



# Effect of canopy thickness and canopy saturation on the amount and kinetic energy of throughfall: an experimental approach

著者	Nanko Kazuki, Onda Yuichi, Ito Akane, Moriwaki Hiromu
journal or publication title	Geophysical research letters
volume	35
number	5
page range	L05401
year	2008-03-01
権利	Copyright 2008 by the American Geophysical Union.
URL	<a href="http://hdl.handle.net/2241/98485">http://hdl.handle.net/2241/98485</a>

doi: 10.1029/2007GL033010

1 **Title**

2 Effect of canopy thickness and canopy saturation on the amount and kinetic energy of  
3 throughfall: an experimental approach

4 **Authors' names and Affiliations**

5 Kazuki Nanko,<sup>1</sup> Yuichi Onda,<sup>1</sup> Akane Ito,<sup>1</sup> Hiromu Moriwaki<sup>2</sup>

6 <sup>1</sup> Graduate School of Life and Environmental Sciences, University of Tsukuba, Japan

7 <sup>2</sup> National Research Institute for Earth Science and Disaster Prevention (NIED), Japan

8 Tel: +81-3-5841-6878; Fax: +81-3-5841-6878

9 E-mail: [knanko@geoenv.tsukuba.ac.jp](mailto:knanko@geoenv.tsukuba.ac.jp); [nanko-kazuki@gi.main.jp](mailto:nanko-kazuki@gi.main.jp)

10 **Abstract**

11 To investigate how canopy thickness and canopy saturation affect the amount and  
12 kinetic energy of throughfall, we conducted indoor experiments using a 9.8-m-tall  
13 transplanted Japanese cypress (*Chamaecyparis obtusa*) and a large-scale rainfall  
14 simulator with spray nozzles at a height of 16 m. The amount of throughfall and  
15 raindrop sizes and velocities were measured at twenty-four points under four canopy  
16 structures generated by staged branch pruning. Decreasing the canopy thickness  
17 resulted in increases of the initial throughfall amount, volume proportion of large  
18 throughfall drops, the number of drops with high velocities, and throughfall kinetic  
19 energy. Compared to a saturated canopy, a canopy undergoing wetting had lower  
20 throughfall amounts and volume proportion of large drops, but higher mean drop  
21 velocity. Canopy thickness affected throughfall generation by affecting the processes  
22 of canopy saturation and drop generation within the canopy.

## 1 **1. Introduction**

2 The impact of a raindrop is an important first step toward soil loss and subsequent  
3 sediment transport [e.g., *Ellison*, 1944]. Many studies have proposed the use of the  
4 kinetic energy of rainfall as an indicator of rainfall erosivity [e.g., *Mihara*, 1951;  
5 *Brandt*, 1990; *Morgan et al.*, 1998; *Kinnel*, 2005]. The forest canopy can modify the  
6 amount of rainfall and kinetic energy that reaches the ground surface by altering two  
7 main factors [*Wainwright et al.*, 1999]: raindrop characteristics and interception.  
8 Throughfall consists of three drop components: free throughfall, drips, and splash  
9 water droplets [*Nanko et al.*, 2006]. Compared to open rainfall, throughfall drops are  
10 larger in size because of the dripping effect [*Chapman*, 1948; *Nanko et al.*, 2004] and  
11 generally have lower velocity because they have a shorter fall height [*Laws*, 1941;  
12 *Wang and Pruppacher*, 1977; *Zhou et al.*, 2002]. However, in mature forests,  
13 throughfall drips may have sufficient fall heights to reach terminal velocity and greater  
14 kinetic energy compared to open rainfall [*Chapman*, 1948; *Mosley*, 1982; *Zhou et al.*,  
15 2002; *Nanko et al.*, 2004]. The Brandt model [1990] is widely used to calculate  
16 throughfall kinetic energy, but *Nanko et al.* [2008] demonstrated that the Brandt model  
17 may overestimate kinetic energy because it does not incorporate the water splash  
18 component of throughfall. The water splash component is generated by rain-splash on  
19 the canopy, and thus various canopy structures should have different effects on the  
20 water splash component.

21 Furthermore, the canopy dissipates the amount of rainfall, resulting in substantial  
22 temporal and spatial variability of interception [*Keim et al.*, 2005; *Staelens et al.*,  
23 2006]. *Levia and Frost* [2006] reviewed the findings of 163 studies and concluded that  
24 the temporal and spatial variability of throughfall was affected by both meteorological  
25 factors (e.g., event magnitude and wind speed) and canopy factors (e.g., interception

1 storage and the leaf area index [LAI]). Studies of rainfall magnitude have indicated  
2 that the temporal stability of spatial throughfall patterns differs between small and  
3 large rainfall events [*Staelens et al.*, 2006] and that throughfall variability decreases  
4 with increases in the rainfall amount [*Bouten et al.*, 1992]. Therefore, the process of  
5 throughfall generation may vary by the degree of canopy saturation. The effects of  
6 canopy factors on the amount of throughfall and characteristics of raindrops can only  
7 be estimated by excluding the effects of varying meteorological factors.

8 In this study, we conducted indoor experiments using a 9.8-m-tall stand of  
9 transplanted Japanese cypress (*Chamaecyparis obtusa*) and a large-scale rainfall  
10 simulator with spray nozzles at a height of 16 m. Soil erosion in unmanaged Japanese  
11 cypress plantations poses a serious problem [e.g., *Akenaga and Shibamoto*, 1933;  
12 *Miura et al.*, 2002; *Fukuyama et al.*, accepted, 2007; *Gomi et al.*, submitted, 2007;  
13 *Mizugaki et al.*, submitted, 2007], and solutions will require estimations of the amount  
14 and kinetic energy of throughfall. Our goal was to evaluate how canopy thickness (an  
15 easy-to-measure canopy factor) affects the amount and kinetic energy of throughfall  
16 under varied canopy saturation by measuring the amount, drop size, and drop velocity  
17 of throughfall.

18

## 19 **2. Experimental Setup and Procedure**

### 20 **2.1. Experimental Facility**

21 All experiments were conducted from September to October 2005 in the  
22 large-scale rainfall simulator at the National Research Institute for Earth Science and  
23 Disaster Prevention (NIED), Japan. This simulator is 49 m wide, 76 m long, and 21 m  
24 high; it sits on a bare-earth floor surface and has retractable sidewalls. Artificial  
25 rainfall is sprayed from 544 nozzles at a height of 16 m, which is sufficient for

1 raindrops to reach the floor at terminal velocity. Four different types of wide-cone  
2 nozzles (D1/8G-4.3W, D1/8G-8W, D1/4G-14W, and D3/8G-20W, Spraying Systems  
3 Co., Wheaton, IL, USA) are capable of producing rainfall rates of 15 to 200 mm h<sup>-1</sup>  
4 with pressures of 0.069 to 0.490 MPa [Maki *et al.*, 2005]. The total area can be divided  
5 into four quarters that can be sprinkled separately.

## 6 **2.2. Applied Rainfall Event**

7 Rainfall was applied with a rate of 39.8 mm h<sup>-1</sup> for 15 min using the D1/4G-14W.  
8 This rainfall had smaller drop sizes and kinetic energy than natural rainfall of the same  
9 rate under field conditions. In the applied rainfall, the median volume diameter,  $D_{50}$ ,  
10 and the maximum drop size were 1.10 and 2.66 mm, respectively. The applied rainfall  
11 took place in the simulator with the doors and windows closed under constant  
12 meteorological conditions of no wind, little canopy vibration, low raindrop impact to  
13 canopy surface, and little evaporation.

## 14 **2.3. Transplanted Tree**

15 One Japanese cypress (*Chamaecyparis obtusa*) was transplanted in the simulator.  
16 This tree was 21 years old, 9.8 m tall, and 22.6 cm in diameter at breast height (DBH).  
17 Four types of canopy structure were created using staged branch pruning to estimate  
18 the effect of canopy thickness. Each canopy had a first branch height of 2, 3, 4, and 5  
19 m, canopy thickness of 7.8, 6.8, 5.8, and 4.8 m and LAI of 11.1, 10.1, 8.0, and 6.3;  
20 these respective canopies are referred to as T1, T2, T3, and T4 (Figure 1).

## 21 **2.4. Data Collection**

22 Thirty-two measuring points were established under the canopy, positioned  
23 radially in eight directions (Figure 1). We placed four points 40, 100, 150, and 200 cm  
24 from the trunk in each direction (Point-40, Point-100, Point-150, and Point-200,  
25 respectively). The amount of rainfall was measured using 32 tipping-bucket rain

1 gauges (tipping at 0.254 mm; RC-10; Davis Instruments Corp., Hayward, CA, USA)  
2 set on a platform located 50 cm above the ground surface to prevent the backsplash of  
3 drops from the surface; the tip time was recorded with 0.5-s accuracy using a data  
4 logger (HOBO Event; Onset Computer Corp., Bourne, MA, USA). The throughfall  
5 rate was simultaneously measured at all the measuring points for each canopy  
6 structure; rainfall for each event was applied more than 2 days after the previous  
7 application.

8 Drop sizes and velocities were measured using four laser drop-sizing (LD) gauges  
9 consisting of a paired laser transmitter and receiver, as described by *Nanko et al.*  
10 [2006]. When a raindrop passes through the laser sheet, the output voltage from the  
11 receiver is reduced in proportion to the intercepted area. Detailed explanations of the  
12 measurement principle, the calculation procedure, and the reliability estimation have  
13 been provided by *Nanko et al.* [2008]. A screen net was affixed to the platform 15 cm  
14 above the ground surface to prevent backsplash of drops from the surface. The LD  
15 gauges were placed on a platform located 30 cm above the ground surface, lower than  
16 the rain gauges. The measurements of throughfall drops at all 32 points under T1, T2,  
17 and T3, respectively, and 24 points except for Point-200 under T4 were achieved by  
18 relocating the LD gauges. Even if the process of throughfall drop generation changed  
19 over time and among rainfall events, a given rainfall event would result in the same  
20 effects on the amount of throughfall and throughfall drops because of the constant  
21 meteorological conditions in the indoor simulator.

22 In this study, we did not use the data collected at Point-200. We mainly use mean  
23 value among 24 points under each canopy structure. Individual measured data at each  
24 point were shown as online auxiliary materials (Dataset 1 – 6).

### 25 **3. Results and Discussion**

### 1 **3.1. Temporal Variation in Throughfall Rates**

2 The throughfall amount and rate differed with canopy thickness (Figure 2, Dataset  
3 1 and 2). At the beginning of an event, the throughfall rate was lower than the applied  
4 rainfall rate because water was used to saturate the canopy. The time lag required to  
5 stabilize the throughfall rate shortened and mean throughfall amount increased as  
6 canopy thickness decreased: 8, 5, 4, and 4 min, and 6.9, 7.9, 8.1 and 10.5 mm for T1,  
7 T2, T3, and T4, respectively. T4 resulted larger mean throughfall amount than the  
8 applied rainfall (=10.0 mm) because two measuring points near the trunk (Point-40-7  
9 and Point-40-8) yielded more than twice the amount of applied rainfall, 24.1 and 23.1  
10 mm, respectively. Previous studies found that concentrated throughfall was often  
11 observed close to a tree trunk [*Ford and Deans, 1978; Robson et al., 1994*]. The mean  
12 throughfall amount without the two points under T4 was 9.3 mm.

13 To estimate the effect of canopy saturation on throughfall generation, two phases  
14 were defined: the initial phase (0–5 min), when the canopy was moving toward  
15 saturation, and the stable phase (10–15 min), when the canopy was saturated.

### 16 **3.2. Drop Size of Throughfall**

17 Throughfall drops were larger in size than the applied raindrops. Drops with  
18 diameters  $> 3$  mm were considered to be generated by dripping because the applied  
19 rainfall consisted of the drops with diameters of less than 2.66 mm.

20 The mean volume of throughfall drops, which is the volume of all drops with  
21 diameter  $> 3$  mm divided by the total throughfall drops, differed with canopy  
22 thickness and canopy saturation (Figure 3). Drops had a lower volume proportion  
23 during the initial phase than during the stable phase for all canopy structures; for  
24 example, the mean proportion for T1 was 24% during the initial phase and 33% during  
25 the stable phase. This indicates that throughfall generated from a saturated canopy was

1 composed of larger drops than throughfall generated from a canopy that was not yet  
2 saturated. In addition, differences in drop volume proportions between the phases  
3 decreased as the canopy thickness decreased. As canopy thickness decreased, drop  
4 volume proportions increased during both phases: from T1 to T4, volume proportions  
5 increased from 24 to 36% during the initial phase and from 33 to 40% during the  
6 stable phase, respectively.

7 The volume proportion varied spatially among the measuring points, particularly  
8 during the initial phase. The coefficients of variation (CVs) were 57, 53, 35, and 44%  
9 during the initial phase and 33, 37, 29, and 32% during the stable phase for T1, T2, T3,  
10 and T4, respectively. While undergoing wetting, large spatial variability was found on  
11 the ease of large drop generation.

### 12 **3.3. Drop Velocity of Throughfall**

13 Few drops came close to reaching terminal velocity (Figure 4). During the stable  
14 phase under all canopy structures, > 90% of the drops had < 95% terminal velocity.  
15 The fall height was insufficient for drops to gain terminal velocity. Raindrops with  
16 diameters > 3 mm must fall at least 12 m to reach terminal velocity [*Wang and*  
17 *Pruppacher, 1977*].

18 The drop velocity distribution differed with canopy thickness. During the stable  
19 phase, the mean drop velocity increased as canopy thickness decreased (Figure 4),  
20 with values of 6.7, 7.0, 7.3, and 7.6 m s<sup>-1</sup> for T1, T2, T3, and T4, respectively. The  
21 number of drops with lower velocities decreased, whereas the number of drops with  
22 higher velocities increased. Slow drops (velocity < 6 m s<sup>-1</sup>), which are expected to be  
23 generated by a canopy height of < 2 m, represented 32, 20, 8, and 4% of the total  
24 drops under T1, T2, T3, and T4, respectively. The falling height of drops generated  
25 from the lowest canopy layers increased as the first branch height increased. In



1 contrast, fast drops (velocity  $> 7.5 \text{ m s}^{-1}$ ), which are expected to be generated by a  
2 canopy height of  $> 5 \text{ m}$ , represented 27, 32, 41, and 53% of the total drops under T1,  
3 T2, T3, and T4, respectively. The estimated mean drop fall height was close to the  
4 height of the first branch, following the assumption set out by Brandt [1990], at 3.8,  
5 4.5, 5.1, and 5.4 m for T1, T2, T3, and T4, respectively.

6 The distribution of drop velocity differed with canopy saturation, particularly  
7 under thicker canopy structures. For all canopy structures, the initial phase yielded  
8 fewer drops than the stable phase, but throughfall in the initial phase was consisted of  
9 faster drops than the stable phase. The differences of the number proportion of the fast  
10 drops between the initial phase and the stable phase were 15, 10, 7, and 5% under T1,  
11 T2, T3, and T4, respectively. During the initial phase, the lower canopy layers were  
12 not saturated and thus the upper canopy layers allowed more drops to be generated  
13 compared to the lower canopy layers.

#### 14 **3.4. Drop Generation and Kinetic Energy of Throughfall**

15 Canopy thickness affected the amount and drop generation of throughfall because  
16 the canopy water storage and re-interception possibilities of the drops changed. A  
17 thinner canopy like T4 requires less water to saturate the canopy and allows for a  
18 higher volume proportion of large drops. In contrast, a thicker canopy like T1 requires  
19 more water for canopy saturation, allowing a lower volume proportion of large drops.  
20 As the canopy thickness increased, the possibility for re-interception by the lower  
21 canopy layers increased, and the drops were splashed into fine droplets via impact  
22 with the foliage of the lower canopy layers. Consequently, the drop volume proportion  
23 decreased as the canopy thickness increased (Figure 3).

24 The experiments revealed that the kinetic energy of throughfall was greater than  
25 that of the applied rainfall. Furthermore, thinner canopies generated greater kinetic

1 energy than thicker canopies and sufficiently saturated canopies generated greater  
2 kinetic energy than canopies undergoing wetting. The applied rainfall yielded a unit  
3 kinetic energy (i.e., the kinetic energy per unit area and unit volume of precipitation)  
4 of  $12.7 \text{ J m}^{-2} \text{ mm}^{-1}$ . The throughfall yielded mean unit kinetic energies of 15.9, 16.5,  
5 17.7, and  $19.6 \text{ J m}^{-2} \text{ mm}^{-1}$  during the initial phases and of 16.6, 17.6, 19.2, and  $20.7 \text{ J}$   
6  $\text{m}^{-2} \text{ mm}^{-1}$  during the stable phases of T1, T2, T3, and T4, respectively. The initial  
7 phases yielded lower unit kinetic energy than the stable phases.

#### 8 **4. Conclusion**

9 This experimental study revealed the effect of canopy thickness and canopy  
10 saturation on the amount and kinetic energy of throughfall. Decreasing the canopy  
11 thickness resulted in increases in the initial throughfall amount, volume proportion of  
12 large drops, the number of drops with higher velocities, and kinetic energy. Compared  
13 to saturated canopies, canopies undergoing wetting had lower throughfall amounts and  
14 volume proportions of large drops, but higher mean drop velocities. Canopy thickness  
15 affected the amount, drop generation, and kinetic energy of throughfall due to the  
16 change in canopy water storage and re-interception possibilities of drops within the  
17 canopy.

#### 18 **Acknowledgements**

19 This study was partially supported by a grant from the Japan Science and  
20 Technology Agency (JST) to the Core Research for Evolutional Science and  
21 Technology (CREST) research project “Field and modeling studies on the effect of  
22 forest devastation on flooding and environmental issues.” We appreciate Prof.  
23 Masakazu Suzuki, the University of Tokyo, for his useful comments. We also thank  
24 the technical staff in charge of the NIED rainfall simulator.

#### 25 **Reference list:**

- 1 Akenaga, H., and T. Shibamoto (1933), Effect of soil elements in cypress forest  
2 plantations in the Owase region, *J. Jpn. For. Soc.*, *15*, 19–26 (in Japanese).
- 3 Bouten, W., T. J. Heimovaara, and A. Tiktak (1992), Spatial patterns of throughfall  
4 and soil-water dynamics in a Douglas-fir stand, *Water Resour. Res.*, *28*,  
5 3227–3233.
- 6 Brandt, C. J. (1990), Simulation of the size distribution and erosivity of raindrops and  
7 throughfall drops, *Earth Surf. Process. Landf.*, *15*, 687–698.
- 8 Chapman, G. (1948), Size of raindrops and their striking force at the soil surface in a  
9 red pine plantation, *Trans. AGU*, *29*, 664–670.
- 10 Ellison, W. D. (1944), Studies of raindrop erosion, *Agric. Engineer.*, *25*, 131–136.
- 11 Ford, E. D., and J. D. Deans (1978), The effect of canopy structure on stemflow,  
12 throughfall and interception loss in a young Sitka spruce plantation, *J. Applied*  
13 *Ecol.*, *15*, 905–917.
- 14 Fukuyama, T., Y. Onda, C. Takenaka, and D. E. Walling (accepted, 2007),  
15 Investigating erosion rates within a Japanese cypress plantation using Cs-137  
16 and Pb-210<sub>ex</sub> measurements, *J. Geophys. Res. – Earth Surface*.
- 17 Gomi, T., C. S. Roy, M. Ueno, S. Miyata, and K. Kosugi (submitted, 2007),  
18 Characteristics of overland flow generation on steep forested hillslopes of  
19 central Japan, *J. Hydrol.*
- 20 Keim, R. F., A. E. Skaugset, and M. Weiler (2005), Temporal persistence of spatial  
21 patterns in throughfall, *J. Hydrol.*, *314*, 263–274.
- 22 Kinnel, P. I. A. (2005), Raindrop-impact-induced erosion processes and prediction: a  
23 review, *Hydrol. Process.*, *19*, 2815-2844.
- 24 Laws, J. O. (1941), Measurements of the fall-velocity of water and raindrops, *Trans.*  
25 *AGU*, *22*, 709–721.

- 1 Levia, D. F., and E. E. Frost (2006), Variation of throughfall volume and solute inputs  
2 in wooded ecosystems, *Prog. Phys. Geogr.*, 30, 605–632.
- 3 Maki, M., H. Moriwaki, T. Sato, M. Schönhuber, and T. Harimaya (2005), Raindrop  
4 size distribution in NIED rain simulator, *Geophys. Bull. Hokkaido Univ. Jpn.*,  
5 68, 31–50 (in Japanese with an English summary).
- 6 Mihara, Y. (1951), Raindrops and soil erosion, *Bull. Nat. Inst. Agri. Sci., A-1*, 1-59 (in  
7 Japanese with an English summary).
- 8 Miura, S., K. Hirai, and T. Yamada (2002), Transport rates of surface materials on  
9 steep forested slopes induced by raindrop splash erosion, *J. For. Res.*, 7,  
10 201–211.
- 11 Mizugaki, S., Y. Onda, T. Fukuyama, S. Koga, H. Asai, and S. Hiramatsu (submitted,  
12 2007), Estimation of suspended sediment sources using  $^{137}\text{Cs}$  and  $^{210}\text{Pb}_{\text{ex}}$  in  
13 unmanaged Japanese cypress plantation watersheds of southern Japan, *Hydrol.*  
14 *Process.*
- 15 Morgan, R. P. C., J. N. Quinton, R. E. Smith, G. Govers, J. W. A. Poesen, G.  
16 Auerswald, G. Chisci, D. Torri, and M. E. Styczen (1998), The European soil  
17 erosion model (EUROSEM): a dynamic approach for predicting sediment  
18 transport from fields and small catchments, *Earth Surf. Process. Landf.*, 23,  
19 527-544.
- 20 Mosley, M. P. (1982), The effect of a New Zealand beech forest canopy on the kinetic  
21 energy of water drops and on surface erosion, *Earth Surf. Process. Landf.*, 7,  
22 103–107.
- 23 Nanko, K., N. Hotta, and M. Suzuki (2004), Assessing raindrop impact energy at the  
24 forest floor in a mature Japanese cypress plantation using continuous  
25 raindrop-sizing instruments, *J. For. Res.*, 9, 157–164.

- 1 Nanko, K., N. Hotta, and M. Suzuki (2006), Evaluating the influence of canopy  
2 species and meteorological factors on throughfall drop size distribution, *J.*  
3 *Hydrol.*, 329, 422–431.
- 4 Nanko, K., S. Mizugaki, and Y. Onda (2008), Estimation of soil splash detachment  
5 rates on the forest floor of an unmanaged Japanese cypress plantation based on  
6 field measurements of throughfall drop sizes and velocities, *Catena*, 72,  
7 328-361.
- 8 Robson, A. J., C. Neal, G. P. Ryland, and M. Harrow (1994), Spatial variations in  
9 throughfall chemistry at the small plot scale, *J. Hydrol.*, 158, 107–122.
- 10 Staelens, J., A. De Schrijver, K. Verheyen, and N. E. C. Verhoest (2006), Spatial  
11 variability and temporal stability of throughfall water under a dominant beech  
12 (*Fagus sylvatica L.*) tree in relationship to canopy cover, *J. Hydrol.*, 330,  
13 651–662.
- 14 Wainwright, J., A. J. Parsons, and A. D. Abrahams (1999), Rainfall energy under  
15 creosotebush, *J. Arid. Environments*, 43, 111–120.
- 16 Wang, P. K., and H. R. Pruppacher (1977), Acceleration to terminal velocity of cloud  
17 and raindrops, *J. Appl. Meteorol.*, 16, 275–280.
- 18 Zhou, G., X. Wei, and J. Yan (2002), Impacts of eucalyptus (*Eucalyptus exserta*)  
19 plantation on sediment yield in Guangdong Province, Southern China - a  
20 kinetic energy approach, *Catena*, 49, 231–251.

21 **Figure Captions:**

22 Figure 1. The experimental canopy structures (T1 and T4) and 32 measuring points  
23 under the canopy.

24 Figure 2. Temporal variation in the mean rainfall rate for the applied rainfall and  
25 throughfall under four canopy structures. A broken line of T4 is the mean

1 throughfall rate without two measuring points (Point-40-7 and Point-40-8),  
2 yielded more than twice larger throughfall amount than the applied rainfall.

3 Figure 3. Mean volume proportion of drops with diameters  $> 3$  mm under four canopy  
4 structures during the initial and stable phases.

5 Figure 4. Distribution of drop velocity under four canopy structures during the initial  
6 (upper) and stable (lower) phases. Plots represent mean velocity. Broken lines  
7 indicate the expected velocity of drops with a diameter of 4 mm falling from  
8 heights of 2, 5, and 10 m, and the terminal velocity, respectively, as set out by  
9 *Zhou et al.* [2002].









