

Groundwater arsenic and fluoride in Rajnandgaon District, Chhattisgarh, northeastern India

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Abstract The groundwater of Ambagarh Chouki, Rajnandgaon, India, shows elevated levels of As and F⁻, frequently above the WHO guidelines. In this work, the concentrations of As, F⁻, Na⁺, Mg²⁺, Ca²⁺, Cl⁻, SO₄²⁻, HCO₃⁻, Fe, dissolved organic carbon (DOC) and dissolved inorganic carbon (DIC) in the groundwater of Ambagarh Chouki are described. The sources of dissolved components in the groundwater are investigated using the cluster and factor analysis. Five factors have been identified and linked to processes responsible for the formation of groundwater chemistry. High concentrations of dissolved As seems to be linked to high concentrations of DOC, suggesting reductive dissolution of ferric oxyhydroxides as arsenic mobilization process. Fluoride is found in shallow depth water, presumably as a consequence of evaporation of water and removal of Ca²⁺ by precipitation of carbonates.

Keywords Arsenic · Fluoride · Trace element · Groundwater

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Introduction

Arsenic and fluoride are toxic elements presenting serious hazards for human health (Hileman 1988; Flanagan et al. 2012). The elevated concentrations of As in groundwater observed in different regions of South Asia is of an increasing concern because a large population is exposed to inorganic As through drinking water, contaminated food, dust, etc. (Chatterjee and Mukherjee 1999; Chakraborti et al. 2003; Acharyya et al. 2005; Patel et al. 2005; Agusa et al. 2006; Baig et al. 2009; Chauhan et al. 2009; Chatterjee et al. 2010; Kar et al. 2010; Kumar et al. 2010; Shukla et al. 2010; Winkel et al. 2011; Rodríguez-Lado et al. 2013). The fluoride contamination of groundwater was reported in several parts of India (Dutta et al. 2006; Meenakshi and Maheshwari 2006; Singh et al. 2007; Sankararamakrishnan et al. 2008; Suthar et al. 2008; Kundu and Mandal 2009; Yadav et al. 2009; Jha et al. 2010; Reddy et al. 2010). The environment of Ambagarh Chouki, Rajnandgaon, India has high levels of As (Acharyya et al. 2005; Patel et al. 2005; Shukla et al. 2010). Consumption of the contaminated water is the principal exposure pathway of As and F⁻ for humans in this region.

The groundwater As contamination in the Ambagarh Chouki block situated in the Rajnandgaon district of the state of Chhattisgarh was first reported by Chakraborti et al. (1999). Groundwater samples from 146 sites, Public Health Engineering Department (PHED) tube wells and dug wells in 22 villages of this block were analyzed. Arsenic concentration in tube well water was elevated above the WHO guideline by a factor of more than 10 with concentrations reaching 520 µg l⁻¹, whereas in the dug wells (general depth less than 50 m) As concentrations were as high as 880 µg l⁻¹. Medical study in that village indicated 42 % of adults and 9 % of children are suffering from arsenical skin

lesions. Also, epidemiological studies indicated high concentrations of As in urine, hair and nail in respective 89, 75 and 91 % of the village population. In the same area, i.e., at village Kourikasa, Rajnandgaon, district Chhattisgarh, Central India, an extensive study on the groundwater As over a period of 3 years has been conducted (1998–2000) based on the monitoring of As concentrations in hand pumps and dug wells in the region (Pandey et al. 2002).

Arsenic contamination in groundwater at variable concentration levels occurs at small patches in the Ambagarh Chouki block of Rajnandgaon district of the state of Chhattisgarh in Central India (Fig. 1a). Tube well waters are affected by high concentrations of As in some areas confined to N–S trending Dongargarh–Kotri rift zone which exposes the early Proterozoic metavolcanic rocks and co-magmatic granite (Sarkar 1957; Krishnamurthy et al. 1988; Acharyya et al. 2001; Acharyya 2002).

In the Ambagarh Chouki area, there is a rocky basement covered by alluvium, colluvium, and soil. Tube wells tap groundwater in the fracture system in bedrock. Previous studies (Pandey et al. 1999; Deshpande et al. 2010), mentioned panning for gold in the area and the presence of mine dumps around Bodal (20°40'N and 80°46'E), which are located close to Kaurikasa (20°43'9"N and 80°44'7"E), village where severe As–contamination occurs. It was assumed that these factors may have some connection to the source of As in groundwater. Some works also postulated that water in the principal river Sivnath, which flows through the As affected groundwater area, might be contaminated. Thus, Raipur, the capital city of Chhattisgarh state and townships like Rajnandgaon and Drug, located downstream, are likely to be at risk (Pandey et al. 1999). Later studies (Acharyya et al. 2001; Acharyya 2002) found that As pollution of groundwater is restricted to small isolated patches at village scale, and the Sivnath river water has As concentration below $10 \mu\text{g l}^{-1}$. The Chhattisgarh basin, located to the east of the Dongargarh rift zone, and exposing mid-late Proterozoic platform-type clastic and carbonate sediments, is totally free of As contamination. Recent investigations (Pandey et al. 2006; Deshpande et al. 2010; Shukla et al. 2010) in the Kourikasa, Rajnandgaon, Chhattisgarh, found elevated concentrations of As in the wells located in acid volcanic rocks, close to shear zones as well as in granites.

The aim of the study is to consolidate the information and database on the groundwater resources for exploitation of As and F^{-} -safe drinking water in the affected hotspots in selected areas of Rajnandgaon district, Central India. The specific goal of this project was to develop an approach for targeting As and F^{-} -safe aquifers for drinking purposes to: (1) identify As and other trace element baselines in the region, and (2) to develop a scientific database for the identification of As and F^{-} -safe aquifers. Within the framework of this project, we also aim to use the

hydrogeological and geochemical data as a basis for the evaluation of the risks associated with these aquifers, and thereby their sustainability as safe drinking water source.

Materials and methods

Study area

The Kourikasa area covers over $\approx 500 \text{ km}^2$ (20°43'9"N and 80°44'7"E). The sampling locations are shown in Fig. 1b. Total 20 ground water samples in each season from this region were collected in post-monsoon period, November 2010 and pre-monsoon period, May, 2011 as described in the literature (Nielsen and Nielsen 2006).

Analysis

The field parameters, i.e., pH, temperature (T), electrical conductance (EC), dissolved oxygen (DO) and reduction potential (RP) were measured on site. The total hardness (TH) and alkalinity values were analyzed by titration methods (Nollet and De Gelder 2007). The total dissolved solid (TDS) value was determined by evaporation of filtered samples (through glass fiber filter) and drying (APHA 2005).

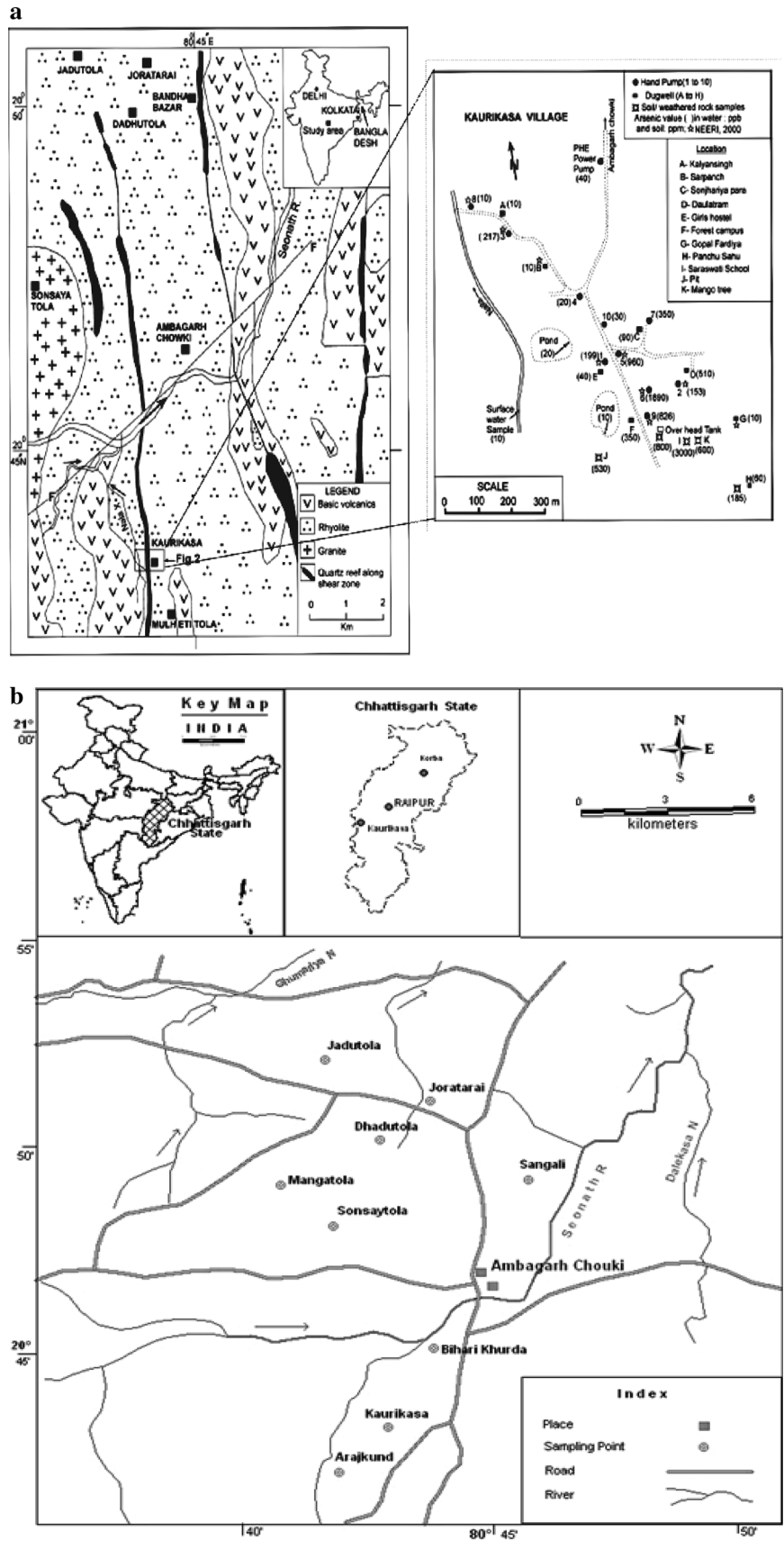
GBC AAS932 equipped with the HG3000 hydride generator was used for the determination of total As. Detection limit of $1 \mu\text{g l}^{-1}$ As was routinely achieved and precision for three replicate measurements was better than $\pm 1.0 \%$.

The CHNSO-IRMS Analyzer by SV Instruments Analytica Pvt. Ltd. was used for analysis of the dissolved organic carbon (DOC) and dissolved inorganic carbon (DIC). The total dissolved carbon (TDC) was oxidized with O_2 at 1020 °C with constant helium flow using WO_3 as catalyst. The resulting CO_2 was monitored by the conductivity detector. The DOC content was determined by oxidation of the prior HCl-treated water sample as described above. The DIC content was computed by subtracting the DOC from the TDC ($\text{DIC} = \text{TDC} - \text{DOC}$).

The Dionex DX120 Ion Chromatograph (Dionex Corporation, Sunnyvale, CA, USA) equipped with anion separation column (AS9-HC, $250 \times 4 \text{ mm}$), cation separation column (CS12A, $250 \times 4 \text{ mm}$) and conductivity detector was used for analysis of the principal ions, i.e., Na^+ , K^+ , Mg^{2+} , Ca^{2+} , Cl^- , SO_4^{2-} and NO_3^- .

Fluoride was determined by Orion Ion Analyzer-720 equipped with the ion-selective electrode using the buffer. The buffer (TISAB-III) solution was added to both samples and standards. TISAB-III was prepared by dissolving 300 g sodium citrate (FW = 294), 22 g of 1,2-cyclohexanediamine N,N,N,N-tetra-acetic acid (CDTA) and 60 g of NaCl in a volume of 1000 ml of water (pH 5.0–5.5). The

Fig. 1 a Geological map of Rajnandgaon District in Chhattisgarh state of Central India with the location of Kourikasa village (*right*). **b** Sampling locations of groundwater in Kourikasa area, Rajnandgaon, CG, India



TISAB-III regulated, adjusted and masked the ionic strength of the solutions, the pH and interferences by polyvalent cations such as Al(III), Fe(III) and Si(IV), respectively. In the solution, CDTA formed stable complexes with polyvalent metal cations which were more stable than metal-fluoride complexes. The electrode was selective for the F^- ions by several orders of magnitude.

The exchangeable sodium ratio (ESR), sodium hazard (SH), magnesium hazard (MH) and water quality index (WQI) were computed for rating of the water quality. The ESP, and SH values are calculated as follows:

$$ESP = \{ (Na) / (Ca + Mg + K + Na) \} \times 100$$

$$SH = \{ \{ [Na] + [K] \} / \{ [Na] + [K] + [Mg] + [Ca] \} \} \times 100$$

where all ions are expressed in $meq\ l^{-1}$.

Magnesium plays important role in irrigation water, and magnesium hazard (MH) is calculated using following equation.

$$MH = \{ [Mg^{2+}] / [Mg^{2+}] + [Ca^{2+}] \} \times 100$$

where all ions are expressed in $mol\ l^{-1}$.

Ten parameters, i.e., pH, EC, DO, hardness, alkalinity, Mg^{2+} , Ca^{2+} , Cl^- , NO_3^- and SO_4^{2-} were selected for evaluation of WQI using the standard values recommended by BIS and WHO. The weighed arithmetic method was used for calculation of the WQI of the groundwater with the help of the following expression.

$$WQI = \sum q_n W_n / \sum W_n$$

where

$$q_n = 100(V_o - V_{io}) / (V_s - V_{io})$$

q_n = Quality rating of the n th water quality parameter, V_n = Estimated value of the n th parameter of a given water, S_n = Standard permissible value of the n th parameter, V_{io} = Ideal value of the n th parameter of pure water (i.e., 0 for all other parameters except pH and dissolved oxygen (7.0 and 14.6 $mg\ l^{-1}$, respectively), W_n = Unit weight for the n th parameter, K = Proportionality constant.

Multivariate statistics (i.e., Cluster and Factor analysis) was used to differentiate the water samples on the basis of their composition and origin (Simeonov et al. 2003; Boyacioglu and Boyacioglu 2008). The software STATISTICA 7.1 window was used for the statistical analysis.

Results and discussion

Physical characteristics of tube well and groundwater

The physico-chemical characteristics of the groundwater are summarized in Table 1. All tube wells in the study area

are deeper, ranging from 52 to 90 m. The age of tube wells was in the range of 1–20 Years. Temperature was ranging from 27.1 to 27.6 °C with a mean value of 27.3 ± 0.2 °C.

In this study, pH values of groundwater ranged from 6.4 to 7.9 with a mean value of 6.8 ± 0.1 . The EC and TDS values ranged from 940 to 1427 $\mu S\ cm^{-1}$ and 574 to 5373 $mg\ l^{-1}$ with mean values of $1192 \pm 84\ \mu S\ cm^{-1}$ and $2323 \pm 595\ mg\ l^{-1}$, respectively. The EC value correlated well ($r = 0.88$) with the sum of total concentration of groundwater ions, Σ_{ions} (i.e., Na^+ , K^+ , Mg^{2+} , Ca^{2+} , F^- , Cl^- , NO_3^- , SO_4^{2-} and HCO_3^-). Among 20 tube wells studied, the highest EC and TDS values were recorded in tube well no. 19. The mean EC and TDS value of groundwater in the studied area was found to be two- and fourfolds higher than the respective permissible limit of 300 $\mu S\ cm^{-1}$ and 500 $mg\ l^{-1}$ (BIS 2003; WHO 2011). The RP and DO value of the water ranged from 240 to 392 mV and 4.1 to 7.3 $mg\ l^{-1}$ with mean values of $334 \pm 17\ mV$ and $4.8 \pm 0.3\ mg\ l^{-1}$, respectively. The alkalinity of the water ranged from 209 to 742 $mg\ l^{-1}$ with a mean value of 514 $mg\ l^{-1}$ as $CaCO_3$. The alkalinity value of water in all tube wells was found to be higher than the recommended value of 120 $mg\ l^{-1}$ (BIS 2003; WHO 2011).

Hardness is a term relating to the concentrations of certain metallic ions in water, particularly Mg and Ca, and is usually expressed as a concentration of dissolved calcite ($CaCO_3$). Durfor and Becker (1964) developed the following classification for water hardness: soft water, 0–60 $mg\ l^{-1}$; moderately hard water, 61–120 $mg\ l^{-1}$; hard water, 121–180 $mg\ l^{-1}$; and very hard water, over 180 $mg\ l^{-1}$. The total hardness of the water ranged from 103 to 529 $mg\ l^{-1}$ with a mean value of $250 \pm 57\ mg\ l^{-1}$. The water of 65 % tube wells was found to be very hard ($>180\ mg\ l^{-1}$) due to higher Mg^{2+} and Ca^{2+} concentrations.

Chemical characteristics of groundwater

The chemical characteristics of the groundwater are summarized in Table 2. The DOC is a very important component in biogeochemical cycling of elements being an indicator for the groundwater pollution. The DOC and DIC concentration were relatively high, ranging from 107 to 1640 and 110 to 3120 $mg\ l^{-1}$ with mean values of 695 ± 198 and $958 \pm 318\ mg\ l^{-1}$, respectively. The higher DIC value in the groundwater of this region was observed, due to oxidation of DOC into CO_2 to form carbonic acid. The DOC value was observed several folds higher than the recommended value of 2.0 $mg\ l^{-1}$.

Very high concentration of F^- in the groundwater of the studied area was observed due to dissolution of the fluoride minerals associated to the bed rocks. The F^- concentration was found to be several folds higher than permissible limit of 1.5 $mg\ l^{-1}$, ranging from 3.7–27 with a mean value of

Table 1 Physico-chemical characteristics of tube well and groundwater, November 2010

S. No.	Location	Age	Depth	T (°C)	pH	EC	TDS	RP	DO
1	Bharī khurda	11	65	27.3	6.8	1215	945	323	4.6
2	Bharī khurda	12	72	27.2	6.5	1150	670	368	4.9
3	Kaudī kasa	13	60	27.2	7.2	967	1116	352	4.3
4	Kaudī kasa	20	70	27.3	6.8	1321	1062	324	4.1
5	Kaudī kasa	17	60	27.3	7.0	1204	3567	307	5
6	Kaudī kasa	15	52	27.3	7.9	1105	1112	311	4.1
7	Arajkund	15	60	27.3	7.2	1406	2615	319	4.6
8	Arajkund	7	67	27.1	6.8	1280	4678	343	4.7
9	Kaudī kasa	13	75	27.3	6.7	1130	1007	364	4.4
10	Kaudī kasa	8	65	27.2	6.9	940	1989	271	4.4
11	Kaudī kasa	17	67	27.3	6.5	1040	3272	327	4.1
12	Sansaytola	17	52	27.4	6.4	1148	3016	368	4.3
13	Sansaytola	14	70	27.4	6.4	1070	3961	392	4.5
14	Sansaytola	9	64	27.4	6.8	1167	1065	293	4.5
15	Dhadītola	8	60	27.5	6.6	1204	2472	240	6.2
16	Joratarāī	8	90	27.6	6.7	1160	2655	377	4.4
17	Jadutola	1	75	27.4	7.1	1320	2492	352	4.8
18	Mangatola	7	60	27.3	6.7	1427	2543	317	5.5
19	Mangatola	9	75	27.5	6.8	1267	3110	368	7.3
20	Sangalī	1	54	27.3	7.1	1310	1100	373	4.4

Table 2 Chemical characteristics (mg l^{-1}) of groundwater, November 2010

S. No.	OC	CC	F ⁻	Cl ⁻	SO ₄ ²⁻	NO ₃ ⁻	HCO ₃ ⁻	Na ⁺	K ⁺	Mg ²⁺	Ca ²⁺	Fe	As*
1	310	310	4.5	17	16	3.1	207	16	1	11	38	1.32	220
2	210	110	3.7	46	30	15	192	17	0.7	12	34	1.12	217
3	460	160	5.7	43	22	17	281	36	0.9	26	64	1.68	510
4	190	190	10	103	51	37	301	42	1.2	49	88	1.78	419
5	1204	1604	19.5	89	60	20	366	42	11.8	47	104	2.25	985
6	310	310	24	39	15	5.2	323	23	2.5	26	34	1.2	179
7	970	970	11	78	62	4.5	348	40	1.1	42	88	1.14	208
8	1020	3120	15	96	59	12	201	53	0.3	32	70	1.56	489
9	330	330	12	50	30	14	146	24	1.4	22	48	1.62	376
10	870	870	16	14	14	4.3	140	21	1.8	16	22	1.45	315
11	1260	1360	8.7	78	45	17	329	80	15.3	13	66	1.8	620
12	1070	1270	19.3	71	29	62	393	34	0.6	21	46	1.72	575
13	1640	1640	12	67	37	51	385	51	0.4	28	50	1.9	826
14	405	405	23.2	28	14	9.2	116	25	0.4	13	26	0.9	148
15	110	1310	14	135	84	134	412	76	0.5	58	138	2.16	906
16	560	1560	7.4	39	45	24	277	57	3	27	56	1.92	721
17	1010	1010	8.3	21	43	2.7	327	27	0.7	12	30	2.06	805
18	756	1015	14.6	128	72	73	299	31	0.5	64	90	1.65	587
19	1103	1103	25.4	195	62	34	366	43	1.2	73	104	2.1	795
20	107	507	27	39	17	17	305	27	0.7	15	38	1.25	210

* = $\mu\text{g l}^{-1}$

14.1 ± 3.1 (BIS 2003; WHO 2011). In 65 % tube wells, the F^- concentration was found above 10.0 mg l^{-1} .

Chloride in ground water may originate from various sources including: the dissolution of halite. The recommended tolerance limit for Cl^- reported is 250 mg l^{-1} (BIS 2003; WHO 2011). The concentration of Cl^- in the groundwater ranged from 14 to 195 mg l^{-1} with a mean value of $69 \pm 20 \text{ mg l}^{-1}$. The remarkably higher concentration of Cl^- observed in the water of tube well nos. 4, 15, 18 and 19 may be due to sewage contamination. The concentration of Na^+ in the groundwater ranged from 16 to 80 mg l^{-1} with a mean value of $38 \pm 8 \text{ mg l}^{-1}$. The ionic ratio of $[Na^+]/[Cl^-]$ in the water ranged from 0.3 to 2.3 with a mean value of 1.1 ± 0.3 . It indicates that sodium concentrations exceed chloride concentrations in many samples (45 %), suggesting dissolution of silicate minerals. The K^+ concentration in the water ranged from 0.3 to 15.3 mg l^{-1} with mean value of $2.3 \pm 1.8 \text{ mg l}^{-1}$. The ionic ratio of $\{Na^+]/[K^+]$ in the water ranged from 9 to 78 with a mean value of 39 ± 10 , similar to natural water.

Sulfate (SO_4^{2-}) is commonly observed in ground water. The established permissible limit for sulfate is 150 mg l^{-1} (BIS 2003; WHO 2011). The concentration of SO_4^{2-} in the groundwater of the studied area ranged from 14 to 84 mg l^{-1} with a mean value of $40 \pm 9 \text{ mg l}^{-1}$. The SO_4^{2-} content in water has a fair correlation ($r = 0.80$) with the NO_3^- and Cl^- content, suggesting origin from a similar sources.

Nitrate (NO_3^-) is the most frequently detected drinking-water contaminant in this region, and nitrate levels of less than 0.2 mg l^{-1} are considered to represent background levels. Concentrations ranging from 0.21 to 3.0 mg l^{-1} are considered transitional. Concentrations between 3.1 and 10 mg l^{-1} represent elevated concentrations due to human activities. The NO_3^- concentration in the water of the studied area ranged from 2.7 to 134 mg l^{-1} with a mean value of $28 \pm 14 \text{ mg l}^{-1}$. The water of 20 % tube wells was contaminated with NO_3^- beyond the permissible limit of 45 mg l^{-1} (BIS 2003; WHO 2011).

Iron (Fe) concentrations commonly exceeded the recommended limit of 0.3 mg l^{-1} in all tube wells (BIS 2003; WHO 2011). The Fe concentration ranged from 0.90 to 2.25 mg l^{-1} with mean value of $1.63 \pm 0.20 \text{ mg l}^{-1}$.

Arsenic concentrations in water of all tube wells were above the WHO drinking water guideline value ($10 \mu\text{g l}^{-1}$), ranging from 148 to $985 \mu\text{g l}^{-1}$ with a mean value of $506 \pm 118 \mu\text{g l}^{-1}$ (WHO 2011).

Seasonal variations

The pH value in the pre-monsoon period (summer season) was found to increase slightly in the range of 6.7–8.3 with a mean value of 7.2 ± 0.2 due to slight increase of alkaline nature of the ground water. The concentrations of As, F^- ,

Cl^- , SO_4^{2-} , Na^+ , K^+ , Mg^{2+} , Ca^{2+} and Fe increased up to 41, 36, 14, 38, 26, 39, 43, 23 and 69 %, respectively, in the pre-monsoon period 2011, possibly due to lowering of the water table and increasing water temperature ($\approx 27\text{--}29 \text{ }^\circ\text{C}$) (Fig. 2). However, the concentration of the NO_3^- in the pre-monsoon period was decreased (21 %) probably due to limited anthropogenic inputs.

Water quality and toxicity

The concentrations of F^- , As and Fe were several folds higher than the values recommended by the WHO. Several cases of arsenic poisoning, i.e., hypopigmentation, melanosis, keratosis, etc., in people of Kaurikasa area have been reported. Chronic exposure to fluoride contaminated water has led to development of skin, dental and skeleton fluorosis in humans and domestic animals. The Water Quality Index (WQI) value for 20 groundwater tube wells ranged from 90 to 317 with a mean value of 172 ± 28 , indicating unsuitability for drinking purpose.

The value of exchangeable sodium ratio (ESR), sodium hazard (SH) and magnesium hazard (MH) values of the water ranged from 12 to 42 %, 12 to 47 % and 25 to 56 % with a mean value of $24 \pm 3 \%$, $25 \pm 4 \%$ and $44 \pm 3 \%$, respectively. The water of the studied area was found to be strongly sodic with moderate values of SH and MH. These

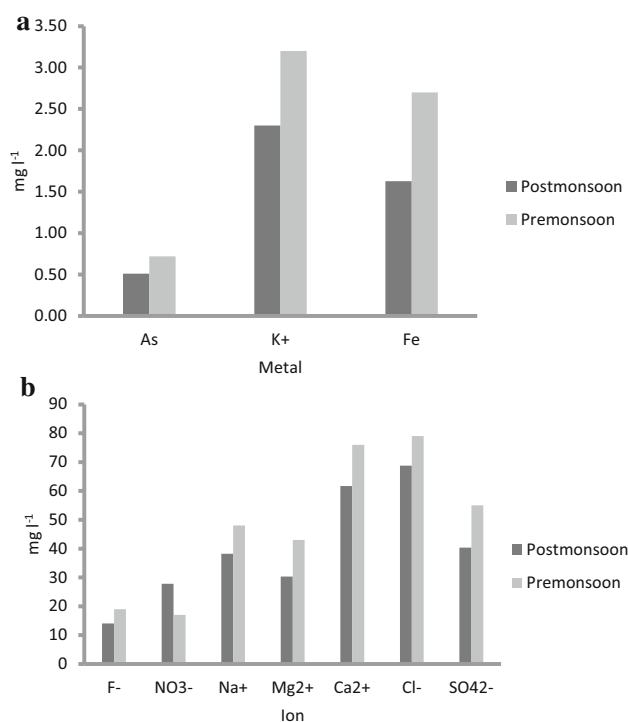


Fig. 2 Seasonal variation of arsenic and other elements in the groundwater

data reveal that groundwater could be used for irrigation purposes.

Cluster analysis

Cluster analysis was performed on the dataset by Ward’s method using square Euclidean distances as similarity measure (Davies 1986). Figure 3 shows the dendrogram of the physico-chemical characteristics of groundwater in the Ambagarh Chouki. Two groups of parameters are observed. The first group (Group-I) includes major cations (i.e., Ca^{2+} , Mg^{2+} , Na^+ and K^+), anions (i.e., SO_4^{2-} , Cl^- , NO_3^- and F^-) and the physical parameters (i.e., pH, T, depth and age), exhibiting strong mutual correlation. The second group (Group-II) contains parameters EC, TDS, HCO_3^- , As, DOC and CC, which have some relationship with carbon species in groundwater. This suggests that high As concentration in groundwater is linked to DOC and, thus, to its redox status.

The dendrogram of groundwater samples also presents two clusters (Fig. 4). Cluster-I comprises the samples from the tube wells TW5, TW7, TW8, TW10–13, TW15–19 (12 tube wells). Cluster-II contains the samples from TW1–4, TW6, TW9, TW14 and TW20 (8 tube wells). All the parameters differentiate the two groups of tube well water, except pH, T, K^+ and F^- . The parameters such as RP, DO, EC, TDS, As, Fe, OC, CC, Na^+ , Mg^{2+} , Ca^{2+} , Cl^- , SO_4^{2-} , NO_3^- and HCO_3^- have higher median values in cluster-I than cluster-II (Fig. 4).

Factor analysis

This analysis is a multivariate statistical method, aiming to explain the general relationship between measured variables by computing multivariate factors that may help to

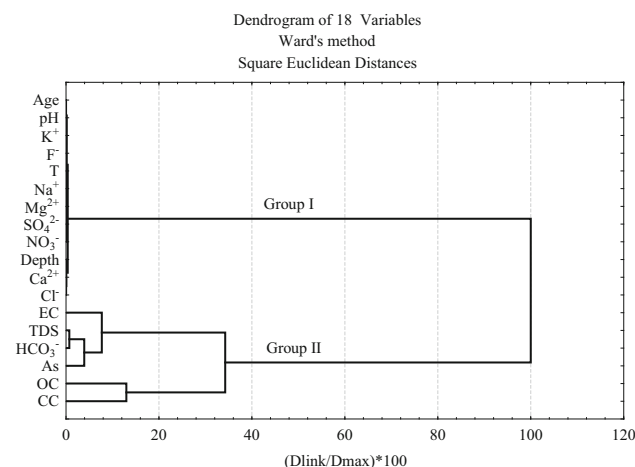


Fig. 3 Dendrogram of the physico-chemical characteristics of groundwater in Ambagarh Chouki

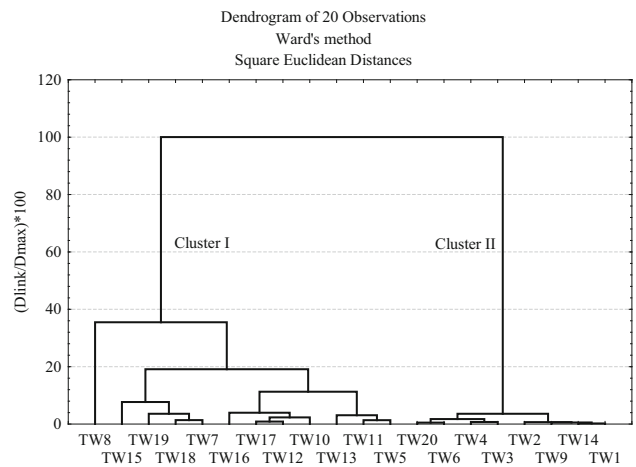


Fig. 4 Dendrogram of sampling sites in Ambagarh Chouki (TW Tube Well)

better classify the original data. Eighteen variables were included in this analysis. First, the raw data were standardized which gave the mean and variance of the variables equal to zero and one, respectively. In reference to the eigenvalues, five factors were extracted after varimax rotation as they have eigenvalues greater than 1 (Kaiser 1958). Table 3 presents the eigenvalues, the percentage of variance and the cumulative percentage of variance associated with each factor. The terms “strong”, “moderate” and “weak” were applied to factor loadings, referring to absolute loading values of >0.70 , $0.70-0.50$ and $0.50-0.30$, respectively. Factor-1 explains 42.37 % of the total variance (Table 3) and has strong positive loadings on EC, TDS, Mg^{2+} , Ca^{2+} , Cl^- and SO_4^{2-} . The EC values are positively well correlated with the content of anions and cations. Thus, Factor-1 can reflect the mineralization of the groundwater.

Factor-2 accounts for 13.56 % of the total variance and has high positive loading on depth and high negative loading on F^- . These values show that the two parameters are in opposite relation, i.e., F^- concentration decreases with depth of wells.

Factor-3 accounts for 10.24 % of the total variance and has strong positive loadings on DOC, CC and a moderate loading on As. A moderate correlation exists between As and Na^+ , Mg^{2+} , Ca^{2+} , Cl^- , SO_4^{2-} , NO_3^- , HCO_3^- , OC, and CC ($0.48 \leq r \leq 0.63$). Factor-3 is related to DOC, a controlling factor of CC content in the groundwater. Complex processes could be related to Factor-3, such as mobilization of As by reductive dissolution (Liu et al. 2003).

Factor-4 accounts for 7.98 % of the total variance and has a strong loading on age of wells and a moderate loading on K^+ . No significant correlation exists between K^+ and others major ions. This suggests that K^+ is derived mostly from K-feldspar dissolution (Jalali 2007). Factor-4 may

Table 3 Eigen values, percent of variance, cumulative percent of variance for the factor analysis of hydro chemical data in Central India

Parameter	Factor 1	Factor 2	Factor 3	Factor 4	Factor 5
Age	0.09	0.03	-0.09	-0.87	-0.04
Depth	-0.06	0.78	0.17	0.26	0.04
T	0.11	0.12	0.04	0.18	0.85
pH	-0.07	-0.66	-0.18	0.05	-0.31
EC	0.94	-0.18	0.13	-0.08	0.11
TDS	0.95	-0.18	0.13	-0.08	0.12
As	0.41	0.22	0.56	-0.06	0.51
Na ⁺	0.47	0.27	0.47	-0.30	0.35
K ⁺	0.03	-0.01	0.48	-0.68	-0.06
Mg ²⁺	0.93	-0.08	0.02	0.07	0.14
Ca ²⁺	0.93	0.06	0.08	-0.18	0.19
F ⁻	0.14	-0.80	0.14	0.28	0.14
Cl ⁻	0.92	-0.02	0.15	0.00	0.16
SO ₄ ²⁻	0.86	0.23	0.28	0.01	0.15
NO ₃ ⁻	0.57	0.10	-0.11	0.01	0.64
HCO ₃ ⁻	0.45	-0.27	0.27	-0.30	0.61
OC	0.02	0.00	0.86	-0.17	0.05
CC	0.27	0.13	0.84	0.16	0.01
Eigen value	7.60	2.53	1.83	1.41	1.17
% Variance	42.23	14.06	10.19	7.85	6.51
% Cumulative	42.23	56.29	66.48	74.33	80.84

Absolute loading values >0.70, significant at $p < 0.05$

The bold values indicates the stronger factor loading (> 0.7)

represent the dissolution of K-minerals (or K-feldspars) which is related to the age of tube wells.

Factor-5 accounts for 6.68 % of the total variance and has a strong loading on temperature, T. The correlation between T and As is weak ($r = 0.48$, $p < 0.05$). Temperature is a parameter which could control the rate of some processes such as the release of As by reductive dissolution into groundwater.

Conclusions

The toxic elements, i.e., As and F⁻ occur in several folds higher concentrations than the WHO permissible limits. The ground water mineralization also comprises species such as Mg²⁺, Ca²⁺, Cl⁻ and SO₄²⁻. In addition, the groundwater is strongly sodic in nature with high WQI value. These hydrochemical data indicate that the groundwater is unsuitable for human consumption. Factor analysis yielded five factors, which were linked to water chemistry-forming processes. Fluoride contamination of the groundwater seems to be related to the tube wells installed at shallow depth and, thus, to enrichment of shallow ground water by evaporation. The presence of As

in groundwater seems to be caused by reductive dissolution linked to the presence of dissolved organic matter.

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