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The molecular structure of melts along the carbonatite–kimberlite–basalt compositional joint: CO$_2$ and polymerisation

Yves Moussallam, Pierre Florian, Dario Corradini, Yann Morizet, Nicolas Sator, Vuilleumier, Bertrand Guillot, Giada Iacono-Marziano, Burkhard C. Schmidt, Fabrice Gaillard

Abstract

Transitional melts, intermediate in composition between silicate and carbonate melts, form by low degree partial melting of mantle peridotite and might be the most abundant type of melt in the asthenosphere. Their role in the transport of volatile elements and in metasomatic processes at the planetary scale might be significant yet they have remained largely unstudied. Their molecular structure has remained elusive in part because these melts are difficult to quench to glass. Here we use FTIR, Raman, $^{13}$C and $^{29}$Si NMR spectroscopy together with First Principle Molecular Dynamic (FPMD) simulations to investigate the molecular structure of transitional melts and in particular to assess the effect of CO$_2$ on their structure. We found that carbon in these glasses forms free ionic carbonate groups attracting cations away from their usual ‘depolymerising’ role in breaking up the covalent silicate network. Solution of CO$_2$ in these melts strongly modifies their structure resulting in a significant polymerisation of the aluminosilicate network with a decrease in NBO/Si of about 0.2 for every 5 mol% CO$_2$ dissolved. This polymerisation effect is expected to influence the physical and transport properties of transitional melts. An increase in viscosity is expected with increasing CO$_2$ content, potentially leading to melt ponding at certain levels in the mantle such as at the lithosphere–asthenosphere boundary. Conversely an ascending and degassing transitional melt such as a kimberlite would become increasingly fluid during ascent hence potentially accelerate. Carbon-rich transitional melts are effectively composed of two sub-networks: a carbonate and a silicate one leading to peculiar physical and transport properties.

Keywords : CO$_2$; polymerisation; speciation; transitional melt; kimberlite; glass

1. Introduction

Carbon dioxide (CO$_2$) is typically the second-most abundant volatile in terrestrial melts. In common silicate melts, found near the Earth’s surface and in its crust, its concentration is typically greatly inferior to that of water and its influence on the melt physical properties secondary. In the Earth’s upper mantle however, carbon-rich and typically silica-poor melts, generated by low degree partial melting of the mantle are probably widespread (e.g. Wyllie and Huang, 1976, Eggler, 1978, Dalton and Presnall, 1998, Gudfinnsson and Presnall, 2005, Dasgupta and Hirschmann, 2006 and Dasgupta et al., 2013; Massuyeau et al., in press). Very low degree partial melting of carbonated peridotite at ~250 km depth produces carbonatite liquid (e.g. Dasgupta and Hirschmann, 2006). With either increasing temperature, decreasing pressure or with
the addition of CO₂ or H₂O, the liquid produced by peridotite partial melting will become increasingly SiO₂-rich, forming kimberlite, melilitite and eventually basalt (e.g. Wyllie and Huang, 1976; Eggerl, 1978; Wyllie, 1989). These melts – carbonatite, kimberlite and melilitite – transitional between pure carbonatite and silicate liquids (∼10 to ∼40 wt% SiO₂), can contain up to several tens of weight percent of CO₂ (Brey and Green, 1976; Moussallam et al., 2014). Understanding the effect of CO₂ on the structure of these melts/glasses is therefore a cornerstone to predict their physical characteristics and transport properties.

The effect of CO₂ on melt polymerisation has remained unclear. Eggerl (1978) suggested that the dissolution of CO₂ as carbonate ion (CO₂⁻) should result in melt polymerisation via the reaction:

$$CO₂ + 2Q^n(M^{n+}) = CO₂^{2-}(M^{n+}) + 2Q^{n+1}$$

Where Qⁿ denotes a silica tetrahedron linked by bridging O atoms to n adjacent tetrahedra (n = 0 correspond to an isolated SiO₄ tetrahedron, n = 4 correspond to a fully connected tetrahedron with four bridging oxygens), Mⁿ⁺ denotes a metal cation in network modifying or charge balancing role. In Eq. (1), the increase in polymerisation is denoted by the building of Si-O-Si bonds.

This depiction has gained popularity, being referred to as a largely admitted concept in review articles (e.g. Mysen, 2013, 2012) and books (e.g. Frost and Frost, 2013; Mysen and Richet, 2005). Solid evidence of the effect of CO₂ dissolution on the aluminosilicate network structure however has been lacking. Mysen and Virgo (1980a, 1980b) found that CO₂ depolymerises albite and anorthite melts while polymerising diopside and NaCaAlSi₂O₇ melts. Konschak (2008) investigated the albite – diopside joint and found no effect of CO₂ on polymerisation from Ab₅₀Di₅₀ to Ab₇₅Di₂₅ but found that a slight depolymerisation is associated with the addition of CO₂ to Ab₉₀Di₁₀. As noted by Ni and Keplekker (2013) these evidence are all based on very subtle differences in Raman spectra, which interpretation can be controversial. All these studies have focused on the incorporation mechanisms of CO₂ and its impacts on the molecular structure of silicate melts. Here we look at CO₂ in transitional melts, typically produced by very low-degree partial mantle melting and with silica content being about half of that of basals.

The principal question we target in this study is: Does carbon dioxide influence the degree of polymerisation of transitional melts? We present results from the first Infrared, Raman and Nuclear Magnetic Resonance (NMR) spectroscopy investigation of transitional glasses and compare them with results from Ab-initio First Principle Molecular Dynamic simulations applied to melt conditions. We explore a compositional range from ~44 to 23 mol% SiO₂ (on a volatile-free basis) with 0 to 26 mol% CO₂. For all compositions investigated, we show that (1) CO₂ is present in the glass and in the melt dominantly as carbonate ion (2) the degree of polymerisation of the glass/melt increases with increasing CO₂ content. We conclude that the physical properties of transitional melts such as viscosity, electrical conductivity and sound velocity are expected to be greatly affected by their CO₂ content with implications regarding melt mobility in the upper mantle.

We note that polymerisation, in this manuscript is strictly defined as the process by which oxygen atoms are shared between silica (or alumina) tetrahedra and is expressed as the ratio of non-bonding oxygens to tetrahedral cations (NBO/T) or as the ratio of non-bonding oxygens per silicon (NBO/Si). We further note that the NBO/T ratio is only a calculated 'statistical average' and various combinations of O species can give the same NBO/T. The NBO/T concept fails past NBO/T = 4 but Q units at greater values can still join up in isolated 'polymerised units'.

2. Methods

2.1. Experimental methodology

2.1.1. Starting material

Starting materials were produced by mixing powders from a natural lamproite from Torre Alfina, Italy (Pecceirillo et al., 1988) with various amounts of synthetic powders of pure oxides (SiO₂, Al₂O₃, MgO, CaCO₃, Na₂CO₃, K₂CO₃) and natural dolomite. The strategy was to produce a series of progressively more silica and alumina poor compositions additionally low in iron in order for the final product to be analysable by NMR spectroscopy. The source of CO₂ in experiments was mainly CaCO₃. The Torre Alfina rock, was fused twice in air at 1400°C and quenched to glass in order to ensure homogeneity and remove any volatile present. The composition of all mixtures used as starting materials is reported in Table 1.

<table>
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<tr>
<th>Name</th>
<th>SiO₂</th>
<th>TiO₂</th>
<th>Al₂O₃</th>
<th>FeO</th>
<th>MgO</th>
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<td>5.4</td>
<td>1.0</td>
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<td>3.9</td>
<td>0.9</td>
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<td>0.3</td>
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<tr>
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<td>23.1</td>
<td>0.3</td>
<td>5.4</td>
<td>1.5</td>
<td>8.3</td>
<td>30.2</td>
<td>0.4</td>
<td>3.2</td>
<td>0.1</td>
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<td>1.6</td>
<td>8.3</td>
<td>29.7</td>
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<td>0.6</td>
<td>5.2</td>
<td>0.3</td>
<td>15.2</td>
<td>100</td>
</tr>
</tbody>
</table>

2.1.2. HP-HT experiments

Experiments were performed in the pressure range 0.1 to 1500 MPa. High temperature experiments at 0.1 MPa were performed in a high-temperature furnace. The sample was contained in a Pt crucible and heated to 1500°C for 30 min and then quenched to glass by complete immersion in cold water.

Experiments at high pressure were performed by Moussallam et al. (2014) in the pressure range 50 to 350 MPa at relatively constant temperature (1225 to 1270°C) with no added water and under oxidised conditions (log fO₂ = FMQ + 3). Experiments were performed in internally heated pressure vessels at the ISTO-CNRS laboratory in Orléans which can reach pressures of up to 400 MPa (±3 MPa) under controlled temperature up to 1300°C (±5°C) (see supplementary information).

Experiments at 1500 MPa and 1350°C were performed in a piston-cylinder apparatus with in a 3/4 inch (1.9 cm) assemblies. Experimental charges consisted of natural anhydrous sample powder (30 mg) loaded in sealed gold–palladium (Au80Pd20) capsules (1 cm in length, 2.5 mm inner diameter and 2.9 mm outer diameter). The capsules were introduced in a talc–pyrex–graphite furnace assembly and surrounded by MgO. A B-type thermocouple was located at ~1 mm atop of the capsule and the run temperature should be considered as a minimum value. Uncertainties are of 10% in relative pressure and of 15°C for temperature. We used a modified perforated anvil through which cold water was circulated in order to maximise quenching rate (>200°C s⁻¹).

A pure carbonate glass of composition K₂Mg(CO₃)₂ was synthesised in internally heated pressure vessel at 803°C and 1083°C.
bar under relatively oxidised conditions (log fO2 = FMQ + 3) with no added water. This composition was shown previously to be quenchable to carbonate glass without crystallisation (Ragone et al., 1966; Genge et al., 1995). The sample was enriched in 13C by using K213CO3 (99% 13C) as starting material. Drop quench at the end of the experiment resulted in pure carbonate glass as indicated by the lack of birefringence in polarisation microscopy.

2.2. Analytical techniques

2.2.1. Microbeam analyses

All experimental products were examined by optical microscope and scanning electron microscope (SEM) to check for the presence of quench crystals. Electron microprobe analyses (EMPA) were performed on a Cameca SXFive at the ISTO-CNRS laboratory in Orléans. Glasses were analysed using an accelerating voltage of 15 kV, a beam current of 6 nA and a defocused beam of 10 μm with a count time of 10 s. Na was analysed first in order to minimise its loss during analyses. No K loss with time was observed at our operating conditions.

2.2.2. Infrared spectroscopy

Infrared spectra were collected using a Nicolet 6700 FTIR spectrometer attached to a Continuum microscope. We used a Globar internal IR source with a KBr beam splitter and a liquid nitrogen cooled MCT/A detector. The spectral resolution was set to 4 cm⁻¹, and spectra were accumulated for 128 scans. Background spectra were acquired by accumulating 256 scans and used to correct for background. Spectra were obtained directly on small glass fragments (unpolished) in transmission mode.

2.2.3. Micro-Raman spectroscopy

An Innova 300-5W Argon ion laser (Coherent®) operating at 514 nm was used as the exciting source to produce Raman scattering. Spectra were collected by a Jobin–Yvon Labram spectrometer (focal distance = 300 mm) equipped with a grating with 2400 grooves/mm and a CCD detector. The spectral frequency position was calibrated using the emission lines of Ne- and Hg-lamps with an accuracy within ±1 cm⁻¹. Analyses were performed in confocal mode (hole = 500 μm, slit = 200 μm) and using a ×50 Olympus objective resulting in an analysed volume of a few μm³. Spectra were acquired in the 700–1300 cm⁻¹ spectral region which corresponds to the vibrational region for aluminosilicate framework symmetric stretch (ν1 for Q14 species; e.g. Mysen et al., 1982) and also the symmetric stretch for CO2− molecular groups (ν1 for CO2−) at around 1080 cm⁻¹. The acquisition time was of 10 × 60 s and the focus depth was optimised in order to obtain the highest Raman signal. At least three spectra were acquired for each sample.

2.2.4. NMR spectroscopy

All NMR experiments were performed on a 9.4 T Avance Bruker Spectrometer operating at 79.5 MHz (29Si) and 100.6 MHz (13C). We used a 4 mm diameter rotor spinning at 14 kHz and applied an Hahn–Echo acquisition with a radio-frequency field of 50 kHz (29Si) and 35 kHz (13C) and an inter-pulse delay of 1 rotor period. The relaxation time are found to be around 75 ms for both nuclei because of the presence of iron (1 to 4 wt% on volatile free basis) in the sample. The recycle delays were set to 250 ms, while accumulating between 300000 and 900000 scans, depending on the amount of sample available. A Teflon spacer was used to position the sample in the middle of the rotor when using small amounts.

2.3. First principles MD simulations

First-Principles Molecular Dynamics (FPMD) simulations were performed on two systems. The first one is a simplified TA10 system (32 wt% SiO2; expressed on a volatile-free basis; cf. Table 1) composed of SiO2 (x = 0.30), Al2O3 (x = 0.07), MgO (x = 0.14) CaO (x = 0.44) and K2O (x = 0.05) (were x refers to the mole fraction of each component). A simulation box with a side length L = 14.7 Å was used. This box consisted of 238 atoms (135 O, 31 Si, 7 Al, 15 Mg, 45 Ca and 5 K). This gave a density of 2.92 g/cm³. The system was run at T = 1773 K. The second system is a simplified TA10 composition containing 13.8 wt% of CO2, which translates in terms of mole fractions to CO2 (x = 0.18), SiO2 (x = 0.25), Al2O3 (x = 0.05), MgO (x = 0.12), CaO (x = 0.36) and K2O (x = 0.04). In this case, the simulation box contained 304 atoms (179 O, 31 Si, 7 Al, 15 Mg, 45 Ca, 5 K and 22 C) and its side length measures L = 15.7 Å, corresponding to a density of 2.83 g/cm³. For this system, we fixed the temperature at T = 1498 K. The adopted temperature conditions represent melt conditions almost identical to glass synthesis conditions. The experimental glass structure however is that preserved at the glass transition temperature, which is lower than the synthesis temperature, while the structure observed by FPMD is that of the melt. The simulations required the use of about 215 000 single CPU hours, using the high performance of the IDRIS supercomputer. Additional simulation details are given in the supplementary information.

Theoretical NMR spectra were obtained from calculations performed on MD boxes, details are given in the supplementary information.

3. Results

Five starting compositions were saturated with CO2, at pressures from 0.1 to 1500 MPa. This allowed us to create samples of similar compositions with variable CO2 content. All experiment are reported in Table 2. Analytical data from pressure between 100 to 350 MPa are from the experiments reported in Moussallam et al. (2014). The FPMD ab-initio molecular dynamic simulations were performed on a single composition (with 32 wt% SiO2 on a volatile free basis) equilibrated at a single pressure, with and without CO2.

3.1. Infrared spectroscopy

Infrared spectra for CO2-bearing glasses ranging in composition from 24 to 44 wt% SiO2 (expressed on a volatile-free basis) are reported in Fig. 1 together with an infrared spectrum obtained on a pure carbonate glass of composition 50:50 mol% MgCO3–K2CO3. In all compositions investigated CO2 in the glass is in the form of carbonate, in contrast to recent findings in water-saturated carbonate melts (Foustoukos and Mysen, 2015). The absence of absorption peak at ~2350 cm⁻¹ in all spectra shows the absence of molecular CO2 in the quenched glass, although it may have been present in the melt (see section 3.4). The carbonate v3 doublet has a similar position across composition (mid-point at ~1460 cm⁻¹) with a splitting of Δv3 ~ 70 cm⁻¹. In detail, small variations between peak position occur from the most silica-rich (TA12; 44 wt% SiO2 on a volatile free basis) to the most silica-poor glass (TA6; 24 wt% SiO2 on a dry basis) with the mid-point shifting from 1465 to 1455 cm⁻¹ and the Δv3 split shifting from 80 to 60 cm⁻¹. The spectra obtained on pure carbonate glass shows a mid-point at 1450 cm⁻¹ and a splitting of Δv3 ~ 60 cm⁻¹.

3.2. Raman spectroscopy

Raman spectra of quenched glasses ranging in composition from 24 to 44 wt% SiO2 (expressed on a volatile-free basis) and containing 0 to 22 wt% CO2 are presented in Fig. 2 in the frequency range 700 to 1200 cm⁻¹. All glasses containing CO2 show a strong peak at ~1080 cm⁻¹ corresponding to the stretching of C=O vibration in CO2− configuration while no evidence
for CO$_{2\text{mol}}$ was observed in the Raman spectra at $\sim$1400 cm$^{-1}$ (Mysen and Virgo, 1980a; Sharma, 1979; Verweij et al., 1977; White, 1974) consistent with FTIR spectra shown in Fig. 1. Besides the 1084 cm$^{-1}$ peak, the Raman spectra shown in Fig. 2 can be considered as a complex convolution of individual Gaussian peaks; each corresponding to a given Q$^n$ species (e.g. Mysen, 1990, 2007; Mysen and Cody, 2005; Mysen and Richet, 2005 and references therein). It is clear from Fig. 2 that the addition of CO$_2$ strongly modifies the shape of the Raman spectra in all investigated compositions. Deconvolution of the high-frequency envelope (800–1200 cm$^{-1}$) in which 1st-order Raman scattering of (Si, Al)–O stretching vibration occurs is shown in Fig. 3. Deconvolution of Raman spectra is an inherently interpretative process with non-unique solution, all results are hence treated in a purely qualitative sense. Measured Raman spectra are simulated by four individual Gaussian peaks for volatile-free glasses and with a fifth peak at $\sim$1080 cm$^{-1}$ for the v$_1$ CO$_2$ vibrational signature in CO$_2$-bearing glasses. Assigning (Si, Al)–O stretching frequency to particular Q-species is not straightforward as the frequency

for of each band varies as a function of melt chemistry. There is a general consensus however that higher-order (more polymerised) Q-species show (Si, Al)–O stretching at higher frequency than lower-order (less polymerised) Q-species. Here we assign the band near 968 cm$^{-1}$ (marked in orange in Fig. 3) to (Si, Al)–O– stretching vibrations in Q$^3$-species (in accordance with Brawer and White, 1975; Furukawa et al., 1981; Mysen, 2007; Mysen et al., 1982; Mysen and Cody, 2005), the band near 910 cm$^{-1}$ (marked in green in Fig. 3) to (Si, Al)–O– stretch vibrations in Q$^4$-species (in accordance with Mysen, 2007) and the band near 860 cm$^{-1}$ (marked in cyan in Fig. 3) to (Si, Al)–O– stretch vibrations in Q$^5$-species (in accordance with Mysen et al., 1982). The band at $\sim$1025 cm$^{-1}$ is assigned to Si–O$^0$ vibration in any structural unit with bridging oxygen (in accordance with Mysen et al., 1982). In the fits reported in Fig. 3, band positions were fixed for a given composition while bandwidth, and band intensity were treated as independent variables. Minimum $\chi^2$ was used as the principal convergence criterion. It is clear from Figs. 2 and 3 that for all investigated compositions, the addition of CO$_2$ results in a change of the glass structure towards more polymerised configurations.

### 3.3. NMR spectroscopy

$^{13}$C NMR spectra of compositions ranging from 24 to 44 wt% SiO$_2$ (expressed on a volatile-free basis), produced at 350 MPa and containing 0 to 22 wt% CO$_2$ are shown in Fig. 4 together with the $^{13}$C NMR spectra of a pure carbonate glass of composition 50:50 mol% MgCO$_3$–K$_2$CO$_3$. Peak parameters are given in Table S3 in the supplementary information. All spectra show a single symmetric peak at $\sim$168 ppm indicating that CO$_2$ is dissolved as carbonate. No evidence of CO$_{2\text{mol}}$ was observed in the $^{13}$C spectra at 125 ppm.

$^{29}$Si NMR spectra of the same glasses are presented in Fig. 5 (Note that the CO$_2$-free TA9 sample is slightly affected by quench crystal nuclei, slightly distorting the NMR spectra). The solution of $\sim$2 wt% CO$_2$ to the most silica-rich composition (44 wt% SiO$_2$) on a volatile-free basis) and from 17 to 22 wt% CO$_2$ in the most silica-poor composition (24 wt% SiO$_2$) on a volatile-free basis) results in no discernible change to the $^{29}$Si chemical shift in glasses. For compositions with 38, 32 and 31 wt% SiO$_2$ (expressed on a volatile-free basis), the solution of 7.5, 13.8 and 14.9 wt% CO$_2$ (corresponding to 10, 18 and 19 mol% CO$_2$ respectively) results in a significant shift of the peak maximum of the $^{29}$Si spectra to 3 to 5 ppm, respectively.
5 ppm towards more negative values and an even more drastic shift of the centre of mass from 4 to 7 ppm towards more negative values. The asymmetry of these peaks also changes becoming skewed towards more negative values of \(^{29}\text{Si}\) chemical shift with increasing CO\(_2\) content. The \(^{29}\text{Si}\) spectra shown in Fig. 5 can be considered a complex convolution of individual peaks; each corresponding to a given \(Q^n(p\text{Al})\) species (where \(n\) denotes as usually the number of bridging oxygen on the \(Q\) unit and \(p\) the number of connected aluminium atoms (hence a \(Q^n\) species with \(p\) Si–O–Al bridges and \((n-p)\) Si–O–Si ones) (e.g. Hiet et al., 2009; Lee and Stebbins, 1999; Lippmaa et al., 1981, 1980; Maekawa et al., 1991; Magi et al., 1984; Engelhardt and Michel, 1987; Morizet et al., 2015, 2014a). The resonance frequency of \(Q^n(p\text{Al})\)-species however varies as a function both \(n\) and \(p\) but in opposite direction resulting in ambiguities in the band assignment in \(^{29}\text{Si}\) NMR spectra across composition. For this reason we restrain from using NMR spectra to quantify the observed chemical shift in terms of changes in \(Q^n\)-species.

3.4. First principle MD simulations

Two first principles molecular dynamic (FPMD) simulations were performed on a single composition (32 wt% SiO\(_2\) expressed on a volatile-free basis). Simulations were performed on the base composition with 13.8 wt% and 0 wt% CO\(_2\) respectively. Fig. 6 shows a snapshot of the last configuration extracted from both FPMD simulations (Figs. 6A and 6B). FPMD simulation on the CO\(_2\)-bearing composition shows that CO\(_2\) is dissolved dominantly as carbonate species with very few CO\(_2\)\(^{\text{mol}}\) (Fig. 6F). The silicate structure from both simulations is shown in Fig. 6C and Fig. 6D. The silicate network from both simulations appears clearly distinct with the CO\(_2\)-bearing simulation being more polymerised than the CO\(_2\)-free simulation. The average number of NBO/Si decreases from 3.4 in the CO\(_2\)-free simulation to 2.8 in the simulation with 13.8 wt% CO\(_2\). When considering also NBO-carb the number of NBO decreases further to 2.7 (Fig. 6E).

Simulations of the \(^{29}\text{Si}\) NMR spectra were produced using results from FPMD simulations (see supplementary information). Fig. 7 compares the simulated spectra with the spectra obtained on the corresponding experimentally quenched glasses. The agreement between the simulated FPMD-based NMR spectra and the spectra obtained on the corresponding experimental glasses is excellent (considering the limited number of atoms – and hence statistics – on which the NMR calculations could be performed), confirming that the molecular structure resulting from FPMD simulations reproduces the structure of the silicate network of the experimental system and that the experimental glasses preserved the structure present in the melt conditions. The agreement between FPMD-simulated and measured NMR spectra is best for the composition with 13.8 wt% CO\(_2\) as four FPMD simulations were used to calculate the NMR spectra compared to two FPMD simulations used for the composition with 0 wt% CO\(_2\). The same procedure has been applied to the \(^{13}\text{C}\) spectra, and as can also be seen from Fig. 7 the result is satisfactory. The asymmetry displayed by the calculated spectra indicates the small but significant abundance of CO\(_3\) groups linked to the aluminosilicate network with a bridging oxygen (NBO-carb.; Brooker et al., 2001b) in the FPMD simulation and which are not apparent (or not preserved during quenching) in the experimental glass.

4. Discussion

4.1. Carbon dioxide speciation and environment in transitional melts

4.1.1. Carbon dioxide speciation

Infrared spectra show that carbon dioxide in the investigated compositions is present in the glass (quenched from melt) as
carbonate. Raman spectra confirm that carbon dioxide is only present as CO$_2^-$ in all investigated glasses. $^{13}$C NMR spectra also show that CO$_2$ is dissolved in all investigated glasses as carbonate group. We found no evidence for the presence of CO$_2^{\text{mol}}$ at 125 ppm.

In mafic silicate glass, carbon dioxide typically dissolves as carbonate ion (CO$_3^{2-}$) (e.g. Fine and Stolper, 1986). An increasing amount of evidence however suggest that carbon dioxide may actually not entirely dissolve as CO$_3^{2-}$ in the more common silicate melt compositions, including basalts. Annealing ex-
Fig. 4. $^{13}$C NMR spectra obtained on quenched glasses ranging in composition from 24 to 44 wt% SiO$_2$ (expressed on a volatile-free basis) and containing 0 to 22 wt% CO$_2$. The NMR spectra of a pure carbonate glass of composition 50:50 melt MgCO$_3$–K$_2$CO$_3$ is also shown for reference. All spectra can be described as a single Gaussian peak, the centre of which is reported (in ppm) beside each spectra. At the resolution considered there is no difference in the $^{13}$C chemical shift across investigated compositions. The peak at 110 ppm apparent in some spectra is an artefact resulting from the use of Teflon as a sample holder. Peak parameters are given in Table S3.

Experiments – performed below the glass transition temperature ($T_g$) – on CO$_2$–bearing jadeite (Morizet et al., 2001), albite and dacitic (Nowak et al., 2003) glasses show that the equilibrium between molecular and carbonate species shifts towards molecular CO$_2$ with increasing temperature. Molecular dynamics simulations by Guillot and Sator (2011) and Morizet et al. (2015) confirmed that the mole fraction of molecular CO$_2$ in basaltic melt increases with temperature, while carbonate ions are the only species observed in basaltic glasses, a significant amount of molecular CO$_2$ is present in the melt. A recent in-situ infrared spectroscopy study by Konschak and Keppler (2014) studied the speciation of carbon dioxide in diamond anvil cell experiments up to 1000 °C, confirming that the CO$_2$ + O$_2$ = CO$_2^-$ equilibrium in the melt shift towards CO$_2$ with increasing temperature in dacite and phonolite melts (at this time no similar experiments on basaltic composition have been performed).

Given the silica-poor composition investigated here, the speciation of carbon dioxide as CO$_2^-$ in the glass is fully consistent with previous studies (e.g. Brooker et al., 2001b; Ni and Keppler, 2013). The speciation of carbon dioxide in the melt could be different to that preserved in the glass (which preserves the structure at the glass transition temperature) and was investigated by FPMD simulations, showing that ~90% of carbon in the investigated composition is stable as carbonate (Fig. 6F). This is consistent with previous studies suggesting that for the low-silica composition investigated most of the CO$_2$ should also be dissolved as CO$_2^-$ in the melt (e.g. Konschak and Keppler, 2014; Ni and Keppler, 2013; Vuilleumier et al., 2015).

4.1.2. Carbon dioxide environment

The mid-point of the carbonate $v_3$ doublet at 1460 cm$^{-1}$ and split of $\Delta v_3 = 60$–80 cm$^{-1}$ in the infrared spectra is close to that found in most Ca-bearing silicate melts (Brooker et al., 2001b). Blank and Brooker (1994) and Brooker et al. (2001b) have suggested that NBO-carbonate species (in which a carbonate group is linked to a Si tetrahedra via a non-bridging oxygen) can produce $v_3$ splitting of the order $\Delta v_3 = 70$–100 cm$^{-1}$. We show however that $\Delta v_3$ split of 60 cm$^{-1}$ can be observed in pure carbonate glass were carbonate group are not linked to any silica tetrahedra, and consistently with Sharma and Simons (1980) and White (1974). This illustrates, that the $v_3$ splitting is only function of the degree of distortion of the carbonate group, its absolute value therefore does not necessarily imply chemical bonding to a specific ion. Another possible interpretation is that the observed $v_3$ ‘doublet’ might represent two distinct symmetrical site with different surrounding cations (in this case Mg or K) rather than the lifting of $v_3$ degeneracy by distortion. Indeed Geige et al. (1995) suggested than in pure carbonate melts, CO$_3^{2-}$ groups interact strongly with Ca and Mg cations and less strongly with alcalies. $^{13}$C NMR spectra show that in all investigated compositions a single symmetric peak centred at ~168 ppm is present, suggesting a single environment for CO$_3^{2-}$ or that NMR is insensitive to any difference. This peak position was found in CO$_2$-bearing nepheline, albite and sodalmellite melt by Kohn et al. (1991), in nepheline, jadeite and albite melt by Brooker et al. (1999), in haplo-phonolites by Morizet et al. (2002) in amorphous calcium carbonate by Nebel et al. (2008) in nephelinite by Morizet et al. (2014a, 2014b) and diopside and Ca-mellilitite by Morizet et al. (2015) who attributed it to NBO-carbonate species (where a CO$_3^{2-}$ molecule is connected to a NBO and charge compensated by a surrounding cation). The presence of the same peak position in pure carbonate glass (Fig. 4), close to that of common carbonate minerals (e.g. Papenguth et al., 1989), however argues against such definite designation. The $^{13}$C chemical shift we identify in transitional glasses is indistinguishable from that found in
Fig. 6. Snapshot of the last configuration extracted from first principles molecular dynamic (FPMD) simulations on a TA10 composition (32 wt% SiO₂; expressed on a volatile-free basis) equilibrated with 0 wt% CO₂ (A) and 13.8 wt% CO₂ (B). Carbon atoms are in blue, oxygen atoms are in red, silicon atoms are in yellow, aluminum atom are in light brown, Ca atoms are in green (large balls), Mg atoms are in light green (smaller balls), K atoms are in purple. The atomic radii drawn are proportional to the van der Waals radii. Snapshot of the silicate network solely is shown for both simulations with 0 wt% CO₂ (C) and 13.8 wt% CO₂ (D) with silica tetrahedra coloured based on the number of shared oxygen as follow: blue for no shared oxygen (i.e. Q^0-species), green for one oxygen shared (i.e. Q^1-species), orange for two oxygen shared (i.e. Q^2-species), and purple for three oxygen shared (i.e. Q^3-species). E: Plot of the number of non-bridging oxygen divided by the number of Si as a function of simulation time for simulations performed with 0 and 13.8 wt% CO₂. F: Probability (calculated over the 20 ps of simulation time) of having a fraction of C atoms involved in CO\textsuperscript{mol} (black) or CO\textsuperscript{3−} (red). (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

a pure carbonate glass suggesting that CO\textsuperscript{3−} groups are located in a similar environment (i.e. as free carbonate). Moreover, the barycentre of the 13C NMR component of NBO-carbonate species found in the MD simulation resonate at −1.4 ppm for CO\textsuperscript{3−}-Al and −4.3 ppm for CO\textsuperscript{3−}-Si with respect to that of isolated CO\textsubscript{3} units (see Fig. 7). In other words if the presence of iron broadens the NMR line to the point where we cannot expect to resolve those species, those shifts are big enough that a significant amount of NBO-CO\textsubscript{3} units should produce an asymmetry in the 13C NMR line on its right tail, in contradiction with the fully symmetric observed line. Bridging carbonate species (Si–carb–Si), also referred to as “network carbonate” are particularly unlikely to be present in the observed glass as such highly distorted species would have been clearly identifiable in the infrared (characterised by large Δν\textsubscript{3} split
Fig. 2. Upper two panels: $^{29}$Si NMR spectra obtained on quenched glasses with 32 wt% SiO$_2$ (expressed on a volatile-free basis) and containing 0 and 13.8 wt% CO$_2$ compared to simulated $^{29}$Si NMR spectra obtained from FPMD simulations on the same composition. The deconvolution of the simulated $^{29}$Si NMR spectra is shown in terms of simulated contribution from $Q^2$-species. Note that simulated $^{29}$Si NMR spectra are based on four and two FPMD simulations for the compositions with 13.8 and 0 wt% CO$_2$ respectively explaining the lesser spectrum quality for the composition with 0 wt% CO$_2$. Lower panel: $^{13}$C NMR spectra obtained on quenched glasses with 32 wt% SiO$_2$ (expressed on a volatile-free basis) and 13.8 wt% CO$_2$ compared to simulated $^{13}$C NMR spectra obtained from FPMD simulations. Note that the discrepancy between the two spectra lies in the asymmetry displayed by the calculated spectra which is due to the non-negligible presence of CO$_2$ groups linked to the alumino-silicate network (NBO-carb) in the FPMD simulation. NBO-carb groups are not apparent (or not preserved during quenching) in the experimental glass. The signal of the experimental glass spectrum is broader than the simulated spectrum due to the presence of Fe.

$\sim$150–300 cm$^{-1}$; Brooker et al., 1999, 2001b) and $^{13}$C NMR spectra (characterised by chemical shift near 155–165 ppm; Morizet et al., 2002, 2014a, 2015; Brooker et al., 1999).

FPMD simulations of the melt predict that most dissolved carbon atoms form free CO$_2$$_2$– groups linked to alkaline earth cations (i.e. are in the same environment than in a pure carbonate; not linked to the silicate network). Few NBO-carbonate are predicted to be stable and no network carbonate species was found to be stable. These findings are consistent with previous molecular dynamic simulation performed by Vuilleumier et al. (2015) on Kimberlite melt and with our infrared and $^{13}$C NMR spectroscopic results and suggest that the environment of CO$_2$$_2$ in transitional glasses and melts is dominantly as free carbonates. Those results are also consistent with the current knowledge suggesting that network carbonate are stable in more polymerised (more silica rich) compositions.

4.2. Effect of carbon dioxide on transitional melts polymerisation

It is clear from both Raman and NMR investigations that the addition of CO$_2$ has a drastic effect on the glass structure, systematically modifying it to more polymerised configuration by sequence metal cations from their usual depolymerising role (Figs. 3 and 5). FPMD simulations and the resulting calculation of NMR parameters also are consistent with CO$_2$ dissolution modifications of the network structure to a more polymerised configuration (Figs. 6 and 7). This effect appears to be minimal to insignificant for the addition of up to $\sim$3 wt% CO$_2$ in the most silica-rich composition investigated and for changes from 17.2 to 21.8 wt% CO$_2$ in the most silica-poor composition investigated. Using the relative area defined by each band used in the Raman spectra deconvolution, the relative abundance of each $Q$-species as a function of CO$_2$ content can be expressed. We stress that such estimate of $Q$-species proportion from Raman spectra can only be qualitative.

We also stress that the change seen in both Raman and $^{29}$Si NMR spectra could be interpreted in different ways from the increase in glass polymerisation corresponding to topological changes (Massiot et al., 2013). For instance, both Raman and $^{29}$Si NMR spectra could be affected by the electrical charge distribution due to the nearby presence of CO$_2$$_2$ group therefore distorting the geometry of the $Q^2$ cluster (Massiot et al., 2013). Those geometrical changes (change in bond lengths and bond angles) have been recognised to affect in a significant way the
the measured signal. We note finally that the null-hypothesis; that dissolution of CO$_2$ does not affect the degree of polymerisation of the melt and resulting glass would predict that both Raman and NMR spectra should result in line shapes more consistent with an increasingly depolymerised structure with the addition of CO$_2$ (due to dilution of the aluminosilicate component with increasing amount of CO$_2$). The fact that we observe the exact opposite (Raman and NMR spectra shifting to line shape more consistent with an increasingly polymerised structure with the solution of CO$_2$) argues for a polymerising role of CO$_2$ on the molecular structure of the compositions investigated. Fig. 8 shows the evolution of the apparent degree of polymerisation of the glass (expressed as NBO/Si) as a function of mol% CO$_2$ as calculated from FPMD simulation. The extent of the polymerisation of glasses increases with increasing amount of CO$_2$ (the amount of non-bridging oxygen per silica decreases). The relative change with increasing CO$_2$ content is apparent in the Raman spectra (Fig. 3), NMR spectra (Fig. 5) and FPMD results (Fig. 6 and Fig. 8). The relationship between NBO/Si and CO$_2$ content calculated from FPMD simulation (see also Fig. 6E) is of $\sim$−0.2 for every 5 mol% CO$_2$ added. If we consider that CO$_2$ in the melt forms free CO$_2^{−}$ groups by taking a non-bridging oxygen from the silicate network we can calculate the expected effect on the melt NBO/Si. This theoretical effect is shown for two of the investigated composition in Fig. 8 alongside measurements from FPMD simulations. The predicted effect of CO$_2$ on the melt NBO/Si matches fairly well that measured by FPMD simulations.

As discussed in section 4.1, our spectroscopic investigations and FPMD simulations suggest that most of the carbon dioxide in transitional melts is dissolved as carbonate ions, in form of free carbonates. This prediction is fully consistent with the polymerisation effect of CO$_2$ we measure in experimental transitional glasses and observe in FPMD simulations.

### 4.3. Implications for melt physical properties and mobility

The degree of polymerisation of a melt often can be linked to its physical properties. The density, viscosity and compressibility of silicate melts depend on their degree of polymerisation (Gaskell, 1982; Mysen et al., 1982; Mysen and Richet, 2005). Other properties such as element diffusivity (e.g. Eyring, 1935; Tinker et al., 2004) electrical conductivity (e.g. Pommier et al., 2013) and sound velocity (Sakamaki et al., 2014) can also be related to melt polymerisation.

Our results show that carbon dioxide dissolved in transitional melts increases their degree of polymerisation implying that the physical properties of these melts will be highly dependent on their CO$_2$ content. The viscosity and density of transitional melt are then expected to be impacted by high CO$_2$ contents. Whether carbonated melt will ascend, sink or pond in the upper mantle, is dictated by their transport properties and hence, we argue, by their CO$_2$ content. Geophysical investigations of the upper mantle have defined a laterally continuous layer, the low velocity zone (LVZ), marked by a sharp reduction in seismic wave velocity and elevated electrical conductivity. Both features are sometimes related to melt ponding at the boundary between the rigid lithosphere and ductile asthenosphere (Sakamaki et al., 2013; Schmerr, 2012; Sifre et al., 2014). If the LVZ is indeed a region where melt ponds, there must be a physical reason explaining why such melt would accumulate and not ascend through the lithosphere. A conceivable hypothesis is that carbonated melts, with their higher than expected degree of polymerisation might be more viscous than expected, potentially impeding their rise and making them stall at the LVZ.

Conversely, the removal of CO$_2$ from a transitional melt will also strongly influence its structure. Ascending transitional melts...
such as those generating kimberlites will see their molecular structure change tremendously as a result of CO$_2$ degassing. Considering these melts could contain several tens of wt% CO$_2$ originally, they will become increasingly depolymerised as CO$_2$ is exsolved from the melt. The viscosity of these melts is therefore expected to decrease which should increase their mobility, potentially speeding up their ascent.

The molecular picture of transitional melt we have depicted here, is one in which the addition of CO$_2$ leads to the formation of free carbonate and an ever more polymerised aluminosilicate network. In essence, this view is one in which two sub-networks, a carbonated and a silicate one are cohabiting but getting increasingly disconnected from each other. This view is consistent with the realisation that many experiments in transitional melt composition result in immiscibility between a silicate and a carbonate liquid (Brooker et al., 2011; Brooker and Kjarsgaard, 2011). Liquid immiscibility might take place once a critical proportion of isolated carbonate groups has been reached.

CO$_2$ solubility in silicate melts is clearly related to the amount of non-bridging oxygen (e.g. Holloway et al., 1976; Brooker et al., 2001a; Ni and Keppeler, 2013), yet the aforementioned effect on CO$_2$ in transitional melts is to decrease the number of NBO. As transitional melts incorporate CO$_2$, their structure changes (number of NBO decreases) hence their capacity to dissolve CO$_2$ decreases. This structural approach can help explaining the very peculiar solubility law of CO$_2$ in transitional melts highlighted by Moussallam et al. (2014) who found a strongly non-linear dependence of CO$_2$ solubility to pressure. With increasing pressure CO$_2$ solubility first increases rapidly but after a certain amount of CO$_2$ is dissolved, the rate of CO$_2$ dissolution with increasing pressure diminishes and the pressure-solubility relationship becomes linear and comparable to that of a silicate melt (at least in the pressure range investigated).

5. Summary

We have investigated the structure of transitional (carbon-rich; silica-poor) melts quenched to glass by means of infrared, Raman, $^{13}$C and $^{29}$Si NMR spectroscopy and first principles molecular dynamic simulations. We have found that carbon in transitional glasses is present as CO$_2$-, most likely as free carbonate groups and that the addition of CO$_2$ strongly modifies the structure of transitional glasses resulting in a strong polymerisation of the aluminosilicate network. The key conclusions we draw from our findings are:

i. In transitional melts the addition of CO$_2$ results in the formation of free carbonate groups, effectively removing NBOs from the aluminosilicate network, hence forcing polymerisation.

ii. The resulting expected effect on transitional melt properties is an increase of their viscosity with increasing CO$_2$ content although this effect might be balanced by shear dilution of the aluminosilicate component. Ascent of carbonated melts in the mantle might be impeded by higher than expected melt viscosities and lead to melt ponding at certain levels within the mantle such as at the lithosphere–asthenosphere boundary.

iii. Carbon-rich transitional melts are effectively composed of two sub-networks; a carbonate and a silicate one. This peculiar molecular structure might lead to physical properties of the melt being determined by either sub-network. Electrical conductivity will be high as determined by the purely ionic structure of the carbonate sub-network while diffusivity of network-forming species will be low as determined by the polymerised structure of the silicate sub-network.

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Appendix A. Supplementary material

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References


