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Lunar Tidal Acceleration Obtained From Satellite-Derived Ocean Tide Parameters

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INTRODUCTION

Lambeck [1975] gives the ocean tide, ξ , as

$$\xi_{\mu} = \sum_{s=0}^{\infty} \sum_{t=0}^s P_{st}(\sin\phi).$$

$$(C_{st}^{\pm})_{\mu} \sin[\sigma(\tau) \pm t\lambda + (\epsilon_{st}^{\pm})_{\mu}]$$

where the $(C_{st}^{\pm})_{\mu}$ and $(\epsilon_{st}^{\pm})_{\mu}$ are functions of the spherical harmonic expansions of the μ tide component, and $\sigma(\tau)$ represents a linear combination of the solar and lunar angular quantities [Goad, 1977]. Lambeck [1975] also gives for the rate of change of the semi-major axis of the moon due to ocean tides

$$\dot{a} = 2 k'_{stuv} (s-2u+v) \begin{matrix} (s-t) \text{ even} \\ \left[\begin{matrix} \cos \\ \sin \end{matrix} \right] (\epsilon_{st}^{\pm})_{\mu} \\ (s-t) \text{ odd} \end{matrix},$$

where k'_{stuv} is a function of the orbital semi-major axis, inclination, eccentricity, Earth and lunar masses, mean density of the Earth and oceans, the load deformation coefficient, and particularly the coefficient C_{st}^+ as a linear factor. From elementary celestial mechanics, the secular change in the mean motion, n , of the moon is given by,

$$\dot{n} = -\frac{3}{2} \frac{n}{a} \dot{a}.$$

The critical factor in computing the tidal acceleration \dot{n} of the moon is shown to be the quantity

$$C_{22}^+ \cos \epsilon_{22}^+$$

and Lambeck [1975] notes that for three different global M_2 ocean tide models this quantity varies by only 10% from the mean of the three. Taking this mean value for the dominant M_2 tide effect, the results by Dietrich for the O_1 tide, and estimating the N_2 tide as proportional to the corresponding tide raising M_2 and N_2 potentials, Lambeck [1975] obtains the estimate

$$\dot{n} = -35 \pm 4 \text{ arc-sec}/(100 \text{ yr})^2$$

for the lunar tidal acceleration if it is assumed that there is no contribution from a solid tide phase lag. He further notes that astronomically obtained values at the time of his publication ranged from -37 to -52 arc-sec/(100 yr)². Subsequently, the values estimated from astronomical data $\dot{n}_t = -27.2 \pm 1.7$ and $\dot{n}_t = -26 \pm 2$ have been obtained by Muller [1976] and Morrison and Ward [1976], respectively.

Artificial satellites can help resolve this problem because results obtained from them are independent of lunar or planetary data and do not require knowledge of any tidal mechanisms. However, until very recently there were no reliable estimates of M_2 ocean tide parameters from satellite orbit perturbations.

Previously, we published a constraint equation on the M_2 ocean tide coefficients that was in reasonable agreement with numerical ocean tide models [Goad and Douglas, 1977]. In the present paper we obtain additional equations from an analysis of the evolution of the orbit of GEOS-3 that enable an explicit solution for $C_{22}^+ \cos \epsilon_{22}^+$ for the M_2 tide and hence an estimate for the tidal acceleration of the moon. Our value of (-27.6 arc-sec/(100 yr)²) is in agreement with the recent solution of Muller [1976] and Morrison and Ward [1975].

THE ORBIT OF GEOS-3

Table 1 gives the orbital specifications of GEOS-3. Osculating orbit elements for 100 successive two-day arcs were obtained from the Doppler tracking done in support of the GEOS-3 program by the Naval Surface Weapons Center (NSWC), Dahlgren, Virginia. The mean elements corresponding to the osculating elements are given in Table 2. All long periodic and secular perturbations are present in these elements. We have analyzed extensively only the mean inclination and node. The precision of the mean inclination is about 0".02. The mean node is less accurate (about 0".04) primarily due to the uncertainty of the prediction of UT1 time used in the data reductions.

Table 3 gives the perturbation on the inclination and node of GEOS-3 computed from our solution for the $(2,2)^+$ and $(4,2)^+$ coefficients. A detailed discussion of the tidal perturbations on satellite orbits is given by Lambeck et al., [1974]. We have used the results of that work for our own computations with the exception of eq. 6, p. 424 for the perturbation of the node. We are unable to verify this equation. Our application of the differential equation for the time rate of change of the node to the tidal potential developed by Lambeck et al. [1974] is

$$\Delta\Omega_{n\ell mpq}^{\pm} = \frac{4\pi(1+k'_{\ell})}{2\ell+1} G a_e \left(\frac{a_e}{a}\right)^{\ell+1} \frac{\rho_w}{Na^2(1-e^2)^{\frac{1}{2}} \sin i} C_{n\ell m}^{\pm} G_{\ell pq}(e)$$

$$\dot{\gamma}_{\pm} \left\{ \frac{\partial F_{\ell mp}(i)}{\partial i} - \frac{3}{2} N \left(\frac{a_e}{a}\right)^2 \frac{C_{20} \sin i}{(1-e^2)^2} F_{\ell mp}(i) \frac{[(\ell-2p) \cos i - m]}{\dot{\gamma}_{\pm}} \right\}$$

$$\begin{bmatrix} + \cos \\ - \sin \end{bmatrix} \begin{matrix} \ell-m \text{ even} \\ \gamma_{n\ell mpq}^{\pm} \\ \ell-m \text{ odd} \end{matrix} .$$

Table 1. GEOS-3 Orbital Characteristics

Epoch MJD 42525.0	April 23, 1975 0 hrs.
\bar{a} 7219 km	
\bar{e} .0005	
\bar{i} 114°99	
Perigee Altitude	837 km
M ₂ Tidal Period	17. ^d 2

Table 3. Amplitude and phase of inclination and node perturbations due to the M_2 ocean tide on GEOS-3 satellite.

Source	Inclination		Node	
	Amplitude	Phase	Amplitude	Phase
Observed	0".040	327°	0".027	291°
Pekeris and Accad*	0".051	339°	0".030	352°
Hendershott (Model 1)*	0.061	317°	0".030	286°
Hendershott (Model 2)*	0".065	276°	0".025	244°
Bogdanov and Magarik*	0".050	328°	0.049	290°

* As reported by Lambeck et al. [1974].

The difference appears in the indirect term arising from the interaction with the Earth's oblateness [Kaula, 1966, p. 49].

Note in Table 3 that the ocean tide effects on GEOS-3 are rather different on the inclination and node. The perturbation on the inclination comes almost entirely from the (2,2) term and that of the node from the (4,2) term. Using the methods employed in our previous paper [Goad and Douglas, 1977] to remove known perturbations, we obtain the observation equations for the M_2 ocean tides:

Inclination:

$$\begin{aligned} & (3''.99 \pm 0.4) \times 10^{-2} \sin [\sigma(\tau) + 327^\circ \pm 4^\circ] \\ & = \frac{1''.26}{\text{cm}} \times 10^{-2} C_{22}^+ \sin [\sigma(\tau) + \varepsilon_{22}^+] - \frac{0''.32}{\text{cm}} \times 10^{-2} C_{42}^+ \sin [\sigma(\tau) + \varepsilon_{42}^+] + \dots \end{aligned}$$

Node:

$$\begin{aligned} & (2''.73 \pm 0.7) \times 10^{-2} \cos [\sigma(\tau) + 291^\circ \pm 13^\circ] \\ & = - \frac{0''.24}{\text{cm}} \times 10^{-2} C_{22}^+ \cos [\sigma(\tau) + \varepsilon_{22}^+] - \frac{3.38}{\text{cm}} \times 10^{-2} C_{42}^+ \cos [\sigma(\tau) + \varepsilon_{42}^+] + \dots \end{aligned}$$

These equations can be used for a GEOS-3 solution alone for $(2,2)^+$ and $(4,2)^+$ quantities because the node and inclination on GEOS-3 have negligible correlation [Anderle, 1977, private communication]. However, a combined solution with 1967-92A yields more precise results. To demonstrate the precision of the GEOS-3 inclination data, we have prepared Figure 1 to show the observed and calculated values of the M_2 ocean tide perturbation of the inclination of GEOS-3.

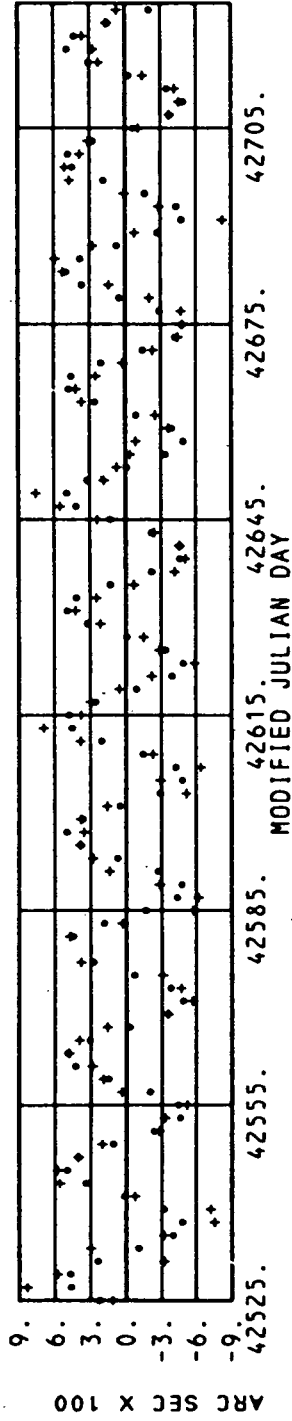


Figure 1. Observed and Calculated Inclination Values
at M₂ Frequency for GEOS-3

- - calculated
- + - observed

COMPUTATION OF TIDAL PARAMETERS AND THE LUNAR TIDAL ACCELERATION

As shown by Lambeck et al. [1974], computation of ocean tide parameters from satellite orbit perturbations requires assumption of values for the solid tide Love number k_2 and the lag angle. Table 4 gives the results for ocean tidal parameters for various values of the solid tide phase lag with k_2 assumed to be 0.30, a value often quoted from seismic measurements.

Note in Table 4 that increasing the solid phase lag results in a large reduction in the value of C_{22}^+ , until at 1° the value of either the GEOS-3 or the combined GEOS-3/1967-92A solution falls to less than 1/2 of the value obtained by Hendershott (quoted in Lambeck et al. [1974]). Thus, although solid and ocean tide effects on satellite orbits cannot be separated, assumption of a solid tide phase lag as large as 1° leads to unacceptable values for the ocean tide coefficients with $k_2 = 0.30$.

Using the combined GEOS-3/1967-92A results for the M_2 tide coefficients in the equations for \dot{a} , \dot{n} and using the values given by Lambeck [1975] for the O_1 and N_2 values, we obtain the value $\dot{n} = -27.6 \pm 3$ arc sec/(100 yr)². Muller [1976] obtained -27.2 ± 1.7 ; and Morrison and Ward [1976] obtained -26 ± 2 . All of these values are outside the range of -35 to -52 arc sec/(100 yr)² reported in Lambeck [1975].

Table 4. M_2 Tidal Estimates Obtained from GEOS-3 and Combined GEOS-3/1967-92A Orbit Perturbations for Various Solid Tide Phase Lags and $k_2 = .30$

Phase Lag (assumed)	C_{22}^+		ϵ_{22}^+		C_{42}^+		ϵ_{42}^+	
	Combined	GEOS-3	Combined	GEOS-3	Combined	GEOS-3	Combined	GEOS-3
0	3.23	2.86	331°	332°	0.87	1.05	113°	115°
0.5	2.76	2.39	325°	326°	0.87	1.05	113°	115°
1.0	2.32	1.95	318°	317°	0.87	1.05	113°	116°

DISCUSSION

The formal uncertainty of our solution for the M_2 tide parameters corresponds to ± 1.8 arc sec/(100 yr)² in \dot{n}_t . However, the O_1 and N_2 tides also contribute about -4.4 arc sec/(100 yr)² total lunar tidal acceleration. Because of the very approximate way in which parameters for these latter tidal components were estimated by Lambeck [1975], an uncertainty of perhaps 25% in the O_1 and N_2 contributions is conceivable, leading to an additional source of uncertainty of at least one arc sec/(100 yr)². Therefore, we estimate the total uncertainty of our value of the tidal \dot{n} for the moon to be ± 3 arc sec/(100 yr)².

Another significant matter is concerned with the effect of a non-zero but small solid tide phase lag δ_2 . Again assuming that the Love number k_2 has the value of 0.30, the total solid and fluid lunar tidal acceleration is given by

$$\dot{n}_{\text{total}} = -1040 k_2 \sin(2\delta_2) - 8.32 C_{22}^+ \cos \epsilon_{22}^+ - 4.4(N_1 + O_2)$$

arc sec/(100 yr)² where the last term is that given by Lambeck [1975]. However, the effect of a non-zero value of δ_2 in a satellite solution for ocean tide parameters results in a compensating change in $C_{22}^+ \cos \epsilon_{22}^+$ that maintains a nearly constant value for the total \dot{n} . For example, assumption of $\delta_2 = 0.5^\circ$ changes the value of the total solid/fluid \dot{n} by only 1 arc sec/(100 yr)². Thus our value for the lunar tidal acceleration is insensitive to any small future adjustments of the value of δ_2 .

Finally, some consideration is required of the method used by Lambeck to obtain the 2nd degree spherical harmonic coefficients from numerical ocean tide models. Lambeck [1975] notes that the three M_2 tide models he used to estimate the lunar \dot{n} vary by 10% about their mean value for $C_{22}^+ \cos \epsilon_{22}^+$. However, his procedure for obtaining the coefficients can introduce considerable uncertainty into the process. For two of the models he obtained values of

the tidal amplitude and phase at a 10° grid interval and then expanded the values in a series of harmonics. We decided to test the effect of grid-size on the coefficient estimates. E. Schwiderski (NSWC), provided us with the M_2 tidal phase and amplitude at each 1° of latitude and longitude based upon his latest numerical solution of the Laplace tidal equations. This model is being used in the analysis of satellite altimeter data at the NSWC (Schwiderski, 1977). Table 5 shows that the value of the critical $C_{22}^+ \cos \epsilon_{22}^+$ term does depend on the grid interval by (in this case) an amount corresponding to as much as $5.4 \text{ arc sec}/(100 \text{ yr})^2$. Thus it is critical in future analyses of numerical ocean tide models to compute spherical harmonic coefficients from data gridded at much finer than 10° intervals. It is also worth noting that the \dot{n}_t obtained with zero solid phase lag and the Schwiderski ocean tide parameters is $-28.9 \text{ arc sec}/(100 \text{ yr})^2$. No information is available to estimate the uncertainty of this result.

In spite of new results for the lunar tidal accelerations smaller in absolute value than $\dot{n}_t = -30 \text{ arc sec}/(100 \text{ yr})^2$, the source of this acceleration can be accounted for by the ocean tides alone. The $(2,2)^+$ term in the expansion of the ocean tides is equivalent to an equatorial bulge that leads the moon by 55° so that even though the amplitude is small compared to solid tide, the torque produced is large.

The results from perturbations of artificial satellites are especially satisfying because no knowledge of tidal mechanisms is required to obtain the principal component of the expansion of the ocean tides in spherical harmonics that is responsible for the lunar tidal acceleration. New, accurate numerical tide models for the O_1 and N_2 tides would be useful in further refining the results presented here.

Table 5. Effect of Grid Size on Estimates of Tidal Parameters

Grid Interval, deg.	C_{22}^+	ϵ_{22}^+	$C_{22}^+ \cos \epsilon_{22}^+$
10 x 10	3.26	325	2.67
8 x 8	3.81	331	3.33
5 x 5	3.38	325	2.76
3 x 3	3.65	325	2.99
1 x 1	3.59	326	2.98

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<p>16. Abstract Analysis of 100 sets of mean elements of GEOS-3 computed at 2-day intervals had yielded observation equations for the M_2 ocean tide from the long periodic variations of the inclination and node of the orbit. If the 2nd degree Love number is given the value $k_2 = 0.30$ and the solid tide phase angle is taken to be 0°, the values are</p> <p>Inclination: $(3^{\circ}99 \pm 0.4) \times 10^{-2} \sin [\sigma(\tau) + 327^\circ \pm 4^\circ]$ $= \frac{1^{\circ}26}{\text{cm}} \times 10^{-2} C_{22}^+ \sin [\sigma(\tau) + \epsilon_{22}^+] - \frac{0^{\circ}32}{\text{cm}} \times 10^{-2} C_{42}^+ \sin [\sigma(\tau) + \epsilon_{42}^+] + \dots$</p> <p>Node: $(2^{\circ}73 \pm 0.7) \times 10^{-2} \cos [\sigma(\tau) +]91^\circ \pm 13^\circ]$ $= - \frac{0^{\circ}24}{\text{cm}} \times 10^{-2} C_{22}^+ \cos [\sigma(\tau) + \epsilon_{22}^+] - \frac{3^{\circ}38}{\text{cm}} \times 10^{-2} C_{42}^+ \cos [\sigma(\tau) + \epsilon_{42}^+] + \dots$</p> <p>where $\sigma(\tau) + 2\Omega - 2M^* - 2\omega^* - 2\Omega^*$ and the starred quantities are lunar orbit elements. Combining these equations with the result obtained by Goad and Douglas [1977] for the satellite 1967-92A gives the M_2 ocean tide parameter values</p> $C_{22}^+ = 3.23 \pm .25 \text{ cm}, \epsilon_{22}^+ = 331^\circ \pm 6^\circ$ $C_{42}^+ = .87 \pm .19 \text{ cm}, \epsilon_{42}^+ = 113^\circ \pm 6^\circ$ <p>Under the same assumption of zero solid tide phase lag, the lunar tidal acceleration is mostly (85%) due to the C_{22} term in the expansion of the M_2 tide with additional small contributions from the O_1 and N_2 tides. Using Lambeck's [1975] estimates for the latter we obtain for the tidal acceleration in lunar longitude the value $\dot{n}_t = -27.6 \pm 3 \text{ arc sec}/(100 \text{ yr})^2$, in excellent agreement with the most recent determinations from ancient and modern astronomical data.</p> <p>The mean elements of GEOS-3 are also presented in tabular form.</p>			
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