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### DETERMINATION AND ERROR ANALYSIS OF EMITTANCE AND SPECTRAL EMITTANCE MEASUREMENTS BY REMOTE SENSING

#### R. Kumar\*

Instituto de Pesquisas Espaciais (INPE) Conselho Nacional de Desenvolvimento Científico e Tecnológico (CNPq) 12.200 - S. J. dos Campos, SP, Brasil

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A part of this work was done at the Laboratory for Applications of Remote Sensing (LARS), Purdue University. W. Lafayette, Indiana and sponsored by the National Aeronautics and Space Administration under Grant No. 15-005-112.

\*Research Scientist in the Space Computer Science Division.

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#### Introduction

All objects above absolute zero radiate energy by virtue of their temperature and emittance. At temperatures normally exhibited by natural objects at or near the earth's surface, this radiation is almost entirely in the infrared wavelength region from approximately 4 µm to 100 µm.

Planck's radiation law states that at any given temperature, T, and wavelength,  $\lambda$ , the maximum possible spectral radiance of emitted radiation is the blackbody spectral radiance at that temperature, T, and wavelength,  $\lambda$ . No natural surface emits like a perfect blackbody due to reflection of rays by the surface due to discontinuity in the index of refraction. Therefore the spectral emittance or spectral emissivity of a target is defined as the ratio of the radiation emitted by the target to the radiation emitted by the blackbody at the same temperature and wavelength. The band emittance and emittance of a target are defined similarly except that the radiation is integrated over a certain wavelength interval  $\lambda$  to  $\lambda$  and the entire wavelength region respectively. Thus, no instruments can yield a correct estimate of surface temperature by remote sensing if the emittance of the surface is not taken into account. Gates and

emittance of the surface is not taken into account. Gates and Tantraporn<sup>1</sup> (1952) have pointed out that accurate knowledge concerning the infrared reflectance, absorptance, and emittance of leaves in the 1.0 to 15.0 µm region is essential for a detailed understanding of the energy exchange in the biosphere. Accurate knowledge of the infrared emittance of plants is required in many diverse ecological applications. Most researchers recognized this problem but neglected to apply the needed corrections, arguing that all plant surfaces have a longwave emittance of 0.95 or greater, most leaf emittances being 0.97 to 0.98.

Monteith and Szeicz<sup>2</sup> (1962) and Gates<sup>3</sup> (1963) estimate that assuming the emittance of the plant surfaces equal to one may cause errors of at most  $0.2^{\circ}$ C in the measurement of their temperature. Fuchs and Tanner<sup>4</sup> (1966) show, from a simple calculation, that if the incident thermal radiation from the sky and surroundings were 300 watts per square meter, corresponding to an apparent sky temperature of  $-4^{\circ}$  C, and if the surface temperature were 25° C, a change of emittance from 0.95 to 0.98 would cause a measurement error of 2.2° C. For many detailed investigations, such discrepancies are intolerable. Kumar<sup>5</sup> (1972) has done a thorough review of the general area of "reflection and emission from plants".

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#### Emittance of Leaves

The emittance of a surface depends on its roughness as well as its temperature, in general. Emittance of leaves given by Falckenberg<sup>6</sup> (1928), Gates<sup>7</sup> (1964) and Gates et. al.<sup>8</sup> (1965) range from 0.95 to 0.98. Gubareff et. al.9 (1960), p. 293, indicated the emittance of the leaves to be more than 0.90. Turrell and Austin<sup>10</sup> (1965) reported the emittance of citrus leaves to be about 0.87, using the data of Gates and Tantraporn<sup>1</sup> (1952). Gates and Tantraporn<sup>1</sup> measured the reflectance of upper and lower surfaces of leaves of deciduous trees and herbaceous plants at 3.5, 7.5, 15 and 25  $\mu$ m, at an angle of incidence of 20° and 65°. The infrared radiant source was the Globar whose radiation was reflected off the leaf surface at the desired angle by means of spherical frontsurface mirrors and then focussed upon the entrance slit of the infrared spectrometer. It can be seen from their data the reflectance of most leaves in the infrared wavelength region beyond 2 µm lies between 0 and 10% and its value varies from leaves of one plant to another.

Fuchs and Tanner<sup>4</sup> (1966) found the emittance of single leaves of snap beans and tobacco to be 0.96 and 0.97, respectively. They covered the temperature-controlled anodized surface by the base of a "skewed" aluminum cone with a highly polished reflecting surface. The leaf was placed at the base of the cone, and the apex of the cone was cut to fit the entrance pupil of the infrared thermometer, so that the leaf behaves like a blackbody. Idso et. al.11 (1969) described a method for determining the infrared emittance of individual plant leaves which is relatively simpler than Fuchs and Tanner (F - T) method. Measurements on a wide variety of plant surfaces by this method indicated that significant differences of emittance exist among the various species. The emittance values ranged from about 0.94 to 0.995. They pointed out that the values of emittance, determined by infrared thermometers of different spectral sensitivity, are not uniquely related, thereby emphasizing the importance of emittance measurements.

The emittance of a single leaf is not representative of the emittance of a plant canopy because of the multiple internal reflections resulting from the plant geometry; therefore, the emittance of a plant canopy has to be determined separately<sup>4</sup>. For determining the emittance of a plant canopy, one has to define a meaningful temperature of a plant. For example, the temperature of a plant canopy can be defined as the average of the temperature of its leaves. It is extremely difficult to measure the temperature of the plant surfaces in

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the field conditions because the surface temperature varies quite rapidly, especially when the wind is blowing. Fuchs and Tanner<sup>4</sup> determined the radiation emitted by the vegetal surface. Then, they covered the vegetal surface with a bottomless, hemispherical "pop tent" covered on the inside with aluminum foil so that vegetal surface behaves approximately like a blackbody<sup>12</sup>. They determined the emittance as the ratio of radiant flux density emitted by a vegetal surface to the radiant flux density emitted by it, when it is covered with the pop tent. Covering the vegetal surface by the tent changes its energy balance and thus its surface temperature. During daytime, the tent cuts off the solar radiation and sky radiation. It affects the convective heat exchange and may also modify the transpiration pattern. They found the emittance of dense canopies of alfalfa and sudangrass to be between 0.97 and 0.98 on clear, calm nights when surface temperatures are more steady. The author believes that the emittance of a wide variety of leaves has not been measured carefully, in the natural environment in which the plants grow. The effect of orientation, temperature, moisture content, plant diseases, nutrient deficiency, etc., on the emittance of the leaves should be studied.

#### Emittance of Soils

With the advent of radiation data now available from the infrared measurements made by the satellite, there is an increasing need to know accurately the values of terrestrial surface emittance in order to interpret the data<sup>13</sup>. Most of the investigators (for example, Wark et. al. <sup>14</sup>) have assumed black earth in interpreting data of the satellite. Buettner et. al.<sup>13</sup> (1964) have pointed out that the other data and literature are quite conflicting and confusing. For example, Falckenberg<sup>6</sup> (1928) shows that sand of the Baltic has an emittance of 0.89 at 300° K. While Kruse et.al.<sup>15</sup> (1962) give a value for "terrain of 0.35 (it is not clear here whether this is the emittance of the whole spectrum or just a portion of it, but presumably the entire spectrum). Barnes<sup>16</sup> (1963) lists values of emittance of 0.28 to 0.44 for gravel, plowed field, and granite. Falckenberg<sup>6</sup> (1928) also lists the value for snow emittance as 0.995, while Miller<sup>17</sup> (1963) lists the window value (water vapor window) as 0.35.

Buettner et. al.<sup>13</sup> also reported that infrared signals received in the 8 to 12 µm water vapor window by weather satellites and aircraft are dependent on surface temperature, surface emittance and atmospheric interference. Nowhere can variations of surface emittance be neglected in order to evaluate the correct surface temperatures. They presented three methods of determining surface emittance: (1) reflection data from polished samples ran on a spectrophotometer, (2) a device constructed by Buettner and Kern<sup>13</sup> called an emissivity box, and (3) emittance as inferred from the TIROS satellite data. They gave geological interpretations of the TIROS satellite data along with the values of emittance determined by the above three methods. Emittance of quartz, granite, feldspar, obsidian, basalt, dunite, dolomite, sand, water, etc., ranged from about 0.7 to 0.97.

In the emissivity box (Buettner and Kern (B - K) Method<sup>13</sup>), the radiosity of the soil is first measured when it is exposed to a highly reflecting enclosed surface. The soil is then exposed to a highly emitting surface. The emittance of the soils is determined as the ratio of the two measurements. Because of the large size of the B - K apparatus, this method is used almost exclusively for determining infrared emittance of bare soils and rocks. On the other hand, the method presented by Fuchs and Tanner4 (F - T Method) is well adapted to this use and, in addition, is suitable for measuring the infrared emittance of plant leaves, but it requires that the base of the cone needed to create the blackbody cavity should be larger than the individual leaves of most plants. Fuchs and Tanner18 (1968) recommended that aluminum cones with apex angles of 120° or larger, or shallow cylindrical cavities which have smaller apparent emissivities be used to improve accuracy of the measurements. Idso et. al. 11 (1969) proposed a method for determining emittance of plants and soils which overcomes the dimensional limitations of the methods of Fuchs and Tanner as well as Buettner and Kern and it gives results in good agreement to both of them.

Hovis, Jr.<sup>19</sup> (1966) has given <u>spectral</u> reflectance of some common minerals like carbonates, sulfates, nitrates, salt, silica, etc. from 0.5 to 6 µm. He pointed out that the infrared reflectance spectra of these minerals exhibit spectral absorption band patterns that can be detected in reflection from surface minerals.

Hovis, Jr.<sup>20</sup> (1966) made the spectral reflectance measurements in the 0.5 to 22 µm, from a number of soils including some beach sands and some common surface minerals - calcium carbonate, calcium sulfate or gypsum, sodium chloride, sodium carbonate, and sodium nitrate. He reported that if the measurements are restricted to the windows in the atmosphere of earth, the 10 µm to 12 µm interval seems to be the best choice for radiometry sin e, in this interval, minerals are most uniformly black. Lorenz<sup>21</sup> (1966) found the gray

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reflectances of some natural surfaces -- sand, concrete, coarse gravel, brick roof tile, fine basaltic gravel, asphalt and lawn to be in the range 0.027 to 0.062. Gray emittance can be found from gray reflectance using Kirchhoff's law.

#### Emittance Changes

Gates22 (1965) has pointed out that plants, like animals, must regulate their temperature in order to function at optimum physiological efficiency. This is accomplished through three mechanisms: radiation, transpiration and convection. The question arises: does a leaf adjust its emittance in order to regulate its temperature? The emittance of a leaf can be changed by changing its surface geometry and/or its orientation. A change in environmental conditions can change the leaf moisture content, which in turn changes its surface geometry. The author believes, yet to be confirmed by experiments, that it is unlikely that a leaf adjusts its surface geometry in order to regulate its emittance and hence temperature, because it is much easier for a leaf to change its orientation in order to adjust its heat load and hence temperature. For example, on a hot sunny day, it can change its orientation with respect to direct sunlight to reduce solar radiation absorbed by it.

Conaway and Van Bavel<sup>23</sup> (1966) reported that their radiometrically determined temperatures of the plant canopy could be explained by postulating a widely varying emittance of the plant canopy. They said that the possibility of widely varying emittance of the plant canopy exists since many changes occur in a plant canopy, as factors such as wind speed, solar radiation and water availability vary, but no attempt was made to study these.

#### Error Analysis of Emittance and Spectral Emittance Measurements

Fuchs and Tanner<sup>4</sup> (1966) have discussed the theory of remote sensing of surface temperatures using infrared radiometer, a part of which is summarized here. They made the following assumptions for remote sensing of plants.
1.) Gray emittance in the wavelength range sensed by the radiometer (i.e., 8 µm to 20 µm) or

$$\varepsilon(\lambda) = \varepsilon (8 \ \mu m < \lambda < 20 \ \mu m) \tag{1}$$

2.)  $\varepsilon(\lambda)$  is independent of temperature

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3.) Kirchhoff's Law is valid.

$$(\lambda) = \alpha(\lambda) = \mathbb{1} - \rho(\lambda)$$
(2)  
(transmission = 0 in 8 µm <  $\lambda$  20 µm)

where

#### $\rho(\lambda)$ = the reflectance at wavelength $\lambda$

Under these assumptions, the total apparent outward radiant flux density measured by the radiometer is given by F.

$$F_{o} = \varepsilon f(T) \sigma T^{4} + (1 - \varepsilon) f(T_{S}) F_{S}$$
(3)

radiant flux radiant flux density emitted density reflected by the plant from the plant

T = coptact temperature of the plant . For example, temperature of a plant can be defined as the average temperature of its constituents.

$$f(T) = \int_{0}^{\infty} f(\lambda) F_{\lambda}, \quad b(T) d\lambda / \int_{0}^{\infty} F_{\lambda, b}(T) d\lambda$$
(4)

where

 $F_{\lambda,b}$  (T) = the blackbody radiant flux density at temperature T and wavelength  $\lambda$ 

- $F_{S}$  = environment integrated radiant flux density at temperature  $T_{S}$
- $f(\lambda)$  = filter characteristics for the infrared radiometer (for example, for Barnes infrared radiometer:  $f(\lambda) \neq 0$ , 8 µm  $\leq \lambda \leq 20$  µm,  $f(\lambda) = 0$ , otherwise).

Equation (3) is the fundamental equation describing the quantity of radiation received at the detector of the radiometer. Since the absolute value of f(T) is not important, but only its form, an assumption is now made that f(T) is constant for the range of biological temperatures encountered. The radiometer is calibrated according to the quantity.

$$F_{d} \equiv F_{o} / f(T)$$
(5)

where

E denotes that the quantity is defined.

Hence,

$$f(T_{S}) = \varepsilon \sigma T^{4} + (1 - \varepsilon) \frac{f(T_{S})}{f(T)} F_{S} = \sigma T_{e}^{4}$$

where

# s = full radiator surface temperature as sensed by the radiometer.

The radiometer interprets the contact surface temperature, T, to be full radiator surface temperature  $T_s$ . If  $\varepsilon = 1$ , then  $T_s = T$ . Also, if the integrated full radiator temperature of the surroundings  $(T_s)$  is equal to the contact surface temperature (T), then  $T_s = T$ . If the surface has an emittance substantially less than unity, and if the surface images sun or cold sky, the full radiator surface temperature as measured with the IR radiometer may be quite different from the contact temperature. Hence, there is a need for determining emittance and environment integrated full radiator temperature,  $T_s$ , accurately. Equation (6) can be written as

$$F_{A} = \epsilon \sigma T^{4} + (1 - \epsilon) F_{S}^{*}$$
(7)

where

 $F_{S}^{*} = \frac{f(T_{S})}{f(T)} F_{S}$ 

Fuchs and Tanner assumed that  $F_{X}^{*}$  is a constant (i.e., independent of temperature T), which in turn assumes that f(T)is constant. f(T) = constant is a good approximation for small variations of T. Equation (7) can be rearranged to give

$$= \frac{F_d - F_S^*}{\sigma T^4 - F_S^*}$$
(8)

 $F_d$  and T are experimentally measured values and contain measurement error. Let  $\delta F_d$  and  $\delta T$  be the measurement errors in  $F_d$  and T respectively. Then Eq. (8) can be rewritten as

$$\varepsilon + \delta \varepsilon = \frac{F_d + \delta F_d - F_S^*}{\sigma (T + \delta T)^4 - F_S^*}$$
(9)

where

 $\delta \varepsilon$  = error in the determination of emittance

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(6)

From Eqs. (8) and (9), assuming 
$$\left|\frac{\delta T}{T}\right| < 1$$
,  $\left|\frac{4\delta T}{T} \frac{\sigma T^4}{\sigma T^4} - F_S^{\star}\right| < 1$ 

and neglecting second order and higher orders of  $\frac{\delta T}{T}$  (i.e.,  $\left(\frac{\delta T}{T}\right)^2$ ,  $\left(\frac{\delta T}{T}\right)^3$ , etc.), we get

$$\delta \varepsilon = \delta F_{d} - \frac{4\delta T}{T} \frac{\sigma T^{4} (F_{d} - \frac{*}{S})}{\sigma T^{4} - F_{S}^{*}}$$
(10)

It should be pointed out that Eq. (7) is not an exact equation because of a number of assumptions made in deriving the equation. A complete error analysis should also include another term in  $\delta F_d$  representing the error due to the assumptions made. Equation (10) can be rewritten as

$$\left|\delta\varepsilon\right| = \frac{1}{\left|\sigma T^{4} - F_{S}^{*}\right|} \left|\delta F_{d} - \frac{4\delta T}{T} \frac{\sigma T^{4}}{\left(\sigma T^{4} - F_{S}^{*}\right)} \left(F_{d} - F_{S}^{*}\right)\right| (11)$$

where

or

= denotes the absolute value

$$|\delta\varepsilon| \leq \frac{1}{|\sigma T^4 - F_S^*|} \left[ |\delta F_d| + \frac{\left|\frac{4\delta T\varepsilon}{T}\right|}{|\sigma T^4 - F_S^*|} \right] (using the triangle (12) inequality)$$

It should be pointed out that when the environment integrated radiant flux density,  $F_{S}^{*} = \sigma T^{4}$ , the radiant flux density coming from the target is equal to  $\sigma T^{4}$  (blackbody radiant flux density) irrespective of the emittance of the target. Thus, the radiometer cannot be used to measure the emittance of the target in this particular case.

Using Eq. (12),  $|\delta\varepsilon|$  is plotted against  $F_S^*$  with the parameters  $\varepsilon$ , t,  $|\delta T|$  and  $|\delta F_d|$  in Figs. 1 to 4. It is clear from Figs. 1 to 4 that the upper bound of  $|\delta\varepsilon|$  (absolute error in the determination of emittance) increases with the increase in  $F_S^*$ . Thus, for accurate determination of emittance of the plant surfaces in the field, the experiment should be conducted when the sky is relatively clear to have a small value of  $F_S^*$  and hence small  $|\delta\varepsilon|$ . Fig. 1 shows that the emittance of



emittance for the thermal channels which have been used in multispectral scanners of Environmental Research Institute of Michigan, formerly called Willow Run Laboratory, University of Michigan (4.5 to 5.5 µm, 8 to 13.5 µm and 9.3 to 11.7 µm), and Skylab (10.2 to 12.5 µm). Fig. 5 shows that for a given value of the emittance, the difference between temperature and band radiance temperature is least in the 4.5

the target, say a leaf, has a very little influence on  $|\delta\varepsilon|$ .

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Determination of Temperature and Spectral Emittance

It is quite important to know the temperature of the natural target -- say plants -- to detect the subtle changes in their temperature due to stresses by insects, plant diseases, physiological disorders, nutrient deficiency and adverse environment effects. Optimum temperature exists for certain biological activity in the organisms. It is also necessary to know the temperature of a target for energy balance calculations.

#### (A) Spectral Radiance Temperature and Average Temperature

If the temperature of a natural target is not constant over the whole target, its meaningful temperature can be defined as the average temperature of the target over its entire area. It was shown by Kumar and Silva<sup>24</sup> that, for a target having a non uniform temperature distribution and emittance = 1, the difference between the spectral radiance temperature of a target, as measured by the spectroradiometer, and its average temperature is rather small and within the accuracy of the spectroradiometer.

#### (A) Spectral Emittance and Temperature Determination

Let a spectroradiometer whose radiance and wavelength scales have been calibrated, record the spectral radiance coming from the natural target at wavelengths  $\lambda_1$ ,  $\lambda_2$  ...  $\lambda_N$ . Let the target be opaque (i.e., transmission = 0) in the wavelength region  $\lambda_1$  to  $\lambda_N$ . Throughout this section, i will refer to integers 1, 2, ... N; j and k will refer to integers lying between 1 and N. The basic assumption made in this analysis is that the reflected radiation from the target is negligible as compared to the radiation emitted from it. If the sky is clear, this is true for the most natural targets for wavelength longer than approximately 4.6  $\mu m$ <sup>25</sup>. Neglecting interference by atmosphere and using Planck's Law,

$$T_{g}(\lambda_{i}) = \frac{hc}{k\lambda_{i}} \qquad \frac{1}{\log_{e} \left\{1 + \frac{2hc^{2}}{\lambda_{i}^{5}L_{\lambda_{i}}}\right\}}$$
(13)

i = 1, 2, ... N

where  $T_{a}(\lambda_{i}) = \text{spectral radiance temperature at wavelength } \lambda_{i}$ 

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 $L_{\lambda}$  = spectral radiance of radiation emitted by the target at wavelength  $\lambda_i$ 

The temperature of a target, for applications to spectroradiometric measurements, 's defined using Planck's Law, as follows.

$$T = \frac{hc}{k\lambda_{i}} \frac{1}{\log_{e} \left[1 + \frac{(2hc^{2} \epsilon (\lambda_{i}))}{\lambda^{5}L_{\lambda_{i}}}\right]}$$
(14)

It should be pointed out that in actual practice, the value of temperature T, at the wavelengths  $\lambda_1, \lambda_2 \dots \lambda_N$  may not be equal because of the experimental errors involved. Throughout this Section, the word "temperature" of a natural target will be used to mean temperature, T, defined by eq. (14)

There are N equations in eq. (13) and N unknowns,  $T_i(\lambda_i)$ , which can be determined.

$$L_{\lambda_{i},b}(T_{s}(\lambda_{i})) = L_{\lambda_{i}}(T) = \varepsilon(\lambda_{i}) L_{\lambda_{i},b}(T)$$
(15)

where

$$L_{\lambda_{i}, b} \begin{pmatrix} T_{s}(\lambda_{i}) \end{pmatrix} = blackbody spectral radiance at wavelength \lambda_{i} and temperature T_{s}(\lambda_{i}) \\ \vdots \\ \varepsilon (\lambda_{i}) = spectral emittance of the target at wave-$$

From Eqs. (14) and (15), one obtains

length

$$T_{s}(\lambda_{i}) \leq T, T_{s}(\lambda_{i}) = T \text{ if and only if } \epsilon(\lambda_{i}) = 1$$
 (16)

There are N equations in eq. (15) and N + 1 unknowns -- $\varepsilon(\lambda_i)$ , i=1,2,... N and T. Therefore, Eq. (15) cannot be solved, in general, i.e., the temperature, T, of the target cannot be determined without any knowledge about its spectral emittance.

Ai

The variables of eq. (15) are plotted vs. wavelength in Figs. 6 to 11. Although the wavelength range in these Figures is taken to be 0.5 to 14 µm for illustrative purposes, we are concerned only with the wavelength range about 4 to 14 µm in this Section. Figures 6 and 7 show that the difference between the temperature, T, and spectral radiance temperature,

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EMITTANCE AND SPECTRAL EMITTANCE



 $T_{g}(\lambda)$ , increases almost linearly as the wavelength increases for a given value of spectral emittance. Figure 11 shows that  $[T - T_{s}(\lambda)]/\lambda$  vs. wavelength is almost a constant for a given value of spectral emittance of about 0.8 to 1. To find a good approximation of T, one minimizes  $[T - T_{s}(\lambda)]$  over the wavelength range of the spectroradiometer. If the target were a gray body (i.e., spectral emittance is independent of wavelength), min  $[T - T_{s}(\lambda)]$  will be found at the shortest wavelength of the wavelength range of the spectroradiometer  $-\lambda_1$ . In actual practice, however, the natural target may not be a gray body 1 in the wavelength range 4 to 14 µm<sup>1</sup>, 19, 20. Therefore, in general, min  $[T - T_{s}(\lambda)]$  can lie anywhere between  $\lambda_1$ and  $\lambda_N$  depending on spectral emittance and spectral radiance temperature of the target. Figures 8 and 9 show respectively that the accuracy of the experimentally determined spectral emittance of a target using a spectroradiometer increases with



increasing wavelengths; and the accuracy is weakly dependent upon the temperature of the target. As pointed out earlier, eq. (15) cannot be solved in general. However, the solutions of some of the special cases of eq. (15) are discussed as follows.

(a) If the target is a blackbody at one or more wavelengths  $\lambda_j$ ,  $1 \le j \le N$ , one gets the CALL Line wavelengths  $\lambda_j$ .

$$T = T_{s} (\lambda_{j}) = \max_{\lambda_{i}} [T_{s} (\lambda_{i})], i = 1, 2, ... N$$
 (17)

(b) The spectral emittance at some wavelength  $\lambda_j$  is a known function of the spectral emittance at some other wavelength  $\lambda_k$ , i.e.,

$$\varepsilon (\lambda_{j}) = f \left[\varepsilon(\lambda_{k})\right]$$
(18)

Equations (15) and (18) have (N+1) equations and (N+1) unknowns  $-\epsilon$  ( $\lambda_1$ ), i = 1, 2, ... N and T. Thus, the unknowns can be calculated.

It should be pointed out that if  $T_s(\lambda) = T = constant$ (i.e., independent of wavelength), it does not imply that the target is necessarily a blackbody, for its spectral emittance,  $\varepsilon(\lambda)$ , could be given by

$$e^{hc/\lambda kT} - 1) / (e^{hc/\lambda kT'} - 1)$$
(19)

In actual practice, however, it is very unlikely that the spectral emittance of a natural target is given by eq. (19). In actual practice, if  $T_{s}(\lambda) = \text{constant } \pm e$ , in the wave-



one could say that the target is probably almost a blackbody with spectral emittance greater than approximately 0.98; where e = average of absolute value of the error involved in determination of  $T_s$  ( $\lambda$ ) (Figure 7). A typical value of e is 0.5° C.

(c) Let the upper and/or lower bounds on the spectral emittance of the target be known in some known wavelength intervals  $\lambda_j \leq \lambda \leq \lambda_k$ . These bounds may be determined from the available data of spectral emittance in the literature and/or researcher's own experience and knowledge.

(i) Upper Bounds

min. 
$$\varepsilon$$
  $(\lambda_i) \leq m_1$ , max.  $\varepsilon$   $(\lambda_i) \leq m_2$  (20)

Let

U

max.  $T_s(\lambda_i) = T_i$  (found from spectroradiometric data) (21)

or

$$T \ge T_1$$
 (eq. (16)) (22)

Let { } denote the assumed value. Assume  $\{T\}_1 = T_1$ . Determine values of { $\epsilon$  ( $\lambda_1$ ) $\}_1$  using eq. 15 and check if min.  $\epsilon(\lambda_1) \leq m_1$ . If not, then continue taking  $\{T\}_2 = T_1 + \Delta T$ ,  $\{T\}_2 = T_1 + 2\Delta T$ ... until at some {T} = T<sub>2</sub>,min.  $\epsilon$  ( $\lambda_1$ ) =  $m_1$  within certain prechosen suitable accuracy. Here  $\Delta T$  is some suitable increment in temperature. For example, one can take  $\Delta T$  as the average of absolute error involved in the determination of spectral radiance temperature by spectroradiometer. Similarly, one can determine  $T_3$  corresponding to the condition max.  $\epsilon$  ( $\lambda_1$ ) =  $m_2$ . Let

max.  $T_s(\lambda_i) = T_1$  (known value)

max. 
$$[T_1, T_2, T_3] = T_{h, 1, b}$$
 (23)

where

 $T_{h.1.b.}$  = high lower bound of the temperature T

If any other lower bound on T is known, it should also be included in eq. (23) to determine  $T_{h,1,b}$ .

$$c(\lambda_{i})_{1.u.b.} = \frac{L_{\lambda_{i},b}}{L_{\lambda_{i},b}} (\lambda_{i}, T_{s})$$
(24)

Where

 $\epsilon (\lambda_i)$  1.u.b. = low upper bound on the spectral emittance  $\epsilon (\lambda_i)$ 

(ii) Lower Bounds

The same procedure used in "Upper Bounds" can be followed to determine  $T_{1.u.b.}$  and  $\varepsilon (\lambda_i)_{h.1.b.}$  from the known lower bounds on spectral emittance.

Let the wavelengths at which min.  $\{\varepsilon(\lambda_i)_{1.u}\}$ , min.  $\{\varepsilon(\lambda_i)_{h.1.b}\}$ , max.  $\{\varepsilon(\lambda_i)_{1.u.b}\}$  and max.  $\{\varepsilon(\lambda_i)_{1.u.b}\}$ , minimized/maximized over the wavelength range  $\lambda_1$  to  $\lambda_N$  occur be  $\lambda_a$ ,  $\lambda_b$ ,  $\lambda_c$  and  $\lambda_d$ , respectively. Then in.  $\{\varepsilon(\lambda_i)\}$  lies between  $\lambda_a$  and  $\lambda_b$  and max.  $\{\varepsilon(\lambda_i)\}$  lies be weight  $\lambda_c$  and  $\lambda_d$  respectively. If  $T_s(\lambda_j) \ge T_s(\lambda_k)$  for all  $\lambda_j \ge \lambda_R$ , where  $1 \ge j \ge N$ ,  $1 \le k \le N$ , it implies that min.  $\varepsilon(\lambda_i)$  and max  $\varepsilon(\lambda_i)$  would occur at wavelengths  $\lambda_a = \lambda_b$ , and  $\lambda_c = \lambda_d$ , respectively.

The author emphasizes that for most of the plant canopies with good ground cover (say, ground cover > 0.70), one can take  $T_{h,1,b}$  as the approximation to the temperature, T, of the plant defined by Eq. (14) because the spectral emittance of most plant canopies with good ground cover probably lies between about 0.95 and 1.00 in 7 to 14 µm wavelength range. If no more information on the spectral emittance of the target is available other than upper and/or lower bounds of its spectral emittance which have been used to calculate  $T_{h,1,b}$ , and  $T_{1,u,b}$ , respectively, one can take T (estimated) =  $(T_{h,1,b}, + T_{1,u,b})/2$ .

One practical way of determining the temperature of a natural target — say a plant canopy is to take spectroradiometric data on the plant canopy at two different times when the sky is clear (cloud free) so that the reflected radiation from the target is negligible as compared to the radiation emitted by it. These times should be selected close enough so that the geometry (orientation of the leaves, stems, etc. percent ground cover) of the plant canopy can be assumed to be practically the same at both times. Wind should be steady at both of these times so that the wind does not change the geometry of the plant. Also, these two times should

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#### EMITTANCE AND SPECTRAL GREEPANCE

be such that the temperature of the plant at one time is likely to be significantly (at least 20 C) different from ita temperature at the other time. This can be approximately checked by making a few measurements of contact temperature of the leaves directly visible from the spectroradiometer by theraistor. Now, an assumption can be made that the emittance of the plant canopy is equal at both of these times since its geometry is practically the same at both times. This gives one additional equation (eq. (18)) and hence temperature, T, of the plant canopy can be determined. However, if the temperature of the plant canopy at one time is close to its temperature at another time within the experimental accuracy, then the spectral radiance of the plant canopy recorded by spectroradioacter will be equal at two times within experimental accuracy, i.e., two equations of eq. (15) are the same within experimental accuracy. So, there are only N independent equations in (15) and N+1 unknowns  $\neg = \varepsilon(\lambda_{+}), i = 1, 2, \dots$  N, and T. Temperature, T, cannot be determined without any knowledge of the spectral emittance of the target.

The same procedure can be followed to determine the temperature of a plant canopy by taking spectroradiometric data on two plant canopies of the same crop, maturity, geometry (orientation of leaves, stems, etc.) soil background, percent ground cover etc. These plant canopies can be assumed to have equal spectral emittance in the wavelength range of the spectroradiometer. Thus if their temperatures are different, one can determine their emittance and temperature.

To create an illustration, previously acquired calibrated data of the Exotech Model 20 C spectroradiometer in the wavelength range 2.8 to 5.6 µm on Russell Silt Loam Soil were available<sup>24</sup>. Maximum value of spectral radiance temperature from this data was found to be 42° C at 3.64 µm (eq. (21)). So, the contact temperature of the soil should be at last 42° C (eq. (22)). Assuming contact temperature of the soil to be 42° C, its spectral emittance was calculated avoiding the wavelength region close to the 4.3 µm CO, band and is shown in Table I. Obviously spectral emittance at 3.64 µm = 1. Note that Table I gives the upper bound of the spectral emittance.

#### Table 1 Upper Bound of Spectral Emittance

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### EMITTANCE AND SPECTRAL EMITTANCE

be such that the temperature of the plant at one time is likely to be significantly (at least 2º C) different from its temperature at the other time. This can be approximately checked by making a few measurements of contact temperature of the leaves directly visible from the spectroradiometer by thermistor. Now, an assumption can be made that the emittance of the plant canopy is equal at both of these times since its geometry is practically the same at both times. This gives one additional equation (eq. (18)) and hence temperature, T, of the plant canopy can be determined. However, if the temperature of the plant canopy at one time is close to its temperature at another time within the experimental accuracy, then the spectral radiance of the plant canopy recorded by spectroradiometer will be equal at two times within experimental accuracy, i.e., two equations of eq. (15) are the same within experimental accuracy. So, there are only N independent equations in (15) and N+1 unknowns --  $c(\lambda_i)$ , i = 1, 2, ... N, and T. Temperature, T, cannot be determined without any knowledge of the spectral emittance of the target.

The same procedure can be followed to determine the temperature of a plant canopy by taking spectroradiometric data on two plant canopies of the same crop, maturity, geometry (orientation of leaves, stems, etc.) soil background, percent ground cover etc. These plant canopies can be assumed to have equal spectral emittance in the wavelength range of the spectroradiometer. Thus if their temperatures are different, one can determine their emittance and temperature.

To create an illustration, previously acquired calibrated data of the Exotech Model 20 C spectroradiometer in the wavelength range 2.8 to 5.6  $\mu$ m on Russell Silt Loam Soil were available<sup>24</sup>. Maximum value of spectral radiance temperature from this data was found to be 42° C at 3.64  $\mu$ m (eq. (21)). So, the contact temperature of the soil should be at last 42° C (eq. (22)). Assuming contact temperature of the soil to be 42° C, its spectral emittance was calculated avoiding the wavelength region close to the 4.3  $\mu$ m CO, band and is shown in Table 1. Obviously spectral emittance at 3.64  $\mu$ m = 1. Note that Table I gives the upper bound of the spectral emittance.

## Table 1 Upper Bound of Spectral Emittance

Wavelength in µm 3.4 3.6 3.8 4.6 4.8 5.0 5.5

Upper Bound of Spectral Emittance 0.73 0.8

0.73 0.82 0.81 0.71 0.72 0.74 0.75