Annual and semi-annual S_q variations at 96° MM MAGDAS I and II stations in Africa

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The quiet daily variations of the geomagnetic field (S_q) have been studied by using the geomagnetic data of the 10 International Quietest Days. Geomagnetic data from MAGDAS I and II stations—ten stations were selected around the 96° magnetic meridian (MM)—over Africa have been analyzed from September 2008 to August 2009. The analysis of this unique data set clarifies the characteristics of the S_q variations over the African sector for the first time by observational geomagnetic data. In this paper, we found that $S_q(H)$ and $S_q(D)$ show a predominantly annual variation. The most interesting result, and a new finding in the African sector, is the appearance of two vortices in the day lit southern hemisphere during spring. Further data analysis also revealed many other interesting characteristics for the S_q over the African sector.

Key words: S_q , northward current, ionospheric currents, ground-based magnetometers.

1. Introduction

On days which are quiet with respect to solar terrestrial activity phenomena, the geomagnetic field variations are called solar quiet, or quiet daily, variations ' S_q ' (Campbell, 1982). The system of horizontal electric currents in the ionosphere could produce this S_q field (Chapman and Bartels, 1940). It is considered that the S_q current system consists of four sets of vortices, two in each hemisphere; in both the northern or southern hemisphere, one vortex is situated in the dayside, while the other extends over the nightside. The dayside vortex represents the S_q equivalent current. The dayside or ' S_q current' vortex of the northern hemisphere is counter clock-wise, while the vortex of the southern hemisphere is clock-wise (Takeda, 2002).

The S_q field changes are caused by the dynamo effect of diurnal and semidiurnal tidal winds in the lower ionosphere around an altitude of 100 km, that are generated in situ by solar heating during the daylight hours (Matsushita, 1977). On the other hand, numerical simulation for S_q fields has shown that fluctuations in the neutral wind propagated below can cause significant day-to-day variations in S_q fields even though in-situ solar heating is in a steady state (Miyahara and Ooishi, 1997). The distribution of S_q mainly depends on the latitude and the local time (Chapman and Bartels, 1940).

The study of electric currents in the ionosphere is not easy because of the difficulty in making direct in situ measurements of these currents, which is why collecting data from ground-based magnetometers is so important. The magnetic field perturbations are usually resolved along a geomagnetically north-south (positive north), east-west (positive east), and parallel to B, directions and are denoted as H, D, and Z components, respectively. Sometimes a geographic coordinate system is used, in which case the symbols X, Y, and Z denote the magnetic perturbations in the north, east, and vertical (positive down) directions. It is impossible to derive the true horizontal ionospheric current distribution uniquely from ground magnetic perturbations, since they are a superposition of contributions from the hor-

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MAGDAS/CPMN

(MAGnetic Data Acqusition System/Circum-pan Pacific Magnetometer Network)



Fig. 1. A station map of MAGDAS stations used in the study.

izontal ionospheric currents, field-aligned currents, distant currents in the magnetosphere, and currents induced in the Earth's surface. For these reasons, the ground magnetic perturbations are usually expressed in terms of 'equivalent' ionospheric currents.

It should be noted that determination of the datum level of the geomagnetic quiet daily variation is difficult because of a lack of knowledge of the electromagnetic state of the night-side ionosphere. However, some researchers (e.g. Matsushita and Maeda, 1965) assumed that the zero level is the daily mean of S_q field variations, allowing for some amount of night-side S_q currents. Other researchers (e.g. Matsushita, 1967) preferred to use the mean of nocturnal values as the zero level. In that case, a derived S_q current system exists only in the day-side hemisphere. In-situ observations of the S_q current by rocket measurements did not produce evidence supporting the existence of a night-side S_q current system (Onwumechili, 1992).

The S_q variations were studied all over the world by using observational data in the sectors of the world that have stations, but the African sector was studied only by using the simulation results, because of the absence of a magnetometer array covering the African sector. In 2008, MAGnetic Data Acquisition System (MAGDAS) magnetometers were installed over Africa (Yumoto *et al.*, 2009), and (Yumoto *et al.*, 2010), which has enabled the study of the S_q over Africa. In the present paper, by using this unique data set, S_q variations over the African sector will be discussed for the first time as observational results. The geomagnetic quiet interval from September 2008 to August 2009 has been analyzed by using this data set, and the annual and semiannual S_q variations and other features have been studied.

2. The Data and Analysis Method

The data set analyzed in this study was obtained from ten ground-based magnetometer stations in Africa along the 96° MM [magnetic meridian] longitudinal sector. The stations are L'aquila (LAQ, G.M.Lat = 36.25° , G.M.Lon = 87.56° ; Fayum (FYM, 25.76° , 112.65°); Aswan (ASW, 15.20°, 104.24°); Khartoum (KRT, 5.69°, 103.8°); Nairobi (NAB, -10.65°, 108.18°); Dalelsalam (DES, -16.26°, 110.59°); Lusaka (LSK, -26.06°, 98.32°); Maputo (MPT, -35.98°, 99.1°); Durban (DRB, -39.21°, 96.1°); Hermanus (HER, -42.29°, 82.20°). Further information about the stations is available at the website of the MAGDAS/CPMN project http://www.serc.kyushu-u.ac.jp/ index_e.html. The coordinates of the stations used in this study are listed in Table 1 and their corresponding locations on the world map are shown in Fig. 1. The stations are arranged in Table 1 according to their latitudes. The data for solar quiet days were selected according to the 10 International Quietest Days per month from the German Research Center for Geosciences (GFZ): http://www-app3. gfz-potsdam.de/kp_index/qddescription.html.

In this paper, we have analyzed only the H and D components of the geomagnetic field. The daily variation of the H component at each station can be represented by a Fourier series as follows:

$$\Delta H(t) = a_0 + \sum_{m=1}^{4} \left(a_m \cos \frac{2\pi}{T} m t + b_m \sin \frac{2\pi}{T} m t \right), \ (1)$$

where a_m and b_m are Fourier coefficients which are calculated by using a least-square fitting. *T* is equal to 24, which is the total number of hours per day. *t* is the sample time (local time). Δ represents the deviation of the geomagnetic field from the baseline. The baseline was calculated in this paper as the average value for the night-time hours from 2300 LT to 0300 LT for each station, as we assumed that

Station	Code	G. G Lat.	G. G Lon.	G. M Lat.	G. M Lon.
L'Aquila	LAQ	42.38	13.32	36.25	87.56
Fayum	FYM	29.18	35.50	25.76	112.65
Aswan	ASW	23.59	32.51	15.20	104.24
Khartoum	KRT	15.33	32.32	5.69	103.8
Nairobi	NAB	-1.16	36.48	-10.65	108.18
Dal Es Salaam	DES	-6.47	39.12	-16.26	110.59
Lusaka	LSK	-15.23	28.20	-26.06	98.32
Maputo	MPT	-25.57	32.36	-35.98	99.57
Durban	DRB	-29.49	30.56	-39.21	96.1
Hermanus	HER	-34.34	19.24	-42.29	82.20
Addis Ababa	AAB	9.04	38.77	0.18	110.47

Table 1. Geographic and geomagnetic coordinates of MAGDAS stations used in the study.



Fig. 2. The equivalent S_q current system for the yearly average of all the stations of the African sector.

there is no night time current. Similarly, the daily variation of the D component, $\Delta D(t)$ was represented by a Fourier series. We analyzed the mean hourly values of all-day solar daily variations of $S_q(H)$ by using Eq. (1) as a function of local time, and Fourier coefficients with order m = 0-4. The daily variations of S_q (D) were analyzed by a similar method. The longitudinal (i.e., local time) variations can be described by a Fourier analysis from Eq. (1), while the latitudinal distribution of ΔH and ΔD can be estimated by fitting Legendre polynomials with degree n = 0-6 for each local time (t). The magnetic field vectors derived from the magnetometer data of each station were rotated clockwise through 90° to indicate a current flow direction so that the equivalent current system of the S_q could be plotted as a function of local time and latitude. The stations used in this paper are located between geomagnetic latitudes from 36°N to 42°S. The data were divided into four seasonal groups: Fall includes the months from September to November; S-Summer from December to February; Spring from March to May; and N-Summer from June to August.

3. Results of the Data Analysis

In this paper, the results will represent the characteristics of the solar quiet geomagnetic variations in the African sector and the seasonal, semi-annual and annual variations of the equivalent S_q current system. Our observational results show that:

(1) Figure 2 shows the S_q equivalent current system plotted for a one-year average of the magnetic field data from September 2008 till August 2009. In this figure,

there is asymmetry in shape between the northern vortex and the southern vortex of the current system. An asymmetry in intensity was also found in the yearly average of the horizontal component intensity of the S_q variations in the following way: two different groups of conjugate stations were used to calculate the yearly average of the horizontal component intensity of the magnetic field. The horizontal component intensity was calculated with Eq. (2).

horizontal component intensity =
$$\sqrt{H^2 + D^2}$$
. (2)

The horizontal component intensity for each hemisphere was taken as the average of the data from the selected three stations which represent each hemisphere at each local time. The first group of these conjugate stations (see Table 2) consisted of three pairs of 'geographic' conjugate stations: three representing the northern hemisphere, while the other three represent the southern hemisphere. From the data of these three geographic conjugate stations, the yearly average of the horizontal component intensity shows an asymmetry between the northern hemisphere and the southern hemisphere as shown in Fig. 3, where the black line (N-H) represents the intensity of the horizontal component of the northern hemisphere, while the grey line (S-H) represents the intensity of the horizontal component intensity of the southern hemisphere. The other group of conjugate stations consists of three pairs of 'geomagnetic' conjugate stations (see Table 3), and the intensity of the horizontal component tends to be



Fig. 3. Comparison between the northern and the southern horizontal component intensity for geographic conjugate stations.



The horizontal component intensity of geomagnetic conjugate stations

Fig. 4. Comparison between the northern and the southern horizontal component intensity for geomagnetic conjugate stations.

Table 2. Station names and their geographic latitudes for the three pairs of geographic conjugate stations used in this study.

Northern hemisphere	Southern hemisphere
Fayum (29.18°)	Durban (-29.49°)
Aswan (23.59°)	Maputo (-25.57°)
Khrtoum (15.33°)	Lusaka (-15.23°)

symmetrical between the northern and southern hemispheres (see Fig. 4). Here, also, the black line (N– H) represents the intensity of the horizontal component of the northern hemisphere, while the grey line (S–H) represents the intensity of the horizontal component intensity of southern hemisphere. From this result, one can conclude that the geomagnetic equator in the African is very close to the real axis of S_q .

Table 3. Station names and their geomagnetic latitudes for the three pairs of magnetic conjugate stations used in this study.

Northern hemisphere	Southern hemisphere
L'Aquila (36.25°)	Maputo (-35.98°)
Fayum (25.76°)	Lusaka (-26.06°)
Aswan (15.20°)	Dal Es Salaam (-16.26°)

(2) Another interesting result is the disappearance of the northern hemisphere vortex during daytime hours (see Fig. 5). The absence of the northern vortex was accompanied by the appearance of the northward current in the northern hemisphere in the morning, which then reversed in the afternoon. Normally, the current system in the northern hemisphere is directed southward in the morning and northward in the afternoon. The



Fig. 5. The equivalent S_{q} current system for February which shows clearly the northward currents.



Fig. 6. The monthly average shown in Panel (a) of the *H* component and Panel (b) of the *D* component for February 2009.

disappearance of the northern hemisphere vortex was observed from January to March. The corresponding daily variations of this current can be found in Fig. 6. Panel (a) of Fig. 6 represents the H component of the magnetic field, and Panel (b) represents the D component of the magnetic field. In Fig. 6, the H component shows a normal behavior while the D component shows an abnormal behavior as the westward field in the morning hours dominates at all the stations. Normally, the $S_{a}(D)$ in the northern hemisphere represents a well-defined eastward field in the morning hours, changing to a westward field during the noon hours (i.e. normally, the $S_q(D)$ in the northern hemisphere is higher than the zero level in the morning sector and lower than the zero level in the afternoon). Our results in Fig. 6 show that the variations of the D component are lower than the zero level in the morning



Fig. 7. The monthly average shown here in Panel (a) of the H component and Panel (b) of the D component for February 2009 after adding the AAB station.

sector, which implies the existence of the northward current. By using incoherent scatter sound, Mazaudier and Venkateswaran (1985) also observed a northward current over Saint-Santin.

(3) A dip equatorial station—Addis Ababa (0.57 dip latitude)—belonging to the 96° MM was added to the stations used in this study so that the equatorial current system in Africa could also be discussed. Figure 7 represents the monthly average of the magnetic field for February at all the African chain stations. Panel (a) represents $S_q(H)$, while Panel (b) represents $S_q(D)$. According to Chapman (1951), the amplitude of $S_q(H)$ near the dip equator is higher than the amplitude of $S_q(H)$ at other latitudes due to large eastward ionospheric currents in the dynamo region over the dip equator (i.e., equatorial electrojet). In Panel (a), the daily maximum of the $S_q(H)$ at Addis Ababa appears



Fig. 8. The equivalent S_q current system for May.



Fig. 9. Focus on the morning vortex with a different local time scale.

between 1030 to 1130 LT, while at other low latitude stations the maximum appears around noon. According to Rastogi and Iyer (1976), during the minimum sunspot years, the maximum amplitude of $S_{a}(H)$ appears before noon in the range between 1030 to 1130 LT. On the other hand, during the maximum sunspot years, the maximum of the amplitude of $S_{a}(H)$ appears around noon. Since the results we represent here are for a minimum sunspot year, our results from Addis Ababa show such phase shift. One explanation for this phase shift in the maximum value of $S_q(H)$ to occur before noon was suggested by Rastogi and Iyer (1976). The integrated current in the dynamo region can be expressed as the product of the electron number density (N_e) , the electronic charge (e), and the westward electron drift velocity (v), so the current should be a maximum between the time of the maximum of the electron number density (N_e) and the westward electron drift velocity (v). The electron density becomes maximum around noon while the electric field (i.e. The westward electron drift velocity) has its maximum value near 9–10 LT. Thus, the suggested product maximum is around 11 LT. In Panel (b), $S_q(D)$ at the station near the dip equator (i.e. AAB) shows a variation from the night-time level during 0600–1600 LT.

- (4) Figure 8 shows the appearance of two vortices on the sunlit side of the southern hemisphere in spring, which is the most interesting result and a new finding of the S_q current system in the African sector. The first vortex started early in the morning around 0600 LT and lasted until 0900 LT, while the other vortex had a focus around noon. Figure 9 shows the first vortex more clearly in a different time scale. A possible explanation is that it may be caused as a result of the counter electrojet (Gurubaran, 2002), but further research needs to be done to define this phenomena.
- (5) From the data of the seasonal variations of both northern and southern hemispheres, it was found that there



Fig. 10. The seasonal variation of all stations of the northern hemisphere for the H component, and the D component.



Fig. 11. The seasonal variation of all stations of the southern hemisphere for the H component, and the D component.

is an annual variation with the maximum amplitude of $S_q(H)$ and $S_q(D)$ occurring around the local summer of each hemisphere. In Figs. 10 and 11, the daily range averaged for different seasons is plotted. As shown in Fig. 10, in the northern hemisphere, the highest peak between the four seasons appears in the Northern Summer, for both H and D components. On the other hand, as shown in Fig. 11, the highest peak for the southern hemisphere appears in the Southern Summer. The absence of a semiannual variation in the seasonal variations and the domination of the annual variation for both H and D components were noticed.

4. Summary and Conclusion

- (1) There is an asymmetry in the shape of the S_q current pattern between the southern and northern hemispheres based on the yearly average calculations. The asymmetry in intensity appears in the yearly average of the horizontal component intensity of the S_q variations.
- (2) The absence of the northern current vortex and the appearance of the northward current in the northern hemisphere during morning were found.
- (3) At the dip equatorial station (AAB), the daily maximum value of the H component has been shifted to

the range between 1030 to 1130 LT during a minimum sunspot year.

- (4) The most interesting result and a new finding from the analysis of the data from the African sector is the appearance of two vortices on the sunlit side of the southern hemisphere in spring.
- (5) Domination of the annual variation for both H and D components and the absence of semiannual variations were noticed.

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