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DEVELOPMENT OF A THEORY OF THE
SPECTRAL REFLECTANCE OF MINERALS

PART IV

FINAL REPORT

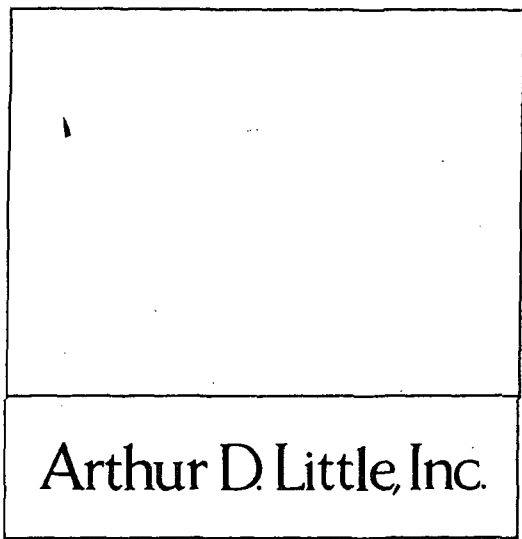
by

J. R. ARONSON, A. G. EMSLIE, E. M. SMITH

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Arthur D. Little, Inc.

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Abstract

This report describes the results of a research program directed toward the development of a theory of the spectral reflectance or emittance of particulate minerals. The theory is expected to prove invaluable in the detailed interpretation of the remote infrared spectra of planetary surfaces. The principal results are described in two manuscripts that have been submitted for publication (Sections II and III). Some further improvements and other results are discussed in other sections.

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I. INTRODUCTION

The results of our research program directed at the development of a theory of the spectral reflectance of minerals have been discussed in a series of reports^{1,2,3} as the work has proceeded. These progress reports have continually updated the theory as successive refinements were made. During the course of the work certain lines of investigation were pursued while others were abandoned without perhaps being sufficiently explored. Frequently, improvements in the fit to certain spectral features were made at the expense of other details, but as the improvements were almost always in the nature of a less approximate treatment, and as the apparent fit showed distinct overall improvements, the authors believe that the theory has continually improved. A thorough description of the details of the work, as of the summer of 1972, has been submitted for publication in Applied Optics. This description takes the form of two papers. The first provides a coherent description of the entire theory and the second shows the results of comparison of that theory with experimental data. For the most part these are up-to-date and are therefore herein presented as the bulk of the final report on this work. Nonetheless in the few months remaining after their submission some further improvements were made in the theory. They will be discussed along with some topics not treated in the papers. It is the authors' intention to write a further paper to discuss these topics and improvements when they have sufficiently matured.

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3. J. R. Aronson, A. G. Emslie, L. H. Roach, E. M. Smith and P. C. von Thuna, "Development of a Theory of the Spectral Reflectance of Minerals, Part III," Report to NASA, Manned Spacecraft Center, January 1972 (Contract NAS 9-10875).

II. The Spectral Reflectance and Emittance
of Particulate Materials,

Part I - Theory

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Abstract

The infrared spectral reflectance of a semi-infinite medium composed of irregular particles of different materials is calculated in terms of the sizes, shapes, and complex refractive indices of the particles. For particles larger than the wavelength the scattering and absorption are computed mainly by geometrical optics but with important wave-optical corrections for the additional absorption caused by edges and asperities, which are represented by dipoles distributed over the surface of the particle.

For particles smaller than the wavelength a Lorentz-Lorenz model is used to derive the average complex index of the medium, the particles being treated as ellipsoids with a wide range of shapes. The average scattering of an individual ellipsoidal particle is then found from the relative refractive index of the particle with respect to the Lorentz-Lorenz medium. For both large and small particles the single-particle scattering is represented by six discrete beams. Calculation of the reflectance is then facilitated by a radiative transfer method that also involves six beams. For particles of intermediate size a suitable formula bridging the results for large and small particles is found to be satisfactory.

INTRODUCTION

Many theoretical attempts have been made to explain the reflection spectra of clouds and powders.^{1,2} Most authors calculate the scattering and absorption of the individual particles by applying the well-known Mie³ theory to spheres of the same volume as the actual particles.⁴ The reflectance of the cloud or powder is then calculated by a two-beam theory such as that of Schuster⁵ or Kubelka and Munk,⁶ by a many-beam method usually based on the general theory of Chandrasekhar,⁷ or by a Monte Carlo technique.⁸ Blevin and Brown⁹ have demonstrated that multiple-beam calculations, in which scattering and absorption are regarded as spatially continuous, remain approximately valid even when the particles are almost in contact. Non-continuum models have, however, been proposed by Lathrop¹⁰ and Melamed.¹¹

This paper presents a theory of spectral reflectance in which two additional physical mechanisms are taken into account, namely, the effect of particle shape and, in the case of closely-spaced

fine particles, the effect of coherent interaction between the particles. Inclusion of these two effects is essential to provide good agreement between theoretical and experimental spectra, as shown in Part II.¹² In addition, the theory incorporates a number of mathematical innovations.

In order to allow for shape features such as edges and surface asperities on large particles ($d \geq \lambda$) we abandon Mie theory and revert to a combination of geometrical optics to obtain the main scattering and absorption, and wave optics to obtain the effect of the edges and asperities, both regarded as induced dipoles. This we will refer to as the coarse particle theory.

In the case of fine closely-spaced particles ($d \ll \lambda$), we also do not use Mie theory but regard the particles as randomly oriented ellipsoidal dipoles which cause both coherent scattering (Fresnel component) and incoherent scattering (Rayleigh component). This we call the fine particle theory. The coarse and fine theories are connected by an empirical bridging formula.

COARSE PARTICLE THEORY

As a model for the coarse particle theory we represent the large-scale shape of a particle by a sphere of the same volume as the actual particle and the small-scale shape, including edges and other surface asperities, as a distribution of dipoles spread uniformly over the surface of the sphere. This model evolved from the practical requirement of mathematical tractability and from the consideration that, for chunky particles at least, the scattering and absorption crosssections depend critically on the small-scale shape but only slightly on the large-scale shape. Final justification of the model appears in the good agreement between theory and experiment. One might think that an equivalent sphere would be a poor representation of a platelet-shaped particle but even in this extreme case the theory gives results in good agreement with observation.

1. Calculation of Single-Particle Scattering and Absorption

Crosssections

We first consider a parallel beam of radiation incident on the equivalent sphere which represents a given particle. Fig. 1 shows reflected and refracted rays produced by an incident ray that strikes the sphere at an angle of incidence θ .

In order to allow for the effect of inhomogeneous wave propagation that occurs owing to the complex index of refraction

$$m = n - ik \quad (1)$$

we define the effective angle of refraction ψ as the direction of the Poynting vector in the particle, given by the formula

$$\tan \psi = \frac{\sin \theta}{\text{Re} (m^2 - \sin^2 \theta)^{1/2}} \quad (2)$$

where Re signifies "real part". Eq. (2) reduces to the usual law of refraction when m is real, i.e., when the absorption index k is zero.

The angle of deviation ϕ of the transmitted ray is given by the relation

$$\phi = 2 \cdot | \theta - \psi | \quad (3)$$

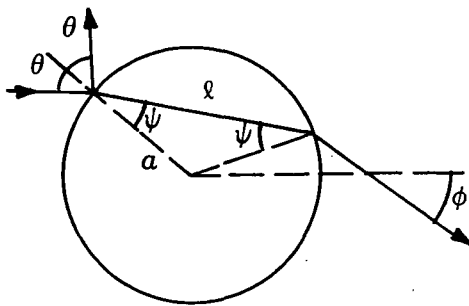


FIGURE 1 RAY TRACING GEOMETRY

where the absolute value signs allow for the case where $n < 1$.

The transmission path l is given by the equation

$$l = 2a \cos\psi \quad (4)$$

where a is the radius of the sphere.

Let I_0 be the intensity of the unpolarized incident parallel beam. Then the power incident on the sphere in the angular range $d\theta$ is

$$dP_0 = \pi a^2 I_0 \sin 2\theta \, d\theta \quad (5)$$

The corresponding refracted power is

$$(dP)_{\text{refr}} = \pi a^2 I_0 T_0^2 T \sin 2\theta \, d\theta \quad (6)$$

where T_0 is the surface transmittance which includes the effects of Fresnel reflection and absorption by asperities and edges, and T is the transmission factor given by

$$T = e^{-\frac{4\pi k l}{\lambda}} \quad (7)$$

where λ is the free-space wavelength of the radiation. An expression for T_0 is derived later.

The angular distribution (6), considered as function of the angle of deviation ϕ given by (3), is the contribution of refraction to the single particle phase function. To facilitate radiative transfer calculations we represent the continuous distribution (6) approximately by a set of six mutually perpendicular discrete beams comprising a forward-scattered beams, a backward-scattered beam and four transversely-scattered beams. For this purpose we need a method for dividing the power dP of a ray in the forward hemisphere between the forward and transverse discrete beams. Suitable weighting factors are $\cos^2\phi$ for the forward direction and $\sin^2\phi$ for the combined transverse directions since this division clearly conserves energy and is exact for $\phi = 0, 45^\circ$, and 90° . For a ray in the backwards ($\phi \geq 90^\circ$) hemisphere the backwards and transverse weighting factors are also $\cos^2\phi$ and $\sin^2\phi$. In the case of isotropic scattering the weighting scheme results in equal intensities in the six discrete beams as it should.

Using this weighting method we find from (6) that the cross-

sections for forward, transverse, and backward scattering due to refraction are

$$\sigma_{\text{refr, f}} = \pi a^2 \int_{\phi \leq \pi/2} T_o^2 T \cos^2 \phi \sin 2\theta \, d\theta \quad (8)$$

$$\sigma_{\text{refr, t}} = \pi a^2 \int_{\theta=0}^{\pi/2} T_o^2 T \sin^2 \phi \sin 2\theta \, d\theta \quad (9)$$

$$\sigma_{\text{refr, b}} = \pi a^2 \int_{\phi \geq \pi/2} T_o^2 T \cos^2 \phi \sin 2\theta \, d\theta \quad (10)$$

The crosssections for externally reflected radiation are given similarly by

$$\sigma_{\text{refl, f}} = \pi a^2 \int_{\pi/4}^{\pi/2} R_o \cos^2 2\theta \sin 2\theta \, d\theta \quad (11)$$

$$\sigma_{\text{refl, t}} = \pi a^2 \int_0^{\pi/2} R_o \sin^2 2\theta \sin 2\theta \, d\theta \quad (12)$$

$$\sigma_{\text{refl, b}} = \pi a^2 \int_0^{\pi/4} R_o \cos^2 2\theta \sin 2\theta \, d\theta \quad (13)$$

where R_o , the surface reflectance, like T_o , includes the effects of both Fresnel reflection and absorption by asperities and edges,

and will be calculated later. These two quantities, as well as T, vary with wavelength since they depend on m which is a function of the wavelength. The weighting factors $\cos^2\phi$ and $\sin^2\phi$ in this case become $\cos^2 2\theta$ and $\sin^2 2\theta$ since $\phi = \pi - 2\theta$ for reflection.

A third contribution to the crosssections arises from radiation that undergoes one or more internal reflections in the particle. For a given incident ray the fraction of the power that emerges from the particle after undergoing internal reflections is the sum of the series $R_o T_o^2 T^2 + R_o^2 T_o^2 T^3 + \dots$ which has the value $R_o T_o^2 / (1 - R_o T)$. We assume for simplicity that this power is distributed isotropically so that the weighting factors are the same for all six discrete beams. The crosssections for internally reflected radiation are therefore

$$\sigma_{\text{int refl, f}} = \frac{1}{6} \pi a^2 \int_0^{\pi/2} \frac{R_o T_o^2 T^2}{1 - R_o T} \sin 2\theta \, d\theta \quad (14)$$

$$\sigma_{\text{int refl, t}} = \frac{4}{6} \pi a^2 \int_0^{\pi/2} \frac{R_o T_o^2 T^2}{1 - R_o T} \sin 2\theta \, d\theta \quad (15)$$

$$\sigma_{\text{int refl, b}} = \frac{1}{6} \pi a^2 \int_0^{\pi/2} \frac{R_o T_o^2 T^2}{1 - R_o T} \sin 2\theta \, d\theta \quad (16)$$

To calculate the absorption crosssection we subtract the sum of the factors $T_o^2 T$, R_o , and $R_o T_o^2 T^2 / (1 - R_o T)$ from unity and integrate over θ . Therefore

$$\sigma_{\text{abs}} = \pi a^2 \int_0^{\pi/2} \left(1 - R_o - \frac{T_o^2 T}{1 - R_o T} \right) \sin 2\theta \, d\theta \quad (17)$$

Some comments are in order on how our procedure compares with current methods. The scheme of applying weighting factors directly to the phase function in order to derive a set of discrete beams differs from the usual method in which the phase function is first expanded in a series of Legendre polynomials which is then truncated, generally after two terms. It is to be noted that the reconstituted phase function obtained from the truncated series may differ very considerably from the original phase function, especially in the case of strong forward scattering. For complete forward scattering the truncated phase function is proportional to $1 + 3 \cos \phi$ which actually yields a negative intensity in the backward direc-

tion ($\phi = \pi$). We believe that our procedure is preferable since it cannot give rise to any such non-physical intensity distribution.

The Mie theory, which is widely applied at present for phase function calculations of both spherical and non-spherical particles, cannot be made to include the important effects of edges and other asperities. The reason is that R_0 and T_0 , which contain these effects, occur in different ways in the expressions for the cross-sections for reflection, refraction, internal reflection, and absorption.

In our treatment we omit the effect of diffraction around a particle on the ground that for powders, on which most of our experimental work has been done, the particles are so close together and so randomly distributed that the conditions for Fraunhofer diffraction do not exist. In the case of clouds, however, a correction for diffraction should be made.

2. Effect of Edges and Asperities

The surface reflection and transmission factors R_0 and T_0 .

depend on the effect of edges and asperities on the particles as well as on the ordinary Fresnel reflectance. In Part II¹² we show experimentally that very fine particles adhering to the larger particles in a coarse powder cause a substantial reduction in reflection level. Another experiment shows that a similar loss occurs in the reflectance of a single crystal when the surface of the crystal is abraded. Moreover, the spectral reflectance of the abraded crystal has a broad feature that closely resembles a feature in the spectrum of a coarse powder of the same material. Microscopic examination of this powder shows that the particles have, in addition to the expected sharp edges, facets that are covered with fine steps, ridges, and other asperities. Another powder composed of spherical particles which have surface asperities but no edges has a smaller reflection loss than a powder of jagged particles. These facts prove that any perturbation of the smooth surface of a particle, whether in the form of an edge, an asperity, or an adhering smaller particle, causes enhanced absorption.

We have considered several possible mechanisms for the enhanced absorption. One idea is that asperities and edges act like a transition layer with a graded refractive index that provides a partial impedance match between vacuum and the interior of the particles. Another concept is that in spectral regions of near-total external reflection, which occurs when $n < 1$, the inhomogeneous wave traveling tangentially just under the surface of the particle is scattered by asperities and edges into the interior of the particles, where it is absorbed. A third possibility is that the asperities and edges act as induced dipoles and absorb some of the incident radiation directly.

In view of the fact that the phenomenon of excessive absorption is observed when the asperities or adhering particles are much smaller than the wavelength we conclude that direct dipole absorption is the most likely mechanism. The reason is that under these conditions dipole absorption is directly proportional to the volume of an asperity while scattering varies as the square of the

volume. Impedance matching requires a transition layer whose thickness is an appreciable fraction of a wavelength. In this paper, we will therefore confine our calculations to the case of direct absorption by surface dipoles. This subject has been investigated in some depth by Strachan¹³ and Berreman.¹⁴ We find, however, that a simpler theory, in which we ignore interactions of the dipoles with each other and with the underlying main particle, serves our purposes very well.

We assume that edges and asperities can be represented as ellipsoidal particles. An ellipsoidal sub-particle of electric susceptibility χ placed in an electric field E acquires a polarization P given by the well-known result of electrostatics:

$$P_x = \frac{\chi}{1+L\chi} E_x \quad (18)$$

$$P_y = \frac{\chi}{1+M\chi} E_y \quad (19)$$

$$P_z = \frac{\chi}{1+N\chi} E_z \quad (20)$$

where the components P_x , P_y , P_z and E_x , E_y , E_z of P and E are re-

ferred to coordinates aligned with the principal axes of the ellipsoid. The depolarization factors L, M, N are determined by the axial ratios of the ellipsoid and satisfy the relation

$$L + M + N = 4\pi \quad (21)$$

In the case of a sphere $L = M = N = 4\pi/3$.

From (18)-(20) one finds for the components of the induced dipole moment μ :

$$\mu_x = \frac{V(K-1)E_x}{4\pi + L(K-1)} \quad (22)$$

$$\mu_y = \frac{V(K-1)E_y}{4\pi + M(K-1)} \quad (23)$$

$$\mu_z = \frac{V(K-1)E_z}{4\pi + N(K-1)} \quad (24)$$

where V is the volume of the particle and K is its dielectric constant.

We assume that the particle is small enough so that (22)-(24) remain approximately valid for an alternating field. Then the power absorbed by the particle is

$$\begin{aligned}
W &= \frac{1}{2} \operatorname{Re} \left(\mathbf{E} \cdot \frac{d\boldsymbol{\mu}}{dt} \right) \\
&= -\frac{\omega}{2} \operatorname{Im} (E_x \mu_x + E_y \mu_y + E_z \mu_z) \\
&= -\frac{\omega}{2} \operatorname{Im} \left\{ \frac{(K-1)|E_x|^2}{4\pi+L(K-1)} + \frac{(K-1)|E_y|^2}{4\pi+M(K-1)} + \frac{(K-1)|E_z|^2}{4\pi+N(K-1)} \right\} \quad (25)
\end{aligned}$$

We next assume that the particle can have all orientations relative to the field E. Then the average power absorbed is

$$\bar{W} = -\frac{\omega V |E|^2}{6} \operatorname{Im}(m^2-1) \left\{ \frac{1}{4\pi+L(m^2-1)} + \frac{1}{4\pi+M(m^2-1)} + \frac{1}{4\pi+N(m^2-1)} \right\} \quad (26)$$

where ω is the angular frequency of the radiation and we have replaced K by the square of the complex refractive index, m.

It is to be noted that for a given shape of ellipsoid, specified by particular values of L, M, and N, \bar{W} becomes very large at wavelengths for which m is near the poles of the expression in curly brackets in (26), i.e., when $n \ll 1$ and k is near $[(4\pi/L)-1]^{1/2}$, $[(4\pi/M)-1]^{1/2}$, or $[(4\pi/N)-1]^{1/2}$. Such values of m actually occur in strong reststrahlen spectral regions in the infrared. Therefore reflection spectra calculated for edges or asperities repre-

sented by a particular choice of L , M , and N show very sharp decreases in reflectance at wavelengths for which m is close to these special values of m , while experimental spectra show much broader absorption maxima in the same regions of the spectrum. It is clear therefore that in the theory one must allow for the observed variability in the shapes of edges and asperities by averaging over a range of values L , M , and N . We have tried averaging Eq. (26) over various ranges of L , M , and N subject, of course, to condition (21). The result is that the widths of the absorption maxima now become similar to those observed experimentally, provided that the range of integration over L , M , and N is wide enough. Apart from this requirement one still has considerable freedom in the choice in the range of ellipsoid shapes included in the average, as determined by L , M , and N values covered. Different ellipsoid shape distributions produce subtle differences in the shape of the theoretical reflectance spectrum.

In the case of edges we choose ellipsoids that are very elongated in the direction of the edge, corresponding to $N = 0$, and with a complete range of cross-sectional shapes. Under these conditions, $M = 4\pi - L$ and L varies from 0 to 4π . On averaging (26) over all values of L with equal weighting we obtain for the power absorbed by an edge:

$$\bar{W}_e = - \frac{\omega V_e |E|^2}{24\pi} \text{Im} (m^2 - 1 + 4 \log m) \quad (27)$$

In the case of surface asperities where the shapes may vary from ridges to mounds we select ellipsoids of revolution ($L = M$) and let N vary from 0 (needle shape) to $4\pi/3$ (sphere). The average value of \bar{W} for asperities is then

$$\bar{W}_a = - \frac{\omega V_a |E|^2}{24\pi} \text{Im} \left(12 \log \frac{m^2+1}{2} - 9 \log \frac{m^2+2}{3} \right) \quad (28)$$

We must next consider the value to be assigned to the volume V of asperities or edges. In the case of asperities we could take V_a to be twice the actual volume of a ridge or bump on the sur-

face in order to allow approximately for the image field. In practice, however, since the details of the size distribution of the asperities are difficult to obtain we regard V as an adjustable constant.

In the case of edges V can be expressed as

$$V_e = \lambda_e \delta b \quad (29)$$

where λ_e , δ , and b are respectively the edge length, the effective depth of penetration of the radiation, and the effective width of the region perturbed by the presence of the edge.

We assume that both δ and b are of the order of $\lambda/4\pi$, where λ is the free-space wavelength of the radiation, except that in spectral regions of high absorption δ should become considerably smaller than b owing to reduced penetration of the radiation. This behaviour in absorbing regions is obtained if we equate δ to the $1/e$ penetration depth of the Poynting vector associated with the refracted wave, which, for an average angle of incidence of 45° , is

given by

$$\delta = \frac{\lambda}{4\pi \operatorname{Re}(0.5 - m^2)^{1/2}} \quad (30)$$

However, since this expression diverges when the medium is transparent and $n > 1/\sqrt{2}$, we replace (30) by the modified expression

$$\delta = \frac{\lambda}{4\pi} \left(1 - e^{-\frac{1}{\operatorname{Re}(0.5 - m^2)^{1/2}}} \right) \quad (31)$$

which reduces to (30) in highly absorbing regions and is approximately $\lambda/4\pi$ in transparent regions.

We now calculate the surface absorptance of a particle. The power density I_0 of the incident parallel beam is given by

$$I_0 = \frac{c |E|^2}{8\pi} = \frac{\omega \lambda |E|^2}{16\pi^2} \quad (32)$$

The average surface absorptance caused by edges and asperities for locations on the equivalent spherical particle corresponding to the angle of incidence θ is given by the expression

$$A_0 = 1 - e^{-X} \quad (33)$$

where

$$X = \frac{N_e \bar{W}_e + N_a \bar{W}_a}{I_0 \cos \theta} \quad (34)$$

Here N_e and N_a are the average number of edges and asperities per unit area of the spherical surface. The obliquity factor $1/\cos\theta$ allows for the increase in the number of edges and asperities per unit crosssection of the beam as θ increases. The exponential form of (33) ensures that A_o always lies between 0 and 1 and has the correct value when \bar{W}_e and \bar{W}_a are small.

The surface reflectance and transmittance are related to A_o by the equations

$$R_o = \left(\frac{R_s + R_p}{2} \right) (1 - A_o) \quad (35)$$

$$T_o = \left(1 - \frac{R_s + R_p}{2} \right) (1 - A_o) \quad (36)$$

where R_s and R_p are the Fresnel reflectances at the angle of incidence θ for the two states of polarization of the incident radiation. It will be noticed that (35) and (36) reduce to the usual expressions in the absence of edges or asperities. The equations also satisfy energy conservation since

$$R_o + T_o + A_o = 1 \quad (37)$$

From (27), (28), (29) and (32), Equation (34) becomes

$$X = \frac{2\pi}{3\lambda \cos\theta} \left\{ -N_e \ell_e \delta b \operatorname{Im}(m^2 - 1 + 4 \log m) \right. \\ \left. - N_a V_a \operatorname{Im}\left(12 \log \frac{m^2 + 1}{2} - 9 \log \frac{m^2 + 2}{3}\right) \right\} \quad (38)$$

In this expression we consider $N_a V_a$, the total volume of asperities per unit area of particle surface, as an adjustable parameter.

Likewise b , the effective width of an edge, is adjustable to some extent from a nominal value of $\lambda/4\pi$.

The product $N_e \ell_e$, which is the total length of all the edges of a particle divided by the surface area of the equivalent sphere, depends on the geometry of the particle. Values of this quantity, expressed in terms of the diameter d of the equivalent sphere, are shown in Table I for the five regular polyhedra. It is to be noted that for any given particle shape the edge absorption depends inversely on d . As one would expect, the proportionality constant increases, albeit erratically, as the number of edges increases. Since the dodecahedron and icosahedron have unlikely geometries

we take N_{ele} to be the mean for the first three polyhedra:

$$N_{ele} = \frac{3.4}{d} \quad (39)$$

Table I. Total Edge Length per Unit Area of Equivalent Sphere

Particle Shape	Number of Edges	N_{ele}
Tetrahedron	6	3.14/d
Cube	12	3.08/d
Octahedron	12	3.96/d
Dodecahedron	30	3.90/d
Icosahedron	30	5.93/d

FINE PARTICLE THEORY

Closely spaced fine particles of size much smaller than a wavelength interact in two ways with a beam of radiation. First, many particles scatter coherently in the beam direction and thereby give rise to an average Lorentz-Lorenz (L.L.) index of refraction

tion for the particulate medium, which causes Fresnel reflection at the surface of the medium. Second, since the particles are not regularly spaced as in a crystal each particle also scatters the radiation incoherently in all directions (Rayleigh Scattering) with an intensity that depends on the contrast between the intrinsic refractive index of a given particle and the L.L. index of the averaged medium in which it is embedded. The fine particle theory includes both of these effects.

1. The Lorentz-Lorenz Index of Refraction

Our model for a fine-particle powder consists of an assembly of ellipsoids of the same volume as the actual particles and covering a wide range of shapes. The reasons for this choice of model are that ellipsoids are mathematically tractable by the L.L. method and that an average over a range of ellipsoid shapes avoids "resonance" difficulties of the kind already mentioned in connection with edge and asperity dipoles. Such difficulties previously led us^{15, 16, 17} to calculate excessively high Fresnel reflectances

when the L.L. method was used for spherical shapes alone.

By analogy with the original L.L. method we consider a particular ellipsoid to be enclosed in a coaxial ellipsoidal cavity of the same shape carved out of a continuous medium that has the average L.L. index \bar{m} . If \bar{E} is the average electric field in the medium then the field E_c in the cavity has the components

$$E_x^c = \bar{E}_x (1 + L\bar{\chi}) \quad (40)$$

$$E_y^c = \bar{E}_y (1 + M\bar{\chi}) \quad (41)$$

$$E_z^c = \bar{E}_z (1 + N\bar{\chi}) \quad (42)$$

where $\bar{\chi}$ is the electric susceptibility of the medium and L, M, and N are the depolarization factors of the cavity and of the ellipsoid. The axes (x, y, z) are aligned with the axes of the ellipsoid.

The field E^c in the cavity produces a polarization P in the particle with components

$$P_x = \frac{\chi E_x^c}{1 + L\chi} = \frac{\chi(1 + L\bar{\chi})\bar{E}_x}{1 + L\chi} \quad (43)$$

$$P_y = \frac{\chi E_y^c}{1 + M\chi} = \frac{\chi(1 + M\bar{\chi})\bar{E}_y}{1 + M\chi} \quad (44)$$

$$P_z = \frac{\chi E_z^c}{1 + N\chi} = \frac{\chi(1 + N\bar{\chi})\bar{E}_z}{1 + N\chi} \quad (45)$$

We now assume that the ellipsoid has random orientation with respect to the field \bar{E} in the medium and calculate the average polarization of the ellipsoid in the direction of \bar{E} . The result is

$$\bar{P} = \frac{\langle P \cdot \bar{E} \rangle}{|\bar{E}|} = \frac{\chi |\bar{E}|}{3} \left(\frac{1 + L\bar{\chi}}{1 + L\chi} + \frac{1 + M\bar{\chi}}{1 + M\chi} + \frac{1 + N\bar{\chi}}{1 + N\chi} \right) \quad (46)$$

The average electric moment induced in the ellipsoid is

therefore

$$\bar{\mu} = \frac{\chi |\bar{E}| V}{3} \left(\frac{1 + L\bar{\chi}}{1 + L\chi} + \frac{1 + M\bar{\chi}}{1 + M\chi} + \frac{1 + N\bar{\chi}}{1 + N\chi} \right) \quad (47)$$

where V is the volume of the ellipsoid.

Next we average $\bar{\mu}$ with respect to ellipsoid shape. For this purpose we assume that all values of L , M , and N from 0 to 4π are equally probable subject to the constraint that $L + M + N = 4\pi$.

The result of the averaging is

$$\bar{\bar{\mu}} = |\bar{E}| V \left\{ \bar{\chi} + \frac{1}{2\pi} \left(1 - \frac{\bar{\chi}}{\chi} \right) \left[\left(1 + \frac{1}{4\pi\chi} \right) \log \left(1 + 4\pi\chi \right) - 1 \right] \right\} \quad (48)$$

We can also write the relation

$$\bar{\chi} = \frac{f\bar{\mu}}{V|\bar{E}|} \quad (49)$$

since the whole medium can be considered to be composed of ellipsoids of average dipole moment $\bar{\mu}$ packed at the volume fraction f .

On substituting for $\bar{\mu}$ from (48) into (49) and solving for $\bar{\chi}$

we obtain the result

$$2\pi\bar{\chi} = \frac{f \left[\left(1 + \frac{1}{4\pi\chi}\right) \log (1 + 4\pi\chi) - 1 \right]}{1 - f + \frac{f}{2\pi\chi} \left[\left(1 + \frac{1}{4\pi\chi}\right) \log (1 + 4\pi\chi) - 1 \right]} \quad (50)$$

Since $4\pi\bar{\chi} = \bar{m}^2 - 1$ and $4\pi\chi = m^2 - 1$ we obtain the following relation between the L.L. index \bar{m} and the particle index m :

$$\bar{m}^2 = 1 + \frac{2f \left[\left(1 + \frac{1}{m^2-1}\right) \log (m^2) - 1 \right]}{1 - f + \frac{2f}{m^2-1} \left[\left(1 + \frac{1}{m^2-1}\right) \log (m^2) - 1 \right]} \quad (51)$$

It is worth noting that $\bar{m} \rightarrow m$, as it should, when $f \rightarrow 1$.

This formula can readily be generalized for the case of a mixture of fine particles of different materials with volume fractions f_j and indices m_j . The result is

$$\bar{m}^2 = 1 + \frac{2 \sum_j f_j \left[\left(1 + \frac{1}{m_j^2-1}\right) \log (m_j^2) - 1 \right]}{1 - \sum_j f_j + \sum_j \frac{2f_j}{m_j^2-1} \left[\left(1 + \frac{1}{m_j^2-1}\right) \log (m_j^2) - 1 \right]} \quad (52)$$

2. Rayleigh Scattering Crosssection

The Raleigh scattering by an ellipsoidal particle of index m immersed in a medium of L.L. index \bar{m} depends on the polarizibility of the ellipsoid, which is related to the relative index m/\bar{m} by the formula

$$\alpha_x = \frac{V\left(\frac{m^2}{\bar{m}^2} - 1\right)}{4\pi + L\left(\frac{m^2}{\bar{m}^2} - 1\right)} \quad (53)$$

with similar expressions for the polarizibilities α_y and α_z along the two other principal axes of the ellipsoid.

The total scattering crosssection for an incident parallel beam of radiation in the medium is given by¹⁸

$$\sigma = \frac{8\pi}{3} \kappa^4 (\ell^2 |\alpha_x|^2 + m^2 |\alpha_y|^2 + n^2 |\alpha_z|^2) \quad (54)$$

where ℓ , m , n are the direction cosines of the electric field in the medium and

$$\kappa = \frac{2\pi}{\lambda_{\text{medium}}} = \frac{2\pi \bar{n}}{\lambda} \quad (55)$$

Here \bar{n} is the real part of the L.L. index \bar{m} .

On averaging over all orientations of the ellipsoid relative

to the electric field direction we find, from (53) and (54):

$$\sigma = \frac{8\pi}{9} \kappa^4 V^2 \left| \frac{m^2}{\bar{m}^2} - 1 \right|^2 \left\{ \left| \frac{1}{4\pi + L \left(\frac{m^2}{\bar{m}^2} - 1 \right)} \right|^2 + \left| \frac{1}{4\pi + M \left(\frac{m^2}{\bar{m}^2} - 1 \right)} \right|^2 + \left| \frac{1}{4\pi + N \left(\frac{m^2}{\bar{m}^2} - 1 \right)} \right|^2 \right\} \quad (56)$$

We next average over all ellipsoid shapes as in the derivation of the L.L. index. Under these conditions the three terms containing L, M, and N are equal and (56) reduces to

$$\sigma = \frac{8\pi}{3} \kappa^4 V^2 \left| \frac{m^2}{\bar{m}^2} - 1 \right|^2 \left\langle \left| \frac{1}{4\pi + L \left(\frac{m^2}{\bar{m}^2} - 1 \right)} \right|^2 \right\rangle \quad (57)$$

where $\langle \rangle$ indicates averaging over L.

On carrying out the averaging for $0 \leq L \leq 4\pi$, and allowing for the fact that m and \bar{m} are complex numbers, we obtain

$$\sigma = \frac{\kappa^4 V^2 \left| \frac{m^2}{\bar{m}^2} - 1 \right|^2}{3\pi} \left\{ \frac{\text{Im}(\log \frac{m^2}{\bar{m}^2}) + \text{Im} \left(\frac{\log \frac{m^2}{\bar{m}^2}}{\frac{m^2}{\bar{m}^2} - 1} \right)}{\text{Im} \left(\frac{m^2}{\bar{m}^2} \right)} \right\} \quad (58)$$

where Im signifies "imaginary part".

One can generalize this expression to include the case of a mixture of mineral powders (or of particles of the same material but of different sizes) by multiplying (58) by the particle volume fraction f_j of material j and summing over j .

RADIATIVE TRANSFER THEORY

The radiative transfer theory presented in this paper is restricted to the calculation of the reflectance of a parallel beam incident normally on the surface of a semi-infinite powder or cloud. This case is equivalent, by Kirchhoff's law, to the calculation of the normal emittance of an isothermal powder or cloud, which is approximately the measured quantity discussed in Part II.

(1) Six-beam Model

Multiple scattering of the incident parallel beam produces diffuse radiation in the particulate medium. We represent the combination of diffuse and collimated radiation at any point in the medium by means of six mutually orthogonal beams, one of which

is in the direction of the collimated incident beam which we take to be the x-direction.

Let I and J be the flux densities of the beams in the x and -x directions and I_1, I_2, I_3, I_4 the fluxes in the transverse directions y, -y, z, -z. Then the radiative transfer equations are

$$\frac{dI}{dx} = -(K + S_b + S_t)I + S_b J + \frac{1}{4} S_t (I_1 + I_2 + I_3 + I_4) \quad (59)$$

$$-\frac{dJ}{dx} = -(K + S_b + S_t)J + S_b I + \frac{1}{4} S_t (I_1 + I_2 + I_3 + I_4) \quad (60)$$

$$\frac{dI_1}{dy} = -(K + S_b + S_t)I_1 + S_b I_2 + \frac{1}{4} S_t (I + J + I_3 + I_4) \quad (61)$$

$$-\frac{dI_2}{dy} = -(K + S_b + S_t)I_2 + S_b I_1 + \frac{1}{4} S_t (I + J + I_3 + I_4) \quad (62)$$

$$\frac{dI_3}{dz} = -(K + S_b + S_t)I_3 + S_b I_4 + \frac{1}{4} S_t (I + J + I_1 + I_2) \quad (63)$$

$$-\frac{dI_4}{dz} = -(K + S_b + S_t)I_4 + S_b I_3 + \frac{1}{4} S_t (I + J + I_1 + I_2) \quad (64)$$

where K, S_b , S_t are the absorption, backscattering, and transverse scattering coefficients, respectively, for any of the beams.

These coefficients will later be expressed in terms of the single-

particle crosssections derived above.

Owing to the symmetry of the situation we have:

$$I_1 = I_2 = I_3 = I_4 = \frac{1}{4} P \quad (65)$$

where P is the sum of the transverse fluxes. Therefore, on

adding Equations (61)-(64) we obtain the result

$$P = \frac{S_t}{K + \frac{1}{2} S_t} (I + J) \quad (66)$$

On substituting this value of P into (59) and (60) and rearranging the resulting equations we find that they can be put in the form

$$\frac{dI}{dx} = -(K' + S')I + S'J \quad (67)$$

$$-\frac{dJ}{dx} = -(K' + S')J + S'I \quad (68)$$

where

$$S' = \frac{4KS_b + 2S_b S_t + S_t^2}{4K + 2S_t} \quad (69)$$

$$K' = \frac{4K^2 + 6KS_t}{4K + 2S_t} \quad (70)$$

Thus the radiative transfer problem here reduces essentially to a two-beam problem with equations which are identical in form

to the well-known equations of Schuster⁵ or Kubelka and Munk.⁶

The interpretation of our coefficients K' and S' is different, however, since, from (69) and (70), each involves both absorption and scattering terms which arise from the interaction of the I and J beams with the reservoir of transverse radiation. In the Schuster-Kubelka-Munk equations, on the other hand, K' and S' refer to absorption and backscattering, respectively, of two oppositely-directed diffuse beams.

We have also tried an alternative six-beam model, similar to that used by Conel,⁴ in which three of the orthogonal beams are equally inclined with respect to the +x direction and the other three beams equally inclined with respect to the -x direction. This model is very convenient if the incident radiation is diffuse, but leads to difficulties for collimated incident radiation. To solve the radiative transfer problem for this orientation of the six beams one must derive coefficients S_1 and S_2 (from the single-

particle scattering coefficients) that describe respectively the rate of scattering of the incident collimated beam into the forward and backward sets of beams. In performing this calculation we found that S_1 becomes negative if the single-particle phase function has a sufficiently large backward-to-forward asymmetry factor. For this reason we abandoned this model.

(2) Coarse-particle Case

In the case of coarse particles the boundary condition at the surface of the particulate medium ($x = 0$) is that $I(o)$ is matched to the flux I_o of the incident collimated beam, while $J(o)$ is matched to the flux J_o of the outgoing diffuse radiation. At $x = \infty$ both I and J are zero. With these boundary conditions Equations (67) and (68) give for the reflectance J_o/I_o the expression

$$R_V = 1 + \frac{K'}{S'} - \sqrt{\left(\frac{K'}{S'}\right)^2 + 2\left(\frac{K'}{S'}\right)} \quad (71)$$

It is worth noting that our six-beam transfer model does not involve a gradual conversion of the incident collimated beam

into diffuse radiation in the medium as in the model of Duntley¹⁹ which is similar to an earlier version of our theory. While this change made little difference,²⁰ our present model is much easier to apply and can readily be extended to the case of oblique incidence.

The scattering coefficients S_t and S_b in (69) and (70) are given by

$$S_t = N \sigma_t \quad (72)$$

$$S_b = N \sigma_b \quad (73)$$

Here N is the number of particles per unit volume and σ_t , σ_b are the total single-particle crosssections for transverse and backward scattering:

$$\sigma_t = \sigma_{\text{refr},t} + \sigma_{\text{refl},t} + \sigma_{\text{int.refl},t} \quad (74)$$

$$\sigma_b = \sigma_{\text{refr},b} + \sigma_{\text{refl},b} + \sigma_{\text{int.refl},b} \quad (75)$$

The absorption coefficient K is given by

$$K = N \sigma_{\text{abs}} \quad (76)$$

The particle density N is related to the particle volume-

fraction f and the diameter d of the equivalent sphere by the formula

$$N = \frac{6f}{\pi d^3} \quad (77)$$

In the case of a mixture of particles, or of a distribution of sizes, we attach a subscript j to f and d in (77) and sum (72), (73) and (76) over j to obtain average values of S_t , S_b and K .

(3) Fine-Particle Case

In the case of fine particles the boundary condition at the surface of the medium is that both the incident collimated beam I_0 and the outgoing internal beam J undergo Fresnel reflection due to the L.L. index \bar{m} of the medium. The surface reflectance R_s is given by

$$R_s = \left| \frac{\bar{m} - 1}{\bar{m} + 1} \right|^2 \quad (78)$$

With this boundary condition the reflectance R derived from Equations (67) and (68) now becomes

$$R = R_s + \frac{(1 - R_s)^2 R_v}{(1 - R_s R_v)} \quad (79)$$

where R_v is the volume reflectance given as before by (71).

As in the coarse-particle case, a comment is in order here.

Since our model involves only collimated rather than diffuse beams the phenomenon of trapped radiation¹⁹ due to total internal reflection does not arise since all beams are either perpendicular or parallel to the surface of the medium.

The coefficients S_t , S_b and K for the fine-particle case are

$$S_t = \frac{4}{6} N \sigma \quad (80)$$

$$S_b = \frac{1}{6} N \sigma \quad (81)$$

$$K = \frac{4\pi\bar{K}}{\lambda} \quad (82)$$

where σ is given by (58). For mixtures we again use average values of S_t , S_b and K obtained by attaching a subscript j to f and d in (77) and summing over j .

COMBINATION OF COARSE AND FINE PARTICLE THEORIES

For particles in the size range between a few tenths of a wavelength and one wavelength neither the coarse nor the fine par-

ticle theory would be expected to give accurate results. The laws of geometrical optics used in the ray tracing part of the coarse particle theory begin to break down for particles smaller than the wavelength and the laws of wave optics gradually take over. Again, for particles larger than a few tenths of a wavelength the concept of the average L.L. index used in the fine particle theory becomes inexact because neighboring particles are too widely spaced to produce strong coherent scattering. Moreover, the dipole approximation used in the Rayleigh scattering part of the fine-particle theory becomes a progressively poorer approximation as the particles become larger.

It would clearly be impractical to derive a comprehensive theory for particles of any size which would include the effects of coherence, edges, asperities, and particle shape. Instead, we use an empirical bridging formula that reduces to the coarse-particle theory for large particles and to the fine-particle theory for very

small particles. The main justification for the bridging formula is that for particles in the intermediate size range the coarse- and fine-particle theories give calculated reflection spectra that resemble each other.

We assume a bridging formula of the type

$$R = F R_{\text{coarse}} + (1 - F) R_{\text{fine}} \quad (83)$$

where F is a weighting factor that is a function of the size parameter $x = \pi d/\lambda$. A simple expression for F in terms of x that approaches 1 when x is large and 0 when x is small is:

$$F = \frac{Ax^p}{1 + Ax^p} \quad (84)$$

where p is a positive integer and A is a parameter that determines the suddenness of the transition from coarse-particle theory to fine-particle theory. We have found by trial that the values $p = 1$, $A = 6$ produce a merging of the two theories that gives good agreement with experiment over a wide range of particle sizes. Under these conditions (84) becomes

$$F = \frac{6x}{1 + 6x} \quad (85)$$

COMPUTATION

The theory described above has been programmed using Fortran IV for high speed computations. The input variables are the particle sizes, d_j , the complex refractive indices, $m_j(\lambda)$, the mineral volume fractions, f_j and two adjustable parameters b and NV which are involved in the effects of dipole absorption by edges and surface asperities respectively.

The complex refractive indices are provided as needed in the program by means of a dispersion formula containing a small number of Lorentz resonance terms and a high frequency dielectric constant. Each resonance term is defined by the usual three classical oscillator parameters; line frequency, width and strength. The values for the oscillator parameters are taken from published results where available.^{21,22} Otherwise we derive them ourselves using a similar least squares method of fitting reflectance spectra,

calculated from the dispersion formula, to the observed reflectance spectra of polished samples.

In the case of uniaxial birefringent minerals, we treat the indices for the $E \perp C$ and $E \parallel C$ orientations with suitable weighting factors as though they represent different species. While this is not an exact method, it has proven to be a very satisfactory approximation.

The output of the program consists of the reflectance as calculated by either the coarse-particle theory or the fine-particle theory or a combination of the two theories as described above, both in the form of printed outputs and spectral plots. The results of a number of intermediate calculations of quantities such as the various K 's and S 's are also produced and these are very useful in developing an understanding of the complicated processes involved.

The program which has not yet been fully optimized takes about 1 minute for a 300 point spectrum on an IBM 360/65 machine,

for a two component mixture and single particle size. Provision for running a particle size dispersion is provided but the running time is not simply proportional to the number of particle sizes employed as repetitive calculations are avoided. The times involved in performing these calculations can thus be seen to be quite modest.

DISCUSSION AND CONCLUSION

The theory presented above contains a number of innovations in the treatment of the interaction of radiation with a particulate medium. As mentioned earlier, we consider that the most important innovations are the inclusion of the additional absorption caused by edges and asperities on coarse particles. Without allowance for these effects one cannot achieve a good fit between theory and experiment in the reststrahlen regions of the spectrum.

Other new features are involved in the calculation of the single particle crosssections in the coarse-particle theory, including the use of the Poynting vector rather than the wave normal

in coarse particle ray tracing; the separate evaluation of the crosssections for external reflection, multiple internal reflection, and refraction, which provides considerable insight into the relative importance of the three mechanisms; and the use of the $\cos^2\phi$, $\sin^2\phi$ weighting factors to convert the single-particle phase function into the mathematically convenient form of six discrete beams.

In the fine-particle theory the recognition of the "resonances" associated with ellipsoid shape appears to be new, as well as the method of averaging over a wide range of shape, both for the Lorentz-Lorenz index and for the Rayleigh scattering. The immersed-particle model used to calculate the Rayleigh scattering of closely spaced fine particles is also novel.

The use of separate theories for coarse and fine particles along with a bridging formula for intermediate particle sizes seems to be a different approach. This method is seen to be a

necessary procedure as soon as one recognizes the essential difference between the regimes of coherent and incoherent scattering by the particles.

Finally, the six-beam radiative transfer method, in which the four transverse beams act as a radiation reservoir that interacts with the forward and backward beams, is new and well matched to the six-beam single-particle phase function.

The development of the theory has proceeded by a process of successive refinement of approximations as the results of the computations were compared with experimental data. A number of refinements remain to be made. Among the most important of these is a suitable allowance for the fact that opaque particles in densely packed powders shadow each other in a way not allowed for in any continuum model. We have attempted to solve this problem by replacing the derivatives on the left-hand sides of Equations (59)-(64) by the finite intensity differences ΔI , ΔJ , etc., that

exist across an individual particle. On solving the difference equations for a monolayer of particles of diameter d we find for the reflectance and absorptance of the monolayer:

$$\rho = \left(\frac{6f}{\pi d^3} \right)^{\frac{2}{3}} \left(\frac{4 \sigma_{abs} \sigma_b + 2 \sigma_b \sigma_t + \sigma_t^2}{4 \sigma_{abs} + 2 \sigma_t} \right) \quad (86)$$

$$\alpha = \left(\frac{6f}{\pi d^3} \right)^{\frac{2}{3}} \left(\frac{4 \sigma_{abs}^2 + 6 \sigma_{abs} \sigma_t}{4 \sigma_{abs} + 2 \sigma_t} \right) \quad (87)$$

where f is the volume fraction of the particles and σ_{abs} , σ_t , σ_b are given, as before, by (17), (74), and (75). The reflectance of a semi-infinite stack of monolayers is

$$R_v = \left(\frac{1 + \rho^2 - \tau^2}{2\rho} \right) - \left\{ \left(\frac{1 + \rho^2 - \tau^2}{2\rho} \right)^2 - 1 \right\}^{\frac{1}{2}} \quad (88)$$

where τ , the monolayer transmittance, is given by

$$\tau = 1 - \rho - \alpha \quad (89)$$

A comparison of spectral reflectances calculated by the non-continuum formula (88) and the continuum formula (71) shows that the non-continuum model gives significantly higher reflectance in

strong reststrahlen regions and is thereby in better agreement with experimental spectra. As one would expect, the two theories give identical results in regions of the spectrum where ρ and $1-\tau$ are small. Unfortunately, the non-continuum model, like that of Melamed,¹¹ cannot be easily extended to the important case of a medium composed of a mixture of particles of different materials with a wide range of sizes, such as occurs on planetary surfaces. The present continuum model, on the other hand, is able to treat such a medium, although with some loss of accuracy in the general spectral level.

An assumption in the coarse-particle theory is that the rays enter each material from the void rather than directly from one material into another. While this is a good approximation for relatively low density powders, it will break down in the case of compact agglomerates. In principle we could extend the computational techniques to handle this case. Another difficult case

would be the suspension of coarse particles of some materials in a medium composed principally of fine-particles of other materials. The present theory is capable of treating this case as well, again with some modifications of the computer programs.

We intend to apply the theory described in this paper to transmission spectra of atmospheric clouds and aerosols in the near future. In the aerosol case immersed-particle scattering reduces to ordinary Rayleigh scattering owing to the small volume fraction of the fine particles involved.

Our method of separately calculating the various scattering crosssections of a single particle allows relatively simple modifications of the crosssection formulas to be made, so that one can treat the case of particles of non-spherical gross shape, e.g., platelets or cylinders. Such particles are often encountered in ice-crystal clouds. It is to be noted that thin platelets give little refractive scattering although in any theory based on a

spherical model (e.g. Mie theory) a refractive contribution is automatically included.

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III. The Spectral Reflectance and Emittance
of Particulate Materials,

Part II Application and Results

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Abstract

Experimental data on the spectral emittance of particulate minerals, obtained using a Michelson Interferometer operating between 300 and 1400 cm^{-1} are compared with the results of a new theory of the spectral reflectance (emittance) of such materials. The comparisons show that the theory predicts the infrared spectra of minerals quite well both for single substances and mixtures, over the wide particle size range between 0.3 μ and 330 μ . The good agreement suggests that the theory can be used in the application of remote infrared spectroscopy to such problems as the compositional analysis of the surface of a planet.

INTRODUCTION

For many years it has been recognized that infrared spectroscopy provides unique information as to the composition of minerals. This results from the rich vibrational spectra that may be obtained in the infrared region. While most workers have concentrated their efforts on laboratory transmission spectra, such

potentially important remote sensing techniques as emission and reflection were in fact investigated by Coblentz¹ almost three quarters of a century ago. The advent of the space program re-awakened an interest in this topic and a number of groups proposed remote infrared spectroscopy as a very valuable tool for the evaluation of planetary surface mineralogy from orbiting or flyby spacecraft. The instrumental capability for obtaining excellent remote data has continued to improve as is demonstrated by recent results from the Mariner Mars² and Nimbus³ orbiting spacecraft.

While the reflectance spectra of polished mineral samples provide many strong features suitable for relatively simple interpretation, the spectra of roughened or particulate samples change in important ways with the degree of roughness or the particle size and packing density. For some time, it was believed by many investigators that the spectra of finely divided minerals approximate blackbody spectra and thus contain very little or no information as to composition. We showed⁴ both experimentally and theoretically that this was not true but that the problem was simply one of the signal-to-noise ratio. Others^{5,6} examined the effects of particle size on spectral shape and found that different spectral features had widely differing behaviors as a function of particle size. In some cases⁷ peaks could even turn into troughs. As the greatest difficulties inherent in the prospects for remote mineral analysis by infrared spectroscopy were clearly those of

interpretation, we began the development of a new theory of the reflectance or emittance of particulate solids several years ago⁸. The current status of the theory is detailed in the preceding paper⁹ (hereafter referred to as I).

An earlier version of the theory produced relatively good agreement with experimental data¹⁰ in the coarse-particle regime. The agreement in all particle size ranges has continued to improve through a process of successive refinement of the assumptions as discrepancies appeared in the comparison of theoretical predictions with experimental measurements.

APPARATUS

Our measurements have been made on a 7-35 μ (1400-300 cm^{-1}) Michelson interferometer spectrometer system, which uses a germanium coated CsI beamsplitter and a Barnes TGS pyroelectric detector. Figure 1 shows the Interferometer. Figure 2 shows the sample chamber in which the emittance spectra of mineral powders and comparison blackbody sources are measured. The samples are heated from below and mounted on a lazy susan turntable so that alternate measurements can be made without breaking the vacuum or inert gas atmosphere. Vacuum measurements can be made to investigate the effects of high thermal gradients on the spectra. Figure 3 shows the entire system including a digital data acquisition system based on a Hewlett-Packard 2116B minicomputer.

The powders are generally supported in a sample tray 0.6 cm deep with a total volume of 3.04 cm^3 . This tray is instrumented

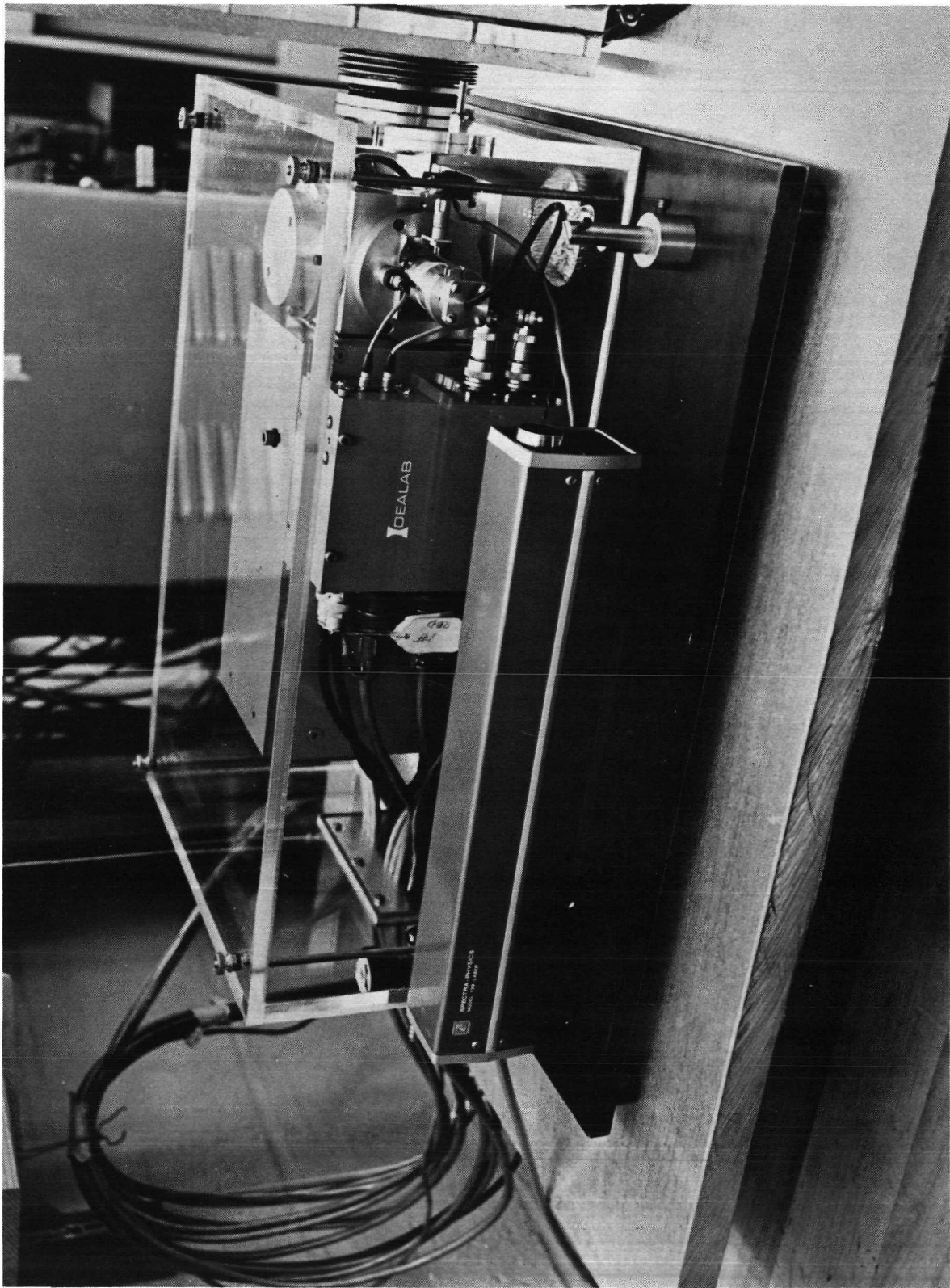


FIGURE 1 MICHELSON INTERFEROMETER

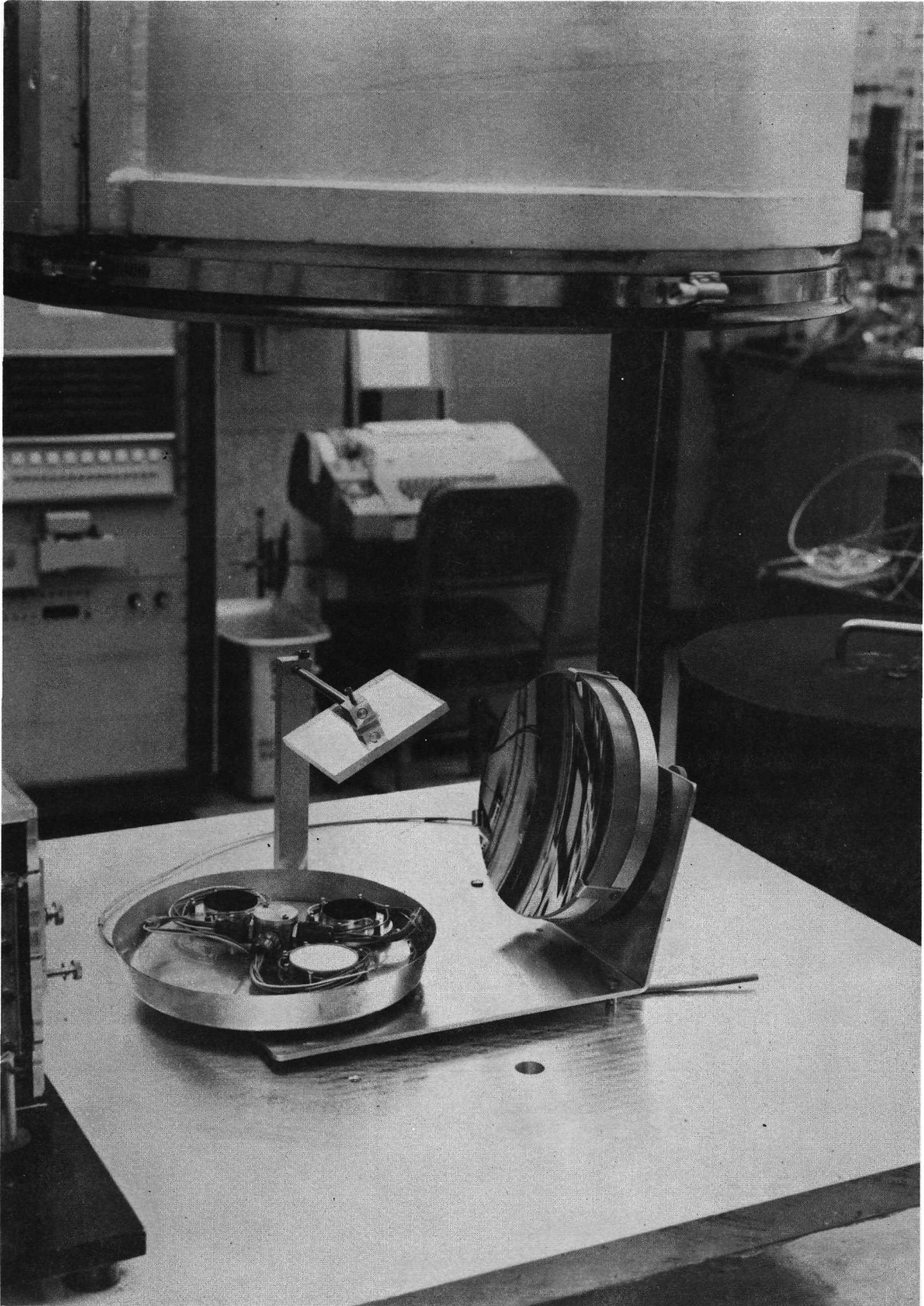


FIGURE 2 EMITTANCE OPTICS

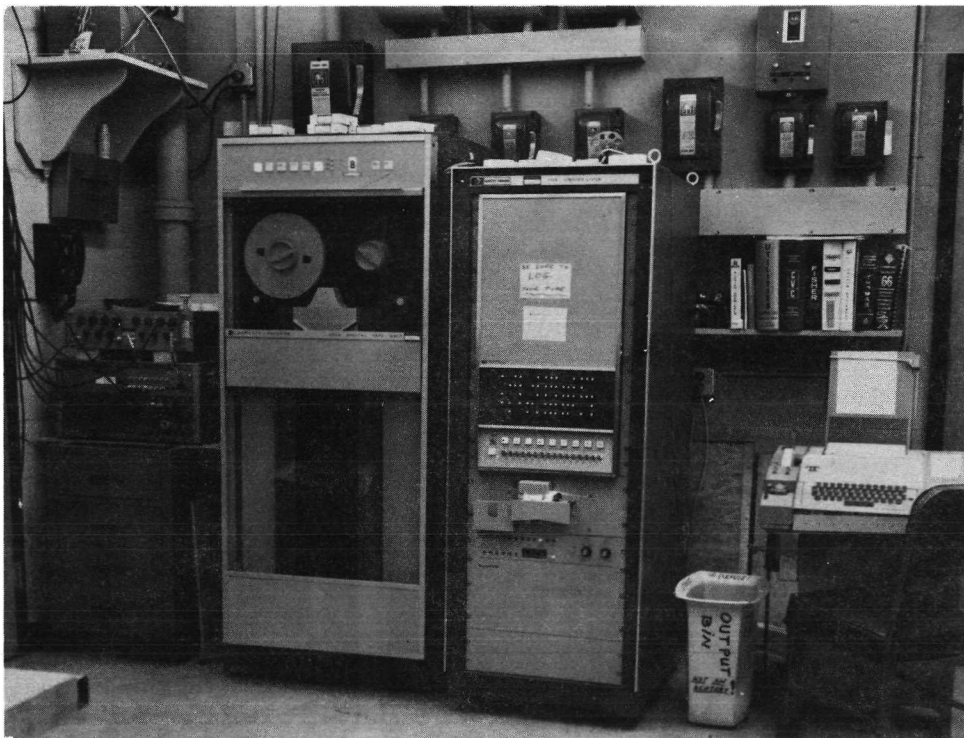
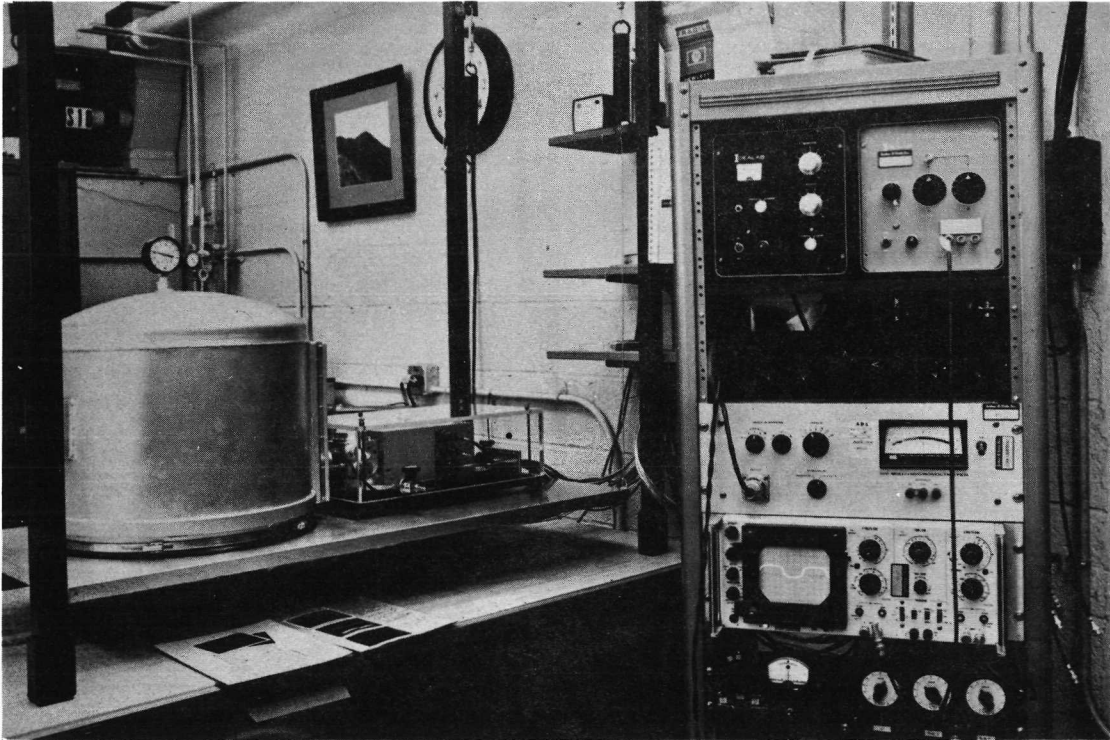


FIGURE 3 INTERFEROMETER SYSTEM

for temperature measurements by imbedding a thermistor in the base and stringing two sets of 0.05 cm differential copper-constantan thermocouples 0.2 and 0.5 cm above the base (Figure 4). The gradient between the upper two stations is extrapolated to the nominal surface in order to provide a first estimate of the "surface" temperature. Thermistors located on the interferometer detector case, on the black shield that surrounds the sample chamber, and on the experimental blackbodies provide the other temperatures required for data reduction.

The experimental blackbodies are constructed of aluminum and have concentric 30° V-grooves cut into the surface. They are coated with Parsons black paint (Eppley Laboratory) as previous results using Nextel black paint (Minnesota Mining and Manufacturing Co.) indicated the presence of spectral features near 1090 cm^{-1} and 460 cm^{-1} , owing to the presence of small glass beads in the Nextel paint. The Parsons black standard was run against a Cabot Corporation Carbolac-1 black powder sample (9m μ carbon particles) as well as a Nextel-coated standard and the data indicate a spectrally flat emittance for the Parsons standard from 1400 cm^{-1} to about 500 cm^{-1} . The emittance of the Parsons standard then appears to fall monotonically reaching apparent values of about .97 near 350 cm^{-1} .

DATA REDUCTION

The data reduction procedure involves Fourier transformation of the interferograms of the sample and the two experimental

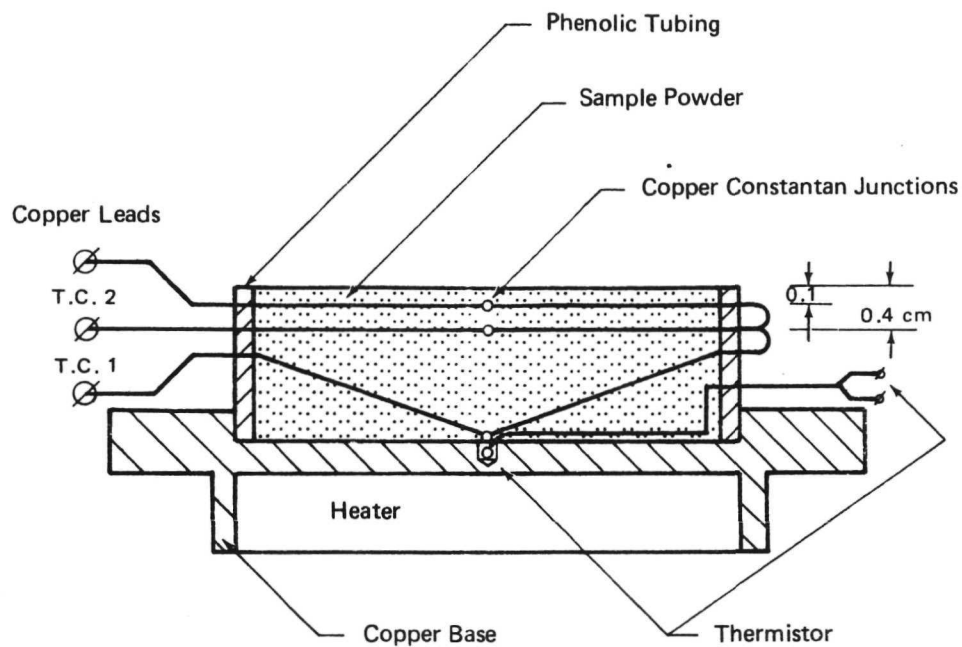


FIGURE 4 SAMPLE TRAY (not to scale)

blackbodies which are run at temperatures close to that of the sample and differ by about 5°C. High signal-to-noise data are obtained with a sample surface temperature about 15°C above the near-ambient interferometer and shield temperatures in about 20 minutes with a spectral resolution of about 3 cm⁻¹.

The spectral emittance $\epsilon_s(\nu)$ of the sample is the spectral radiance $N_s(\nu)$ of the sample divided by the spectral radiance $B_s(\nu)$ of a blackbody at the same temperature as the sample. We calculate $B_s(\nu)$ from the Planck function using the inferred surface temperature of the sample. We determine $N_s(\nu)$ from the measured output spectrum $E_s(\nu)$ of the sample by means of the instrument transfer function and a correction for radiation from the shield that is reflected from the sample.

The transfer function relates radiance $N(\nu)$ to spectral output voltage $E(\nu)$ obtained by Fourier transformation of the interferogram. We assume that the formula is linear and of the form

$$N = N_o + \frac{E}{R} \quad (1)$$

where N_o , the input radiance that produces zero output signal, is equal to the radiance of the instrument in the direction towards the sample. To determine N_o and R we replace the sample successively by two blackbodies at temperatures T_1 and T_2 . Then, from (1)

$$\epsilon_1 B_1 = N_o + \frac{E_1}{R} \quad (2)$$

$$\epsilon_2 B_2 = N_o + \frac{E_2}{R} \quad (3)$$

where ϵ_1, ϵ_2 are the emittances of the "blackbodies," which would allow for small departures from perfect blackbody behavior.

From (2) and (3) we obtain for $R(\nu)$ and $N_o(\nu)$:

$$R(\nu) = \frac{E_1 - E_2}{\epsilon_1 B_1 - \epsilon_2 B_2} \quad (4)$$

and

$$N_o(\nu) = \frac{1}{2} [\epsilon_1 B_1 + \epsilon_2 B_2 - \frac{1}{R} (E_1 + E_2)] \quad (5)$$

With the sample in place we determine the net radiance N_s of the sample from the measured spectral output E_s by means of Eq (1)

$$N_s = N_o + \frac{E_s}{R} \quad (6)$$

and where R and N_o are known from (4) and (5).

On combining (4) - (6) we can express N_s directly in terms of the measured quantities E_1, E_2 and E_s :

$$N_s = \frac{1}{2} (\epsilon_1 B_1 + \epsilon_2 B_2) + \frac{1}{2} (\epsilon_1 B_1 - \epsilon_2 B_2) \left(\frac{2E_s - E_1 - E_2}{E_1 - E_2} \right) \quad (7)$$

Finally, the emittance ϵ_s of the sample is computed from the formula

$$\epsilon_s = \frac{N_s - B_c \left(1 - \frac{\Omega}{\pi}\right) - B_i \frac{\Omega}{\pi}}{B_s - B_c \left(1 - \frac{\Omega}{\pi}\right) - B_i \frac{\Omega}{\pi}} \quad (8)$$

which is based on simple radiative transfer considerations. Here B_c is the radiance of the cavity or shield (assumed to be a blackbody) at temperature T_c surrounding the sample, B_i is the

radiance of the interferometer aperture (also assumed to be a blackbody at temperature T_1), B_s is the Planck function for the temperature T_s of the sample, and Ω is the solid angle subtended at the sample by the interferometer aperture. The cavity walls are coated with 3M Nextel black paint which, together with multiple reflections, should make it an almost perfect blackbody. In the data reduction up to the present we have assumed that the blackbodies are indeed black, so that $\epsilon_1 = \epsilon_2 = 1$.

As the surface of the powder cannot be microscopically smooth (we draw a spatula edge across the powder surface to prevent gross roughness) and as the depth of origin of the radiation is a function of the optical constants at the various frequencies, we cannot expect the definition of the surface to be very exact. Further, the position of the stretched thermocouple wires has some error and the possibility of a poor equilibrium between the radiation field and the shiny thermocouples is quite likely.

For all of these reasons we chose to estimate our "surface" temperature by a radiative expedient based on what is generally known as the principal Christiansen frequency^{11,12}. When the refractive index, n of a substance approaches unity and the absorption index, k is quite small ($\leq 10^{-2}$) the amount of surface reflection is negligible and the refractive scattering is minimal. Thus the emittance is maximum and close to unity. This combination of the optical constants occurs at slightly higher frequencies than the first principal reststrahlen band.

It cannot in principle occur for a mixture of materials as the location of the Christiansen frequency will be different for each species. Even two orientations of birefringent minerals have slightly different Christiansen frequencies. We take the value for the average of the two orientations of corundum to be 1020 cm^{-1} and for quartz to be 1360 cm^{-1} from data of Barker¹³ and Spitzer and Kleinman¹⁴. We assume that the only significant temperature error in our data reduction scheme is that of the "surface temperature". We then use the extrapolated "surface temperature" as a first estimate of the true surface temperature and calculate an emittance spectrum. The computer program is constructed so that this temperature may be varied and the procedure is to make the required changes in this temperature so as to provide a unit emittance at the principal Christiansen frequency.

We have found that the extrapolated surface temperature and the temperature derived by the Christiansen frequency technique differ by a few tenths of a degree in cases of relatively low temperature gradients (ca. $5^\circ/\text{cm}$ to $10^\circ/\text{cm}$ as measured by our upper differential thermocouple) which occur in atmospheric pressure runs. Under conditions of high temperature gradients (vacuum) the procedure is quite inaccurate and we plan to discuss this in a future communication. All of the data are plotted as reflectance ($R = 1 - \epsilon$) for convenience in comparison with the theory. The validity of this procedure which is simply an application of Kirchhoff's Law for "infinitely" thick media has been previously shown¹¹.

COMPARISON OF THEORY AND EXPERIMENT

As discussed in (I), our theory consists of two subtheories. The first, which is based on geometrical optics supplemented by certain wave optics considerations, is applicable in the coarse-particle region which extends down to particle sizes roughly comparable with the free space radiation wavelength. The second, a fine-particle theory, is based on the Lorentz-Lorenz theory of dielectrics, and Rayleigh scattering by particles immersed in the medium whose optical properties are set by the Lorentz-Lorenz field. In order to bridge the two theories for those particle sizes falling in the ill-defined region between the regions of clear applicability of either theory, we have at present resorted to a simple empirical bridging formula that gradually merges the results of the two subtheories.

Figure 5 shows a comparison of the experimental and theoretical results for corundum powders. The 120μ sample was obtained from the Norton Company, the 0.3μ sample from Adolf Meller Co., and the other samples are LWA powders obtained from the Microabrasives Corporation of Westfield, Massachusetts. The particle sizes shown on the figure for the LWA powders are Microabrasives designations except for the 3.5μ sample. This is Microabrasives LWA 3 but our particle size counts indicated a value of about 3.8μ . The shape of the spectrum between 630 and 900 cm^{-1} and the particle size effect in regions of the spectrum where the particles are opaque was correctly predicted

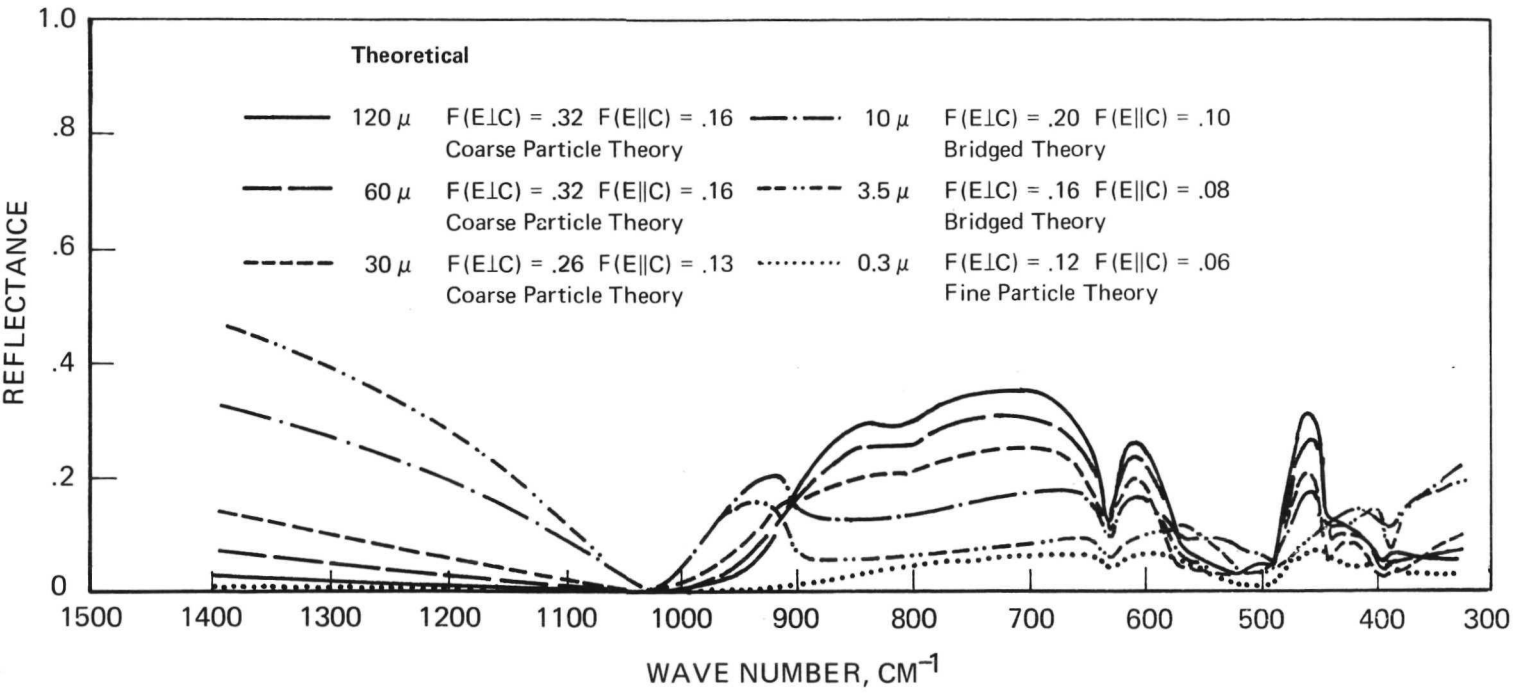
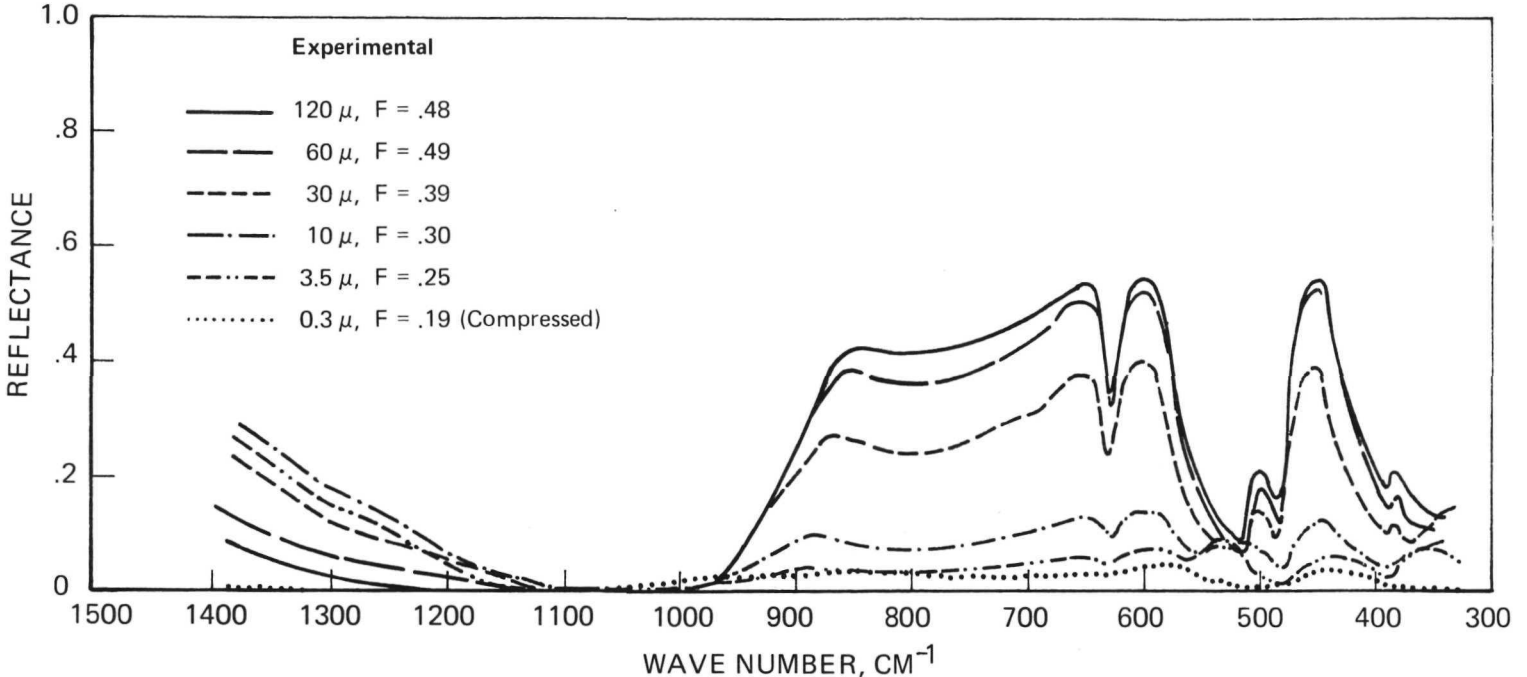
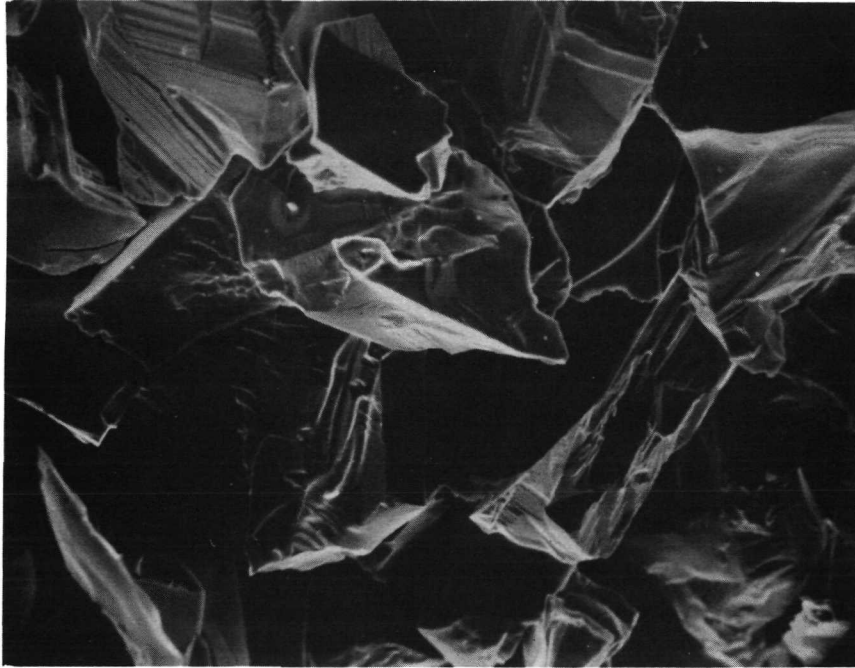


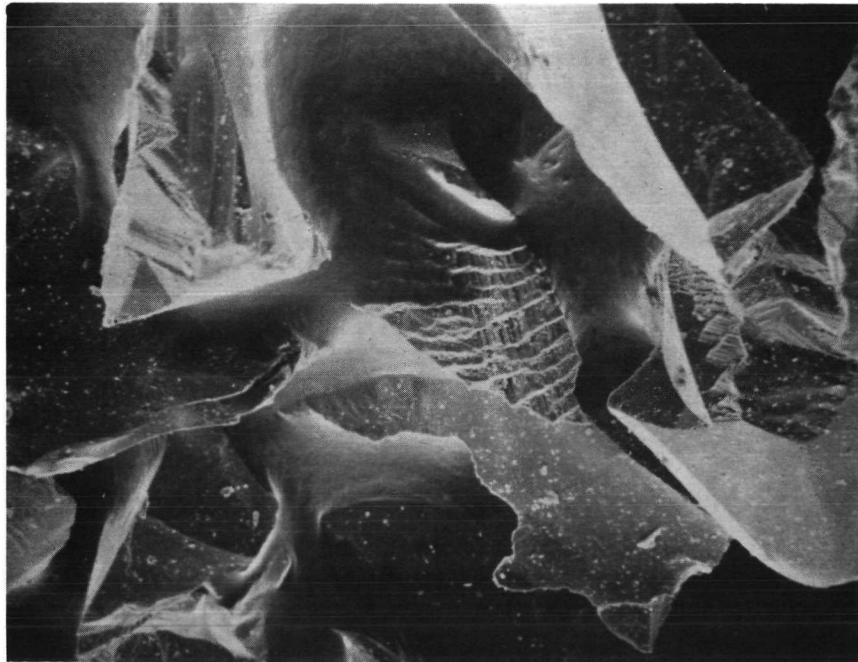
FIGURE 5 COMPARISON OF THEORETICAL AND EXPERIMENTAL REFLECTANCE OF CORUNDUM POWDERS

only after we included the wave optics additions to the basic geometrical optics model to account for the effects of particle edges¹⁰ and surface asperities. The edge effect which dominates at the lower particle sizes (d's) of the coarse-particle theory produces the $1/d$ dependence of the reflectance spectra in this region of high absorption by the particles. The fact that the $630 - 900 \text{ cm}^{-1}$ feature persists to relatively large particle sizes (120μ) was explicable only after scanning electron micrographs (Figure 6) revealed the presence of small step-like surface asperities for this material. If the number of these asperities is proportional to the surface area of the particles they will produce no d -dependent effect (see Figure 7). The surface density of these asperities is an adjustable parameter in the theory. The scale factor for the $1/d$ effect of the particle edges has been chosen, as discussed in (I), by means of reasonable, but somewhat arbitrary physical assumptions. We have used the same factor for all jagged particles. The general level observed in the reststrahlen regions for large particles in all of our work is higher than predicted by the theory. We believe this to be the result of deficiencies in our continuum-type radiative transfer model as the particles in such regions are highly opaque and therefore give rise to a large change in the radiative fluxes in a single layer of particles, which violates the assumptions of continuum model. Correction of the theory by means of a discrete layer model is required.



a.

**SCANNING ELECTRON MICROGRAPH OF STEPLIKE ASPERITIES.
30 μ CORUNDUM, 1000X**



b.

**SCANNING ELECTRON MICROGRAPH OF STEPLIKE ASPERITIES
AND ADDITIONAL FINES. 120 μ CORUNDUM, 500X**

FIGURE 6

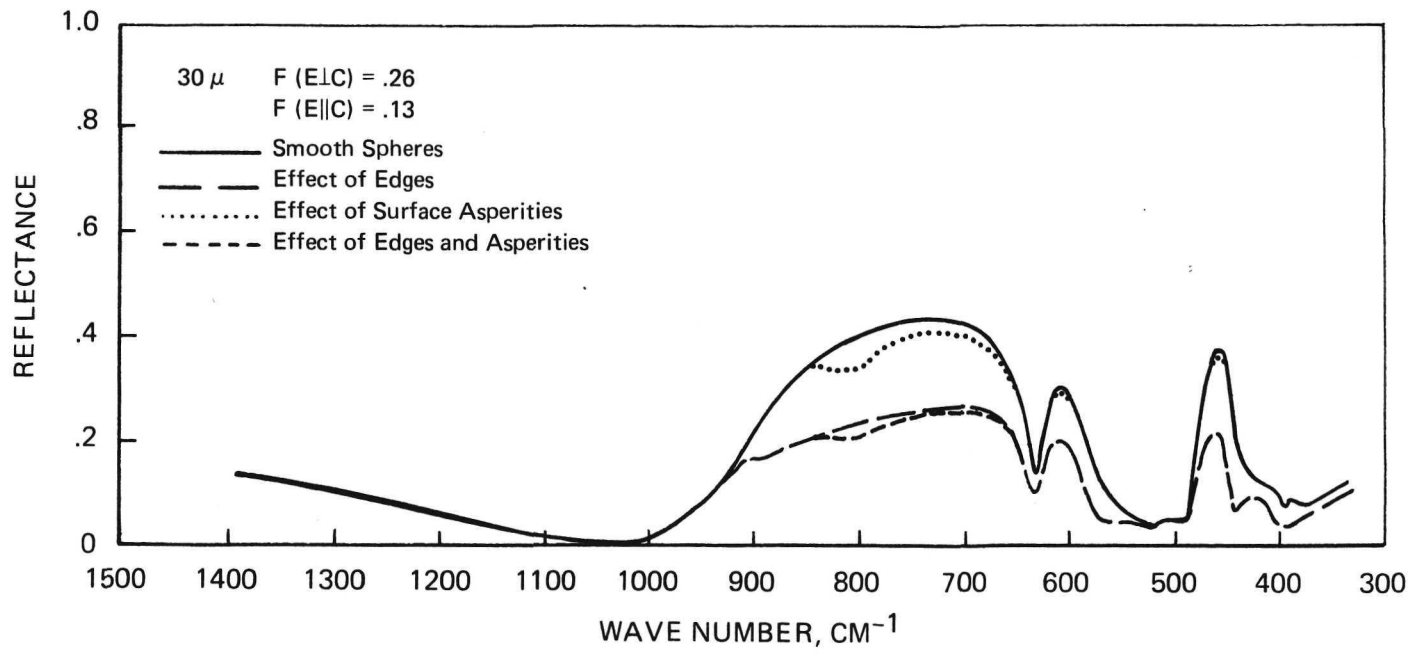
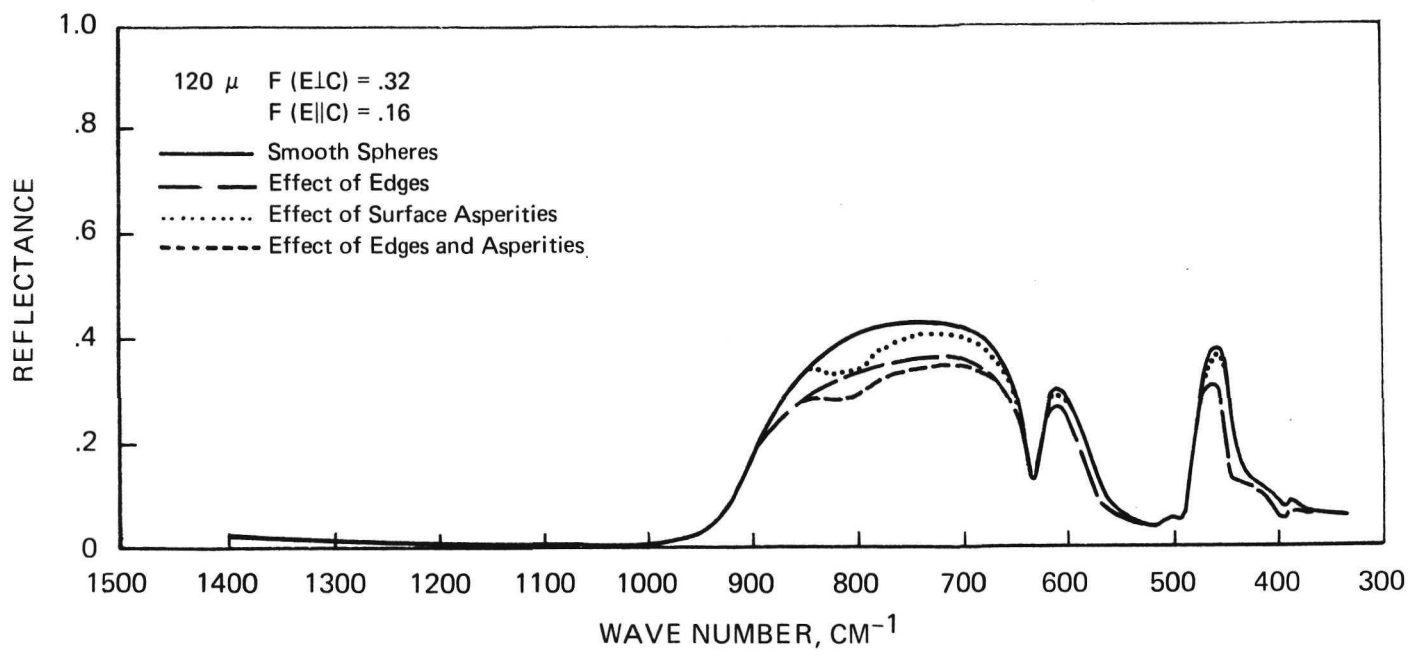


FIGURE 7 THEORETICAL SPECTRA OF CORUNDUM POWDERS DEMONSTRATING THE EFFECTS OF EDGES AND SURFACE ASPERITIES

As reduction of particle size generally results in increasing porosity (decreased packing density) and the fine-particle theory shows a decided porosity dependent effect owing to the dilution of the optical constants by high void fractions, we have carried out several experiments with the 0.3μ powder as a function of packing density. The results showed that as packing density is increased, the principal effect is to increase the spectral contrast in the two features at 450 cm^{-1} and 580 cm^{-1} . The data shown in Figure 5 for the 0.3μ sample were obtained by compressing it gently, thus increasing the volume fraction of material to 0.19 from 0.09 which is its state as prepared by our usual techniques.

The feature observed near 930 cm^{-1} in the theoretical spectra for the 3.5μ and 10μ cases would be considerably reduced if a particle size distribution had been used. It is to be noted that the peak in the experimental spectrum in this general region shifts to higher frequencies as the average particle size is reduced. This may very well be the effect of a contribution by the feature in question. The feature is caused by the onset of a refractive contribution to the scattering as the particles become more transparent for frequencies approaching the Christiansen frequency at 1020 cm^{-1} . This contribution falls off rapidly as the frequency approaches still closer to Christiansen point. A similar explanation applies to the theoretically predicted convex shape of the spectra at frequencies above 1020 cm^{-1} contrasted with the concave shape shown by the experimental data.

Figure 7 shows the theoretical corundum spectra as calculated for the 120 μ and 30 μ particle sizes by options in our program that either, 1) treat the particles as smooth spheres, 2) include both the effects of edges and surface asperities, 3) include edges, but no asperities, and 4) include asperities but no edges. As can be seen both edges and asperities are necessary to account for the particle size dependence in the reststrahlen regions and the shape of the 630 - 900 cm^{-1} feature in the 120 μ particle size sample. The amount of surface asperities used for these plots is $NV = 3 \times 10^{-7}$ cm (see I). This value, which corresponds to the "average" thickness of the asperity layer, indicates that a substantial change in reflectance is caused by a relatively small amount of these asperities. The effect is due to the high efficiency of absorption by ellipsoidal particles having $n \ll 1$ and $k \gg 1$ such as occur in reststrahlen bands. The amount of asperities appears compatible with photomicrographs of the particles. In a similar fashion the scale factor for the d dependence described in (I) by the value $b\sqrt{\lambda}/4\pi$ is somewhat arbitrary and might be slightly adjusted for a better fit to the data. At the present time $b = 0.9$ is used for all of our theoretical computations. However, as these factors are unlikely to be known precisely for arbitrary unknown samples, we have not yet optimized the fit.

The phenomenon of additional absorption by the needle-like dipoles that represent the edges and the more general spheroidal

dipoles that represent surface asperities was an important innovation in our coarse-particle theory. We therefore attempted to prove the reality of these effects by carrying out several critical experiments. The first of these was a comparison of a single crystal emittance spectrum of "randomly oriented sapphire" with the emittance spectrum obtained after substantial surface abrasion with 15 μ diamond paste. The results are shown in Figure 8 and a stereo pair of scanning electron micrographs of a replica of the surface are shown in Figure 9. Figure 8 shows that the anomalous shape of the powder spectrum in the 630 - 900 cm^{-1} region can be produced in a single crystal spectrum by providing surface asperities.

The second critical experiment was carried out by running two samples of corundum beads (Figure 10). These beads were produced for us by D. Spooner and R. Bechtold of the Lockheed Electronics Geophysics section. The mechanism of bead production is to feed small particles of corundum into a hydrogen-oxygen flame where partial melting occurs. After air cooling, small amounts of a number of other Al_2O_3 phases were observed by X-ray diffraction, but annealing at 1400°C for more than 4 hours removed the other phases, resulting in pure $\alpha\text{-Al}_2\text{O}_3$. Figure 11 shows scanning electron micrographs of the sapphire beads. Surface asperities similar to those shown for most samples of crushed corundum (Figure 6) still occur on these beads. Some non-spheroidized material still exists, principally in the smaller

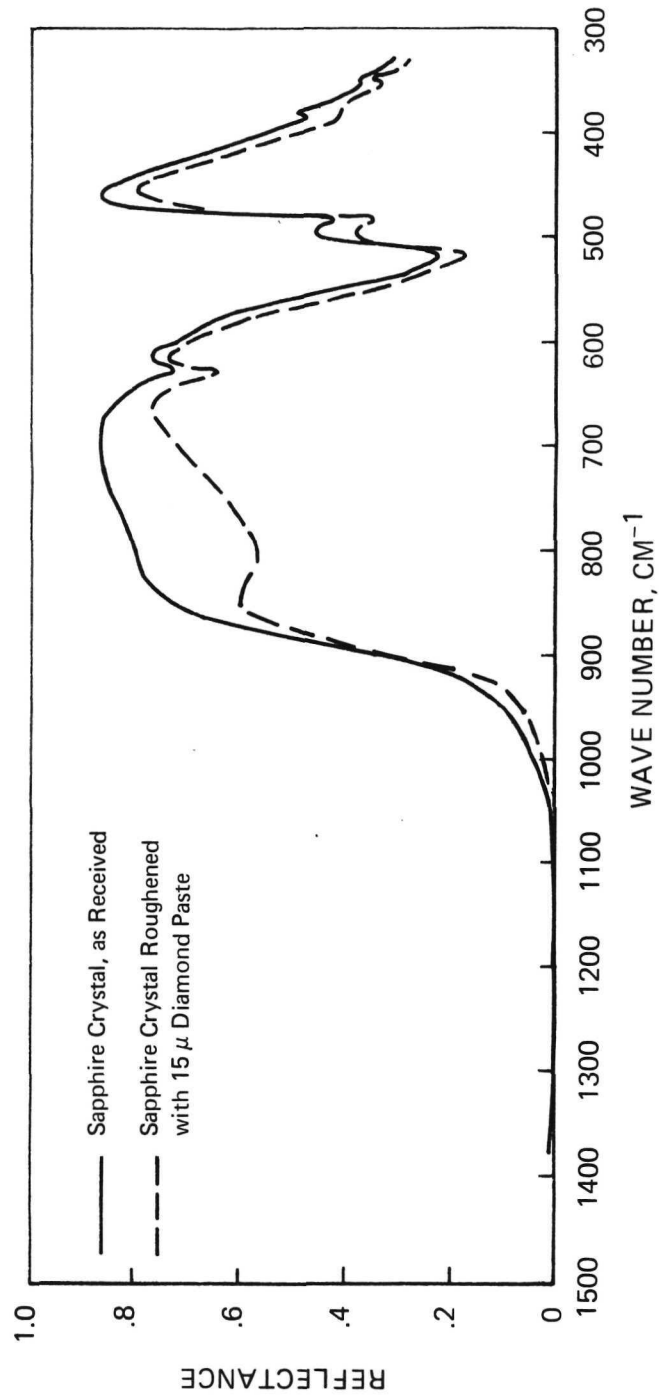


FIGURE 8 EXPERIMENTAL RESULTS OF ABRADING SAPPHIRE

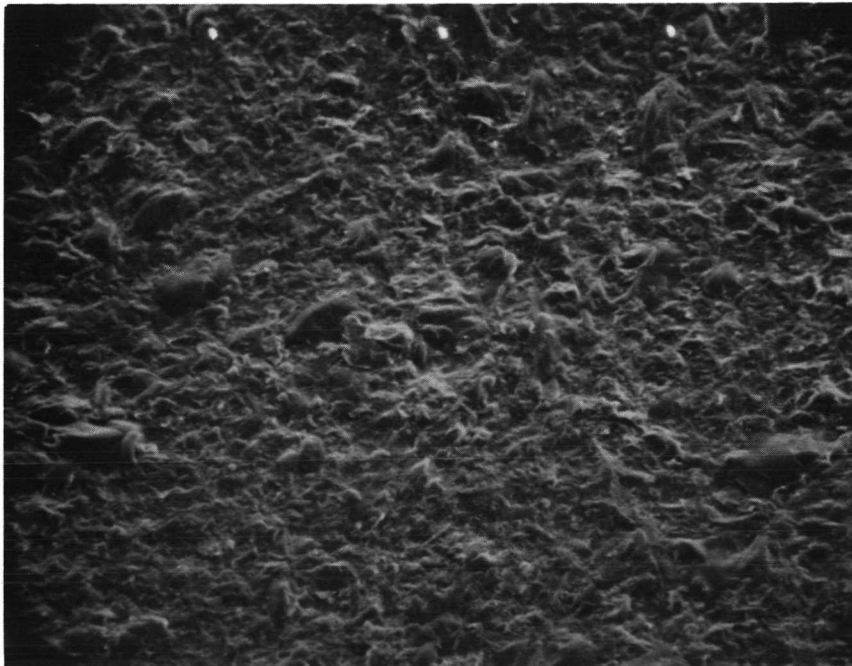
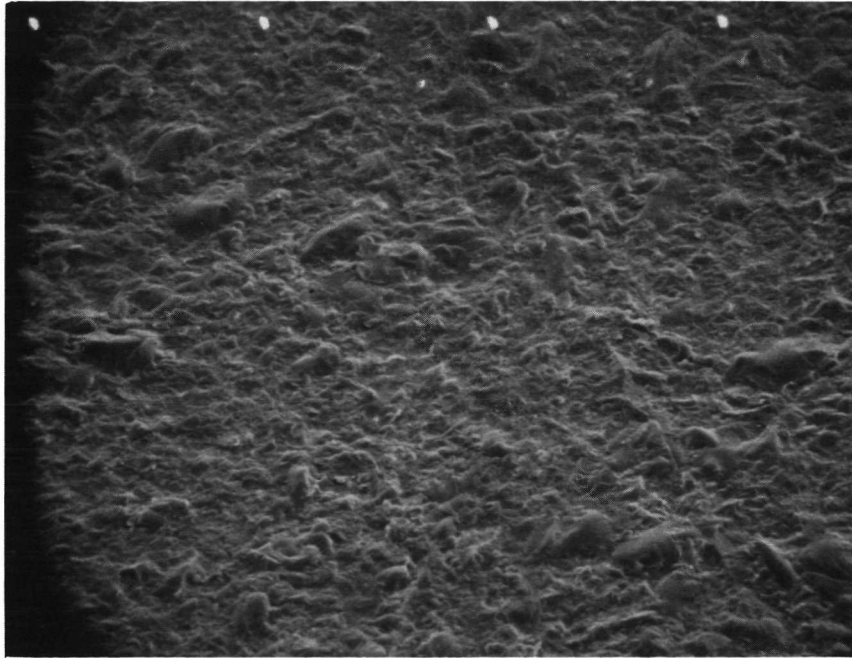


FIGURE 9 SCANNING ELECTRON MICROGRAPHS OF REPLICA OF ABRADED SAPPHIRE CRYSTAL.
PHOTOGRAPHS ROTATED TO SIMULATE THE ACTUAL SURFACE (470X).

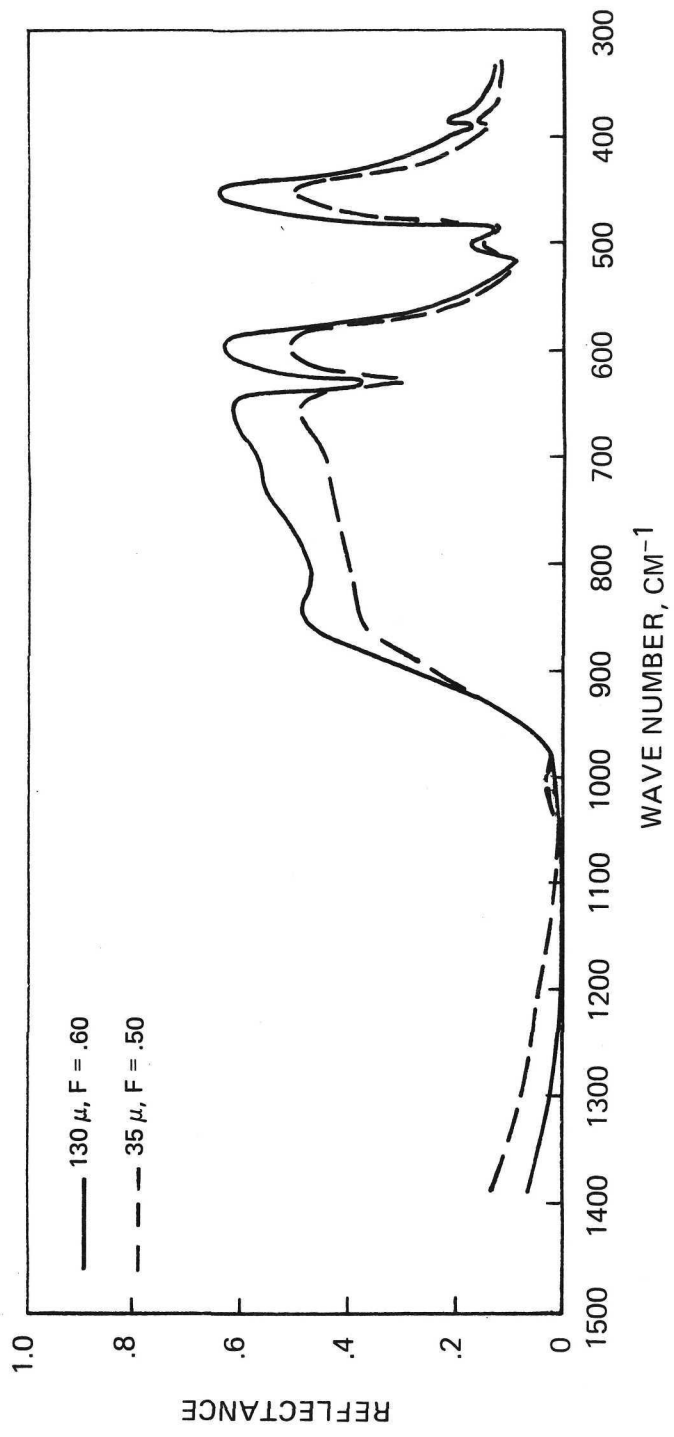
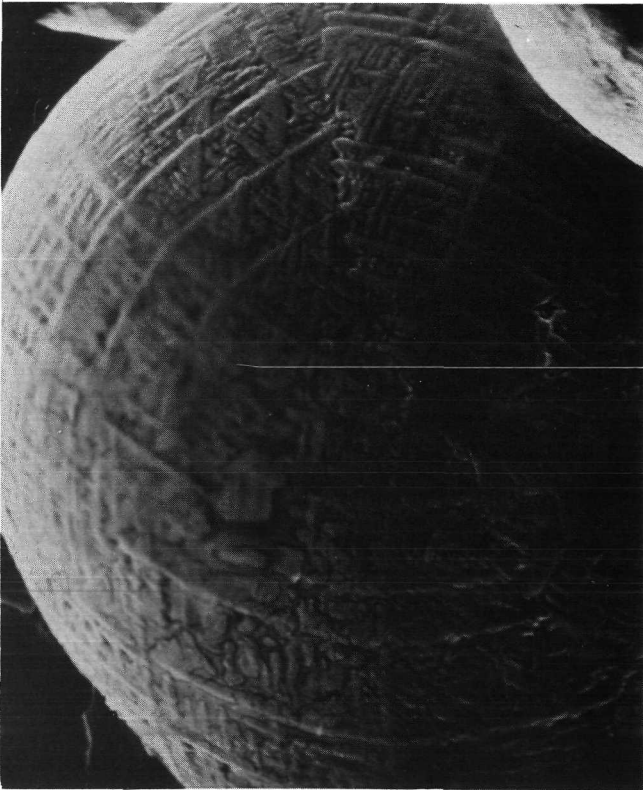
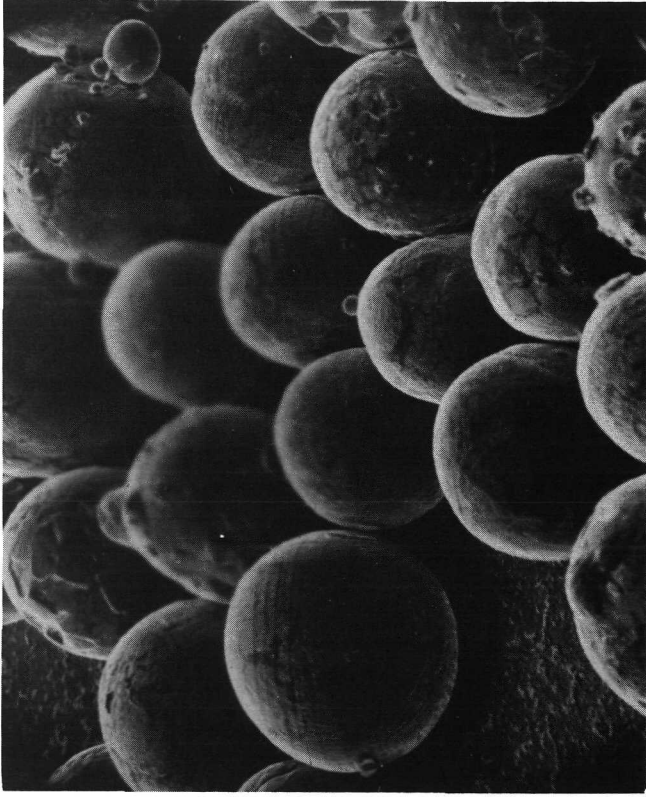


FIGURE 10 EXPERIMENTAL SPECTRA OF CORUNDUM BEADS



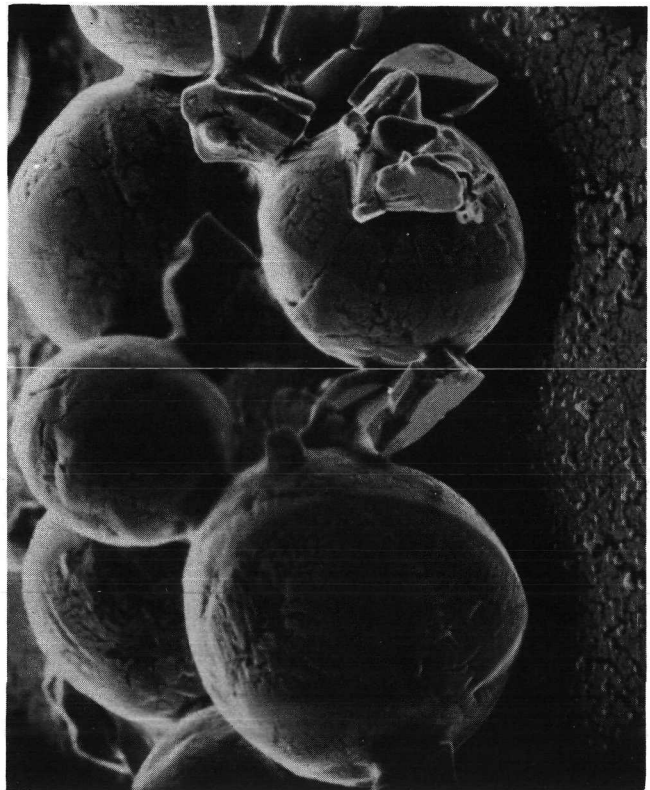
1080X

130 μ



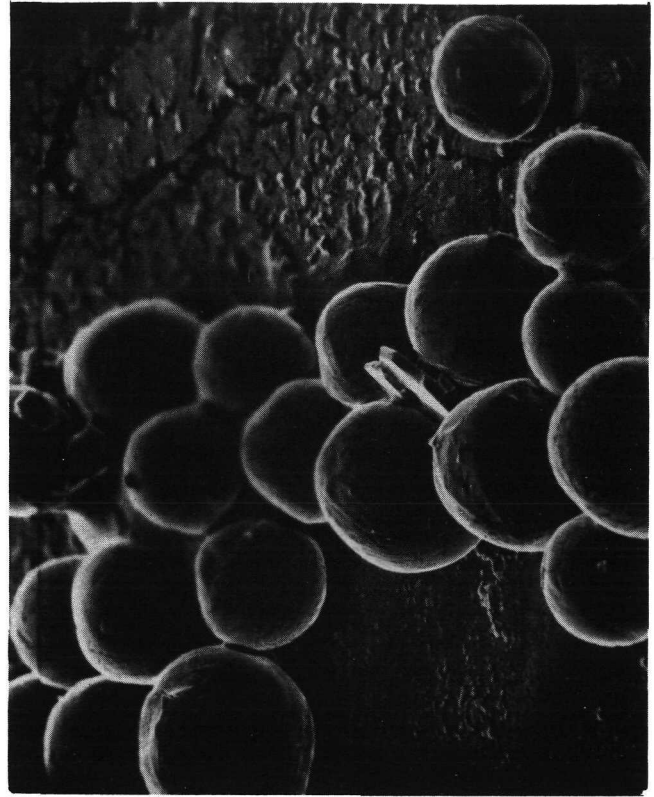
215X

130 μ



1000X

35 μ



500X

35 μ

FIGURE 11 SCANNING ELECTRON MICROGRAPHS OF CORUNDUM BEADS

sized sample but the particles have been largely de-edged. The changed level of the spectra in the $575 - 900 \text{ cm}^{-1}$ region and the noticeably smaller particle size effect when contrasted with the data for comparable sized particles in Figure 5 both confirm the assumed mechanism of edge absorption. We believe that the residual edges in the 35μ beads are largely responsible for the fact that the spectrum of this sample is still somewhat below the spectrum of the 130μ beads in the $575 - 900 \text{ cm}^{-1}$ region. We note that the level of the spectrum of the 35μ beads lies between the level of the 60μ and 120μ corundum powders. A dendritic growth can be seen in the scanning electron micrographs of the bead surfaces. This material is apparently responsible for the shape variation in the $630 - 900 \text{ cm}^{-1}$ region. Such a result is in accord with a number of theoretical experiments we have conducted by varying the shapes of the asperities (I).

The empirical bridging formula used in the region between 0.3μ and 12μ where both theories are being extended beyond their range of applicability appears to provide better results than either theory alone. This can be seen in Figure 12 where the results of both individual theories and the bridged theory are compared with experimental results for 3.5μ corundum powder. The relatively greater strength of the feature near 535 cm^{-1} with smaller particle size can be seen in the experimental data for the 3.5μ and 10μ particles (Figure 5). This is apparently a residue of the strong fine-particle feature shown by the theory (Figure 12) to peak near 570 cm^{-1} and it was used in an attempt to

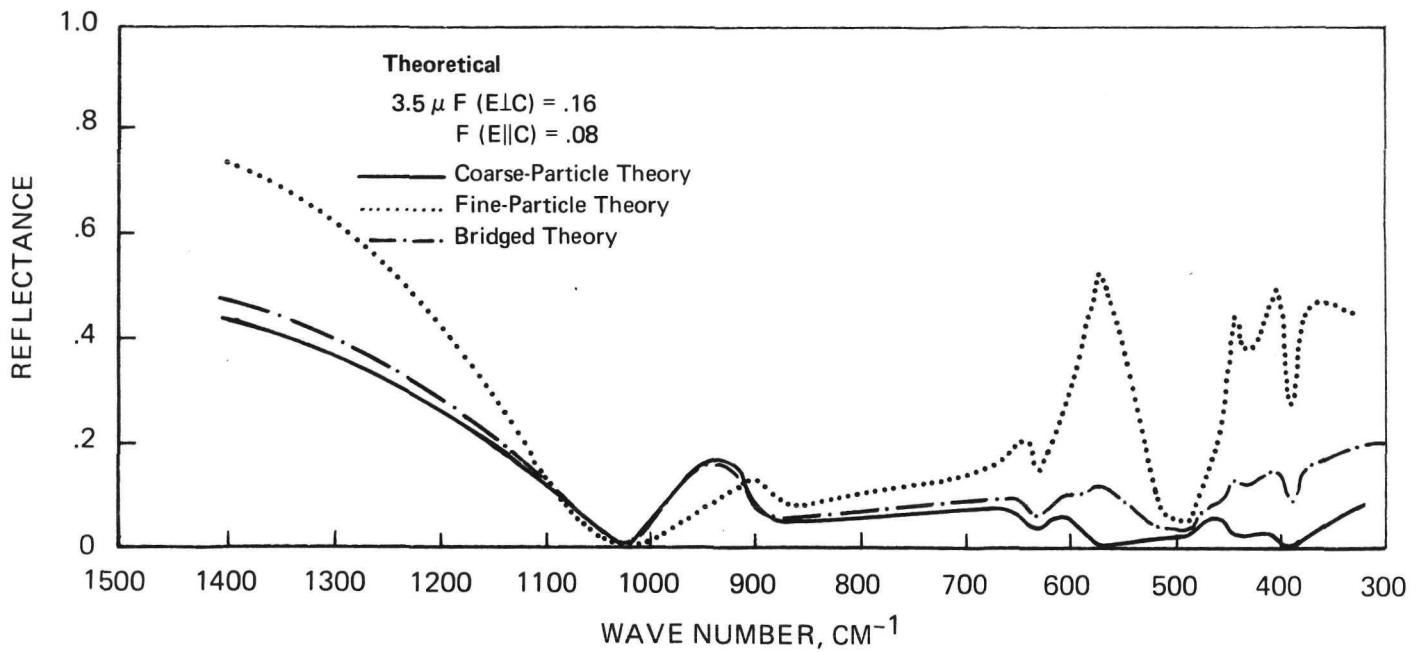
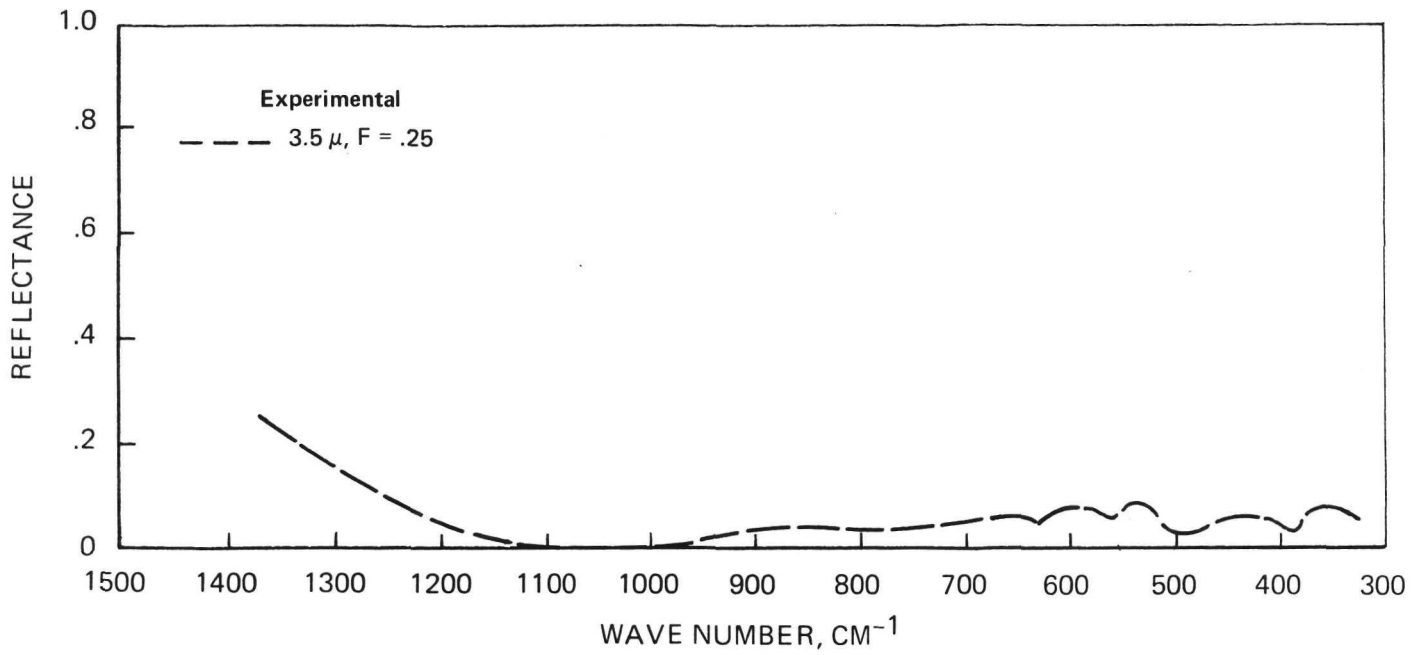


FIGURE 12 COMPARISON OF THE BRIDGED THEORY WITH THE COARSE-PARTICLE AND FINE-PARTICLE THEORIES FOR CORUNDUM

ascertain a reasonable bridging relationship. This relationship can probably be improved upon, but its optimization is somewhat dependent on the optimization of the other adjustable factors as well as the replacement of the monodisperse assumption with a more realistic particle size distribution function. The real particles, of course, have a size distribution, but the theoretical results shown in this paper are all for monodisperse particles.

Figure 13 shows the theoretical and experimental results for quartz powders. Once again the theoretical spectra for the coarser particles do not attain values as high as those observed in the experimental data in the reststrahlen regions. For the theoretical runs we assumed no surface asperities, since the $1100 - 1200 \text{ cm}^{-1}$ feature in quartz, which is similar to the previously discussed $630 - 900 \text{ cm}^{-1}$ feature in corundum, is only observed at the smallest particle sizes. However, clinging fine particles are commonly observed to be present in samples of larger particle size. In our theory they would act in much the same way as surface asperities. The data shown for the sample marked $0 - 20\mu$ provide a good example of the effect of such fines. The average particle size for this sample is near 10μ but the effect of the large number (but small volume fraction (0.007) of particles less than 2.85μ) is such as to give a more pronounced feature in the $1100 - 1200 \text{ cm}^{-1}$ region than is observed for the 4.5μ sample. In fact, the latter sample was obtained only after

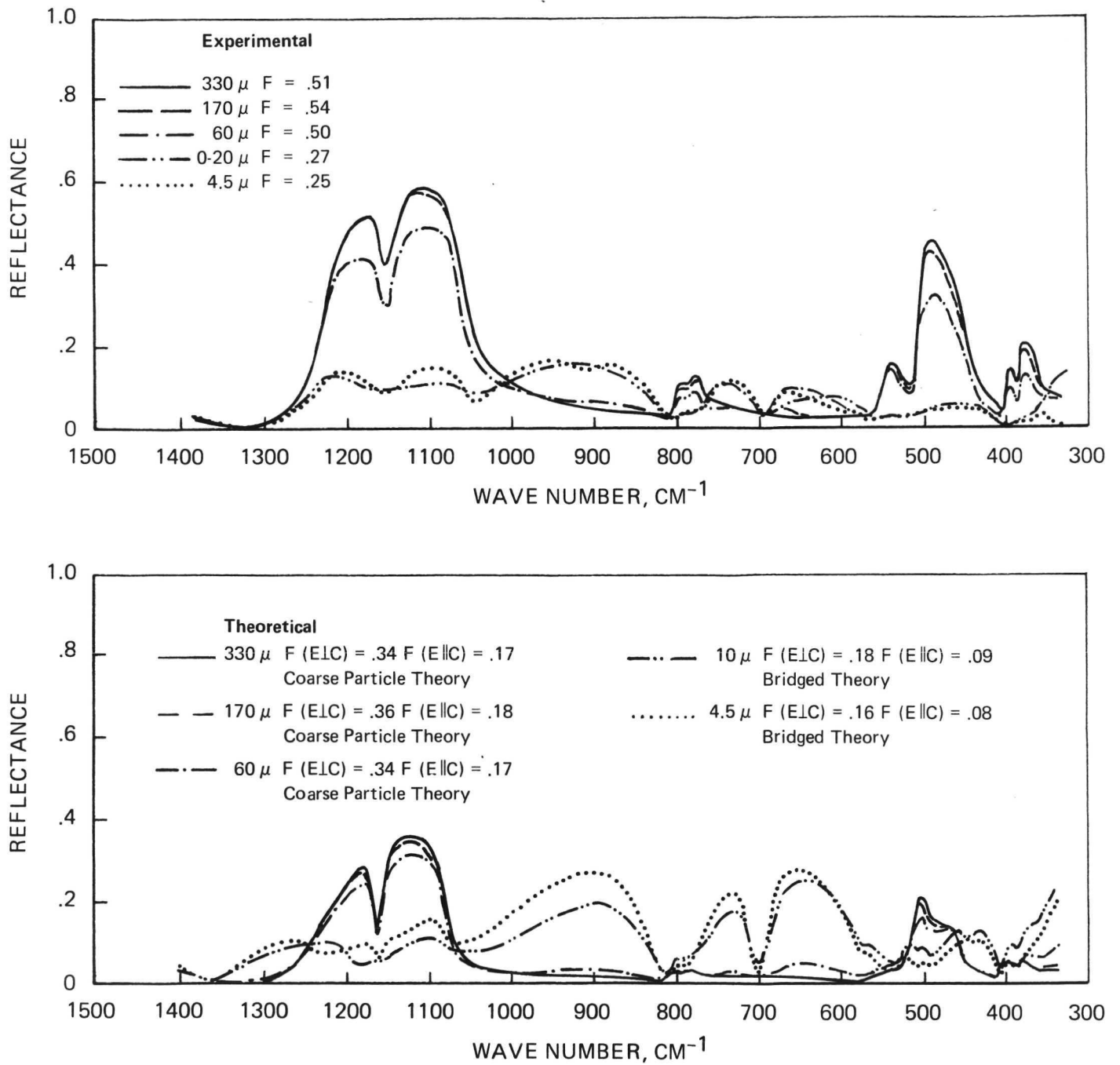


FIGURE 13 COMPARISON OF THEORETICAL AND EXPERIMENTAL REFLECTANCE OF QUARTZ POWDERS

a discrepancy was observed between the experimental particle size and that necessary to produce the feature, using the same value of b , the scale factor for the $1/d$ edge effect as was used for corundum. That discrepancy led us to reexamine the particle distribution in the 0 - 20 μ sample and the wide distribution found suggested that a narrower size range for this important experiment was in order. The 4.5 μ sample was then obtained from Duke Standards of Palo Alto, California. They indicate a 1 - 4.5 μ range for the sample. Our particle counts give 4.6 μ for the volume averaged size of this material, and it has an obviously narrower particle size distribution than the 0 to 20 μ sample. The 0 - 20 μ sample can be reasonably well fitted by the theory, if we use an asperity factor of $NV = 1.5 \times 10^{-5}$ cm to represent the clinging fine particles. This fit tends to confirm our idea that a small volume of clinging fines represents the same kind of extra absorption as do surface asperities. This effect was first observed in some of our early work, when some large quartz particle samples were run as received and after a wet-sieving procedure was used to remove the small volume fraction, but large numbers of clinging fine particles. Figure 14 shows photomicrographs of the samples of quartz powders of large particle size (170 μ) as received and after removal of the fines. Figure 15 shows the resultant spectra. The very large effect shown in Figure 15 was produced by a quite small volume fraction of the fines. As with the surface asperities, the effect appears most



As Received



After Removal of Fines

FIGURE 14 SCANNING ELECTRON MICROGRAPHS (95X) OF 170 μ QUARTZ PARTICLES

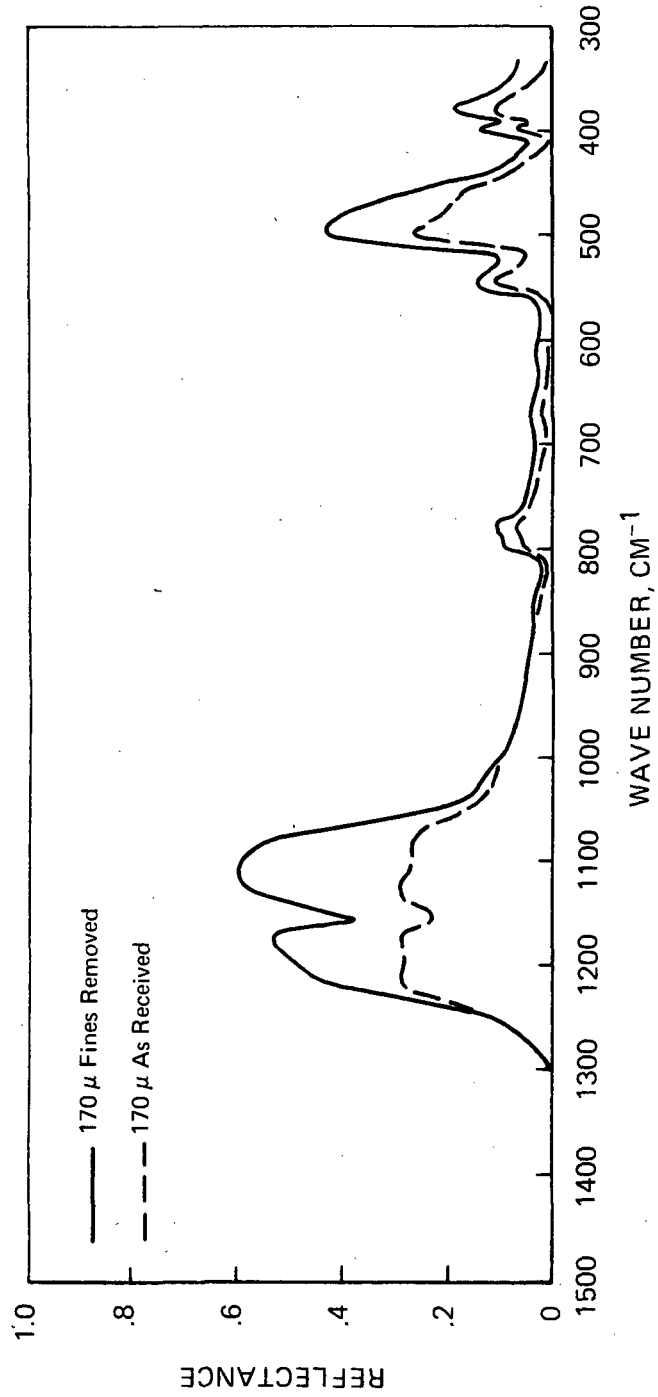


FIGURE 15 EXPERIMENTAL REFLECTANCE OF QUARTZ POWDER:
THE EFFECT OF A SMALL VOLUME OF FINES

significant in regions of high absorption.

As with corundum we note a theoretical feature (Figure 13) slightly to the high frequency side of the principal reststrahlen band for samples of intermediate particle size (4.5μ and 10μ). This feature does not occur in the experimental data which as before is actually for a particle size distribution. We again note an apparent shift of the reststrahlen feature toward higher frequencies as the particle size is reduced, and believe that the theoretical feature contributes to the shift. It is significant that the feature in question is also produced theoretically in Mie theory calculations¹² for the intermediate particle sizes.

Figure 16 shows experimental and theoretical results for garnet powders obtained from Barton Mines of North Creek, New York. The particle sizes used to calculate the theoretical spectra for these samples were obtained by microscopic techniques. They are slightly larger than the analyses given by Barton Mines. No Surface asperities were invoked in computing the theoretical spectra as microscopic examination showed no need for such a factor. The optical constants used as input parameters for the garnet data were obtained by fitting the spectrum of a large polished sample with a set of classical oscillator parameters, using the Lorentz line shape, by well known techniques¹⁵. The minor discrepancy observed near 500 cm^{-1} is thought to be due to a slightly deficient set of dispersion parameters, since an insufficient number of points were taken in the spectrum of the single crystal to resolve the small feature by least squares analysis.

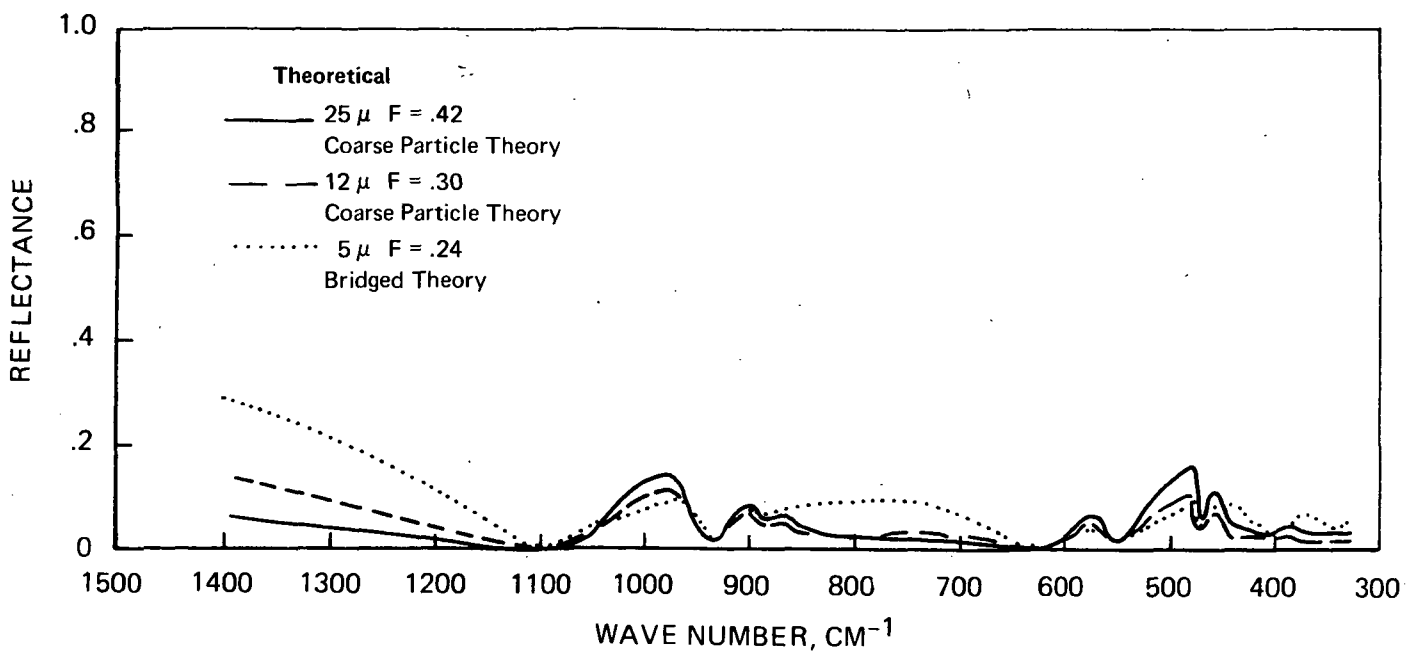
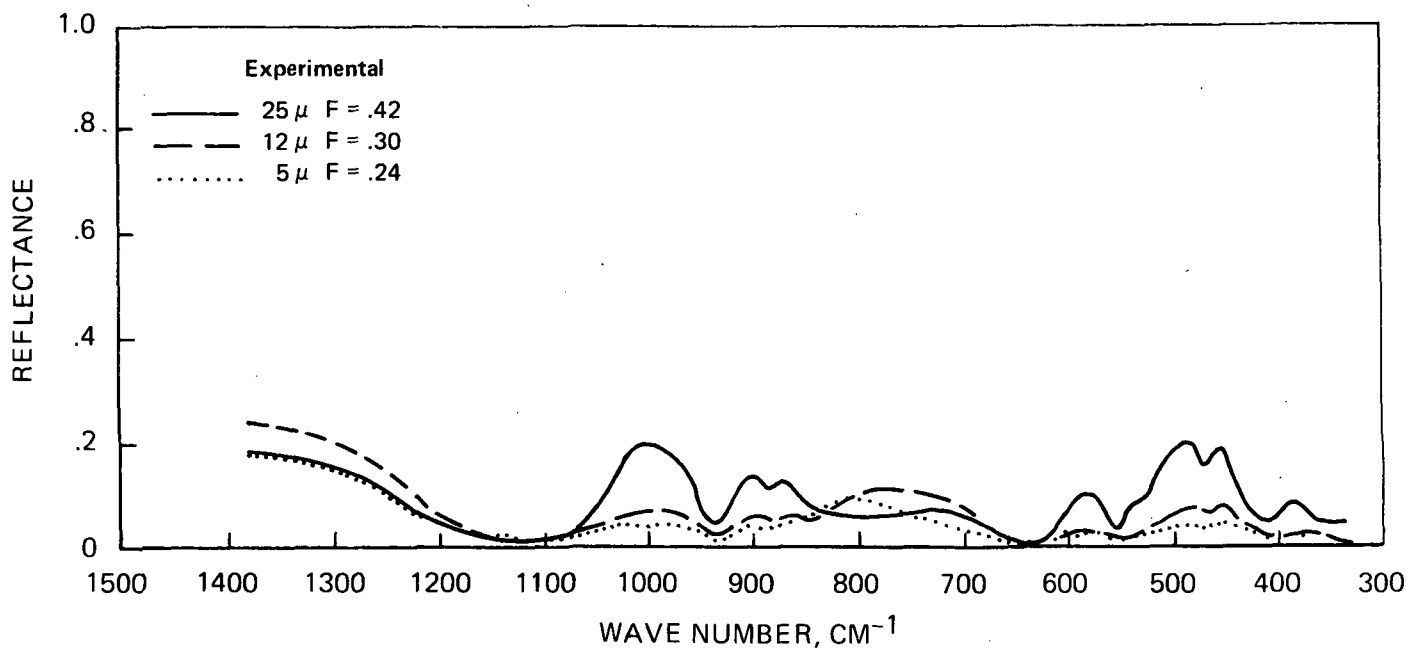


FIGURE 16 COMPARISON OF THEORETICAL AND EXPERIMENTAL REFLECTANCE OF GARNET POWDERS

The experimental and theoretical spectra of two mixtures of quartz and corundum powders are shown in Figure 17. Each mixture actually represents a mixture of four components as quartz and corundum are both uniaxial crystals so that two sets of optical constants are required to represent each of them. In these cases the experimental data were processed using the measured sample temperatures instead of applying the usual Christiansen technique as this technique cannot be justified for the case of mixtures. It is clear from the theoretical results shown in this figure that there is no frequency where the reflectance is zero such as would occur at the principal Christiansen frequency for a pure sample. The experimental results can be seen to confirm this theoretical conclusion. We note that the departure from zero reflectance is substantially greater for the finer particles. Thus arbitrarily setting the lowest point in the spectrum of a mixture at zero reflectance or unit emittance would make a substantially greater error for the finer particle data than for the coarser particle data. It is also quite significant that for the finer particle sample the lowest point in the spectrum occurs near 1160 cm^{-1} . This corresponds to the Christiansen frequency of neither component. Thus an identification technique based on the principal Christiansen frequency would fail for this mixture.

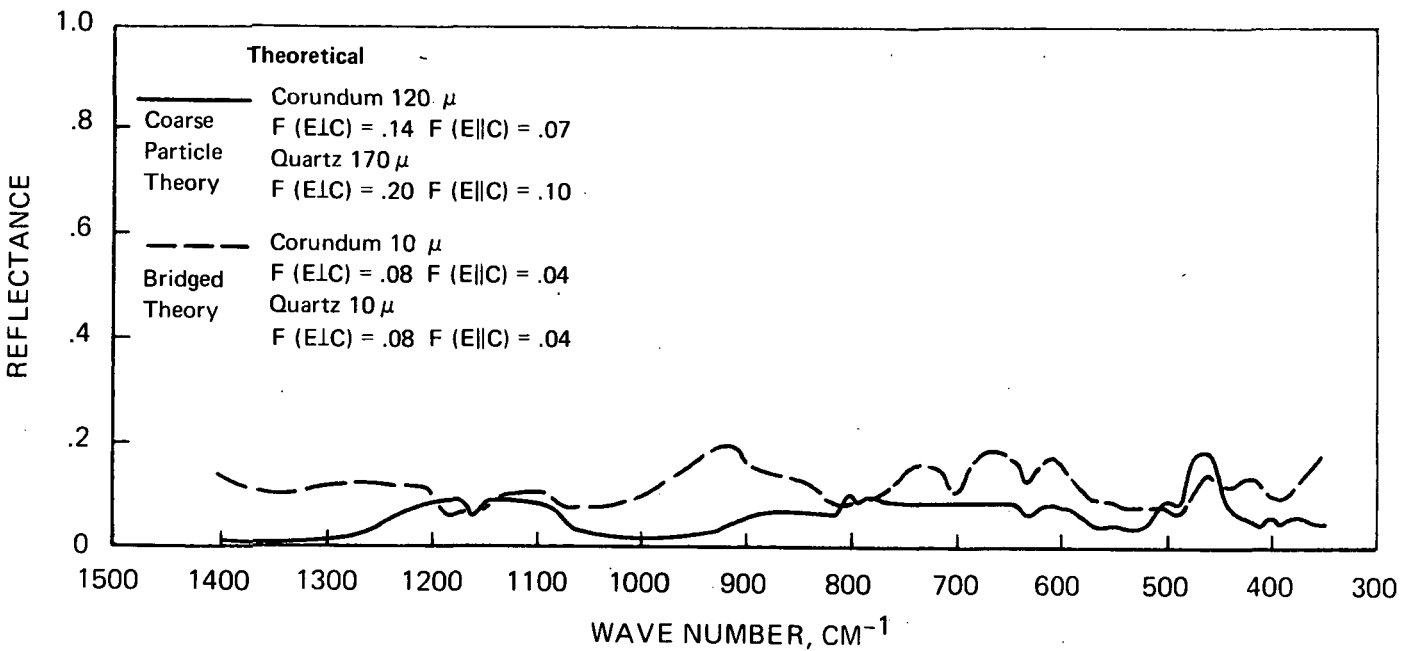
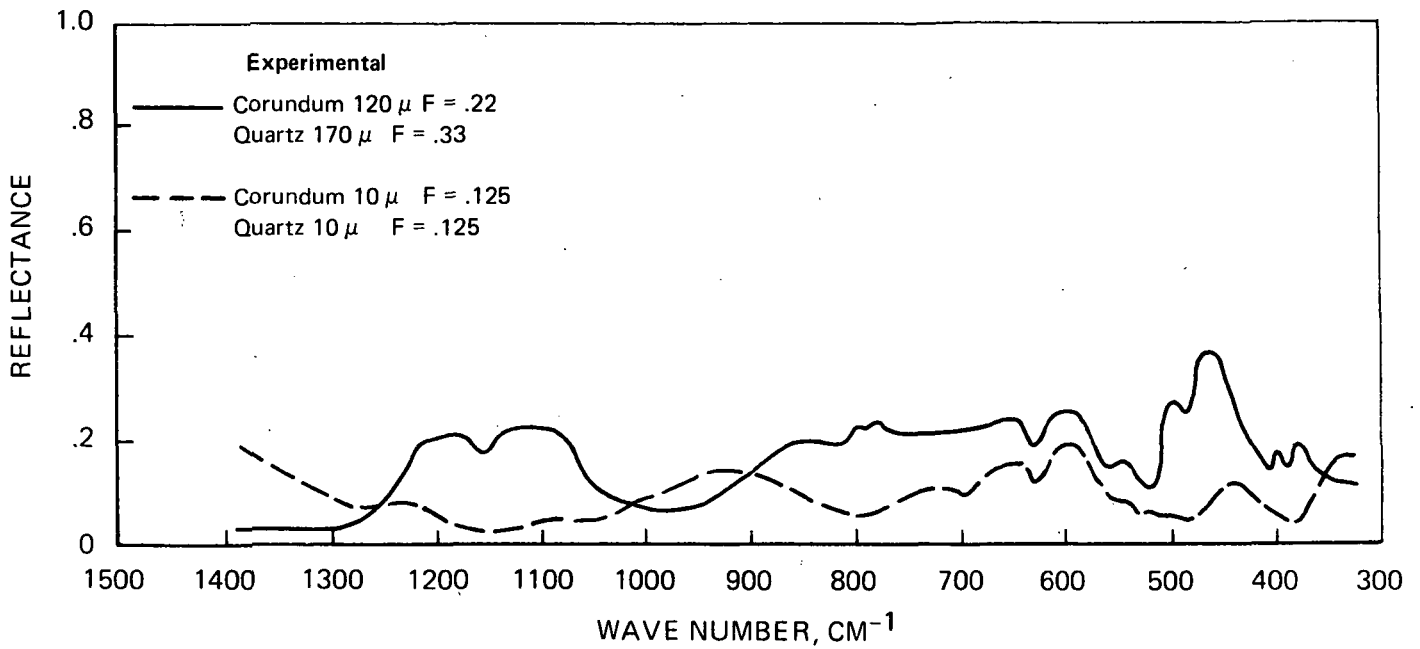


FIGURE 17 COMPARISON OF THE THEORETICAL AND EXPERIMENTAL REFLECTANCE OF MIXTURES OF QUARTZ AND CORUNDUM

For both mixtures the shapes of the theoretical spectra quite adequately represent the experimental data. While the general level of the theoretical spectrum of the mixture of finer size powders is also in good agreement with the experimental spectrum, the same cannot be said for the spectrum of the mixture of coarse powders. Once again we believe this discrepancy is due to our use of a continuum model rather than a discrete layer model.

It is worth noting, however, that for remote sensing applications where the problem is the inverse one of determining mineral compositions from spectral data, the general spectral level is much less important than the detailed structure of the spectral features.

ACKNOWLEDGMENT

The authors are indebted to E. M. Smith and P. C. von Thuna for assistance in the experimental program, and to I. Simon for a number of helpful discussions. This work was supported in part by NASA under Contracts NAS 9-10875 and NAS 9-8396 (Contract Monitor, W. Mendell) and the Air Force under Contract F19628-C-0353 (Contract Monitor, T. Rooney).

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IV. MODIFICATIONS IN THE THEORY

DISCRETE-LAYER MODEL

In Section II, pp 46-48, we outlined a discrete-layer model that describes the radiative transfer in media composed of closely-packed coarse particles much more realistically than can be done by a continuum model in which scattering and absorption are regarded as continuously distributed parameters. Since writing the paper we have extended the discrete-layer model to include the effect of diffuse scattering by asperities on the surfaces of the particles.

We consider a parallel beam of radiation of intensity I_o incident at an angle θ on an element of area dA of the surface of a particle. Then the reflected power is

$$P = R_o I_o dA \cos\theta \quad (1)$$

where R_o is the diffuse reflectance.

For simplicity we assume that the reflected radiation is Lambertian, so that the angular distribution has the form

$$dP = \frac{P}{\pi} \cos\chi d\omega \quad (2)$$

where χ is measured from the normal to dA and $d\omega$ is an element of solid angle.

We now resolve the diffuse radiation into discrete beams along the normal and tangential directions by means of our $\cos^2 \chi$, $\sin^2 \chi$ weighting factors. We find that the radiation splits equally between these directions:

$$P_{\text{normal}} = \frac{1}{2} R_o I_o \cos \theta \, dA \quad (3)$$

$$P_{\text{tangential}} = \frac{1}{2} R_o I_o \cos \theta \, dA \quad (4)$$

Here Eq. (4) represents the sum of four discrete tangential beams.

We next resolve these beams by the same method into back, transverse, and forward components relative to the original incident beam. The results, expressed as differential scattering crosssections, are

$$d\sigma_{\text{back}} = \frac{1}{2} R_o \, dA \left(\frac{1}{4} \cos \theta \sin^2 \theta + \cos^3 \theta \right) \quad (5)$$

$$d\sigma_{\text{transverse}} = \frac{1}{2} R_o \, dA \left(\frac{1}{2} \cos \theta + \frac{1}{2} \cos^3 \theta + \cos \theta \sin^2 \theta \right) \quad (6)$$

$$d\sigma_{\text{forward}} = \frac{1}{2} R_o \, dA \left(\frac{1}{4} \cos \theta \sin^2 \theta \right) \quad (7)$$

The sum of these crosssections is $R_o \, dA \cos \theta$, as it should be.

To obtain the total particle crosssections for transverse and backward external reflection we replace dA by $2\pi a^2 \sin \theta \, d\theta$ in (6) and (5) and integrate with respect to θ :

$$\sigma_{\text{refl,t}} = \pi a^2 \int_0^{\pi/2} R_o \left(\frac{1}{4} + \frac{1}{4} \cos^2 \theta + \frac{1}{2} \sin^2 \theta \right) \sin 2\theta \, d\theta \quad (8)$$

$$\sigma_{\text{refl,b}} = \pi a^2 \int_0^{\pi/2} R_o \left(\frac{1}{8} \sin^2 \theta + \frac{1}{2} \cos^2 \theta \right) \sin 2\theta \, d\theta \quad (9)$$

These formulas replace the specular reflection formulas (12) and (13) in Section II. The diffuse-reflection values of $\sigma_{\text{refl},b}$ and $\sigma_{\text{refl},t}$ are used in (74) and (75) of Section II in the calculation of σ_t and σ_b , which in turn enter into the evaluation of the reflectance and absorptance ρ and α of a monolayer of particles through (86) and (87) of Section II. Eq. (88) then gives the reflectance of the powder.

In the case of a mixture of particles of approximately equal size we calculate the quantities ρ_j and α_j from (86) and (87) for each kind of particle separately as though it were present alone at the volume fraction f . Then we determine the average values $\bar{\rho}$ and $\bar{\alpha}$ by the formulas

$$\bar{\rho} = \sum f_j \rho_j \quad (10)$$

$$\bar{\alpha} = \sum f_j \alpha_j \quad (11)$$

where f_j is the volume fraction of mineral j , excluding vacuum.

We have not yet generalized the discrete-layer model to the case of particles with a wide distribution of sizes.

CONTACT FACTOR

In our report of April 1, 1971, pp 9-12 we derived an expression for the effective fraction G of the surface area of a particle available for scattering after allowance is made for the effect of optical bonding near the points of contact with neighboring particles.

The "contact factor" G was given by

$$G = 1 - \frac{0.6 f \lambda_o}{d} \quad (12)$$

where f , λ_o , and d are the volume fraction of particles, the free-space wavelength, and the particle diameter, respectively.

In using this formula to correct the scattering coefficient S we encountered two difficulties. The first is that G drops abruptly to zero when d falls below a certain critical value. The second is that (12) does not allow for the fact that the concept of optical bonding breaks down in spectral regions where the particles are opaque.

We can avoid both difficulties by modifying (12) to read

$$G = e^{-0.6 f \lambda_o T/d} \quad (13)$$

where

$$T = e^{-4\pi kd/\lambda} \quad (14)$$

is the transmission factor along a particle diameter. One sees that in spectral regions where the particles are opaque, $T = 0$ and $G = 1$. Therefore the contact factor has no effect. As the particle diameter decreases the exponent of (13) increases and G tends to zero continuously.

~~We include the factor G in the theory as a multiplicative factor~~ which modifies the scattering crosssections σ_t and σ_b . The effect is to reduce the scattering, and therefore the reflectance, for transmitting particles only.

COMPARISON WITH EXPERIMENT

Figures 1, 2 and 3 show the results of using the discrete-layer model and the new contact factor to fit the experimental data of Figures 5, 13, and 17 of Section III. A considerable improvement can be seen in the reflectance level in the reststrahlen bands for the coarser particles. Some increase in the level resulting from the discrete-layer model for those regions where the particles are semi-transparent was effectively nullified by the contact factor.

The use of the coarse particle theory instead of the bridged theory for the 10μ particles would have resulted in a slightly lower spectral level but only at the expense of introducing some minor shape discrepancies. No attempt was made to reoptimize the bridging formula. The comparison between these three figures and their analogs in Section III demonstrates the essential validity of the discussion of Section II concerning the failings of a continuum model.

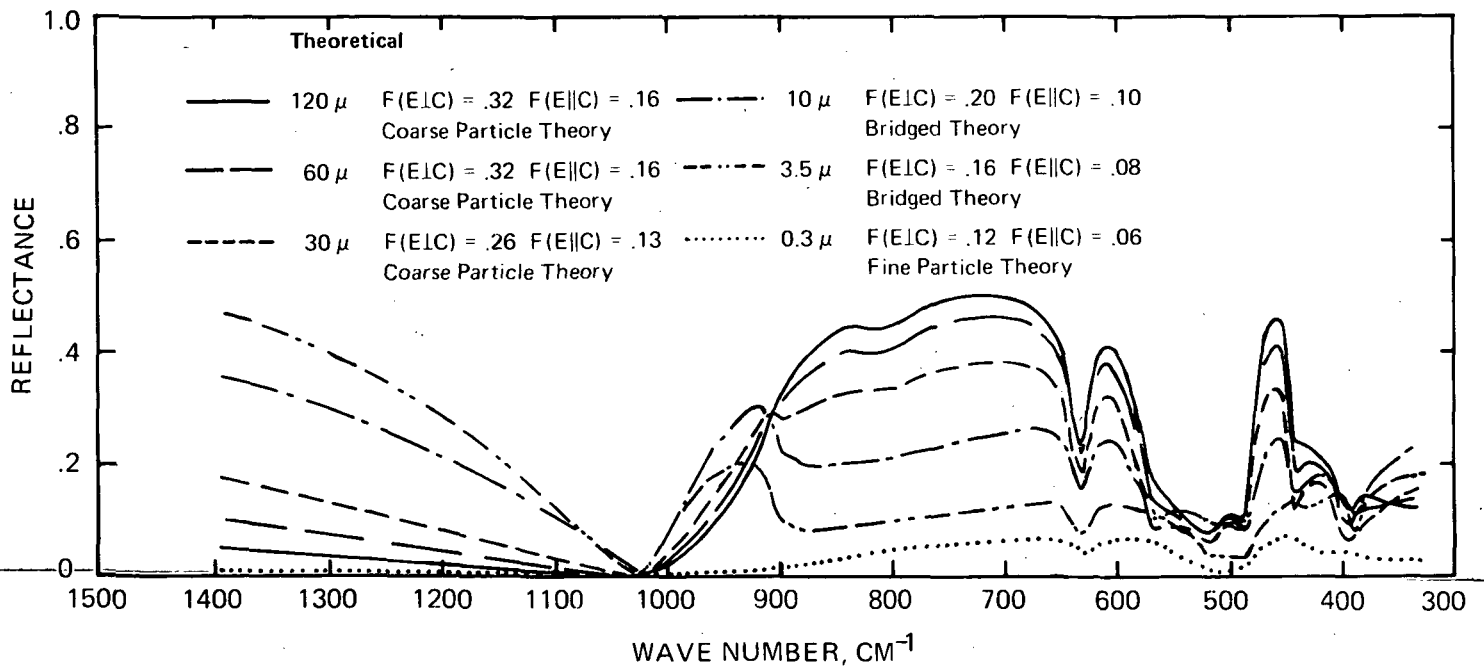
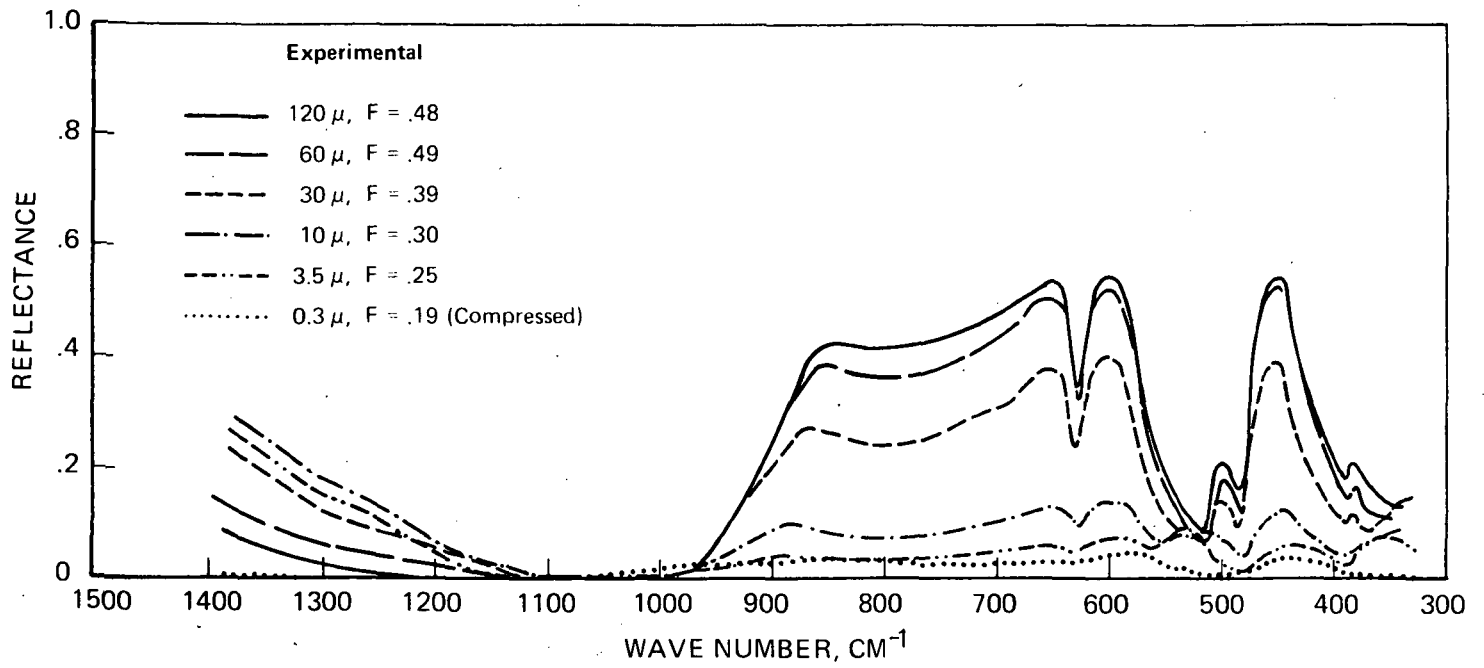


FIGURE 1 COMPARISON OF THEORETICAL AND EXPERIMENTAL REFLECTANCE OF CORUNDUM POWDERS

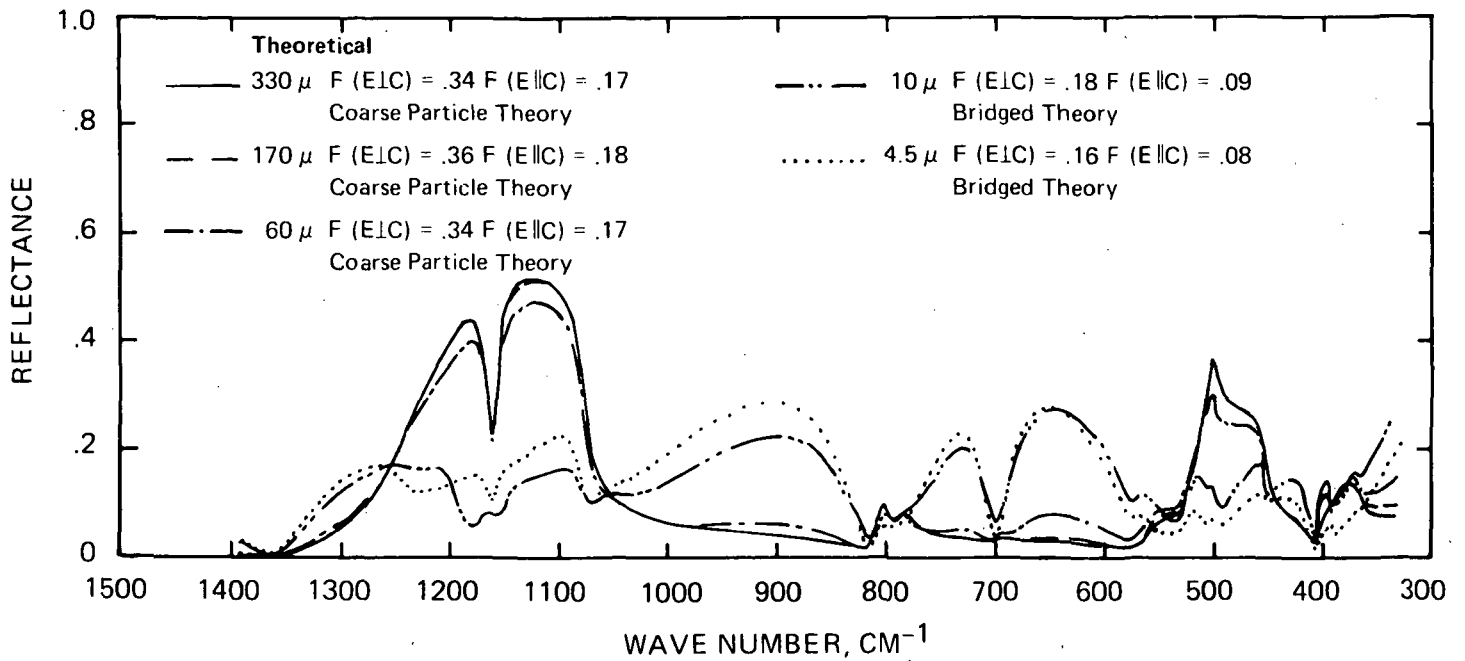
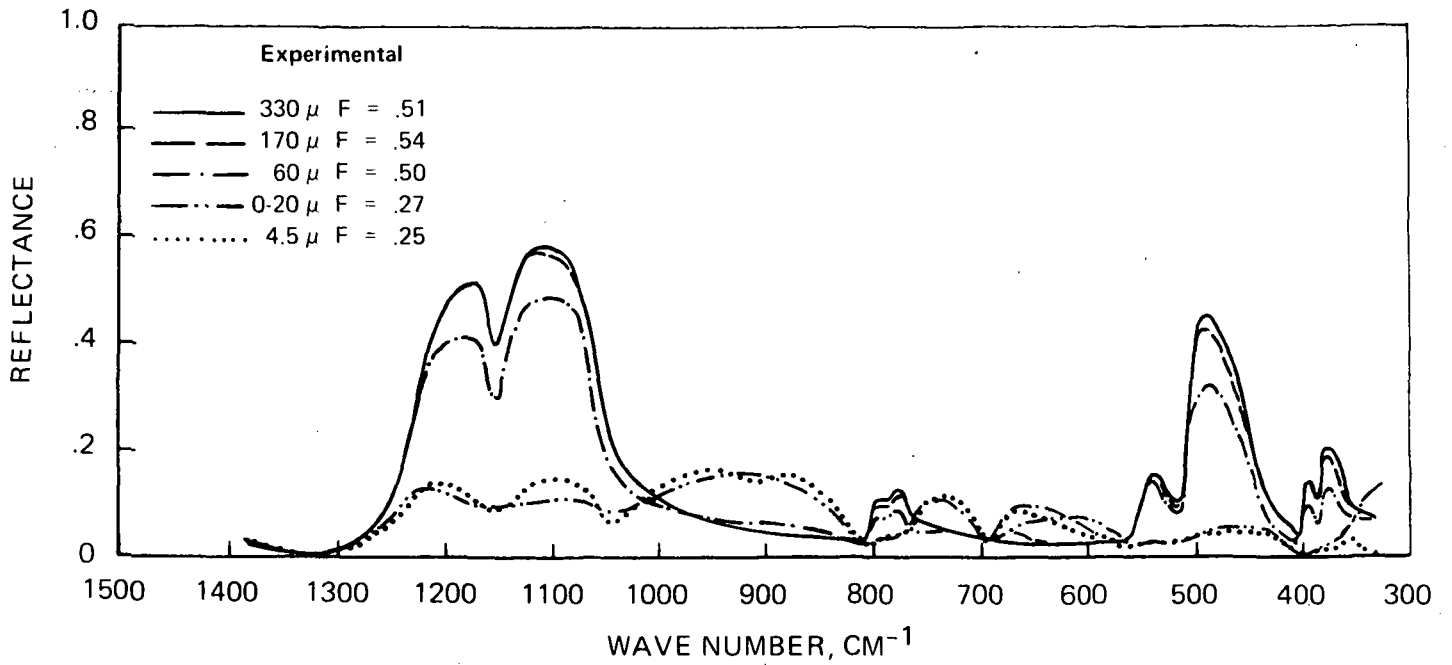


FIGURE 2 COMPARISON OF THEORETICAL AND EXPERIMENTAL REFLECTANCE OF QUARTZ POWDERS

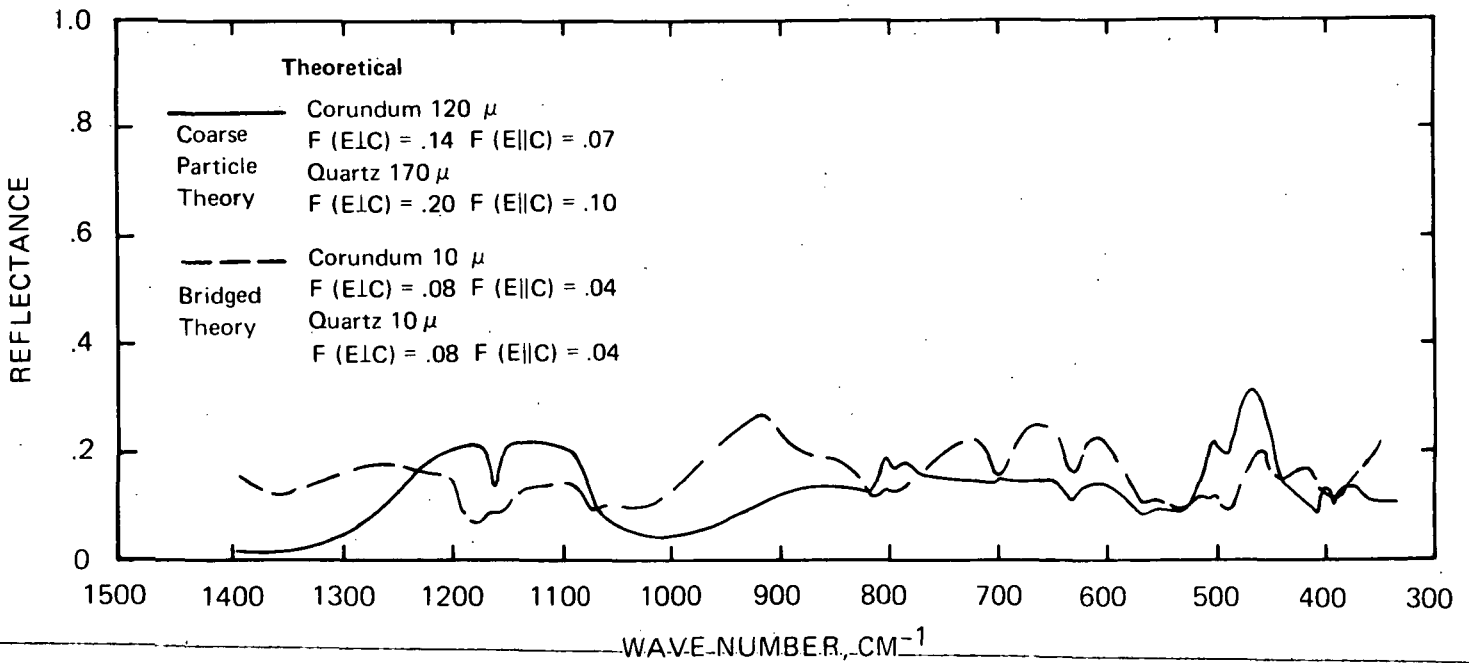
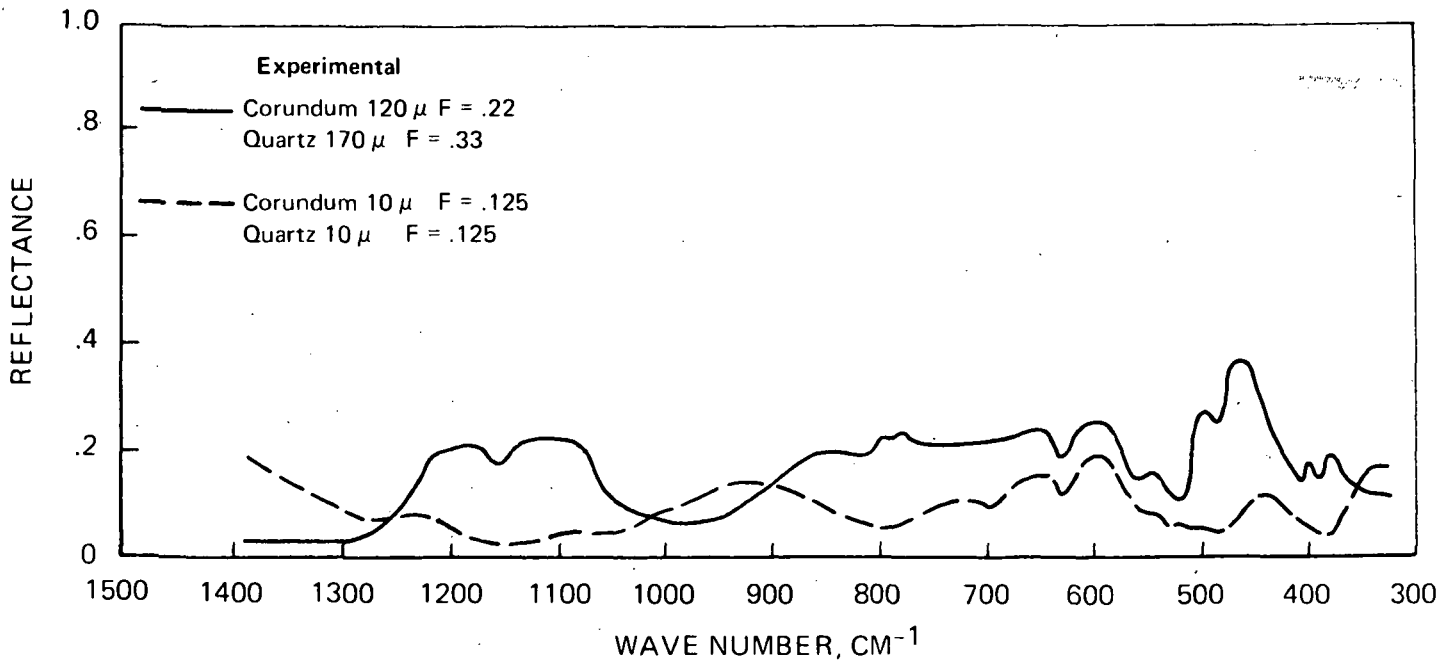


FIGURE 3 COMPARISON OF THE THEORETICAL AND EXPERIMENTAL REFLECTANCE OF MIXTURES OF QUARTZ AND CORUNDUM

V. THE INVERSE PROBLEM

The inverse problem consists finding a suitable way to use the theory of reflectance to analyze the experimental spectrum of an unknown mixture of powdered minerals in order to derive quantitative information about volume fractions (f_j) and particle sizes (d_j) of the minerals (j). It is assumed, of course, that the mixture contains a reasonably small number of candidate minerals of known optical constants.

Our theory indicates quite strongly that a simple linear mixing rule for reflectances is not a good basis for the inverse problem. For example, Figure 1 shows the variation of reflectance predicted by the continuum model, described in Sections II and III, as a function of the relative proportions of the two components in quartz-corundum mixtures. The five frequencies chosen illustrate examples of rather different behavior. The important point is the significant non-linearity of the curves, the extreme example being that for 630 cm^{-1} .

A further consideration is that no theory could possibly include all of the parameters needed to describe a real mixture, such as the particle size distributions of all component minerals, the nature of the surface asperities on each type of particle, the degree of aggregation, etc. This means that the theoretical spectrum will always differ

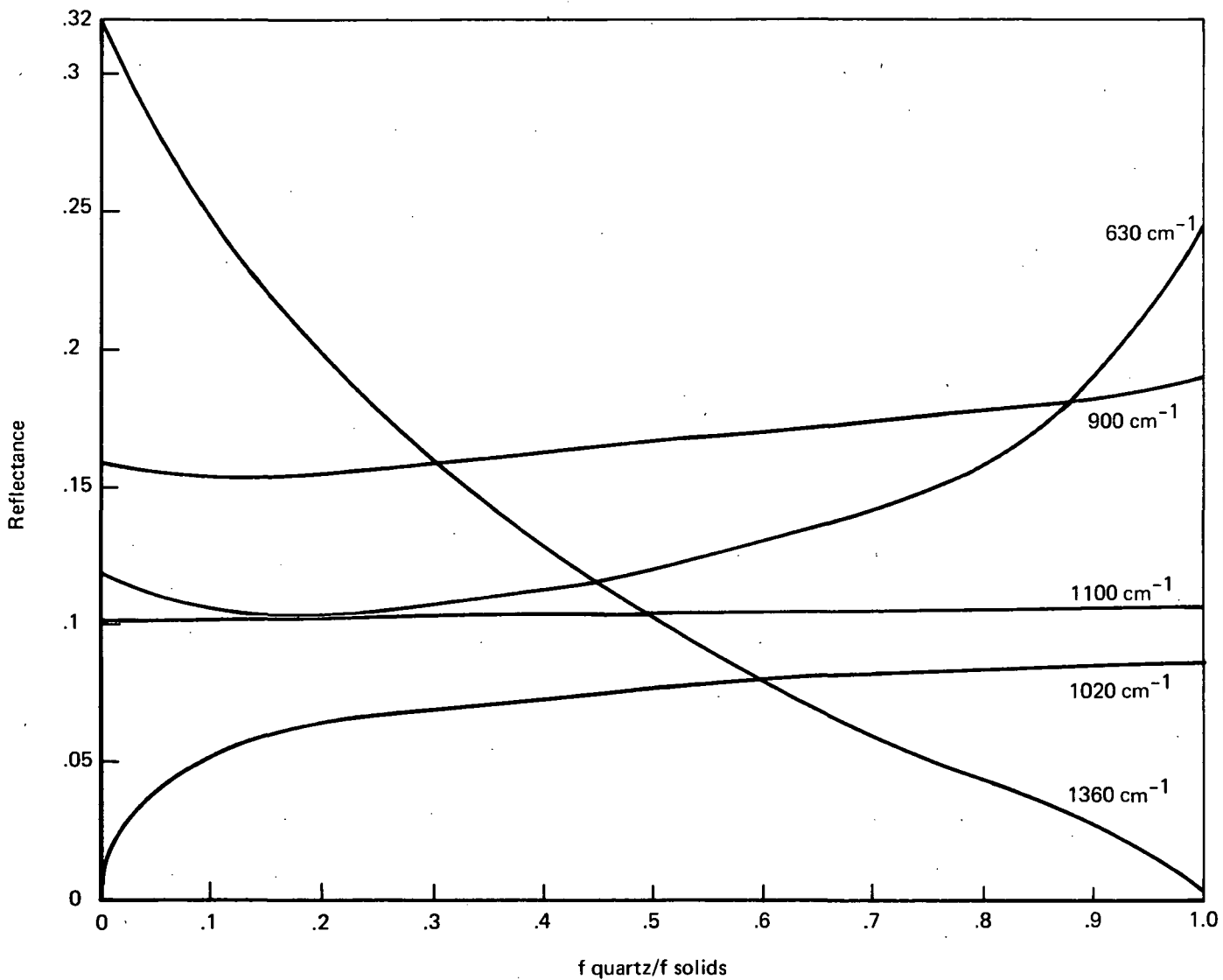


FIGURE 1 THEORETICAL VARIATIONS IN REFLECTANCE WITH THE MIXING RATIO IN MIXTURES OF CORUNDUM AND QUARTZ (10μ powders)

quantitatively from the experimental spectrum even with the optimum choice of the adjustable theoretical parameters. Such differences will occur in spectral level and in the detailed shape of the spectral bands. It appears, however, that the locations of the maxima and minima of the experimental and theoretical spectra agree quite well, for optimum values of the theoretical parameters, and that the ordering of the peaks and valleys with respect to amplitude is about the same for both spectra.

On the basis of these considerations we have developed a new quantitative method for comparing a theoretical spectrum with an experimental one. The method pays attention only to the ordering of the amplitudes in each spectrum at a number of frequencies selected from the experimental spectrum on the basis of features of interest such as peaks, valleys, abrupt changes of slope, etc. Some twenty to thirty frequencies are often sufficient to characterize an infrared spectrum.

After selection of N frequencies we compare the amplitude at the first frequency with the amplitudes at the other $N-1$ frequencies, and record a +, 0, or - at each of these frequencies depending on whether the amplitude is greater than, equal to, or less than that at the first frequency. We then repeat the comparison starting with the second frequency and comparing its amplitude with those of the succeeding $N-2$ frequencies. On continuing the process we obtain a triangular matrix of $N(N-1)/2$ elements which have the symbols +, 0, or -. This

matrix characterizes the characteristic shape of the experimental spectrum. The symbol O is assigned when the amplitudes at two frequencies are equal within some noise level.

We now compute a theoretical spectrum, for a set of assumed values of the f_j and the d_j , and characterize it in the same way as for experimental spectrum, using the same set of N selected frequencies.

Next we compare the theoretical and experimental matrices and count the number of coincidences in the two sets of $N(N-1)/2$ elements. The final score, which lies between 0 and $N(N-1)/2$, is a quantitative measure of the degree of fit between the theoretical and experimental spectra and can be used as the criterion for determining the optimum theoretical parameters.

A computer program was constructed to carry out such an analysis of our data. It was used for two mixtures of quartz and corundum. The coarse-particle theory results were used for the 170μ quartz- 120μ corundum mixture, and the bridged theory for the 10μ mixture. The procedure used the continuum model theory and so its facility for ignoring the spectral level was more severely tested than would presently be the case. The analysis was carried out in each example after choosing 20 spectral points from the experimental data. Figure 2 shows the coincidence scores for each of the chosen ratios of the components. The arrows indicate the measured ratios in the real mixtures, but it must

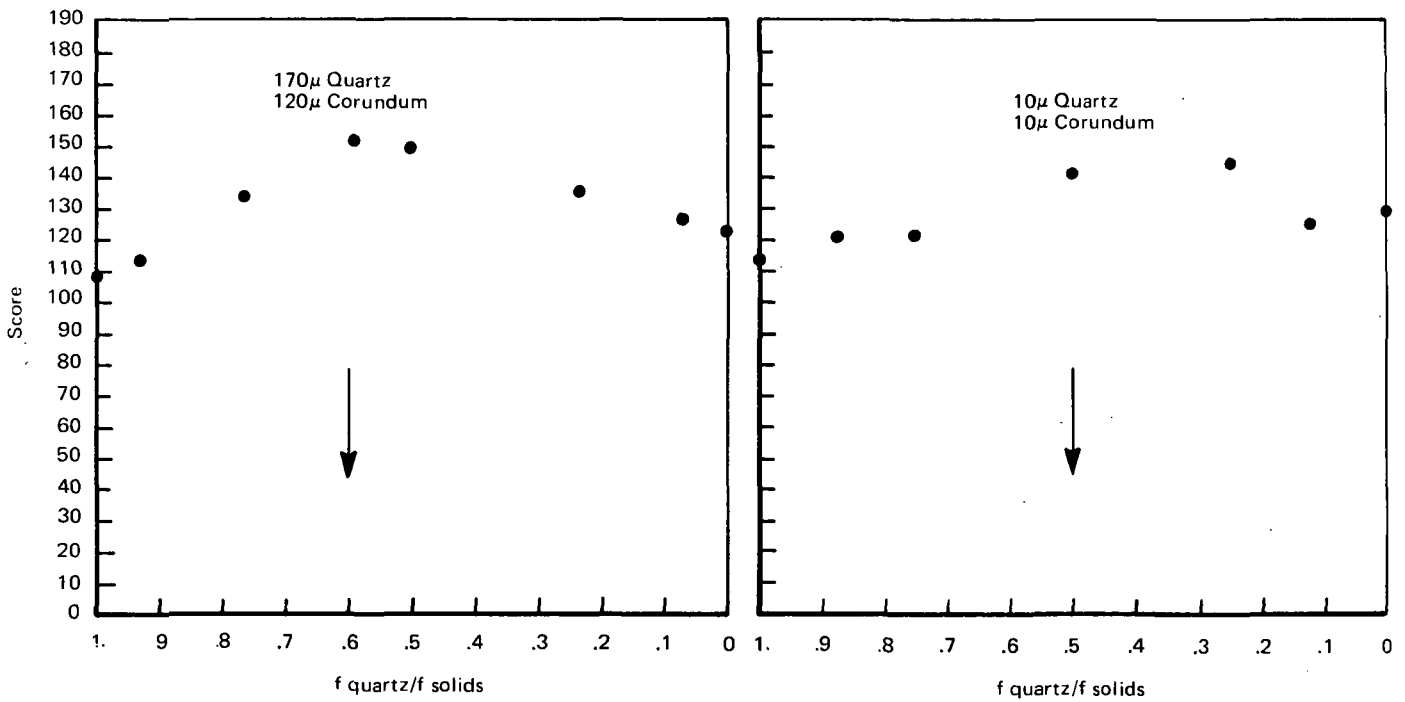


FIGURE 2 ESTIMATE OF THE RELATIVE AMOUNTS IN MIXTURES OF CORUNDUM AND QUARTZ

be recognized that this ratio is the average for the entire sample, and may not describe the top layer where the measured radiation originates. No attempt has yet been made to interpolate between the individual theoretical results although this could clearly be done. An increase in the number of spectral points would also almost certainly improve the technique as 20 points did not sufficiently describe all features in the spectrum. We consider that this initial success should be readily extendable to mixtures having more components.

VI. CONCLUSIONS

The research program described in the current and preceding^{1,2,3} reports has resulted in the development of a comprehensive new theory of the spectral reflectance (or emittance) of particulate materials. The theory has been shown to give excellent agreement with experimental measurements of the infrared emittance spectra of particulate minerals and mixtures of minerals over a wide range of particle sizes and packings.

The theory consists of two subtheories. The coarse particle theory is based mainly on geometrical optics, involving the effects of reflection, refraction and absorption. However, certain wave optical considerations are required as the model consists essentially of an assemblage of randomly contacting large spheres that have a degree of surface roughness. The surface roughness is represented by a distribution of ellipsoidal sub-particles for which the absorption and scattering are calculated by wave optics. The loss of scattering efficiency due to optical bonding at the points of contact is also calculated by wave optics. For fine particles, the theory includes a modification of the Lorentz-Lorenz theory to the case of ellipsoidal particles and the use of Rayleigh scattering by such particles considered to be immersed in the Lorentz-Lorenz medium. The two theories are empirically bridged to cover the intermediate region of particle sizes.

In addition, a new method has been devised to treat the inverse problem of analysis of the spectrum of an unknown mixture. This method correlates the characteristic spectral features with those predicted by the theory and searches for a best fit. At the same time it minimizes the effects of discrepancies in the absolute levels of the spectra. We anticipate that this method will prove very useful in the analysis of the spectra of planetary surfaces.

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VII. SUGGESTIONS FOR FURTHER IMPROVEMENTS

Further work could profitably be undertaken in a number of areas.

A. Methods of Treating Particle Size Distributions

This is especially important with regard to the discrete layer model which at present is valid only for mixtures of particles of comparable diameters.

B. Extension of the Method of Dealing with The Inverse Problem to More Complicated Mixtures

C. Further Study of the Effects of Thermal Gradients

We have carried out some work in this area^{1,2}, but a better understanding of the experimental data in this regard is still required.

D. The Effects of Particle Aggregations

E. Methods for Handling Biaxial Crystallites

F. Extension of the Theory to Particulate Media Having Near Zero Porosity

The coarse particle theory presently assumes that any ray passes from one material to the next via the void space. This may not always be true.

G. Resolution of a Minor Discrepancy

For intermediate size particles, there is an apparent discrepancy between theory and experiment near 930 cm^{-1} for corundum and near 1280 cm^{-1} for quartz. While we believe that the problem relates to the monodisperse assumption used in the theoretical computations (see Section III) the problem requires further investigation.

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