ROCK PROPERTIES OF AN ANDESITE AFTER FREEZING AND THAWING IN WATER

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Freezing and thawing damages of rock materials in many rock slopes have been well known to be severe under climatic cold region of Fukui Prefecture. Here, several weathering tests for hard rock material of andesite have been executed by freezing and thawing in water from -30° C to $+50^{\circ}$ C temperature. As the result, it is cleared that many rock properties tend to fall into decay gradually in linear relationship to the number of repeating cycles within 600 cycles.

1. Introduction

The change processes in quality of rock material due to weathering action are very complex to analyse, and these problems must be considered synthetically from both physical and chemical fields. In general, almost all weathering actions are divided into physical weathering due to freezing, insolation, swelling etc. and chemical weathering due to solution, oxidation and deoxidation, hydration etc .. Here, as one of the physical weathering actions, the degradation of rock properties after freezing and thawing in water has been considered experimentally. The volume of water expands about 9 % in the freezing time at 0°C. Such a large volume change of pore water will raise an enormous pressure among confined rock-forming minerals , and the porosity of rock material increase gradually by means of heaving and wedging of ice. This weathering action of repeating freezing and thawing in water is one of the most severe physical weathering.[1]

In winter, Fukui prefecture has been often covered with snow and ice. Freezing and thawing action occurs easily in this cold region. And, many rock slope failures due to degradation of rock or rock mass have often occured at the thawing time. In this region, the

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freezing and thawing action is considered to be one of the most important factors influencing on the weathering of rock mass or rock materials. And the rock properties, that is, density, absorption, porosity, longitudinal elastic wave velocity, compressive strength. Shore hardness, and so on are estimated to degrade gradually with repeating freezing and thawing, because the joints, cracks or pores of rock materials will expand and the rock-forming minerals are separated together due to expansive pressure of pore water at freezing time.

As laboratory tests, freezing and thawing tests in water against andesite have been executed under the condition of minimum freezing temperature -30° C and maximum thawing temperature $+50^{\circ}$ C. Then. the relations between several rock properties mentioned above and the number of repeating freezing and thawing cycle are analysed, and the effects of freezing and thawing action influenced on the engineering properties of rock materials are considered.

2. Preparation of Rock Samples and Rock Properties

As rock samples, hard rock materials of andesite have been selected. which are distributed for a wide range of Fukui prefecture.[2] And. two groups of fresh rock blocks of andesite have been sampled from construction site of new Takefu tunnel along National Road, route No.8. These rock blocks have been drilled perpendiculary to their sedimentary direction by use of core-cutter, and about one hundred of test pieces have been prepared as cylindrical pieces of diameter 50 mm and height 100 mm respectively. And these test pieces have been strictly selected ascertaining them to be sound without any crack or microfracture with the naked eye. Furthermore, the determination of required minimum number of test pieces has been done Photo.l Polarization microin the manner described in "Richt-



(a) Single Nikol $(\times 87)$



(b) Cross Nikol (×87) photograph of Andesite

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linien für Festigkeitsuntersuchungen an Gesteinen --- Entwurf 1963 des Internationalen Büros für Gebirgsmechanik ---"[3]. The variation of rock properties is considered though all of the two groups of rock blocks were sampled from the same wall rocks respectively. That is, the necessary number of test pieces is determined to be more than 6 test pieces as an unit for each weathering test. The relative deviation of uniaxial compressive strength for each group of test specimens have been calculated to be about 23 % alike from the test results examined for arbitrarily selected 10 fresh test pieces. And the uniaxial compressive strength has been modified in the manner described in the same standard, considering the shape effect of test specimen.

The name of these rock samples is "Altered Andesite". Photo.1 shows the microphotograph of mineral composition of the altered andesite belonging to group X, as an example. And this polarization microphotograph is taken at the initial stage before the weathering test starts. The external appearance shows green and black fine uniform grained rock-forming minerals containing white fine stripes, but there are not any porphyritic crystals in the naked eye. Under the microscope, porphyritic texture is observed, the size of phenocryst is smaller than 1.0 mm diameter and there is a little amount of phenocrysts. All phenocrysts are perfectly substituted by some carbonate minerals "Calcite". And the groundmass has pilotaxitic texture which is constituted from columnar structure crystal "Plagioclase" of $0.1 \sim 0.2$ mm length and chlorite minerals. The content of phenocrysts is only 1 %, and the other parts are constituted of groundmass. And the groundmass is divided into 44 % plagioclase, 36 % chlorite minerals, 8 % zeolite minerals, 7 % carbonate minerals and 4 % oxidized iron.

Table 1 shows the mean value and the standard deviation of initial rock properties for both X and Y group of rock samples. The number of test specimens measured and calculated here is 50 respectively for all rock properties except the uniaxial compressive strength. For X and Y group of rock samples, the density, specific gravity, uniaxial compressive strength and longitudinal elastic wave velocity have nearly almost the same value between X and Y group respectively, but the water content, absorption, porosity and Shore hardness have a different value between X and Y group respectively. The porosity of rock specimen n was calculated as n = $(G_s-G)/G_s$, G_s is specific gravity and G is apparent specific gravity, including the porosity of inner parts of rock which is entirely

Rock samp	Group X	Group Y	
Wet density	Y ₊ (g/cm ³)	2.637+0.007	2.637 <u>+</u> 0.028
Saturated density	$\chi_{gat}(g/cm^3)$	2.638 <u>+</u> 0.007	2.644+0.027
Dry density	$\chi^{\rm Sac}_{\rm g/cm^3}$	2.632 <u>+</u> 0.007	2.600 <u>+</u> 0.038
Specific gravity	G	2.669 <u>+</u> 0.004	2.712 <u>+</u> 0.025
Natural water conten	t w (%)	0.207 <u>+</u> 0.058	1.477 <u>+</u> 0.565
Absorption	W (%)	0.267 <u>+</u> 0.082	1.711 <u>+</u> 0.580
Wet porosity	n ₊ (%)	1.208 + 0.343	2.444 <u>+</u> 0.664
Saturated porosity	n(%)	1.151 <u>+</u> 0.333	2.238 <u>+</u> 0.637
Dry porosity	n _d (%)	1.411 <u>+</u> 0.341	3.575 <u>+</u> 0.815
Shore hardness	H	70.82 <u>+</u> 3.33	47.93 <u>+</u> 5.76
Uniaxial compressive strength	$O_{\rm D}^{\rm S}({\rm kg/cm}^2)$	1523.7 <u>+</u> 223.5	1589.8 <u>+</u> 273.2
Longitudinal elastic	wave velocity		
Wet	V _t (m/sec)		3473.8 <u>+</u> 222.6
Saturated	V _{sat} (m/sec)	3985.3 <u>+</u> 123.3	4059.8 <u>+</u> 294.8
Dry	V _d (m/sec)	3320.7 <u>+</u> 167.6	3020.5 <u>+</u> 154.4
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Table 1 Initial Rock Properties

isolated from the outside world.[4] And absorption W was calculated as (Saturated weight - Dry weight)/Dry weight. The modified uniaxial compressive strength $\mathscr{O}_{ extsf{D}}$ was calculated by the equation $\mathcal{O}_{D}=9$ $\mathcal{O}_{d}/(7+2d/1)$, \mathcal{O}_{d} is a measured value, d and l is a diameter and length of cylindrical rock specimen.[3] And the longitudinal elastic wave velocity was measured under about 15 kg compressive force by use of an apparatus for measuring the velocities of ultrasonic waves in rocks [5].

To investigate the thermal properties of these rock samples, specific heat and thermal conductivity have been measured respectively. The average specific heat has been given as 0.338 cal/g.deg by use

of water calorimeter, and the average thermal conductivity of saturated rock samples has been given as 1.13×10^{-2} cal/cm.deg.sec at normal room temperature. This thermal conductivity test has been executed in the manner described in LEE's Method.[6]

3. Experimental method and apparatus The weathering tests due to freezing thermal cycle test apparatus



Photo.2 General view of

and thawing of rock materials in water have been executed by use of a thermal cycle test apparatus, of which general view is shown in Photo.2. This thermal cycle test apparatus consists of 47cm*50cm*60 cm test chamber and thermal exchange box which is drived by 2 P.S. electric refrigerator and 2 KW heater. The temperature of test chamber is able to decrease until the minimum value -30°C by use of freon gas as an unfrozen liquid, while the temperature is able to increase until the maximum value +100°C by the heater. The operation of this apparatus is automatically controlled by relay controlled system, and is able to continue systematically for a long time. Therefore, freezing and thawing action is able to repeat automatically for the rock test specimens, as both upper and lower limit of temperature are set for the test chamber and for a center core of rock sample respectively. Here, the temperature of center core of rock sample is set between -30°C and +50°C, while the temperature of the test chamber is set between -30°C and +80°C. During the freezing and thawing repeated tests, the temperatures of test chamber, center core of rock sample and water in vessel are continuously recorded on a thermal recorder. For an example, Fig.l shows the one cycle record of their 3 thermal paths. In our freezing and thawing tests, 7.5 hours freezing process, 3 hours thawing process and 10.5 hours total process per one cycle are required. In this case, the temperature of center core of rock sample is lower than 0°C for 5 hours and higher than 0°C for 5.5 hours. And it is cleared that the temperature of water in vessel precedes slightly the temperature of center core of rock samples as shown in this figure.

All rock specimens are set and submerged into the vessel made of stainless steel having low specific heat and high thermal conductivity, of which capacity is 33cm*43cm*15cm, and thickness is 1.4mm. The freezing and thawing tests for group X and Y of rock samples



Fig.l Example of Thermal Path Curves



Photo.3 Example of Surface Exfoliation at 400 cycles

have been executed to 400 cycles and 600 cycles respectively, and the variations of several rock properties are investigated for every 50 cycles and 200 cycles respectively.

4. Experimental results and considerations

Table 2 shows the relations between several rock properties and number of repeated freezing and thawing cycles for rock group X and Y respectively. The mean values and the standard deviations of dry density χ_d , saturated density χ_{sat} , absorption W, saturated porosity n_{sat} , dry porosity n_d , Shore hardness H_s , longitudinal elastic wave velocity at saturated state V_{sat} and at dry state V_d are calculated from the whole data which have been measured for all rock specimens. But for uniaxial compressive strength O_D , they are calculated from only 6 data for every time.

For rock group X, the correlations between rock properties and number of repeated freezing and thawing cycles n are summarized as follows ;

Y d	= -	2.18 *	• 10 ⁻⁵	n + 2.633	(r=-0.86) (1))
Y sat	= -	5.53	• 10 ⁻⁶	n + 2.640	(r=-0.63) (2))
W	=	4.86 🕯	• 10 ⁻⁴	n + 0.286	(r= 0.62) (3))
nsat	=	3.67 4	• 10 ⁻⁴	n + 0.996	(r=0.27)) (4))
nd	=	8.52	* 10 ⁻⁴	n + 1.278	(r= 0.66) (5))
Hs	= -	7.96	• 10 ⁻³	n + 71.40	(r=-0.70)) (6))
$\sigma_{ m D}$	= -	3.49	n +	3208	(r=-0.70) (7)
_				for Rock	Туре А		
0- _D	= _	2.15	n +	1435	(r=-0.89) (8))
D				for Rock	Туре В		
V _{sat}	= _	1.15	n +	3916	(r=-0.93) (9)
Vd	= -	0.348	n +	3483.7	(r= 0.38) (10))

Here, r is the coefficient of correlation. Both dry and saturated densities decrease gradually with cycle number as shown in Fig.2 and Fig.3. Especially, the relation for dry density has high coefficient of correlation. It means that the dry density expresses more delicately the expansion or microfracture of pore structure due to weathering of repeated freezing and thawing than the satulated density. Absorption increases considerably with cycle number as shown in Fig.4. And both saturated and dry porosities tend to increase gradually with cycle number as shown in Fig.5 and Fig.6. These rock properties mentioned above are degraded owing to the failure of frame structure of rock materials due to the freezing expansive

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pressure of pore water, the difference of volume strain of each crystals or other factors. Shore hardness on the surface of rock specimen decreases gradually with cycle number as shown in Fig.7. As the weathering starts from the surface of rock specimen, the surface hardness drops down and surface exfoliation may be occured as shown in Photo.3. And, the uniaxial compressive strength decreases remarkably with cycle number as shown in Fig.8. In general, the compressive strength of weathered rock is controlled by the volume and shape of pores, and materials which occupy the space of pores.[7] The porosity n is divided generally into porosity of fissure n_f and the other parts of porosity n_{p} .[8] The larger the porosity of fissure n_{p} of rocks, the easier the structure tends to degrade due to freezing of pore water. In this figure, rock type A has a high uniaxial compressive strength, and there is not any



Photo.4 Break to pieces(Rock Type A)



Photo.5 Break in Two (Rock Type B) hair-crack in the specimen. Almost all the porosities of rock specimen are constructed of porosity $n_{\rm p}^{}$, and the rock specimen has crushed into fragments at the uniaxial compression test as shown in Photo.4. On the other hand, rock type B has a comparatively low uniaxial compressive strength, and there is some porosity of fissure n, which is very difficult to observe in the naked eye. The rock specimen has broken in two along some surface of crack as shown in Photo.5. Next, both longitudinal elastic wave velocity at saturated state V_{sat} and at dry state V_d decrease gradually with cycle number

Rock sample Cycle Number		Group X			
		103	149	200	255
Dry density	$\chi_{d}(g/cm^{3})$	2.631+0.008	2.631+0.006	2.629 <u>+</u> 0.006	2.628 + 0.007
Saturated density	$\int_{at} (g/cm^3)$	2.639+0.008	2.641 + 0.006	2.643 <u>+</u> 0.005	2.638 <u>+</u> 0.008
Absorption	W (%)	0.304+0.087	0.429 <u>+</u> 0.090	0.494 <u>+</u> 0.080	0.366 <u>+</u> 0.099
Saturated porosity	n _{sat} (%)	0.988 <u>+</u> 0.312	0.877 <u>+</u> 0.025	0.804 <u>+</u> 0.179	1.365 <u>+</u> 0.270
Dry porosity	n _d (%)	1.291 <u>+</u> 0.300	1.296 <u>+</u> 0.269	1.299 <u>+</u> 0.234	1.713 <u>+</u> 0.240
Shore hardness	H		69.82 <u>+</u> 2.15	70.79 <u>+</u> 2.37	70.60 <u>+</u> 4.91
Uniaxial compressive strength	$\mathcal{O}_{D}(kg/cm^{2})$		A 2241+423 B 1123 <u>+</u> 50	A 2985+617 B 1026 1 126	в 646 <u>+</u> 152
Longitudinal elastic	Saturated V _{sat} (m/sec)	3813.3 <u>+</u> 152.7	3647.5 <u>+</u> 99.3	3596.5 <u>+</u> 83.7	3670.7 <u>+</u> 115.2
wave velocity	Dry V _d (m/sec)	3677.2 <u>+</u> 176.3	3402.4+177.8	3395.8 <u>+</u> 111.2	3491.2 <u>+</u> 86.0

Table 2 Alteration of rock properties with cycle number of freezing and thawing for rock group X and Y

				Group Y		
303	350	409	125	200	400	600
2.628 <u>+</u> 0.008	2.628+0.004	2.621+0.005	2.582 <u>+</u> 0.037	2.604 <u>+</u> 0.037	2.617 <u>+</u> 0.039	2.608 <u>+</u> 0.052
2.637 <u>+</u> 0.007	2.638 <u>+</u> 0.004	2.637 <u>+</u> 0.005	2.633 <u>+</u> 0.024	2.643 <u>+</u> 0.026	2.652 <u>+</u> 0.027	2.647 <u>+</u> 0.039
0.332 <u>+</u> 0.116	0.372 <u>+</u> 0.097	0.585 <u>+</u> 0.178	2.080 <u>+</u> 0.661	1.467 <u>+</u> 0.504	1.368 <u>+</u> 0.053	1.810 <u>+</u> 0.904
1.202 <u>+</u> 0.271	1.106 <u>+</u> 0.143	1.125 <u>+</u> 0.200	2.654 <u>+</u> 0.603	2.253 <u>+</u> 0.687	1.969 <u>+</u> 0.752	2.478 <u>+</u> 0.846
1.539 <u>+</u> 0.328	1.470 <u>+</u> 0.162	1.710 <u>+</u> 0.169	4.291 <u>+</u> 0.812	3.490 <u>+</u> 0.775	3.117 <u>+</u> 0.815	3.708 <u>+</u> 1.237
69.62 <u>+</u> 2.62	66.58 <u>+</u> 2.58	68.32 <u>+</u> 2.72		45.30 <u>+</u> 4.96	45.15 <u>+</u> 4.89	51.19 <u>+</u> 4.49
A 2451+722 B 956 <u>+</u> 82	A 1990+611 B 805 1 12	A 1692+425		1329.0 <u>+</u> 524.7	1125.5 <u>+</u> 108.8	1351.8 <u>+</u> 606.6
3587.2 <u>+</u> 71.7	3527.9 <u>+</u> 77.1	3460.0 <u>+</u> 66.9	3913.0 <u>+</u> 203.8	4480.1+298.3	4268.3 <u>+</u> 367.0	4298.5 <u>+</u> 241.9
3345.2 <u>+</u> 75.1	3305.4 <u>+</u> 94.6	3316.0 <u>+</u> 130.8	3537.3 <u>+</u> 268.2	3409.4 <u>+</u> 208.9	3146.2 <u>+</u> 241.3	3001.9 <u>+</u> 210.9



Fig.4 Variation of Absorption W with cycle n





Fig.6 Variation of Dry porosity n_d with cycle n

as shown in Fig.9 and Fig.10 respectively. In this case, the longitudinal elastic wave velocity at saturated state V_{sat} should be more higher evaluated for these weathering tests than the other rock properties, because the relation between V_{sat} and number of cycle n has a very high coefficient of correlations.



Fig.8 Variation of Uniaxial compressive strength \mathcal{O}_{D} with cycle n





Fig.9 Variation of Saturated longitudinal elastic wave velocity V_{sat} with cycle n



For rock group Y, the correlations between rock properties and number of repeated freezing and thawing cycles n are summarized as follows :

=	$3.22 \times 10^{-5} \text{ n} + 2.594$	(r = 0.55)	(11)
=	1.64 * 10 ⁻⁵ n + 2.639	(r = 0.54)	(12)
=	$-2.18 \times 10^{-4} n + 1.745$	(r = -0.18)	(13)
=	$-6.73 \times 10^{-5} n + 2.336$	(r = -0.06)	(14)
=	$-4.99 \times 10^{-4} n + 3.768$	(r = -0.28)	(15)
=	4.82 * 10 ⁻³ n + 45.95	(r = 0.44)	(16)
=	- 0.459 n + 1486.7	(r = -0.62)	(17)
=	0.451 n + 4084.3	(r = 0.48)	(18)
=	- 0.387 n + 3325.4	(r = -0.38)	(19)
	= = = = = =	= $3.22 \times 10^{-5} n + 2.594$ = $1.64 \times 10^{-5} n + 2.639$ = $-2.18 \times 10^{-4} n + 1.745$ = $-6.73 \times 10^{-5} n + 2.336$ = $-4.99 \times 10^{-4} n + 3.768$ = $4.82 \times 10^{-3} n + 45.95$ = $-0.459 n + 1486.7$ = $0.451 n + 4084.3$ = $-0.387 n + 3325.4$	$= 3.22 * 10^{-5} n + 2.594 \qquad (r = 0.55)$ $= 1.64 * 10^{-5} n + 2.639 \qquad (r = 0.54)$ $= -2.18 * 10^{-4} n + 1.745 \qquad (r = -0.18)$ $= -6.73 * 10^{-5} n + 2.336 \qquad (r = -0.06)$ $= -4.99 * 10^{-4} n + 3.768 \qquad (r = -0.28)$ $= 4.82 * 10^{-3} n + 45.95 \qquad (r = 0.44)$ $= -0.459 n + 1486.7 \qquad (r = -0.62)$ $= 0.451 n + 4084.3 \qquad (r = 0.48)$ $= -0.387 n + 3325.4 \qquad (r = -0.38)$

Generally speaking, these rock properties of specimens belonging to this rock group have a comparatively large standard deviation. And, these relations between rock properties and cycle number have not any distinct tendency within 600 cycles. Photo.6 shows the

state of surface exfoliation of this rock group after 600 cycles have passed. And, Photo.7 shows the general view of break in pieces after uniaxial compressive strength has executed at 600 cycles.

Now, the variations of 9 rock properties due to the repeated freezing and thawing in water have been investigated. For defining a degree of weathering, it Photo.6 The state of surface

is difficult to determine what kind of



exfoliation at 600 cycles

property in many engineering properties of rocks is important. The degree of weathering for weak rocks, for example, for mud stone has already been defined by their grain-sizes [9]. For hard or middle hard rocks, the longitudinal elastic wave velocity is one of the most convenient properties to investigate in order to define the degree of weathering. The relations between elastic wave velocity, that is, "Rock Quality Index" and porosity have already studied by Fourmaintraux.D et al



[8],[10]. And there are some examples which Photo.7 Break in the elastic wave velocity has been utilized pieces at 600 cycles for evaluating the degree of weathering of cutting slopes [11], and weathered zone in soft rock [12]. In these cases, it is important to clear the correlations between the elastic wave velocity and the other engineering properties of rocks.[13]

5. Conclusions

Freezing and thawing tests in water have been executed for andesite by use of thermal cycle test apparatus. The temperature of center core of rock sample is controlled to repeat between -30° C and +50°C. The correlations between several rock properties and number of repeated freezing and thawing cycles are experimentally determined. As the results, both dry and saturated densities decrease gradually with cycle number. Especially, the relation for dry density has high coefficient of correlation. It means that the dry density expresses more delicately the expansion or microfracture of pore structure due to repeated freezing and thawing than the saturated density. And absorption, saturated and dry porosities increase gradually with cycle number. Shore hardness measured on the surface of rock specimen decreases gradually with cycle number due to surface exfoliation. And, the uniaxial compressive strength decreases remarkably with cycle number. Both longitudinal elastic wave velocity at saturated state and at dry state decrease gradually with cycle number. In these experiments, the longitudinal elastic wave velocity at saturated state should be more higher evaluated for these weathering tests than the other rock properties, because the relation between the velocity and number of cycle has a very high coefficient of correlations.

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References

- [1] C.D.Ollier : Weathering, Oliver and Boyd.Edinburgh, 1969, pp.6 -36.
- [2] Fukui Prefecture : An explanatory note of geological map of Fukui prefecture, 1969. (In Japanese)
- [3] G.Everling : Richtlinien für Festigkeitsuntersuchungen an Gesteinen, Entwurf 1963 des Internationalen Büros für Gebirgsmechanik, Bergbau Archiv, 25(1964), S.85-91.
- [4] "Engineering properties of rock materials and its application to design and construction"published from the Japanese Soceity of Soil mechanics and Foundation Engineering, 1974, pp.153-242. (In Japanese)
- [5] S.Kunori, M.Abe and T.Shimazaki : An apparatus for measuring the velocities of ultrasonic waves in rocks, geophysical exploration, Vol.22, No.5 (1969) pp.9-15. (In Japanese)
- [6] "Handbook of industrial measurement" published from the Japanese Soceity of Precision Machinery and the Soceity of Automatic control of Measurement, Corona, 1972, pp.746-750. (In Japanese)
- [7] K.Tanuma, M.Fukuda : A preliminary study about some properties of weathered rocks in cold regions, Journal of geography, 80, 5 (1971) pp.18-26. (In Japanese)
- [8] Fourmaintraux.D : Quantification des discontinuités des roches et des massifs rocheux, Rock Mechanics, Vol.7 No.2 (1975) pp.83 -100.
- [9] T.Yamano : Applications of weak rocks and rock materials used for filldam, Soil mechanics and Foundation Engineering, Vol.21 No.3, 181 (1973) pp.25-31. (In Japanese) [10] S.Hata, T.Muro and Y.Kaneko : Estimation of life of rippertip
- due to wear, Proc. of the JSCE, No.268, 1977.
- [11] Y.Nakazawa, S.Okuzono and T.Shimada : Study on weathering characteristics of cut slopes, Proc. of the 9th Symposium on Rock Mechanics, 1975, pp.111-115. (In Japanese)
- [12] Y.Tanaka : Characteristics of Weathering zone in soft rock, Memoirs of the Faculty of Eng., Toyo University, No.11, 1975, pp.73-82. (In Japanese)
- [13] T.Muro : Correlations between several properties of rock specimens, Memoirs of the Faculty of Eng., Fukui University, Vol.24 No.1 (1976) pp.1-18.
- [14] T.Muro, T.Irie and M.Nagai : Weathering tests of rock materials due to freezing and thawing, Proc. of 32th Annual Meeting of JSCE, Vol.3, 1977, pp.362-363. (In Japanese)

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