623E 10	10-37	1.000E 08	4.571E 10						
11-30	1.000E 08	5.623E 10	11- 7	1.000E 08	4.571E 10						
12- 0	1.000E 08	5.623E 10	11-45	1.000E 08	4.571E 10						
12-15	1.000E 08	5.623E 10	12- 7	1.000E 08	4.571E 10						
12-45	1.000E 08	5.623E 10	12-37	1.000E 08	4.571E 10						
13-30	1.000E 08	5.623E 10	13- 7	1.000E 08	4.571E 10						
14- 0	1.000E 08	5.623E 10	13-45	1.000E 08	4.571E 10						
14-30	1.000E 08	5.623E 10	14- 7	1.000E 08	4.571E 10						
15- 0	1.000E 08	5.623E 10	14-37	1.000E 08	4.571E 10						
15-30	1.000E 08	5.623E 10	15- 7	1.000E 08	4.571E 10						
16- 0	1.000E 08	5.623E 10	15-45	1.000E 08	4.571E 10						
16-30	1.000E 08	5.623E 10	16- 7	1.000E 08	4.571E 10						
17- 0	1.000E 08	5.623E 10	16-37	1.000E 08	4.571E 10						
17-15	1.000E 08	5.623E 10	17- 7	1.000E 08	4.571E 10						
17-45	1.000E 08	5.623E 10	17-45	1.000E 08	4.571E 10						
18-30	1.000E 08	5.623E 10	18- 7	1.000E 08	4.571E 10						
19- 0	1.000E 08	5.623E 10	18-37	1.000E 08	4.571E 10						
19-15	1.000E 08	5.623E 10	19- 7	1.000E 08	4.571E 10						
19-33	1.000E 08	5.623E 10	19-33	1.000E 08	4.571E 10						
20- 1	4.786E 10	7.780E 09	20- 1	2.951E 10	1.549E 10						
20-33	4.786E 10	7.745E 09	20-33	2.951E 10	1.549E 10						
21- 3	4.786E 10	7.798E 09	21- 1	2.951E 10	1.549E 10						
21-37	4.898E 10	7.962E 09	21-31	3.020E 10	1.585E 10						
22- 7	5.012E 10	8.147E 09	22- 1	3.090E 10	1.660E 10						
22-33	5.129E 10	8.318E 09	22-33	3.236E 10	1.698E 10						
23- 1	5.248E 10	8.511E 09	23- 3	3.311E 10	1.738E 10						
23-31	5.370E 10	8.730E 09	23-31	3.388E 10	1.778E 10						

## 2. COMPUTATIONAL CONSIDERATIONS

The entire computer program used in numerical integration of the O and O<sub>3</sub> differential equations is given in Tables A38 through A42. Many important notes are contained on the comment cards within each subprogram. All computations directly involved in the numerical integration of the differential equations were done in double precision to decrease round-off errors. The program is written in Fortran IV for the IBM System/360. The actual numerical integration is done by the DRKGS Subroutine which is provided for the IBM System/360 by IBM. The integration was done by means of the fourth order Runge-Kutta formulae in the modification due to Gill. The comment cards in Table A42 contain more information on this method. The maximum absolute error in concentration per integration was limited to a factor of roughly .0005 times the initial value of the concentration for each constituent O and O<sub>3</sub>. It is unlikely that such errors would accumulate significantly as the integration proceeds because an excess or dearth of a species results in a corresponding increase or decrease of its rate of destruction. During most of the calculations this error limit was satisfied while integrating in 15-minute jumps, but at sunrise and sunset, the integration interval dropped to 1 second or less. The typical IBM System/360 computer running time for a 6-day profile of O and O<sub>3</sub> was about 3 minutes.

Following is a list of notes pertaining to the programs. Additional information is contained on comment cards within each program.

Main Program Table A38

ZØ2 = integral of Ø2 from 300 km to altitude of interest, LAT = latitude, DEC = declination TXPLOT, TYX, TYPLOT, SIZEX, SIZES are control vectors for the "calcomp" plctter. All statements of the form CALL CCP . . . are plotting

Table A38

```

C      MAIN PROGRAM - OXYGEN ATMOSPHERE
ISN 0002      EXTERNAL FCT,OUTP
ISN 0003      DIMENSION Y(2),DERY(2),PRMT(5),AUX(8,2),X1(7,8),X3(7,8)
ISN 0004      DIMENSION T(7,8),Z02(7,8),R(87),S2(87),S3(87),VA(13),VB(13)
ISN 0005      DIMENSION TXPLOT(2),TYPLCT(2)
ISN 0006      DIMENSION WAVE(87)
ISN 0007      DIMENSION TYX(2)
ISN 0008      DOUBLE PRECISION Y,DERY,PRMT,AUX,X,X2,Z02,B1,B2,B3,X1,X3,Y,W,S,
              1RATE,VA,VB,A,F2,F3,LAT,DEC,S2,S3,R,H
ISN 0009      COMMON K,M
ISN 0010      COMMON/BFCT/ B1(7,8),B2(7,8),B3(7,8),X2(7,8),A
ISN 0011      COMMON/BOUPT/W,S,NIN
ISN 0012      COMMON /BFCCU/ RATE(5)
ISN 0013      COMMON /BFCL/ F2(7,8),F3(7,8)
ISN 0014      COMMON /BQFL/ LAT(8),DEC(8)
ISN 0015      COMMON/BFLUX/ R,S2,S3,Z02,VA,VB,T
ISN 0016      COMMON/BCHI/COSCHI,YW(300),YW3(300),XW(300),YS(49),YS3(49),XHR(49)
ISN 0017      DOUBLE PRECISION COSCHI
ISN 0018      READ(5,1)X2
ISN 0019      READ(5,1)Z02
ISN 0020      READ(5,1)T
ISN 0021      READ(5,1)X1
ISN 0022      READ(5,1)X3
ISN 0023      READ(5,1)B1
ISN 0024      READ(5,1)B2
ISN 0025      READ(5,1)B3
ISN 0026      READ(5,1)LAT
ISN 0027      READ(5,1)DEC
ISN 0028      1
              FORMAT(1P8D10.3)
              C      X1=(01) X2=(02) X3=(03) T=TEMP B1,B2,B3=K1,K2,K3 RATE COEFFICIENTS (DASA)
              C      A=(FRACTIONAL COMPOSITION OF O2 TO (M))**-1
              C      YS,YS3,XHR ARE VALUES OF O1,O3,&TIME FOR THE LAST 24 HRS (PUNCHED ON CARDS)
              C      YW,YW3,XW ARE PRINTED VALUES OF O1,O3,&TIME USED BY PLOTTER
              C      SEE SUBROUTINE DRKGS FOR VALUES OF PRMT,Y,AND DERY
              C      ALL CALCULATIONS INVOLVING DIFF. EQNS. USE DOUBLE PRECISION
ISN 0029      READ(5,22) WAVE
ISN 0030      22
              FORMAT(1P8E10.3)
ISN 0031      NDIM=2
ISN 0032      A=4.7824
ISN 0033      PRMT(1)=0.0
ISN 0034      PRMT(2)=518400.
ISN 0035      PRMT(3)=1799.99
ISN 0036      TXPLOT(1)=0.0
ISN 0037      TXPLOT(2)=6.0
ISN 0038      TYX(1)=7.0
ISN 0039      TYX(2)=1.0
ISN 0040      TYPLCT(1)=8.0
ISN 0041      TYPLCT(2)=1.0
ISN 0042      CALL CCP1PL(1.5,0.5,-3)
ISN 0043      WRITE(6,5)
ISN 0044      5
              FORMAT (1H1,6X,8HINTERVAL,11X,1HR,14X,2HS2,14X,2HS3,10X,
              1 10HWAVELENGTH,7)
ISN 0045      DO 6 I=1,87
ISN 0046      6
              WRITE(6,7)I,R(I),S2(I),S3(I),WAVE(I)
ISN 0047      7
              FORMAT(10X,I2,10X,1PD9.2,2(7X,1PD9.2),8X,0PF5.0)
ISN 0048      DO 2 M=3,7
ISN 0049      FLAT=LAT(M)*57.295
ISN 0050      FDEC=DEC(M)*57.295
ISN 0051      DO 3 K=1,2
ISN 0052      IF (M-1) 31,30,31
ISN 0053      31
              IF (M-7) 32,30,32
ISN 0054      32
              IF (K-2) 30,40,33
ISN 0055      33
              IF (K-4) 30,40,34
ISN 0056      34
              IF (K-6) 30,40,30
ISN 0057      30
              CONTINUE
ISN 0058      NIN=0
ISN 0059      Y(1)=X1(K,M)
ISN 0060      Y(2)=X3(K,M)
ISN 0061      PRMT(4)=.001*Y(2)
ISN 0062      DERY(1)=Y(2)/(Y(1)+Y(2))
ISN 0063      DERY(2)=1.0-DERY(1)
ISN 0064      W=0.0
ISN 0065      S=0.0
ISN 0066      WRITE(6,8)
ISN 0067      8
              FORMAT(1H1,7X,1H0,11X,2HC2,11X,2H03,8X,9H02 INTEG.,6X,4HTEMP,10X,
              12HB1,11X,2HB2,11X,2HB3,8X,3HALT ,/)

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Table A38 (Continued)

ISN 0068		DD 9 I=1,7
ISN 0069		JALT=95-5*I
ISN 0070	5	WRITE(6,10)X1(I,M),X2(I,M),X3(I,M),Z02(I,M),T(I,M),B1(I,M),B2(I,M), 1,B3(I,M),JALT,M
ISN 0071	10	FORMAT('P8D13:3,5X;I2,3X;I1)
	C	CHAPMAN FUNCTION COEFFICIENTS
	C	CHAPMAN VALUES ARE LINEAR EXTRAPOLATION OF DATA BETWEEN X=900 TO 1000
	C	AND THIS DATA IS EXTENDED FOR ALL OTHER X'S
	C	DATA IS FROM M.V. WILKES TABULATIONS.
ISN 0072		VB(1)=.00018
ISN 0073		VB(2)=.00023
ISN 0074		VB(3)=.00031
ISN 0075		VB(4)=.00043
ISN 0076		VB(5)=.00063
ISN 0077		VB(6)=.00095
ISN 0078		VB(7)=.00150
ISN 0079		VB(8)=.00253
ISN 0080		VB(9)=.00461
ISN 0081		VB(10)=.00916
ISN 0082		VB(11)=.02033
ISN 0083		VB(12)=.05143
ISN 0084		VB(13)=.132
ISN 0085		VA(1)= 5.590 -1000.*VB(1)
ISN 0086		VA(2)= 6.164 -1000.*VB(2)
ISN 0087		VA(3)= 6.867 -1000.*VB(3)
ISN 0088		VA(4)= 7.745 -1000.*VB(4)
ISN 0089		VA(5)= 8.870 -1000.*VB(5)
ISN 0090		VA(6)= 10.352 -1000.*VB(6)
ISN 0091		VA(7)= 12.378 -1000.*VB(7)
ISN 0092		VA(8)= 15.277 -1000.*VB(8)
ISN 0093		VA(9)= 19.670 -1000.*VB(9)
ISN 0094		VA(10)=26.830 -1000.*VB(10)
ISN 0095		VA(11)=39.648 -1000.*VB(11)
ISN 0096		VA(12)=65.505 -1000.*VB(12)
ISN 0097		VA(13)=120.00 -1000.*VB(13)
ISN 0098		CALL DRKGS(PRMT,Y,DERV,NDIM,IHLF,FCT,OUTP,AUX)
ISN 0099		FALT=95-5*K
ISN 0100		N=k
ISN 0101		SIZEX=N/12+1
ISN 0102		SIZES=SIZEX+1.
	C	PLOT RESULTS ON CALCCMP
ISN 0103		CALL CCP1PL(0.0,4.4,-3)
ISN 0104		CALL CCP5AX(0.0,0.0,11,HTIME(HOURS),-11,SIZEX,0.0,TXPLOT)
ISN 0105		CALL CCP5AX(0.0,0.0,26,HLCG(10) OF O CONCENTRATION,26,4.7,90.0, 1,TYPLOT)
ISN 0106		CALL CCP2SY(1.00,4.6,.15,4HLAT=,0.0,4)
ISN 0107		CALL CCP3NR(0.0,0.0,-.15,FLAT,0.0,2)
ISN 0108		CALL CCP2SY(2.75,4.6,.15,4HDEC=,0.0,4)
ISN 0109		CALL CCP3NR(0.0,0.0,-.15,FDEC,0.0,2)
ISN 0110		CALL CCP2SY(4.50,4.6,.15,4HALT=,0.0,4)
ISN 0111		CALL CCP3NR(0.0,0.0,-.15,FALT,0.0,2)
ISN 0112		CALL CCP6LN(XW,YW,N,1,TXPLCT,TYPLOT)
ISN 0113		CALL CCP1PL(0.0,-4.4,-3)
ISN 0114		CALL CCP5AX(0.0,0.0,33H LOG(10) OF O3 CONCENTRATION, 33,3.9, 1,90.0,TYX)
ISN 0115		CALL CCP5AX(0.0,0.0,11,HTIME(HOURS),-11,SIZEX,0.0,TXPLOT)
ISN 0116		CALL CCP6LN(XW,YW3,N,1,TXPLCT,TYX)
ISN 0117		CALL CCP1PL(SIZES,0.0,-3)
ISN 0118	40	CONTINUE
ISN 0119	3	CONTINUE
ISN 0120	2	CONTINUE
ISN 0121		STOP
ISN 0122		END

commands, W and S are used in controlling print-out. R = photon flux; S2 = O<sub>2</sub> absorb. cross-x; S<sub>3</sub> = O<sub>3</sub> absorb. cross-x, the 13 VA values are the y intercept values for the Chapman function approximations for 1° intervals from 80° to 93°. The 13VB values are the slopes of the Chapman function linear approximations for the 1° interval from 80° to 93°. PRMT, DERY, NDIM, AUX, Y, X are described in Table A42.

The values of R, S2, and Se were incorporated into the program in a Bloch DATA subprogram which was in common with the labelled common /BFLUX/.

#### Subroutine FCT TABLE A39

This subroutine is called by the integration routine DRKGS. It calculates the time rate of change of O<sub>3</sub> (DERY(2)) from the values of O(Y(2)) which have been calculated in subroutine DRKGS. Subroutine FCT calls Subroutine Flux to calculate the dissociation rate Jo<sub>2</sub>(F<sub>2</sub>) and Jo<sub>3</sub>(F<sub>3</sub>).

#### Subroutine Flux TABLE A40

This subroutine calculates the photodissociation rates Jo<sub>2</sub> and Jo<sub>3</sub> (F<sub>2</sub> and F<sub>3</sub>) as the solar zenith angle  $\chi$  changes. COSCHI = cos  $\chi$  where  $\chi$  is the solar zenith angle. COSCHI is calculated from the latitude, declination, and hour angle in subroutine flux. For  $\chi > 90^\circ$ , COSCHI is less than zero in ISN 13 so all dissociation rates are set to zero in ISN 14 and 15. For  $\chi > 90^\circ$ , the values are checked in ISN 17 to determine whether the Chapman function approximation is required. From ISN 22 to ISN 29 new values of the dissociation rates F<sub>2</sub> and F<sub>3</sub> are summed over contributions from all 87 wavelength intervals in which solar flux and cross section data were tabulated. ISN 26 checks for large negative arguments and sets contributions from such terms to zero to avoid an exponential underflow.

Table A39

ISN 0002	SUBROUTINE FCT(X,Y,DERY)
ISN 0003	DIMENSION Y(2),DERY(2),PRMT(5),AUX(8,2)
ISN 0004	DOUBLE PRECISION Y,DERY,PRMT,AUX,X,X2,Z02,B1,B2,B3,X1,X3,T,W,S, IRATE,VA,VB,A,F2,F3,LAT,DEC,S2,S3,R,H
ISN 0005	DOUBLE PRECISION DABS
ISN 0006	COMMON R,P
ISN 0007	COMMON/BFCT/ B1(7,8),R2(7,8),B3(7,8),X2(7,8),A
ISN 0008	COMMON /BFCCU/ RATE(5)
ISN 0009	COMMON /BFCFL/ F2(7,8),F3(7,8)
ISN 0010	COMMON /BOUFL/ LAT(8),DEC(8)
ISN 0011	COMMON/ BFLUX/ R(87),S2(87),S3(87),Z02(7,8),VA(13),VB(13),T(7,8)
	C RATE(1)=K1*O1*O2*M
	C RATE(2)=K3*O1*O3
	C RATE(3)=2.*K2*O1**2*M
	C RATE(4)=2.*J2*O2
	C RATE(5)=J3*O3
	C Y(1)=(O1)
	C Y(3)=(O3)
	C DERY(1)=RATE OF CHANGE OF (O1)
	C DERY(2)=RATE OF CHANGE OF (O3)
ISN 0012	CALL FLUX(R,S2,S3,Z02,F2,F3,X,DEC,LAT,K,M,VA,VB,T)
ISN 0013	RATE(1)=B1(K,M)*Y(1)*A*X2(K,M)**2
ISN 0014	RATE(2)=B3(K,M)*Y(1)*Y(2)
ISN 0015	RATE(3)=2.*B2(K,M)*Y(1)**2*A*X2(K,M)
ISN 0016	RATE(4)=2.*F2(K,M)*X2(K,M)
ISN 0017	RATE(5)=F3(K,M)*Y(2)
ISN 0018	DERY(1)=-RATE(1)-RATE(2)-RATE(3)+RATE(4)+RATE(5)
ISN 0019	DERY(2)=RATE(1)-RATE(2)-RATE(5)
ISN 0020	RETURN
ISN 0021	END

Table A40

ISN 0002		SUBROUTINE FLUX(R,S2,S3,ZC2,F2,F3,X,DEC,LAT,K,M,VA,VB,T)
ISN 0003		COMMON/BCHT/COSCHI,YW(300),YW3(300),XW(300),YS(49),YS3(49),XHR(49)
	C	SUBROUTINE FOR COMPUTING PHOTODISSOCIATION RATES J2=F2 J3=F3
	C	RT(J)=INCIDENT PHOTON FLUX IN THE JTH INTERVAL AT INFINITY
	C	S2=02 ABSORB. COEFF. S3=03 ABSORB. COEFF.
	C	Z02=INTEGRAL OF O2 FROM 300 KM TO ALTITUDE UNDER STUDY-US STAND ATM. SUPP.
ISN 0004		DIMENSION Y(2),DERY(2),PRMT(5),AUX(8,2)
ISN 0005		DOUBLE PRECISION Y,DERY,PRMT,AUX,X,X2,ZC2,B1,B2,B3,X1,X3,T,W,S, IRATE,VA,VB,A,F2,F3,LAT,DEC,S2,S3,R,H
ISN 0006		DOUBLE PRECISION SED,VX,CCSCHI,ARG,HANG
ISN 0007		DOUBLE PRECISION DSIN,DCCS,DATAN,DEXP
ISN 0008		DIMENSION T(7,8),Z02(7,8),R(87),S2(87),S3(87),VA(13),VB(13)
ISN 0009		DIMENSION F2(7,8),F3(7,8)
ISN 0010		DIMENSION LAT(8),DEC(8)
	C	DEC= SOLAR DECLINATION IN RADIANS LAT=LATITUDE IN RADIANS
	C	HANG= LOCAL HOUR ANGLE OF SUN MEASURED FROM NOON CHI= SOLAR ZENITH ANGLE.
ISN 0011		HANG=7.272205*10.**(-5)*X
ISN 0012		COSCHI=DSIN(LAT(M))*DSIN(DEC(M))+DCOS(LAT(M))*DCOS(DEC(M))* 1CCOS(HANG)
ISN 0013		IF(COSCHI)1,2,2
	C	ALL FLUX IS SET = 0.0 FOR CHI GREATER THAN 90.0 DEGREES
ISN 0014		F2(K,M)=0.0
ISN 0015		F3(K,M)=0.0
ISN 0016		GO TO 7
ISN 0017	2	IF (COSCHI-.17)3,3,4
	C	COSCHI LESS THAN .17 MEANS CHAPMAN FUNCTION REQUIRED
ISN 0018	3	COMPCH=DATAN(COSCHI)
ISN 0019		I=11.5-COMPCH*57.29578
ISN 0020		VX =220310./T(K,M)
	C	LINEAR APPROX. TO CHAP. FXN. AS FXN. OF X IS MADE
	C	DIFFERENT APPROX. IS USED FOR EACH ANGLE FROM 80 TO 92 DEGREES
ISN 0021		COSCHI = 1.7*(VA(I)+VB(I)*VX)
ISN 0022	4	F2(K,M)=0.0
ISN 0023		F3(K,M)=0.0
	C	COMPUTATION OF J2=F2 J3=F3
ISN 0024		DO 5 J=1,87
ISN 0025		ARG=S2(J)*Z02(K,M)/CCSCHI
ISN 0026		IF (ARG-90.)6,5,5
ISN 0027	6	SED =R(J)*DEXP(-ARG)
ISN 0028		F2(K,M) = F2(K,M)+S2(J)*SED
ISN 0029		F3(K,M) = F3(K,M)+S3(J)*SED
ISN 0030	5	CONTINUE
ISN 0031	7	CONTINUE
ISN 0032		RETURN
ISN 0033		END



Subroutine ØOTP TABLE A41

This subroutine controls the printing of the relevant data.

NIN = number °/ time that output has been called by the equation solving Subroutine DRKGS. Hence, it represents the number of integrations that have been required to that point.

X = TIME in seconds

CHI = Solar zenith angle in degrees calculated from the cosine of chi calculated in Subroutine Flux. The differential equation routine integrate for 6 days under the values of O<sub>3</sub> concentration are with 5% at successive noons in which case the program terminates by setting PRMI(5) = 5.0. This check and resultant termination results in ISN 29 through ISN 33.

ALAT, ADEC are the latitude and declination in degrees. A typical sheet of output is shown in Table A43.

H and SH = integration increment.

Subroutine DRKGS TABLE A42

This is a library subroutine furnished by IBM. It is the subroutine which actually integrates the differential equation. The subroutine has been modified slightly by including the variable H in the arguments of all Subroutine Output statements to transmit the integration interval. Also the card DRKG 0216 has been replaced by the card, 25 CONTINUE, to permit unlimited bisections of the initial integration interval.

Table A41

ISN 0002		SUBROUTINE OUTP(X,Y,DERY,IHLF,NDIM,PRMT,H)
ISN 0003		DIMENSION Y(2),DERY(2),PRMT(5),AUX(8,2)
ISN 0004		DIMENSION YA(8)
ISN 0005		DOUBLE PRECISION Y,DERY,PRMT,AUX,X,X2,ZOZ,B1,B2,B3,XI,X3,T,W,S, IRATE,VA,VB,A,F2,F3,LAT,DEC,S2,S3,R,H
ISN 0006		COMMON K,M
ISN 0007		COMMON/BCUTP/W,S,NIN
ISN 0008		COMMON/BCOOU/RATE(5)
ISN 0009		COMMON/BCUFL/LAT(8),DEC(8)
ISN 0010		COMMON/BCCHI/COSCHI,YW(300),YW3(300),XW(300),YS(49),YS3(49),XHR(49)
ISN 0011		DOUBLE PRECISION CCSCHI
ISN 0012		KIK=NIN+1
	C	RESULTS ARE PRINTED EVERY 1800 SECONDS (HALF HOUR)
ISN 0013		IF (X=1800.*W) I,2,2
ISN 0014	2	W=W+1.
ISN 0015		K=W
ISN 0016		CHI=57.3*DARCCS(CCSCHI)
	C	STORAGE OF RESULTS FOR CALCCMP
ISN 0017	35	YW(N)=DLG10(Y(1))
ISN 0018		YW3(N)=DLG10(Y(2))
ISN 0019		XW(N)=X/3600.
ISN 0020		IF (YW(N)=8.0) 33,34,34
ISN 0021	33	YW(N)=8.0
ISN 0022	34	IF (YW3(N)=7.0) 42,43,43
ISN 0023	42	YW3(N)=7.0
ISN 0024	43	CCONTINUE
	C	NEW PAGE EVERY 86400 SECS (24 HOURS)
ISN 0025		IF (X=86400.*S) 3,4,4
ISN 0026	4	S=S+1.
ISN 0027		IF (X=PRMT(2)) 41,23,23
ISN 0028	41	CCONTINUE
	C	IF X IS WITHIN 15 MINUTES OF TWELVE NOON, THE RESULTS OF SUCCESSIVE NOON VALUES ARE COMPARED
ISN 0029		MP=S
ISN 0030		YA(MP)=Y(2)
ISN 0031		IF (MP=1) 21,21,22
ISN 0032	22	IF (ABS(YA(MP)-YA(MP-1))-0.05*YA(MP)) 23,23,21
ISN 0033	23	PRMT(5)=5.0
ISN 0034		MSAC=XW(N-48)/24.
ISN 0035		TAC=MSAC
	C	STORAGE OF LAST 24 HOUR RESULTS FOR PUNCHING
ISN 0036		DO 30 I=1,49
ISN 0037		YS(I)=YW(N-49+I)
ISN 0038		YS3(I)=YW3(N-49+I)
ISN 0039		XHR(I)=XW(N-49+I)-24.*TAC
ISN 0040		IF (XHR(I)=24.) 31,31,32
ISN 0041	32	XHR(I)=XHR(I)-24.
ISN 0042	31	CCONTINUE
ISN 0043	20	CONTINUE
ISN 0044	21	CONTINUE
ISN 0045		J=95-5*K
ISN 0046		ALAT=LAT(M)*57.3
ISN 0047		ADEC=DEC(M)*57.3
ISN 0048		WRITE(6,7) J,ALAT,ADEC
ISN 0049	7	FORMAT(1H1,5X,4HALT=,I2,2HKM,5X,4HLAT=,OPF5.1,3HDEG,5X,4HDEC=,F4.1 1,3HDEG,777)
ISN 0050		WRITE(6,5)
ISN 0051	5	FORMAT(4X,4HTIME,9X, 1H0,9X,2HC3,7X,7HDERY(0),4X,8HDERY(03),3X,7HRATE(1),3X,7 2HRATE(2),3X,7HRATE(3),3X,7HRATE(4),3X,7HRATE(5),3X,3HCHI,3X, 35HDEL T,2X,5H# INT,2X,4HIHLF,7)
ISN 0052	3	CCONTINUE
	C	TIME IS GIVEN AS DAYS-HOURS-MINUTES
	C	CONVERSION OF TIME TO HOURS-MIN-SEC
ISN 0053		LX=X
ISN 0054		LDAY=LX/86400
ISN 0055		LHR=LX/3600-24*LDAY
ISN 0056		LMIN=LX/60-1440*LDAY-60*LHR
ISN 0057		SH=H
ISN 0058		WRITE(6,6) LDAY,LHR,LMIN,Y,DERY,RATE,CHI,SH,NIN,IHLF
ISN 0059	6	FORMAT(2X,I3,IH=,I2,IH=,I2,2(2X,1PD9.3), 11P2D11.2,1P5D10.2,2X,OPF5.1,2X,OPF5.0,2X,I5,2X,I3)
ISN 0060	1	CCONTINUE
ISN 0061		RETURN
ISN 0062		END

Table A42

C	SUBROUTINE DRKGS (PRMT, Y, DERY, NDIM, IHLF, FCT, OUTP, AUX)	DRKG 1070
C	.....	DRKG 10
C	.....	DRKG 20
C	.....	DRKG 30
C	SUBROUTINE DRKGS	DRKG 40
C	.....	DRKG 50
C	CARD DRKG 0216 WAS REPLACED BY A CARD (#25 CONTINUE) TO ALLOW UNLIMITED	
C	BISECTIONS OF THE INITIAL INTEGRATION INTERVAL IN ORDER TO REACH THE	
C	DESIRED ACCURACY	
C	PURPOSE	DRKG 60
C	TO SOLVE A SYSTEM OF FIRST ORDER ORDINARY DIFFERENTIAL	DRKG 70
C	EQUATIONS WITH GIVEN INITIAL VALUES.	DRKG 80
C	.....	DRKG 90
C	USAGE	DRKG 100
C	CALL DRKGS (PRMT, Y, DERY, NDIM, IHLF, FCT, OUTP, AUX)	DRKG 110
C	PARAMETERS FCT AND OUTP REQUIRE AN EXTERNAL STATEMENT.	DRKG 120
C	.....	DRKG 130
C	DESCRIPTION OF PARAMETERS	DRKG 140
C	PRMT - DOUBLE PRECISION INPUT AND OUTPUT VECTOR WITH	DRKG 150
C	DIMENSION GREATER THAN OR EQUAL TO 5, WHICH	DRKG 160
C	SPECIFIES THE PARAMETERS OF THE INTERVAL AND OF	DRKG 170
C	ACCURACY AND WHICH SERVES FOR COMMUNICATION BETWEEN	DRKG 180
C	OUTPUT SUBROUTINE (FURNISHED BY THE USER) AND	DRKG 190
C	SUBROUTINE DRKGS. EXCEPT PRMT(5) THE COMPONENTS	DRKG 200
C	ARE NOT DESTROYED BY SUBROUTINE DRKGS AND THEY ARE	DRKG 210
C	PRMT(1)- LOWER BOUND OF THE INTERVAL (INPUT),	DRKG 220
C	PRMT(2)- UPPER BOUND OF THE INTERVAL (INPUT),	DRKG 230
C	PRMT(3)- INITIAL INCREMENT OF THE INDEPENDENT VARIABLE	DRKG 240
C	(INPUT),	DRKG 250
C	PRMT(4)- UPPER ERROR BOUND (INPUT). IF ABSOLUTE ERROR IS	DRKG 260
C	GREATER THAN PRMT(4), INCREMENT GETS HALVED.	DRKG 270
C	IF INCREMENT IS LESS THAN PRMT(3) AND ABSOLUTE	DRKG 280
C	ERROR LESS THAN PRMT(4)/50, INCREMENT GETS DOUBLED.	DRKG 290
C	THE USER MAY CHANGE PRMT(4) BY MEANS OF HIS	DRKG 300
C	OUTPUT SUBROUTINE.	DRKG 310
C	PRMT(5)- NO INPUT PARAMETER. SUBROUTINE DRKGS INITIALIZES	DRKG 320
C	PRMT(5)=0. IF THE USER WANTS TO TERMINATE	DRKG 330
C	SUBROUTINE DRKGS AT ANY OUTPUT POINT, HE HAS TO	DRKG 340
C	CHANGE PRMT(5) TO NON-ZERO BY MEANS OF SUBROUTINE	DRKG 350
C	OUTP. FURTHER COMPONENTS OF VECTOR PRMT ARE	DRKG 360
C	FEASIBLE IF ITS DIMENSION IS DEFINED GREATER	DRKG 370
C	THAN 5. HOWEVER SUBROUTINE DRKGS DOES NOT REQUIRE	DRKG 380
C	AND CHANGE THEM. NEVERTHELESS THEY MAY BE USEFUL	DRKG 390
C	FOR HANDLING RESULT VALUES TO THE MAIN PROGRAM	DRKG 400
C	(CALLING DRKGS) WHICH ARE OBTAINED BY SPECIAL	DRKG 410
C	MANIPULATIONS WITH OUTPUT DATA IN SUBROUTINE OUTP.	DRKG 420
C	Y - DOUBLE PRECISION INPUT VECTOR OF INITIAL VALUES	DRKG 430
C	(DESTROYED). LATERON Y IS THE RESULTING VECTOR OF	DRKG 440
C	DEPENDENT VARIABLES COMPUTED AT INTERMEDIATE	DRKG 450
C	POINTS X.	DRKG 460
C	DERY - DOUBLE PRECISION INPUT VECTOR OF ERROR WEIGHTS	DRKG 470
C	(DESTROYED). THE SUM OF ITS COMPONENTS MUST BE	DRKG 480
C	EQUAL TO 1. LATERON DERY IS THE VECTOR OF	DRKG 490
C	DERIVATIVES, WHICH BELONG TO FUNCTION VALUES Y AT	DRKG 500
C	INTERMEDIATE POINTS X.	DRKG 510
C	NDIM - AN INPUT VALUE, WHICH SPECIFIES THE NUMBER OF	DRKG 520
C	EQUATIONS IN THE SYSTEM.	DRKG 530
C	IHLF - AN OUTPUT VALUE, WHICH SPECIFIES THE NUMBER OF	DRKG 540
C	BISECTIONS OF THE INITIAL INCREMENT. IF IHLF GETS	DRKG 550
C	GREATER THAN 10, SUBROUTINE DRKGS RETURNS WITH	DRKG 560
C	ERROR MESSAGE IHLF=11 INTO MAIN PROGRAM. ERROR	DRKG 570
C	MESSAGE IHLF=12 OR IHLF=13 APPEARS IN CASE	DRKG 580
C	PRMT(3)=0 OR IN CASE SIGN(PRMT(3)).NE.SIGN(PRMT(2)-	DRKG 590
C	PRMT(1)) RESPECTIVELY.	DRKG 600
C	FCT - THE NAME OF AN EXTERNAL SUBROUTINE USED. THIS	DRKG 610
C	SUBROUTINE COMPUTES THE RIGHT HAND SIDES DERY OF	DRKG 620
C	THE SYSTEM TO GIVEN VALUES X AND Y. ITS PARAMETER	DRKG 630
C	LIST MUST BE X, Y, DERY. SUBROUTINE FCT SHOULD	DRKG 640
C	NOT DESTROY X AND Y.	DRKG 650
C	OUTP - THE NAME OF AN EXTERNAL OUTPUT SUBROUTINE USED.	DRKG 660
C	ITS PARAMETER LIST MUST BE X, Y, DERY, IHLF, NDIM, PRMT.	DRKG 670
C	NONE OF THESE PARAMETERS (EXCEPT, IF NECESSARY,	DRKG 680
C	PRMT(4), PRMT(5), ...) SHOULD BE CHANGED BY	DRKG 690
C	SUBROUTINE OUTP. IF PRMT(5) IS CHANGED TO NON-ZERO,	DRKG 700
C	SUBROUTINE DRKGS IS TERMINATED.	DRKG 710
C	AUX - DOUBLE PRECISION AUXILIARY STORAGE ARRAY WITH 8	DRKG 720
C	ROWS AND NDIM COLUMNS.	DRKG 730
C	.....	DRKG 740

Table A42 (Continued)

C	REMARKS	DRKG 750
C	THE PROCEDURE TERMINATES AND RETURNS TO CALLING PROGRAM, IF	DRKG 760
C	(1) MORE THAN 10 BISECTIONS OF THE INITIAL INCREMENT ARE	DRKG 770
C	NECESSARY TO GET SATISFACTORY ACCURACY (ERROR MESSAGE	DRKG 780
C	IHLF=11);	DRKG 790
C	(2) INITIAL INCREMENT IS EQUAL TO 0 OR HAS WRONG SIGN	DRKG 800
C	(ERROR MESSAGES IHLF=12 OR IHLF=13),	DRKG 810
C	(3) THE WHOLE INTEGRATION INTERVAL IS WORKED THROUGH,	DRKG 820
C	(4) SUBROUTINE OUTP HAS CHANGED PRMT(5) TO NON-ZERO.	DRKG 830
C		DRKG 840
C	SUBROUTINES AND FUNCTION SUBPROGRAMS REQUIRED	DRKG 850
C	THE EXTERNAL SUBROUTINES FCT(X,Y,DERY) AND	DRKG 860
C	OUTP(X,Y,DERY,IHLF,NDIM,PRMT) MUST BE FURNISHED BY THE USER.	DRKG 870
C		DRKG 880
C	METHOD	DRKG 890
C	EVALUATION IS DONE BY MEANS OF FOURTH ORDER RUNGE-KUTTA	DRKG 900
C	FORMULAE IN THE MODIFICATION DUE TO GILL. ACCURACY IS	DRKG 910
C	TESTED COMPARING THE RESULTS OF THE PROCEDURE WITH SINGLE	DRKG 920
C	AND DOUBLE INCREMENT.	DRKG 930
C	SUBROUTINE DRKGS AUTOMATICALLY ADJUSTS THE INCREMENT DURING	DRKG 940
C	THE WHOLE COMPUTATION BY HALVING OR DOUBLING. IF MORE THAN	DRKG 950
C	10 BISECTIONS OF THE INCREMENT ARE NECESSARY TO GET	DRKG 960
C	SATISFACTORY ACCURACY, THE SUBROUTINE RETURNS WITH	DRKG 970
C	ERROR MESSAGE IHLF=11 INTO MAIN PROGRAM.	DRKG 980
C	TO GET FULL FLEXIBILITY IN OUTPUT, AN OUTPUT SUBROUTINE	DRKG 990
C	MUST BE FURNISHED BY THE USER.	DRKG1000
C	FOR REFERENCE, SEE	DRKG1010
C	RALSTON/WILF, MATHEMATICAL METHODS FOR DIGITAL COMPUTERS,	DRKG1020
C	WILEY, NEW YORK/LONDON, 1960, PP.110-120.	DRKG1030
C		DRKG1040
C	.....	DRKG1050
C		DRKG1060
C		DRKG1080
C		DRKG1090
ISN 0003	DIMENSION Y(1),DERY(1),AUX(8,1),A(4),B(4),C(4),PRMT(1)	DRKG1100
ISN 0004	DOUBLE PRECISION PRMT,Y,DERY,AUX,A,B,C,X,XEND,H,AJ,BJ,CJ,R1,R2, IDFLT	DRKG1110
ISN 0005	DOUBLE PRECISION DABS	DRKG1120
ISN 0006	DO I=1,NDIM	DRKG1130
ISN 0007	1 AUX(8,1)=.066666666666666667D0*DERY(I)	DRKG1140
ISN 0008	X=PRMT(1)	DRKG1150
ISN 0009	XEND=PRMT(2)	DRKG1160
ISN 0010	H=PRMT(3)	DRKG1170
ISN 0011	PRMT(5)=0.D0	DRKG1180
ISN 0012	CALL FCT(X,Y,DERY)	DRKG1190
C		DRKG1200
C	ERROR TEST	DRKG1210
ISN 0013	IF(H*(XEND-X))38,37,2	DRKG1220
C		DRKG1230
C	PREPARATIONS FOR RUNGE-KUTTA METHOD	DRKG1240
ISN 0014	2 A(1)=.5D0	DRKG1250
ISN 0015	A(2)=.29289321881345248D0	DRKG1260
ISN 0016	A(3)=1.7071067811865475D0	DRKG1270
ISN 0017	A(4)=.166666666666666667D0	DRKG1280
ISN 0018	B(1)=2.D0	DRKG1290
ISN 0019	B(2)=1.D0	DRKG1300
ISN 0020	B(3)=1.D0	DRKG1310
ISN 0021	B(4)=2.D0	DRKG1320
ISN 0022	C(1)=.5D0	DRKG1330
ISN 0023	C(2)=.29289321881345248D0	DRKG1340
ISN 0024	C(3)=1.7071067811865475D0	DRKG1350
ISN 0025	C(4)=.5D0	DRKG1360
C		DRKG1370
C	PREPARATIONS OF FIRST RUNGE-KUTTA STEP	DRKG1380
ISN 0026	DO I=1,NDIM	DRKG1390
ISN 0027	AUX(1,I)=Y(I)	DRKG1400
ISN 0028	AUX(2,I)=DERY(I)	DRKG1410
ISN 0029	AUX(3,I)=0.D0	DRKG1420
ISN 0030	3 AUX(6,I)=0.D0	DRKG1430
ISN 0031	IREC=0	DRKG1440
ISN 0032	H=H+H	DRKG1450
ISN 0033	IHLF=-1	DRKG1460
ISN 0034	ISTEP=0	DRKG1470
ISN 0035	IEND=0	DRKG1480
C		DRKG1490
C		DRKG1500
C	START OF A RUNGE-KUTTA STEP	DRKG1510
ISN 0036	4 IF((X+H-XEND)*H)7,6,5	DRKG1520
ISN 0037	5 H=XEND-X	DRKG1530
ISN 0038	6 IEND=1	DRKG1540

Table A42 (Continued)

	C	RECORDING OF INITIAL VALUES OF THIS STEP	DRKG1550
ISN 0039	7	CALL OUTP(X,Y,DERY,IREC,NDIM,PRMT,H)	DRKG1560
ISN 0040		IF(PRMT(5))40,8,40	DRKG1570
ISN 0041	8	ITEST=0	DRKG1580
ISN 0042	9	ISTEP=ISTEP+1	DRKG1590
	C		DRKG1600
	C	START OF INNERMOST RUNGE-KUTTA LOOP	DRKG1610
ISN 0043		J=1	DRKG1620
ISN 0044	10	AJ=A(J)	DRKG1630
ISN 0045		BJ=B(J)	DRKG1640
ISN 0046		CJ=C(J)	DRKG1650
ISN 0047		DO 11 I=1,NDIM	DRKG1660
ISN 0048		R1=H*DERY(I)	DRKG1670
ISN 0049		R2=AJ*(R1-BJ*AUX(6,I))	DRKG1680
ISN 0050		Y(I)=Y(I)+R2	DRKG1690
ISN 0051		R2=R2+R2+R2	DRKG1700
ISN 0052	11	AUX(6,I)=AUX(6,I)+R2-CJ*R1	DRKG1710
ISN 0053		IF(J-4)12,15,15	DRKG1720
ISN 0054	12	J=J+1	DRKG1730
ISN 0055		IF(J-3)13,14,13	DRKG1740
ISN 0056	13	X=X+.500*H	DRKG1750
ISN 0057	14	CALL FCT(X,Y,DFRY)	DRKG1760
ISN 0058		GOTO 10	DRKG1770
	C	END OF INNERMOST RUNGE-KUTTA LOOP	DRKG1780
	C		DRKG1790
	C	TEST OF ACCURACY	DRKG1800
ISN 0059	15	IF(ITEST)16,16,20	DRKG1810
	C		DRKG1820
	C	IN CASE ITEST=0 THERE IS NO POSSIBILITY FOR TESTING OF ACCURACY	DRKG1830
ISN 0060	16	DO 17 I=1,NDIM	DRKG1840
ISN 0061	17	AUX(4,I)=Y(I)	DRKG1850
ISN 0062		ITEST=1	DRKG1860
ISN 0063		ISTEP=ISTEP+ISTEP-2	DRKG1870
ISN 0064	18	IHLF=IHLF+1	DRKG1880
ISN 0065		X=X-H	DRKG1890
ISN 0066		H=.500*H	DRKG1900
ISN 0067		DO 19 I=1,NDIM	DRKG1910
ISN 0068		Y(I)=AUX(1,I)	DRKG1920
ISN 0069		DERY(I)=AUX(2,I)	DRKG1930
ISN 0070	19	AUX(6,I)=AUX(3,I)	DRKG1940
ISN 0071		GOTO 9	DRKG1950
	C		DRKG1960
	C	IN CASE ITST=1 TESTING OF ACCURACY IS POSSIBLE	DRKG1970
ISN 0072	20	IMOD=ISTEP/2	DRKG1980
ISN 0073		IF(ISTEP-IMOD-IMOD)21,23,21	DRKG1990
ISN 0074	21	CALL FCT(X,Y,DERY)	DRKG2000
ISN 0075		DO 22 I=1,NDIM	DRKG2010
ISN 0076		AUX(5,I)=Y(I)	DRKG2020
ISN 0077	22	AUX(7,I)=DERY(I)	DRKG2030
ISN 0078		GOTO 9	DRKG2040
	C		DRKG2050
	C	COMPUTATION OF TEST VALUE DELT	DRKG2060
ISN 0079	23	DELT=0.00	DRKG2070
ISN 0080		DO 24 I=1,NDIM	DRKG2080
ISN 0081	24	DELT=DELT+AUX(8,I)*DABS(AUX(4,I)-Y(I))	DRKG2090
ISN 0082		IF(DELT=PRMT(4))28,28,25	DRKG2100
	C		DRKG2110
ISN 0083	25	CONTINUE	DRKG2120
	C	ERROR IS TOO GREAT	DRKG2130
ISN 0084	26	DO 27 I=1,NDIM	DRKG2140
ISN 0085	27	AUX(4,I)=AUX(5,I)	DRKG2150
ISN 0086		ISTEP=ISTEP+ISTEP-4	DRKG2160
ISN 0087		X=X-H	DRKG2170
ISN 0088		IEND=0	DRKG2180
ISN 0089		GOTO 18	DRKG2190
	C		DRKG2200
	C	RESULT VALUES ARE GOOD	DRKG2210
ISN 0090	28	CALL FCT(X,Y,DERY)	DRKG2220
ISN 0091		DO 29 I=1,NDIM	DRKG2230
ISN 0092		AUX(1,I)=Y(I)	DRKG2240
ISN 0093		AUX(2,I)=DERY(I)	DRKG2250
ISN 0094		AUX(3,I)=AUX(6,I)	DRKG2260
ISN 0095		Y(I)=AUX(5,I)	DRKG2270
ISN 0096	29	DERY(I)=AUX(7,I)	DRKG2280
ISN 0097		CALL OUTP(X-H,Y,DERY,IHLF,NDIM,PRMT,H)	DRKG2290
			DRKG2300

Table A42 (Continued)

ISN 0098		IF (PRMT(5)) 40,30,40	DRKG2330
ISN 0099	30	DO 31 I=1,NDIM	DRKG2340
ISN 0100		Y(I)=AUX(I,I)	DRKG2350
ISN 0101	31	DERY(I)=AUX(2,I)	DRKG2360
ISN 0102		IREC=IHLF	DRKG2370
ISN 0103		IF (IEND) 32,32,39	DRKG2380
	C		DRKG2390
	C	INCREMENT GETS COUPLED	DRKG2400
ISN 0104	32	IHLF=IHLF-1	DRKG2410
ISN 0105		ISTEP=ISTEP/2	DRKG2420
ISN 0106		H=H+H	DRKG2430
ISN 0107		IF (IHLF) 4,33,33	DRKG2440
ISN 0108	33	IMOD=ISTEP/2	DRKG2450
ISN 0109		IF (ISTEP-IMOD-IMOD) 4,34,4	DRKG2460
ISN 0110	34	IF (DELT-.02DD*PRMT(4)) 35,35,4	DRKG2470
ISN 0111	35	IHLF=IHLF-1	DRKG2480
ISN 0112		ISTEP=ISTEP/2	DRKG2490
ISN 0113		H=H+H	DRKG2500
ISN 0114		GO TO 4	DRKG2510
	C		DRKG2520
	C	RETURNS TO CALLING PROGRAM	DRKG2530
ISN 0115	36	IHLF=11	DRKG2540
ISN 0116		CALL FCT(X,Y,DERY)	DRKG2550
ISN 0117		GO TO 39	DRKG2560
ISN 0118	37	IHLF=12	DRKG2570
ISN 0119		GO TO 39	DRKG2580
ISN 0120	38	IHLF=13	DRKG2590
ISN 0121	39	CALL OUTP(X,Y,DERY,IHLF,NDIM,PRMT,H)	DRKG2600
ISN 0122	40	RETURN	DRKG2610
ISN 0123		END	DRKG2620
			DRKG2630

Table A43.

ALT=85KM    LAT=-45.0DEG    DEC=23.5DEG														
TIME	0	03	DERV(0)	DERV(03)	RATE(1)	RATE(2)	RATE(3)	RATE(4)	RATE(5)	CHI	DEL T	# INT	TRFL	
5-0-7	6.485D 10	5.795D 07	1.50D 06	-1.87D 03	5.91D 05	3.08D 03	1.38D 04	1.52D 06	6.16D 05	68.6	450.	1156	2	
5-0-33	6.724D 10	5.972D 07	1.48D 06	1.28D 03	6.10D 05	3.12D 03	1.47D 04	1.50D 06	6.05D 05	69.0	225.	1164	3	
5-1-7	7.020D 10	6.265D 07	1.44D 06	-1.88D 03	6.39D 05	3.58D 03	1.62D 04	1.44D 06	6.63D 05	70.6	450.	1170	2	
5-1-33	7.241D 10	6.431D 07	1.37D 06	1.18D 03	6.56D 05	3.61D 03	1.70D 04	1.39D 06	6.51D 05	72.0	225.	1178	3	
5-2-7	7.507D 10	6.700D 07	1.25D 06	-2.20D 03	6.82D 05	4.08D 03	1.84D 04	1.24D 06	7.07D 05	75.1	450.	1184	2	
5-2-33	7.694D 10	6.834D 07	1.11D 06	5.56D 02	6.96D 05	4.07D 03	1.92D 04	1.12D 06	6.91D 05	77.3	225.	1192	3	
5-3-7	7.856D 10	7.053D 07	8.66D 05	-2.95D 03	7.16D 05	4.51D 03	2.03D 04	8.19D 05	7.44D 05	81.7	450.	1198	2	
5-3-33	8.015D 10	7.125D 07	6.35D 05	5.53D 02	7.24D 05	4.41D 03	2.07D 04	5.85D 05	7.19D 05	84.5	225.	1206	3	
5-4-7	8.100D 10	7.170D 07	3.26D 05	3.76D 03	7.31D 05	4.49D 03	2.12D 04	2.96D 05	7.25D 05	87.9	225.	1212	3	
5-4-33	8.043D 10	7.769D 08	-7.94D 05	6.77D 05	7.24D 05	5.73D 04	2.07D 04	0.0	0.0	93.1	225.	1228	3	
5-5-14	7.837D 10	2.313D 09	-8.66D 05	5.87D 05	7.00D 05	1.68D 05	1.94D 04	0.0	0.0	101.4	900.	1234	1	
5-5-44	7.678D 10	3.272D 09	-9.05D 05	4.99D 05	6.85D 05	2.17D 05	1.86D 04	0.0	0.0	106.4	900.	1236	1	
5-6-29	7.427D 10	4.499D 09	-9.45D 05	4.13D 05	6.55D 05	2.90D 05	1.69D 04	0.0	0.0	116.7	1800.	1238	0	
5-6-59	7.256D 10	5.198D 09	-9.62D 05	3.64D 05	6.55D 05	2.90D 05	1.69D 04	0.0	0.0	116.7	3600.	1239	0	
5-7-29	7.082D 10	5.814D 09	-9.72D 05	3.22D 05	6.23D 05	3.38D 05	1.54D 04	0.0	0.0	127.3	1800.	1240	0	
5-7-59	6.906D 10	6.360D 09	-9.76D 05	2.85D 05	6.23D 05	3.38D 05	1.54D 04	0.0	0.0	127.3	3600.	1241	0	
5-8-29	6.730D 10	6.843D 09	-9.76D 05	2.53D 05	5.91D 05	3.67D 05	1.38D 04	0.0	0.0	137.8	1800.	1242	0	
5-8-59	6.555D 10	7.272D 09	-9.72D 05	2.24D 05	5.91D 05	3.67D 05	1.38D 04	0.0	0.0	137.8	3600.	1243	0	
5-9-29	6.381D 10	7.653D 09	-9.65D 05	2.00D 05	5.60D 05	3.82D 05	1.24D 04	0.0	0.0	147.5	1800.	1244	0	
5-9-59	6.208D 10	7.992D 09	-9.54D 05	1.78D 05	5.60D 05	3.82D 05	1.24D 04	0.0	0.0	147.5	3600.	1245	0	
5-10-29	6.037D 10	8.296D 09	-9.42D 05	1.59D 05	5.29D 05	3.87D 05	1.11D 04	0.0	0.0	155.3	1800.	1246	0	
5-10-59	5.869D 10	8.567D 09	-9.28D 05	1.42D 05	5.29D 05	3.87D 05	1.11D 04	0.0	0.0	155.3	3600.	1247	0	
5-11-29	5.702D 10	8.810D 09	-9.12D 05	1.28D 05	5.00D 05	3.85D 05	9.88D 03	0.0	0.0	158.5	1800.	1248	0	
5-11-59	5.541D 10	9.028D 09	-8.95D 05	1.15D 05	5.00D 05	3.85D 05	9.88D 03	0.0	0.0	158.5	3600.	1249	0	
5-12-29	5.381D 10	9.224D 09	-8.77D 05	1.03D 05	4.71D 05	3.78D 05	8.79D 03	0.0	0.0	155.3	1800.	1250	0	
5-12-59	5.225D 10	9.401D 09	-8.58D 05	9.33D 04	4.71D 05	3.78D 05	8.79D 03	0.0	0.0	155.3	3600.	1251	0	
5-13-29	5.072D 10	9.560D 09	-8.39D 05	8.43D 04	4.44D 05	3.68D 05	7.80D 03	0.0	0.0	147.5	1800.	1252	0	
5-13-59	4.922D 10	9.705D 09	-8.20D 05	7.64D 04	4.44D 05	3.68D 05	7.80D 03	0.0	0.0	147.5	3600.	1253	0	
5-14-29	4.777D 10	9.836D 09	-8.00D 05	6.93D 04	4.18D 05	3.55D 05	6.91D 03	0.0	0.0	137.8	1800.	1254	0	
5-14-59	4.635D 10	9.955D 09	-7.80D 05	6.30D 04	4.18D 05	3.55D 05	6.91D 03	0.0	0.0	137.8	3600.	1255	0	
5-15-29	4.497D 10	1.006D 10	-7.60D 05	5.73D 04	3.93D 05	3.41D 05	6.12D 03	0.0	0.0	127.3	1800.	1256	0	
5-15-59	4.361D 10	1.016D 10	-7.41D 05	5.23D 04	3.93D 05	3.41D 05	6.12D 03	0.0	0.0	127.3	3600.	1257	0	
5-16-29	4.230D 10	1.025D 10	-7.21D 05	4.78D 04	3.70D 05	3.26D 05	5.42D 03	0.0	0.0	116.7	1800.	1258	0	
5-16-59	4.102D 10	1.033D 10	-7.02D 05	4.38D 04	3.70D 05	3.26D 05	5.42D 03	0.0	0.0	116.7	3600.	1259	0	
5-17-29	3.977D 10	1.041D 10	-6.83D 05	4.01D 04	3.48D 05	3.11D 05	4.79D 03	0.0	0.0	106.4	1800.	1260	0	
5-17-59	3.856D 10	1.048D 10	-6.64D 05	3.68D 04	3.48D 05	3.11D 05	4.79D 03	0.0	0.0	106.4	3600.	1261	0	
5-18-29	3.738D 10	1.054D 10	-6.45D 05	3.39D 04	3.27D 05	2.96D 05	4.23D 03	0.0	0.0	96.6	1800.	1262	0	
5-18-59	3.624D 10	1.060D 10	-6.27D 05	3.12D 04	3.27D 05	2.96D 05	4.23D 03	0.0	0.0	96.6	3600.	1263	0	
5-19-14	3.588D 10	1.063D 10	-6.18D 05	3.00D 04	3.17D 05	2.88D 05	3.97D 03	0.0	0.0	92.0	900.	1264	1	
5-19-33	3.495D 10	1.066D 10	-6.07D 05	2.85D 04	3.14D 05	2.86D 05	3.91D 03	0.0	0.0	90.8	225.	1266	3	
5-20-1	4.551D 10	4.076D 07	4.22D 05	-1.08D 03	4.12D 05	1.43D 03	6.70D 03	4.31D 05	4.10D 05	86.3	112.	1322	4	
5-20-33	4.664D 10	4.138D 07	7.31D 05	2.15D 03	4.22D 05	1.50D 03	7.05D 03	7.43D 05	4.19D 05	82.6	225.	1330	3	
5-21-7	4.839D 10	4.308D 07	1.00D 06	6.84D 02	4.41D 05	1.64D 03	7.68D 03	1.07D 06	4.40D 05	78.2	450.	1338	2	
5-21-33	5.011D 10	4.456D 07	1.17D 06	9.88D 02	4.54D 05	1.74D 03	8.17D 03	1.20D 06	4.52D 05	75.8	225.	1344	3	
5-22-7	5.264D 10	4.713D 07	1.32D 06	-2.27D 03	4.80D 05	2.06D 03	9.12D 03	1.36D 06	5.08D 05	72.6	450.	1350	2	
5-22-33	5.479D 10	4.869D 07	1.40D 06	1.20D 03	4.97D 05	2.08D 03	9.77D 03	1.43D 06	4.94D 05	71.0	225.	1358	3	
5-23-7	5.770D 10	5.159D 07	1.47D 06	-1.99D 03	5.26D 05	2.46D 03	1.10D 04	1.49D 06	5.52D 05	69.3	450.	1364	2	
5-23-33	6.002D 10	5.333D 07	1.49D 06	1.28D 03	5.45D 05	2.49D 03	1.17D 04	1.51D 06	5.41D 05	68.7	225.	1372	3	

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