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## 13TH ORDER RESONANCE FROM NAVY TRACKING ON A DIADEME 2 FRAGMENT

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## GODDARD SPACE FLIGHT CENTER

 GREENBELT, MARYLAND
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#### Abstract

A strong constraint on 13th order (odd degree) terms in the geopotential has been derived from Navy tracking on a DIADEME 2 fragment (1967-14F). This object (perigee height: 580 km , orbit inclination: $38.9^{\circ}$ ) is presently decaying slowly through perfect commensurability with these terms. The resonance forces will increase its inclination by $0.02^{\circ}$ when the passage is complete by late 1974. The constraint (lumped harmonics), derived by adjustment of a pair of harmonic coefficients to the Navy inclination data (principally) is: $10^{9}(14.8 \pm 0.8,48.3 \pm 0.7)=0.023(\mathrm{C}, \mathrm{S})_{13,13}-0.172(\mathrm{C}, \mathrm{S})_{15,13}+0.505(\mathrm{C}, \mathrm{S})_{17,13}$ $-0.884(\mathrm{C}, \mathrm{S})_{19,13}+(\mathrm{C}, \mathrm{S})_{21,13}-0.673(\mathrm{C}, \mathrm{S})_{23,13}+0.099(\mathrm{C}, \mathrm{S})_{25,13}+0.295(\mathrm{C}, \mathrm{S})_{27,13}$ $$
-0.279(\mathrm{C}, \mathrm{~S})_{29,13}+0.018(\mathrm{C}, \mathrm{~S})_{31,13}+\ldots
$$


There should be a significant contribution to this result from terms as high as 29th degree. But current geopotential solutions (for 13 th order terms) to this degree are about $20 \%$ in error when judged by this independent data.

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## INTRODUCTION

Among the thousands of Earth satellites launched in the past 15 years, hundreds must have suffered or will suffer measurable effects from poorly known resonances with the Earth's geopotential (King-Hele, 1973a; Gabbard and Wackernagel, 1971). But less than 50 of these objects have actually been used geodetically. The major difficulty of course is the lack of adequate tracking for the majority of the objects which are (or were) rocket bodies or fragments (debris) of larger satellites. However, in the case of deep resonance or close commensurability to the Earth's rotation, the effects are so large that even tracking with crude instruments can reveal them (King-Hele, 1973b). Many of the observations made in support of the tracking on decaying 15 th order resonant orbits were accomplished with amateur observers using binoculars and stopwatches.

In these deep resonances the change of the inclination of the orbit is of the order of $0.01^{\circ}$. To achieve overall inclination accuracies of a tenth of this figure, a topocentric angle good to only about a minute of arc is necessary. KingHele's results were often much better than this, but his amateurs were backed up by more precise camera observations. (See also Winterbottom and KingHele, 1974.) In addition, King-Hele's group has utilized minute-of-are data from the U.S. Navy skin track Naval Space Surveillance (NAVSPASUR) system,
of which there is a great abundance on low altitude satellites. In fact since 1972, the NAVSPASUR data and mean elements derived from them (by the Navy) dominate the orbits used in the 15th order resonance analysis (i.e., King-Hele and Walker, 1972). This data consists of (direction cosine) pointing angles from a fence of receiving stations across the United States. The signals are reflections from the satellite of a continuous high power radar fan illuminating space across the U.S. at about $28^{\circ}$ latitude. The system has tracked objects as high as $18,000 \mathrm{~km}$ altitude. Providing the inclination of the orbit is greater than about $28^{\circ}$ and the object is large enough to be tracked, data on each revolution will be received.

The existence of this good data on a large number of objects opens up the possibility of examining in detail all the deep geopotential resonances besides the one's of 15th order. In this study, the Navy's own mean elements were used directly. They proved sufficiently accurate to reveal the strong effect on the inclination of the decaying (13 revolutions/day) orbit of a DIADEME 2 fragment (1967-14F).

ANALYSIS

In Figure 1, the Navy Mean Inclinations are shown from February 1972 to January 1974. Also in Figure 1 is a plot of the mean (primary) resonant longitude rate for this 13 revs/day orbit:

$$
\dot{\psi}_{13,0}=\dot{\omega}+\dot{M}+13\left(\dot{\Omega}-\dot{\theta}^{\prime}\right)
$$

where $\omega, \mathrm{M}$ and $\Omega$ are the orbit's argument of perigee, mean anomaly and right ascension of the ascending node, and $\dot{\theta}$ is the Earth's rotation rate.


Figure 1. U.S. Navy Orbit Inclinations for 1967-14F

It is recalled that the longitude rates which determine the frequency of gravitational perturbations on an orbit in Kaula's development of the geopotential (Kaula, 1966; especially p. 40, 49 and 55) are given as:

$$
\dot{\psi}=(\ell-2 p) \dot{\omega}+(\ell-2 p+q) \dot{M}+m(\dot{\Omega}-\dot{\theta})
$$

where $\ell$ and $m$ are the degree and order of a geopotential harmonic, and $p$ and $q$ are additional indices related to the inclination and eccentricity of the orbit. The resonances $(\dot{\psi}=0, m \neq 0)$ will occur for $(\ell-2 p+q) \dot{\mathrm{M}}$ close to $\mathrm{m} \dot{\theta}$, since $\dot{\omega}$ and $\dot{\Omega}$ are small. The primary resonances (strongest) occur when $\ell-$ $2 p+q=1$, since then $m$ will be minimum, the closest integer to $\dot{M}$ in revolutions/day. Secondary resonances (sub harmonics) for a given $\dot{\mathrm{M}}$ (near a rational number of revs/day) will occur for $\ell-2 p+q=2,3,4 \ldots$, but in any case all the resonant longitude rates can be characterized by the order m and the $q$ index, by writing

$$
\dot{\psi}_{\mathrm{m}, \mathrm{q}}=-\mathrm{q} \dot{\omega}+(\ell-2 \mathrm{p}+q)(\dot{\omega}+\dot{\mathrm{M}})+\mathrm{m}(\dot{\Omega}-\dot{\theta}) .
$$

For a given resonant order $m$, there will be a series of "side band" resonances characterized by $q$ around the (generally) dominant one for $q=0$. The $q=0$ resonance is also the mean of the series since q can take on all positive or negative integers.

It is seen in Figure 1 that the longitude rate $\dot{\psi}_{13,0}$ for $1967-14 \mathrm{~F}$ goes to zero over the period of record as the inclination of the orbit suffers a progressive oscillation of increasing period and amplitude. The increasing period closely matches the period of $\psi_{13,0}$. There is clearly a strong resonant
perturbation of this orbit entirely analogous to the decaying 15 th order orbits first analyzed by Gooding (1971) and the 11th order orbit examined by Wagner (1973).

But there is more significant detail in Figure 1 than resonance. The "raw" mean inclinations given by the Navy (see also Table 1) are actually Brouwer elements (determined over independent 7 day arcs) with both short and long period zonal effects removed (Brouwer, 1959). Only $J_{2}$ to $J_{5}$ are used in the NavyBrouwer model and the coefficients are not up to date. The period of the principal odd zonal inclination effect $(2 \pi / \dot{\omega})$ is 65 days. The amplitude, with the correct model, is about $0.005^{\circ}$. The error in the Navy model is certainly seen in the raw inclination data. But, in addition, there is an 84 day lunar perturbation with amplitude $0.0014^{\circ}$ which is also observable.

To clarify the quality of this "signal" I have added back the long period zonal effects used in the Navy-Brouwer model to produce mean elements free of this bias (also Table 1). Then I compared these (less smooth) mean elements to values calculated from a trajectory which includes all significant (but nonresonant) long period effects on the orbit, in particular the zonal perturbations from the Smithsonian (SAO) Standard Earth 2 (Gaposchkin and Lambeck, 1971), radiation pressure, atmospheric drag and direct lunar-solar gravity. The comparison (observed minus calculated values) in Figure 2, shows the resonant signal much more strongly and also reveals what appears to be a much reduced residual effect due to odd zonal error in the Smithsonian field.






NOTE: Unprimed vaiues are original Navy Mean Elements, the equivalent of Brouwer double primed elements. The primed valubs (used in the orbit determinations) are elements with Navy-used long period terms added back in, the equivalent of Brouwer single primed values. The Navy-used zonal coefficients are: $10^{6} J_{2}=1082.63,10^{6} J_{3}=-2.55$,
$10^{6} J_{4}=-1.61,10^{6} \mathrm{~J}_{5}=-0.19$.


Figure 2. Resonant Variation of the Inclination for the Orbit of 1967-14F

The residuals in inclination (rms) from this trajectory (and others including those with resonant effects) are shown in Table 2. The calculated trajectories were "fit" to the mean element "observations" by differential correction of initial elements and other model parameters (see i.e., Morrison, 1970 or Wagner and Douglas, 1970) under the conditions stated in Table 2. The principal difficulty in these orbit determinations was in following the mean anomaly of the satellite which (with a perigee of 580 km ) underwent fluctuations of tens of degrees from drag error. The level of this error was less than $10 \%$ after correction for a single drag coefficient. Yet it was necessary to include additional accelerations in the semimajor axis and mean anomaly to permit the critical resonant longitude $\psi_{13,0}$ to be calculated to better than $5^{\circ}$ (rms). (This error is consistent with the formal accuracy of the resonance determination.)

A more readily available data source for satellite objects, the North American Air Defense Command's SPADATS elements were also evaluated for 196714F by the same mean element program. (See run 2 of Table 2.) Figure 3 shows the inclination residuals from a nonresonant trajectory with this data over a somewhat shorter are in 1972-1973. For most of the period no detail at all can be seen. At the end of the arc some definition of the resonance appears. The NORAD-SPADA TS data quality is clearly not uniform in this arc. The small "acceptable" portion could not be used for adequate resonance recovery.

Table 2
Results of Orbit and Coefficient Determinations for 1967-14F Using Navy and NORAD Data

| Run | Field Used | $\begin{aligned} & \text { Residuals* } \\ & \text { in } \\ & \text { Inclination } \\ & \text { (rms) } \\ & \left(10^{-3} \mathrm{deg}^{\prime} \mathrm{s}\right) \end{aligned}$ | Data Used | Comments |
| :---: | :---: | :---: | :---: | :---: |
| 1 | SAO SE 2 (non resonant) | 5.41 | NAVSPASUR - <br> All elements <br> MJD 41303 - <br> 41989 | Drag, radiation, $\ddot{a}$ and $\ddot{M}$ coefficients solved from data. Commensurability at 41985. Resonance it "I" clear, residuals about $0.001^{\circ}$. |
| 2 | SAO SE 2 (non resonant) | 6.25 | SPADATS - <br> All elements <br> MJD 41302 - <br> 41936 | Same as above. Resonance in I not seen except possibly after MJD 41700. Residuals about $0.01^{\circ}$. |
| 3 | SAO SE 2 (non resonant) | 6.83 | Same as run \#1, but to MJD 42060 | Same as run \#1. |
| 4 | SAO SE $2+3,0$ and $(23,13)$ solved from data $\begin{aligned} & 10^{9}(\mathrm{C}, \mathrm{~S})_{23,13}= \\ & (-20.7 \pm 1.4,-69.8 \\ & \pm 1.2) \end{aligned}$ | 1.13 | Same as run \#1. | Same as run \#1. <br> Weight on I: <br> $0.0002^{\circ}$. Maximum along track error: $6^{\circ}$. |

[^0]Table 2 (continued)

| Run | Field Used | $\begin{gathered} \text { Residuals* } \\ \text { in } \\ \text { Inclination } \\ \text { (rms) } \\ \left(10^{-3} \mathrm{deg}^{\prime} \mathrm{s}\right) \end{gathered}$ | Data Used | Comments |
| :---: | :---: | :---: | :---: | :---: |
| 5 | SAO SE $2+3,0$ and <br> $(2.3,13)$ solved from data $\begin{aligned} & 10^{9}(\mathrm{C}, \mathrm{~S})_{23.13}= \\ & (-22.0 \pm 1.2,-71.8 \\ & \pm 1.1) \end{aligned}$ | 1.29 | NAVSPASUR All elements to MJD 42060 | Drag, radiation, ä, ä, $\mathrm{M}^{[3]}, \mathrm{M}^{[4]}$, $\mathrm{M}^{[5]}$ coefficients solved from data. Short period lunar terms used. Maximum along track error: $10^{\circ}$. Correlation coefficient: $(\mathrm{C}, \mathrm{S})_{23,13}=-0.38$. Weight on I: $0.0002^{\circ}$. |
| 6 | SAO SE $2+3,0$ and <br> $(23,13)$ solved from data $\begin{aligned} & 10^{9}(\mathrm{C}, \mathrm{~S})_{23,13}= \\ & (-21.2 \pm 3.4,-71.3 \\ & \pm 2.4) \end{aligned}$ | 1.29 | NAVSPASUR - <br> Inclination data only to MJD 42060 | All state elements except "I" and radiation coefficient fixed from solution in run \#5. Correlation coefficient $(\mathrm{C}, \mathrm{S})_{23,13}=0.34$. |

*Observed - calculated values from converged mean element trajectory


Figure 3. Inclination Residuals from a Nonresonant Trajectory for 1967-14F Using NORAD-SPADATS Observations

## RESONANCE RECOVERY

The geopotential indices for this (mean) primary resonance ( $\ell-2 p+q=1$ ) are $\mathrm{m}=13, \mathrm{q}=0$ and all $\ell \geq 13$ for which

$$
\ell=2 \mathrm{p}+1, \mathrm{p} \leq \ell .
$$

Thus the $\ell$ are all odd and any one of the $(\ell, m)$ terms $(13,13),(15,13),(17,13)$ . . . can be recovered from the resonance in inclination on this orbit.

It is noted that the first of the side band primary resonances $(q= \pm 1)$ which are affected by the even degree geopotential term (and $m=13$ ) have only slight effect in this data span since $\dot{\omega}=5.6^{\circ} /$ day (relatively large compared to $\psi_{13,0}$ in this period). The $q=-1$ resonance has a period of 180 days at 4300 MJD declining to 60 days at 42060 MJD. The $q=+1$ resonance has a period of only 40 days at the beginning of the data span, increasing to 70 days at the end. In addition, the amplitudes of these terms are of order $e^{|q|}$, where $e$ is the orbit's eccentricity, [see Kaula, 1966, p. 37] or reduced by a factor of 0.08 with respect to the $q=0$ resonance for $1967-14 \mathrm{~F}$.

The fact that (essentially) only a single geopotential term can be recovered from each resonance is a consequence of the fact that only a single harmonic perturbation is responsible for it (see King-Hele, 1973a or Wagner, 1973). The amplitude of this term changes only as a consequence of the change in frequency through the resonance. The fundamental amplitude (a weighted sum of the resonant geopotential coefficients) remains constant. The scale of this sum is determined by the actual resonance perturbation. However, even for the same
fundamental frequency $(\psi)$ this weighted sum is different for 4 of the Kepler element variations (a, e and " I " have the same sum, but $\omega, \Omega$ and M are distinct).

Why (physically) only 4 of the 6 sums (or amplitudes) yield independent information is unclear. It may be due to the choice of the classical Kepler elements to express the perturbations. In particular, the choice of the mean anomaly instead of the true anomaly introduces an infinite number of frequencies to express the effects of a single geopotential harmonic on an orbit. In some sense this must dilute the information content of any single frequency. But in any case, for the dragged orbits, only the inclination variation (essentially free of drag error) has provided unambiguous recovery of resonance information. However, recently, King-Hele (1973c) and Winterbottom and King-Hele (1974) have shown that near circular orbits can suffer significant resonances of $\Omega$, e and $\omega$ apparently quite distinct from drag (and other) effects. A strong resonance in e was also seen by Wagner (1973) in the very slow decay of the Vanguard 3 orbit. On the other hand no significant resonance of other elements on the DIADEME 2 fragment has yet been seen, though the full passage will not be over till late 1974.

I chose the geopotential harmonic $(23,13)$ to absorb the inclination resonance on 1967-14F and at the same time made a $5 \%$ adjustment of $(3,0)$ to correct odd zonal model error (see Table 2, runs 4-6). The same differential correction program was used in these adjustments as previously to reveal the
resonance. But added weight was given to the inclination data in these adjustments and additional secular accelerations were used to reduce the along track or phase error of the resonance. Full data correction runs were made for 686 and 757 day arcs (runs 4 and 5). The inclination residuals were significantly larger in the longer arc, possibly showing the influence of the secondary resonance with $\mathrm{m}=26$. But this conclusion must wait till the passage is complete and the smaller (higher frequency) effect can be well separated. In any case, the $(23,13)$ recovery is not substantially altered over the longer span. A final run was made using the inclination data only, to confirm that the other elements, influenced by drag, were not distorting this result. The $(23,13)$ recovery in this correction (run 6) was between the values in the two full data analyses.

The solid curve in Figure 2 shows the computed resonance in inclination from the best (most representative) results to date (run 5). The values for the harmonic coefficients themselves are somewhat larger than Kaula's rule [Kaula. 1966, p. 98]. But this is not significant since it is only a linear sum of resonant terms which is well determined.

## RESONANT CONSTRAINT (LUMPED COEFFICIENTS)

Following the method developed by Gooding (1971) and elaborated by Wagner (1973) the resonant inclination variation for $1967-14 \mathrm{~F}\{\mathrm{a}=1.88 \mathrm{e} . \mathrm{r} ., \mathrm{e}=0.082$, $\mathrm{I}=38.92^{\circ}$ ) is determined by the lumped sine and cosine terms:

$$
\begin{gather*}
(\mathrm{C}, \mathrm{~S})_{13,0}=0.023(\mathrm{C}, \mathrm{~S})_{13,13}-0.172(\mathrm{C}, \mathrm{~S})_{15,13}+0.505(\mathrm{C}, \mathrm{~S})_{17,13} \\
-0.884(\mathrm{C}, \mathrm{~S})_{19,13}+1.000(\mathrm{C}, \mathrm{~S})_{21,13}-0.673(\mathrm{C}, \mathrm{~S})_{23,13}+0.099(\mathrm{C}, \mathrm{~S})_{25,13} \\
+0.295(\mathrm{C}, \mathrm{~S})_{27,13}-0.279(\mathrm{C}, \mathrm{~S})_{29,13}+0.018(\mathrm{C}, \mathrm{~S})_{31,13}+0.156(\mathrm{C}, \mathrm{~S})_{33,13} \\
-0.105(\mathrm{C}, \mathrm{~S})_{35,13}-0.036(\mathrm{C}, \mathrm{~S})_{37,13} \\
+0.085(\mathrm{C}, \mathrm{~S})_{39,13}-0.021(\mathrm{C}, \mathrm{~S})_{41,13}+\ldots \tag{1}
\end{gather*}
$$

The terms in this series are for fully normalized geopotential harmonics [Kaula, 1966, p. 7]. The weights are just the fundamental amplitudes (without the rate denominator) of the linear perturbation of the inclination due to a fully normalized harmonic [Kaula, 1966, p. 40]. The sum of these fundamental amplitudes (with $C$ and $S$ coefficients) are merely the coefficients of the cos $\left(\psi_{13,0}\right)$ and $\sin \left(\psi_{13,0}\right)$ terms determining the rate of the resonance variation of the inclination. It is these two (lumped) coefficients which are actually "observed" in this resonance. The reason $(23,13)$ was chosen to absorb the effect was because it had a high weight in (1). It also was the lowest degree 13th order term not present in the SAO SE 2 field. Using the $(23,13)$ weight of -0.673 , the "observed" lumped coefficients are (best results):

$$
\begin{equation*}
10^{9}(\mathrm{C}, \mathrm{~S})_{13.0}=(14.8 \pm 0.8,48.3 \pm 0.7), \tag{2}
\end{equation*}
$$

with a correlation coefficient of -0.38 .
With regard to "lumped coefficients", if their perturbations can be properly identified, the weights of their constituents can serve to extend the information in them to any degree. Such "lumped coefficients" have been reported many times from "shallow resonant" satellite orbit analyses 「i.e., Yionoulis, 1965;

Murphy and Cole 1968; Gaposchkin and Veis 1967; Douglas and Marsh, 1970]. One difficulty with "shallow resonant coefficients" is that except for near circular orbits, the side band resonances (i.e., $q= \pm 1$ ) will also be observable. They will not be separable unless long ares of data are analyzed spanning at least a rotation of perigee. But if the "Iumped coefficients" are well determined (and identified), as here and in the previous 15th and 11th order analyses, they can serve as absolute benchmarks for geopotential determinations. An example of this use has been given by Wagner (1973) for a previously poorly observed 11th order resonance. Here, I calculate the "lumped coefficient" for the 1967-14F resonance from Equation (1) with a number of recent fields which are well represented by 13 th order terms. The results are presented in Figure 4. The merit of each of these fields is now immediately apparent when compared to the "observed" coefficients.

## DISCUSSION OF RESULTS

The "point" in Figure 4 representing the "observed" lumped coefficients has two uncertainty circles about it. The inner one represents the formal $1 \sigma$ uncertainty of the "best" solution. The small solution shift from the abbreviated data analysis is also seen. The larger circle represents the uncertainty in the lumped coefficient if terms of degree higher than 30 are ignored (as they generally are in current global solutions). This expected "truncation error" is calculated as the root sum of squares for the neglected terms in Equation (1) using Kaula's rule ( $10^{-5} / \ell^{2}$ ) for the harmonic coefficients (Table 3). It is also

## CALCULATED FROM

- PGS 11
$\triangle$ GEM 5
- GEM 6

】 PGS 62
4 SAO SE 2

- SAO SE 3

OBSERVED FROM 1967-14F RESONANCE

- BEST RESULTS
x RESULTS FOR DATA THRU 41989 MJD
ESTIMATED EFFECT FOR ALL TERMS
-     - OF DEGREE GREATER THAN 29

Figure 4. Lumped Harmonic for 1967-14F Resonance

Table 3

Estimated Cumulative Effect on Lumped Harmonic for 1967-14F,
From Geopotential Terms

| $\ell$ | 13 | 15 | 17 | 19 | 21 | 23 | 25 | 27 | 29 | 31 | 33 | 35 | 37 | 39 | 41 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| $10^{9}$ RSS* (all terms $\left.\geq \ell\right)$ | 40.9 | 40.9 | 40.2 | 36.2 | 26.6 | 13.9 | 5.8 | 5.5 | 3.8 | 1.9 | 1.9 | 1.3 | 0.9 | 0.8 | 0.6 |

*Using Kaula's rule and Equation (1).
of interest to calculate the a priori variance of the total lumped coefficient, as its square root can be thought of as a reasonable ( $1 \sigma$ ) value for it. Using Kaula's rule again [over all the coefficients in Equation (1)] this value is indicated by the circular arc centered on the origin in Figure 4.

The convergence of the series in Equation (1) is actually stronger than it appears. The weights can be shown to decrease (on average) at least as fast as $1 / \ell$ while the coefficients behave as $10^{-5} / \ell^{2}$. Comparison of the observed lumped harmonic with its a priori standard deviation shows its value is slightly greater than Kaula's rule would have predicted. In terms of Kaula's rule, the formal statistics say the lumped harmonic is determined to better than 1 part in 40 .

The fields chosen for comparison with the "observed" harmonic all used basically the same shallow 13th order resonant orbits but with different amounts and kinds of tracking data. They all extend to at least the 22nd degree in 13th order terms. The purely satellite data fields are GEM 5 and PGS62 (F. Lerch, Personnal Communication, 1973 and 1974). The others are combination solutions with surface gravimetry data. The Smithsonian fields (Gaposchkin and Lambeck, 1971; Gaposchkin, 1973) used analytic techniques for both orbit and geopotential determination. The satellite data in these consisted only of camera and laser observations. The fields originating at Goddard Space Flight Center employed numerical integration to rationalize the orbital data. GEM 5, 6 and PGS-62 contain significant amounts of electronic data (radar range and range rate, and

Doppler observations) on the 13th order shallow resonant orbits. PGS 11 (F. Lerch, Private Communication, 1974) contains only camera data for the satellite observations.

As for the truncation of 13 th order terms, GEM 5, 6 and SAO SE 2 extend to $(22,13)$, SAO SE 3 goes to $(23,13)$, and PGS 11 and 62 include all 13th order terms to $(29,13)$.

Yet in spite of all these differences, the fields are clustered fairly closely in Figure 4, showing the dominance of the similar 13 th order satellite information. The Goddard fields (with substantially more data) are somewhat closer to the observed harmonic, but the truncation is significantly different for them. In fact the (probable) truncation error alone would account for all of the distance of the GEM 5 and 6 models from the observation. On the other hand the SAO SE 2 (at the same truncation) is significantly farther from the observed value. (It is somewhat beyond the probable truncation error for all the terms of degree greater than 22.) The SAO SE 3 (with 23,13 terms) is marginally closer than SAO SE 2 to the observation but the added terms should have made a much greater improvement still. The same situation holds (and even more strongly) for the more recent Goddard fields (PGS 11 and 62) which extend to $(29,13)$. Here, no improvement is seen over the earlier GEM solutions but the truncation error (given by the dotted circle about the observation) should be much reduced.

The simplest way to interpret these comparisons is to say that the higher degree terms (i.e., above about 21) for this resonance are not well determined.

The conclusion then follows that these terms will be significantly improved with the use of the lumped harmonic (i.e., Equation (2)) for the DIADEME 2 fragment. The simplest example of such use would be to add a single high degree term to an existing field. Of course this would not be a realistic solution on two counts. First, it would upset the 13th order ties (correlation) in the existing field. Secondly, it would ignore the contributions of still higher degree terms. But it does give a representative value (exact for 1967-14F) and a first approximation of the realistic numbers. Using GEM 6 , I find the added $(23,13)$ term $(\mathrm{C}, \mathrm{S})$ to be $10^{-9}(-16.9,0.3)$ which has an rms of $12.0 \times 10^{-9}$ compared to 18.9 x $10^{-9}$ for Kaula's rule.

## SUMMARY AND CONCLUSIONS

A strong 13th order resonance has been observed and analyzed from U.S. Navy Tracking Data on the slowly decaying orbit of a DIADEME 2 fragment (1967-14F). The exact commensurability for the orbit occurred in late 1973 and the major changes due to the resonance will be over by late 1974. Nevertheless, apparently stable and well determined values of a lumped harmonic for this resonance have been found which should significantly improve 13th order geopotential terms to about as high as degree 33. There is fairly close agreement of calculated values from recent fields with the "harmonic" observation (within $20 \%$ ). The major part of the discrepancy is probably due to poorly known coefficients above degree 21 in these fields.

Substantial improvement of 13 th order and high degree terms will be seen with use of the lumped values (and the linear constraint) in combination solutions with other data.

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[^0]:    *Observed - calculated values from converged mean element trajectory

