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Validation of the AIEM through correlation length parameterization at field scale using radar imagery in a semi-arid environment

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Abstract—This work aimed to validate the Advanced Integral Equation Model (AIEM) through different correlation length parameterizations using radar imagery for field-scale studies in a semi-arid environment. The study compared backscattering coefficients simulated from the AIEM and retrieved from SAR imagery of a study site in Sardinia. Two treatments for correlation length were adopted: *in situ* measurements and empirically based correlation length estimation. The results showed an overestimation of backscattering coefficients of 2.5 dB with an RMSE of 3.1 dB for HH and VV polarizations and an underestimation of 27.7 dB and an RMSE of 31.0 dB for HV polarization from the AIEM parameterized by *in situ* measurements. When using the AIEM with empirical correlation length, a bias of less than 1.0 dB was found with an RMSE of 1.7 dB for HH and VV polarizations and an overestimation of 1.1 dB and an RMSE of 5.1 dB for HV polarization. Better results were obtained with surface soil moisture (SSM) measured at 10 cm than at 5 cm. Promising soil moisture data retrieval from SAR imagery is expected from using the empirical correlation length-parameterized AIEM for field-scale purposes in semi-arid environments.

Index Terms-Correlation length, Soil moisture, SAR images, C-band, AIEM model, Semi-arid environment

I. INTRODUCTION

C urface soil moisture (SSM) distribution plays an important Drole in agricultural management tasks, such as irrigation modeling, at the field scale. Radar remote sensing has been used for retrieving and mapping SSM for decades. Various models have been developed for this purpose, ranging from site-dependent empirical models rooted in extensive databases [1-4] to semi-empirical models [5-9] to site-independent theoretical backscatter models [10-14]. For bare soil studies, the Integral Equation Model (IEM) is one the most widely used analytical physical backscatter models thanks to its large roughness validity domain and several further developments, such as the AIEM (Advanced IEM). Numerical models are also used, such as the Kirchhoff Approximation (KA) and the Numerical Maxwell Model of 3-D Simulations (NMM3D) [14, 15]. Few applications are reported for the AIEM [12, 16, 17].

Theoretical models require accurate surface roughness parameterization, which is labor-intensive for field measurements and is problematic when, for example, estimating correlation length. A semi-empirical calibration for the IEM based on a large *in situ* database was introduced [18-21] and extended to cross-polarization in Baghdadi et al. [22]. This approach relates the empirical correlation length to incidence angle, polarization, radar frequency and standard deviation of surface height (RMS height), restricting the surface roughness parameter to RMS height only. Good agreement has been found between simulated backscatter and SAR images for incidence angles between ca. 20°-50° for C-band signal.

This work aimed to validate the AIEM model for use in a field-scale operational approach in development through the CLIMB project [23]. An introduction of the study site and a database description are given in Section II, followed by a description of the backscatter model in Section III. The AIEM is evaluated in Section IV. The findings and results are

discussed and conclusions are offered in the last section.

II. STUDY SITE AND DATASET

A. Azienda San Michele

The *Campidano Plain* is the agricultural heartland of Sardinia and is located in the southern part of the island. The study site, *Azienda San Michele* (Figure 1), has a total area of 435 hectares and is located on the east edge of the *Campidano Plain* between the villages *Ussana* and *Donori*, with central coordinates of 39°25'N, 9°06'E. The *Azienda* is one of the research-based farms operated by the Agricultural Research Center of Sardinia, the AGRIS (*Agenzia per la ricerca in agricoltura*). Several reference fields of the experimental farm ranging from 1.7 to 4.4 hectares were kept in bare conditions for a two-month period (May – June in both 2008 and 2009).



Figure 1 Land-use map of selected reference fields of *Azienda San Michele* in UTM coordinate system zone 32N (indicating field boundaries, corner reflectors, bare fields and tested crop fields, e.g., bean, wheat and canola fields).

B. Radar Data

A total of 13 dual-polarization ASAR APS images and 11 Radarsat-2 Fine-Quad-Polarization images, including both ascending and descending modes, were acquired over the campaign periods (Table 1). ASAR images covering a swath width of up to 100 km were normalized to a spatial resolution of approximately 30 m. Radarsat-2 images covering a swath width of approximately 25 km were normalized to a spatial resolution of approximately 12.5 m. The radiometric precision is within 0.5 dB in both sensors [24]. An average weekly coverage was guaranteed during the campaign period over the study site.

Date (YYMMDD)	Sensor	Incidence angle (°)	Polarizations
080503	ASAR	23	HH, VV
080522	ASAR	23	HH, VV
080524	Radarsat-2	22	HH,VV, HV, VH
080527	Radarsat-2	21	HH,VV, HV, VH
080528	ASAR	19	HH, VV
080531	ASAR	29	HH, VV
080606	ASAR	41	HH, VV
080613	Radarsat-2	27	HH,VV, HV, VH
080616	ASAR	23	HH, VV
080617	Radarsat-2	22	HH,VV, HV, VH
080624	Radarsat-2	28	HH,VV, HV, VH
090421	ASAR	19	HH, VV
090425	Radarsat-2	22	HH,VV, HV, VH
090427	ASAR	23	HH, VV
090428	Radarsat-2	21	HH,VV, HV, VH
090504	ASAR	29	HH, VV
090513	ASAR	19	HH, VV
090516	ASAR	29	HH, VV
090519	Radarsat-2	22	HH,VV, HV, VH
090522	Radarsat-2	21	HH,VV, HV, VH
090523	ASAR	23	HH, VV
090601	ASAR	23	HH, VV
090612	Radarsat-2	22	HH,VV, HV, VH
090615	Radarsat-2	21	HH,VV, HV, VH

C. Field Data

Extensive field measurements were collected from 3rd May to 25th June 2008 and from 21st April to 16th June 2009. The SSM was measured by an FDR (frequency domain reflectometer) at depths of 5 cm and 10 cm (Figure 2) and the gravimetric method was adopted as a reference. The values at each depth were averaged for each reference field to a single representative value for the reference field at the corresponding depth. The SSM measurements were taken each day in accordance with the satellite pass. Surface roughness was measured using the close-range photogrammetric technique with a Rollei 7 camera for each sample point on bare fields [25]. Digital surface models (DSMs) were derived from pairs of photos, where each photo covered a surface area of 1 m by 1 m. The RMS height was calculated from the DSM. Both exponential and Gaussian semivariogram functions were used in ArcGIS 9.x for correlation estimation

from point data of stereo image pairs [26, 27]. Values from different sample points (SPs) were averaged to generate a single representation of each study field. Table 2 demonstrates the minimum and maximum values of field-averaged SSM, RMS height (s) and Gaussian correlation length (l). The SSM at both depths ranged from extremely dry (~3.0 vol. %) to nearly 30 vol. % during the campaign periods. On average, the SSM at a 10 cm depth was 5.9 vol. % higher than the measurement at a 5 cm depth. Tables 3 and 4 showed the infield variability, i.e., the variability between SPs within a field, of these parameters for two sampling dates: 6th June 2008 (a wetter period) and 15th June 2009 (a drier period). Note that: 1) between different reference fields within the experimental farm, the field-averaged SSM can reach a difference higher than 10.0 vol. % and 2) the SSM-values of different sample points showed low variations (standard deviation lower than 4 vol. % for all reference fields).



Figure 2 Comparison of SSM at 5 cm and 10 cm depths in 2008 and 2009 showing an RMSE of 4.3 vol. % and a bias of 5.9 vol. %.

TABLE 2 IN-SITU GEOPHYSICAL CHARACTERISTICS AND MEASURED DAT	ΓES
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(YYMMDD).					
	SSM_5 cm	SSM_10	cm	s (cm)	l (cm)
	(vol. %)	(vol. %)			
Min.	1.3	3.1		1.7	15.3
Date	090513	090615		090516	090513
Max.	25.6	29.2		3.2	28.2
Date	080606	090427		080617	080503

TABLE 3 IN-FIELD VARIABILITY OF SURFACE GEOPHYSICAL PARAMETERS FOR 6^{TH} June 2008 (mean value±standard deviation).

Field	SSM_5 cm (vol. %)	SSM_10 cm (vol. %)	s (cm)	<i>l</i> (cm)
F11	25.6 ± 2.1	25.5 ± 2.8	1.6 ± 0.2	20.3 ± 4.9
F21	11.3 ± 2.5	17.8 ± 2.5	$2.0\!\pm\!0.4$	17.6 ± 4.8
F31	6.9 ± 0.9	8.3 ± 1.5	1.9 ± 0.1	26.8 ± 14.3

TABLE 4 IN-FIELD VARIABILITY OF SURFACE GEOPHYSICAL PARAMETERS FOR 15^{TH} June 2009 (MEAN VALUE + STANDARD DEVIATION)

10				
Field	SSM_5 cm (vol. %)	SSM_10 cm (vol. %)	<i>s</i> (cm)	<i>l</i> (cm)
F11	10.6 ± 5.3	21.3 ± 4.1	$2.1\!\pm\!0.2$	20.3 ± 5.2
F21	3.0 ± 1.5	13.5 ± 6.3	$1.7\!\pm\!0.3$	18.0 ± 5.2
F32	2.6 ± 1.0	3.2 ± 1.2	$1.6 {\pm} 0.1$	16.0 ± 1.0

III. AIEM

The IEM was developed and introduced by Fung et al. [28, 29]. The IEM has been well described (e.g., [18, 30]) and the equations are not reprinted here. The IEM is a widely used analytical physical model and is known for its comprehensive validity under roughness conditions. Wu and Chen [11] updated and improved the IEM by replacing Fresnel reflection coefficients with a transition function that considers both surface roughness and permittivity. Because both the estimation and sensitivity analysis of the surface roughness information in the backscatter model can be problematic, especially concerning the correlation length l, roughness parameters must be treated carefully. Brogioni et al. [16] assessed the AIEM validity for roughness parameterization resulting in a slight extension from the IEM validity [31]. They found that valid RMS heights and correlation lengths range from ~0.77 to 11.1 cm and ~0.11 to 110 cm, respectively, for the C-band signal of the AIEM.

Baghdadi et al. [19, 22] proposed replacing l with a fitting parameter "Lopt" with a Gaussian function for the IEM. The Lopt depends on a given radar wavelength, incidence angle θ , RMS height *s* and polarization. The empirical correlation length Lopt is expressed as in equations 1-3 for the C-band:

$$L_{opt}(s,\theta,HH) = 0.162 + 3.006 (\sin 1.23\theta)^{-1.494} s$$
(1)

$$L_{out}(s,\theta,VV) = 1.281 + 0.134 (\sin 0.19\theta)^{-1.59} s$$
(2)

$$L_{out}(s,\theta,HV) = 0.9157 + 1.2289 (\sin 0.1543\theta)^{-0.3139} s$$
(3)

Note that *s* is described in cm and θ is described in degrees. Good results were found in [18, 19, 22], where simulated backscattering coefficients and those retrieved from SAR agreed well for soil moistures between 5% and 47%. Equations 1-3 will be applied in this study to measurements collected in a semi-arid environment that is drier than that of the study sites in [18, 19, 22].

IV. MODEL VALIDATION

The empirical correlation length is calculated from the *in* situ RMS height and incidence angle in accordance with the polarization (equations 1-3). The *Lopt* values were lower in the HV than the HH and VV polarizations. The comparison between *in situ* l and *Lopt* showed RMSEs of approximately 7 cm and 12 cm for the co-polarizations (HH and VV) and the cross-polarization, respectively (Figure 3). Moreover, l was larger than *Lopt* by approximately 2 cm, 1 cm and 11 cm for the HH, VV and HV polarizations, respectively.

Figure 4 showed a comparison between the backscattering coefficients retrieved from SAR and those simulated from the AIEM, with different ranges of the incidence angle. The simulations were conducted based on SSM_5 cm and SSM_10 cm. The empirical correlation length was also calculated according to equations 1-3. Figures 4a, 4b and 4c were plotted from *in situ* experiments Gaussian correlation length and SSM_10 cm for the HH, VV and HV polarizations, respectively and were used to exemplify the backscatter simulations from the *in situ* measurements. Similarly, for the

HH, VV and HV polarizations, Figures 4d, 4e and 4f were plotted simulations from the empirical correlation length and the SSM at 5 cm, while Figures 4g, 4h and 4i were plotted simulations from the empirical correlation length and the SSM at 10 cm. Note that only 33 samples were available for cross-polarization from Radarsat-2 data against 70 samples for HH and VV.



Figure 3 Comparison of empirical correlation length *Lopt* for the HH, HV and VV polarizations and *in situ* correlation length *l* in cm. Each dot represents a correlation value for one reference field and one image acquisition date.

For *in situ* correlation length *l*, the backscattering coefficients generated from SSM_10 cm fit the HH and VV polarizations better than did those from SSM 5 cm. For the HV polarization, the selections of l at 5 cm or 10 cm provided similar simulations. By comparing the backscattering coefficients simulated from SSM_5 cm and those retrieved from the SAR imagery, large underestimations of over 14.0 dB and RMSEs over 27.0 dB were found for both the HH and VV polarizations for the Gaussian correlation function. For HV polarization, an underestimation of 31.0 dB and an RMSE exceeding 33.0 dB were observed. Figures 4 a-c showed that the AIEM simulated from SSM_10 cm results in an overestimation of approximately 2.5 dB and an RMSE of approximately 3.1 dB for the HH and VV polarizations, a major improvement. However, a significantly large underestimation of over 27.7 dB and an RMSE of 31.0 dB were found again for the HV polarization.

For the empirical correlation length *Lopt*, a comparison between figures 4 d-f and 4 g-i showed clear improvement in both RMSEs and bias for the simulations from the SSM_10 cm depth over those from SSM_5 cm. Overestimations of approximately 2.7 dB, 2.1 dB and 2.2 dB and RMSEs of approximately 3.1 dB, 2.7 dB and 6.0 dB were found for the AIEM simulations from SSM_5 cm for the HH, VV and HV polarizations, respectively. For the simulations from SSM_10 cm, the biases were restricted to no more than 1.1 dB and the RMSEs were reduced to 1.6 dB, 1.8 dB and 5.1 dB for the HH, VV and HV polarizations, respectively.

The effect of incidence angle on the performance of the AIEM model was not analyzed because the available samples were acquired mainly with incidence angles between 20° and 30° . The few samples with incidence angles of less than 20° or greater than 30° were not sufficient to analyze the influence of

the incidence angle on the quality of the AIEM simulations.



(i)

Figure 4 Comparison between backscattering coefficients estimated from the AIEM model and those calculated from the SAR images for the HH, VV and HV polarizations: s and l using *in situ* SSM_10 cm (a, b, c); s and *Lopt* using in situ SSM_5 cm (equation 1) (d, e, f); and s and *Lopt* (equation 1) using *in situ* SSM_10 cm (g, h, i). Each dot represents a

correlation value for one reference field and one image acquisition date. Bias is defined as AIEM-SAR values.

The better performances of the AIEM based on the SSM at a 10 cm depth could be due to the significant penetration of the radar wave in this semi-arid environment, i.e., the radar signal is more sensitive to deep layers in dry soil.

For a more detailed study of the model performance at different depths, the dataset was separated into two parts: (1) samples with SSM_5 cm < 10 vol. % and (2) samples with SSM_5 cm > 10 vol. %, independent of the SSM at 10 cm. For SSM_5 cm < 10 vol. %, the difference between the AIEM with Lopt and SAR data (bias) decreased from +2.3 dB (SSM_5 cm) to -0.5 dB (SSM_10 cm) for HH, from +1.80 dB (SSM_5 cm) to +0.7 dB (SSM_10 cm) for VV and from -5.4 dB (SSM_5 cm) to -1.1 dB (SSM_10 cm) for HV. For SSM_5 cm > 10%, the bias decreased from +3.4 dB (SSM_5 cm) to +0.3 dB (SSM_10 cm) for HH and from +2.7 dB (SSM_5 cm) to +1.5 dB (SSM_10 cm) for VV. The statistics for HV were not calculated because only 10 samples were available for SSM_5 cm > 10%.

In the totality of samples with all polarizations, the statistics showed that the difference between the AIEM with Lopt and the SAR data decreased from +0.7 dB (SSM_5 cm) to -0.2 dB (SSM_10 cm) for SSM_5 cm < 10 vol. % and from +3.3 dB (SSM_5 cm) to +1.8 dB (SSM_10 cm) for SSM_5 cm > 10 vol. %. The RMSE decreased from 3.7 dB (SSM_5 cm) to 2.4 dB (SSM_10 cm) for SSM_5 cm < 10 vol. % and from 3.7 dB (SSM_5 cm) to 3.2 dB (SSM_10 cm) for SSM_5 cm > 10 vol. %.

V. DISCUSSION AND CONCLUSIONS

This study confirmed the applicability of Lopt in the AIEM model. Moreover, these results showed that the use of Lopt ensures better agreement between the AIEM model and the SAR data when using a dataset acquired in a semi-arid environment with an SSM 10 cm between 3.1 vol. % and 29.2 vol. %. The RMSE was lower for the HH and VV polarizations than for the HV polarizations. However, these results were obtained with fewer samples for HV (33 points for HV and 70 points for HH and VV). This approach can help to minimize the uncertainties introduced by correlation length measurement. Comparably, Huang et al. [14] reported a promising confidence level of approximately 2.0 dB in bias for the AIEM performance in a total of 34 samples, while this study showed a bias within 1.0 dB for HH and VV polarizations from 70 samples and a bias of 1.1 dB for HV polarization from 33 samples. Additional theoretical models may need to be considered to assess the empirical correlation length.

The empirical correlation length solves the scale effect on the parameterizing correlation length. When accurate inversions of both RMS height and SSM from remote sensingbased approaches are available, such as when using full polarimetric SAR [34], the AIEM coupled with Lopt can be used as an operational approach. Promising soil moisture data retrieval from SAR imagery is expected from the use of

empirical correlation length-parameterized AIEM for fieldscale purposes in semi-arid environments.

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