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Documentation of the Fourth Order Band Model

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Fourth Order Band Model

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November 1979

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Preface

We have decided to compile a preliminary documentation of the new GLAS Fourth Order General Circulation Model. The present documentation has not been subjected to a careful editing process; we hope that its possible usefulness will compensate for some of its defects. The model dynamics (COMP0, COMP1 and COMP2) is still undergoing minor improvements, especially in the time differencing scheme which we hope will improve its efficiency. The "physics" routine (COMP3) has not been documented because it is being thoroughly revised. The present version of COMP3, similar to the one used in the GLAS/GISS models (see the documentation by Tsang and Karn), with modifications introduced by Y. Sud (1979) is included in the code. Criticisms and suggestions for improvement will be greatly appreciated, since a final documentation will be prepared in 1980.

We are very grateful to all the people that have helped us generously. In particular, Dr. N. Rushfield had a major impact in the process of making the model operational. W. Connelly, D. Edelmann, D. Han, S. Breining, P. Anolick and M. Almeida were very helpful in the development of the model. The documentation was expertly and cheerfully typed by S. Mathis; D. Edelmann and D. Rosen have also cooperated in its compilation. We want to express our special gratitude toward Dr. Y. Sud, who offered us generously both his advice and his time in the development of the "physics" routine, and most especially to Dr. M. Halem without whose many useful suggestions, constant encouragement, and long patience, this work would not have been finished.

Eugenia Kalnay-Rivas

November 1979

I. Introduction

The band fourth order model is a GCM which uses quadratically conservative, fourth order horizontal space differences on an unstaggered grid and second order vertical space differences with a Matsuno (forward-backward) or a smooth leapfrog time scheme to solve the primitive equations of motion.

This program numerically solves these equations one latitude band at a time which greatly reduces the amount of computer core storage needed to run the program. It also uses the same variable names, order of computations, I/O, post-processing as the standard second order GCM. Appropriate modifications have been made for the fourth order differences and leapfrog scheme. (See the 1978 Goddard Modeling and Simulation Research Review for an overview of the fourth order band model.)

The main feature of this model is that fourth order approximations are used for all the horizontal derivatives. The derivative $\frac{\partial q}{\partial x}$ is approximated by

$$\frac{4}{3} \left(\frac{q(x+\Delta x) - q(x-\Delta x)}{2\Delta x} \right) - \frac{1}{3} \left(\frac{q(x+2\Delta x) - q(x-2\Delta x)}{4\Delta x} \right)$$

and the derivative $\frac{\partial}{\partial x} (qT)$ by

$$\frac{4}{3} \left[\frac{(T(x)+T(x+\Delta x))(q(x)+q(x+\Delta x)) - (T(x)+T(x-\Delta x))(q(x)+q(x-\Delta x))}{4\Delta x} \right] \\ - \frac{1}{3} \left[\frac{(T(x)+T(x+2\Delta x))(q(x)+q(x+2\Delta x)) - (T(x)+T(x-2\Delta x))(q(x)+q(x-2\Delta x))}{8\Delta x} \right]$$

The second approximation is derived by averaging the flux qT to yield a conservative form of the dynamic equations. Note that if T is equal to 1 the second equation reduces to the first.

The primary variables are the horizontal components of the wind velocity, $W=(u,v)$, the temperature, T , the specific humidity, q , and the shifted surface pressure, π , ($\pi=P_s-P_{top}$, $P_{top}=10$ mb).

The secondary variables are the geopotential, ϕ , the vertical wind velocity, $\dot{\sigma}$, and the pressure, p .

The following pages give the differential equations of motion for the GCM model with the initial and boundary conditions. This is followed by the equations with the corresponding fourth order approximations which use the same notation as the current second order model. A complete description of the primitive equations with the σ coordinate system is found in the Arakawa UCLA notes (1976).

II. Primitive Equations of Motion

1.&2. Horizontal momentum equations

$$\begin{aligned} \mathbf{V} \frac{d\pi}{dt} + \pi \frac{d\mathbf{V}}{dt} &= \frac{\partial \pi \mathbf{V}}{\partial t} + \nabla \cdot (\pi \mathbf{V} \mathbf{V}) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma} \mathbf{V}) \\ &= -\pi \nabla \phi - \pi \sigma \frac{RT}{p} \nabla \pi - (f + u \frac{\tan \phi}{a}) \mathbf{k} \times \pi \mathbf{V} + \pi \mathbf{F} \end{aligned}$$

3. Continuity equation

$$(3.1) \quad \frac{\partial \pi}{\partial t} + \nabla \cdot (\pi \mathbf{V}) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma}) = 0, \quad \text{or}$$

$$(3.2) \quad \frac{\partial \pi}{\partial t} = - \int_0^1 \nabla \cdot (\pi \mathbf{V}) d\sigma = - \nabla \cdot \int_0^1 \pi \mathbf{V} d\sigma$$

4. Equation of state

$$\alpha = \frac{RT}{p}$$

5. First law of thermodynamics

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma} T) = \frac{\pi \omega \alpha}{C_p} + \frac{\pi Q}{C_p} \quad (\omega = \frac{dp}{dt})$$

From $\theta = T/p^k$, $p = p_T + \sigma \pi$, $\omega = \dot{\sigma} \pi + \dot{\pi} \sigma$, $\dot{\pi} = \frac{\partial \pi}{\partial t} + \mathbf{V} \cdot \nabla \pi$, $k = R/C_p$ we get

$$\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} = p^k \frac{\partial \pi \dot{\sigma} \theta}{\partial \sigma} + \pi \frac{\dot{\sigma} \alpha \pi}{C_p} \quad \text{Replacing in 5,}$$

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + p^k \frac{\partial \pi \dot{\sigma} \theta}{\partial \sigma} = \frac{\pi \sigma k T}{p} \left(\frac{\partial \pi}{\partial t} + \mathbf{V} \cdot \nabla \pi \right) + \frac{\pi Q}{C_p}$$

6. Humidity equation

$$\frac{\partial \pi q}{\partial t} + \nabla \cdot (\pi \mathbf{V} q) = 0$$

7. Hydrostatic equation

$$\frac{\partial \phi}{\partial p} = -C_p \theta \quad \left(\text{from } \frac{\partial p}{\partial \phi} = -\rho = -\frac{1}{\alpha} \right)$$

Of the variables $\pi, u, v, T, q, \phi, \alpha, \dot{\sigma}$ we update the 5 primary variables $\pi, u, v, T,$ and q using equations 1, 2, 3.2, 5 and 6. From equations 3.1, 4, and 7 we can obtain $\phi, \alpha,$ and $\dot{\sigma}$ which are our secondary variables. Note that $p = \sigma\pi + p_{\text{top}}$.

Sea level pressure (used only in the smoothing routine SMSHAP)

$$\text{Hydrostatic eq. } \Rightarrow \frac{\partial p}{\partial \phi} = -p = -\frac{1}{\alpha} = -\frac{p}{RT} \therefore \log\left(\frac{\text{SLP}}{p}\right) = -\int_{\phi_s}^0 \frac{d\phi}{RT}$$

$$\therefore \text{SLP} = p(\sigma=1) \exp\left(\frac{\phi_s}{RT}\right)$$

DERIVATION OF THE EQUATIONS AT THE POLES

Consider the continuity equation

$$\frac{\partial \pi}{\partial t} + \nabla \cdot (\pi \mathbf{V}) + \frac{\partial \pi \dot{\sigma}}{\partial \sigma} = 0$$

coupled with a conservation equation $\frac{dT}{dt} = S$ which can be expanded into

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} = \pi S$$

If we integrate this equation over a polar cap of radius $\Delta\phi$

$$\begin{aligned} \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi T}{\partial t} a^2 \cos\phi d\phi d\lambda &= - \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \nabla \cdot (\pi \mathbf{V} T) a^2 \cos\phi d\phi d\lambda \\ &- \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} a^2 \cos\phi d\phi d\lambda - \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \pi S a^2 \cos\phi d\phi d\lambda \end{aligned}$$

and we assume the value of $\frac{\partial \pi T}{\partial t}$ to be approximately constant over the polar cap

$$\begin{aligned} \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi T}{\partial t} a^2 \cos\phi d\phi d\lambda &\approx \left(\frac{\partial \pi T}{\partial t} \right) \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} a^2 \cos\phi d\phi d\lambda \\ &= 2\pi a^2 (1 - \cos\Delta\phi) \left(\frac{\partial \pi T}{\partial t} \right)_{NP} \end{aligned}$$

The first term in the rhs is, using Gauss' theorem

$$- \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \nabla \cdot (\pi \mathbf{V} T) a^2 \cos\phi d\phi d\lambda = - \int_0^{2\pi} \pi \mathbf{v} T a \sin\Delta\phi d\lambda$$

This can be approximated as

$$-\frac{2\pi}{IM} \sum_{i=1}^{IM} \pi_i v_i T_i a \sin \Delta \phi$$

The third term on the rhs is

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta \phi} \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} a^2 \cos \phi d\phi d\lambda \approx 2\pi a^2 (1 - \cos \Delta \phi) \left(\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} \right)_{NP}$$

and similarly with the source term.

From these equations we obtain

$$\left(\frac{\partial \pi T}{\partial t} \right)_{\text{Pole}} = -(-1)^m \frac{\cot \Delta \phi / 2}{aIM} \sum_{i=1}^{IM} (\pi_i v_i T_i) \frac{\pi}{2} - \Delta \phi - \left(\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} \right)_{\text{Pole}} + S$$

$$\text{since } \frac{2\pi a \sin \Delta \phi / 2}{IM} \frac{1}{2\pi a^2 (1 - \cos \Delta \phi)} = \frac{2 \sin \Delta \phi / 2 \cos \Delta \phi / 2}{IM a 2 \sin^2 \frac{\Delta \phi}{2}} = \frac{\cot \frac{\Delta \phi}{2}}{aIM}$$

This formulation is used in the continuity, momentum temperature and moisture equations. Note that the first term changes sign in the South Pole ($m=1$). In the momentum equations we make use of the transformation

$$U_i = -\sin \lambda u_i - (-1)^m \cos \lambda v_i$$

$$V_i = (-1)^m \cos \lambda u_i - \sin \lambda v_i$$

where $m=1$ for the South Pole and $m=2$ for the North Pole. U_i, V_i are the "cartesian" velocities at longitude λ_i , and u_i, v_i the corresponding spherical velocities.

The pressure gradient terms are computed making use of Green's

$$\text{Theorem: } \iint \frac{\partial Q}{\partial x} dx dy = \oint Q dy; \iint \frac{\partial P}{\partial y} dx dy = - \oint P dx$$

For example

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2-\Delta\phi} \frac{\partial \phi}{\partial x} a^2 \cos \phi d\phi d\lambda = -\phi \phi a \Delta\phi (-\cos \lambda d\lambda)$$

$$\phi = \frac{\pi}{2} - \Delta\phi$$

$$\approx + \frac{a\Delta\phi 2\pi}{IM} \sum_{i=1}^{IM} \phi_i \cos \lambda_i$$

In the U-momentum equation we have then the following pressure terms

$$\left(\frac{\partial \pi U}{\partial t}\right)_P = - \frac{\Delta\phi}{a(1-\cos\phi) IM} \sum_{i=1}^{IM} (\phi_i + \left(\frac{\sigma RT}{\rho}\right)_P \pi_i) \cos \lambda_i$$

and similarly for πV .

In the model we have approximated $\Delta\phi \approx 2 \sin \frac{\Delta\phi}{2}$

$$\text{Then } \frac{\Delta\phi}{a(1-\cos\phi) IM} \sim \frac{1}{a \sin \frac{\Delta\phi}{2} IM}$$

Based on this formulation we construct the fourth order scheme at the Poles by taking $\frac{4}{3}$ of the differences evaluated at $\Delta\phi$ from the Poles (as expanded here), minus $\frac{1}{3}$ of the differences at $2\Delta\phi$ from the Poles.

This formulation is not conservative at the Poles. However we have found that this has had no noticeable effect in the conservation of mass or energy in the model. In our shallow water experiments we studied a set of equations that were quadratically conservative, but inconsistent at the Pole, and another scheme analogous to the GFDL scheme, which is both quadratically conservative and consistent

at the Poles, but suffers from a serious truncation error near the Poles in the pressure gradient term. The scheme that we chose gave better results than the other two (Kalnay-Rivas, 1976).

Computation of the horizontal pressure gradient as suggested by N. A. Phillips

(1) Let $\theta = \bar{\theta} + \theta'$, $\bar{\theta} = 280^\circ\text{K}/1000^k$ i.e. constant

(2) $\phi = \bar{\phi} + \phi'$, $k = R/C_p = .286$

$$\frac{\partial \bar{\phi}}{\partial p^k} = -C_p \bar{\theta} \quad \text{and} \quad \bar{\phi} = \phi_0 - C_p \bar{\theta} p^k \quad \text{with} \quad \phi_0 = C_p \bar{\theta} 1000^k$$

(3) $T = \bar{T}(p) + T'$, $\bar{T}(p) = \bar{\theta} p^k$

Thus our new dependent variables are

$$\phi' = \phi + C_p \bar{\theta} (p^k - 1000^k)$$

$$T' = T - \bar{\theta} p^k$$

In this way $\pi(\nabla\phi + \frac{\sigma RT}{p} \nabla\pi)$, the pressure gradient in the momentum equations gets transformed into

$$\begin{aligned} (4) \quad & \pi(\nabla_{\phi'} \bar{\phi} + \nabla_{\phi'} \phi' + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi + \frac{\sigma RT'}{p} \nabla\pi) \\ & = \pi(\nabla_{\phi'} \phi' + \frac{\sigma RT'}{p} \nabla\pi) + \pi(\nabla_{\phi'} \bar{\phi} + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi) \end{aligned}$$

But the second parenthesis is zero:

$$\nabla_{\phi'} \bar{\phi} + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi = -C_p \bar{\theta} \nabla p^k + R \bar{\theta} \frac{p^k}{p} \nabla p = 0$$

In regions of steep orography, the second parenthesis in (4) is much larger than the first. When the horizontal pressure gradient terms are computed in their original form, the near cancellation of the two terms introduces large truncation errors. The procedure suggested by Phillips greatly reduces this truncation error. We have chosen a simpler definition of ϕ_0 than the one suggested by Phillips.

III. Finite Difference Variables and Grid

The notation used in fourth band model is the same as the standard GLAS(GISS) second order GCM except that we use a non-staggered horizontal grid. A complete description of the variables can be found in the TSANG-KARN documentation of the GISS 9 level model.

Let $u_{ijk}, v_{ijk}, T_{ijk}, q_{ijk}, \pi_{ij}$ be the finite difference approximations to the primary variables $u, v, T, q,$ and π at the mesh point $(i\Delta\lambda, j\Delta\phi, (k-1/2)\Delta\sigma)$. Also the scaled variables $\pi u, \pi v, \dots$ are approximated by $\pi u_{ijk}, \pi v_{ijk}, \dots$.

The finite difference equations also use the following geometric arrays:

$$DXP(j) = m_j = a \cos \phi_j \Delta\lambda, \quad DYU(j) = n_j = a \Delta\phi_j,$$

We use a factor of 12 in DXYP to make our scaled fourth order differences simpler.

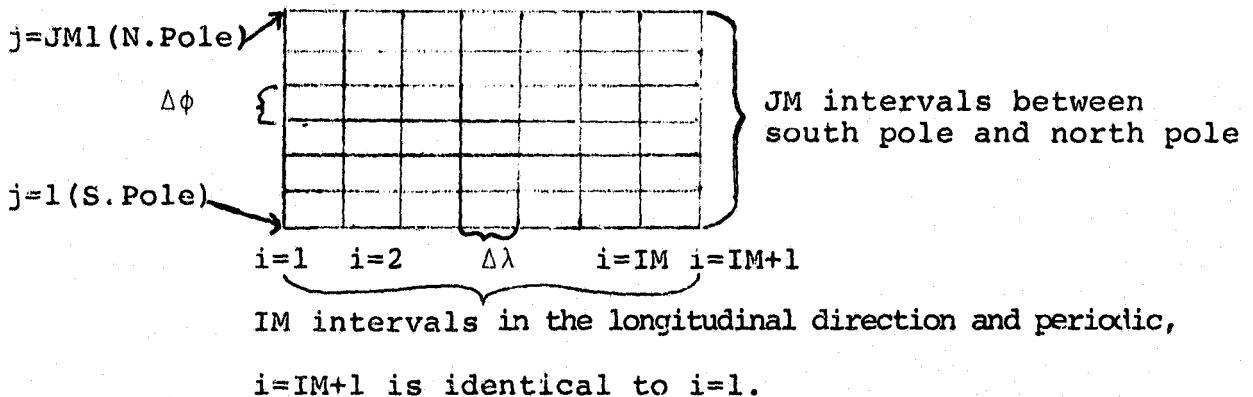
$$DXYP(j) = 12 \cdot n_j \cdot n_j, \quad \Pi_{ij} = DXYP(j) \cdot \pi_{ij}, \quad U_{ijk} = n_j \pi_{ij} U_{ijk}$$

$$\dot{S}_{ijk} = \Pi_{ij} \cdot \dot{\sigma}_{ijk}, \quad ADLDP = 12 \cdot a \cdot \Delta\lambda \Delta\phi, \quad V_{ijk}^* = m_j \pi_{ij} V_{ijk}$$

$$F_{ijk} = DXYP(j) f_j + ADLDP \cdot \sin \phi_j U_{i\&j}$$

Horizontal Grid

The fourth order band model uses an unstaggered grid in the horizontal direction.



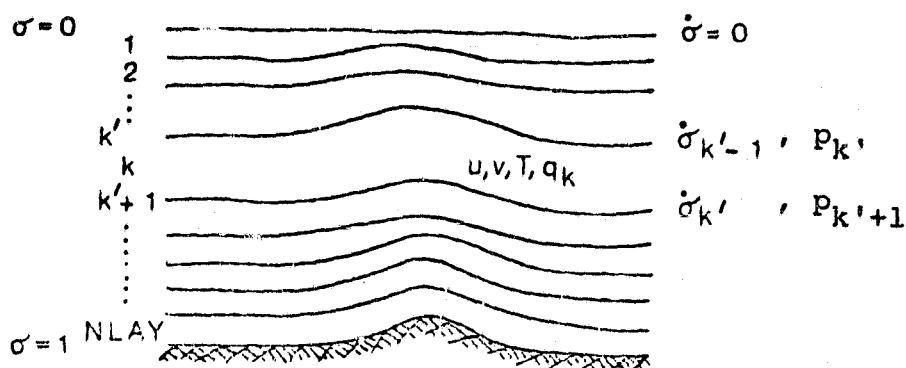
$$\Delta\lambda = \text{DLON} = \frac{\pi}{JM} \quad \Delta\phi = \text{DLAT} = \frac{\pi}{JM}$$

$$\lambda_i = (i-1)\Delta\lambda - \pi = (i-1)5^\circ - 180^\circ \quad \phi_j = (j-1)\Delta\phi - \frac{\pi}{2} = (j-1)4^\circ - 90^\circ$$

$$\text{JSP} = 1 \quad \text{JNP} = \text{JM}1 = \text{JM}+1$$

Vertical Grid

The vertical grid is staggered; the values of all the variables u, v, T, q, π, \dots except $\dot{\sigma}$ are computed at the center of each layer. The values of $\dot{\sigma}$ are computed at the edges of the layers.



In the case of uniform vertical resolution and $\text{NLAY}=9$ vertical layers,

$$\sigma_k = \text{SIG}(K) = \frac{k-1}{9} (0, 1/9, \dots, 1)$$

$$\sigma_{k'} = \text{SIGE}(K) = \frac{2k-1}{18} = \left(\frac{1}{18}, \dots, \frac{17}{18}\right)$$

$$\text{and } \Delta\sigma = 1/9.$$

$\dot{\sigma}_{ijk}$ and its scaled version \dot{S}_{ijk} are defined at the eight interior edges, i.e. for $\sigma_{k'}$ with $k=1$ to 8 since $\dot{\sigma}(0)=\dot{\sigma}(1)=0$ from the boundary conditions. The pressure $p_{k'} = \pi\sigma_{k'} + p_{\text{TOP}}$ is defined at the same level as $\dot{\sigma}_{k'-1}$. Note for level k we need \dot{S}_{ijk} and \dot{S}_{ijk-1} to form the second order vertical differences:

$$\dot{S}_{ijk} (v_{ijk} + v_{ijk+1}) - \dot{S}_{ijk-1} (v_{ijk} + v_{ijk-1})$$

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III-2 Periodic Filtering of Short Waves

An integral part of the numerical scheme is the periodic application (every ISMTH steps, generally 2 hours) of a 16th order Shapiro filter. This has the effect of removing waves shorter than $4\Delta x$, which are not resolved in the model, while waves longer than $4\Delta x$, which are accurately computed by the difference scheme, are not affected by the filter (Kalnay-Rivas and Hoitsma, 1979).

The filter is applied to an array q_j in the following way:

$$\text{Let } d_+ (q_j) = q_{j+1} - q_j, \quad d_- (q_j) = q_j - q_{j-1}$$

Then a Shapiro filter of order $2N$ is given by

$$\bar{q}_j = q_j - (-1)^N (d_+ d_-)^N q_j$$

The response of the filter applied to a wave of the form

$$q_j = Q \exp(i \frac{2\pi}{L} \Delta x j) \text{ is}$$

$$\bar{q}_j = (1 - \sin^{2N} \frac{\pi \Delta x}{L}) q_j$$

The 2-dimensional filter is applied as a product of 1-dimensional filters (first in longitude, then in latitude). In latitude we filter the fields on great circles formed by meridians of longitude λ and $\lambda + \pi$, where $0 \leq \lambda < \pi$. We are presently filtering only potential temperature and sea level pressure. These fields were chosen because they are not very affected by orography. Winds are not currently filtered, because the adjustment between mass and velocity fields does not allow the development of short waves in the winds alone. However, in the tropics, where the adjustment of winds to the mass field is minimal, the winds are somewhat noisy, and we may opt to filter them too.

IV. Boundary Conditions

Periodicity in the Zonal (East-West) direction

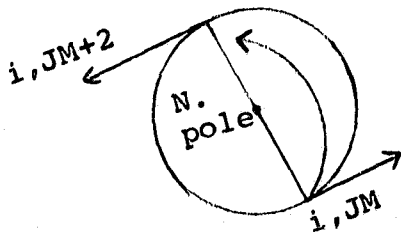
$$\pi_{IM+mj} = \pi_{mj} \quad m=1,2,3,\dots, \quad j=2,\dots,JM$$

$$Q_{IM+mjk} = Q_{mjk} \quad m=1,2,3,\dots, \quad j=2,\dots,JM, \quad k=1,\dots,NLAY$$

for $Q=u,v,T,q,\phi,\delta,\pi u,\pi v,\dots$

Boundary conditions at the north and south poles. Define the array

INDEX as follows:



$$INDEX(i) = i + \frac{IM}{2} \quad i=1,2,\dots,\frac{IM}{2}$$

$$INDEX(i + \frac{IM}{2}) = i \quad i=\frac{IM}{2}+1,\dots,IM$$

i.e., INDEX: $\frac{IM}{2}+1, \frac{IM}{2}+2, \dots, IM, 1, 2, \dots, \frac{IM}{2}$

Then we can define

$$\pi_{iJM+2} = \pi_{INDEX(i)JM} \quad \text{for the points needed "beyond" the North Pole,}$$

and $\pi_{i0} = \pi_{INDEX(i)2}$ for the points "beyond" the South Pole

and similarly for T,q,ϕ .

For the horizontal velocity $V=(u,v)$ we have

$$V_{iJM+2k} = -V_{iJMk}$$

$$V_{i0k} = -V_{i2k}$$

V. Finite Difference Equations

THE ZONAL (U) MOMENTUM EQUATION

$$\frac{\partial \pi u}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial (\pi u \cdot u)}{\partial \lambda} + \frac{\partial (\pi v \cos \phi \cdot u)}{\partial \phi} \right] - \frac{\partial (\pi \dot{\sigma} u)}{\partial \sigma}$$

$$- \frac{\pi}{a \cos \phi} \left[\frac{\partial \phi}{\partial \lambda} + \frac{\sigma RT}{p} \frac{\partial \pi}{\partial \lambda} \right] + \left(f + \frac{u \tan \phi}{a} \right) \pi v + \pi F_x$$

$$\frac{\partial \pi U_{ijk}}{\partial t} = \left\{ \begin{array}{l} \text{COMP1: HA: } \overbrace{PU1_{i-1/2j}}^{(U_{ijk}^* + U_{i-1jk}^*)} (U_{ijk} + U_{i-1jk}) - \overbrace{PU1_{i+1/2j}}^{(U_{ijk}^* + U_{i+1jk}^*)} (U_{ijk} + U_{i+1jk}) \\ - .5 \left[\overbrace{PU2_{i-1}}^{(U_{ijk}^* + U_{i-2jk}^*)} (U_{ijk} + U_{i-2jk}) - \overbrace{PU2_{i+1}}^{(U_{ijk}^* + U_{i+2jk}^*)} (U_{ijk} + U_{i+2jk}) \right] \\ + 4 * \left[\overbrace{PV1}^{(V_{ijk}^* + V_{ij-1k}^*)} (U_{ijk} + U_{ij-1k}) - \overbrace{(V_{ijk}^* + V_{ij+1k}^*)}^{PV1} (U_{ijk} + U_{ij+1k}) \right] \\ - .5 * \left[\overbrace{(V_{ijk}^* + V_{ij-2k}^*)}^{PV2} (U_{ijk} + U_{ij-2k}) - \overbrace{(V_{ijk}^* + V_{ij+2k}^*)}^{PV2} (U_{ijk} + U_{ij+2k}) \right] \\ \text{VA: } \\ + .5 * \left[\dot{S}_{ijk-1} (U_{ijk} + U_{ijk-1}) - \dot{S}_{ijk} (U_{ijk} + U_{ijk+1}) \right] / \Delta \sigma_k \\ \text{COMP2: PG: } \\ + (\pi_{ij} * n_j * \left\{ 8 * [\phi'_{i-1jk} - \phi'_{i+1jk} + \frac{\sigma_k RT'_{ijk}}{p_k} (\pi_{i-1j} - \pi_{i+1j})] + \right. \\ \left. [\phi'_{i+2jk} - \phi'_{i-2jk} + \frac{\sigma_k RT'_{ijk}}{p_k} (\pi_{i+2j} - \pi_{i-2j})] \right\}) \\ \text{C: } \\ + F_{ijk} * \pi_{ij} * V_{ijk} \end{array} \right\} + (\pi F_x)$$

Note: COMP1, COMP2, COMP3 are the names of the three subroutines where the different terms are computed. HA: Horizontal advection terms. VA: Vertical advection terms. PG: Pressure gradient terms. C: Coriolis term.

Also note that PV2 is set equal to zero for $j=0$ and $j=JM$, i.e., there is no transport of mass across the poles.

THE MERIDIONAL (V) MOMENTUM EQUATION

$$\frac{\partial \pi V}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial (\pi u \cdot v)}{\partial \lambda} + \frac{\partial (\pi v \cos \phi v)}{\partial \phi} \right] - \frac{\partial \pi \dot{\sigma} v}{\partial \sigma}$$

$$- \frac{\pi}{a} \left[\frac{\partial \phi}{\partial \phi} + \frac{\sigma RT}{p} \frac{\partial \pi}{\partial \phi} \right] - \left(f + \frac{u \tan \phi}{a} \right) \pi u + \pi F_y$$

COMP1: HA:

$$\frac{\partial \pi V_{ijk}}{\partial t} = \left\{ 4. * [(U_{ijk}^* + U_{i-1jk}^*) (V_{ijk} + V_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (V_{ijk} + V_{i+1jk})] \right.$$

$$- .5 [(U_{ijk}^* + U_{i-2jk}^*) (V_{ijk} + V_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (V_{ijk} + V_{i+2jk})]$$

$$+ 4. * [(V_{ijk}^* + V_{ij-1k}^*) (V_{ijk} + V_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (V_{ijk} + V_{ij+1k})]$$

$$- .5 * [(V_{ijk}^* + V_{ij+2k}^*) (V_{ijk} + V_{ij-2k}) - (V_{ijk}^* + V_{ij+2k}^*) (V_{ijk} + V_{ij+2k})]$$

VA:

$$+ .5 * [\dot{s}_{ijk-1} (V_{ijk} + V_{ijk-1}) - \dot{s}_{ijk} (V_{ijk} + V_{ijk+1})] / \Delta \sigma_k \}$$

COMP2: P:

$$+ (\pi_{ij} * m_j * \{ 8. * [\phi_{ij-1k}^* - \phi_{ij+1k}^* + \frac{\sigma_k RT_{ijk}^*}{p_k} (\pi_{ij-1} - \pi_{ij+1})]$$

$$+ [\phi_{ij+2k}^* - \phi_{ij-2k}^* + \frac{\sigma_k RT_{ijk}^*}{p_k} (\pi_{ij+2} - \pi_{ij-2})] \}$$

C:

COMP3:

$$- F_{ijk} * \pi_{ij} * U_{ijk} + (\pi F_y)$$

THE THERMODYNAMIC ENERGY EQUATION

$$\frac{\partial \pi T}{\partial t} = -\frac{1}{a \cos \phi} \left[\frac{\partial \pi U T}{\partial \lambda} + \frac{\partial \pi V \cos \phi \cdot T}{\partial \phi} \right] - p^k \frac{\partial \pi \dot{\sigma} T / p^k}{\partial \sigma}$$

$$+ \frac{\pi \sigma K T}{p} \left(\frac{\partial \pi}{\partial t} + \frac{U}{a \cos \phi} \frac{\partial \pi}{\partial \lambda} + \frac{V}{a} \frac{\partial \pi}{\partial \phi} \right) + \frac{\pi Q}{C_p}$$

COMP1: HA:

$$\frac{\partial \pi T_{ijk}}{\partial t} = \{ 4 * [(U_{ijk}^* + U_{i-1jk}^*) (T_{ijk} + T_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (T_{ijk} + T_{i+1jk})]$$

$$- .5 * [(U_{ijk}^* + U_{i-2jk}^*) (T_{ijk} + T_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (T_{ijk} + T_{i+2jk})]$$

$$+ 4 * [(V_{ijk}^* + V_{ij-1k}^*) (T_{ijk} + T_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (T_{ijk} + T_{ij+1k})]$$

$$- .5 * [(V_{ijk}^* + V_{ij-2k}^*) (T_{ijk} + T_{ij+2k}) - (V_{ijk}^* + V_{ij+2k}^*) (T_{ijk} + T_{ij+2k})] \}$$

VA:

$$+ .5 p_{ijk}^k \left[\dot{s}_{ijk-1} \left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk-1}}{p_{ijk-1}^k} \right) - \dot{s}_{ijk} \left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk+1}}{p_{ijk+1}^k} \right) \right] / \Delta \sigma_k \}$$

COMP2:

$$\left(+ \frac{\sigma_k K T_{ijk}}{p_{ijk}} \left[\pi_{ij} \frac{\partial \pi}{\partial t} + U_{ijk}^* \{ 8 * [\pi_{i+1j} - \pi_{i-1j}] + \pi_{i-2j} - \pi_{i+2j} \} \right. \right.$$

$$\left. \left. + V_{ijk}^* \{ 8 * [\pi_{ij+1} - \pi_{ij-1}] + \pi_{ij-2} - \pi_{ij+2} \} \right] \right)$$

COMP3:

$$+ \left(\frac{\pi Q_{ijk}}{C_p} \right)$$

THE MOISTURE EQUATION

$$\frac{\partial \pi q}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial \pi U \cdot q}{\partial \lambda} + \frac{\partial \pi V \cos \phi \cdot q}{\partial \phi} \right] - \frac{\partial \pi \delta q}{\pi \sigma} + \pi (E - C)$$

COMP1: HA

$$\begin{aligned} \frac{\partial \pi q_{ijk}}{\partial t} = & \{ 4. * [(U_{ijk}^* + U_{i-1jk}^*) (q_{ijk} + q_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (q_{ijk} + q_{i+1jk})] \\ & - .5 * [(U_{ijk}^* + U_{i+2jk}^*) (q_{ijk} + q_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (q_{ijk} + q_{i+2jk})] \\ & + 4. * [(V_{ijk}^* + V_{ij-1k}^*) (q_{ijk} + q_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (q_{ijk} + q_{ij+1k})] \\ & - .5 * [(V_{ijk}^* + V_{ij-2k}^*) (q_{ijk} + q_{ij-2k}) - (V_{ijk}^* + V_{ij+2k}^*) (q_{ijk} + q_{ij+2k})] \\ & + .5 * [\dot{S}_{ijk-1} (q_{ijk} + q_{ijk-1}) - \dot{S}_{ijk} (q_{ijk} + q_{ijk+1})] / \Delta \sigma_k \} \end{aligned}$$

COMP3:

$$+ \pi (E_{ijk} - C_{ijk})$$

Note: The current transport scheme for the moisture field is being modified.

THE PRESSURE-TENDENCY EQUATION

$$\frac{\partial \pi}{\partial t} = - \sum_{\ell=1}^L \frac{(\Delta \sigma)_{\ell}}{a \cos \phi} \left[\frac{\partial \pi u}{\partial \lambda} + \frac{\partial \pi v \cos \phi}{\partial \phi} \right]$$

COMPl:

$$\begin{aligned} \text{CONV}_{i\ell j} = & \Delta \sigma_{\ell} \{ 8. * (U_{i-1j\ell}^* - U_{i+1j\ell}^*) + U_{i+2j\ell}^* - U_{i-2j\ell}^* \\ & + 8. * (V_{ij-1\ell}^* - V_{ij+1\ell}^*) + V_{ij+2\ell}^* - V_{ij-2\ell}^* \} \end{aligned}$$

COMPl:

$$\frac{\partial \pi_{ij}}{\partial t} = \left(\sum_{\ell=1}^L \text{CONV}_{i\ell j} \right)$$

THE VERTICAL VELOCITY EQUATION

$$\frac{\partial \pi}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial \pi u}{\partial \lambda} + \frac{\partial \pi v \cos \phi}{\partial \phi} \right] - \frac{\partial}{\partial \sigma} (\pi \dot{\sigma})$$

Thus

$$\Delta \sigma_k \frac{\partial \pi}{\partial t} = \text{CONV}_{ikj} - \Delta \dot{S}_{ijk}$$

giving:

COMPl:

$$\dot{S}_{ijk} = \{ \dot{S}_{ijk-1} + \text{CONV}_{ikj} - \Delta \sigma_k \frac{\partial \pi}{\partial T} \}$$

VI. The Forecast Equations at the Poles

The fourth order band model uses a spherical cap at the poles, and the finite difference approximations to the equations of motion must be derived for this spherical region. Stereographic projection is used to give us a well-defined velocity vector at the poles.

From trigonometry, (Fig. VIa) if the vector (U, V) in X - Y coordinates is represented by (U', V') in X' - Y' coordinates where the prime coordinate axes are rotated by an angle λ then,

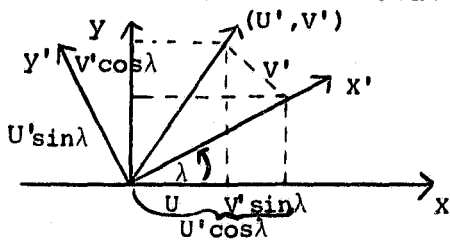


Fig. VIa

From Fig. VIb, which shows a unit vector in both coordinate systems, we see that, at the North Pole, spherical coordinates are rotated by an angle $(\frac{\pi}{2} + \lambda)$ with respect to the polar stereographic coordinates. Therefore,

$$U_{NP} = U \cos(\lambda + \frac{\pi}{2}) - V \sin(\lambda + \frac{\pi}{2})$$

$$V_{NP} = U \sin(\lambda + \frac{\pi}{2}) + V \cos(\lambda + \frac{\pi}{2})$$

or

$$\begin{aligned} (3) \quad U_{NP} &= -U \sin \lambda - V \cos \lambda \\ V_{NP} &= U \cos \lambda - V \sin \lambda \end{aligned}$$

Similarly, for the South Pole:

$$\begin{aligned} (4) \quad U_{SP} &= -U \sin \lambda + V \cos \lambda \\ V_{SP} &= -U \cos \lambda - V \sin \lambda \end{aligned}$$

also,

$$\begin{aligned} (5) \quad U &= -U_{NP} \sin \lambda + V_{NP} \cos \lambda \\ V &= -U_{NP} \cos \lambda - V_{NP} \sin \lambda \end{aligned}$$

$$\begin{aligned} (6) \quad U &= -U_{SP} \sin \lambda - V_{SP} \cos \lambda \\ V &= U_{SP} \cos \lambda - V_{SP} \sin \lambda \end{aligned}$$

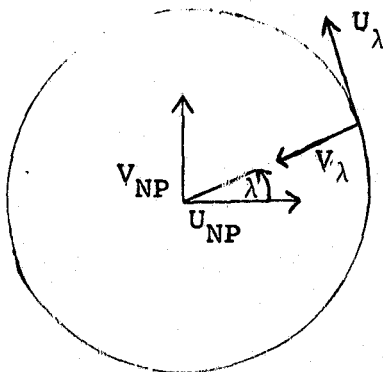


Fig. VIb

The initial values for U_{NPK} , V_{NPK} , U_{SPK} , V_{SPK} are obtained from equations 3 and 4 by averaging in the zonal direction on the line of latitude $j=45$ for the north pole and $j=2$ for the south pole. In the forecast stereographic velocities are advected and transformed back into spherical velocities after each time step by equations 5 and 6.

In the following equation we will denote the polar velocities by $VP_{k,m}$, $VP_{k,m}$ and the temperature and other variables in a similar way, $TP_{k,m}$, $\pi P_{k,m}$, ... where $m=1$ for the south pole and $m=2$ for the north pole. The following constants are used,

$$COEF1 = (-1)^m \quad COEF2 = -COEF1 \quad RADIM = 3 \cdot a \cdot IM$$

$$CON1 = 4COT(.5\Delta\phi) / RADIM$$

$$CON2 = -COT(\Delta\phi) / RADIM$$

$$CON3 = 4DT / (RADIM * SIN(.5\Delta\phi))$$

$$CON4 = DT / (RADIM * SIN(\Delta\phi))$$

$$JPOL(K,M) = \begin{pmatrix} 2 & JM \\ 3 & JM-1 \end{pmatrix}$$

$r = JPOL(1,m)$ = first interior value of j (2 for the S. Pole, JM for the N. Pole)

$s = JPOL(2,m)$ = second interior value of j (3 for the S. Pole, $JM-1$ for the N. Pole).

ZONAL (U) MOMENTUM EQUATION (POLES)

$$\begin{aligned}
 \frac{\partial \pi UP_{k,m}}{\partial t} = & \text{COMP1:} \\
 & \text{IM} \\
 & \{ \text{COEF1} * [\text{CON1} * \sum_{i=1}^{\text{IM}} -\pi_{ir} V_{irk} (U_{irk} \text{Sin} \lambda_i + \text{COEF1} * V_{irk} \text{COS} \lambda_i) \\
 & + \text{CON2} * \sum_{i=1}^{\text{IM}} -\pi_{is} V_{isk} (U_{isk} \text{Sin} \lambda_i + \text{COEF1} * V_{isk} \text{COS} \lambda_i)] \\
 & + .5 [\dot{S}P_{k-1} (UP_k + UP_{k-1}) - \dot{S}P_k (UP_k + UP_{k+1})] / \Delta \sigma_k \} \\
 & \text{COMP2:} \\
 & + \pi P_m \cdot [-\frac{\text{CON3}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{irk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{ir}) \text{COS} \lambda_i \\
 & + \frac{\text{CON4}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{isk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{is}) \text{COS} \lambda_i] \\
 & \text{COMP3:} \\
 & + f_m \pi P_m \cdot VP_{k,m} + (\pi P_m \cdot FP_{k,m}
 \end{aligned}$$

MERIDIONAL (V) MOMENTUM EQUATION (POLES)

$$\begin{aligned}
 \frac{\partial \pi VP_{k,m}}{\partial t} = & \text{COMP1:} \\
 & \text{IM} \\
 & \{ \text{COEF1} * [\text{CON1} * \sum_{i=1}^{\text{IM}} -\pi_{ir} V_{irk} (\text{COEF2} * U_{irk} * \text{COS} \lambda_i + V_{irk} \text{Sin} \lambda_i) \\
 & + \text{CON2} * \sum_{i=1}^{\text{IM}} -\pi_{is} V_{isk} (\text{COEF2} * U_{isk} * \text{COS} \lambda_i + V_{isk} \text{Sin} \lambda_i)] \\
 & + .5 [\dot{S}P_{k-1} (VP_k + VP_{k-1}) - \dot{S}P_k (VP_k + VP_{k+1})] / \Delta \sigma_k \} \\
 & \text{COMP2:} \\
 & + \pi P_m \cdot [-\frac{\text{CON3}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{irk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{ir}) \text{Sin} \lambda_i \\
 & + \frac{\text{CON4}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{isk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{is}) \text{Sin} \lambda_i] \\
 & \text{COMP3:} \\
 & - f_m \pi P_m \cdot UP_{k,m} + (\pi P_m \cdot SP_{k,m}
 \end{aligned}$$

THE THERMODYNAMICS ENERGY EQUATION (POLES)

COMP1:

$$\frac{\partial \pi TP_{k,m}}{\partial t} = \{ \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} T_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk} T_{isk})$$

$$+ .5 p_{k,m}^k * \{ \dot{S}_{k-1,m} \left(\frac{TP_{k,m}}{p_{k,m}^k} + \frac{TP_{k-1,m}}{p_{k-1,m}^k} \right)$$

$$- \dot{S}_{k,m} \left(\frac{TP_{k,m}}{p_{k,m}^k} + \frac{TP_{k+1,m}}{p_{k+1,m}^k} \right) \} / \Delta \sigma_k \}$$

COMP2:

$$+ \left(\frac{\pi_{m}^k \sigma_k K T P_{k,m}}{p_{k,m}^k} \right) \left\{ \frac{\partial \pi_{m}}{\partial t} + \text{CON5} * \sum_{i=1}^{\text{IM}} [(\text{COEF1} * U P_{k,m} \cos \lambda_i + V P_{k,m} \sin \lambda_i) * \right.$$

COMP3:

$$\left. (8 \pi_{ir} - \pi_{is}) \right] \}$$

$$+ \left(\frac{\pi_{m}}{C_p} Q P_k \right)$$

THE MOISTURE BALANCE EQUATION (POLES)

COMP1:

$$\frac{\partial \pi q P_{k,m}}{\partial t} = \{ \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} q_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk} q_{isk})$$

$$+ .5 [\dot{S}_{k-1,m} (q_{k,m}^P + q_{k-1,m}^P) - \dot{S}_{k,m} (q_{k,m}^P + q_{k+1,m}^P)] / \Delta \sigma_k \}$$

COMP3:

$$+ (\pi_{m} * E P_k)$$

THE PRESSURE TENDENCY EQUATION (POLES)

$$\text{CONVPL}_{k,m} = \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk}) \Delta\sigma_k$$

$$\frac{\partial \pi P_m}{\partial t} = \sum_{k=1}^{\text{NLAY}} \text{CONVPL}_{k,m}$$

THE VERTICAL VELOCITY EQUATION (POLES)

$$\dot{S}P_{k,m} = \dot{S}P_{k-1,m} + \text{CONVPL}_k - \Delta\sigma_k \frac{\partial \pi P_m}{\partial t}$$

VII. Diagnostic Equations ($\phi, \dot{\sigma}, p$)

Once the updated values of πU , πV , πT , πq are found we unscale:

$$U_{ijk}^{n+1} = \pi U_{ijk}^{n+1} / \pi_{ij}^{n+1} \quad \text{for all } i, j, k.$$

Similarly for V, T, q . We also filter the fields near the poles to prevent linear instability (see subroutine AVRX). $\dot{\sigma}$ is obtained from \dot{S} by unscaling also.

We determine from π_{ij}^{n+1} and σ_k , p_{ijk}^{n+1}

$$p_{ijk}^{n+1} = \sigma_k \pi_{ij}^{n+1} + p_{TOP} \quad p_{TOP} = \text{constant.}$$

ϕ_S is the surface geopotential (a function only of latitude and longitude)

For the Phillips geopotential we define $(p)^k$ at the center of the layer in the following way:

$$(p)_{ijk}^k = \frac{p_{ijk+1}^{k+1} - p_{ijk}^{k+1}}{(k+1)(p_{ijk+1}^{k+1} - p_{ijk}^{k+1})} \quad \left(p^k = \frac{1}{k+1} \frac{\partial p^{k+1}}{\partial p} \right)$$

$k=1, 2, \dots, NLAY$

where

$$p_{ijk}^{k+1} = (\text{SIGE}(k) * \pi_{ij} + p_{TOP})^{k+1} \quad (\text{SIGE}(k) = \sigma_k'), \text{ i.e.}$$

p^{k+1} is obtained by exponentiation and $(p)^k$ by differences.

The following equations represent the geopotential calculations used in the old fourth order model.

Let

$$C_{ijk} = \frac{\pi_{ij}^k \sigma_k R \Delta \sigma_k}{p_{ijk}} - \frac{C_p}{2} \left[\sigma_k \left(\frac{p_{ijk+1}^k}{p_{ijk}^k} - 1 \right) + \sigma_{k-1} \left(1 - \frac{p_{ijk-1}^k}{p_{ijk}^k} \right) \right]$$

for $k=1, \dots, NLAY$

with

$$p_{ij0}^k = p_{ij1}^k \text{ and } p_{ijNLAY+1}^k = p_{ijNLAY}^k$$

An optimized version of C_{ijk} is:

$$C_{ijk} = \frac{\pi_{ij}^k \sigma_k R \Delta \sigma_k}{p_{ijk}} - \frac{.5 C_p}{p_{ijk}^k} \left\{ \Delta \sigma_k (p_{ijk+1}^k - p_{ijk}^k) + \sigma_{k-1} (p_{ijk+1}^k - p_{ijk-1}^k) \right\}$$

Rather than compute ϕ and then subtract $\bar{\phi}$ we do everything at once:

$$(1) \quad \phi'_{ijNLAY} = \phi'_S - CPTH * (PSKAPA - p_{ijNLAY}^k) + \sum_{\ell=1}^{NLAY} C_{ij\ell} T_{ij\ell}$$

$$(2) \quad \phi'_{ij\ell} = \phi'_{ij\ell+1} + \frac{C_p}{2} (p_{ij\ell+1}^k - p_{ij\ell}^k) \left(\left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk+1}}{p_{ijk+1}^k} \right) - 2\bar{\phi} \right)$$

where

$$CPTH = C_p \cdot \bar{\phi} \quad PSKAPA = 1000^k = p_S^k$$

VIII. The Time Differencing Scheme

The fourth order band model has the option of using the Matsuno time scheme or the smooth leapfrog scheme (see MWR-Vol 100 (487-490) R. Asselin).

Let Q^n represent a typical variable that is to be updated to time $n+1$, and let $D(Q^n)$ represent the nonlinear space differences. The Matsuno (Euler-backward) scheme is as follows:

$$\begin{aligned}\tilde{Q} &= Q^n + \Delta t D(Q^n) \\ Q^{n+1} &= Q^n + \Delta t D(\tilde{Q})\end{aligned}$$

The standard leapfrog scheme is given by

$$(1) \quad \frac{Q^{n+1} - Q^{n-1}}{2\Delta t} = D(Q^n)$$

For the smooth leapfrog scheme we replace Q^{n-1} by \bar{Q}^{n-1}

$$(2) \quad Q^{n+1} = \bar{Q}^{n-1} + 2\Delta t D(Q^n)$$

with

$$(3) \quad \bar{Q}^n = (1-\nu)Q^n + .5\nu(\bar{Q}^{n-1} + Q^{n+1})$$

Equation (3) represents a simple time filter except \bar{Q}^{n-1} is used instead of Q^{n-1} in order to save core storage. The above equations (2) and (3) represent the order in which the smooth leapfrog scheme is evaluated. For $n=1$ we define $\bar{Q}^0 = Q^0$ then we update in equation (2) followed by the filtering in equation (3) which is needed for the next time step.

The smoothing step introduces dissipation with respect to time, controlled explicitly by the parameter ν , as compared to the implicit dissipation in the Matsuno scheme. The amplification factor can be found in the paper by Asselin.

A further modification must be made to the smooth leapfrog scheme when source terms are included. Essentially the idea is that we must include the source term effect over two steps rather than one. If we do not do this, then the source effects (COMP3) are included only in every other step which will introduce large discretization errors. (For details see the attached report, Appendix B.)

If the source terms are called every NCOMP3 steps, then for step $n = \text{NCOMP3}$

$$Q_*^n = \bar{Q}^{n-2} + 2\Delta t D(Q^{n-1})$$

$$\bar{Q}_*^{n-1} = (1-\nu) Q^{n-1} + .5\nu(\bar{Q}^{n-2} + Q^n)$$

Then compute the source terms S^n and include in both steps n and $n-1$,

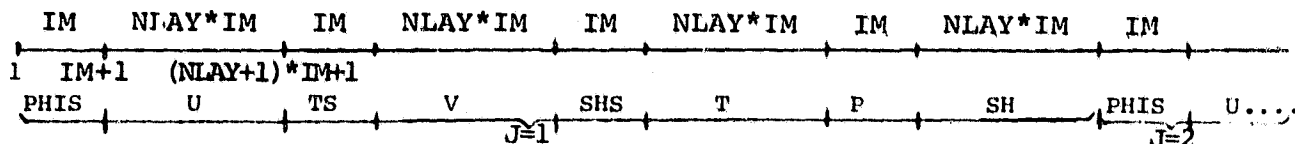
$$Q^n = Q_*^n + \text{NCOMP3} \cdot \Delta t S^n$$

$$\bar{Q}^{n-1} = \bar{Q}_*^{n-1} + \text{NCOMP3} \cdot \Delta t S^n$$

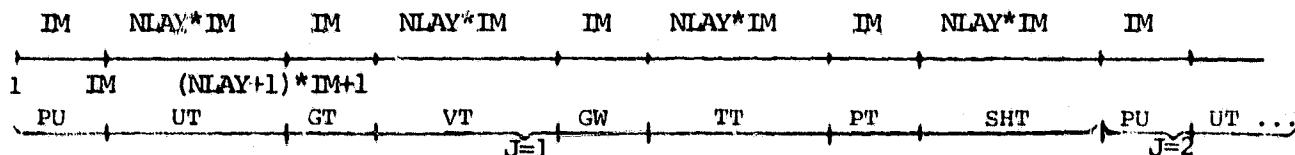
The actual code is complicated by the fact that we actually use scaled and unscaled variables, but the generalization is straightforward.

IX. Documentation of the Code (Preliminary)

The band fourth model uses special equivalences and nonstandard dimensions in order to have the variables P,U,V,T,SH stored in contiguous locations for each line of latitude. Thus, we desire to have the variables stored as follows.



The scaled variables PT, UT, VT, TT, SHT are to be stored as follows:



The above storage designation is accomplished by dimensioning PHIS(4*(NLAY+1)*IM,JNP) instead of PHIS(IM,JNP) (for a fine grid we have PHIS(2880,46), for ultrafine we have PHIS(4800,72)). Then we equivalence NLAY*IM locations of the first line of latitude of U with PHIS(IM+1,1) to PHIS((NLAY+1)*IM,1) (the first IM locations of PHIS are used for the quantity PHIS itself).

Similarly, we equivalence TS(1,1) to TS(IM,1) with PHIS((NLAY+1)*IM+1,1) to PHIS((NLAY+2)*IM,1) and so forth. (See the enclosed computer codes for the exact values in the fine and ultrafine versions.)

In order to have the successive lines of latitude arranged properly in storage we dimension our variables to achieve this purpose. Thus, we have $U(IM, 4*(NLAY+1), 1)$ instead of $U(IM, NLAY, JM)$, similarly for V, T, SH, UT, VT, TT , and SHT . For P we have $P(4*(NLAY+1)*IM, 1)$ instead of $P(IM, JNP)$ and similarly for TS, SHS, GT, GW , and PT . Note that the variables U, UT, \dots are only computed at $IM*NLAY*JNP$ points, the special dimensions are needed to properly align the variables in storage. It is important to recall that we are using two properties of the FORTRAN Compiler:

First, by equivalencing two elements of two different arrays we implicitly equivalence the other elements of the arrays. Second, that last array dimension can be left as 1 as long as it is dimensioned properly in the calling routine or it is equivalenced to a properly dimensioned array. Note we could have dimensioned U as $U(IM, 4*(NLAY+1), JNP)$ we would use the same amount of storage. It is crucial to have the $IM*(NLAY+1)$ dimension because it would cause the computations on U for the successive lines of latitude to be shifted $IM*(NLAY+1)$ locations where the next line of latitude of U are stored. For clarity and simplicity in programming we use the standard equivalence of U, V, T, SH with $Q(I, L, N, J)$ with U equivalent to $Q(I, L, 1, J)$ and so forth. Since we want P, U, V, T , and SH in contiguous storage locations and the equivalence of Q with U, V, T , and SH , then we must dimension Q as $Q(IM, NLAY+1, 4, 1)$ instead of $Q(IM, NLAY, 4, 1)$. We fill in the extra locations by including $PHIS, TS, SH$, and P .

COMPØ Description

This subroutine contains the time schemes and controls the calling sequences of the routines COMP1, COMP2, COMP3, and the polar filtering routine AVRX.

The logic of this subroutine involves three considerations. First, it permits one to choose either the Matsuno or the smooth leapfrog time scheme. Second, the latter scheme involves the use of storage arrays PSM and QSM. Third, the unscaling, smoothing of the updated variables QT and the call of subroutine COMP3 occur at the value $JS2=J-2$ when J is the value of the current line of latitude which is being computed. The routine COMP1 and COMP2 require the values of U,V,... at JS2 in order to compute the updated values at J because fourth order differences are used in the meridional (J) direction. Only after the COMP1 and COMP2 are called for value J can we unscale, smooth, and finish processing the variables at JS2.

The code for the poles is identical in format with the code for the other J values except the variables are scaled by the pressure PPOl only.

The main program contains two calls to COMPØ which cannot be treated independently because of the calling sequence: COMPØ(Q,QT), COMPØ(QT,Q). If the first call is a leapfrog step (LF), the second one must be also LF. If the first call is Matsuno predictor (MP), the second one must be Matsuno corrector (MC). This is represented symbolically by $LF \rightarrow LF$, or $MP \rightarrow MC$. The second call can be followed by either MP or LF. Each of these combinations requires different transfers.

Description of the Time Step Sequence Parameters

NSTEP: Counts the time steps. Starts and restarts both begin with **NSTEP=0**.

NSEQ: The number of steps (combined matsuno and leapfrog) in each (repeated) sequence of time steps.

MLF(I): **MLF(I)=0** or **1** according to whether the **I**th step in the sequence is Leapfrog or Matsuno, respectively. First step is always Matsuno (**MLF(1)=1**).

ISMTH,) Smoothing routine (**SMSHAP**) is called **MOD(NSTEP-NSM1,**
NSM1:) **ISMTH)**; if **ISMTH=0**, there is no smoothing.

NCOMP3,) Physics routine (**COMP3**) is called **MOD (NSTEP-NCM1,**
NCM1:) **NCOMP3)**.

Sample Runs

Matsuno only:

NSEQ=1, MLF(1)=1, MATSUN=1, DT=750., NSM1=0, NCM1=0
BCINO3=4, ISMTH=8

Leapfrog only:

NSEQ=1, MLF(1)=1, MATSUN=0, DT=600., NSM1=0, NCM1=0
NCOMP3=5, ISMTH=10

1 Matsuno, 4 Leap-Frog:

NSEQ=5, MLF=(1,0,0,0,0), MATSUN=not needed, NSM1=0,
NCM1=0, DT=600., NCOMP3=5, ISMTH=10

COMPl Description

The COMPl subroutine contains the horizontal and vertical advection differences. The DO loops over I are arranged to make use of the periodicity of the variables in the zonal (I) direction. For example, suppose we are to compute $D(I)=Q(I+1)-Q(I)$ for $I=1,\dots,IM$. Then the corresponding code is

```

I=IM
DO 10 IP1=1,IM
D(I)=Q(IP1)-Q(I)
I=IP1
10 CONTINUE
    
```

Where we used the periodicity $Q(IM+1)=Q(1)$.

In the meridional (J) direction we compute our difference approximations in stages in order to make maximum use of each line of latitude of a typical variable when it is in core. The fourth order difference approximation to $\frac{\partial Q}{\partial \phi} (i\&l j\Delta\phi)=D(i,\ell,j)$

$$(1) \quad \frac{4}{3} \left(\frac{Q_{i\&l j+1} - Q_{i\&l j-1}}{2\Delta\phi} \right) - \frac{1}{3} \left(\frac{Q_{i\&l j+2} - Q_{i\&l j-2}}{4\Delta\phi} \right)$$

Thus we see that $Q_{i\&l j}$ will be needed in the difference approximations to $\frac{\partial Q}{\partial \phi}$ for $j-2, j-1, j+1, \text{ and } j+2$. The corresponding code is

```

DO 20 I = 1, IM
QFLUX1= 4*(Q(I,L,JP1)+Q(I,L,J))
QFLUX2=-.5*(Q(I,L,JP2)+Q(I,L,J))
D(I,L,JP2)= D(I,L,JP2)-QFLUX2
D(I,L,JP1)= D(I,L,JP1)-QFLUX1
D(I,L,J) = D(I,L,J)+QFLUX1+QFLUX2
20 CONTINUE
    
```

For simplicity the array D is initialized to zero and Q is scaled so that (1) contains no divisions.

At the poles ($j=1$ or $J=JNP=46$) we use the values given in section IV on boundary conditions. We have special code for these cases denoted $J=2$ or $J=JM$ corrections.

COMP2 Description

This routine contains the Coriolis force term, the geopotential calculation (which should be made into a separate subroutine), and the pressure gradient and energy term calculations.

The geopotential PHI is dimensioned PHI(72,9,5) since we only need at most five storage locations for any computation. We use modular arithmetic (MOD5) to compute the indices JMOD, JP1MOD, JP2MOD which correspond to the standard index values of J, J+1, and J+2. Thus PHI(I,L,6) is stored in PHI (I,L,1), and we avoid shifting array values by using the JMOD index as a pointer.

For the south pole calculation we need geopotential values at J=2 and 3, thus, for J=1 we compute PHI for J=1,2, and 3. For successive values of J we need only compute PHI at JP2 which is needed in the pressure gradient calculation at J. Therefore, the calculation of PHI and the associated array PK are coded for calculation at JP2. Except for the first J value we are only computing the geopotential at one latitude value for each pressure gradient calculation.

X: Flow Charts

COMPO (Q,QT)

```

JS1=1
JS2=1

((MAIN LOOP))
DO 10 J=1, JM

((COMPUTE ALL J-PARAMETERS))

IF (JP2>JM)                                GO TO 25
IF (J>1)                                    GO TO 18

JP2=JP2MOD=2
((SAVE QTPOL(M) IN QSMPOL(M), M=1,2))

18 ((SAVE QJ(JP2) IN QSM(JP2MOD)))

IF (JP2>2)                                GO TO 25
((INCREMENT JP2, JP2MOD))                  GO TO 18

25 CALL COMP1(Q,QT,J)
CALL COMP2(Q,QT,J)

((ELIMINATE NEGATIVE HUMIDITIES))

IF (PT<(400,1100.))                        STOP
IF (J<3)                                    GO TO 200
IF (J=3)                                    GO TO 70

29 ((UNSCALE QT(JS2)))
CALL AVRX(QT,JS2)

((LEAP FROG, MATSUNO AS RELATED TO THE
SEQUENCE OF 2 CALLS TO COMPO:
1st CALL ⇒ LF ⇒ LF or MP ⇒ MC,
2nd CALL ⇒ LF ⇒ LF or LF ⇒ MP or MC ⇒ MP))

IF (LF→LF)                                GO TO 45
IF (LF→MP.OR.MC→MP)                       GO TO 58

((STATUS: MP→MC OR MC→LF))
((P(I,JS2)+PSM(I,JS2MOD), Q+QSM))

IF (MC→LF)                                GO TO 63

(STATUS: MP→MC))

```

(contin.)

COMPJ (Q,QT)

```

58      ((P(I,JS2)+PT(I,JS2),Q-QT*DXYP*PT))  GO TO 64
45      ((P(I,JS2)+ *P(I,JS2)+α*(PSM+PT),
        Q-β*DXYP*P*Q+α*(QSM+QT*DXYP*PT)))
63      CONTINUE
        ((SOURCE TERM CORRECTION FOR LEAP FROG))
      IF (NOT COMP3 CALL)                      GO TO 67
        ((Q-Q-QT*DXYP*PT))
64      IF (NOT COMP3 CALL)                      GO TO 67
      IF (MATSUNO PREDICTOR STEP)              GO TO 67
        CALL COMP3 (QT,JS2)
        ((COMPLETE SOURCE TERM CORRECTION,
        Q-Q+QT*DXYP*PT))
67      IF (J<JM)                               GO TO 200
        ((INCREMENT JS2,JS2MOD))
      IF (JS2<JM)                               GO TO 29
70      ((POLES))
200     JS2=JS1
        JS1=J
10      CONTINUE
        ((END OF J LOOP))
      RETURN

```

COMP1 (Q,QT,J)

```

        ((COMPUTE JP2-JS2MOD))
      IF (J>2)                                  GO TO 2150
        JS1=JS2-JS1MOD=JS2MOD=1

```

(contin.)

COMPL (Q,QT,J)

```

2150 IF (J=JM)                                GO TO 2158
      ((COMPUTE PV1, PV2 = Vj* + Vj+2*))
      ----- ● -----
IF (J=1)                                      GO TO 2225

2158      ((COMPUTE PU1, PU2 = U1* + U1+2*))
      ((COMPUTE HORIZ.ADVEC.-IN LONG. DIREC.))
IF (J = JM)                                  GO TO 2237

2225      ((COMPUTE HORIZ.ADV.-IN LAT. DIREC.))
IF (J>1)                                      GO TO 2290
      ((J=2 CORREC. TO HORIZ. ADVEC & CONV.))
      ((CONV. CALC. FOR CONT. EQ.))

2290 IF (J<JM)                                GO TO 2405

2237      ((J=JM CORREC. TO HORIZ. ADVEC. & CONV.))
      ----- ● -----

2405      ((CONV + CONV * DSIG(L)))
IF (L<NLAY)                                  GO TO 2150
IF (J=1)                                      GO TO 2600
      ((COMPUTE SIGDOT,PT))
      ((COMPUTE VERTICAL ADVEC.))

IF (J<JM)                                    RETURN

2600      ((POLES, M=1 or 2))
      ((HORIZ. ADVEC.))
      ((CONT. E.Q.))
      ((SIGDOT PL, PTPOL))
      ((VERT. ADVEC.))
      RETURN

```

END OF COMPL

Alternative code from -----●----- to -----●-----

```

      -----●-----
IF (J.EQ.1)                                  GO TO 2222
      ((COMPUTE PU1, PU2))
      ((COMPUTE HA(I) ))
      GO TO 2235

2222      ((J=2 CORREC ))

2235 IF (J<JM)                                GO TO 2225
      ((J=JM CORREC))
      GO TO 2405

2225      ((HORIZ ADVEC(J) ))
      ((CONV.CALC.FORCONT.EQ.))

2405      ((CONV CONV*DSIG(L)))
      -----●-----

```

COMP2 (Q,QT,J)

```

((COMPUTE JP2=... ))
IF(J>1) ((JP2=JP2MOD=JPKP2=1)) GO TO 3001
:3005
3001 ((CORIOLIS ))
IF(J>JM) GO TO 3032
((FIRST MAIN LOOP IN L))
3005 ((DO 3030 LX=1,NLAY ))
((COMPUTE ENERGY TERM IN
THERMODYNAMIC EQN ))
((L=NLAY+1-LX ))
IF(L<NLAY) GO TO 3055
3007 ((COMPUTE PK(I,LL,JPKP2)FORLL=1,NLAY))
((COMPUTE PHI'(I,NLAY,JP2MOD)) :3065
3055 ((COMPUTE PHI' AT JP2MOD FOR L<NLAY))
3065 IF(JP2=1 or JP2=JNP) GO TO 3030
((COMPUTE W(JP2MOD)))
3030 ((INCREMENT LX ))
((IF J=1 GO TO 3005 AND COMPUTE PHI(2),W(2)))
(( 2nd MAIN LOOP IN L))
3032 ((DO 3031 LX=1, NLAY))
IF(J=1) GO TO 3111 ( POLES)
IF (J > 2) GO TO 3085
((J=2: VT,TT CORREC.))
3085 (((P) FOR U EQ.))
IF (J=2orJM) GO TO 3135
(((P) FOR V EQ.))
3135 IF (J < JM) GO TO 3031
(( CORREC. TO (P) J=JM))
3111 ((POLES))
3031 ((CONTINUE))
((RETURN))
END

```

A FOURTH-ORDER FORECASTING MODEL*

(E. Kalnay-Rivas, D. Hoitsma, and P. Anolick)

The GISS fourth-order model (Kalnay-Rivas, et al., 1977), which is a fourth-order, energy-conserving GCM on an unstaggered grid, had shown promising capabilities. It produced forecasts that showed an improvement over the second-order GISS forecasts with the same fine grid ($4^\circ \times 5^\circ$) resolution, but that were somewhat inferior to the "ultrafine" forecasts. However, the first version of the model required excessive amounts of computer memory and time for execution.

The model has been reprogrammed into the "fourth-order band model." The new program solves the primitive equations one latitude band at a time. The arrays are stored in an interlaced way, with all arrays being updated at the same latitude stored contiguously, and similarly for all arrays used in the computation of the time derivatives. This design of the program makes effective use of the virtual memory capability of the Modeling and Simulation Facility's IBM 370/165 or Amdahl computers. The virtual memory facility permits the execution of programs whose core size is larger than the one available, by placing the excess on disk and reading in those pages of information needed in the current calculations. The band fourth-order data structure and computations were constructed to optimize this virtual I/O process; a possible improvement may be to interlace also the arrays being updated with those used to compute the time derivatives. The use of the virtual memory facility avoids the explicit I/O used in the current Kern model, and yields a simpler program. (The band structure was suggested by G. Russell.)

The band fourth-order model computations have been optimized so that each time step is computed in the Amdahl in half the time required by the old model. The array structure has also been designed to reduce the amount of overall storage and high-speed memory by a factor of two (see Table 1).

Table 1. Comparison of Fourth-Order Model Computing Requirements.

	Original 4th-Order Model	Band 4th-Order Model
Core (bytes)	3500K	1500K
CPU time per step (COMP1, COMP2)	34 sec	17 sec

*Reprint from Atmospheric and Oceanographic Research Review-1978, G-AS, NASA Technical Memorandum 80253.

The model has been programmed so that it can use both the Matsuno and the leapfrog time schemes, the latter with the Robert time smoother. If Q^n and $D(Q^n)$ denote the fields at time $n\Delta t$ and the corresponding time derivative computed from the space differences, the smoothed leapfrog scheme is

$$Q^{n+1} = \bar{Q}^{n-1} + 2\Delta t D(Q^n)$$

$$\bar{Q}^n = Q^n + .5v(Q^{n-1} + Q^{n+1} - 2Q^n)$$

with $\bar{Q}^0 = Q^0$. The use of a smoothing coefficient requires a slightly smaller time step. For example, with $v = .1$, the model is marginally stable with $\Delta t = 288$ sec., compared to $\Delta t = 300$ sec. for the Matsuno and leapfrog schemes.

The smoothed leapfrog scheme is further modified to include the subgrid "physics" terms, and scaling and spatial smoothing procedures. Since the "physics" is called every few time steps, unless the leapfrog scheme is restarted after every call to the "physics," only one of the two consecutive fields will be affected. The restarting procedure is time-consuming, so it has been replaced by the following algorithm

$$Q^{n-1} = Q^{n-1} - Q^n$$

$$\tilde{Q}^n = Q^n + S^n$$

$$Q^{n-1} = Q^{n-1} + \tilde{Q}^n$$

where S^n corresponds to the "physics" terms. This procedure, which ensures that the physics is applied to two consecutive time steps, has been tested with good results.

The model can be extended into an "ultrafine" version in a straightforward way.

Preliminary Results. The first numerical integrations performed with the new band fourth-order model show dramatic improvements over forecasts made with the GLAS model with the same resolution ("fine grid" or 4° latitude by 5° longitude). The quality of the forecasts is now comparable with those produced with the "ultrafine" (2.5° latitude by 3° longitude) version of the model. Results from a 3-day numerical integration are presented in Figure 1. During an extended 8-day integration of the new fourth-order model, the atmospheric systems remained remarkably smooth, exhibiting a realistic behavior both with respect to position and intensities.

References

Kalnay-Rivas, E., A. Bayliss, and J. Storch, 1977: The fourth-order GISS model of the global atmosphere. Contributions to Atmos. Phys., 20, pp. 299-311.

The 4th Order GISS Model of the Global Atmosphere*

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Abstract: The new GISS 4th order model of the global atmosphere is described. It is based on 4th order quadratically conserving differences with the periodic application of a 16th order filter on the sea level pressure and potential temperature equations, a combination which is approximately enstrophy conserving. Several short range forecasts indicate a significant improvement over 2nd order forecasts with the same resolution (~ 400 km). However the 4th order forecasts are somewhat inferior to 2nd order forecasts with double resolution. This is probably due to the presence of short waves in the range between 1000 km and 2000 km, which are computed more accurately by the 2nd order high resolution model. An operation count of the schemes indicates that with similar code optimization, the 4th order model will require approximately the same amount of computer time as the 2nd order model with the same resolution.

It is estimated that the 4th order model with a grid size of 200 km provides enough accuracy to make horizontal truncation errors negligible over a period of a week for all synoptic scales (waves longer than 1000 km).

1. Introduction

It is generally accepted that the use of 4th order finite differences is more efficient in reducing space truncation errors than the use of higher resolution on a 2nd order model (KREISS and OLIGER, 1972). Linear analyses and shallow water type of experiments give *an upper limit* of the improvement that can be expected from the reduction of errors in the horizontal differences. For example, Table 1 corresponds to a linear wave equation with phase speed $c = 11 \text{ ms}^{-1}$, typical of atmospheric motions. It provides a measure of the "computational predictability period" after which the errors introduced by horizontal truncation alone become very serious. The table suggests that for waves longer than 2000 km a 4th order -- 400 km grid model is preferable to a 2nd order -- 200 km grid model. Waves shorter than 2000 km are forecast more accurately by a 2nd order -- 200 km grid model. In order to insure that horizontal truncation errors are small in the 1--3 week period of atmospheric predictability for all synoptic scale waves it is necessary to use either 4th order differences with a grid resolution of the order of 200 km or 2nd order differences with a grid of the order of 100 km. Numerical experiments with simple nonlinear models (WILLIAMSON and BROWNING, 1973, KALNAY-RIVAS 1976a, from now on I) also indicate dramatic reduction of errors by the use of 4th order differences.

All these studies indicate that a considerable improvement in forecasting skill is to be expected from the use of 4th order differences *if a substantial portion of the forecasting errors is due to horizontal truncation errors*. In actual numerical forecasts there are several other important sources of errors: vertical truncation errors, errors in the initial conditions, and poor "physics" (i.e. parameterization of physical processes like radiation, dissipation, subgrid transports, cumulus convection, boundary layers, etc.)

* (Paper presented during the DMG-AMS Meeting, Hamburg 1976; see Preface to Issue 1 - 2/1977)

■ **Table 1.** Flapsed time T after which a wave of wavelength L moving with a phase speed $c = 11 \text{ ms}^{-1}$ lags by more than 100 km due to space truncation errors.

	$\Delta x = 400 \text{ km}$	$\Delta x = 200 \text{ km}$
2nd order differences	L = 2000 km, T = 0.4 days L = 4000 km, T = 1.6 days	L = 2000 km, T = 1.6 days L = 4000 km, T = 6.4 days
4th order differences	L = 2000 km, T = 1.5 days L = 4000 km, T = 21 days	L = 2000 km, T = 21 days L = 4000 km, T = 328 days

At the Goddard Institute for Space Studies (GISS), New York, we have developed a 4th order general circulation model (GCM) with the expectation that it will yield more accurate short range forecasts, and probably more realistic climate simulations, but with the improvement limited by the other sources of errors. The model is described in Section 2. The results of several experimental short range forecasts are discussed in Section 3. A sample 36-hour forecast is presented and compared with forecasts by the standard 2nd order GISS model, which has the same resolution as the 4th order model, and by the "ultrafine" 2nd order GISS model, which has twice the horizontal resolution. Section 4 contains the conclusions and a discussion of future work.

2. Description of the model

The 4th order GISS global atmospheric model is a primitive equation model using longitude (λ), latitude (ϕ) and sigma (σ) coordinates. The basic equations and the parameterizations of physical processes are the same as those of the standard 2nd order GISS global model (Somerville et al, 1974). The finite difference scheme is quite different and is described in the following subsections.

2.1. Finite-difference scheme

The horizontal grid is uniform (constant $\Delta\lambda$ and $\Delta\phi$) and non-staggered. There are two ways to define such a grid, depending on whether variables are defined at the poles, as done by WILLIAMSON and BROWNING (1973), or half a grid size away from the poles (HOLLOWAY et al, 1973). We chose the former method because in the absence of smoothing near the poles it allows a time step twice as large as the latter. The singularity that spherical coordinates have at the poles is explicitly avoided by the use of a polar cell where "stereographic" velocities are used.

In the vertical direction we use a staggered grid with σ , the vertical velocity, defined at the boundaries of each layer, and all other variables in the center of the layer. The vertical grid can be non uniform although experiments have been made so far with constant $\Delta\sigma$. Experiments reported in I indicated that no significant improvement in accuracy was obtained when a 4th order staggered conservative scheme was used in the vertical direction. Therefore the model has 2nd order vertical differences. The computation of the geopotential Φ is performed as indicated by ARAKAWA (1972), with a modification suggested by PHILLIPS (1974).

a) Forecast equation away from the poles

Several systems of horizontal difference were tested as reported in I. The scheme chosen for the basis of these experiments consists of the simplest possible quadratically conservative 4th order differences.

We define the finite difference horizontal divergence operator at the vertical level k as

$$D_k(\mathbf{g}) = -\frac{1}{a \cos \phi} \left[\frac{4}{3} \delta_\lambda (\Pi u^\lambda \mathbf{g}^\lambda) - \frac{1}{3} \delta_{2\lambda} (\Pi u^{2\lambda} \mathbf{g}^{2\lambda}) + \frac{4}{3} \delta_\phi (\Pi v \cos \phi^\phi \mathbf{g}^\phi) - \frac{1}{3} \delta_{2\phi} (\Pi v \cos \phi^{2\phi} \mathbf{g}^{2\phi}) \right]_k \quad (2.1)$$

and the finite difference gradient operator as

$$\nabla \mathbf{g} = \frac{\mathbf{i}}{a \cos \phi} \left[\frac{4}{3} \delta_\lambda \mathbf{g}^\lambda - \frac{1}{3} \delta_{2\lambda} \mathbf{g}^{2\lambda} \right] + \frac{\mathbf{j}}{a} \left[\frac{4}{3} \delta_\phi \mathbf{g}^\phi - \frac{1}{3} \delta_{2\phi} \mathbf{g}^{2\phi} \right] \quad (2.2)$$

Here $\Pi = p_S - p_T$ is the difference between the pressure at the surface and the constant pressure at the top of the model, a is the radius of the earth, $\mathbf{w} = u\mathbf{i} + v\mathbf{j}$ is the horizontal velocity vector in spherical coordinates, and we use the finite difference notation

$$\delta_{n\lambda} \mathbf{g} = [\mathbf{g}(\lambda + n\Delta\lambda/2, \phi, \sigma, t) - \mathbf{g}(\lambda - n\Delta\lambda/2, \phi, \sigma, t)] / (n\Delta\lambda), \quad (2.3)$$

$$\mathbf{g}^{n\lambda} = [\mathbf{g}(\lambda + n\Delta\lambda/2, \phi, \sigma, t) + \mathbf{g}(\lambda - n\Delta\lambda/2, \phi, \sigma, t)] / 2,$$

and similar formulas for the other independent variables. With this notation, the continuity equation away from the poles is

$$\frac{\partial \Pi}{\partial t} = -D_k(\mathbf{w}) - \Pi \delta_\sigma \dot{\sigma}_k \quad (2.4a)$$

or, if we integrate it in the vertical and make use of the boundary condition

$$\dot{\sigma} = 0 \text{ at } \sigma = 0, 1,$$

$$\frac{\partial \Pi}{\partial t} = - \sum_{k=1}^K D_k(\mathbf{w}) \Delta \sigma_k \quad (2.4b)$$

where K is the number of vertical layers. Equation (2.4b) is the forecast equation for Π . The momentum equation is

$$\left(\frac{\partial \mathbf{w}}{\partial t} \right)_k = -D_k(\mathbf{w}) - \Pi \delta_\sigma (\dot{\sigma} \mathbf{w}^\sigma)_k + \left(f + \frac{u_k \tan \phi}{a} \right) \Pi \mathbf{k} \times (\Pi \mathbf{w})_k - \Pi \left[\nabla \Phi' + \frac{\sigma RT'}{p} \nabla \Pi \right]_k + \mathbf{F}_k, \quad (2.5)$$

The first law of thermodynamics is

$$\left(\frac{\partial \Pi T}{\partial t} \right)_k = -D_k(T) - \{ p^k \Pi \delta_\sigma [\dot{\sigma} (T/p^k)^\sigma] \}_k + \left\{ \frac{\Pi \sigma R T}{p} \left[\frac{\partial \Pi}{\partial t} + \mathbf{w} \cdot \nabla \Pi \right] \right\}_k + \frac{\Pi}{C_p} Q. \quad (2.6)$$

Here T is absolute temperature, \mathbf{F} is the frictional force, $\kappa = R/C_p$, Q is the diabatic heating per unit mass, p is the pressure and p^h is computed as indicated in Subsection 2.1c. The moisture equation for the water vapor mixing ratio q is

$$\left(\frac{\partial \Pi q}{\partial t}\right)_k = -D_k(q) - \Pi \delta_{jk} (\dot{q}^o)_k + \Pi (E - C). \quad (2.7)$$

E and C are the rates of evaporation and condensation.

In the momentum equation (2.5) we follow a device suggested by PHILLIPS (1974) to alleviate the difficulties of the computation of the pressure gradient in σ -coordinates in the vicinity of orography. We define $T' = T - T(p)$, $\Phi' = \Phi - \Phi(p)$, where $T = \theta p^h$, $\theta = 280 \text{ K}/(1000 \text{ mb})^h$, $\Phi = \Phi_0 \cdot C_p \theta p^h$ and $\Phi_0 = [(1000 \text{ mb})^h \cdot p^h]^{-1}$. We have simplified PHILLIPS' expression for Φ_0 since its precise value is not important. With this procedure, in the pressure gradient term there is an exact cancellation of the terms $\nabla \Phi + \frac{\sigma R T}{p} \nabla \Pi$, and this implies a significant reduction of truncation errors in regions with steep orography.

b) Forecast equations at the Poles

We have followed the method used by WILLIAMSON and BROWNING (1973) and define a polar cap of radius $a \Delta \phi$ on which we use "stereographic" (or rather "cartesian") velocity components defined by the transformation

$$\begin{aligned} U_{1p} &= -u \sin \lambda + v \cos \lambda \\ V_{1p} &= +u \cos \lambda + v \sin \lambda \end{aligned} \quad (2.8)$$

with inverse

$$\begin{aligned} u &= -U_{1p} \sin \lambda + V_{1p} \cos \lambda \\ v &= +U_{1p} \cos \lambda + V_{1p} \sin \lambda \end{aligned} \quad (2.9)$$

where the top and bottom signs correspond to the north and south poles respectively. The positive x -axis of the cartesian coordinates coincides with the meridian of longitude $\lambda = 0$. The difference in signs at the south pole is due to the choice of a right handed system of coordinates with the vertical unit vector pointing outwards. Formulas (2.8) and (2.9) are used to define "spherical" or "cartesian" velocity components wherever they are required in the finite difference equations.

The finite difference horizontal divergence operator at the poles is defined by an average over all longitudes λ_i :

$$\begin{aligned} D_{1p, k}(g) &= + \sum_{i=3}^{1+2} \left\{ \frac{4}{3} A_1 (\text{Hvg})_{\lambda_i} + \left(\frac{\pi}{2} - \Delta \phi\right) \right. \\ &\quad \left. - \frac{1}{3} A_2 (\text{Hvg})_{\lambda_i} + \left(\frac{\pi}{2} - 2 \Delta \phi\right) \right\}_k \end{aligned} \quad (2.10)$$

where $\lambda_i = (i-3) \Delta \lambda$, $\Delta \lambda = 2\pi/l$, and $A_n = \Delta \lambda \sin(n \Delta \phi) / [2\pi a^2 (1 - \cos(n \Delta \phi))]$. The finite difference gradient operator at the poles is also defined by an average over all longitudes:

$$\nabla_{1p} g = \frac{2}{1a \Delta \phi} \sum_{i=3}^{1+2} \left\{ \frac{4}{3} g_{\lambda_i} + \left(\frac{\pi}{2} - \Delta \phi\right) - \frac{1}{3} g_{\lambda_i} + \left(\frac{\pi}{2} - 2 \Delta \phi\right) \right\} \cdot \{\cos \lambda_i \mathbf{i} + \sin \lambda_i \mathbf{j}\} \quad (2.11)$$

Then the continuity equation at the poles is

$$\frac{\partial \Pi_{\pm p}}{\partial t} = \mp \sum_{k=1}^K D_{\pm p, k}(1) \Delta \sigma_k, \quad (2.12)$$

the momentum equation is

$$\begin{aligned} \left(\frac{\partial \Pi V}{\partial t} \right)_k &= \mp D_{\pm p, k}(V) - \Pi_{\pm p} \delta_\sigma \left(\dot{\sigma}_{\pm p} \bar{V}_{\pm p}^\sigma \right)_k + f_{\pm p} \mathbf{k} \times (\Pi V)_{\pm p, k} \\ &- \Pi_{\pm p} \left[\nabla_{\pm p} \Phi' + \frac{\sigma R T'}{p} \nabla_{\pm p} \Pi \right]_k + F_{\pm p, k} \end{aligned} \quad (2.13)$$

where $V_{\pm p} = U_{\pm p} \mathbf{i} + V_{\pm p} \mathbf{j}$. The thermodynamic equation is

$$\begin{aligned} \left(\frac{\partial \Pi T}{\partial t} \right)_{\pm p, k} &= \mp D_{\pm p, k}(T) - \{ p_{\pm p}^k \Pi_{\pm p} \delta_\sigma \left[\dot{\sigma}_{\pm p} \overline{\left(T_{\pm p} / p_{\pm p}^k \right)} \right] \}_k \\ &+ \left\{ \frac{\Pi_{\pm p} \sigma \kappa T_{\pm p}}{p_{\pm p}} \left[\frac{\partial \Pi_{\pm p}}{\partial t} + V_{\pm p} \cdot \nabla_{\pm p} \Pi \right] \right\}_k + \frac{\Pi_{\pm p}}{C_p} Q_{\pm p} \end{aligned} \quad (2.14)$$

and the moisture equation is similar.

c) Diagnostic equations

In σ -coordinates, the pressure is defined by $p = \sigma \Pi + p_T$. Following a suggestion by PHILLIPS (1974) we define at the center of a layer $p_k^k = 1/(\kappa + 1) \delta_\sigma p_k^{\kappa+1} / \delta_\sigma p_k$, instead of $p_k^k = (\bar{p}_k^\sigma)^\kappa$ as in ARAKAWA (1972). This formula was derived under the assumption that the potential temperature varies in a σ -layer much less than either the temperature or the pressure. PHILLIPS indicated that a more accurate relationship between temperature and geopotential can be expected from this formula.

The geopotential Φ is determined following ARAKAWA (1972): If Φ_S is the surface geopotential, then the geopotential at the center of the lowest layer is

$$\Phi_K = \Phi_S + \sum_{k=1}^K C_k T_k \quad (2.15)$$

and at other levels

$$\Phi_k = \Phi_{k+1} + C_p \Delta \sigma_k + \frac{1}{2} \delta_\sigma \left(p_k^k + \frac{1}{2} \right) \left(\frac{T_k + 1/2 \sigma}{p_k^k + 1/2} \right) \quad (2.16)$$

The coefficients C_k are determined from

$$C_k = \frac{\Pi \sigma_k R \Delta \sigma_k}{p_k} - \frac{C_p}{p_k^k} \left(\sigma \Delta \sigma \delta_\sigma p^k \right)_k \sigma \quad (2.17)$$

The vertical velocity in σ -coordinate is determined from (2.4a) and (2.4b), and the boundary condition $\dot{\sigma} = 0$ at $\sigma = 0, 1$:

$$\delta_\sigma \dot{\sigma}_k = \frac{1}{\Pi} \left\{ \sum_{k=1}^K D_k(1) \Delta \sigma_k - D_K(1) \right\} \quad (2.18)$$

d) **Boundary conditions**

In the east-west direction we use periodicity: $g(\lambda + 2\pi, \phi) = g(\lambda, \phi)$ for all variables. When the value of a variable is needed "beyond" the poles, we define it by continuation along the same meridian:

$$w\left(\lambda, \pm\left(\frac{\pi}{2} + \Delta\phi\right)\right) = -w\left(\pi + \lambda \pm\left(\frac{\pi}{2} - \Delta\phi\right)\right); g\left(\lambda, \pm\left(\frac{\pi}{2} + \Delta\phi\right)\right) = g\left(\pi + \lambda, \pm\left(\frac{\pi}{2} - \Delta\phi\right)\right),$$

where g represents all variables other than the two horizontal velocity components. In the vertical the top and bottom are material surfaces through which there is no flux ($\dot{\sigma} = 0$ at $\sigma = 0, 1$) except for subgrid boundary layer fluxes of momentum and heat included in F and Q .

2.2. Filtering near the poles and high order filtering

The CFL computational stability condition requires a very small time step unless linearly unstable short waves are filtered out near the poles. For this purpose several alternative procedures have been tried but so far the method found to give most satisfactory results is the fourier filtering of the prognostic variables u, v, T and *the indirect smoothing of Π through the filtering of the sea level pressure (SLP) field.*

The fourier filtering is performed polewards of 66° latitude. The amplitudes of the fourier components of zonal wavenumber n are multiplied by a transfer function which is 1 for $n \leq N$ and decreases linearly to zero between $n = N$ and $n = N + 5$. The number of retained modes is defined by $N(\phi) = \text{integer part of } (90 \cos \phi)$. We have tried filtering the stereographic velocities U, V instead of u, v but no improvement was obtained.

Since both the surface geopotential and the surface pressure fields contain large amplitude short wave components due to the presence of orography, an artificial smoothing of these fields represents a distortion of the real geometry of the boundary. On the other hand, the SLP is an intrinsically smooth field, and its high wavenumber components more closely represent *atmospheric waves*. Therefore in our model *the SLP is filtered near the poles and Π is recovered by solving the transcendental equation used to relate them.*

It was found that the 4th order model forecasts were less smooth than those of the 2nd order GISS model, especially in regions with steep orography. Based on the same considerations we have introduced a periodic application of a high (16th) order filter (SHAPIRO, 1970) on the SLP and potential temperature fields, which are not very affected by orography.

The filter is of the form $\bar{g} = \{1 - (F_\phi^2)^8\} \{1 - (F_\lambda^2)^8\} g$, where $F_\lambda^2(g_{ij}) = (g_{i+ij} - 2g_{ij} + g_{i-ij})/4$ and has a response $F_\lambda^2[\exp(ik\lambda)] = -\sin^2(k\Delta\lambda/2) \exp(ik\lambda)$. This filter eliminates waves shorter than $4\Delta\lambda$ and even after hundreds of applications has negligible damping effect on waves longer than $4\Delta\lambda$.

In the meridional direction there are three simple alternative forms for F_ϕ^2 :

$$F_{\phi I}^2(g_{ij}) = (g_{ij+1} - 2g_{ij} + g_{ij-1})/4$$

$$F_{\phi II}^2(g_{ij}) = [(g_{ij+1} - g_{ij}) \cos \phi_{j+1/2} - (g_{ij} - g_{ij-1}) \cos \phi_{j-1/2}] / (4 \cos \phi_j)$$

and

$$F_{\phi III}^2(g_{ij}) = [F_{\phi I}^2(g_{ij} \cos \phi_j)] / \cos \phi_j$$

FRANCIS (1975) used $F_{\phi II}^2$, which is the only one that conserves the area weighted average of g . However, we have found that the three meridional filters produced virtually identical results. This could

be expected because the waves affected by the filter are too short to be strongly influenced by the convergence of the meridians. Since $F_{\phi I}^2$ and $F_{\phi II}^2$ can be programmed as efficiently as F_{λ}^2 we find them preferable. In the model we use $F_{\phi I}^2$, with g continued "beyond" the poles as indicated in Subsection 2.1d. At the poles $\bar{g}_{\pm p}$ is defined as the average over all longitudes λ_i of the filtered values $\bar{g}_{i, \pm p}$. At the present we are applying the high order filter to the SLP and potential temperature fields once every two hours.

2.3. Enstrophy constraint

BAYLISS and ISAACSON (1975) have developed a simple method to make any finite difference scheme conservative with respect to any quantity. The method has been tested in our model by forcing conservation of enstrophy on the dry adiabatic version of the model, although in such a model it is potential enstrophy that is conserved. The procedure is the following: Let the vectors U and V represent the values of the velocity components discretized over the grid and let the functional $G(U, V)$ denote the consistent 4th order approximation of the mean square vorticity. At each time step a correction U', V' is added to the predicted values \bar{U}, \bar{V} such that $G(\bar{U} + U', \bar{V} + V') = G(U_0, V_0)$ where U_0, V_0 denote the velocities at time $t = 0$. Since this equation cannot be solved explicitly, it is linearized about the predicted values \bar{U}, \bar{V} :

$$\left(\frac{\partial G}{\partial U}\right)_{\bar{U}, \bar{V}} \cdot U' + \left(\frac{\partial G}{\partial V}\right)_{\bar{U}, \bar{V}} \cdot V' = G(U_0, V_0) - G(\bar{U}, \bar{V}) \quad (2.19)$$

If we assume

$$(U', V') = \alpha \left(\frac{\partial G}{\partial U}, \frac{\partial G}{\partial V}\right)_{\bar{U}, \bar{V}} \quad (2.20)$$

then α can be determined from (2.19):

$$\alpha = \frac{G(U_0, V_0) - G(\bar{U}, \bar{V})}{\left|\frac{\partial G}{\partial U}\right|^2 + \left|\frac{\partial G}{\partial V}\right|^2} \quad (2.21)$$

It may be easily shown that the choice (2.20) minimizes $\|U'\| + \|V'\|$, the norm of the correction vector. It has been found that this procedure improves the forecasting skill of the model. However, when it is used in combination with the periodic application of the high order filter the improvement is marginal. Therefore this option is not included in the standard version of our model.

2.4. Summary of the properties of the model

The finite differences of the model have the following properties:

- a) Horizontal differences are performed on a nonstaggered grid and have 4th order truncation errors.
- b) Vertical differences are performed on a staggered grid and have 2nd order truncation errors.
- c) The differences are conservative in the sense that in the absence of diabatic and dissipation, mass and energy are conserved except for marginal terms at the poles. Non-conservative differences similar to those used by WILLIAMSON and BROWNING (1973), *starting from real data*, proved to be nonlinearly unstable after less than one day of integration.
- d) Unlike the "box method", the horizontal differences remain 4th order near the poles (KALNAY-RIVAS, 1976b)

- e) The model contains *no horizontal diffusion* except for the possible use of a dissipative time scheme and for the periodic application of the high order filter. It has been shown in I that the combination of a 4th order quadratically conservative scheme with the *periodic* application of a high order filter *replaces successfully the use of an enstrophy-conserving scheme*. This is because waves shorter than four times the grid size are the ones subject to aliasing and to large truncation errors, and they are removed by the filter *before they attain finite amplitude*. Waves longer than four times the grid size are accurately computed by the 4th order scheme and are not affected by the filter.

3. Numerical forecasts

3.1. Analytical initial conditions

The model was tested using 3 dimensional extensions of both the nontrivial steady state solution and the Rossby-Haurwitz wave initial conditions used in I. The forecasts remained smooth during the several days of interaction. Errors in the steady state case were an order of magnitude larger than in I because of the use of single precision arithmetic on the IBM/360-95 computer. However, it was found necessary to use double precision for the longitudinally averaged pressure gradient at the poles in order to avoid excessive local error growth.

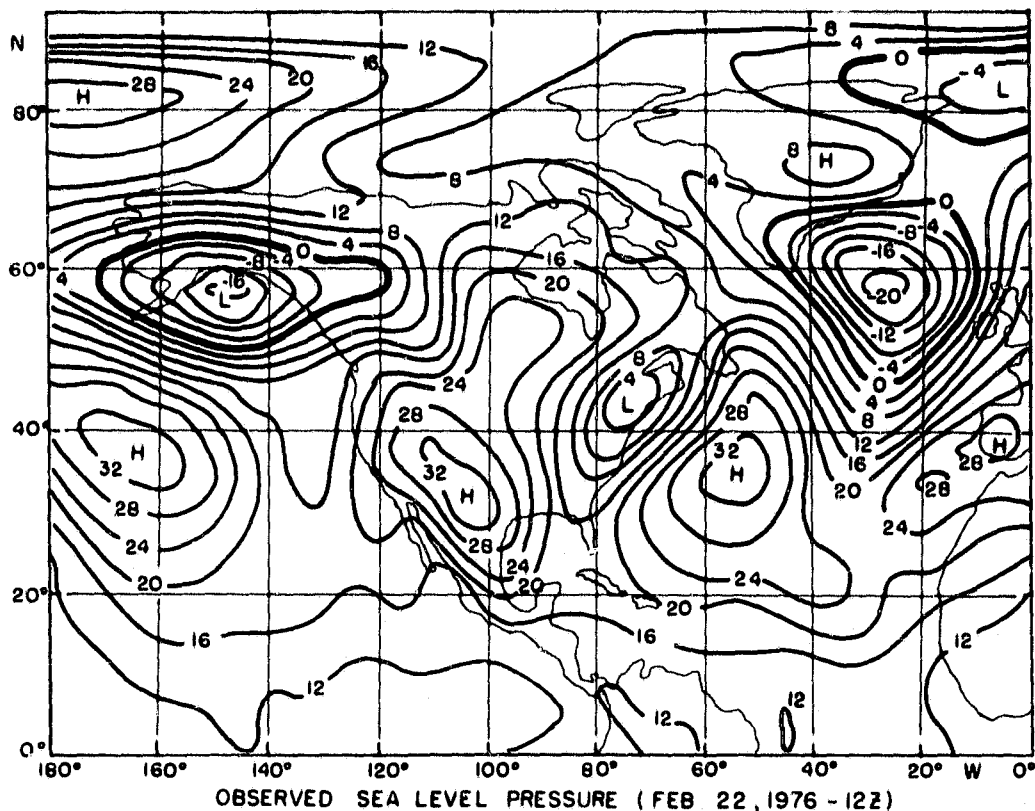
3.1. Forecasts with real initial data

As mentioned before, the parameterization of physical processes included in the terms **F**, **Q**, **C** and **E** of Equations (2.5) -- (2.7) has been adapted from the standard 2nd order GISS model (SOMERVILLE et al, 1974). The only change that has been introduced is the reduction by a factor of 2 of the surface drag coefficient. This was found to be necessary in order to avoid an excessive damping of the pressure systems after one or two day forecasts. We believe that the fact that in the staggered grid the surface winds are obtained by a horizontal average reduces the relative effect of friction in the standard model. At the present time (November 1976), we have performed 5 experimental short range forecasts. Since they have been made with slightly different versions of the 4th order model, we don't yet have a reliable measure of the model's forecasting skill. The initial data so far has been available in the staggered grid used by the standard model so that winds have been linearly interpolated to the nonstaggered grid. This is a source of error which may have adversely affected the 4th order forecasts.

We have performed a single 4-day forecast which indicated that the model remains very stable and that synoptic systems show no tendency to become unrealistically weak.

All the numerical 4th order forecasts show a significant improvement over the 2nd order forecasts performed with the same resolution. This improvement appears as a better estimation of the changes in position and intensity of several pressure systems. This is an encouraging result, especially in view of the study made by BAUMHEFNER and DOWNEY (1976) which indicates that the standard 2nd order GISS model forecasts compare favorably with those made by the NMC and NCAR models with similar resolution.

On the other hand, and contrary to our expectations, the forecasts made with the 2nd order GISS model with double horizontal resolution were found to be either similar or superior to the 4th order forecasts. This is more true in the sea level pressure than in the 500 mb forecasts. We consider that there are several possible reasons for this result. One is the extra errors introduced in the 4th order model initial

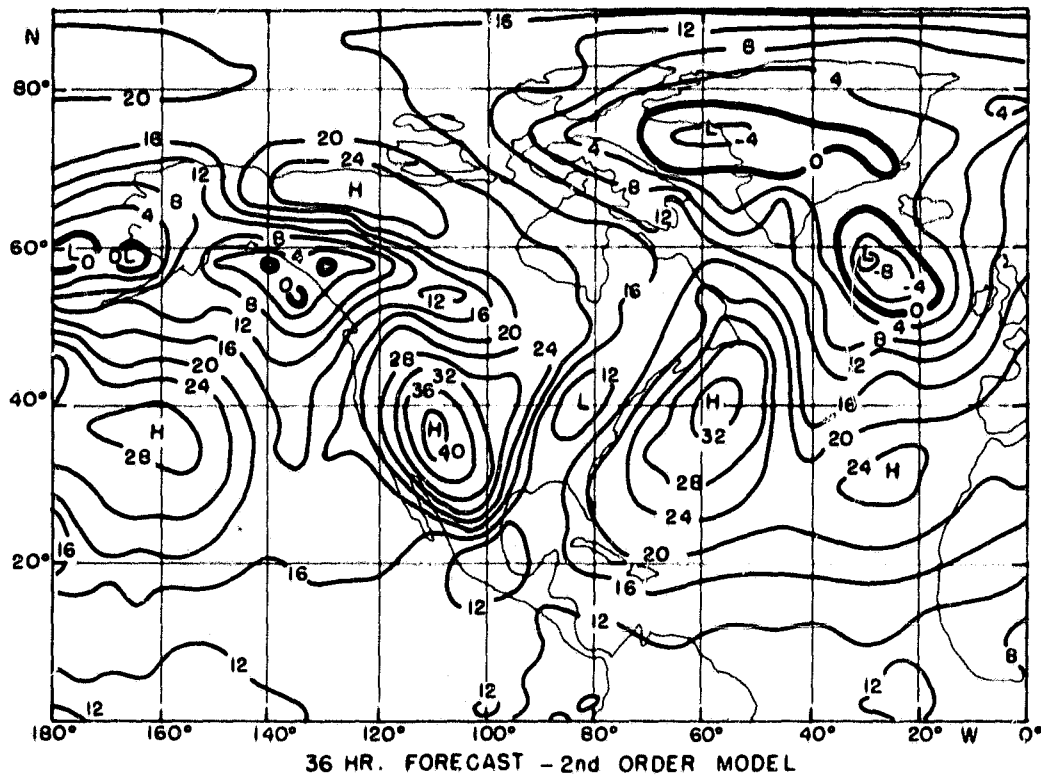


● Figure 1. Observed 36 hr. sea level pressure at 0000 GMT, 22 February 1976

data by the averaging of the winds. A second reason (and probably the most important one) is the existence of waves with significant amplitudes and wavelengths between 1000 and 2000 km. These short waves are better detected in the initial data and more accurately forecast by the 2nd order model with double resolution. The nonlinear interaction of these short waves with longer synoptic waves which according to linear theory are better forecasted by the 4th order model is clearly very important. A third important reason is the fact that errors introduced by the parameterization of subgrid processes become less important as the grid size is reduced.

We present an example of a 36 hour forecast of the sea level pressure corresponding to February 22, 1976, 12Z. Figure 1 is the verification sea level pressure map. Figure 2 is the forecast computed with the standard 2nd order GISS model, which has a resolution $\Delta\lambda = 5^\circ$, $\Delta\phi = 4^\circ$, $\Delta\sigma = 1/9$. Figure 3 is the 4th order forecast computed with the same resolution, and Figure 4 is the forecast obtained using the "ultrafine" 2nd order GISS model which has a resolution $\Delta\lambda = 2.5^\circ$, $\Delta\phi = 2^\circ$, $\Delta\sigma = 1/9$.

The most striking improvement showed by the 4th order forecast is in the position and intensity of the low in the eastern portion of North America. In fact, this low and the high over the Atlantic Ocean have been forecasted better by the 4th order model than by the 2nd order model with double resolution. On the other hand the developing cold high in central Canada and the intense low pressure center south of Greenland have been predicted better by the double resolution 2nd order model. As usual in this

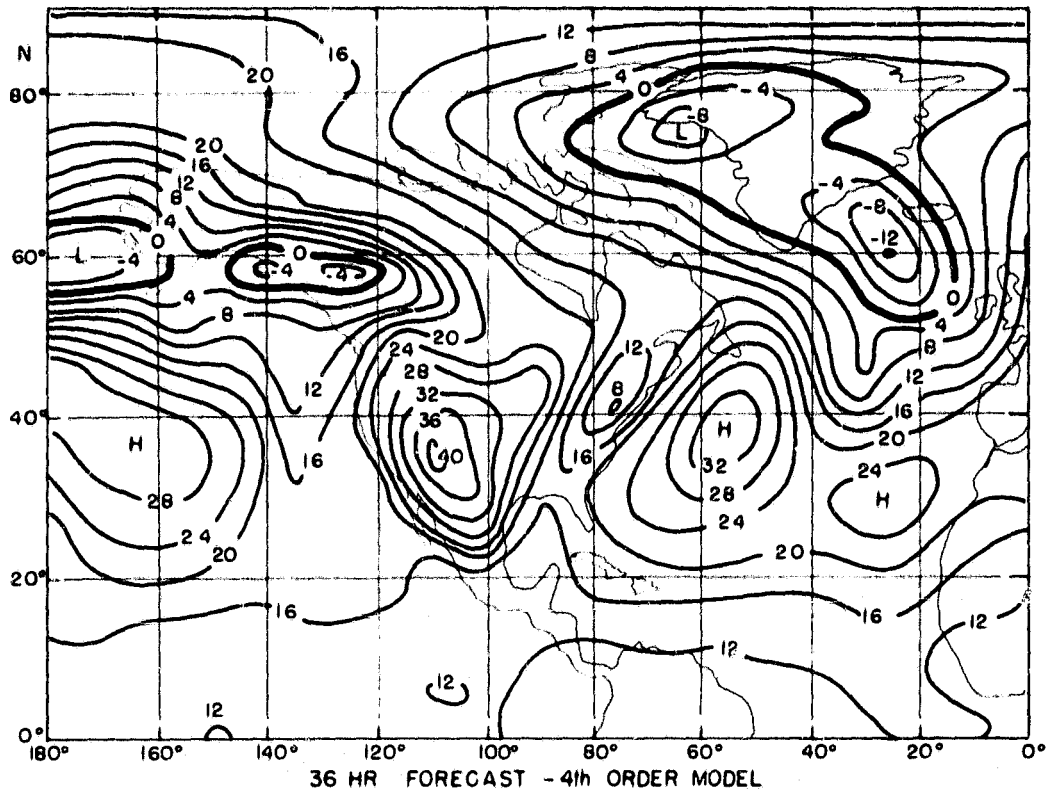


● Figure 2. 36 hr. forecast using the standard 2nd order GISS model

type of comparison, the three forecasts share several important deficiencies. For example the intensity of the high over southwest USA has been overpredicted by the three models, probably because of difficulties introduced by the orography. The low south of Alaska has been erroneously split by the three models. This splitting, which is slightly less pronounced in the 4th order forecast, may be due to both errors in the initial data and orographic and coastal problems.

4. Summary and conclusions

We have described the characteristics of the 4th order GISS model of the global atmosphere. It is based on a quadratically conservative scheme with the periodic application of a 16th order filter on the sea level pressure and potential temperature fields. As shown in I this combination is approximately enstrophy-conserving. An operation count of the numerical schemes indicates that with similar code optimization the 4th order model will require approximately the same amount of computer time as the 2nd order enstrophy-conserving GISS model with the same resolution. We also plan to introduce a simplified semi-implicit scheme, and to study the possibility of using the combination of a nonstaggered vertical grid with the Kreiss 4th order method described in I.

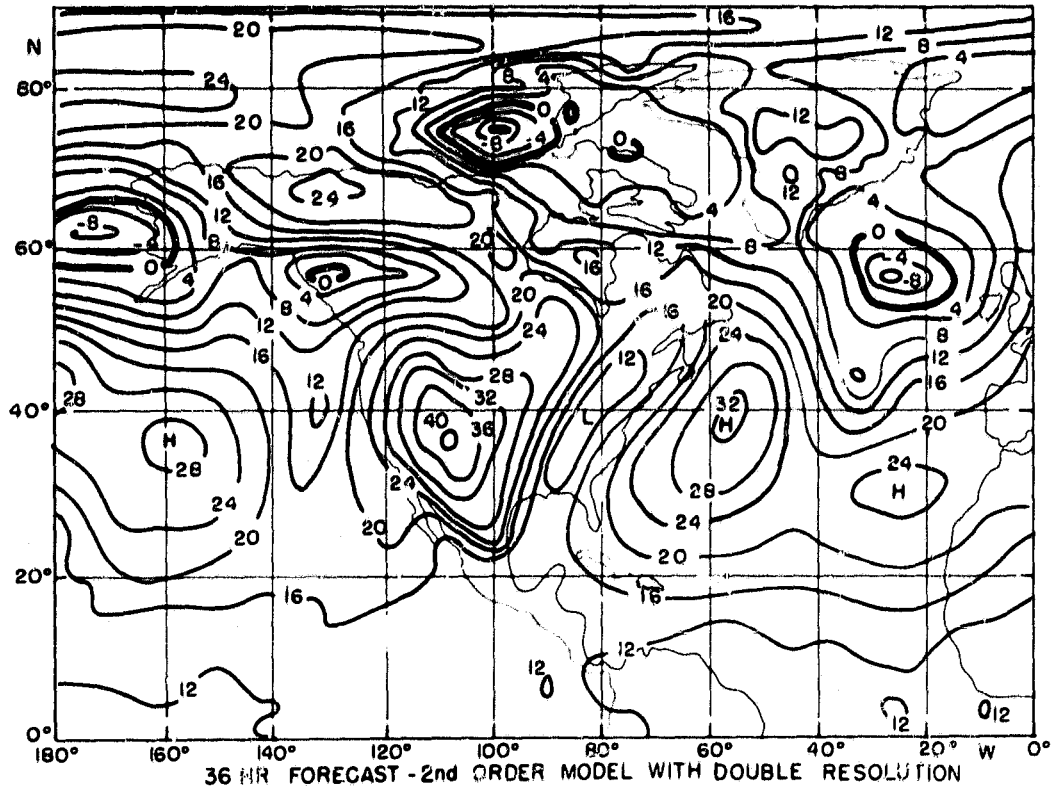


● Figure 3. 36 hr. forecast using the 4th order GISS model with the same resolution as in Figure 2

The results of several short range forecasts indicate a significant improvement over the 2nd order forecast with the same resolution. This improvement is shown in the estimations of changes in position and intensity of several pressure systems. We plan to study the impact that the greater accuracy of 4th order differences has on the forecasting skill of variables of more practical importance, such as temperature and precipitation.

It has been found that the 4th order forecasts are somewhat inferior to the forecasts made with a 2nd order model with double horizontal resolution. We consider that one of the most important reasons for this result are the presence of waves in the range between 1000 km and 2000 km which are computed more accurately by the high resolution 2nd order model than by the 4th order model. Another important reason is that errors introduced by the parameterization of subgrid processes become smaller as the size is reduced.

It is estimated that the 4th order model with a grid size of 200 km provides enough accuracy to make horizontal truncation errors negligible over a period of a week for all synoptic scales (waves longer than 1000 km).



● Figure 4. 36 hr. using the "ultrafine" 2nd order GISS model with double horizontal resolution

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References

- ARAKAWA, A. 1972: Design of the UCLA GCM. Techn. Rept. No. 7, Dept. of Meteorology, UCLA.
- BAUMHEFNER, D., and P. DOWNEY, 1976: Forecast intercomparisons between large-scale numerical weather prediction models. *Annalen der Meteorologie*, 11. Simulation of large scale atmospheric processes (extended abstracts) 205-208.
- BAYLISS, A. and E. ISAACSON, 1975: How to make your algorithm conservative. *Notices of the Am. Math. Soc.*, August 1975.
- FRANCIS, P. E., 1975: The use of a multipoint filter as a dissipative mechanism in a numerical model of the general circulation of the atmosphere, *Q. J. Roy. Meteor. Soc.*, 101, 567-582.

- HOLLOWAY, J. L., M. Spelman and S. Manabe, 1973: Latitude-longitude grid suitable for numerical time integration of a global atmospheric model. Mon. Wea. Rev., 101, 69--78.**
- KALNAY-RIVAS, E. 1976a: Numerical experiments with 4th order conservative finite differences. Annalen der Meteorologie, 11. Simulation of large scale atmospheric processes (extended abstracts). Addendum.**
- KALNAY-RIVAS, E., 1976b: High latitude truncation errors of box-type primitive equation models. Mon. Wea. Rev., 104, 1066--1069.**
- KREISS, H.-O., and J. OLIGER, 1972: Comparison of accurate methods for the integration of hyperbolic equation. Tellus, 24, 199--215.**
- PHILLIPS, N. A., 1974: Application of Arakawa's Energy-Conserving Layer Model to Operational Numerical Weather Prediction. NMC Office Note 104.**
- SHAPIRO, R., 1970: Smoothing, filtering and boundary effects. Rev. Geophys. and Space Phys., 8, 359--387.**
- SOMERVILLE, R., P. H. STONE, M. HALEM, J. E. HANSEN, J. S. HOGAN, L. M. DRUYAN, G. RUSSELL, A. A. LACIS, W. J. QUIRK and J. TENENBAUM, 1974: The GISS model of the global atmosphere. J. of Atmos. Sci., 31, 84--117.**
- WILLIAMSON, D. L. and G. L. BROWNING, 1973: Comparison of grids and difference approximations for numerical weather prediction over a sphere. J. Applied Met., 12, 264--274.**

THE EFFECT OF ACCURACY, CONSERVATION AND FILTERING
ON NUMERICAL WEATHER FORECASTING

ORIGINAL PAGE IS
OF POOR QUALITY

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1. INTRODUCTION

The design of a numerical model for atmospheric simulation is not a straight-forward procedure. Both in the areas of mathematical and numerical analysis, and in the parameterization of physical processes not explicitly resolved, the modeler faces several difficult choices between equally reasonable methods, and sometimes between similarly unsatisfactory methods.

In this paper we discuss the considerations leading to the numerical design of the GLAS Fourth-Order Global Atmospheric Model. This model, which was briefly described in Kalnay-Rivas et al., [1977], has been restructured, and several minor changes were introduced. *The computation time and memory requirements for the 4th order model are now similar to those of the present second order GLAS model with the same 4° latitude, 5° longitude and 9 vertical-level resolution [Somerville et al., 1974]. However, the fourth-order model forecast skill is significantly better than that of the current GLAS model, and after 3 days it is comparable or better than that obtained with the 2.5° by 3° version of the GLAS model.*

A discussion of several of the basic characteristics of the model design is contained in section 2. For each of them we present some of the possible alternatives, their advantages and disadvantages, and the reason for our choice. In section 3, we discuss the effect on numerical forecasts of changes in the accuracy, resolution and conservation properties of the models. ⁽¹⁾ Section 4 contains some final remarks.

(1) As of this writing, several of the numerical experiments are not complete.

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2. DISCUSSION OF THE DESIGN OF THE MODEL

Different numerical analysts and atmospheric modelers often take different approaches in the design of a numerical model for weather prediction. Sometimes they even have different basic philosophies. For example Arakawa [1966, 1972], has been a pioneer in the development and use of numerical schemes that reproduce as closely as possible the conservation properties of the continuous equations of fluid dynamics that determine the motion of the atmosphere. On the other extreme, Kreiss and Olinger [1972, 1973] have advocated the use of more accurate schemes even when they don't formally satisfy any conservation properties.

The current GLAS model is based on Arakawa's [1972] second order scheme using a staggered grid B. The scheme is energy conserving and approximately enstrophy conserving for non-divergent flow. In this section we isolate the areas in which the Fourth-Order model is different from the GLAS model, discuss some of the alternatives, and the justification of our choice.

2.1 Accuracy

There is a consensus among modelers that for finite difference models with second-order accuracy, horizontal resolution of about 400 km, and about 10 vertical levels, horizontal truncation errors are the most important source of errors. The truncation errors can be reduced by any of the following methods:

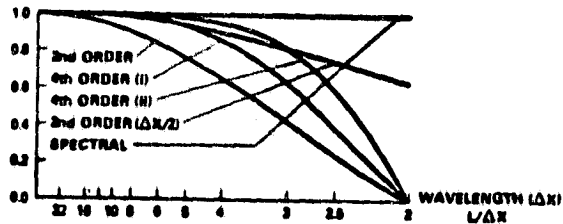
a. Increased horizontal resolution, retaining second-order differences.

Advantages: If the horizontal grid size is reduced by a factor of 2, truncation errors are reduced by a factor of 4. A comparison of the effect of truncation errors on the computational phase speed and group velocity of a linear wave is presented in Fig. 1. Another advantage is that increasing

the resolution allows smaller but possibly important scales to be explicitly included in the model.

Disadvantages: The reduction of error is slow, and the computation time is increased by a factor of 8.

COMPUTATIONAL PHASE SPEED c/c



COMPUTATIONAL GROUP VELOCITY c_g/c

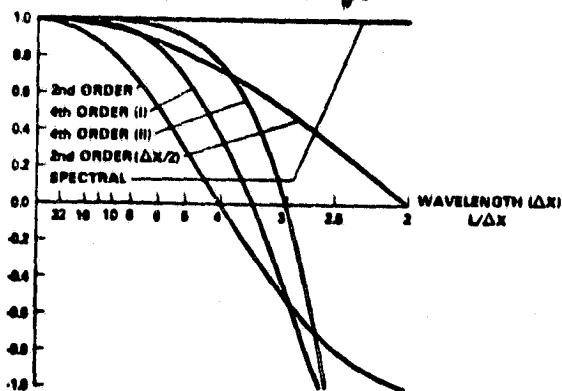


Fig. 1 Computational phase speed and group velocity for a linear wave of constant phase speed c , for different numerical schemes as a function of wavelength L .

b. Fourth order schemes of the first kind. We include in this group the common 5-point explicit fourth order schemes [Kreiss and Olliger, 1972].

Advantages: Phase speed errors are much smaller, especially for long waves (Fig. 1). Waves longer than $5\Delta x$ have smaller errors in their phase speed with fourth order differences than with second order differences with twice the resolution. If we consider the computational group velocity, the crossover occurs at about $7\Delta x$. The computational time is only increased by a factor between 2 and 3. **Disadvantages:** Waves in the range of $2\Delta x$ to $4\Delta x$ are still grossly misrepresented.

c. Fourth order schemes of the second kind. In this group we include linear finite elements, cubic splines and the "compact" or Pade type 3-point 4th order differences introduced by Kreiss [Orszag & Israeli, 1972], all of which have the same fourth order accuracy.

Advantages: These fourth-order differences are considerably more accurate than the fourth-order differences of the first kind (Fig. 1). The crossover with double-resolution second-order differences in the phase speed error occurs at $3\Delta x$, and in the group velocity error at about $4\Delta x$.

Disadvantages: In the simplest case, Kreiss' fourth order differences require the solution of tri-diagonal matrices. Finite element schemes require the solution of at least block tri-diagonal matrices. Even though there are efficient methods to perform these inversions, they are still computationally expensive. The extra accuracy can be compensated with 4th order differences of the first kind by increasing the resolution, which has other advantages, as we mentioned earlier. These schemes, as well as in higher order finite difference schemes, waves in the range of $2\Delta x$ to $4\Delta x$ are still grossly misrepresented.

d. Spectral schemes.

Advantages: If the basis of the spectral expansion are the eigenfunctions of the wave equation, spectral schemes have no phase speed errors. Because of this they require less resolution than finite difference schemes. **Disadvantages:** Because of the large number of computations required for the nonlinear terms, spectral schemes are competitive with finite difference schemes only in combination with the use of less resolution and semi-implicit time schemes.

Our choice: We chose to use fourth order finite differences of the first kind because they are computationally efficient and have small truncation errors except in the range of waves with wavelengths between $2\Delta x$ and $4\Delta x$.

2.2 Type of Grid

Both staggered and unstaggered grids have been widely used by atmospheric modelers.

a. Unstaggered grid: The advantage of this grid is its simplicity. Higher order schemes are easily developed with this grid. Its disadvantage is that all centered differences have to be computed over a distance of $2\Delta x$.

b. Staggered grids: Several staggered grid configurations are possible as reviewed by Arakawa [1972]. The one he called scheme C, which is the most commonly used, has the pressure defined at the center of a grid cell, and the velocity components u and v defined at their corresponding normal walls. This grid has the advantage that the pressure gradient and velocity divergence terms are computed over a distance of only $1\Delta x$, so that inertia gravity waves are computed with double resolution. Therefore, as pointed out by Arakawa, geostrophic adjustment is best represented in this grid. On the other hand, advection terms are computed with no more accuracy than in the unstaggered grid, and in the Coriolis' acceleration term, it is necessary to take horizontal averages of the velocities. The

higher resolution of inertia gravity waves reduce the maximum time step for explicit time schemes by a factor of two. Full fourth order schemes can be developed with staggered schemes but they are very involved [Kalnay-Rivas, 1976]. To date, modelers using staggered grids have introduced fourth order differences *only* in the advection terms.

Our choice: During extended range forecasts, second order errors in the non-advective terms may become important. For this reason we chose to use an unstaggered grid that allows the use of a simple, full fourth-order scheme.

2.3 Conservation Properties and the Use of Horizontal Diffusion

With respect to conservation properties, there are basically 3 types of finite difference schemes for the primitive equations:

a) Nonconservative schemes, the simplest of which is the one based on the advective form of the equations; b) quadratically or energy conserving schemes; and c) enstrophy conserving schemes. Advective and quadratically conservative schemes can be easily developed using staggered or unstaggered grids [Lilly, 1965; Bryan, 1966]. Enstrophy conserving schemes for the primitive equations have been developed on a grid C by Grammelvedt [1969] and Arakawa and Mintz [1974]. Sadourny [1965a, b] constructed a potential enstrophy conserving scheme on grid C, and Arakawa [1978] has recently developed a potential enstrophy and energy conserving scheme also on grid C.

Nonconservative schemes require a procedure to damp waves shorter than $4\Delta x$, which otherwise grow spuriously causing catastrophic nonlinear instability [Phillips, 1959]. This has usually been done by means of linear or nonlinear horizontal diffusion, or by using dissipative numerical schemes such as the Lax-Wendroff, or schemes that contain explicit horizontal averaging.

Quadratically conservative schemes avoid the unbounded growth of the solutions associated with catastrophic nonlinear instability. However, as Arakawa [1972] and Sadourny [1975a, b] have pointed out, in the course of long integrations, there is still a spurious build up of energy in the shortest waves, which appears as an *unbounded growth of the total enstrophy*. In the absence of horizontal diffusion, this type of *slow nonlinear instability* will completely distort the solution. Enstrophy conserving schemes, on the other hand, impose a stronger constraint on the growth of the smallest scales present in the model. For this reason, the UCLA and the GLAS models, which use enstrophy conserving schemes, *do not need to include horizontal diffusion*.

It should be emphasized that conservation of enstrophy does not necessarily imply a more accurate or realistic simulation. In the real atmosphere, the constraint of quasi-geostrophic motion implies that very little of the energy generated in the baroclinically unstable scales can reach the smallest scales

and be eventually dissipated [Charney, 1972]. In a numerical model, the finite resolution imposes an artificial "wall" at the short end of the spectrum, inducing an excessive accumulation of energy in the shortest waves. This problem is worst in nonconservative schemes, but it appears even in alias-free, energy- and enstrophy-conserving spectral models. This justifies using some parameterization of the unresolved subgrid eddies to withdraw energy from the smallest resolved scales.

For this purpose, Leith [1972] suggested the use of nonlinear horizontal diffusion in which the eddy diffusion coefficient is proportional to the local gradient of vorticity. This formulation is consistent with the transfer of energy to subgrid scales in two-dimensional turbulence. In another widely used formulation, suggested by Smagorinsky [1963] and based on a three-dimensional turbulent cascade theory, the diffusion coefficient is proportional to the deformation tensor. These formulations are better than the use of linear diffusion, but they both share the following problems: a) the diffusion coefficient is computed inaccurately for the shortest waves, and, more importantly, b) when the diffusion coefficient is large enough to avoid the spurious growth of the smallest scales, it produces excessive damping of the larger scales [Merileea, 1975; Williamson, 1978]. Furthermore, at the short end of the spectrum (scales of the order of 100 km), neither quasi-geostrophic nor 3-dimensional isotropic turbulence theories are really justified. Williamson [1978] generalized a higher order diffusion of the form $\nabla^2 \kappa \nabla^2$, suggested by Kreiss and Oliger [1972], still using the deformation type of diffusion coefficient. This formulation has the advantage that, because it is more scale dependent, there is less damping of the longer waves.

In our model, we have taken an approach closer to Fourier filtering the shortest waves, as suggested by Phillips [1959]. Our "subgrid parameterization" is based on the following argument: The 4th order scheme is adequate accurate for waves longer than $4\Delta x$, but grossly inaccurate for waves between $2\Delta x$ and $3\Delta x$. Since the shortest waves *cannot provide any useful information in a finite difference scheme*, we filter them out of the system while their amplitude is still small. Even though they are filtered out, the shortest waves still play an important role in the model: they act as a buffer or "sponge layer" in the spectrum domain, allowing energy to trickle down from longer waves and avoiding the accumulation of energy that would otherwise occur at the short wave cutoff.

The elimination of the short waves is performed in the model with the periodic use of a 16th order Shapiro [1970] filter, first introduced in GCMs by Francis [1975]. Figure 2 indicates the response of Shapiro filters of order 4, 8 and 16. It can be observed that in the case of a 16th order filter waves longer than 4 are scarcely affected at all

even after 128 applications, which in our model correspond to a 10 day integration. Waves shorter than 4 are virtually eliminated. Lower order filters introduce too much decay at long scales.

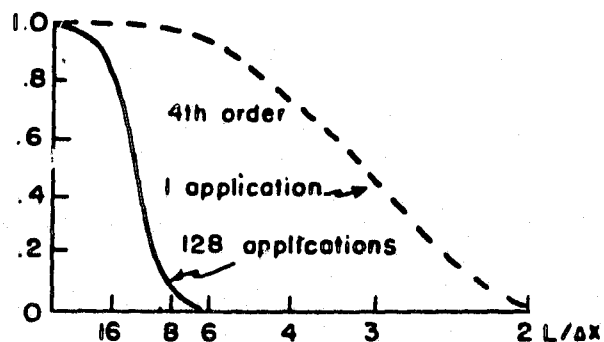
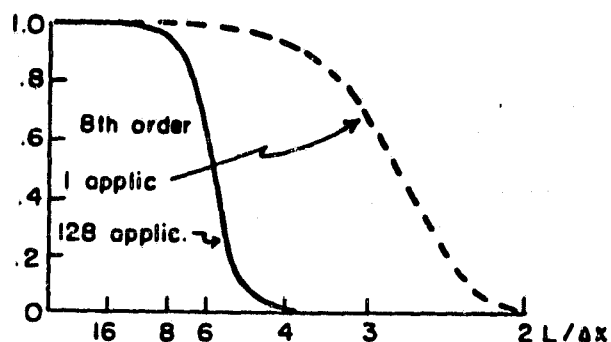
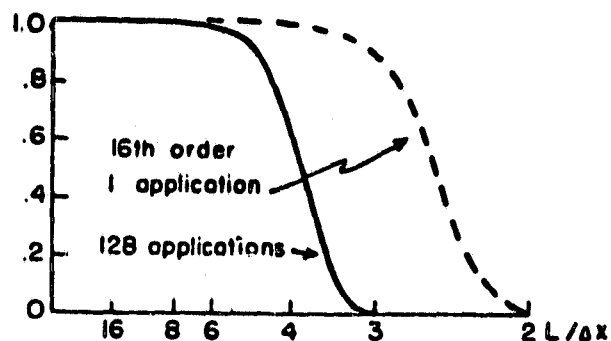


Fig. 2 Response of Shapiro filters after 1 and after 128 applications.

In order to explore the effect of filtering short waves we performed a series of 20 day forecasts with a simple shallow water equation model [Kalnay-Rivas, 1976], using different combinations of schemes, smoothing operators and frequency of application as indicated on Table 1.

A1: 4th order, non-conservative, no smoothing.
 A2: " , " , linear diffusion, $\nu = 10^3 \text{ m}^2/\text{s}$
 A3: " , " , 16th order filter/4 hrs.
 A4: " , " , " /time step.

B1: " , quadrat. conserv., no smoothing.
 B2: " , " , 16th order filter/4 hrs.
 B3: " , " , 8th " " "
 B4: " , " , 4th " " "

C: 2nd order, " , no smoothing

TABLE 1: Characteristics of the different runs in Fig. 3.

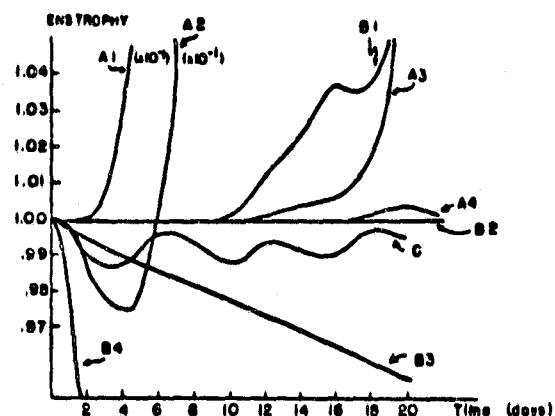
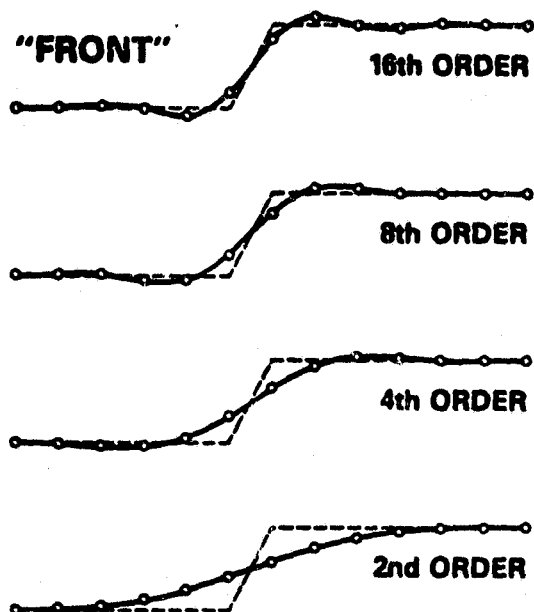


Fig. 3 Variations of total enstrophy during numerical integrations as indicated in Table 1. The scale corresponding to experiments A1 and A2 is multiplied by 10^{-1} .

We found that all stable runs conserved total energy with a high degree of accuracy. Fig. 3 shows the variation in time of the total potential enstrophy, which is conserved exactly in the continuous equations. The results indicate that when the 16th order filter is applied every time step (10 minutes) even a formally nonconservative scheme conserves both total energy and total potential enstrophy during a long integration (Run A4). The quadratically conservative scheme controls better the amount of energy going into the smallest scales, so that in Run B2 it was enough to apply the filter every 4 hours to conserve potential enstrophy within 0.05%, even though such conservation is not formally guaranteed in the scheme.

It may be questioned whether the application of a high order filter eliminates small scale features like frontal zones or the effects of orographic or cumulus convection forcing on small scales. Fig. 4 indicates that this is true for lower order filters. However a 16th order filter eliminates only those components which are not resolved anyway, and still allows for the formation of sharp gradient zones and strong local maxima.

10 PASSES (ONE DAY) OF SHAPIRO FILTER



10 PASSES (ONE DAY) OF SHAPIRO FILTER

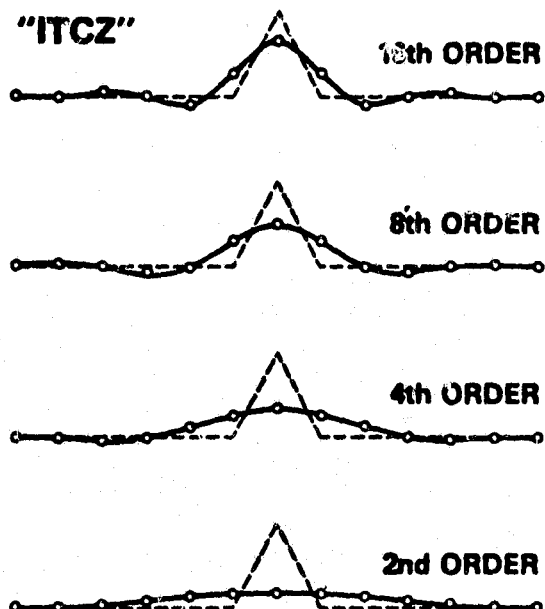


Fig. 4 Effect of 10 applications of a 16th order Shapiro filter on a) a step function b) a spike.

Our choice: We chose to use a quadratically (energy) conserving scheme because it is only slightly less efficient than an advective scheme and it requires the application of a high-order filter only every few hours to avoid the effect of slow nonlinear instability associated with a spurious growth of enstrophy. We apply the filter to the sea level pressure and potential temperature fields in order to compensate for the effect of topography. (3)

3. EXPERIMENTAL FORECASTS: PRELIMINARY RESULTS

We plan to perform an extensive series of forecasts to study the effect of using different schemes and varying resolution on the quality of actual weather forecasts. In this section we present some preliminary results.

We tested the Fourth Order model by making 3-day forecasts from several initial conditions. In every case the model performed much better than the GLAS model with the same 4° by 5° resolution. After 3 days the forecasts were comparable to the 2.5° by 3° version of the GLAS model in the sea level pressure maps, and had slightly less phase errors in the 500 mb maps.

Fig. 5a shows the verification sea level pressure map corresponding to a 3-day forecast with February 19, 1976, 0Z (a case that has been studied in detail by Atlas et al., 1979). Fig. 5b is the 3-day forecast with the Fourth Order 4° by 5° model verifying on the same date. The excessive gradients, especially at high latitudes are due to a faulty computation of the ground temperature by the radiation routine used in that run. Figs. 5c and 5d are the forecasts generated by the GLAS models with 4° by 5° and 2.5° by 3° resolution respectively. Fig. 6 displays the corresponding 500 mb maps.

In Fig. 7a, we present the verification sea level pressure map corresponding to a 3-day forecast with February 1, 1976, 0Z initial conditions. The results of six different 3-day forecasts are shown in Figs. 7b to 7g. The forecast in Fig. 7b was computed with the new full Fourth-Order model, using the same resolution and parameterization of physical processes (except for the long-wave radiation routine) as in the 4° by 5° GLAS model (Fig. 7c). In Fig. 7d, the model was the same as in 7b, but full Second-Order accuracy was used. If we compare these three made with the same resolution, we see that the fourth order model did considerably better, especially in forecasting the development and motion of the cyclone southwest of Greenland. The two second-order forecasts are close to each other.

(3) The idea of filtering these fields is due to Dr. A. Bayliss.

Fig. 7 shows the forecast made with the 2.5° by 3° GLAS model starting from the NASA initial conditions with assimilation of satellite data [Atlas et al., 1979]. All other forecasts were made from NMC's Global Analysis initial conditions. The cyclogenesis forecast was poorer than with the fourth order model, but over continental North America the forecast was better. This may be due to the higher resolution or, possibly, improved initial conditions.

Fig. 7f presents the forecast made with NMC's 6-layer, 380 km resolution, which has resolution comparable to our 4° by 5° grid. For these initial conditions, NMC's forecast was better than the GLAS second order forecast, although it shared some of its errors (such as a spurious anticyclogenesis over the Great Lakes).

Fig. 7g shows a forecast made with a slightly different version of the fourth order model. The differences were as follows: In the 7g forecast we used a Matsuno time step, a surface drag that was .75 of that of the standard model, the Shapiro was applied every two hours, and a slightly different scheme for the vertical advection of moisture was used. In the 7b forecast we used a combined Matsuno-leap-frog scheme, the surface drag was the same as in the standard model and the filter was applied every hour. The two forecasts are quite similar, but from other experiments it seems that the positions of the oceanic lows west of Spain and south of Alaska were somewhat affected by the change in frequency of the Shapiro filter.

4. FINAL REMARKS

Although these are preliminary results, the new Fourth-Order model seems to have very good forecasting skill. Similar excellent results with fourth order schemes were reported by Campana [1978, 1979], and by Williamson (1978). Campana found that most of the improvement over the second order model was obtained just by introducing fourth order accuracy on the horizontal averages performed on the Shuman-Hovermale model. Campana also obtained that for 2-day forecasts, fourth order differences were important in the advection terms but not in the pressure and continuity terms. We made shallow water equation experiments that showed some deterioration of the solution after about 8 days when only the advection terms were written with fourth order accuracy. Since baroclinic models are much more unstable than the SWE, Campana's results may not hold during extended forecasts. We are performing experiments to study this possibility in our model. We are also repeating some of the forecasts using a vector invariant form of the momentum equations. The NMC 6-layer PE model used in Fig. 7f is based on such a formulation, and we want to determine if it affects the solution.

We are performing integrations with a 2.5° by 3° version of the Fourth-Order model. From linear theory, this should make horizontal truncation errors negligible for most synoptic scales for periods of a week or longer. Still, narrow atmospheric features like those due to sharp orographic forcing on the ITCZ will remain poorly resolved.

We have found the fourth order model to be extremely sensitive to the parameterization of physical processes. For example, it became unstable through excessive cooling near the surface, apparently due to a flaw in the radiation routine. None of the second order models was sensitive to this problem.

In this paper we have discussed the use of formally nonconservative schemes coupled with periodic filtering of the shortest waves as an alternative to the use of conservative schemes. It is clear from the experimental results that formal enstrophy conservation has had no beneficial impact on the quality of the forecast (Figs. 7c and 7d). Conservation of potential enstrophy [Arakawa, 1978] might have a more positive effect.

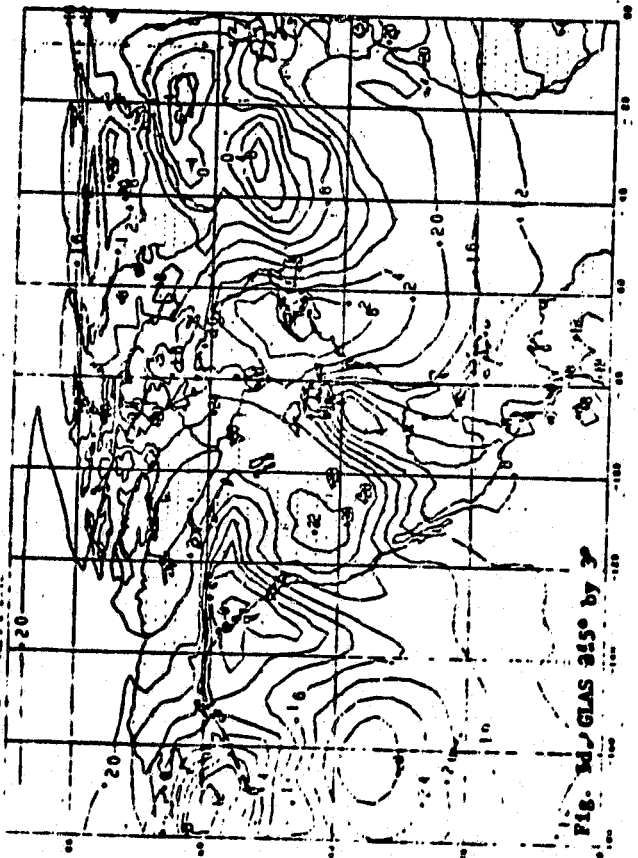
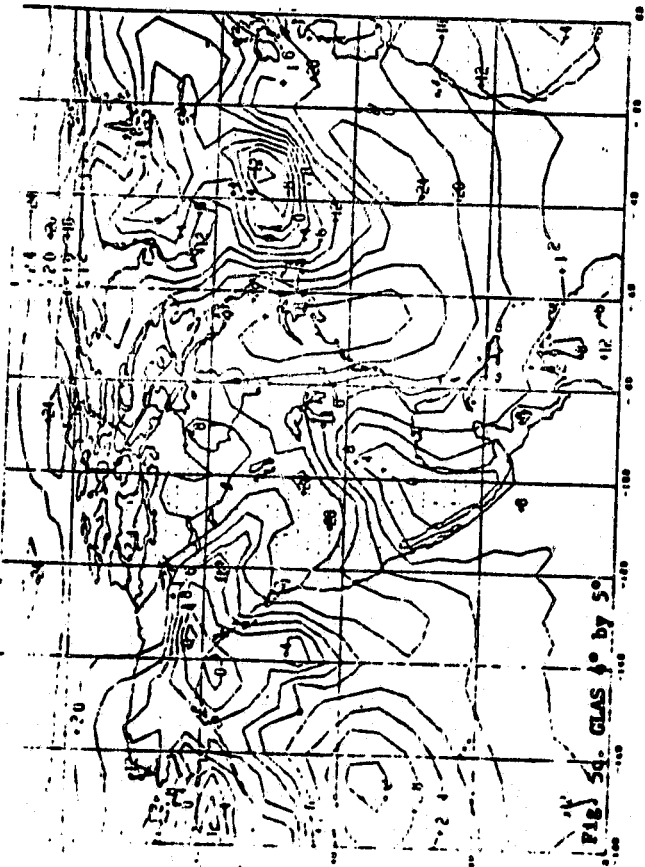
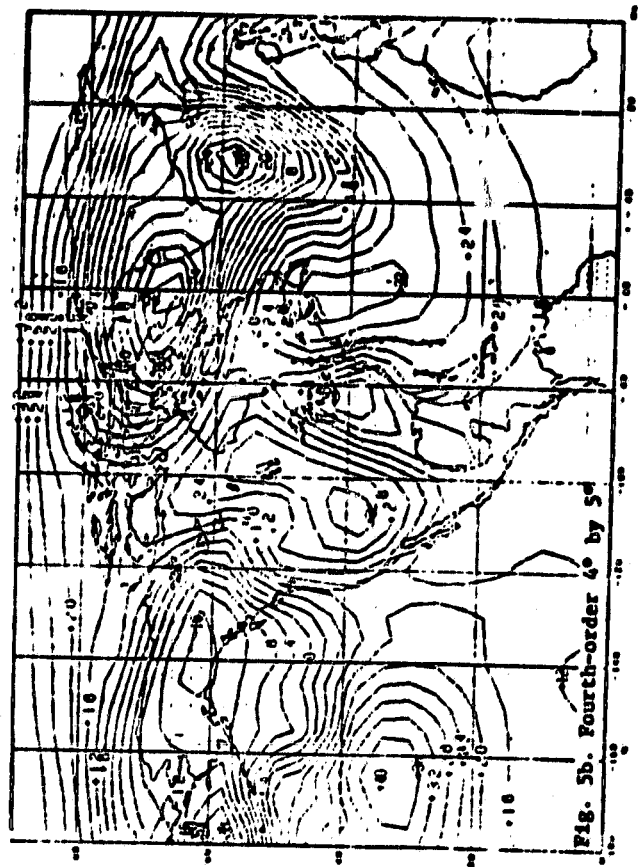
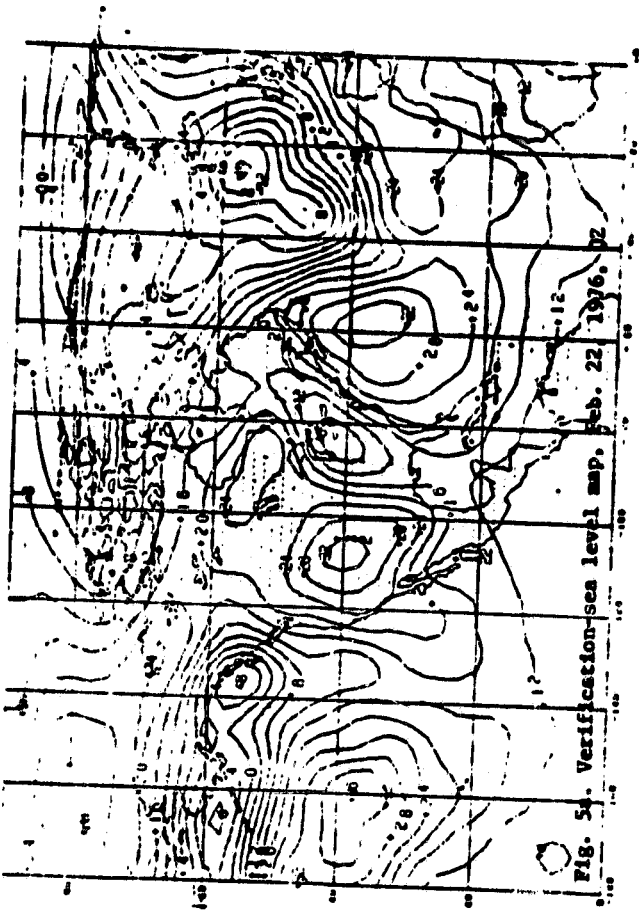
We think that the use of nonconservative schemes with high order periodic filtering is also justified in climate simulations as long as the rates of energy or enstrophy loss due to filtering remain much smaller than the observed rates of generation and dissipation.

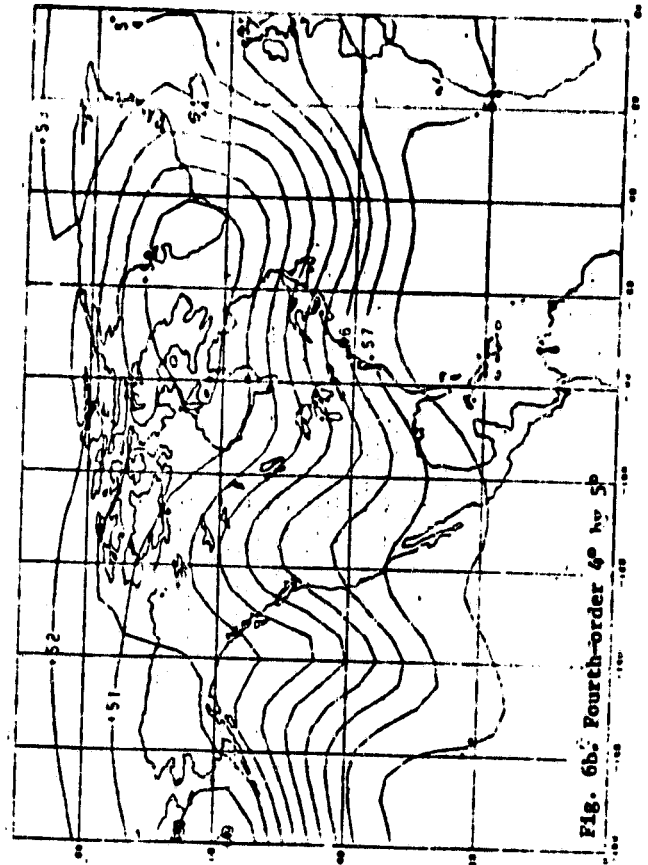
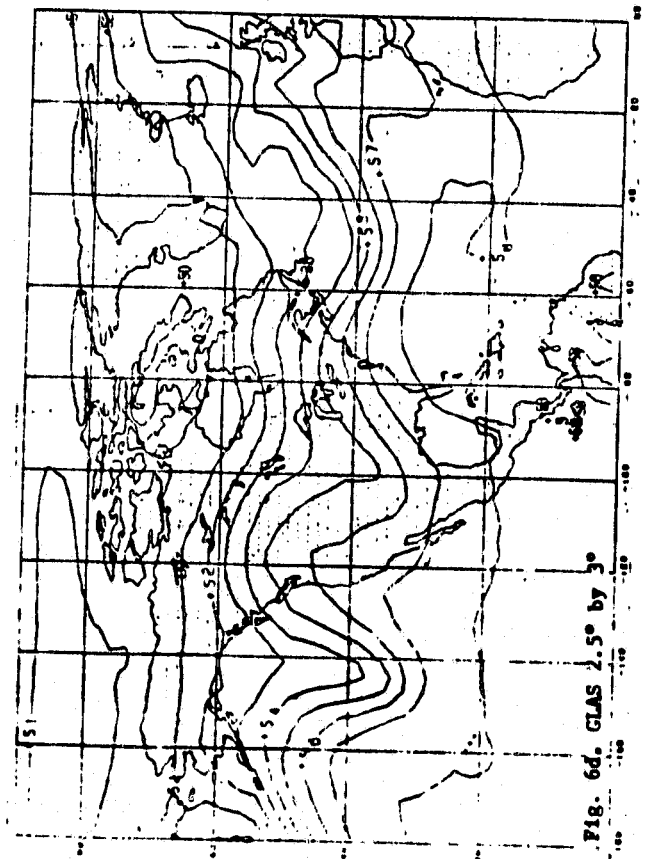
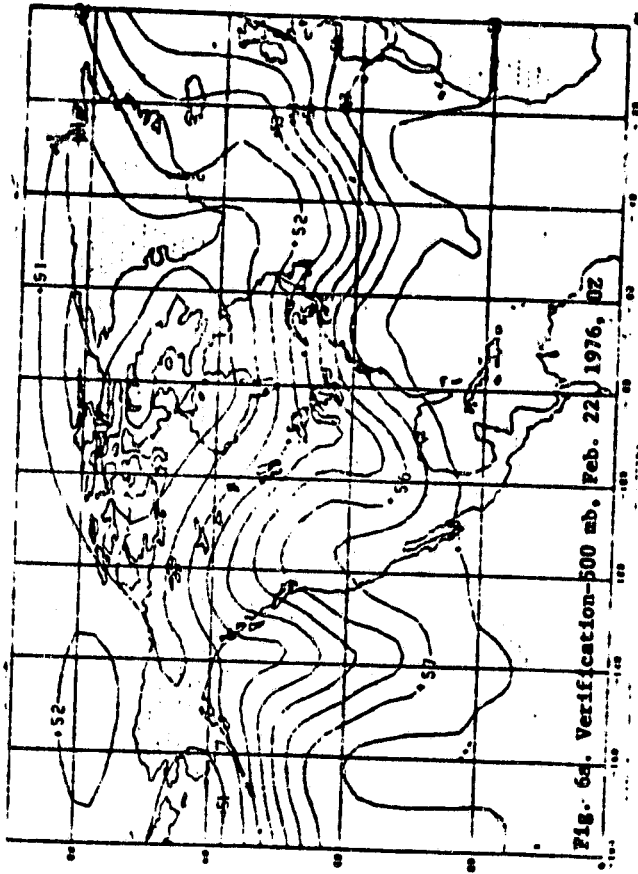
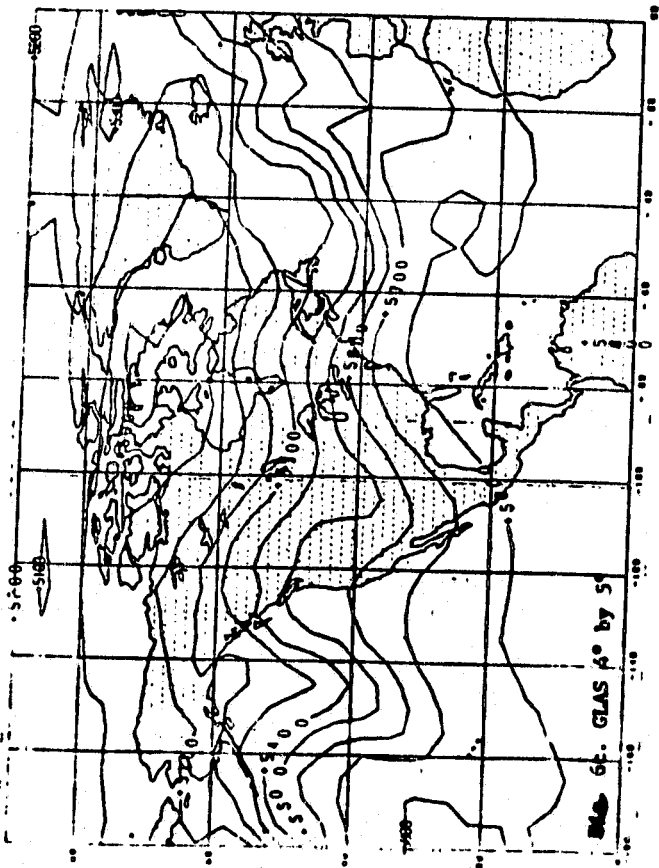
Acknowledgements: Dr. A. Bayliss' contribution was crucial in the development of the model. The authors have benefitted from stimulating discussions with Profs. M. Cane, E. Isaacson and D. Randall. The authors are grateful to Dr. M. Halem for his patience, encouragement and many useful suggestions. Dr. N. Rushfield and Mr. W. Connelly were extremely helpful in the development of the program. Dr. M. Wu provided us with her accurate radiation routine.

Fig. 5 Sea level pressure maps corresponding to 3-day forecasts with February 19, 1976, OZ initial conditions.

Fig. 6 Same as Fig. 5 but 500 mb geopotential heighte.

Fig. 7 Same as Fig. 5 but for February 1, 1976, OZ initial conditions.





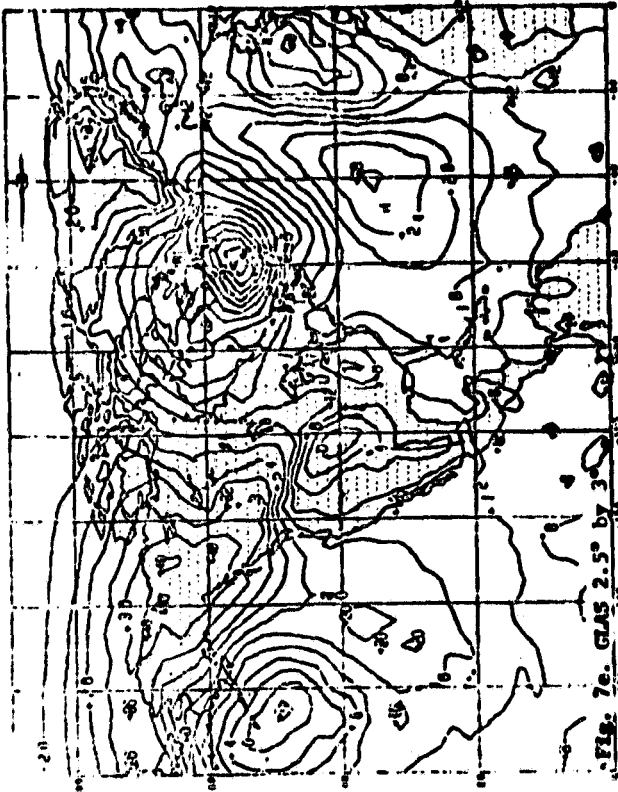


Fig. 7e. GLAS 2.5° by 3°

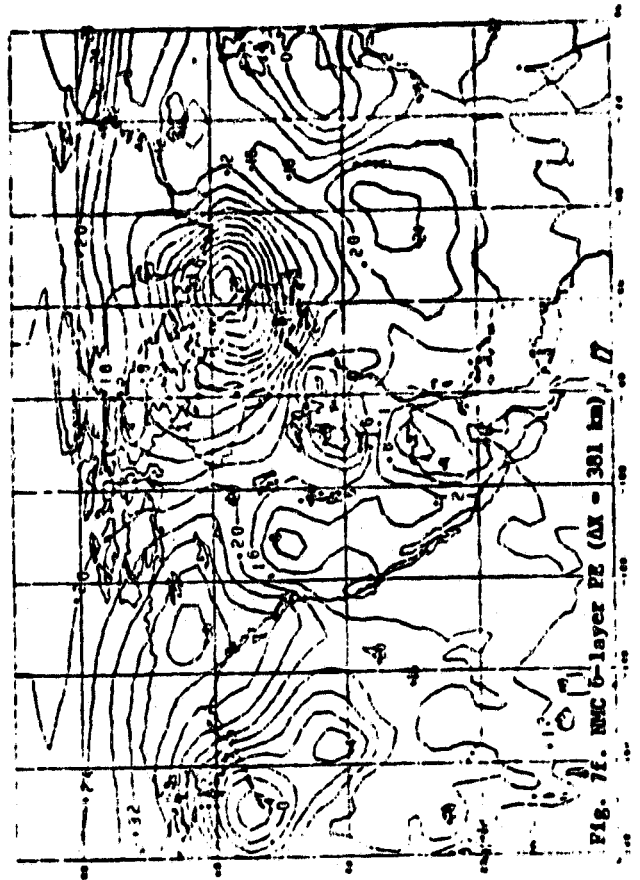


Fig. 7f. NMC 6-layer PE ($\Delta X = 361$ km)

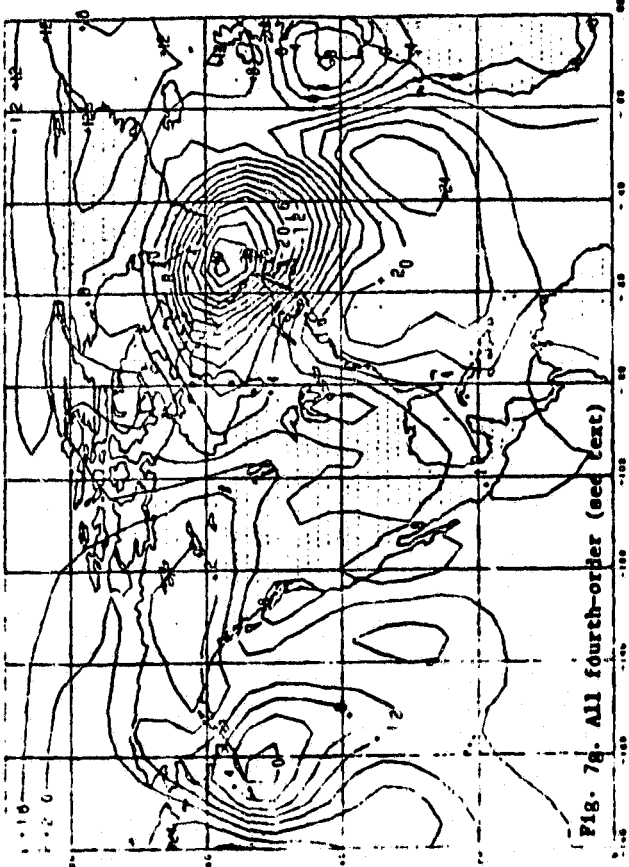


Fig. 7g. All fourth-order (see text)

REFERENCES

- Arakawa, A., 1966: Computational design for long term numerical integration of the equation of fluid motion: Two dimensional incompressible fluid. Part I. J. Compu. Phys., 1, 119-143.
- Arakawa, A., 1972: Design of the UCLA GCM, Report No. 7, Dept. of Meteorology, UCLA.
- Arakawa, A. and Y. Mintz, 1974: The UCLA atmospheric GCM. Workshop Notes, Dept. of Meteorology, UCLA.
- Atlas, R., M. Halem and M. Ghil, 1979: Subjective evaluation of the combined influence of satellite temperature sounding data and increased model resolution on numerical weather forecasting. To be presented at the 4th Conference on Numerical Weather Prediction, Silver Spring, MD, October.
- Campana, K., 1978: Real data experiments with a fourth-order version of the operational seven-layer model. NMC Office Note 188.
- Campana, K., 1979: Higher order finite difference experiments with a semi-implicit model at NMC. J. Atmos. Sci., 107, 363-376.

- Charney, J. G., 1971: Geostrophic turbulence. J. Atmos. Sci., 28, 1087-1095.
- Francis, P. E., 1975: The use of a multi-point filter as a dissipative mechanism in a numerical model of the general circulation of the atmosphere. Q. J. Roy. Met. Soc., 101, 567-582.
- Grammelvedt, A., 1969: A survey of finite difference schemes for the primitive equations for a barotropic fluid. Mon. Wea. Rev., 97, 384-404.
- Kalnay-Rivas, E., 1976: Numerical experiments with fourth-order conservative finite differences. Annalen der Meteorologie, 11. Simulation of large scale atmospheric processes (extended abstracts). Addendum.
- Kalnay-Rivas, E., A. Bayliss and J. Storch, 1977: The 4th order GISS model of the global atmosphere. Cont. to Atmos. Phys., 50, 306-311.
- Kraiss, H.-O. and J. Olinger, 1972: Comparison of accurate methods for the integration of hyperbolic equations. Tellus, 24, 199-215.
- Kraiss, H.-O. and J. Olinger, 1973: Methods for the approximate solution of time dependent problems. GARP Pub. Series No. 10, WMO.
- Leith, C. E., 1972: Internal turbulence and vertical momentum transfer. Parameterization of Sub-grid Processes, GARP Pub. Series No. 8, 40-51.
- Lilly, D. K., 1965: On the computational stability of numerical solutions of time-dependent nonlinear geophysical fluid dynamic problems. Mon. Wea. Rev., 93, 11-26.
- Merilees, P. E., 1975: The effect of grid resolution on the instability of a simple baroclinic flow. Mon. Wea. Rev., 103, 101-104.
- Orszag, S. A. and M. Israeli, 1974: Numerical simulation of viscous incompressible flow. Ann. Rev. of Fluid Mech., 6, 281-318.
- Phillips, N. A., 1959: An example of nonlinear computational instability. The Atmosphere and Sea in Motion, Rockefeller Inst. Press, New York, 501-504.
- Shapiro, R., 1970: Smoothing, filtering and boundary effects. Rev. Geophys. and Space Phys., 8, 359-387.
- Smagorinsky, J., 1963: General circulation experiments with the primitive equations. Mon. Wea. Rev., 91, 99-164.
- Williamson, D. L., 1978: The relative importance of resolution, accuracy and diffusion in short-range forecasts with the NCAR global circulation model. Mon. Wea. Rev., 106, 69-86.


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ISN 0121 YFDAY=FLCAT(ITAU-IDAY*124)/XINT
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ISN 0124 MATSUN=MATSUNX
ISN 0125 MATSNX=MLF(12)
ISN 0126 GO TO 11
ISN 0127 12 NTH=MOD(NSTEP-1,NSEQ)+1
ISN 0128 NEXT=MOD(NSTEP,NSEQ)+1
ISN 0129 MATSUN=MLF(NTH)
ISN 0130 MATSNX=MLF(NEXT)
ISN 0131 CONTINUE
ISN 0132 DT=DT*SAVE
ISN 0133 IF(NATSUN.EQ.0)DT=DT*52
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ISN 0135 $ SMT.P.Q.UT.VPOL.VTPOL.SMPOL.PPOL.OTPOL
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ISN 0137 IF(INDS.EQ.0)GO TO 14
ISN 0138 NSM=MOD(NSTEP-NSM1,(SMTH)
ISN 0139 IF(NATSUN.EQ.1)GO TO 14
ISN 0140 IF(NSM.GT.1)GO TO 14
ISN 0141 CALL SMSHAP(PT.OT.VTPOL.PPOL.PPOL.PPOL.PPOL)
ISN 0142 14 CONTINUE
ISN 0143 N*NS=1
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ISN 0145 NPC=1
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ISN 0148 IF(INSEQ.GT.1)GO TO 15
ISN 0149 MATSUN=MATSUNX
ISN 0150 MATSNX=MLF(12)
ISN 0151 GO TO 18
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ISN 0154 MATSUN=MLF(NTH)
ISN 0155 MATSNX=MLF(NEXT)
ISN 0156 CONTINUE
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ISN 0169 IF(EVENT(24)) CALL DAILY
ISN 0170 GOTO 40
ISN 0171 ***** END OF MAIN LOOP
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ISN 0173 88 CONTINUE
ISN 0174 WRITE(3,904) IDAY,TOFDAY
ISN 0175 STOP
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ISN 0183 * RUN START,F8.2,2X,*,ILL END,F6.2)
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ISN 0171	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0172	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0173	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0174	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0175	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0176	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0177	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0178	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0179	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0180	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0181	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0182	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0183	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0184	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0185	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0186	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0187	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0188	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0189	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0190	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0191	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0192	U	SFA	CE	R04	000E70						R04	NBR					

ISN	NAME TAG	TYPE	ADD.	1	2	3	NAME	TAG	TYPE	ADD.	1	2	3	NAME	TAG	TYPE	ADD.
ISN 0117	C	SFA	CE	R04	000000						R04	000388					
ISN 0118	P	SFA	CE	R04	002E10						R04	NBR					
ISN 0119	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0120	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0121	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0122	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0123	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0124	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0125	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0126	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0127	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0128	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0129	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0130	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0131	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0132	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0133	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0134	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0135	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0136	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0137	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0138	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0139	U	SFA	CE	R04	000E70						R04	NBR					
ISN 0140	U	SFA	CE	R04	000E70												

NAME	TYPE	REL.	ADDR.	ME	INFO
NTH	SF		000300	R04	
PSM	SFA		000300	R04	
STA	SFA		000300	R04	
CON1	SF		000300	R04	
CON2	SF		000300	R04	
CON3	SF		000300	R04	
DATE	SF		000300	R04	
DATG	SF		000300	R04	
PAST	SF		000300	R04	
INVS	SFA		000300	R04	
JAN1	SFA		000300	R04	
JAN2	SFA		000300	R04	
JAN3	SFA		000300	R04	
JAN4	SFA		000300	R04	
JAN5	SFA		000300	R04	
JAN6	SFA		000300	R04	
JAN7	SFA		000300	R04	
JAN8	SFA		000300	R04	
JAN9	SFA		000300	R04	
JAN10	SFA		000300	R04	
JAN11	SFA		000300	R04	
JAN12	SFA		000300	R04	
JAN13	SFA		000300	R04	
JAN14	SFA		000300	R04	
JAN15	SFA		000300	R04	
JAN16	SFA		000300	R04	
JAN17	SFA		000300	R04	
JAN18	SFA		000300	R04	
JAN19	SFA		000300	R04	
JAN20	SFA		000300	R04	
JAN21	SFA		000300	R04	
JAN22	SFA		000300	R04	
JAN23	SFA		000300	R04	
JAN24	SFA		000300	R04	
JAN25	SFA		000300	R04	
JAN26	SFA		000300	R04	
JAN27	SFA		000300	R04	
JAN28	SFA		000300	R04	
JAN29	SFA		000300	R04	
JAN30	SFA		000300	R04	
JAN31	SFA		000300	R04	
JAN32	SFA		000300	R04	
JAN33	SFA		000300	R04	
JAN34	SFA		000300	R04	
JAN35	SFA		000300	R04	
JAN36	SFA		000300	R04	
JAN37	SFA		000300	R04	
JAN38	SFA		000300	R04	
JAN39	SFA		000300	R04	
JAN40	SFA		000300	R04	
JAN41	SFA		000300	R04	
JAN42	SFA		000300	R04	
JAN43	SFA		000300	R04	
JAN44	SFA		000300	R04	
JAN45	SFA		000300	R04	
JAN46	SFA		000300	R04	
JAN47	SFA		000300	R04	
JAN48	SFA		000300	R04	
JAN49	SFA		000300	R04	
JAN50	SFA		000300	R04	
JAN51	SFA		000300	R04	
JAN52	SFA		000300	R04	
JAN53	SFA		000300	R04	
JAN54	SFA		000300	R04	
JAN55	SFA		000300	R04	
JAN56	SFA		000300	R04	
JAN57	SFA		000300	R04	
JAN58	SFA		000300	R04	
JAN59	SFA		000300	R04	
JAN60	SFA		000300	R04	
JAN61	SFA		000300	R04	
JAN62	SFA		000300	R04	
JAN63	SFA		000300	R04	
JAN64	SFA		000300	R04	
JAN65	SFA		000300	R04	
JAN66	SFA		000300	R04	
JAN67	SFA		000300	R04	
JAN68	SFA		000300	R04	
JAN69	SFA		000300	R04	
JAN70	SFA		000300	R04	
JAN71	SFA		000300	R04	
JAN72	SFA		000300	R04	
JAN73	SFA		000300	R04	
JAN74	SFA		000300	R04	
JAN75	SFA		000300	R04	
JAN76	SFA		000300	R04	
JAN77	SFA		000300	R04	
JAN78	SFA		000300	R04	
JAN79	SFA		000300	R04	
JAN80	SFA		000300	R04	
JAN81	SFA		000300	R04	
JAN82	SFA		000300	R04	
JAN83	SFA		000300	R04	
JAN84	SFA		000300	R04	
JAN85	SFA		000300	R04	
JAN86	SFA		000300	R04	
JAN87	SFA		000300	R04	
JAN88	SFA		000300	R04	
JAN89	SFA		000300	R04	
JAN90	SFA		000300	R04	
JAN91	SFA		000300	R04	
JAN92	SFA		000300	R04	
JAN93	SFA		000300	R04	
JAN94	SFA		000300	R04	
JAN95	SFA		000300	R04	
JAN96	SFA		000300	R04	
JAN97	SFA		000300	R04	
JAN98	SFA		000300	R04	
JAN99	SFA		000300	R04	
JAN100	SFA		000300	R04	

***** COMMON INFORMATION *****

NAME OF COMMON BLOCK	SIZE OF BLOCK	COMMON	005624	HEXADECIMAL	BYTES
CON1	R04				
CON2	R04				
CON3	R04				
CPATH	R04				
CP2	R04				
IMD2	R04				
JP2	R04				
INDEX	R04				
PK	R04				
PHISPL	R04				
VARIABLE OFFSETE					
UPOL	R04				
GPOL	R04				
SHPOL	R04				
SMPOL	R04				
TTPOL	R04				
NAME OF COMMON BLOCK	SIZE OF BLOCK	COMMON	00D146	HEXADECIMAL	BYTES
CON1	R04				
CON2	R04				
CON3	R04				
CPATH	R04				
CP2	R04				
IMD2	R04				
JP2	R04				
INDEX	R04				
PK	R04				
PHISPL	R04				
VARIABLE OFFSETE					
UPOL	R04				
GPOL	R04				
SHPOL	R04				
SMPOL	R04				
TTPOL	R04				
NAME OF COMMON BLOCK	SIZE OF BLOCK	COMMON	003E70	HEXADECIMAL	BYTES
CON1	R04				
CON2	R04				
CON3	R04				
CPATH	R04				
CP2	R04				
IMD2	R04				
JP2	R04				
INDEX	R04				
PK	R04				
PHISPL	R04				
VARIABLE OFFSETE					
UPOL	R04				
GPOL	R04				
SHPOL	R04				
SMPOL	R04				
TTPOL	R04				

VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 JMIN I% N.R. JMAX I% N.R. JSUM I% V.R. SMTM R% N.R.

NAME OF COMMON BLOCK # *N3|,E OF BLOCKCOMMON 300C38 HEXADECIMAL BYTES

VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 NSEQ I% 000005 MLF I% 000094 MATSNX I% 00002C NSM1 I% 000030

LABEL ADDR LABEL ADDR LABEL ADDR LABEL ADDR
 \$ 000298 10 0002F4 20 000314 30 000354
 40 0003FE 50 000430 60 00049A 70 0005C4

OPTIONS IN EFFECT NAME= MAIN,OPT=02,L,INECNT=55,SIZE=1000K,
 OPTIONS IN EFFECT SOURCE,EBCDIC,NOLIST,NODECK,LOAD,MAP,NODEIT, ID,NOXREF
 STATISTICS SOURCE STATEMENTS = 71 PROGRAM SIZE= 1660
 STATISTICS NO DIAGNOSTICS GENERATED
 ***** END OF COMPILATION *****
 LEVEL 19.6-APR 71 OS/360 FORTRAN M AT GLSS 868K BYTES OF CORE NOT USED
 DATE 12/12/79-0753.04

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COMPILER OPTIONS - NAME= MAIN,OPT=02,L,INECNT=55,SIZE=1000K,
SUBROUTINE INPUT
LOGICAL FLAGS
COMMON/FLD/FLAGS(5)
COMMON/FOURTH/CON1,CON2,CON3,CON4,CON5,CON6,CON7,CON8,ALPHA,BETA,IMD2P1,
$ ADLDP,JMS2,JMS1,CPTH,PSKAPA,R,RAD,CP,INC,NSTART,CPD2,
$ IND2,RAD1,JEND(2),JPI1,JPI2,MATSUN,JNE XI(2),MODPK(2,2),
$ JPOL(2,2),JPMOD(2,2),SINLON(72),COSLOPK(2,2),
$ PSIGN(4),POLES(148),W(72,9,5),PK(72,9,3),DIFF(9),
$ CONVP(9),SDPOL(9,2),PHISPL(2),SUM(2,5)
EQUIVALENCE (OPOL(1,1),POLES(1,1)),(OTPOL(1,1),POLES(73)),
$ (UPOL(1,1),POLES(1,1)),(VPOL(1,1),POLES(10)),(TPOL(1,1),
$ POLES(19)),(SHPOL(1,1),POLES(28)),(UTPOL(1,1),POLES(73)),
$ (VTPOL(1,1),POLES(62)),(WTPOL(1,1),POLES(91)),(SWTPOL(1,1),
$ POLES(100)),(PPOL(1,1),POLES(149)),(PPDL(1,1),POLES(147))
DIMENS ION PPOL(2),PTPOL(2),UPOL(36,1),VPOL(36,1),TPOL(36,1),
$ SHPOL(36,1),OPOL(9,4,2),OT(72,10,4,1)
DIMENS ION PTPOL(2),UTPOL(36,1),VPOL(36,1),TPPOL(36,1),
$ SHTPOL(36,1),OTPOL(9,4,2),OT(72,10,4,1)
INTEGER*2 ALBEDO
REAL*4 KAPA,LAY
REAL*8 SINLON,COSLON,SUM
REAL*8 ALONG,TWOPI
COMMON/INSMH/F5D6,F17,F98
COMMON/WORK2/PU(2880,46)
COMMON/SMITH/OSM(72,9,4,5),PSM(72,5),OSMPOL(9,4,2),PSMPOL(2)
COMMON/SDOT48/SDOT(72,46,8),OMEGA(72,46,5),MMXSTS,RADT(72,46,1,0)
+ ,COMG
LOGICAL COMG
INTEGER*2 RADT,OMEGA
INTEGER*2 ISDOT(72,46,8)
EQUIVALENCE (SDOT(1,1,5),ISDOT(1,1,1))
COMMON /CNTRL/
$ JSP,JNP,IM,NLAY,PTOP,ISTART,JSPFI,JNPMI,FIM,NLAYM,NLAYPI,
$ JI,JK,KM,TAUT,IKOT,MROT,JTEST,ITEST,
$ NR,JAYS(12),INCS(11),JSB,JNB,OLAT,DLON,
$ OT,TAU,ITAU,XINT,IDAY,JDAY,TDODAY,JDAT,IMONTH(2),JYEAR,NSTEP,
$ NCV,CLC,NCOMP3,NHOGAN,TAUP,TAUI,TAUE,TAUO,DTMULT,
$ PI,GRAY,KGAS,KAPA,PSL,ED,F,PU,NFL,PSF,DRCH,RSDIST,SIND,COSO,
$ RHM,CDX,DUMYTC(18),ALTER,DUMYVAL(99)
$ XLABEL(20),SIG(20),SIGE(21),DLCG(19),
$ JIPS(11),JMS(11),JUS(11),JMU(11),KSES(11),KNBR(11),
$ (LAT(46),DXL(46),DXP(46),DRUT(46),LXP(46),LXP(46),F(46),SINL(46),
$ COSL(46),DUMY(72),PHIS(2880,46),ALBEDO(72,46)
DIMENS ION U(72,40,1),V(72,40,1),T(72,40,1),SH(72,40,1),
$ (2880,1),TS(2880,1),SWS(2880,1),GT(2880,1),GM(2880,1),C(330)
DIMENS ION UT(72,40,1),VT(72,40,1),TT(72,40,1),SHT(72,40,1),
$ PT(2880,1)
COMMON/S/JMIN,JMAX,JSUM,SMTH(37,17)
COMMON/MLF/NSQ,MLF(10),MATSX,NSMI,NCM1
EQUIVALENCE
$ (Q1(1,1),U(1,1,1)),PHIS(73,1)),(TS(1,1),PHIS(72,1)),
$ (T1(1,1),PHIS(79,1)),(SMS(1,1),PHIS(144,1)),
$ (SM1(1,1),PHIS(233,1)),(QI(1,1,1),PHIS(216,1)),
$ (G1(1,1),PU(72,1)),(V(1,1,1),PU(793,1)),
$ (GM(1,1),PU(144,1)),(T(1,1,1),PULS(1,1)),
$ (PT(1,1),PU(216,1)),(SHT(1,1),PU(233,1)),(C(1,1),JSP)
DIMENS ION CI(300)
DIMENS ION INCP(46),ICU(46)
REAL*8 RECORD,ANDENO,DPI
DIMENS ION RECORD(10)
  
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84 M03320
 84 M03330
 84 M03340
 84 M03350
 84 M03360
 84 M03370
 84 M03380
 84 M03390
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 84 M03880
 84 M03890
 84 M03900
 84 M03910
 84 M03920
 84 M03930
 84 M03940
 84 M03950
 84 M03960


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ISN 0102      WRITE (.6,901) IVER
ISN 0103      READ (.5,902) XLABEL
ISN 0104      WRITE (.3,903) XLABEL
C****        COPY INPUTZ NAMELIST CNTC CORE TAPE AND TITLE PAGE
ISN 0105      READ (.5,904) RECORD
ISN 0106      WRITE (.11,904) RECORD
ISN 0107      IF (.RECORD(1).NE.ANDEND) GO TO 20
ISN 0108      REWIND 11
ISN 0109      READ (11,INPUTZ)
ISN 0110      IF (.START.EQ.7) WRITE (.3,938) ISTART
ISN 0111      IF (.START.EQ.7.AND.PAVZE.NE.0.) WRITE (.15,538) ISTART
ISN 0112      READ IN VARIABLE ALBEC
ISN 0113      READ (.40) (:,ALBEC(I,J),J=1,JNPI,I=1,IM)
C****        READ IN SURFACE GEOPOTENTIAL
ISN 0114      READ (.17) (:,PHIS(I,J),J=1,JNPI,I=1,IM)
ISN 0115      READ (.46) (:,PHIS(I,J),J=1,JNPI,I=1,IM)
ISN 0116      WRITE (.6,1900) (J,PHIS(I,J),I=1,IM),J=1,JNPI
ISN 0117      WRITE (.6,1941) (J,ALBEC(I,J),I=1,IM),J=1,JNPI
ISN 0118      FORMAT (1X,46: J=,13, PHIS(I,J)=,9(1X,8E13.6/),1)
ISN 0119      FORMAT (1X,46: J=,13, ALBEC(I,J)=,2(1X,3E13/),1)
ISN 0120      MMXTS = TAUT + XINT + .01
ISN 0121      IF (PAZE.EQ.1) PAUSE , MOUNT TAPES
ISN 0122      CALL SSM5CH (5,MSSWS)
ISN 0123      IF (KSSW5.EQ.1) ISTART=8
ISN 0124      KTR=8
ISN 0125      C**** * ON INITIAL START ENDFILE MODEL OUTPUT TAPE
ISN 0126      C
ISN 0127      C
ISN 0128      C
ISN 0129      C
ISN 0130      IF (.START.EQ.5.AND.ISTART.LE.7) ENDFILE PTR
ISN 0131      GO TO (30,120,120,40,110,110,40,50,90), ISTART
ISN 0132      STOP
ISN 0133      C**** SET DEPENDENT QUANTITIES
ISN 0134      30 DLONG=2.*DPI/IM
ISN 0135      40 DLAT=DPI/JM
ISN 0136      FM=IM
ISN 0137      KM=4*NLAY+5
ISN 0138      LMA=1*N*4
ISN 0139      JSPPI=JSP+1
ISN 0140      JNPMI=JNP-1
ISN 0141      JNPP1 = JNP + 1
ISN 0142      JNPP1D2 = JNP + 1 / 2
ISN 0143      JSB=1
ISN 0144      JSBPI = JSB + 1
ISN 0145      JNB=JM
ISN 0146      JNBMI=JNB-1
ISN 0147      NLAYMI=NLAY+1
ISN 0148      NLAYMI=NLAY+1
ISN 0149      C**** IF (DSIG(I,1).NE.0.) GO TO 60
ISN 0150      50 DSIG(I,1)=.0./NLAY
ISN 0151      60 STGE(I)=0./NLAY
ISN 0152      70 STGE(I,1)=STGE(I)+DSIG(I,1)
ISN 0153      80 DSIG(I,1)=STGE(I)+DSIG(I,1)
ISN 0154      90 DSIG(I,1)=STGE(I)+DSIG(I,1)
ISN 0155      IF (.START.EQ.4) GO TO 120
ISN 0156      IF (.START.EQ.7) GO TO 110
ISN 0157      MACHINE CHECK RESTART, ISTART = E OR 9
ISN 0158      WRITE (.15,905)
ISN 0159      READ (.15,907) TAUP
ISN 0160      TAUP=TAUP-AMOD(TAUP,TAUT)
ISN 0161      WRITE (.3,908) TAUP
ISN 0162      GO TO 120
ISN 0163      C**** INITIAL CONDITIONS FROM UNIT 12. ISTART= 5, 6, OR 7
ISN 0164      110 KTR=12
ISN 0165      TAUT=TAU
ISN 0166      IF (TAUP.LI.0.) TAUP=TAU
ISN 0167      TAPNUM=ISK
ISN 0168      IF (PAZE.EQ.0.) GO TO 120
ISN 0169      WRITE (.15,909)
ISN 0170      READ (.15,902) TAPNUM
ISN 0171      WRITE (.13,916) TAPNUM
ISN 0172      READ TAPE ON UNIT KTR
ISN 0173      C
ISN 0174      C
ISN 0175      C
ISN 0176      C
ISN 0177      C
ISN 0178      C
ISN 0179      C
ISN 0180      120 READ(KTR,ERR=800,END=130)TAUX,C1,(:,PHIS(I,J),J=1,JNPI,I=1,IM),
ISN 0181      (:,ALBEC(I,J),J=1,JNPI,I=1,IM),(:,PHIS(I,J),J=1,JNPI,I=1,IM),
ISN 0182      (:,ALBEC(I,J),J=1,JNPI,I=1,IM)
ISN 0183      CALL (DPI,ALBEC(I,J),J=1,JNPI,I=1,IM)
ISN 0184      IF (.START.LE.4) GO TO 120
ISN 0185      IF (TAUX.LE.TAUP+.01) GO TO 120
ISN 0186      CONTINUE
ISN 0187      WRITE (.6,912) TAUP
ISN 0188      C
ISN 0189      C

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BAND 4850
BAND 4860
BAND 4870
BAND 4880
BAND 4890
BAND 4900
BAND 4910
BAND 4920
BAND 4930
BAND 4940
BAND 4950
BAND 4960
BAND 4970
BAND 4980
BAND 4990
BAND 5000
BAND 5010
BAND 5020
BAND 5030
BAND 5040
BAND 5050
BAND 5060
BAND 5070
BAND 5080
BAND 5090
BAND 5100
BAND 5110
BAND 5120
BAND 5130
BAND 5140
BAND 5150
BAND 5160
BAND 5170
BAND 5180
BAND 5190
BAND 5200
BAND 5210
BAND 5220
BAND 5230
BAND 5240
BAND 5250
BAND 5260
BAND 5270
BAND 5280
BAND 5290
BAND 5300
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BAND 5330
BAND 5340
BAND 5350
BAND 5360
BAND 5370
BAND 5380
BAND 5390
BAND 5400
BAND 5410
BAND 5420
BAND 5430
BAND 5440
BAND 5450
BAND 5460
BAND 5470
BAND 5480
BAND 5490
BAND 5500
BAND 5510
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BAND 5570
BAND 5580
BAND 5590
BAND 5600
BAND 5610
BAND 5620
BAND 5630
BAND 5640
BAND 5650
BAND 5660
BAND 5670
BAND 5680
BAND 5690
BAND 5700
BAND 5710
BAND 5720

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15N 0190
15N 0191
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15N 0224
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15N 0228
15N 0229
15N 0230
15N 0231
15N 0232
15N 0233
15N 0234
15N 0235
15N 0236

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CALL EXIT
C 130 IF(I,START,GE,SIG)GTO 830
BACKSPACE KTR
TAULTAUX
140 TAULTAUX
DO 145 I=1,JM
BACKSPACE KTR
145 BACKSPACE KTR
150 IROT=IROT-1
IF(I,START,EQ,DR,ISTART,EG,7) GC TO 160
C**** COPY C ARRAY FROM CI ARRAY
DO 150 I=1,300
C(I)=C(CI)
150 IF(I,START,EG,2,DR,ISTART,EG,5) GC TO 160
REWIND 11
READ (11,INPUTZ)
IF(I,TR,EG,8) C(16)=C(115)
IF(I,TR,EG,8) ALABEL(20)=C(1220)
IF(I,TR,EG,12) ALABEL(20)=TAPRUM
C**** CALCULATE DISTANCE PROJECTION ARRAYS
150 FJ=90-S*(JSP/JNP)
DO 200 J=1,JNP
LAT(J)=DLAT*(J-FJE0)
DO 210 I=1,JNP
SIN(I,J)=SIN(LAT(I,J))
COS(I,J)=COS(LAT(I,J))
DO 220 J=1,JM
DXP(J)=RAD*DLON*COSL(J)
DO 230 J=2,JM
DYU(J)=RAD * DLAT
DYV(J)=RAD*(LAT(J)-LAT(J-1))
CONTINUE
230 JMPID2 = (JM + 1)/2
JNPI = JNP + 1
DO 240 J = 2,JMPID2
DXP(J) = 12 * DXP(J) * DYU(J)
DXV(JNPI-J) = DXVP(J)
DXV(1) = 0.
DXVP(JNPI) = 3.
WRITE(3,2200)(DXVP(J),J=1,JNP)
FORMAT(1X,*,DXVP=/(1X,SEI6.8))
2200 RETURN
C * * * AUG 23 1977 ADDITIONS HERE TILL RETURN
ENTRY INPUT*

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84M05730
84M05740
84M05750
84M05760
84M05770
84M05780
84M05790
84M05800
84M05810
84M05820
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84M05840
84M05850
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84M05870
84M05880
84M05890
84M05900
84M05910
84M05920
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84M05950
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84M05970
84M05980
84M05990
84M06000
84M06010
84M06020
84M06030
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84M06070
84M06080
84M06090
84M06100
84M06110
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84M06140
84M06150
84M06160
84M06170
84M06180
84M06190
84M06200
84M06210
84M06220
84M06230
84M06240
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84M06340
84M06350
84M06360
84M06370
84M06380
84M06390
84M06400
84M06410
84M06420
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84M06480
84M06490
84M06500
84M06510
84M06520
84M06530
84M06540
84M06550
84M06560
84M06570
84M06580
84M06590
84M06600

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15N 0275
15N 0276

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INC = 1
BETA = 1. - 2. * ALPHA
FSD6 = 15. / 6.
F17 = 17. / 18.
F98 = .325 / 981.
R = RGAS
RADIM = 3 * RAD * IM
CON1 = 4. * COTAN(S*DLAT) / RADIM
CON2 = -COTAN(DLAT) / RADIM
CON3 = 1. / (RADIM * DLAT)
INC(1,1) = 2
JPOL(1,1) = 3
JPOL(1,2) = JM - 1
JPOL(2,2) = JM - 1
JMOD(1,1) = 2
JMOD(2,1) = 3
JMOD(1,2) = MOD(JM-1,S) + 1
JMOD(2,2) = MOD(JM-2,S) + 1
MODPK(1,1) = 2
MODPK(2,1) = 3
MODPK(1,2) = MOD(JM-1,3) + 1
MODPK(2,2) = MOD(JM-2,3) + 1
JNP=JMT
JEND(1) = 1
JEND(2) = JNP
JNEXT(1) = 2
JNEXT(2) = JM
ADLDP = 12. * RAD * DLAT * DLON
IMD2=IM/2
IMD2PI = IMD2 + 1
JMS1 = JM - 1
JMS2 = JM - 2
JMS1-6 = 2 + JSMT
JMIN = JM + JSMT
JMAX = JMAX + JMIN
TROJN = 2. * JM
JLTM = JMIN - 1

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* * *

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ISN 0278
ISN 0279
ISN 0280
ISN 0281
ISN 0282
ISN 0283
ISN 0284
ISN 0285
ISN 0286
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ISN 0296
ISN 0297
ISN 0298
ISN 0299
ISN 0300
ISN 0301
ISN 0302
ISN 0303
ISN 0304
ISN 0305
ISN 0306
ISN 0307

DO 301 J = 1, JLM
JDH = J + 1
IF (J.GE.JMIN) JDM = JMAX + 1 - JMIN
IF (J.LE.JMAX) JDM = JMAX - 1
IF (J.GT.JMIN) JDM = JMAX + 1 - JMIN
DO 310 I = 1, IMD2
SMTH(I, J) = 1.
IF (I.GE.IMAX) AND (I.LE.(IMAX3) SMTH(I, J) = 1 - FLOAT(I - IMAX3) / 6.
IF (I.GT.IMAX3) SMTH(I, J) = 0.
CONTINUE
CONTINUE
PIKAPA = KAPA + 1
DO 320 L = 1, NLAY
DIFF(L) = PIKAPA * DSIG(L)
CONTINUE
PSKAPA = EXPBYK(1000.)
THBAR = 280. / PSKAPA
CP = RGAS / KAPA
CPD2 = 5 * CP
CPH = CP * THBAR
WRITE(6, 9778) PSKAPA, THBAR, RGAS, KAPA, CP, CPTH
PHISPL(1) = PHIS(1, 1)
PHISPL(2) = PHIS(1, 2)
IMD2 = IM / 2
TWOPI = 2 * PI
TWOPI = 6.283185307175587
DO 1000 I = 1, IMD2
ALONG = FLOAT(I - 1) * TWOPI / FLOAT(IM)
C * * * ALONG = FLOAT(I - 1) * TWOPI / FLCAT(IM) - PI
C * * *
SINLON(1) = -DSIN(ALONG)
COSLON(1) = -DCOS(ALONG)
SINLON(IMD2+1) = -SINLON(1)
COSLON(IMD2+1) = -COSLON(1)
INDEX(1) = I + IMD2
INDEX(IMD2+1) = I
CONTINUE
OHM2 = 2 * TWOPI / SDAY
DO 1010 J = 1, JM
FL(J) = OHM2 * SINL(J)
C * * *
FLJNP(J) = OHM2
PSIGN(1) = -1.
PSIGN(2) = -1.
PSIGN(3) = 1.
PSIGN(4) = 1.
C * * * INITIAL VELOCITIES GIVEN ON A STAGGERED GRID.
C * * * INTERPOLATION TO UNSTAGGERED GRID IS USED.
C * * * MAKE USE OF EQUIVALENCING AND OVERLAPPING THE U ARRAY.
C * * *
DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
IF (L.START.EQ.7) GO TO 1017
Q1(L, L, NM2, JEND(M)) = QT(L, L, NM2, JEND(M))
Q1(L, L, N, JEND(M)) = QT(L, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
PL(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q1(I, L, NP2, J) = QT(I, L, NP2, J)
IF (L.START.EQ.7) GO TO 1026
Q1(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J + 1)
+ QT(I, L, N, J - 1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q1(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
PT(I, J) = PT(I, J)
CONTINUE
C * * * SCALE UT
C * * *

ISN 0308
ISN 0309
ISN 0310
ISN 0311
ISN 0312
ISN 0313
ISN 0314
ISN 0315
ISN 0316
ISN 0317
ISN 0318
ISN 0319
ISN 0320
ISN 0321
ISN 0322
ISN 0323

DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
IF (L.START.EQ.7) GO TO 1017
Q1(L, L, NM2, JEND(M)) = QT(L, L, NM2, JEND(M))
Q1(L, L, N, JEND(M)) = QT(L, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
PL(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q1(I, L, NP2, J) = QT(I, L, NP2, J)
IF (L.START.EQ.7) GO TO 1026
Q1(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J + 1)
+ QT(I, L, N, J - 1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q1(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
PT(I, J) = PT(I, J)
CONTINUE
C * * * SCALE UT
C * * *

ISN 0324
ISN 0325
ISN 0326
ISN 0327
ISN 0328
ISN 0329
ISN 0330
ISN 0331
ISN 0332
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ISN 0339
ISN 0340
ISN 0341
ISN 0342
ISN 0343
ISN 0344
ISN 0345
ISN 0347

DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
IF (L.START.EQ.7) GO TO 1017
Q1(L, L, NM2, JEND(M)) = QT(L, L, NM2, JEND(M))
Q1(L, L, N, JEND(M)) = QT(L, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
PL(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q1(I, L, NP2, J) = QT(I, L, NP2, J)
IF (L.START.EQ.7) GO TO 1026
Q1(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J + 1)
+ QT(I, L, N, J - 1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q1(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
PT(I, J) = PT(I, J)
CONTINUE
C * * * SCALE UT
C * * *

ISN 0348
ISN 0349
ISN 0350
ISN 0351
ISN 0352
ISN 0353
ISN 0354

DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
IF (L.START.EQ.7) GO TO 1017
Q1(L, L, NM2, JEND(M)) = QT(L, L, NM2, JEND(M))
Q1(L, L, N, JEND(M)) = QT(L, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
PL(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q1(I, L, NP2, J) = QT(I, L, NP2, J)
IF (L.START.EQ.7) GO TO 1026
Q1(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J + 1)
+ QT(I, L, N, J - 1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q1(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
PT(I, J) = PT(I, J)
CONTINUE
C * * * SCALE UT
C * * *

BAND 6610
BAND 6620
BAND 6630
BAND 6640
BAND 6650
BAND 6660
BAND 6670
BAND 6680
BAND 6690
BAND 6700
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BAND 6970
BAND 6980
BAND 6990
BAND 7000
BAND 7010
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BAND 7480

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ISN 0420

ISN 0421

ISN 0422

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DO 1020 N = 1.4  
DO 1020 L = 1.4  
DO 1020 I = 1.4  
OT(I,L,N,J) = DXP(J) * PI(I,J) * Q(I,L,N,J)  
CONTINUE  
C  
C * * * COMPUTATION OF STEREOGRAPHIC POLAR VELOCITIES: UPOL(L,M).  
C * * * VPOL(L,M), WHERE M=1 CORRESPONDS TO SOUTH POLE  
C * * *  
DO 1040 M = 1.2  
COEF1 = (-1) ** M  
DO 1040 L = 1.4  
IF(I,START,NE,?) GO TO 1052  
S1 = 0.  
S2 = 0.  
DO 1050 I = 1.4  
S1 = S1 - U(I,L,JNEXT(M)) * SIN(LN(I)) - COEF1 * V(I,L,JNEXT(M))  
S2 = S2 + COEF1 * U(I,L,JNEXT(M)) * COS(LN(I)) + COEF1 * V(I,L,JNEXT(M)) * SIN(LN(I))  
VPOL(L,M) = S1 / FLOAT(I,M)  
GO TO 1053  
1052 UPOL(L,M) = -U(I,L,JEND(M)) * SIN(LN(I)) - COEF1 * V(I,L,JEND(M)) * COS(LN(I))  
VPOL(L,M) = -U(I,L,JEND(M)) * COS(LN(I)) + COEF1 * V(I,L,JEND(M)) * SIN(LN(I))  
CONTINUE  
1053 TPOL(L,M) = T(I,L,JEND(M))  
SPOL(L,M) = SH(I,L,JEND(M))  
PPOL(L,M) = PI(I,L,JEND(M))  
WRITE(6,198)M,L,PPOL(M),(OPOL(L,N),M),N(1.4)  
1988 FORMAT(198)M,L,PPOL(M),(OPOL(L,N),M),N(1.4)  
TPOL(M),PPOL(M)  
DO 1045 N = 1.4  
COEF1 = 0.  
DO 1040 M = 1.4  
COEF1 = COEF1 + OPOL(L,N,M) * TPOL(L,M)  
C * * * WITH THE OT, I, L, JEND(M) COMPUTATION BELOW AGREES  
C * * * IF(I,START,NE,?) GO TO 1054  
U(I,L,JEND(M)) = -UPOL(L,M) * SIN(LN(I)) + COEF1 * VPOL(L,M) * COS(LN(I))  
V(I,L,JEND(M)) = -COEF1 * UPOL(L,M) * COS(LN(I)) - VPOL(L,M) * SIN(LN(I))  
CONTINUE  
1054 OT(I,L,N,JEND(M)) = Q(I,L,N,JEND(M)) * PPOL(M)  
CONTINUE  
1040 OT(I,L,N,JEND(M)) = Q(I,L,N,JEND(M)) * PPOL(M)  
CONTINUE  
800 WRITE(7,911)  
C*** ERROR ENCOUNTERED READING TAUC RECORD ON UNIT KTR  
CALL EXIT  
810 WRITE(7,916) J  
C*** ERROR ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR  
CALL EXIT  
C*** END-OF-FILE ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR  
820 WRITE(7,917) J  
C*** LATER PICKUP TAPE NEEDED  
C*** CALL EXIT  
830 WRITE(7,913) TAUP  
STOP  
901 FORMAT(740X,'GISS N LAYER WEATHER MODEL//100X *VERSION* (4//)  
902 FORMAT(20A4)  
903 FORMAT(25X,20A4)  
904 FORMAT(15A8)  
905 FORMAT(15A8)  
906 FORMAT(15X,10A8)  
907 FORMAT(426.2 AS 04026*20*)  
908 FORMAT(78.2)  
909 FORMAT(78.2)  
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1000 FORMAT(78.2)
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IMD2 144 000050 RADIN R44 000054 JEND 144 000058
 JPD2 144 N4R 000069 MATSUN R44 000069 JPI 144 000069
 JPOL 144 000084 JPMOD 144 000094 JNEXT 144 000094
 INDEX 144 000224 PSIGN R44 000404 SINLN R44 000404
 PK 144 N4R 000404 DIFR R44 000404 POLES R44 000414
 PHISPL R44 000574 SUM R44 000574 CONVPL R44 000574
 SDPOL R44 N4R 000574

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 OPOL 000414
 TPOL 00045C
 TTPOL 00057C
 SHTPCL 0005A0
 SHPOL 0005B0
 SHTPCL 0005A0
 SHPOL 0005B0
 SHTPCL 0005A0

NAME OF COMMON BLOCK * 1156, ZE OF BLOCK COMMON 00060C HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 F5D6 R44 000000 F17 R44 000004 F98 R44 000008

NAME OF COMMON BLOCK * 0085, ZE OF BLOCK COMMON 001600 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 PU R44 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 OT 000120
 GW 001680
 TT 0017A0

NAME OF COMMON BLOCK * 0010, ZE OF BLOCK COMMON 0001A8 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 QSH R44 000000 PSN R44 000004 GMPOL R44 000008

NAME OF COMMON BLOCK * 8007, ZE OF BLOCK COMMON 0385A8 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 SDOT R44 000000 OMEGA I42 000004 MNMXTS I44 0286E0
 CONG L44 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 ISDOT 00CF00

NAME OF COMMON BLOCK * 0018, ZE OF BLOCK COMMON 063E30 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JSD R44 000000 JAP R44 000004 JSPR1 I44 000018
 PTD R44 000010 ISTART I44 000014 NLA YPI I44 000024
 FJM R44 000020 NLAYMI I44 000024 ITESI I44 000048
 MROT I44 000040 JTEST I44 000048 JJB I44 000080
 JAVE I44 000080 DLON R44 000088 JDI I44 00009C
 DLAT R44 000094 XINT R44 0000C8 JDAY I44 N4R 0000C0
 TOFLAU R44 N4R 0000C8 JDATE I44 N4R 0000D0
 NSTED R44 N4R 0000D0 N4R 0000D0 JYEAR I44 N4R 0000D4
 INSTEP R44 N4R 0000D0 N4R 0000D0 N4R 0000D4
 DTMUT R44 000108 N4R 0000D0 N4R 0000D4
 KAPA R44 000118 PSL R44 00012C N4R 000124
 NPL9 R44 000128 N4R 00012C N4R 000124
 SIND R44 N4R 000128 N4R 00012C N4R 000124
 DUMNYC R44 N4R 000128 N4R 00012C N4R 000124
 SALS R44 N4R 000128 N4R 00012C N4R 000124
 JIBS R44 N4R 000128 N4R 00012C N4R 000124
 KSB5 R44 N4R 000128 N4R 00012C N4R 000124
 DYE R44 000128 N4R 00012C N4R 000124
 PHIS R44 000050 ALBEDO I42 062350

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 Q 000E70
 SHS 002300
 C 000000

NAME OF COMMON BLOCK * 0010, ZE OF BLOCK COMMON 0005E0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JMIN I44 000000 JMAX I44 000004 JSUN I44 000008
 JSMIN I44 000000 JSMAX I44 000004 JSUN I44 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 Q 000E70
 SHS 002300
 C 000000

NAME OF COMMON BLOCK * 0010, ZE OF BLOCK COMMON 0005E0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JMIN I44 000000 JMAX I44 000004 JSUN I44 000008
 JSMIN I44 000000 JSMAX I44 000004 JSUN I44 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 Q 000E70
 SHS 002300
 C 000000

NAME OF COMMON BLOCK * 0010, ZE OF BLOCK COMMON 0005E0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JMIN I44 000000 JMAX I44 000004 JSUN I44 000008
 JSMIN I44 000000 JSMAX I44 000004 JSUN I44 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 Q 000E70
 SHS 002300
 C 000000

NAME OF COMMON BLOCK * *NSI* OF BLOCKCOMMON 000038 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
NSI L#4 000000 NSI L#4 000004 MATSNX L#4 000018
NSI L#4 00003A NSI L#4 00003A NSI L#4 00003C

NAME OF COMMON BLOCK * *NSM* OF BLOCKCOMMON 000039 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
NSM L#4 000000 NSM L#4 000004 FAST L#4 000018
NSM L#4 00003A NSM L#4 00003A NSM L#4 00003C

NAME OF COMMON BLOCK * *DS_* OF BLOCKCOMMON 000004 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
DS_* L#4 000000 DS_* L#4 000004 DS_* L#4 000018
DS_* L#4 00003A DS_* L#4 00003A DS_* L#4 00003C

NAME OF COMMON BLOCK * *IT* OF BLOCKCOMMON 000008 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
IT L#4 000000 IT L#4 000008 IT L#4 000018
IT L#4 00003A IT L#4 00003A IT L#4 00003C

NAME OF COMMON BLOCK * *S6_* OF BLOCKCOMMON 000008 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
S6_* L#4 000000 S6_* L#4 000008 S6_* L#4 000018
S6_* L#4 00003A S6_* L#4 00003A S6_* L#4 00003C

NAME OF COMMON BLOCK * *T* OF BLOCKCOMMON 000008 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
T L#4 000000 T L#4 000008 T L#4 000018
T L#4 00003A T L#4 00003A T L#4 00003C

LABEL ADDR
20 001430
50 001900
60 001950
100 001A88
130 001D8C
160 001E6C
200 001F5E
300 0023DE
301 0023FE
1017 00276A
1025 002968
1052 002C98
1040 002F30
1830 003050

LABEL ADDR
30 001B40
70 001956
80 001ACC
120 001B54
140 001DAC
190 001EAC
230 001F56
320 002440
1018 002776
1030 0029F6
1053 002CE2
500 002FED

LABEL ADDR
50 0019C8
90 001A4A
121 001D62
150 001E24
210 001F5E
310 0023D2
311 002644
1010 002644
1026 002960
1050 002CCE
1054 002F16
820 003028 NR

OPTIONS IN EFFECT NAME= MAIN,OPT=02,ALINECNT=55,SIZE=1000K.
OPTIONS IN EFFECT SOURCE,EBCDIC,NOLIST,NODECK,LOAD,MAP,NOEDIT,NOXREF
STATISTICS SOURCE STATEMENTS = 421 PROGRAM SIZE= 12544
STATISTICS NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****
LEVEL 19.6-APR 71

05/360 FORTRAN M AT GISS

COMPILER OPTIONS - NAME= MAIN,OPT=02,ALINECNT=55,SIZE=1000K.
SUBROUTINE TWRITE SOURCE,EBCDIC,NOLIST,NODECK,LOAD,MAP,NOEDIT,NOXREF
C**** FINE NINE LAYER GLOBAL CDE BLOCK FOR VM MODEL, FEBRUARY 1976
ISN 0002 INTEGER*2 ALBEDO
ISN 0003 REAL*8 KAPA,LAT
ISN 0004 COMMON /CNTRL/
ISN 0005 * JSP,JNP,M,NLAY,PTOP,I,START,J,SCPI,JNPM1,FIM,NLAYM,NLAY*J,
* JI,JM,KM,TAUT,IRGT,WRGT,JTEST,ITEST,
* NR,JAYS(12),INCS(11),JSB,JNB,CLAT,DOLON,
* DT,TAU,ITAU,XINT,TDAY,JDAY,TOFGAY,JDATÉ,JMONTH(2),JYEAR,NSTEP,
* NCYCLE,NCMPS,KAPA,PSI,TAUT,TAUTALD,DTMULT,
* PI,GRAY,RCAS,KAPA,PSI,ED,FMU,NFLB,PSF,PRCH,RSDIST,SIND,COSD,
* RHMAX,CDX,DUMMYC(18),I,ALTER,DUMMYA(99),
* XLABE(20),SIG(20),DSIG(20),SICE(21),DSIGO(19),
* JIPSI(1),JMP(11),JUS(11),JMS(11),KNS(11),KNS(11),
* LATT(4),DX(46),DXP(46),DY(46),DYP(46),LALBEDO(12,46),
* COSL(46),DUMMY(72),PHIS(2880,46),ALBEDO(12,46),
* DIMENSION U(72,40,11),V(72,40,11),T(72,40,11),SM(72,40,11),
* PL(2880,11),TS(2880,11),S(2880,11),G(12880,11),G(2880,11),C(1300),
* COMMON /WORK2/PU(2880,46)
C + COME
ISN 0008 COMMON/SDOT48/SDOT(72,46,8),GMEGA(72,46,8),MNMTS,RADT(72,46,10)
ISN 0009 LOGICAL COME
C INTEGER*2 RADT,OMEGA
ISN 0010 INTEGER*2 ISDOT(72,46,8)
ISN 0011 C EQUIVALENCE (SDOT(1,1,5),SDOT(1,1,1))
C DIMENSION SDFIL(72,12),RADFIL(72,16),OMF(1,72,9)
ISN 0012 C EQUIVALENCE (SDFIL(72,12),RADFIL(72,16),OMF(1,72,9))
ISN 0013 C EQUIVALENCE
ISN 0014 C (U(1,1),PHIS(73,1)),(TS(5,1),PHIS(73,1)),
(V(1,1),PHIS(73,1)),(S(5,1),PHIS(73,1)),
(LATT(4),DX(46),DXP(46),DY(46),DYP(46)),
(COSL(46),DUMMY(72)),(G(1,1),G(2,1)),(C(1),JSP),
(M,NLAY,PTOP,I,START,J,SCPI,JNPM1,FIM,NLAYM,NLAY*J),
COMMON /WORK1/TAUT,IRGT,WRGT,
EQUIVALENCE(C(100),MACHIN)
ISN 0015 MACHIN = 100
ISN 0016 IMPL = 14 + 1
ISN 0017 IM2 = 2 * IM
ISN 0018 IM3 = 3 * IM
ISN 0019 IM3PI = IM3 + 1
ISN 0020 IM3PI = IM3 + 1
ISN 0021 IM3PI = IM3 + 1
ISN 0022 IM3PI = IM3 + 1
ISN 0023 IM3PI = IM3 + 1
ISN 0024 IM3PI = IM3 + 1
ISN 0025 TAUR = INT(TAU + .5)
ISN 0026 BACKSPACE 8
C

70 WRITE (8,END=80,ERR=90)TAUR,C(1,1),J,J=1,JNP,I=1,IM,
((TS(I,J),J=1,JNP),=(I,IN),(S(I,J),J=1,JNP),=(I,IN),
((G(I,J),J=1,JNP),=(I,IN),((C(I,J),J=1,JNP),=(I,IN),
ISN 0027

772K BYTES OF CORE NOT USED
DATE 12/12/79-0783.20

ISN 0035
ISN 0036
ISN 0037
ISN 0038
ISN 0040
ISN 0042
ISN 0043

ISN 0044
ISN 0045
ISN 0046
ISN 0047
ISN 0048
ISN 0049

ISN 0050
ISN 0051
ISN 0052
ISN 0053
ISN 0054
ISN 0055
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ISN 0112

ISN 0114
ISN 0115
ISN 0116
ISN 0117
ISN 0118
ISN 0119

JMOD = MOD(J - 1.5) + 1
JS1MOD = MOD(JS1 - 1.5) + 1
JS2MOD = MOD(JS2 - 1.5) + 1
IF (JP2 .GT. JMI)
GO TO 25
GO TO 16
JP2 = 2
JP2MOD = 2

C * * * SAVE QTPOL IN QSMPOL
C
DO 15 M = 1.2
PSMPOL(M) = PTPOL(M)
DO 15 N = 1.4
DC 15 L = 1. NLAY
QSMPOL(L,N,M) = QTPOL(L,N,M) CONTINUE

C * * * SAVE OT(JP2) IN QSM(JP2MOD)
C
DO 22 I = INC.IM,INC
PSM(I,JP2MOD) = PT(I,JP2)
CONTINUE

DO 20 N = 1.4
DO 20 L = 1. NLAY
DO 20 I = INC.IM,INC
OSM(I,L,N,JP2MOD) = OT(I,L,N,JP2)
CONTINUE
GO TO 25
JP2 = JP2 + 1
JP2MOD = JP2MOD + 1

25 CALL COMPI(U,V,T,SM,P,Q,UPOL,VPCL,TPOL,SPOL,PPOL,OPOL,UT,VT,TT,
S,SHT,PT,OT,UTPOL,VTPOL,TPOL,SPOL,PPOL,OPOL,UT,VT,TT,
S, SHT,PT,OT,UTPOL,VTPOL,TPOL,SPOL,PPOL,OPOL,UT,VT,TT,
IF (JANEI.AND.JONE.JNP160 TO 210
M=1
IF (JAEQ,INP)=2
DO 204 L=2,NLAY
LM=1
IF (SHTPOL(L,M),GE.0,IGO TO 204
SHTPOL(L,M)=SHTPOL(L,M)+SHTPOL(L,M)*DSIG(LM)/DSIG(L)
SHTPOL(L,M)=0
CONTINUE
204 GO TO 212
IF (SHTPOL(NLAY,M),L.T.0.) SHTPOL(NLAY,M)=0.
210 DO 202 L=2,NLAY
DO 202 I=1,IM
LM=1
IF (SHT(I,LM),GE.0,IGO TO 202
IF (SHT(I,LM),L.T.0.) SHT(I,LM)=SHT(I,LM)+SHT(I,LM)*DSIG(LM)/DSIG(L)
SHT(I,LM)=0
CONTINUE
202 CONTINUE
IF (SHT(I,NLAY),L.T.0.) SHT(I,NLAY)=0.
203 CONTINUE
212 CONTINUE

DO 27 I=1,INC.IM,INC
IF (PST,LE,1100.) GO TO 27
WRITE(3,3021)PT(I,J),S(I,J)
WRITE(3,3021)UT(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)VT(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)TP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)SP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)PP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)OP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)UT(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)VT(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)TP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)SP(I,INC.L,JP1),L=1,NLAY)
WRITE(3,3021)PP(I,INC.L,JP1),L=1,NLAY)
STOP

27 IF (J .LT. 3)
IF (J .EQ. 3)
CONTINUE
GO TO 20
GO TO 70

C * * * UNSCALE
C
DO 60 N = 1.4
DO 60 L = 1,NLAY
DO 60 I = INC.IM,INC
OT(I,L,N,JS2) = OT(I,L,N,JS2) / (DAXP(JS2) * PT(I,JS2))
CONTINUE

60 CALL AVRX(UT,VT,TT,SHT,PT,GT,JS2)
CONTINUE
C * * * NO TIME SMOOTHING FOR MATSUKI STEP
C

84M09990
84M10000
84M10010
84M10020
84M10030
84M10040
84M10050
84M10060
84M10070
84M10080
84M10090
84M10100
84M10110
84M10120
84M10130
84M10140
84M10150
84M10160
84M10170
84M10180
84M10190
84M10200
84M10210
84M10220
84M10230
84M10240
84M10250
84M10260
84M10270
84M10280
84M10290
84M10300
84M10310
84M10320
84M10330
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84M10360
84M10370
84M10380
84M10390
84M10400
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84M10690
84M10700
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84M10770
84M10780
84M10790
84M10800
84M10810
84M10820
84M10830
84M10840
84M10850
84M10860

```

C *** IF(LEAP.AND.NEXT STEP IS LEAP)
C *** IF(MATSUN+MATSX.EQ.0)GO TC 45
C *** IF(NEXT CALL IS TO MATS PRED.)
C *** IF((NPC.EQ.1).AND.(MATSX.EQ.1))GO TO 58
C *** DO 28 I = INC.IM,INC
C *** P(I.JS2) = PSM:I.JS2MOD)
C *** CONTINUE
C *** DO 30 N = 1,4
C *** DO 30 L = 1,NLAY
C *** DO 30 I = INC.IM,INC
C *** Q(I.L.N.JS2) = QSM:I.L.N.JS2MOD)
C *** CONTINUE
C *** DO 61 I = INC.IM,INC
C *** P(I.JS2) = PT(I.JS2)
C *** GO TC 67
C *** CONTINUE
C *** DO 62 N = 1,4
C *** DO 62 L = 1,NLAY
C *** DO 62 I = INC.IM,INC
C *** Q(I.L.N.JS2) = QT(I.L.N.JS2) * (DXYP(JS2) * PT(I.JS2))
C *** CONTINUE
C *** GO TC 64
C *** DO 45 DO 48 I = INC.IM,INC
C *** P(I.JS2) = BETA * P(I.JS2) + ALPHA * (PSM(I.JS2MOD) + PT(I.JS2))
C *** CONTINUE
C *** DO 48 DO 50 N = 1,4
C *** DO 50 L = 1,NLAY
C *** DO 50 I = INC.IM,INC
C *** Q(I.L.N.JS2) = BETA * DXYP(JS2) * P(I.JS2) * Q(I.L.N.JS2)
C *** + ALPHA*(QSM(I.L.N.JS2MOD)+QT(I.L.N.JS2)*DXYP(JS2)*PT(I.JS2))
C *** CONTINUE
C *** DO 63 CONTINUE TERM CORRECTION DUE TO LEAPFROG TIME SCHEME
C *** DO 63 * * * SOURCE
C *** IF(MOD(NSTEP-NCM1,NCOMP3).NE.0)GC TO 67
C *** DO 57 N = 1,4
C *** DO 57 L = 1,NLAY
C *** DO 57 I = INC.IM,INC
C *** Q(I.L.N.JS2) = Q(I.L.N.JS2) - QT(I.L.N.JS2) *
C *** IDXP(JS2)*PT(I.JS2)
C *** CONTINUE
C *** DO 57 IF(MOD(NSTEP-NCM1,NCOMP3).NE.0)GC TO 67
C *** IF(MATSUNO.PRED. STEP)
C *** IF((MATSUNEQ.1).AND.(NPC.EQ.0))GO TO 67
C *** CALL COMP3:UT.VT.T.SHT.PT.UTFOL.VT.POL.UTPOL.SHTPOL.PT.POL.
C *** $ OTPOL.JS2)
C *** IF(MATSUNO.OR NOT COMP3 STEP)
C *** IF((MATSX.NE.0).OR.(MOD(NSTEP-NCM1,NCOMP3).NE.0))GO TO 67
C *** DO 65 N = 1,4
C *** DO 65 L = 1,NLAY
C *** DO 65 I = INC.IM,INC
C *** Q(I.L.N.JS2) = Q(I.L.N.JS2) + DXYP(JS2) * PT(I.JS2) *
C *** $ CONTINUE
C *** CONTINUE
C *** IF (J .LT. JM)
C *** JS2 = JS2 + 1
C *** JS2MOD = MOD(JS2-1,5) + 1
C *** GO TO 29
C *** IF (JS2 .LE. JM)
C *** * * * POLES
C *** DO 70 M = 1
C *** IF (J .EQ. JM)
C *** COEF1 = (-1.) ** M
C *** IF(LEAP.AND.NEXT STEP IS LEAP)
C *** IF(MATSUN+MATSX.EQ.0)GO TO 85
C *** IF(MATS.PRED.OR(MATS CORR. AND NEXT STEP IS LEAP))
C *** IF((MATSUNEQ.1).AND.(NPC.MATSX.EQ.0))GC TO 78
C *** PPOL(M) = PTPOL(M)
C *** DO 75 N = 1,4
C *** DO 75 L = 1,NLAY
C *** OPOL(L,N,M) = OTPOL(L,N,M)
C *** CONTINUE
C *** GO TC 54
C *** DO 78 PPOL(M) = PSMPOL(M)
C *** DO 80 N = 1,4
C *** DO 80 L = 1,NLAY
C *** OPOL(L,N,M) = OSMPOL(L,N,M)
C *** CONTINUE
C *** GO TC 54

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ISN 0120
ISN 0122
ISN 0124
ISN 0125
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ISN 0131
ISN 0132
ISN 0134
ISN 0135
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ISN 0197
ISN 0199

84M 0870
84M 0880
84M 0890
84M 0900
84M 0910
84M 0920
84M 0930
84M 0940
84M 0950
84M 0960
84M 0970
84M 0980
84M 0990
84M 1000
84M 1010
84M 1020
84M 1030
84M 1040
84M 1050
84M 1060
84M 1070
84M 1080
84M 1090
84M 1100
84M 1110
84M 1120
84M 1130
84M 1140
84M 1150
84M 1160
84M 1170
84M 1180
84M 1190
84M 1200
84M 1210
84M 1220
84M 1230
84M 1240
84M 1250
84M 1260
84M 1270
84M 1280
84M 1290
84M 1300
84M 1310
84M 1320
84M 1330
84M 1340
84M 1350
84M 1360
84M 1370
84M 1380
84M 1390
84M 1400
84M 1410
84M 1420
84M 1430
84M 1440
84M 1450
84M 1460
84M 1470
84M 1480
84M 1490
84M 1500
84M 1510
84M 1520
84M 1530
84M 1540
84M 1550
84M 1560
84M 1570
84M 1580
84M 1590
84M 1600
84M 1610
84M 1620
84M 1630
84M 1640
84M 1650
84M 1660
84M 1670
84M 1680
84M 1690
84M 1700
84M 1710
84M 1720
84M 1730
84M 1740

```

ISN 0200      PPOL(M) = BETA * PPOL(M) + ALPHA * (PSNPOL(M) + PTPOL(M))
ISN 0201      DO 87 N = 1,4
ISN 0202      DD 87 L = 1,NLAY
ISN 0203      OPOL(L,N,M) = BETAPPOL(M) * OPOL(L,N,M) + AL * PA
ISN 0204      $ * (OSMPOL(L,N,M) + QTPOL(L,N,M))
C * * * SOURCE TERM CORRECTION DUE TO LEAPFROG TIME SCHEME
C * * *
C *** 12/12/77
ISN 0205      93 CONTINUE
ISN 0206      IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 94
ISN 0207      DD 92 N = 1,4
ISN 0208      DD 92 L = 1,NLAY
ISN 0209      OPOL(L,N,M) = OPOL(L,N,M) - QTPOL(L,N,M)
ISN 0210      QTPOL(L,N,M) = QTPOL(L,N,M)
ISN 0211      C * * * UNSCALE
C * * *
ISN 0212      94 CONTINUE
ISN 0213      DD 96 N = 1,4
ISN 0214      DD 96 L = 1,NLAY
ISN 0215      QPOL(L,N,M) = QPOL(L,N,M) / PTFOLL(M)
ISN 0216      IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 98
ISN 0217      IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 98
ISN 0218      IF (MATSUNO PRED, STEP)
ISN 0219      CALL COMP3(UT, VT, TT, SHT, PT, DT, UTPOL, VTPOL, TTPOL, SHTPOL, PTPOL)
ISN 0220      $ QTPOL(J, S2)
ISN 0221      IF (MATSUNO, NE, 0) GO TO 58
ISN 0222      DD 195 N = 1,4
ISN 0223      DD 195 L = 1,NLAY
ISN 0224      QPOL(L,N,M) = QPOL(L,N,M) + PTPOL(M) * QTPOL(L,N,M)
ISN 0225      QPOL(L,N,M) = QPOL(L,N,M) + QTPOL(L,N,M)
ISN 0226      C * * * COPY QTPOL INTO QT
ISN 0227      C * * *
C * * *
ISN 0228      98 DO 100 I = INC, IM, INC
ISN 0229      PT(I, JEND(M)) = PTPOL(M)
ISN 0230      100 CONTINUE
ISN 0231      DD 110 N = 3,4
ISN 0232      DD 110 L = 1,NLAY
ISN 0233      DD 110 I = INC, IM, INC
ISN 0234      QT(I, L, N, JEND(M)) = QTPOL(L, N, M)
ISN 0235      110 CONTINUE
ISN 0236      DD 120 L = 1,NLAY
ISN 0237      DD 120 I = INC, IM, INC
ISN 0238      UT(I, L, JEND(M)) = UTPOL(L, N, M) * SIN(LN(I) + COEF1 * VTPOL(L, N, M)) * COS(LM(I))
ISN 0239      VT(I, L, JEND(M)) = VTPOL(L, N, M) * COS(LN(I) + COEF1 * UTPOL(L, N, M)) * SIN(LM(I))
ISN 0240      120 JS2 = JS1
ISN 0241      JS2 = J
ISN 0242      10 FORMAT( '0 PRESSURE DIAGNOSTIC. (J,PT,PE=.2I4.2E15.5)
ISN 0243      301
ISN 0244      302 FORMAT(5X, 2E14.4)
ISN 0245      RETURN
ISN 0246      END
ISN 0247

```

NAME TAG TYPE ADD	COMP	SIZE OF PROGRAM	HEXADECIMAL BYTES
K SFA XR	104	000000	000100
L SFA XR	104	000100	000100
V SFA XR	104	000000	000000
ED	104	000000	000000
GM	104	000000	000000
JM	104	000000	000000
PK	104	000000	000000
SH SFA XR	104	000000	000000
VT SFA XR	104	000000	000000
DYP	104	000000	000000
INC SFA C	104	000000	000000
JP2 SFA C	104	000000	000000
JP2 SFA C	104	000000	000000
NPC	104	000000	000000
PST	104	000000	000000
SHT SFA XR	104	000000	000000
AVRX SF	104	000000	000000
CON2	104	000000	000000
COSD	104	000000	000000
DIFD	104	000000	000000
DYVD	104	000000	000000
INCS	104	000000	000000
JDAY	104	000000	000000
JMS1	104	000000	000000
JPS	104	000000	000000
KSB5	104	000000	000000

NAME	VAR.	NAME	TYPE	REL.	ADDR.	HE	NAME	TYPE	REL.	ADDR.	HE
NELW	C	R4	NSM	R4	000000	C	NSM	R4	000000	C	NSM
PHIS	CE	R4	OPOL	R4	000010	CE	OPOL	R4	000010	CE	OPOL
R37	C	R4	SIGE	R4	000030	C	SIGE	R4	000030	C	SIGE
SIND	C	R4	TAUT	R4	N.R.	C	TAUT	R4	N.R.	C	TAUT
TAUD	C	R4	TAUT	R4	N.R.	C	TAUT	R4	N.R.	C	TAUT
ALPHA	SFA	R4	XINT	R4	N.R.	SFA	XINT	R4	N.R.	SFA	XINT
UPOL	XR	R4	COMPO	R4	000000	XR	COMPO	R4	000000	XR	COMPO
COMP2	F	R4	DSIG	R4	000000	F	DSIG	R4	000000	F	DSIG
INDEX	XF	R4	ISDGT	R4	000000	XF	ISDGT	R4	000000	XF	ISDGT
JDATE	C	R4	JTEST	R4	N.R.	C	JTEST	R4	N.R.	C	JTEST
JSPR1	C	R4	OMEGA	R4	N.R.	C	OMEGA	R4	N.R.	C	OMEGA
NSTEP	A	R4	GPOL	R4	000000	A	GPOL	R4	000000	A	GPOL
PTPOL	XR	R4	SHPOL	R4	000000	XR	SHPOL	R4	000000	XR	SHPOL
SDPOL	XR	R4	UTPOL	R4	000000	XR	UTPOL	R4	000000	XR	UTPOL
TTPOL	XR	R4	ALBEDO	R4	N.R.	XR	ALBEDO	R4	N.R.	XR	ALBEDO
STMULT	XF	R4	DUNMYA	R4	N.R.	XF	DUNMYA	R4	N.R.	XF	DUNMYA
SIN	C	R4	DUNMYC	R4	N.R.	C	DUNMYC	R4	N.R.	C	DUNMYC
IBCON#	F	R4	ISTART	R4	N.R.	F	ISTART	R4	N.R.	F	ISTART
JMGNTH	XF	R4	JP2MOD	R4	000000	XF	JP2MOD	R4	000000	XF	JP2MOD
JS2MOD	XF	R4	MATSUM	R4	000000	XF	MATSUM	R4	000000	XF	MATSUM
NCOM#3	A	R4	NMOGAN	R4	N.R.	A	NMOGAN	R4	N.R.	A	NMOGAN
NLAYP1	C	R4	PHISPL	R4	000000	C	PHISPL	R4	000000	C	PHISPL
PSKAP2	C	R4	NSINLN	R4	000000	C	NSINLN	R4	000000	C	NSINLN
SHTPOL	SFA	R4	TOFDAY	R4	000000	SFA	TOFDAY	R4	000000	SFA	TOFDAY

***** COMMON INFORMATION *****

NAME OF COMMON BLOCK	* F:USIZE OF BLOCKCOMMON	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE
CON1	R4	R4	CON3	R4	000008	R4	CON3	R4	000008	R4	CON3	R4	000008
CON5	R4	R4	ALPHA	R4	000018	R4	ALPHA	R4	000018	R4	ALPHA	R4	000018
IMD2P1	R4	R4	JMS2	R4	N.R.	R4	JMS2	R4	N.R.	R4	JMS2	R4	N.R.
CP1	R4	R4	R	R4	N.R.	R4	R	R4	N.R.	R4	R	R4	N.R.
CP	R4	R4	NSTART	R4	000044	R4	NSTART	R4	000044	R4	NSTART	R4	000044
IMD2	R4	R4	JEND	R4	000054	R4	JEND	R4	000054	R4	JEND	R4	000054
JP2	R4	R4	JNEXT	R4	000058	R4	JNEXT	R4	000058	R4	JNEXT	R4	000058
JPOL	R4	R4	MATSUM	R4	000058	R4	MATSUM	R4	000058	R4	MATSUM	R4	000058
INDEX	R4	R4	SINLN	R4	000044	R4	SINLN	R4	000044	R4	SINLN	R4	000044
PK	R4	R4	POLES	R4	N.R.	R4	POLES	R4	N.R.	R4	POLES	R4	N.R.
PHISPL	R4	R4	CONVPL	R4	N.R.	R4	CONVPL	R4	N.R.	R4	CONVPL	R4	N.R.
SUM	R4	R4	SUM	R4	N.R.	R4	SUM	R4	N.R.	R4	SUM	R4	N.R.

***** COMMON BLOCK * =M..ZE OF BLOCKCOMMON *****

NAME OF COMMON BLOCK	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	
OSM	R4	R4	OSMPL	R4	000020	R4	OSMPL	R4	000020	R4	OSMPL	R4	000020

***** COMMON BLOCK * -MTR# OF BLOCKCOMMON *****

NAME OF COMMON BLOCK	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	
JSP	R4	R4	JSP1	R4	000004	R4	JSP1	R4	000004	R4	JSP1	R4	000004
PTOP	R4	R4	ISTART	R4	000004	R4	ISTART	R4	000004	R4	ISTART	R4	000004
FM	R4	R4	NLAYM1	R4	N.R.	R4	NLAYM1	R4	N.R.	R4	NLAYM1	R4	N.R.
FM	R4	R4	KH	R4	N.R.	R4	KH	R4	N.R.	R4	KH	R4	N.R.
JM	R4	R4	JTEST	R4	N.R.	R4	JTEST	R4	N.R.	R4	JTEST	R4	N.R.
JAYS	R4	R4	JNB	R4	N.R.	R4	JNB	R4	N.R.	R4	JNB	R4	N.R.
DLAT	R4	R4	DLON	R4	N.R.	R4	DLON	R4	N.R.	R4	DLON	R4	N.R.
DLAT	R4	R4	JDATE	R4	N.R.	R4	JDATE	R4	N.R.	R4	JDATE	R4	N.R.
ITAU	R4	R4	NYCLE	R4	N.R.	R4	NYCLE	R4	N.R.	R4	NYCLE	R4	N.R.
INSTP	R4	R4	PI	R4	N.R.	R4	PI	R4	N.R.	R4	PI	R4	N.R.
TAUP	R4	R4	PSL	R4	N.R.	R4	PSL	R4	N.R.	R4	PSL	R4	N.R.
OTMULT	R4	R4	PSF	R4	N.R.	R4	PSF	R4	N.R.	R4	PSF	R4	N.R.
KAPA	R4	R4	OSD	R4	N.R.	R4	OSD	R4	N.R.	R4	OSD	R4	N.R.
NEU	R4	R4	IALTER	R4	N.R.	R4	IALTER	R4	N.R.	R4	IALTER	R4	N.R.
SIND	R4	R4	DSIG	R4	000000	R4	DSIG	R4	000000	R4	DSIG	R4	000000
DUNMYC	R4	R4	JMPS	R4	N.R.	R4	JMPS	R4	N.R.	R4	JMPS	R4	N.R.
SIG	R4	R4	KNBS	R4	N.R.	R4	KNBS	R4	N.R.	R4	KNBS	R4	N.R.
JPS	R4	R4	DVU	R4	N.R.	R4	DVU	R4	N.R.	R4	DVU	R4	N.R.
KSBS	R4	R4	SINL	R4	N.R.	R4	SINL	R4	N.R.	R4	SINL	R4	N.R.
DXP	R4	R4	ALBEDO	R4	N.R.	R4	ALBEDO	R4	N.R.	R4	ALBEDO	R4	N.R.
PHIS	R4	R4	PHIS	R4	000050	R4	PHIS	R4	000050	R4	PHIS	R4	000050

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP

VARIABLE	OFFSETE	NAME	TYPE	REL.	ADDR.	HE	VARIABLE	OFFSETE	NAME	TYPE	REL.	ADDR.	HE
TS	001890	TS	C		000000		TS	001890	TS	C		000000	

***** COMMON BLOCK * #DS..ZE OF BLOCKCOMMON *****

NAME OF COMMON BLOCK	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	* NAME	* TYPE	* REL.	* ADDR.	* HE	
ISWTC	R4	R4	ISWTC	R4	000000	R4	ISWTC	R4	000000	R4	ISWTC	R4	000000

VARIABLE OFFSETE

VARIABLE	OFFSETE	NAME	TYPE	REL.	ADDR.	HE
TS	001890	TS	C		000000	

VARIABLE OFFSETE

VARIABLE	OFFSETE	NAME	TYPE	REL.	ADDR.	HE
TS	001890	TS	C		000000	

NAME OF COMMON BLOCK * SROT#9E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
SROT R#4 000003 OMEGA I#2 N#R. MNMXTS I#4 N#R. RADT I#2 N#R.
CORG L#4 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 00CF00 VARIABLE OFFSETE

NAME OF COMMON BLOCK * #56#7E OF BLOCKCOMMON 000006 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
ISMTH I#4 NPC I#4 000004

NAME OF COMMON BLOCK * #M3#1E OF BLOCKCOMMON 000038 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
NSE0 I#4 000034 MLP I#4 N#R. MATSNX I#4 00002C NSM1 I#4 N#R.

NAME OF COMMON BLOCK * #T#7E OF BLOCKCOMMON 0035C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
TOPOG R#4 N#R.

NAME OF COMMON BLOCK * #W6R#E OF BLOCKCOMMON 0101C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
FD R#4 N#R.

NAME OF COMMON BLOCK * #0#5#E OF BLOCKCOMMON 081400 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
PU R#4 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE VARIABLE OFFSETE

LABEL	ADDR	ADDR
15	000B1E	000B1E
16	000D4A	000E2A
17	001FCA	001F04
18	001F4B	001F50
19	0018FA	0019EE
20	001C12	001C6E
21	001F2E	001F72
22	002004	002142
23	002278	00219E
24	002482	002324
25	00278C	0024EC
26		0027CC

LABEL	ADDR	ADDR
27	000B66	000BDE
28	000E2A	000E84
29	001F04	001F5A
30	001F50	001F84
31	0019EE	001A3E
32	001C6E	001D82
33	001F72	002098
34	002142	0022A4
35	00219E	0023A2
36	002324	00254C
37	0024EC	0027E4
38	0027CC	

LABEL	ADDR	ADDR
39	000B66	000BDE
40	000E2A	000E84
41	001F04	001F5A
42	001F50	001F84
43	0019EE	001A3E
44	001C6E	001D82
45	001F72	002098
46	002142	0022A4
47	00219E	0023A2
48	002324	00254C
49	0024EC	0027E4
50	0027CC	

```

*OPTIONS IN EFFECT* NAME= MAIN,OPT=02,L INECLT=55,SIZE=100K.
*OPTIONS IN EFFECT* SOURCE=EBCOIC,NOLIST,NODECK,LCAD,MAP,NCECREDIT,LD,NOXREF
*STATISTICS* SOURCE STATEMENTS = 246 PROGRAM SIZE= 11120
*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILEATION *****
LEVEL 19.6-APR 71
  
```

```

COMPILER OPTIONS - NAME= MAIN,OPT=02,L INECLT=55,SIZE=100K.
SUBROUTINE CO.SPI:UAV,T,SH,P,C,UPCL,VPCL,JPCL,SMPOL,PPCL,QPOL,
$ UT,V7,TT,SH,PT,QT,UTPOL,VTPOL,VPOL,STPOL,STPOL,PTPOL,OTPOL,OTPOL,TJ
$ COMMON/FGURTH/CONI,CON2,CON3,CON4,CONST,PBAR,ALPHA,BETA,IMDZPI,
$ ADLDP,JMS2,JMS1,CPTH,PSKAPA,R,RAD,CP,INC,NSTART,CPO2,
$ IND2,RADIM,JEND(2),JPI,JP2,MATSUN,JNEXT(2),MODPK(2,2),
$ PSIGN(4),POLES(148),W(72,9,5),PK(72,9,3),DIFF(72),
$ CONVPL(9),SDPOL(9,2),PHISPL(2),SUM(2,5)
$ DIMENSION PPOL(2),PIPOL(2),UPOL(36,1),VPOL(36,1),TPOL(36,1),
$ SHPOL(36,1),QPOL(9,4,2),0(72,10,4,1)
$ DIMENSION PIPOL(2),VTPOL(36,1),VTPOL(36,1),TPOL(36,1),
$ SHTPOL(36,1),QPOL(9,4,2),GT(72,10,4,1)
$ COMMON/SMITH/OSM(72,9,4,5),PSMI(72,5),OSMFOL(9,4,2),PSMPOL(2)
INTEGER*2 ALBEDO
REAL*4 KAPA,LAT
COMMON /CNTRL/,
  
```

```

* JSP, JNP, IM,RLAY,PTOP,ISTART,JSPPI,JMPM],FIM,NLAYMI,NLAYPI,
* JI,JM,KM,TAUT,IROT,MROT,JTEST,ITEST,
* NR,JAYS(12),INCS(11),JDB,JNB,DLAT,DLON,
* DT,TAU,ITAU,XINT,TDAY,JDAY,TOFCAY,JDAT,E,MONTH(2),JYEAR,NSTEP,
* NCYCLE,NCOMP3,MDGAN,TAUP,TAUI,FAUE,FALO,DTMUL,
* PI,GRAY,RGAS,KAPA,PL,ED,FAU,NL,P,SP,MRCH,RSDIST,SIND,COSD,
* RHMAX,COX,DUMMYC(18),FALTER,DU,DUYA(99),
* XLABEL(20),SIG(20),OSIG(20),SIG(21),D,SGO(19),
* JIPS(11),JMPS(11),JIUS(11),JMU(11),PKS(11),KNGS(11),
* LAT(46),DXU(46),DAP(46),DYU(46),DVP(46),CXP(46),F(46),SINL(46),
* COSL(46),DUMMY(72),PHIS(2680,46),ALBEDO(72,46)
* DIMENSION U(72,40,1),V(72,40,1),W(72,40,1),SM(72,40,1),
* P(2880,1),TS(2880,1),SHS(2880,1),GT(2680,1),G(2880,1),C(300)
EQUIVALENCE (TS(1,1),PHIS(72,1)),(G(1,1),SHS(1,1)),(C(1,1),JSP)
$ (GT(1,1),PU(72,1))
COMMON/COMP/2/PI,TPOL
COMMON/WORK/CONV(72,5,46)
COMMON/WORK2/PU(2880,46)
DIMENSION UT(72,40,1),VT(72,40,1),PT(640,1),PV(72,2),
*PT(2880,1),SD(72,9,1),PIT(640,1),CONV(1,1,19),
EQUIVALENCE (SD(1,1,1),CONV(1,1,19)),
$ PIT(1,1))
COMMON /DEBUG/ISWTCH
C
COMMON/SDOT48/SDOT(72,46,8),OMEGA(72,46,5),MNXPTS,RADT(72,46,10)
+ ,COMG
LOGICAL COMG
C
INTEGER*2 RADT,OMEGA
INTEGER*2 ISDOT(72,46,8)
C
EQUIVALENCE (SDOT(1,1,5),ISDOT(1,1,1))
C
COMMON/SMO00/ISMTH,NPC
QSAT,IM,PR) = 622 * EXP:21*65604 - 5417.983 / TM) / PR
DATA J272/
JMS2 = JM - 2
FICO = 0 * DT
F2CO = -0.5 * DT
  
```

DATE 12/12/79-0783.48

780K BYTES OF CORE NOT USED

LABEL	ADDR
20	000C6B
202	000F5C
29	00150A
58	0018A4
48	001AC2
64	001D82
75	002098
87	0022A4
96	0023A2
110	00263E

LABEL	ADDR
BAMI 2320	
BAMI 2330	
BAMI 2340	
BAMI 2350	
BAMI 2360	
BAMI 2370	
BAMI 2380	
BAMI 2390	
BAMI 2400	
BAMI 2410	
BAMI 2420	
BAMI 2430	
BAMI 2440	
BAMI 2450	
BAMI 2460	
BAMI 2470	
BAMI 2480	
BAMI 2490	
BAMI 2500	
BAMI 2510	
BAMI 2520	
BAMI 2530	
BAMI 2540	
BAMI 2550	
BAMI 2560	
BAMI 2570	
BAMI 2580	
BAMI 2590	
BAMI 2600	
BAMI 2610	
BAMI 2620	
BAMI 2630	
BAMI 2640	
BAMI 2650	
BAMI 2660	
BAMI 2670	
BAMI 2680	
BAMI 2690	
BAMI 2700	
BAMI 2710	
BAMI 2720	
BAMI 2730	
BAMI 2740	
BAMI 2750	
BAMI 2760	
BAMI 2770	
BAMI 2780	
BAMI 2790	
BAMI 2800	
BAMI 2810	
BAMI 2820	
BAMI 2830	
BAMI 2840	
BAMI 2850	
BAMI 2860	
BAMI 2870	

ORIGINAL PAGE IS
OF POOR QUALITY

```

ISN 0029      L = J + 2
ISN 0030      JP2 = J + 1
ISN 0031      JS1 = J + 1
ISN 0032      JS1 = J - 1
ISN 0033      JS2 = J - 2
ISN 0034      JP2MOD = MOD(JP2-1,S) + 1
ISN 0035      JPMOD = MOD(JP1-1,S) + 1
ISN 0036      JMOD = MOD(J-1,S) + 1
ISN 0037      JS1MOD = MOD(JS1-1,S) + 1
ISN 0038      JS2MOD = MOD(JS2-1,S) + 1
ISN 0039      IF (J .GT. 2)
ISN 0040          GO TC 2150
ISN 0041      JS1 = 1
ISN 0042      JS2 = 1
ISN 0043      JS1MOD = 1
ISN 0044      JS2MOD = 1
ISN 0045      C * * BEGIN OF LAYER LOOP
ISN 0046      C * * CALC OF PV
ISN 0047      C
ISN 0048      C
ISN 0049      2150 IF (J .EQ. JM)
ISN 0050          GO TC 2158
ISN 0051          DO 2154 I = INC,IM,INC
ISN 0052              PV(I,1) = EXP(J) * P(I,J) + DAP(JP1) * P(I,JP1)
ISN 0053              * V(I,L,J)
ISN 0054              PV(I,2) = EXP(J) * P(I,J) + EXP(JP2) * P(I,JP2)
ISN 0055              * V(I,L,JP2)
ISN 0056          CONTINUE
ISN 0057      C * * PU CALCULATION : OMIT FOR J=1, SOUTH POLE)
ISN 0058      C
ISN 0059      C
ISN 0060      2154 IF (J .EQ. 1)
ISN 0061          GO TC 2225
ISN 0062          ISI = IM - 1
ISN 0063          PUISI = DYU(J) * P(ISI,J) + U(ISI,L,J)
ISN 0064          PUIPI = DYU(J) * P(I,J) + U(I,L,J)
ISN 0065          DO 2160 IPI = INC,IM,INC
ISN 0066              PUIPI = DYU(J) * P(IPI,J) + U(IPI,L,J)
ISN 0067              PUISI = PUISI + PUI
ISN 0068              PUIPI = PUIPI
ISN 0069              ISI = I
ISN 0070              I = IPI
ISN 0071          CONTINUE
ISN 0072      C * * HORIZONTAL ADVECTION OF MOMENTUM, TEMPERATURE, MOISTURE
ISN 0073      C * * COMPUTE FLUXES FIRST IN THE LONGITUDINAL (I) DIRECTION
ISN 0074      C
ISN 0075      C
ISN 0076      2160 ISI = IM - 1
ISN 0077          I = IPI
ISN 0078          DO 2223 IPI = INC,IM,INC
ISN 0079              FLUX1 = FICO * PUI(SI,I)
ISN 0080              FLUX2 = F2CC * PUI(SI,2)
ISN 0081              DO 2223 N = 1,4
ISN 0082                  QFLUX1 = FLUX1 + (Q(SI,L,N,J) + Q(I,L,N,J))
ISN 0083                  QFLUX2 = FLUX2 + (Q(SI,L,N,J) + Q(IPI,L,N,J))
ISN 0084                  IF (N.EQ.4) GO TO 110
ISN 0085                  IF (Q(SI,L,4,J).GT.0) GO TO 110
ISN 0086                  IF (QFLUX1.GT.0) QFLUX1=0
ISN 0087                  IF (QFLUX2.GT.0) QFLUX2=0
ISN 0088                  IF (Q(I,L,N,J).LE.0) .AND. (QFLUX1.LT.0) QFLUX1=0
ISN 0089                  IF (Q(IPI,L,N,J).LE.0) .AND. (QFLUX2.LT.0) QFLUX2=0
ISN 0090              CONTINUE
ISN 0091              Q(IPI,L,N,J) = Q(IPI,L,N,J) + QFLUX2
ISN 0092              Q(I,L,N,J) = Q(I,L,N,J) + QFLUX1
ISN 0093              Q(I,SI,L,N,J) = Q(I,SI,L,N,J) - QFLUX1 - QFLUX2
ISN 0094          CONTINUE
ISN 0095          ISI = I
ISN 0096          I = IPI
ISN 0097          IF (J .EQ. JM)
ISN 0098              GO TO 2237
ISN 0099      C * * NOW DO FLUX CALCULATION IN THE LATITUDINAL (J) DIRECTION
ISN 0100      C
ISN 0101      C
ISN 0102      2225 DO 2230 I = INC,IM,INC
ISN 0103              FLUX1 = FICO * PV(I,1)
ISN 0104              FLUX2 = F2CC * PV(I,2)
ISN 0105              DO 2230 N = 1,4
ISN 0106                  QFLUX1 = FLUX1 + (Q(I,L,N,J) + Q(I,L,N,JP1))
ISN 0107                  QFLUX2 = FLUX2 + (Q(I,L,N,J) + Q(I,L,N,JP2))
ISN 0108                  IF (N.EQ.4) GO TO 120
ISN 0109                  IF (Q(I,L,N,JP1).GT.0) GO TO 115
ISN 0110                  IF (QFLUX1.GT.0) QFLUX1=0
ISN 0111                  IF (QFLUX2.GT.0) QFLUX2=0
ISN 0112                  IF (Q(I,L,N,JP1).LE.0) .AND. (QFLUX1.LT.0) QFLUX1=0
ISN 0113                  IF (Q(I,L,N,JP2).LE.0) .AND. (QFLUX2.LT.0) QFLUX2=0
ISN 0114              CONTINUE
ISN 0115              Q(I,L,N,JP2) = Q(I,L,N,JP2) + QFLUX2
ISN 0116              Q(I,L,N,JP1) = Q(I,L,N,JP1) + QFLUX1

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2230 QT(I,L,N,J) = QT(I,L,N,J) - QFLUX1 - QFLUX2
CONTINUE
C * * * FOLLOWING IS CORRECTION FOR CASE J=2
C * * * P(I,0) = P(INDEX(I),2)
C * * * V(I,L,2) = -V(INDEX(I),L,2) ETC.
C * * * PV(I,2) FOR J=0 IS EQUAL TO: -CONV(I,L,2) * SEE ABOVE
C * * *
IF(J.GT.1)
DO 2240 I = INC,IM,INC
CONV(I,L,2)=0
CONTINUE
C * * * CONTINUITY EQUATION
C * * * DUE TO 1 LINE OF LATITUDE AT A TIME THE DEFINITION OF PV,
C * * * AND 4TH ORDER FORM OF CONV, WE COMPUTE IT AS FOLLOWS:
C * * *
2250 I22 = IM - 1
IS1 = IM
DO 2400 I = INC,IM,INC
CONV(I,L,JP2) = -PV(I,I,J2)
CONV(I,L,JP1) = CONV(I,L,JP1) + E * PV(I,I,1)
CONV(I,L,J) = CONV(I,L,J) - 8 * PV(I,I,1) + PV(I,I,2)
\$ 8 * PV(I,I,1) - PV(I,I,1) + (PV(I,2) - PV(I,2,2))
I22 = IS1
IS1 = I
CONTINUE
C * * * FOLLOWING IS CORRECTION FOR J=JM.
C * * * U(I,L,JM2) = -U(INDEX(I),L,JM), ETC.
C * * *
2237 IF (J.EQ.1) JM1
I22 = IM - 2
I = IM - 1
I1 = IM
DC 2250 IP2 = INC,IM,INC
PVLJM = DXPLJM * V(I,L,JM)
PV2JM = 0
CONV(I,L,JM) = CONV(I,L,JM) - 8 * PVLJM
\$ + PV2JM + DYU(I,I) * (P(I,1,JM) * U(I,1,L,JM)
\$ - P(I,1,JM) * U(I,1,L,JM)) + (P(I,2,J) * U(I,2,L,JM)
\$ - P(I,2,JM) * U(I,2,L,JM))
CONTINUE
C * * * CALCULATION OF THE REMAINDER OF QT(I,L,N,JM)
C * * *
FLUX1 = FICO * PVLJM
FLUX2 = F2CO * PVLJM
DO 2255 N = 1,4
DELUX1 = FLUX1 * (I(L,N,JM) + S(I,L,N,NP1))
DELUX2 = FLUX2 * (I(L,N,JM) + S(I,L,N,NP1))
IF(N.EQ.1) GO TO 126
IF(I(L,N,JM).GT.0) GO TO 126
IF(DELUX1.GT.0) DELUX1=0
IF(DELUX2.GT.0) DELUX2=0
IF(I(L,N,NP1).LE.0) AND (QFLUX1.LT.0) QFLUX1=0
IF(I(L,N,NP1).GT.0) AND (QFLUX2.LT.0) QFLUX2=0
1: QFLUX2 = I(L,N,JM) - QT(I,L,N,J) - OFLUX2
QT(I,L,N,JM) = QT(I,L,N,J) - OFLUX2
CONTINUE
C * * * NOTE: MULTIPLY BY DSIG(L) FOR LAT J ONLY.
C * * *
2405 DO 2410 I = INC,IM,INC
CONV(I,L,J) = CONV(I,L,J) * DSIG(L)
CONTINUE
2410 I = L + 1
IF (L.EQ. NLAY) GO TO 2150
C * * * END OF LAYER LOOP
C * * * COMPUTATION OF SIGMA DOT AND NEW SURFACE PRESSURE
C * * * SKIP THIS COMPUTATION IF (J.EG.1) (SOUTH POLE)
C * * * P(I,I,J) IS EQUIVALENT TO CONV(I,9,J)
C * * *
IF (J.EQ.1)
DO 2440 I = INC,IM,INC
\$ P(I,I,J) = CONV(I,9,J)
DO 2420 L = 1,NLAYM1
P(I,I,J) = P(I,I,J) + CONV(I,L,J)
CONTINUE
2420 SD(I,L,J) = CONV(I,I,J) - DSIG(I) * P(I,I,J)
\$ (FICOMG) OMEGA(I,J,1) = SD(I,I,J)/DXPLJM * E6

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GO, TC 2440
IF (NLAY * EQ, 2)
DO 2430 L = 2, NLAYM1
SD(I, L, J) = SD(I, L-1, J) + CONV(I, L, J) - (SIG(L) * PIT(I, J)
IF (COMG) OMEGA(I, J, L) = SD(I, L, J) / DXYP(J) * #10.ES
CONTINUE
CONTINUE
2430
2440
DO 2450 I = INC, IM, INC
DELPT = DT * PIT(I, J) / DXYP(J)
IF (I = 1)
IF (EQ, 1) I = 1
PT(I, J) = PT(I, J) + DT * PIT(I, J) / DXYP(J)
CONTINUE
2450 IF (NOT, COMG) GO TO 2456
DO 2455 L = 1, NLAY
DC 2455 I = I, IM
OMEGA(I, J, L) = OMEGA(I, J, L) + SIG(L) * PIT(I, J) / DXYP(J)
2456 CONTINUE
C * * * VERTICAL ADVECTION OF MOMENTUM AND TEMPERATURE
C * * *
DO 2490 L = 1, NLAYM1
LPI = L + 1
C
C
C FIRST DO MOMENTUM
DO 2470 N = 1, 2
DO 2470 I = INC, IM, INC
SFLUX = F2CO * SD(I, L, J) * (O(I, L, N, J) + O(I, L, P1, N, J))
OT(I, L, P1, N, J) = OT(I, L, P1, N, J) - SFLUX / (SIG(L, P1)
OT(I, L, N, J) = OT(I, L, N, J) + SFLUX / (SIG(L)
CONTINUE
2470
C * * * VERTICAL ADVECTION OF TEMPERATURE
C * * * TRYING GISS CODE FOR PK CALC. INSTEAD OF 4TH ORDER CODE
C * * *
DO 2480 I = INC, IM, INC
PL1 = SIG(L) * P(I, J) + PTOP
PL2 = SIG(L, P1) * P(I, J) + PTOP
PK1 = EXPBK(P, 1)
PK2 = EXPBK(PL2)
TFLUX = F2CO * SD(I, L, J) * (T(I, L, J) / P1 + T(I, L, P1, J) / PK2)
TT(I, L, P1, J) = TT(I, L, P1, J) - PK1 * TFLUX / DSIG(L)
TT(I, L, J) = TT(I, L, J) + PK1 * TFLUX / DSIG(L)
CONTINUE
RETURN
2480 IF (J .LT. JM)
2490
C * * * CALCULATIONS AT THE POLES
C * * *
C * * * WHEN (J=1) OR (J=JM) WE DO THE CALCULATIONS AT THE POLES.
C * * * MOMENTUM, TEMPERATURE, SPECIFIC HUMIDITY AT SOUTH POLE.
C * * * (M = 1) AND NORTH POLE (M = 2)
C * * * JPOL(K, M) = (2, 3, JM, JM-1)
C * * * COM1 = DXPI2) / (3 PI * A * A * SIN(OLAT / 2) ** 2)
C * * * COM2 = -DXPI3) / (12 PI * A * A * SIN(OLAT) ** 2)
C * * * (K = 1) CORRESPONDS TO THE LATITUDE LINE NEAREST TO THE POLE.
C * * *
2600 M = 1
IF (J, EQ, JM)
PITPOL(M) = 0.
COEF1 = (-1.) ** M
COEF2 = -COEF1
DO 2500 L = 1, NLAY
DO 2505 N = 1, 5
SUM(K, N) = 0.
CONTINUE
2505
C * * * HORIZ. ADVECS OF MOMENTUM, TEMPERATURE, SPECIFIC HUMIDITY
C * * * AT THE NORTH AND SOUTH POLES.
C * * * 4/4/78 - PIV TO BE CHANGED TO PIV FOR: -PIEQLONGPIE
C * * *
DO 2510 I = INC, IM, INC
DO 2510 K = 1, 2
IF (M, EQ, 2) JKP = JNP
JK = JPOL(K, M)
PIV = P(I, JK) * V(I, L, JK) + V(I, L, J, K) * SIN(LON(I)
SUM(K, 1) = SUM(K, 1) - PIV * (COEF1 * V(I, L, JK) + COS(LON(I)
SUM(K, 2) = SUM(K, 2) - PIV * (COEF2 * U(I, L, JK) + COS(LON(I)
SUM(I, L, JK) * SIN(LON(I))
  
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JEND	THBAE	R	R	00000				
JMS1	ADLDP	R	R	00000				
JMS2	PSKAPA	R	R	00000				
JMS3	CP	R	R	00000				
JMS4	IMD2	R	R	00000				
JMS5	JP2	R	R	00000				
JMS6	JPOL	R	R	00000				
JMS7	INDEX	R	R	00000				
JMS8	PHISPL	R	R	00000				
JMS9	CON1	R	R	00000				
JMS10	CON2	R	R	00000				
JMS11	THBAE	R	R	00000				
JMS12	ADLDP	R	R	00000				
JMS13	PSKAPA	R	R	00000				
JMS14	CP	R	R	00000				
JMS15	IMD2	R	R	00000				
JMS16	JP2	R	R	00000				
JMS17	JPOL	R	R	00000				
JMS18	INDEX	R	R	00000				
JMS19	PHISPL	R	R	00000				
JMS20	CON1	R	R	00000				
JMS21	CON2	R	R	00000				
JMS22	THBAE	R	R	00000				
JMS23	ADLDP	R	R	00000				
JMS24	PSKAPA	R	R	00000				
JMS25	CP	R	R	00000				
JMS26	IMD2	R	R	00000				
JMS27	JP2	R	R	00000				
JMS28	JPOL	R	R	00000				
JMS29	INDEX	R	R	00000				
JMS30	PHISPL	R	R	00000				
JMS31	CON1	R	R	00000				
JMS32	CON2	R	R	00000				
JMS33	THBAE	R	R	00000				
JMS34	ADLDP	R	R	00000				
JMS35	PSKAPA	R	R	00000				
JMS36	CP	R	R	00000				
JMS37	IMD2	R	R	00000				
JMS38	JP2	R	R	00000				
JMS39	JPOL	R	R	00000				
JMS40	INDEX	R	R	00000				
JMS41	PHISPL	R	R	00000				
JMS42	CON1	R	R	00000				
JMS43	CON2	R	R	00000				
JMS44	THBAE	R	R	00000				
JMS45	ADLDP	R	R	00000				
JMS46	PSKAPA	R	R	00000				
JMS47	CP	R	R	00000				
JMS48	IMD2	R	R	00000				
JMS49	JP2	R	R	00000				
JMS50	JPOL	R	R	00000				
JMS51	INDEX	R	R	00000				
JMS52	PHISPL	R	R	00000				
JMS53	CON1	R	R	00000				
JMS54	CON2	R	R	00000				
JMS55	THBAE	R	R	00000				
JMS56	ADLDP	R	R	00000				
JMS57	PSKAPA	R	R	00000				
JMS58	CP	R	R	00000				
JMS59	IMD2	R	R	00000				
JMS60	JP2	R	R	00000				
JMS61	JPOL	R	R	00000				
JMS62	INDEX	R	R	00000				
JMS63	PHISPL	R	R	00000				
JMS64	CON1	R	R	00000				
JMS65	CON2	R	R	00000				
JMS66	THBAE	R	R	00000				
JMS67	ADLDP	R	R	00000				
JMS68	PSKAPA	R	R	00000				
JMS69	CP	R	R	00000				
JMS70	IMD2	R	R	00000				
JMS71	JP2	R	R	00000				
JMS72	JPOL	R	R	00000				
JMS73	INDEX	R	R	00000				
JMS74	PHISPL	R	R	00000				
JMS75	CON1	R	R	00000				
JMS76	CON2	R	R	00000				
JMS77	THBAE	R	R	00000				
JMS78	ADLDP	R	R	00000				
JMS79	PSKAPA	R	R	00000				
JMS80	CP	R	R	00000				
JMS81	IMD2	R	R	00000				
JMS82	JP2	R	R	00000				
JMS83	JPOL	R	R	00000				
JMS84	INDEX	R	R	00000				
JMS85	PHISPL	R	R	00000				
JMS86	CON1	R	R	00000				
JMS87	CON2	R	R	00000				
JMS88	THBAE	R	R	00000				
JMS89	ADLDP	R	R	00000				
JMS90	PSKAPA	R	R	00000				
JMS91	CP	R	R	00000				
JMS92	IMD2	R	R	00000				
JMS93	JP2	R	R	00000				
JMS94	JPOL	R	R	00000				
JMS95	INDEX	R	R	00000				
JMS96	PHISPL	R	R	00000				
JMS97	CON1	R	R	00000				
JMS98	CON2	R	R	00000				
JMS99	THBAE	R	R	00000				
JMS100	ADLDP	R	R	00000				
JMS101	PSKAPA	R	R	00000				
JMS102	CP	R	R	00000				
JMS103	IMD2	R	R	00000				
JMS104	JP2	R	R	00000				
JMS105	JPOL	R	R	00000				
JMS106	INDEX	R	R	00000				
JMS107	PHISPL	R	R	00000				
JMS108	CON1	R	R	00000				
JMS109	CON2	R	R	00000				
JMS110	THBAE	R	R	00000				
JMS111	ADLDP	R	R	00000				
JMS112	PSKAPA	R	R	00000				
JMS113	CP	R	R	00000				
JMS114	IMD2	R	R	00000				
JMS115	JP2	R	R	00000				
JMS116	JPOL	R	R	00000				
JMS117	INDEX	R	R	00000				
JMS118	PHISPL	R	R	00000				
JMS119	CON1	R	R	00000				
JMS120	CON2	R	R	00000				
JMS121	THBAE	R	R	00000				
JMS122	ADLDP	R	R	00000				
JMS123	PSKAPA	R	R	00000				
JMS124	CP	R	R	00000				
JMS125	IMD2	R	R	00000				
JMS126	JP2	R	R	00000				
JMS127	JPOL	R	R	00000				
JMS128	INDEX	R	R	00000				
JMS129	PHISPL	R	R	00000				
JMS130	CON1	R	R	00000				
JMS131	CON2	R	R	00000				
JMS132	THBAE	R	R	00000				
JMS133	ADLDP	R	R	00000				
JMS134	PSKAPA	R	R	00000				
JMS135	CP	R	R	00000				
JMS136	IMD2	R	R	00000				
JMS137	JP2	R	R	00000				
JMS138	JPOL	R	R	00000				
JMS139	INDEX	R	R	00000				
JMS140	PHISPL	R	R	00000				
JMS141	CON1	R	R	00000				
JMS142	CON2	R	R	00000				
JMS143	THBAE	R	R	00000				
JMS144	ADLDP	R	R	00000				
JMS145	PSKAPA	R	R	00000				
JMS146	CP	R	R	00000				
JMS147	IMD2	R	R	00000				
JMS148	JP2	R	R	00000				
JMS149	JPOL	R	R	00000				
JMS150	INDEX	R	R	00000				
JMS151	PHISPL	R	R	00000				
JMS152	CON1	R	R	00000				
JMS153	CON2	R	R	00000				
JMS154	THBAE	R	R	00000				
JMS155	ADLDP	R	R	00000				
JMS156	PSKAPA	R	R	00000				
JMS157	CP	R	R	00000				
JMS158	IMD2	R	R	00000				
JMS159	JP2	R	R	00000				
JMS160	JPOL	R	R	00000				
JMS161	INDEX	R	R	00000				
JMS162	PHISPL	R	R	00000				
JMS163	CON1	R	R	00000				
JMS164	CON2	R	R	00000				
JMS165	THBAE	R	R	00000				
JMS166	ADLDP	R	R	00000				
JMS167	PSKAPA	R	R	00000				
JMS168	CP	R	R	00000				
JMS169	IMD2	R	R	00000				
JMS170	JP2	R	R	00000				
JMS171	JPOL	R	R	00000				
JMS172	INDEX	R	R	00000				
JMS173	PHISPL	R	R	00000				
JMS174	CON1	R	R	00000				
JMS175	CON2	R	R	00000				
JMS176	THBAE	R	R	00000				
JMS177	ADLDP	R	R	00000				
JMS178	PSKAPA	R	R	00000				
JMS179	CP	R	R	00000				
JMS180	IMD2	R	R	00000				
JMS181	JP2	R	R	00000				
JMS182	JPOL	R	R	00000				
JMS183	INDEX	R	R	00000				
JMS184	PHISPL	R	R	00000				
JMS185	CON1	R	R	00000				
JMS186	CON2	R	R	00000				
JMS187	THBAE	R	R	00000				
JMS188	ADLDP	R	R	00000				
JMS189	PSKAPA	R	R	00000				
JMS190	CP	R	R	00000				
JMS191	IMD2	R	R	00000				
JMS192	JP2	R	R	00000				
JMS193	JPOL	R	R	00000				
JMS194	INDEX	R	R	00000				
JMS195	PHISPL	R	R	00000				
JMS196	CON1	R	R	00000				
JMS197	CON2	R	R	00000				
JMS198	THBAE	R	R	00000				
JMS199	ADLDP	R	R	00000				
JMS200	PSKAPA	R	R	00000				
JMS201	CP	R	R	00000				
JMS202	IMD2	R	R	00000				
JMS203	JP2	R	R	00000				
JMS204	JPOL	R	R	00000				
JMS205	INDEX	R	R	00000				
JMS206	PHISPL	R	R	00000				
JMS207	CON1	R	R	00000				
JMS208	CON2	R	R	00000				
JMS209	THBAE	R	R	00000				
JMS210	ADLDP	R	R	00000				
JMS211	PSKAPA	R	R	00000				
JMS212	CP	R	R	00000				
JMS213	IMD2	R	R	000				

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NAME OF COMMON BLOCK * C:W799E OF BLOCKCOMMON 000C08 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
PITPOL R#4 000000

NAME OF COMMON BLOCK * *W6R9E OF BLOCKCOMMON 01D1C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CONV R#4 000000
EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 000900
SD 000000 PIT VARIABLE OFFSETE

NAME OF COMMON BLOCK * *D599E OF BLOCKCOMMON 081C00 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
PU R#4 000000
EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 001600
GT 000B40

NAME OF COMMON BLOCK * *D5_7E OF BLOCKCOMMON 000C04 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
ISBTR I#4 NoR.

NAME OF COMMON BLOCK * *S0T19E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
SOOT R#4 000003 OMEGA I#2 019E00 VAR. NAME TYPE REL. ADDR. ME
COMG L#4 0389A4 MIMXTS I#4 NoR.
EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 00CF00
ISSOT VARIABLE OFFSETE

NAME OF COMMON BLOCK * *S619E OF BLOCKCOMMON 000C08 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME

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(SN 0029 JSI = J - 1
(SN 0030 JS2 = J - 2
(SN 0031 JP2MOD = MOD(JP2-1.5) + 1
(SN 0032 JPIMOD = MOD(JPI-1.5) + 1
(SN 0033 JMOD = MOD(J-1.5) + 1
(SN 0034 JS1MOD = MOD(JS1-1.5) + 1
(SN 0035 JS2MOD = MOD(JS2-1.5) + 1
C * *
(SN 0036 JPKP2 = MOD(JP2-1.3) + 1
(SN 0037 JPKP1 = MOD(JPI-1.3) + 1
(SN 0038 JPK = MOD(J-1.3) + 1
(SN 0039 DTKAPA = DT * KAPA
(SN 0040 IF(JGT1)
(SN 0041 JP2 = 1
(SN 0042 JP2MOD = 1
(SN 0043 JPKP2 = 1
(SN 0044
(SN 0045
C * * * * * CORIOLIS FORCE FOR 4TH ORDER DICES NOT FOLLOW GISS
C * * * * * METHOD OF COMPUTATION
C * * * * * ADLDP = (12 * RAD * DLAT * DLON = 12. * DYU(J) * DLON)
C
3001 DO 3000 L = 1,NLAY
DO 3000 I = INC,IM,INC
FX = DT * (F(J) * DXYP(J) + ADLDP * SIML(J) * U(I,L,J))
UT(I,L,J) = UT(I,L,J) + FX * P(I,J) * V(I,L,J)
VT(I,L,J) = VT(I,L,J) - FX * P(I,J) * U(I,L,J)
CONTINUE
3000
C * * * * * M A I N L O O P
C
C IF(JGTJMJ) GO TO 3032
C
3005 DO 3030 LX = 1,NLAY
LPI = L + 1
IF(JAE0.1ORJAE0.JNP) GO TO 3141
COMPUTATION OF THE ENERGY TERM FOR THE THERMODYNAMIC
EQUATION FOR J=2,3,....JM-LVJM
SET UP CYCLIC I-INDICES
IS2 = IM - 3
IS1 = IM - 2
I = IM - 1
IPI = IM
SET UP J-INDICES
NDXJP2 = JP2
IF(JAE0.JMJ) NDXJP2 = JM
NDXJS2 = JS2
IF(JAE0.2) NDXJS2 = 2
PERFORM CALCULATIONS
DO 3140 IP2 = INC,IM,INC
SKT = DTKAPA*SIG(L)*(I,L,J)/(PTCP+SIG(L)*P(I,J))
PIV = DXP(J)*V(I,L,J)
SET UP PROPER I-INDEX AND PRESSURE TERMS
NDXI = I
IF(JAE0.JMJ) NDXI = INDEX(I)
GI = P(NDXI,NDXJP2)
NDXI = I
IF(JAE0.2) NDXI = INDEX(I)
G2 = P(NDXI,NDXJS2)
TT(I,L,J) = TT(I,L,J) + SKT*P(I,J)*PIT(I,J) + DYU(J) *
I U(I,L,J)*G2 + P(PIV,J)*P(I,J) - P(IP2,J) +
I P(IS2,J) + PIV*(P(I,JPI-I,JSI) - GI + G2)
RESET CYCLIC INDICES
IS2 = IS1
IS1 = I
I = IPI
IPI = IP2
C
3140 CONTINUE
C

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ISN 0086 3141 CONTINUE
ISN 0087 IF(J.EQ.JM) GO TO 3030
ISN 0089 C * * COMPUTE PK AT LATITUDE JP2 (USING GISS METHOD)
IF(L.LT.NLAY) GO TO 3055
3007 DO 3010 I = 1,IM
PLI = SIGE(I) * P(I,JP2) + PTOP
PKI = PLI * EXPBK(PLI)
LL = I
DO 3010 LLPI = 2,NLAYPI
PL2 = SIGE(LLPI) * P(I,JP2) + PTEP
PK2 = LL * EXPBK(PL2)
PK1(LL,JKPK2) = (PK2 - PK1) / (DIFF(LL) * P(I,JP2))
PK1 = PK2
LL = LLPI
3010 CONTINUE
C * * COMPUTATION OF PHI (GEOCENTRAL) AT JF2 FOR L=NLAY,
C * * PSKAPA=1000*K, SIGE(I)=0,
C * * HERE PHI IS NORMALIZED, I.E., EQUALS (STANDARD PHI)-PHIBAR.
IF (.LT. NLAY) GO TO 3055
DO 3050 I = INC,IM,INC
PHI(LL,JP2MOD) = PHIS(I,JP2) + CPTH*(PK(I,L,JKPK2) - PSKAPA)
LLMI = I
DO 3050 LL = 1,NLAY
LLPI = LL + I
IF(LL.EQ.NLAY) LLPI = LL
DUM1 = DSIG(LL)*(PK(I,LLPI,JKPK2) - PK(I,LLMI,JKPK2))
DUM2 = SIGE(LL)*PK(I,LLPI,JKPK2) - PK(I,LLMI,JKPK2)
WL = SIG(LL) * P(I,JP2)
DUM3 = WL * R * DSIG(LL) / (WL + PTOP)
LLM1 = LL
C * * COMPUTATION OF PHI AT JF2
3050 I : DUM1+DUM2/PK(I,LL,JKPK2); I+1,LL,JP2
CONTINUE
GO TO 3065
C * * COMPUTATION OF PHI AT JF2
3055 DO 3060 I = INC,IM,INC
PHI(I,JP2MOD) = PHI(I,LLPI,JP2MOD) + CFC2 * (PK(I,LLPI,JP2) -
$ PK(I,LL,JP2)) / PK(I,LL,JP2)
$ PK(I,LLPI,JP2) - ITHBAR)
CONTINUE
3060 C * * FOR PHI(I,LL,0) USE INDEX(I)
C * * COMPUTE W(I,LL,JP2MOD) FOR PRESSURE GRADIENT TERM AND SKT.
C * * PIV FOR ENERGY TERM IN THERMODYNAMICS EQ.
3065 IF(JP2.EQ.1.OR.JP2.EQ.JNP) GO TO 3030
DO 3070 I = INC,IM,INC
IHBAR = IHBAR * PK(I,LL,JKPK2)
IPRIME = I(L,JP2) - IHBAR
PRESUR = PTOP + SIGE(I) * R * IPRIME / PRESUR
W(I,LL,JP2MOD) = SIGE(I) * R
CONTINUE
GO TO 3030
C * * IF :J=1) RETURN AND COMPUTE PHI(I,LL,2),W(I,LL,2),.....
C * * FOR SOUTH POLE CALCULATIONS
3030 CONTINUE
GO TO 3032
3074 IF(J.EQ.JM)
JP2 = JP2 + 1
JP2MOD = JP2MOD + 1
JKPK2 = JKPK2 + 1
C PRESSURE GRADIENT (V EQUATION) FOR J=2
3032 DO 3031 LX = 1,NLAY
L = NLAYPI - LX
IF(J.EQ.1) GO TO 3111
3080 IF (J.GT. 2)
DO 3083 I = INC,IM,INC

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BAMI 8000
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BAMI 8120
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BAMI 8420

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C
ISN 0147 PIV = DXP(J) * V(I,L,J)
ISN 0148 P1 = PHI(INDEX(I),L,J)
ISN 0149 G1 = P(INDEX(I),J)
ISN 0150 P2 = PHI(I,L,I)
ISN 0151 G2 = P(I,I)
ISN 0152 W1 = W(I,L,J)
ISN 0153 VT(I,L,J) = VT(I,L,J) + DT * DXP(J) * P(I,J) * (G1 - G2)
$ * P2 - PHI(I,L,J) * P1 + W1 * (G2 - P(I,J))
$ * PHI(I,L,J) * P2 - G1))
CONTINUE
3083
C * * COMPUTATION OF THE PRESSURE GRADIENT FOR ZONAL (U)
C * * MOMENTUM EQ.
3085 IS2 = IM - 3
IS1 = IM - 2
I = IM - 1
IPI = IM
DO 3090 IP2 = INC, IM, INC
W1 = W(I,L,JMOD)
$ * NEXT TWO LINES HAVE ADDED (I)S
UT(I,L,J) = UT(I,L,J) + DT * DU(J) * P(I,J) * (G1 - G2)
$ * PHI(I,L,JMOD) - PHI(IPI,L,JMOD) + W1 * (P(I,S1) - P(I,S2))
$ * P1 * IP2 - P(I,S2))
IS2 = IS1
IS1 = I
I = IPI
IPI = IP2
3090 IF (J.EQ.2.OR.J.EQ.JM) GO TO 3135
CONTINUE
C * * COMPUTATION OF THE PRESSURE GRADIENT FOR
C * * MERIDIONAL (V) MOMENTUM EQ.
C
DO 3100 I = INC, IM, INC
W1 = W(I,L,JMOD)
VT(I,L,J) = VT(I,L,J) + DT * DXP(J) * P(I,J) * (G1 - G2)
$ * PHI(I,L,JMOD) - PHI(I,L,JPI) + W1 * (P(I,S1) - P(I,S2))
$ * (P(I,JPI) + PHI(I,L,JPI) - PHI(I,L,JMOD) + W1 * (P(I,JPI) - P(I,JMOD) - PHI(I,L,JMOD) + W1 * (G3 - P(I,JMOD))))
CONTINUE
3100
C 3125 IF (J.LT.JM) GO TO 3031
C
C PRESSURE GRADIENT FOR J=JM
C
DO 3150 I = INC, IM, INC
G1 = P(I,JM)
G2 = P(I,JNP)
G3 = P(INDEX(I),JM)
W1 = W(I,L,JMOD)
VT(I,L,JM) = VT(I,L,JM) + DT * DXP(JM) * G1 * (G1 - G2)
$ * PHI(I,L,JMOD) - PHI(I,L,JPI) + W1 * (P(I,S1) - P(I,S2))
$ * (G3 - P(I,JMOD))
CONTINUE
3150
C * * CALCULATIONS AT THE POLES
C
C * * MOMENTUM AND TEMPERATURE AT SOUTH POLE (M = 1) AND AT
C * * NORTH POLE (M = 2)
C * * (K = 1) CORRESPONDS TO THE LATITUDE LINE NEAREST THE POLE
C * * JMOD(K,M) = 2,3,MOD(JM-1,5)+1,MOD(JM-1)-1,5+1
C * * CON3 = 4 * DT / (3 * RAD * SIN(SIN(SD(LAT)))
C * * CON4 = DT / (3 * RAD * SIN(SIN(DLAT)))
C * * CON5 = 1. / (3 * RAD * IMPLAT)
3111 M = 1
IF (J.EQ. JM) M = 2
JKP = 1
IF (M.EQ.2) JKP = JNP
JP2 = 1
JP2MOD = 1
JKP2 = 1
C * * CORIOLIS TERM AT THE POLES
C
F1 = F(JEND,M) * PPO(L,M)
UTPOL(L,M) = UTPOL(L,M) + DT * F1 * VPOL(L,M)
VTPOL(L,M) = VTPOL(L,M) - DT * F1 * UPOL(L,M)
COEF = -1) * M
DC 3200 N = 1,3
DO 3200 K = 1,2
SUM(K,N) = 0.
3200

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UPOL	XR	R04	000030	VR	NAME	TYPE	REL.	ADDR.	ME	VR	NAME	TYPE	REL.	ADDR.	ME	VR	NAME	TYPE	REL.	ADDR.	ME
ADLDP	F	R04	000030	F	ALPHA	C	R04	00000C	NrR	VR	CONZ	SF	R04	000170	NrR	VR	XINT	XR	R04	000170	NrR
DSIGO	F	R04	000024	C	DUNNY	C	R04	NrR	NrR	VR	INDEX	SF	R04	000174	NrR	VR	CONZ	XR	R04	000174	NrR
ISMTX	C	R04	NrR	C	ITEST	C	R04	NrR	NrR	VR	JMAD	F	R04	0002E4	NrR	VR	ISDOT	CE	R04	000178	NrR
JNEXT	C	R04	NrR	C	JMAD	C	R04	NrR	NrR	VR	JMAD	S	R04	0001B0	NrR	VR	JMAD	SF	R04	00017C	NrR
JMODP	F	R04	000074	C	JMAD	C	R04	NrR	NrR	VR	JTEST	S	R04	NrR	NrR	VR	JMAD	SF	R04	000180	NrR
MODPR	F	R04	000074	C	JMAD	C	R04	NrR	NrR	VR	JTEST	S	R04	NrR	NrR	VR	JMAD	SF	R04	000184	NrR
POLLS	F	R04	000074	C	JMAD	C	R04	NrR	NrR	VR	JTEST	S	R04	NrR	NrR	VR	JMAD	SF	R04	000188	NrR
OTPOL	SF	R04	000000	C	RADIM	C	R04	00018C	NrR	VR	PSIG	SF	R04	00018C	NrR	VR	OTPOL	XR	R04	000192	NrR
SHPOL	SF	R04	000000	C	RADIM	C	R04	00018C	NrR	VR	RHMAX	SF	R04	000000	NrR	VR	SHPOL	XR	R04	000196	NrR
CSLON	SF	R04	000000	C	FIXPI	C	R04	000014	NrR	VR	TTPOL	SF	R04	000000	NrR	VR	CSLON	XR	R04	000200	NrR
ISMTX	C	R04	000188	C	DKADA	C	R04	00019C	NrR	VR	ALBED	SF	R04	000000	NrR	VR	ISMTX	XR	R04	000204	NrR
JNEXT	C	R04	000188	C	EXPRTK	C	R04	00019C	NrR	VR	DTMULT	SF	R04	000000	NrR	VR	JNEXT	XR	R04	000208	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR	FALTER	SF	R04	000000	NrR	VR	MODPR	XR	R04	000212	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR	JSMOD	SF	R04	00019C	NrR	VR	OTPOL	XR	R04	000216	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR	NCYCLE	SF	R04	00019C	NrR	VR	SHPOL	XR	R04	000220	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR	PLAINT	SF	R04	000000	NrR	VR	CONVPL	XR	R04	000224	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR	SKTPOL	SF	R04	0001B0	NrR	VR	DUNNY	XR	R04	000228	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR	XLABEL	SF	R04	0001B0	NrR	VR	JMAD	XR	R04	000232	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000236	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000240	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000244	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000248	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000252	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000256	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000260	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000264	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000268	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000272	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000276	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000280	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000284	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000288	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000292	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000296	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000300	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000304	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000308	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000312	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000316	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000320	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000324	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000328	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000332	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000336	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000340	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000344	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000348	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000352	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000356	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000360	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000364	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000368	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000372	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000376	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000380	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000384	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000388	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000392	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000396	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000400	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000404	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000408	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000412	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000416	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000420	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000424	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000428	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000432	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000436	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000440	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000444	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000448	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000452	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000456	NrR
SHPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	SHPOL	XR	R04	000460	NrR
CONVPL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	CONVPL	XR	R04	000464	NrR
DUNNY	F	R04	000030	C	ISMTX	C	R04	000000	NrR	VR						VR	DUNNY	XR	R04	000468	NrR
JMAD	SF	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	JMAD	XR	R04	000472	NrR
MODPR	F	R04	000074	C	ISMTX	C	R04	000000	NrR	VR						VR	MODPR	XR	R04	000476	NrR
OTPOL	SF	R04	000000	C	ISMTX	C	R04	000000	NrR	VR						VR	OTPOL	XR	R04	000480	NrR

NAME OF COMMON BLOCK * 005#E OF BLOCKCOMMON 081600 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 PU R#A 000000
 EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE
 GT 000B40 CB 00168C VARIABLE OFFSETE

NAME OF COMMON BLOCK * 005_7E OF BLOCKCOMMON 000604 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 ISWTCB I#A N#R.

NAME OF COMMON BLOCK * S#OT#9E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 SDOT R#A 000000 QNEGA I#2 019E00 MINMXTS I#4
 COMG L#A 0389A4
 EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE
 ISDOT 00CF00 VARIABLE OFFSETE

NAME OF COMMON BLOCK * S#6-#E OF BLOCKCOMMON 000108 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 ISMTH I#4 N#R. NPC I#4

NAME OF COMMON BLOCK * C#M79#E OF BLOCKCOMMON 000C0E HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE


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ISN 0037      DUMNY(I)=Q(I,L,N,J)*ALPHS(Q(IPINC,L,N,J)+Q((IMINC,L,N,J)-
              $ Q(I,L,N,J)-Q(I,L,N,J))
              IMINC=I
              I=IPINC
              DO 20 A=INC,IM,INC
              OVI A JI=DUMNY(I)
              CONTINUE
              I=IM
              I=IM-INC
              DO 35 IPINC=INC,IM,INC
              DUMNY(I)=P(I,J)+ALPHS(P(IPINC,J)+P(IMINC,J)-P(I,J)-P(I,J))
              IMINC=I
              I=IPINC
              P(I,40)=INC,IM,INC
              CONTINUE
              GO TO 60
C * * * FOURIER SMOOTHING NEAR POLES
C 12/12 VERSION 3 SMSHAP ON P AND T
C IF(I,J)GT,JMIN)ANDS(J,L)GT,JMAX))OR(JO)EQ,JO)OR(JO)EQ,JMP) RETURN
C IF(SUBS=J)
C IF(SUBS=JMAX)
C N = 2.0*NSM / 1.4
C DO 650 I = 1,MLAY
C DO 650 J = 1,IM
C DATA(I) = Q(I,L,N,J)
C DO 670 I = 1,IMD2PI
C TRAN(I) = SMTH(I,JSUB) * TRAN(I)
C DO 680 I = 1,IM
C O((L,N,J) = DATA(I) / FLOAT(IM)
C CONTINUE
C CALL SMSHAP
C IF(JO)EQ,JM)
C DO 100 I = 1,IM
C SLP(I,J) = P(I,J)
C * * * TRANSFORM TO SEA LEVEL PRESSURE
C
C DO 110 I = 1,IM
C TSURF = TIO(SLP(I,J),T(L,NLAYM),J,T(L,NLAY,J))
C SLP(I,J) = (SLP(I,J) + PTOP) * SLEXP(PHIS(I,J),TSURF)
C CONTINUE
C 760 IF(IM)GT,01G0 TO 781
C DATA(I)=SLP(I,J)
C CONTINUE
C CALL FOURT2(DATA,IM,1,-1,0)
C DO 770 I=1,IMD2PI
C TRAN(I)=SMTH(I,JSUB)*TRAN(I)
C CONTINUE
C CALL FOURT2(TRAN,IM,1,1,-1)
C DO 790 I=1,IM
C SLP(I,J)=DATA(I)/FLOAT(IM)
C CONTINUE
C 781 CONTINUE
C * * * SMOOTHING ALONG LONGITUDE (SEE SMSHAP)
C
C NSM = 8
C IF(JO)LT,8)OR(JO)GT,JM-6) NSM = 4
C IF(JO)LT,6)OR(JO)GT,JM-4) NSM = 2
C IF(JO)LT,4)OR(JO)GT,JM-2) NSM = 1
C SGN = (-1)**NSM
C CX = 4**NSM
C DO 130 I = 1,IM
C DATA(I) = SLP(I,J)
C DO 140 N = 1,NSM
C I = IM
C DO 145 IP = I,IM
C CATA(I) = DATA(IP) - DATA(I)
C I = IP
C IS = IM
C DO 150 I = 1,IM
C DATA(I) = CATA(I) - CATA(IS)
C IS = I
C DO 160 I = 1,IM

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 INDEX 104 No.R
 PK 104 No.R
 PHISPL 104 No.R
 CUSLJN 104 No.R
 SDPOL 104 No.R

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000414
 VPOL 000430

VARIABLE OFFSETE 000438
 VPOL

SINLUN 104 No.R
 CONPOL 104 No.R
 POLES 104 No.R
 200414
 No.R

VARIABLE OFFSETE 00045C
 TPOL

NAME OF COMMON BLOCK * 083C30 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JSC 104 000000 000004
 P10 104 000010 000004
 P14 104 000020 000024
 MROT 104 000030 000024
 JAVE 104 000040 000024
 QAVT 104 000050 000024
 TITAY 104 000060 000024
 HSTED 104 000070 000024
 DTRULT 104 000080 000024
 KAPA 104 000090 000024
 NCLN 104 000100 000024
 SIND 104 000110 000024
 DUMRVC 104 000120 000024
 SIG 104 000130 000024
 JIPE 104 000140 000024
 K895 104 000150 000024
 OXP 104 000160 000024
 F 104 000170 000024
 PHIS 104 000180 000024
 VAR. NAME TYPE REL. ADDR. ME
 IM 104 000008
 JSPPI 104 000008
 RLAYPI 104 000008
 TAU 104 000008
 ITST 104 000008
 JSB 104 000008
 DT 104 000008
 IDAY 104 000008
 JMONTH 104 000008
 NCOMPN 104 000008
 TAVE 104 000008
 GRAY 104 000008
 ED 104 000008
 MRCM 104 000008
 SHINX 104 000008
 DUMRYA 104 000008
 SIGE 104 000008
 JIUS 104 000008
 LAT 104 000008
 DYP 104 000008
 COSL 104 000008
 VAR. NAME TYPE REL. ADDR. ME
 NLAY 104 00000C
 JNPMI 104 00000C
 JI 104 00000C
 IROT 104 00000C
 NR 104 00000C
 JNB 104 00000C
 TAU 104 00000C
 JDAY 104 00000C
 JYEAR 104 00000C
 NMOGAN 104 00000C
 TAOU 104 00000C
 RGAS 104 00000C
 FMU 104 00000C
 RSD 104 00000C
 CDI 104 00000C
 KLABEL 104 00000C
 DSIGO 104 00000C
 JMUS 104 00000C
 DRU 104 00000C
 DITP 104 00000C
 DUMMY 104 00000C

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000000
 VPOL

NAME OF COMMON BLOCK * 115670 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 F5DE 104 000000 000004
 F17 104 000004
 VAR. NAME TYPE REL. ADDR. ME
 F9B 104 000008

NAME OF COMMON BLOCK * 00570E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 ISATCP 104 000000 000004
 NPC 104 000004

NAME OF COMMON BLOCK * 00570E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 ISMTH 104 000000 000004
 NPC 104 000004

NAME OF COMMON BLOCK * 00570E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JMIN 104 000000 000004
 JMAX 104 000004

NAME OF COMMON BLOCK * 00570E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 PU 104 000000 000004
 JSUM 104 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000840
 GP

NAME OF COMMON BLOCK * 00570E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 DATA 104 000000 000004
 CATA 104 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000840
 GP

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000840
 GP

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000840
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120 DO 110 I = 1,NLAY
    GO TO 140
    POT(I,J,L) = T(I,L,J)
    CONTINUE
110 DO 100 I = 1,IM
    SLP(I,J) = P(I,J)
    CONTINUE
100 ** * TRANSFORM TO SEA LEVEL PRESSURE AND POTENTIAL TEMPERATURE ** *
C ** *
C
140 DO 210 I = 1,IM
    TSURF = T10(SL(I,J),POT(I,J,J.5))
    DO 220 L = 1,NLAY
        PKAPA = EXPB*(SLP(I,J) * SIGIL) + PTOP
        POT(I,J,L) = POT(I,J,L) / PKAPA
        CONTINUE
220 PHISX = PHIS(I,J)
    SLP(I,J) = (SLP(I,J) + PTOP) * SLEXP(PHISX,TSURF)
    IF (J.EQ.1)OR(J.EQ.JNP) GO TO 230
    CONTINUE
    GO TO 200
230 DO 240 I = 2,IM
    DO 250 L = 1,NLAY
        POT(I,J,L) = POT(I,J,L)
    CONTINUE
250 SLP(I,J) = SLP(I,J)
    CONTINUE
240 NF = -1
200 ** * SMOOTHING ALONG LONGITUDE ** *
C ** *
C ** *
C ** *
C ** *
C ** * 16TH ORDER AT ALL J VALUES.
1000 NF = NF + 1
    J = 2,JM
    NSM = 1
    IF (J.LT.8)OR(J.GT.JM-6) NSM = 4
    IF (J.LT.6)OR(J.GT.JM-4) NSM = 2
    IF (J.LT.4)OR(J.GT.JM-2) NSM = 1
    SGN = (-1)**NSM
    CX = NF * GT * NSM
    IF (NF * GT * NSM)
        GO TO 1024
    DATA(I) = SLP(I,J)
    CONTINUE
    GO TO 1028
1020 CONTINUE
    GO TO 1028
1024 DO 1026 I = 1,IM
    DATA(I) = POT(I,J,NF)
    CONTINUE
1026 DO 1030 N = 1,NSM
    I = IM
    DO 1035 IP = 1,IM
        CATA(IP) = DATA(IP) - DATA(I)
        I = IP
    CONTINUE
1035 IS = IM
    DO 1040 I = 1,IM
        DATA(I) = CATA(I) - CATA(IS)
        IS = I
    CONTINUE
1040 IF (NF * GT * 0)
    CONTINUE
1050 DO 1050 I = 1,IM
    SLP(I,J) = SLP(I,J) -
        DATA(I) / CX
    GO TO 1010
1054 DO 1056 I = 1,IM
    POT(I,J,NF) = POT(I,J,NF) -
        DATA(I) / CX
    CONTINUE
1056 CONTINUE
1010 GO TO 1010
C ** * SMOOTHING ALONG LATITUDE
C ** *
C ** * ALSO USE PERIODICITY ALONG COMPLETE MERIDIAN (J=1,2*JM)
C
CY=4**NSM
IM2 = IM / 2
JMT2 = JM * 2
DO 1110 I = 1,IM2
    IF (NF * GT * 0)
        DO 1120 J = 1,JM
            DATA(J) = SLP(I,J)
            DATA(J+JM) = SLP(I+IM2,JM-J+2)
        CONTINUE
    GO TO 1124
1120 CONTINUE
    GO TO 1128
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C * * * SHOULD PUT (-1) FOR WIND SMOOTHING IN 1126
C
1124 DO 1126 J = 1, JM
DATA(J) = POT(I, J, NF)
DATA(J+JM) = POT(I+IMD2, JM-J+2, NF)
CONTINUE
1126
1128 DO 1130 N=1, NSM
J = JMT2
DO 1135 JP = 1, JMT2
CATA(J) = DATA(JP) - DATA(J)
J = JP
CONTINUE
1135 JS = JMT2
DO 1140 J = 1, JMT2
DATA(J) = CATA(J) - CATA(JS)
JS = J
CONTINUE
1140
1130 IF (NF.GT. 0)
DO 1150 J = 1, JM
SLP(I, J) = SLP(I, J) - DATA(J) / CY
SLP(I+IMD2, JM-J+2) = SLP(I+IMD2, JM-J+2) - DATA(J+JM) / CY
CONTINUE
GO TC 1110
1150
1154 DO 1170 J = 1, JM
POT(I, J, NF) = POT(I, J, NF) - DATA(J) / CY
POT(I+IMD2, JM-J+2, NF) = POT(I+IMD2, JM-J+2, NF) - DATA(J+JM) / CY
CONTINUE
1170
1110 IF (NF.GT. 0)
S1 = 0.
S2 = 0.
DO 1160 I = 1, IM
S1 = S1 + SLP(I, JNP)
S2 = S2 + SLP(I, JNP)
CONTINUE
1160 S1 = S1 / IM
S2 = S2 / IM
DO 1165 I = 1, IM
SLP(I, I) = S1
SLP(I, JNP) = S2
CONTINUE
1165
1164 S1 = 0.
S2 = 0.
DO 1173 I = 1, IM
S1 = S1 + POT(I, J, NF)
S2 = S2 + POT(I, JNP, NF)
CONTINUE
1173 S1 = S1 / IM
S2 = S2 / IM
DO 1175 I = 1, IM
POT(I, I, NF) = S1
POT(I, JNP, NF) = S2
CONTINUE
1175
1180 IF (NF.LT. NLAY)
CONTINUE
GO TO 1000
C
C * * * SECOND MAIN J-LOOP
C * * * TRANSFORM BACK TO SURFACE PRESSURE
C
MAXIT = 50
DO 2000 J = 1, JNP
DO 1320 I = 1, IM
PHSX = PHIS(I, J)
PSFC = SLP(I, J) - PTOP
DO 1330 NITX = 1, MAXIT
PI = PSFC
T8 = POT(I, J, 8) * EXPBYK(P5FC * SIG(8) + PTOP)
T9 = POT(I, J, 9) * EXPBYK(P5FC * SIG(9) + PTOP)
TSURF = T10(P5FC, T8, T9)
PSFC = SLP(I, J) / SLEXP(PHISX, TSURF) - PTOP
IF (ABS(P5FC-PI)/1000. .LE. 1.E-5) GO TC 1340
PRINT 1335, MAXIT
FORMAT(IX, ' MORE THAN ', IS, ' ITERATIONS.', I)
PRINT 7990, J
PRINT 7991, I
7990 FORMAT(IX, ' J = ', I5)
7991 FORMAT(IX, ' I = ', I5)
1340 SLP(I, J) = PSFC
IF (J.EQ.1 .OR. J.EQ.JNP) GO TC 1350
CONTINUE
1350 DO 1360 I = 2, IM
CONTINUE
GO TC 1310

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1018 ISN 0108
1019 ISN 0109
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1021 ISN 0111
1022 ISN 0112
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1024 ISN 0114
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1100 ISN 0190

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ISN 0191
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1360 SLP(I,J) = SLP(I,J) CONTINUE  
1370 CONTINUE  
C * * * TRANSFORM BACK TO TEMPERATURE  
C  
DO 1400 L = 1,NLAY  
DO 1420 I = 1,IM  
PKAPA = EXPBKISLP(I,J) * SIGIL) + PTC(I)  
POT(I,J,L) = POT(I,J,L) * PKAPA  
IF (J.EQ.1 .OR. J.EQ.JNP) GO TO 1430  
CONTINUE  
GO TO 1400  
1420 DC 1440 I = 2,IM  
1430 POT(I,J,L) = POT(I,J,L) CONTINUE  
1440 CONTINUE  
1400  
C * * * RETURN SMOOTHED VALUES  
C  
DO 1520 L = 1,NLAY  
DO 1520 I = 1,IM  
T(I,L,J) = POT(I,J,L)  
1520 P(I,J) = SLP(I,J) CONTINUE  
1530 CONTINUE  
2000  
DO 1550 L = 1,NLAY  
TPOL(L,1) = POT(1,1,L)  
TPOL(L,J2) = POT(1,JNP,L) CONTINUE  
1550 PPOL(1) = SLP(1,1)  
PPOL(2) = SLP(1,JNP)  
C * * *  
12714 RUN HERE TILL RETURN  
IF(SATSNX.EQ.0)RETURN  
DO 500 N=1,4  
DO 500 K=1,NLAY  
DO 500 I=1,IM  
GT(I,K,N,J)=P(I,J)*DXYP(J)*Q(I,K,N,J)  
500 CONTINUE  
DO 501 J=2,JM  
DC 501 I=1,IM  
PT(I,J)=G(I,J)  
501 CONTINUE  
DO 502 L=1,NLAY  
DO 502 M=1,4  
DO 502 N=1,2  
QPOL(L,N,M)=PPOL(M)*QPOL(L,N,M)  
502 CONTINUE  
PTPOL(1)=PPOL(1)  
PTPOL(2)=PPOL(2)  
599 FORMAT(1X,' NEW SMOOTHER ENTERED. TAU=',F12.3)  
RETURN  
END
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04M23000
04M23010
04M23020
04M23030
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