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	(E82-10359) COMPARISON OF CHANGES OF GEOMAGNETIC FIE MAGSAT ALTITUDES (TRODUCTO	STCRM-TIME LC AT GROUND AND	N82-26757
	Espaciais, Sao Jose) 17 p	o de Pesquisas HC A02/MF A01 CSCL 04B	Unclas G3/43 00359
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1. Publication N9 INPE-2297-RPE/401	2. Version	3. Date Jan., 1982	5. Distribution			
4. Origin P	rogram					
6. Key words - selected GEOMAGNETIC VARIATI STORM-TIME CHANGES	by the author	r(s)	L			
7. U.D.C.: 550.38	an a					
8. Title	INPE-	-2297-RPE/401	10. N9 of pages: 16			
FIELD AT GROUND AND MAG	ieona givelti c	11. Last page: 15				
			12. Revised by			
<b>9. Authorship</b> Rajaram Nalin Ba	Purushottam Ka bulal Trivedi	me	Haw tau Ivan Jelinek Kantor			
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1st MAGSAT Progress Report for Project M55 - December 1981

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COMPARISON OF STORM-TIME CHANGES OF GEOMAGNETIC FIELD AT GROUND AND AT MAGSAT ALTITUDES

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#### 1. INTRODUCTION

The purpose of this investigation is to compare the storm-time variations of geomagnetic field at ground and at MAGSAT altitudes.

#### 2. TECHNIQUES

Most of the ground data are <u>not yet available to us</u>. We have asked for these from the WDC-A, Boulder, Colorado. However, we know that these will be for the H, D, Z components of geomagnetic field and, in particular, we wish to concentrate on the H component.

From the MAGSAT tapes, data are available for the X, Y, Z components. We converted these (by computer program) to yield the component H =  $(X^2 + Y^2)^{1/2}$ .

MAGSAT goes round the earth in a roughly polar orbit, needing about 1.6 hours. The plane of the orbit seems to be roughly perpendicular to the sun-earth line so that each pass has a south-north swing during local dusk hours and a subsequent north-south swing during local dawn hours. Thus, these two swings have equatorial crossings roughy 0.8 hours apart and at roughly diametrically opposite longitudes. After every complete pass, there is a longitude shift of I

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about 24 degrees (westwards) and, in about 15 complete passes, 24 hours and all longitudes are covered.

Fig. 1 shows the  $\overline{H}$  values for two successive halves of the same pass (379) for the low latitude region (equator to  $\pm$  30<sup>0</sup>) geographically. During the pass, the satellite seems to change its altitude by a few tens of kilometers. We assumed that H is inversely proportional to  $R^3$ , where R = distance of the satellite from the center of the earth, and <u>normalised</u> all H values to a fixed  $R_0 = 6800$ km. What are shown in Fig. 1 are  $\overline{H}$  values i.e. values of H normalised to a constant geocentric distance  $R_0 = 6800$  km. One can see that the latitudinal gradients of H are very large indeed (more than 150 gamma per degree) except near the  $\overline{H}_{max}$ . However  $\overline{H}_{max}$  itself has a very large longitudinal variation. For example, it has values as high as 33000 gamma at longitudes of about  $+100^{\circ}$ , dropping to as low as 23000 gamma at longitudes of about - 70°, a drop of 10000 gamma in 170° i.e. about 60 gamma per degree of longitude. In storm-time studies, the effects we expect are of the order of few tens of gamma. It is obvious therefore, that values of H or  ${\rm H}_{\rm max}$  cannot be used directly in such a study, lest even slight errors in latitudes or longitudes could give errors overwhelming the expected effect.

A reasonable alternative is to subtract out some base values which have already taken into account the gross features of the latitude and longitude variation of H. Such base values are obtained from the geomagnetic field model. The MAGSAT tapes we have contain model values of X, Y, Zviz. XMD, YMD, ZMD. We used these to obtain HMD =  $(XMD^2 + YMD^2)^{1/2}$  and obtained HMD which is HMD normalized to a constant  $R_0 = 6800$  km as before. If the model is adequate, the difference  $\Delta H = \overline{H} - \overline{HMD}$  should be zero. In Fig. 1, we show the latitude variation of  $\Delta H$  for the two swings (N + S) and (S + N) of pass 379. In contrast to the sharp gradients in  $\overline{H}$ , the variation in  $\Delta H$  is very small (few tens of gamma over a large latitude range). This is true for widely differing longitudes. Hence, the major gradients, both latitudinal as well as longitudinal, are taken care of by the model subtraction.

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However, the fact remains that  $\Delta H$  so obtained is not zero. We now examine the utility of  $\Delta H$  for storm-time studies.

#### 3. ACCOMPLISHMENTS

Contraction of the local distribution of the

We obtained the latitudinal variation of  $\Delta H (+30^{\circ} \text{ to } -30^{\circ} \text{ latitude})$  for the two swings of every pass, as shown in Fig. 1. To start with, we concentrated on  $\Delta H_0$  i.e. the values of  $\Delta H$  at the <u>equatorial</u> crossing (geographical latitude zero). Table 1 shows a sample listing for passes Nos 154-169 on Nov. 12, 1979 (Julian day 44189). From the MAGSAT tape, values are available for Mean Local Time (MLT) as well as geographic longitude. The UT in Table 1 is calculated as:

UT = MLT - (Longitude in degrees)/15.

for values at equatorial crossing only.

Fig. 2 is a plot of  $\Delta H_0$  at DAWN (dots) and DUSK (crosses) for six days Nov. 2-7 in the top row, Nov. 8-13 in the middle row and Nov. 14-19 in the botton row. Standard Dst (Sugiura and Poros, 1971 and similar furthur publications) is also p'otted and K<sub>p</sub> histograms are marked. For each day, there are about 15 dots of  $\Delta H_0$  (Dawn) and 15 crosses of  $\Delta H_0$  (Dusk). Judging from Dst, the period contains quiet as well as disturbed days. Thus, Nov. 5, 6, 35 were very quiet (highest 3 hourly  $K_p$  was 1<sup>+</sup>). In contrast, Nov. 13-14 had a geomagnetic storm with Kp exceeding 5 and Dst reaching - 90. It is interesting to note that  $K_p$  does not always match Dst. Fig. 3 shows a plot of  $K_p$ versus Dst for Nov. 2-19, 1979. Whereas very high  $K_p$  are associated with high Dst, the scatter is very large, mainly because of the glaring mismatch on Nov. 15 when K<sub>D</sub> was very low but Dst was still large (about - 30) in the storm recovery and on Nov. 11 when  $K_p$  was large while Dst was very small or positive. Dst is obvicusly a finer index of storm-time variations.

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## 4. SIGNIFICANT RESULTS

The following is noteworthy in Fig. 2:

- (i) Both the  $\Delta H_0$  (Dawn) and  $\Delta H_0$  (Dusk) seem to follow the Dst trend. Thus, these H residuals are representative of storm--time variations, at least qualitatively.
- (ii) In general,  $\Delta H_0$  is non-zero and is negative. The  $\Delta H_0$  (Dusk) (crosses) are more negative. The non-zero values could be an indication of the inadequacy of the model. However, on Nov. 11 when Dst attained positive values, both the  $\Delta H_0$  at Dusk and Dawn became almost zero. This leads us to believe that the non-zero values of  $\Delta H_0$  are indicative mainly of storm-time activity, in smaller or larger degrees.
- (iii) On Nov. 13-14,  $\Delta H_0$  (Dusk) are numerically very much larger than  $\Delta H_0$  (Dawn). Thus, the storm effect is seen more effectively in the Dusk sector. However, on Nov. 14 at about 0600 when the storm was still recovering (Dst about - 60), the Dawn and Dusk  $\Delta H_0$  merge into each other, each having a value of about - 60.
- (iv) Conventionally, the storm-time variation is considered to be composed of an isotropic component Dst and a LT dependent component DS. For the MAGSAT data, an <u>average</u> of  $\Delta H_0$  (Dawn) and  $\Delta H_0$  (Dusk) should be roughly equivalent to Dst while their difference should be equivalent to DS. However, as seen from Table 1, the  $\Delta H_0$  at Dusk and Dawn are not recorded simultaneously (at the same UT) but are recorded about 0.8 hours apart. As a rough approximation, in-between values could be obtained as averages of the previous and succeeding values, separately for  $\Delta H_0$  (Dawn) and  $\Delta H_0$  (Dusk). For example in Table 1,  $\Delta H_0$  (Dusk) at UT = 8.4 and UT = 10.0 are - 42 and - 40 respectively. Hence  $\Delta H_0$  (Dusk), at UT = 9.2 could be

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assigned as - 41. Similarly, for UT = 2.4 and UT = 4.1, the  $\Delta H_0$  (Dawn) are - 18 and - 30 respectively. The in-between value for UT = 3.3 may be assigned as - 24. The complete set with <u>manipulated</u> values so obtained are given in Table 1. Also the average Dst' =  $[\Delta H_0$  (Dusk) +  $\Delta H_0$  (Dawn) ]/2 and the difference DS' =  $[\Delta H_0$  (Dusk) -  $\Delta H_0$  (Dawn) ] are given.

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Fig. 4 shows a plot of Dst versus  $\Delta H_0$  (Dusk) in the upper half, for the four days Nov. 11-15, 1979. Values during the recovery of the storm (Nov. 14-15) are shown as triangles. Whereas the dots fall roughly on one straight line, the triangles seem to lie above this line, indicating lesser  $\Delta H_0$  (Dusk) when Dst is still high. In the lower half of Fig. 4, a similar plot is shown for Dst versus  $\Delta H_0$  (Dawn). The slope of the regression line is <u>lesser</u>, indicating smaller values of Dawn  $\Delta H_0$  compared to  $\Delta H_0$  (Dusk) for similar Dst.

Fig. 5 upper half shows a plot of Dst versus the average  $Dst' = [\Delta H_0 (Dusk) + \Delta H_0 (Dawn)]/2$ . The scatter about the regression line is much smaller as compared to that in Fig. 4 (upper half), indicating that Dst is more similar to this average Dst' than to either  $\Delta H_0$  (Dusk) or  $\Delta H_0$  (Dawn). The lower half of Fig. 5 shows a plot of Dst versus the difference DS' =  $[\Delta H_0 (Dusk) - \Delta H_0 (Dawn)]$ . Here, the scatter is very large. In particular, the DS' is <u>very near</u> <u>zero</u> for a large range of Dst, during the recovery phase (triangles) of the storm. This implies that the LT dependent DS' component vanishes much before the complete recovery of a storm.

(v) If the  $\Delta H_0$  (Dusk) and  $\Delta H_0$  (Dawn) plotted in Fig. 2 are genuine, the implication is that both Dst' and DS' are largely of magnetospheric origin and, Dst' lasts longer than DS'. Also, DS' is composed mainly of larger depressions from zero level of the Dusk values and the DS' values are not well correlated with Dst' values. Since the DS is attributed to partial ring currents associated with field-aligned currents passing through the auroral ionosphere, the lack of

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correlation probably indicates partially independent evolutions of the main equatorial ring currents and its leakages to the auroral regions.

(vi) If the  $\Delta H_{\Omega}$  values have some non-physical origins, corrections for the same may be needed. If the model values HMD are wrong, one will have to await for better estimates. On the other hand, the model does not take care of local anomalies. Since the  $\Delta H_0$  shown in Fig. 2 are  $\Delta H$  values at equatorial crossings, local anomalies at certain longitudes may affect the  $\Delta H_0$  values. For example, Regan et al (1975), indicate the Bangui or Central African anomaly at about + 15<sup>0</sup> longitude. Fig. 6 shows the latitude variation of  $\Delta H$  for a few passes in this region. Whereas a clear depression (anomaly) is seen at about + 6° latitude, the value of  $\Delta H_0$  (at latitude 0°) is - 16, - 30 and - 31 for the passes 264, 49, 380, all of which occurred at very low  $K_{\mbox{\scriptsize D}}$  and Dst. In Fig. 2, successive dots for several passes are consistently above the successive crosses, thus indicating that local anomalies are probably not playing any important role. Nevertheless, a quantitative estimate of  $\Delta H_0$  does need a proper correction for local anomalies, if any. To estimate the nature of these corrections, all passes at given longitudes (or at least longitude zones) will have to be examined for quiet periods and AH versus latitude patterns will have to be established for quiet periods, to be later subtracted from similar patterns for individual passes for disturbed days. Work in this direction is in progress.

In due course,  $\Delta H_0$  values so corrected will be compared with ground data, when these are available, and results will be reported in future reports. We are also keeping in mind the possibility that  $\Delta H_0$  (Dawn) and  $\Delta H_0$  (Dusk) may differ because of Sq effects (Sugiura and Hagan = 1979).

## 5. DATA QUALITY

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From about the first 400 passes that we have examined so far (Nov. 2-27, 1979), a majority shows a reasonable  $\Delta H$  versus latitude variation as shown in Fig. 1. However, a few passes, notably passes Nos., 1-15 on Nov. 2, 1979 (Julian day 44179) show very odd patterns of  $\Delta H$  versus latitude, with values ranging and oscillating from + 100 to - 100 or more. In Fig. 2, Nov. 2 shows very large  $\Delta H_0$  (Dusk), unwarranted by the low values of Dst. A similar discrepancy occurs on Nov. 16. We believe that these passes are not reliable. This may kindly be reexamined and, if true, the MAGSAT investigators be informed corrections, if any.

The quality of the rest of the data seems to be very good.

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Reagan R.D., Cain J.C. and Davis W.M. A global magnetic anomaly map. J. Geophys. Res. 80, 794-802 (1975).

Sugiura M. and Poros D.J. Hourly values of equatorial Dst for the years 1957 to 1970. GSFC Publication X-645-71-278 (1971).

Sugiura M. and Hagan M.P. Geomagnetic Sq variation at satellite altitudes: Is Sq correction important in MAGSAT data analysis? <u>Geophys. Res. Letters 6</u>, 397-400 (1979). 3

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TABLE	1.	UT, Longitude and AH at equatorial crossing, as well as other	A.
		parameters for passes 154-169 on Nov. 12, 1979 (Julian day	
		44189)	

Dage MA		DUSK	DUSK swing DAWN swi		swing			Complete set with Hantpulated values			Average	Difference
rass ny							DUSK 1	wing	UAWN	swing	DUSK Allo	Minus
		Long,	LHO	Long,	AHO		Long,	Allo	Long	allo	DAWN AHO	DAWN AHO
154 154	0.0 0.8	- 96 <sup>0</sup> -	- 27	+ 730	- 15	+ 1 + 1	- 96 <sup>0</sup> - 108 <sup>0</sup>	- 27 - 29	+ 736	- 15	- 22	- 14
165 165	1.6 2.4	- 119º	- 30	+ 490	- 18	+ 1	- 119º - 132º	- 30 - 29	+ 61 <sup>0</sup> + 49 <sup>0</sup>	- 17	- 24 - 24	- 13 - 11
156 156	3.3 4,1	- 1430	- 27	+ 260	- 30	+ 2	- 143 <sup>0</sup> - 155 <sup>0</sup>	- 27 - 26	+ 38 <sup>0</sup> + 26 <sup>0</sup>	- 24	- 26 - 28	- 3 + 4
157 157	5.0 5.9	- 166 <sup>0</sup> -	- 24	+ 30	- 21	+'9	- 166 <sup>0</sup> - 178 <sup>0</sup>	- 24 - 24	+ 15 <sup>0</sup> + 3 <sup>0</sup>	- 26 - 21	- 25 - 23	+ 2
168 158	6,8 7,6	+ 171°	- 23	- - 21 <sup>0</sup>	- 24	0 + 1	+ 171 <sup>0</sup> + 159 <sup>0</sup>	- 23 - 33	+ 9° - 21°	- 23 - 24	- 23 - 29	0 - 9
159 169	8,4 9,2	+ 147 <sup>0</sup>	- 42 -	- - 44 <sup>0</sup>	- 8	0 0	+ 147 <sup>0</sup> + 136 <sup>0</sup>	- 42 - 41	- 33 <sup>0</sup> - 44 <sup>0</sup>	- 16 - 8	- 29 - 25	- 26 - 33
160 160	10.0 10.7	+ ',24 <sup>0</sup> -	- 40 -	- 68 <sup>0</sup>	- 25	0 + 13	+ 124 <sup>0</sup> + 112 <sup>0</sup>	- 40 - 39	- 56 <sup>0</sup> - 68 <sup>0</sup>	- 17 - 25	- 29 - 32	- 23 - 14
161 161	11.4 12.4	+ 100 <sup>0</sup> -	- 37 -	- 910	- - 23	+ 5 + 6	+ 100 <sup>0</sup> + 89 <sup>0</sup>	- 37 - 41	- 80° - 91°	- 24 - 23	- 31 - 32	- 13 - 18
162 162	12,9 13,9	+ 77 <sup>0</sup> -	- 44 -	- - 115 <sup>0</sup>	- 26	+ 6 + 1	+ 77 <sup>0</sup> + 65 <sup>0</sup>	- 44 - 41	- 103 <sup>0</sup> - 115 <sup>0</sup>	- 25 - 26	- 35 - 34	- 19 - 15
163 163	14.4 15.2	+ 53 <sup>0</sup> -	- 37 -	- 138 <sup>0</sup>	- - 24	- 1 - 3	+ 53 <sup>0</sup> + 42 <sup>0</sup>	- 37 - 39	- 127 <sup>0</sup> - 138 <sup>0</sup>	- 25 - 24	- 31 - 32	- 12 - 15
164 164	15,9 16,9	+ 30 <sup>0</sup> -	- 40	- 161 <sup>0</sup>	- - 25	- 3 - 3	+ 30 <sup>0</sup> + 19 <sup>0</sup>	- 40 - 46	- 150 <sup>0</sup> - 161 <sup>0</sup>	- 25 - 25	- 33 - 36	- 15 - 21
165 165	17.5 18.2	+ 7 <sup>0</sup> -	- 51 -	- + 175 <sup>0</sup>	- - 31	0 + 4	+ 7 <sup>0</sup> - 5 <sup>0</sup>	- 51 - 39	- 173 <sup>0</sup> + 175 <sup>0</sup>	- 28 - 31	- 40 - 35	- 23 - 8
166 166	19,0 19,7	+ 170 -	- 26	+ 1520	- - 21	+4 +4	- 17 <sup>0</sup> - 29 <sup>0</sup>	- 26 - 24	+ 164 <sup>0</sup> + 152 <sup>0</sup>	- 26 - 21	- 26 - 23	0 - 3
167 167	20.5 21.2	- 400	- 22	- + 128 <sup>0</sup>	- - 29	1 - 9	- 40 <sup>0</sup> - 52 <sup>0</sup>	- 22 - 25	+ 140 <sup>0</sup> + 128 <sup>0</sup>	- 25 - 29	- 24 - 27	+ 3 + 4
168 168	21 .9 22 . 7	- 64 <sup>0</sup> -	- 27 -	+ 105°	- - 25	- 9 - 9	- 64 <sup>0</sup> - 76 <sup>0</sup>	- 27 - 28	+ 117 <sup>0</sup> + 105 <sup>0</sup>	- 27 - 25	- 27 - 27	0-3
169	23.4	- 87 <sup>0</sup>	- 29	•	•	- 6	- 87 <sup>0</sup>	- 29				-

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Fig. 1 - The latitude variation of  $\overline{H}$  (the H component normalized to a geocentric distance of 6800 km) for the Dawn and Dusk swings of Pass 379 as also of  $\Delta H = \overline{H} - \overline{HMD}$ , where  $\overline{HMD}$  is the normalised value, predicted by geomagnetic model.

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Fig. 2 - Plot of AH<sub>O</sub> (the AH values at equatorial crossings) for Dusk (crosses) and Dawn (dots) swings, as also of Dst for Nov. 2-7 (top), Nov. 8-13 (middle) and Nov. 14-19, 1979 (bottom). Kp is also indicated as histograms.

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Fig. 3 -  $K_p$  versus Dst for Nov. 2-19, 1979.

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Fig. 4 - Dst versus  $\Delta H_0$  (Dusk) (upper half) and  $\Delta H_0$  (Dawn) (lower half) for the storm period Nov. 11-15, 1979. Triangles represent values in the recovery phase (Nov. 14-15).

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Fig. 5 - Dst versus Dst' =  $[\Delta H_0 (Dusk) + \Delta H_0 (Dawn)]/2$  in the upper half and versus DS' =  $[\Delta H_0 (Dusk) - \Delta H_0 (Dawn)]$ in the lower half.

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Fig. 6 - The latitude distribution of  $\Delta H$  for a few passes at longitudes + 13° to + 16° in the Central African region. The Dst and Kp values are indicated and show quiet conditions for Passes 264, 49, 380 and disturbance for Pass 341.

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