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Spectral Analyses of Satellite Geopotential Missions

Abstract

A new, geometrical, first order, non-non-non-time plane of consta been developed based on Orlov's uniforming rotation of constants plane of constants plane of constants plane o $\frac{100}{100}$ relation spectra $\frac{100}{100}$ $\frac{100}{$ $\frac{1}{2}$ OSU86F are shown for the indicated 1984 $\frac{1}{2}$ and $\frac{1}{2}$ is a pair of a pa TRANSIT satellites at 400km altitude with a 93?5 inclination.

Introduction

To perform the integration for a geodetic satellite a new form has been developed for the geopotential on an orbit as

$$
V = \frac{GM}{a} \sum_{k=0}^{\infty} \xi^k \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} e^{i(pu + q\vartheta)} [\overline{R}_{p,q}^{(k)} + \varphi \overline{T}_{p,q}^{(k)}],
$$

where the symbols are defined in $\frac{1}{2}$ and $\frac{1}{$ derived for the generation of the complex coefficients from a spherical harmonic geopotential model.

The geopotential variation

$$
a \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \overline{R}^{(0)} e^{i(pu+q\vartheta)}, \qquad (1)
$$

is the distance of an equiporter of θ and surface from sphere of radius(0.24R)^o θ . The spectrum in Fig. I is a 16 bar graph of the amplitudes 21 log-log bar 10 P, 10 m the frequency is $|p_1 + q_2|$. The along-track deflection evaluation nominal circular orbit is

$$
\frac{1}{a} \frac{\partial V}{\partial u} = \frac{GM}{a^2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} i p \overline{R}_{p,q}^{(0)} e^{i(pu+q\vartheta)}, \qquad (2)
$$

and its spectrum is plotted in Fig. 2.

Position Perturbations

The real forms of the series $\frac{1000L}{M}$ that the formulas are 1988a], but it is apparent from [Melvin, 1988b] whit porturbations are simpler with complex coefficients. In fact the orbit perturbations are of the same form as the disturbing potential

$$
\varphi = \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \varphi_{p,q} e^{i(pu+q\vartheta)}, \qquad \eta = \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \eta_{p,q} e^{i(pu+q\vartheta)}
$$

$$
\xi = \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \xi_{p,q} e^{i(pu+q\vartheta)}, \qquad \chi = \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \chi_{p,q} e^{i(pu+q\vartheta)}
$$

where the coefficients are computed from the algorithm

$$
\varphi_{p,q} = -\frac{n_0^2 \bar{T}_{p,q}^{(0)}}{(pn+q_0^2)^2 - N^2}, \quad (3), \qquad \eta_{p,q} = \frac{n_0^2 p \bar{R}_{p,q}^{(0)}}{\nu (pn+q_0^3)},
$$

$$
\xi_{p,q} = -\frac{n_0^2 \bar{R}_{p,q}^{(1)} + 2\nu^2 \eta_{p,q}}{(pn+q_0^3)^2 - N^2}, \quad (5), \qquad x_{p,q} = -i\nu \frac{\eta_{p,q} - 2\xi_{p,q}}{pn + q_0^3}.
$$
 (7)

After multiplication by 2a, the amplitudes of the coefficients of (3), \langle 3), and (*I*) are plotted in Figs. 3, $\overline{5}$, and $\overline{7}$ from which it is seen that position perturbations accentuate the low and near orbital frequencies and attenuate the high frequencies. Expressions for the cross-track (3), radial (5), and along-track (7) position perturbations from the Kaula-Allan theory are found in [Rosborough and Tapley, 1987].

Velocity Perturbations

By use of coordinates at the nominal satellite position, the cros: track, radial, and along-track velocity components are

$$
v_{\varphi} = a(\dot{\varphi} - \hat{\Omega} \sin \iota \cos u) = a \left[\sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} i(pn+q\dot{\vartheta}) \varphi_{p,q} e^{i(pu+q\theta)} - \hat{\Omega} \sin \iota \cos u \right], (4)
$$

$$
v_r = a\xi = a \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} i(pn + q\dot{\theta})\xi_{p,q} e^{i(pu+q\theta)}, \qquad (6)
$$

$$
v_{u} = av + a(\chi + \nu \xi) = av[1 + \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} (\eta_{p,q} - \xi_{p,q}) e^{i(pu+q\theta)}] . (8)
$$

The amplitudes of the coefficients of (4), (6) and (8) are plotted in Figs. 4, 6, and 8. By comparison of the high frequency fall-off in the position and velocity spectra, it is clear why more geopotential information is obtained from Doppler beacon satellites than from skin tracking even if it is by laser reflectors.

Intersatellite Measurements

For a close satellite pair in same orbit, the intersatellite range is

$$
\rho = r\Delta u = a\delta u(1 + \xi + \frac{\partial x}{\partial u}) = a\delta u[1 + \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} (\xi_{p,q} + ip\chi_{p,q})e^{i(pu+q\vartheta)}],
$$
(9)

A time derivative yields the intersatellite range rate as

$$
\dot{\rho} = a\delta u \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} i(pn + q\dot{\vartheta}) (\xi_{p,q} + ip\chi_{p,q}) e^{i(pu+q\vartheta)} . \qquad (10)
$$

For a nominal separation of $a\delta u=100$ km the spectra of (9) and (10) are plotted in Figs. 9 and i0. A comparison of the high frequency portions of Figs. 2 and i0 and the foregoing formulas show

$$
\dot{\rho} \simeq \frac{a\delta u}{a^2 n} \frac{\partial V}{\partial u} ,
$$

from which the result of [Comfort, 1973] is modified to state that at high frequencies for a pair of satellites flying in formation the intersatellite

range rate mimics the along track deflection. The paucity of spectral
lines near twice per orbit in Fig. 10 indicates that complete geopotential
recovery is not possible with only intercatellite measures The paucity of spectral \sim recovery is not possible with only intersatellite range rate.

Gravity Gradient

Equally simple formulas for the gravity gradient tensor along a
nominal circular orbit in Orlov's plane are found in [Melvin, 1988b] as

$$
\Gamma_{r,r} = 2n_{o}^{2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} \overline{R}_{p,q}^{(2)} e^{i(pu+q\vartheta)} (11), \ \Gamma_{r,u} = in_{o}^{2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} p(\overline{R}_{p,q}^{(1)} - R_{p,q}^{(0)}) e^{i(pu+q\vartheta)} (12)
$$

$$
\Gamma_{r,\varphi} = n_{o}^{2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} (\bar{T}_{p,q}^{(1)} - \bar{T}_{p,q}^{(0)}) e^{i(pu+q\vartheta)} \Gamma_{u,u} = n_{o}^{2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} (-p^{2} \bar{R}_{p,q}^{(0)} + \bar{R}_{p,q}^{(1)}) e^{i(pu+q\vartheta)} \tag{14}
$$

 \sim

$$
\Gamma_{u,\varphi} = in_{\mathbf{0}}^{2} \sum_{p=-\infty}^{\infty} \sum_{q=-\infty}^{\infty} p \overline{T}_{p,q}^{(0)} e^{i(pu+q\vartheta)}, \quad \Gamma_{\varphi,\varphi} = -\Gamma_{r,r} - \Gamma_{u,u} \tag{16}
$$

Figs. 11 through 16. \mathbf{m}

Geopotential Recovery

Although it is subjective and dependent on the sensitivity of the
measuring devices, it became apparent in the generation of the spectrum in
Fig. 10 that the claim of 100th order recovery from SAGE, [Pisacane, et al., 1984], could not be substantiated. From intersatellite range rate, μ appears devices, it is developed appears in the generation of the spectrum in the spectr $\frac{f(t)}{t}$ in the called $\frac{f(t)}{t}$ of $\frac{f(t)}{t}$ order recovery from SAGE, $\frac{f(t)}{t}$ in S information beyond 90th order in the gravity gradient components.

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te Pair

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