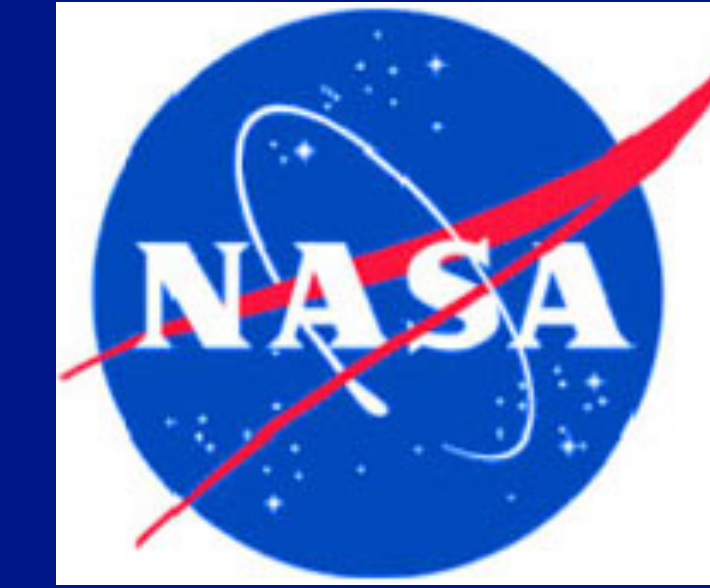




# Evaluating Effects of H<sub>2</sub>O and Overhead O<sub>3</sub> on Global Mean Tropospheric OH Concentration



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## 1. Scientific motivation for studying tropospheric OH

### Sensitivity of OH<sup>TROP</sup> to climate change

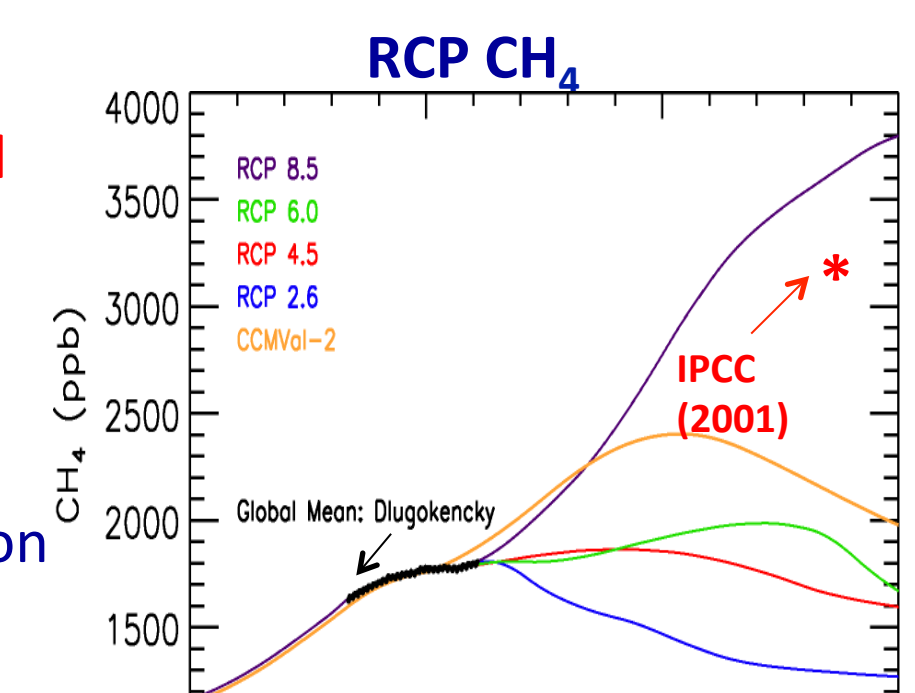
- IPCC (2001) predicted future OH would fall due to rising CH<sub>4</sub>
- Best knowledge of global tropospheric OH (OH<sup>TROP</sup>) comes from analysis of CH<sub>3</sub>CCl<sub>3</sub> observations
- The CH<sub>4</sub> lifetime inferred from OH<sup>TROP</sup> based on CH<sub>3</sub>CCl<sub>3</sub> is 8.9 years; yet many sources use a CH<sub>4</sub> lifetime of ~12 years, which was found by modeling studies described in IPCC (2001)
- The 12 year CH<sub>4</sub> lifetime is central to subsequent IPCC reports and is called the "perturbation lifetime"
- Our work is motivated by understanding:
  - what factors other than rising CH<sub>4</sub> will affect OH<sup>TROP</sup>
  - veracity of notion that future OH<sup>TROP</sup> will decline

### Importance of OH<sup>TROP</sup>

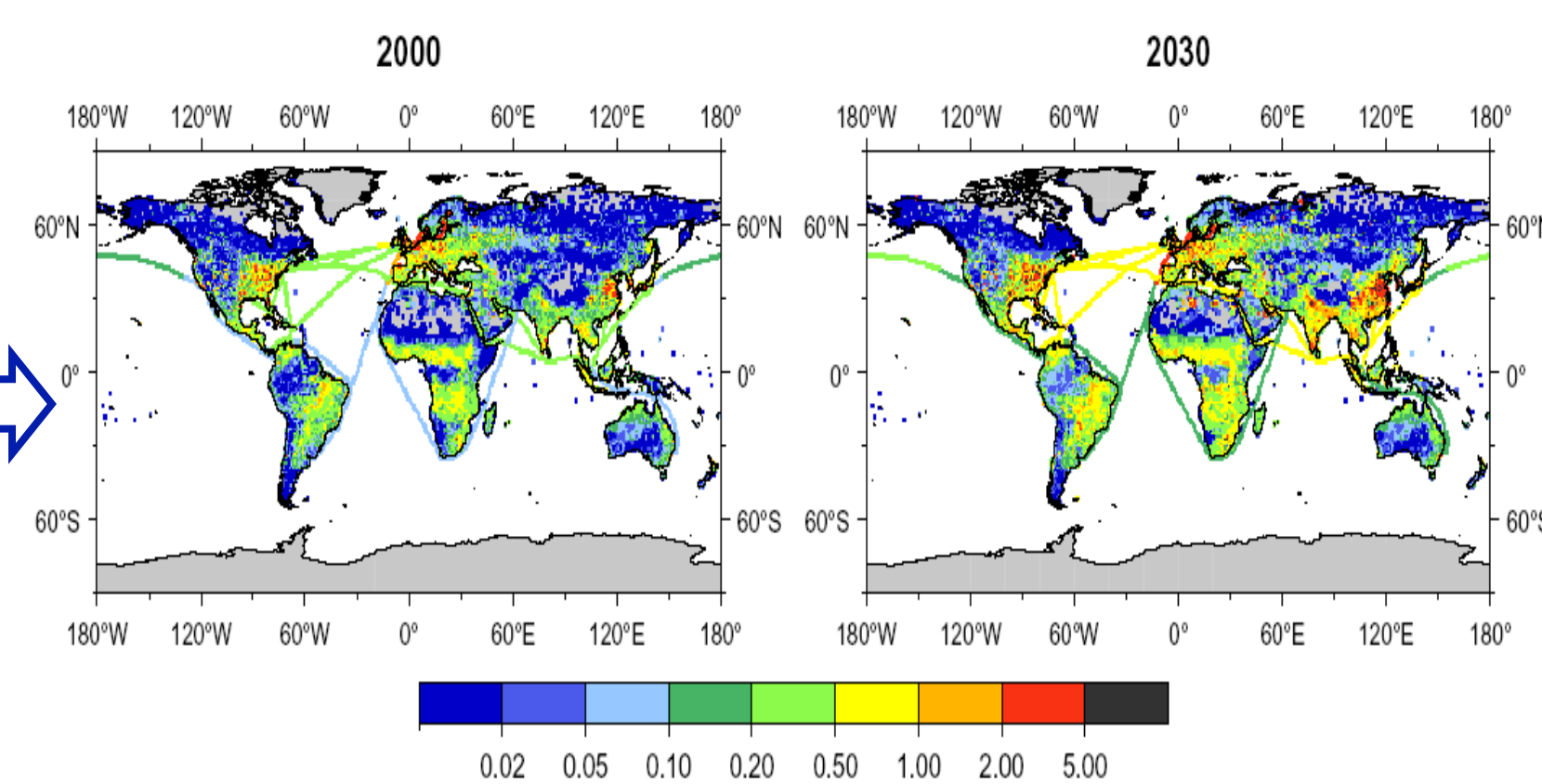
- OH serves as the main sink for many species: CO, CH<sub>4</sub>, CH<sub>3</sub>CCl<sub>3</sub>, all HCFCs and HFCs, biogenic VOCs (isoprene and terpene), anthropogenic VOCs (formaldehyde, benzene, toluene, etc.)
- HO<sub>x</sub> is integral to O<sub>3</sub> formation and destruction pathways; alone, HO<sub>x</sub> depletes O<sub>3</sub>, but with NO<sub>x</sub> and VOCs, HO<sub>x</sub> creates O<sub>3</sub>

### Methane Oxidation and OH

- Under low NO<sub>x</sub>, the oxidation of CH<sub>4</sub> destroys HO<sub>x</sub>
- Under high NO<sub>x</sub>, the oxidation of CH<sub>4</sub> produces HO<sub>x</sub>
- **Loss of methylperoxy radical CH<sub>3</sub>O<sub>2</sub> with NO or HO<sub>2</sub> is critical**
- Future levels of CH<sub>4</sub> and NO<sub>x</sub> are highly uncertain
  - CH<sub>4</sub> between now and 2100 varies dramatically among RCP scenarios
  - NO<sub>x</sub> between now and 2100 will depend on whether the developing world implements selective catalytic reduction on coal power plants and catalytic converters on cars
- Future OH<sup>TROP</sup> will also depend on overhead column O<sub>3</sub>, local humidity, and biogenic emission of VOCs, etc.
- **Given this complexity, we expect CCMs to project a wide range of values for OH<sup>TROP</sup> during the rest of this century**



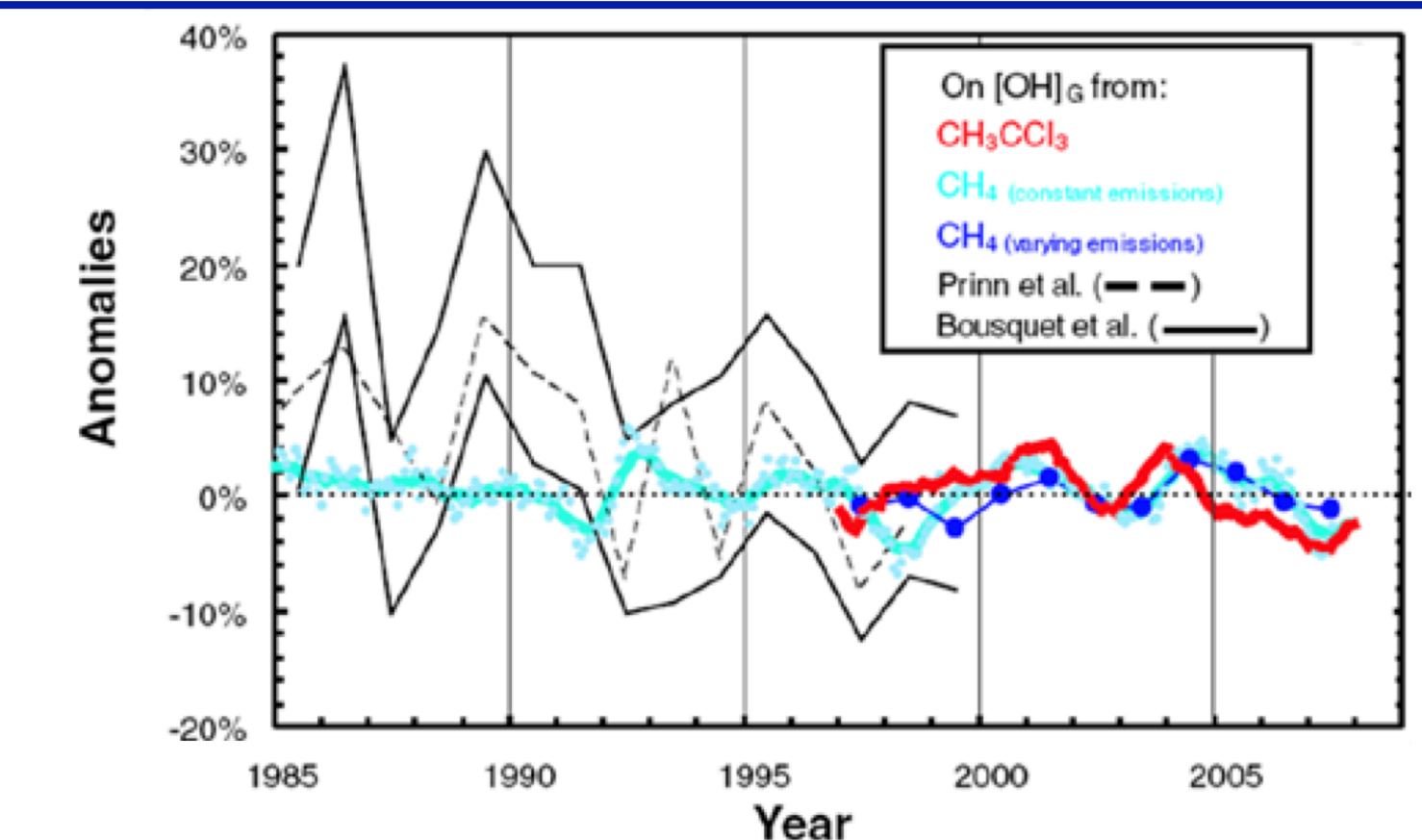
NO<sub>x</sub> emissions from industry, power generation, traffic, domestic heating, and biomass burning used as input for prior CCM calculations, for years 2000 (38.0 Tg N / yr total) and 2030 (67.6 Tg N / yr total). From Eyring et al., ACP, 2007.



## 3. Preliminary results for expected changes in OH<sup>TROP</sup>

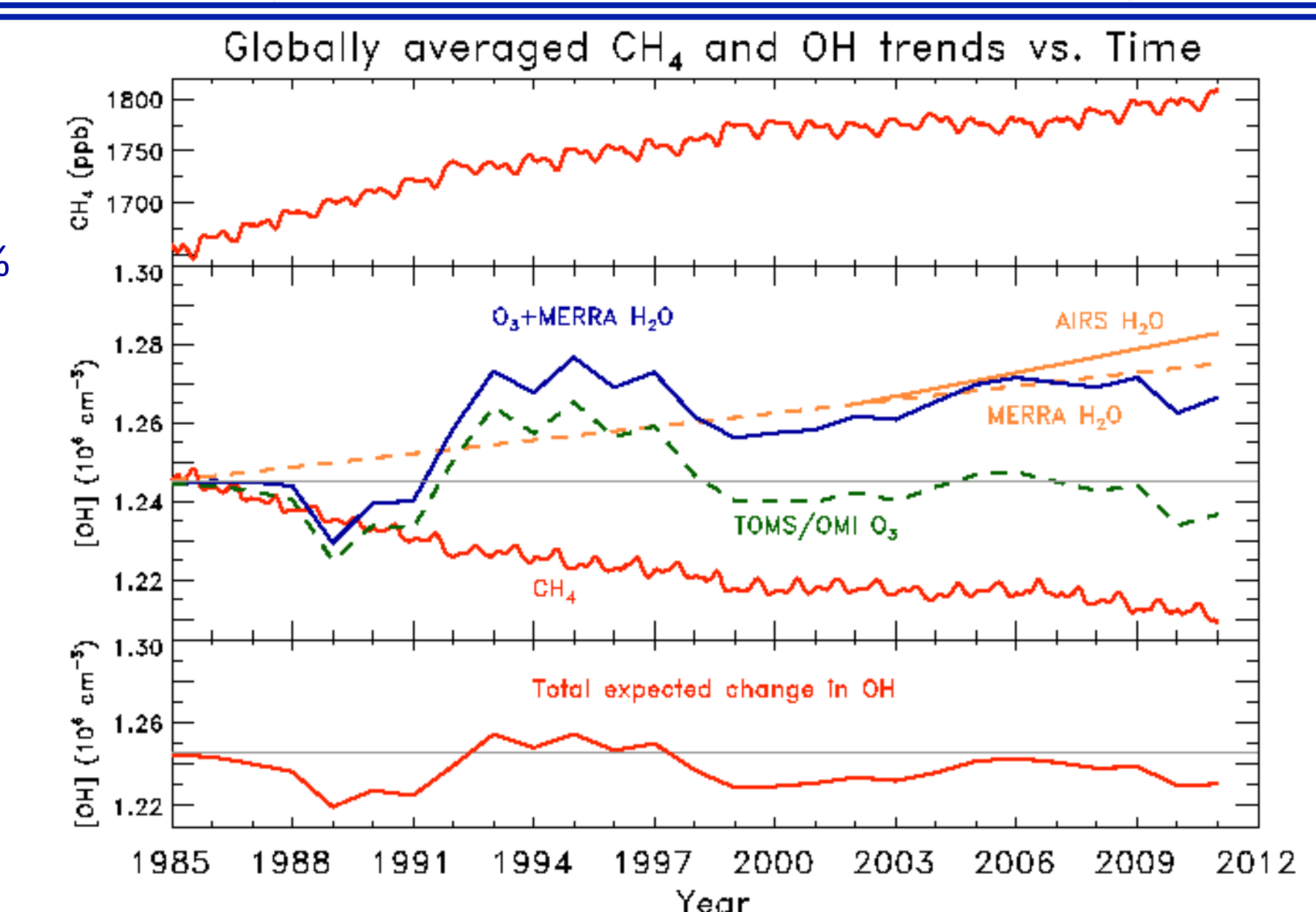
### Previous Work on OH<sup>TROP</sup> Trends

- Montzka et al., 2011 found that OH<sup>TROP</sup> does not vary interannually (from 1997 to present)
- Prior studies by Prinn et al., 2001 and Bousquet et al., 2005 suggest large interannual variability in OH<sup>TROP</sup> (1985-2000)
- We suggest the OH<sup>TROP</sup> behavior in all three studies may be physically possible, based on our preliminary results



### Current Conclusions from OH<sup>TROP</sup> Analysis

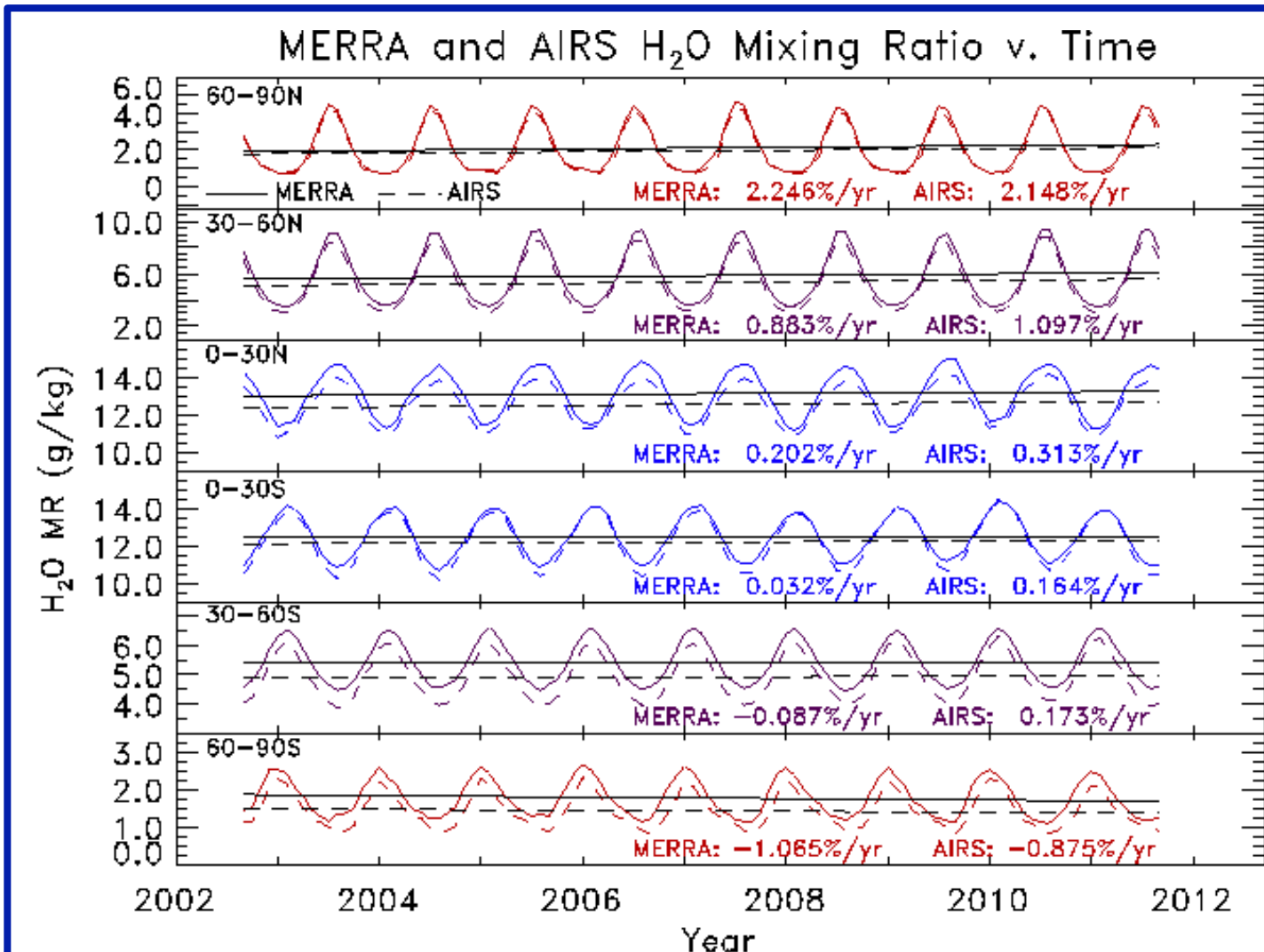
- Effect of CH<sub>4</sub> on OH<sup>TROP</sup> is taken from IPCC (2001), Section 4.2.1.1, which states that "the feedback of CH<sub>4</sub> on tropospheric OH" found using contemporary chemical transport models is -0.32% for every 1% increase in CH<sub>4</sub> (red line, middle panel of figure to right)
- Primary effect of overhead O<sub>3</sub> is rise in OH<sup>TROP</sup> following the 1991 eruption of Mount Pinatubo (green dashed line) due to enhanced removal of stratospheric O<sub>3</sub> by volcanic aerosol
- Rising H<sub>2</sub>O from MERRA and AIRS increases OH<sup>TROP</sup> (orange lines) by an amount comparable to the decrease expected from rising CH<sub>4</sub>
- Overall expected change in OH<sup>TROP</sup> (bottom panel) shows higher level of interannual variability prior to ~1999 and lower variability thereafter



## 2. Methods for estimating changes in OH<sup>TROP</sup>

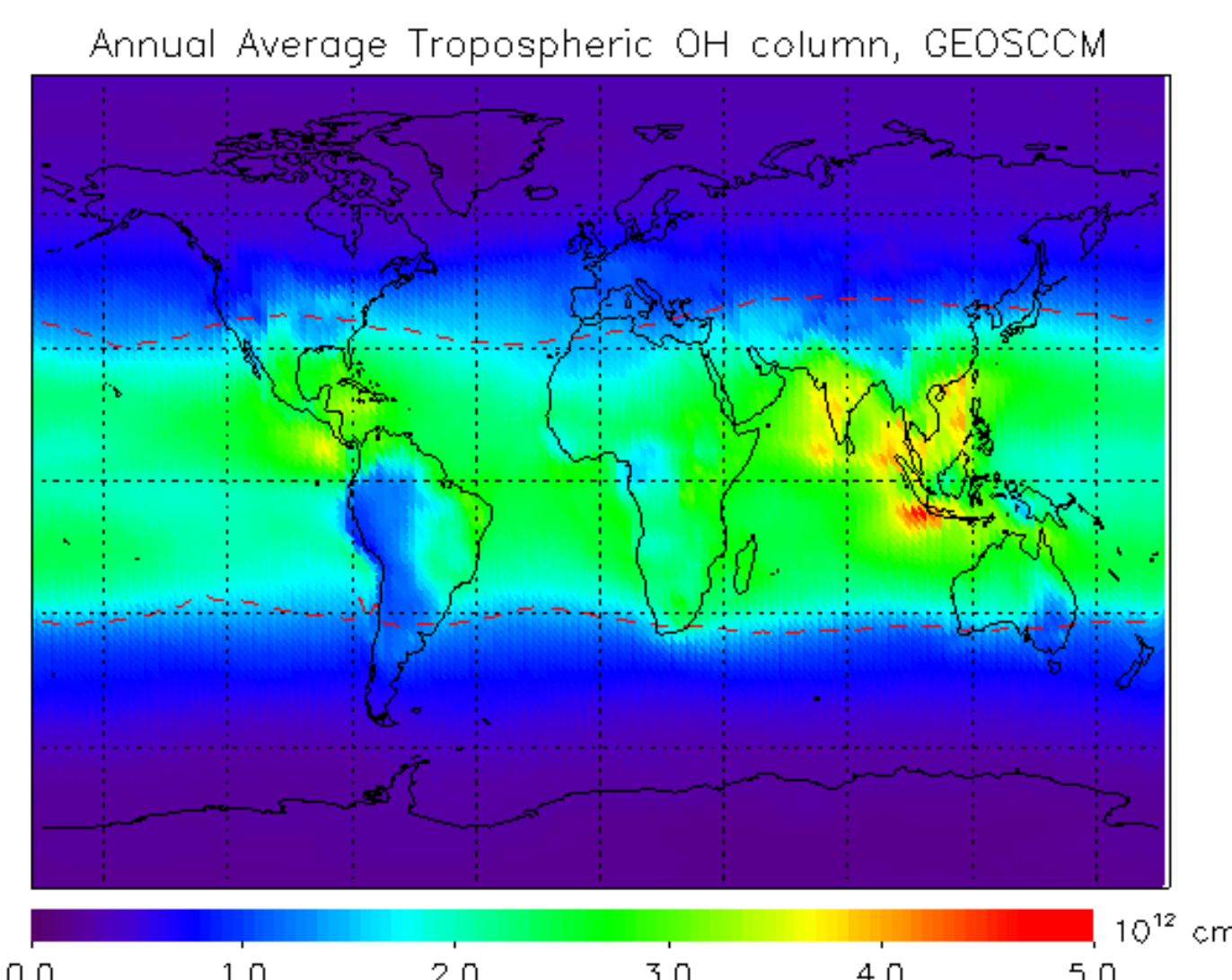
### Base Case OH Conditions

- Initial modern-day conditions for OH are taken from a time-slice run of GEOS CCM using 2005 emissions
- Monthly mean mixing ratios of OH and related species are provided on a 144 longitude, 91 latitude, 72 pressure level grid
- Calculated changes in OH due to H<sub>2</sub>O and overhead O<sub>3</sub> are applied to initial OH field



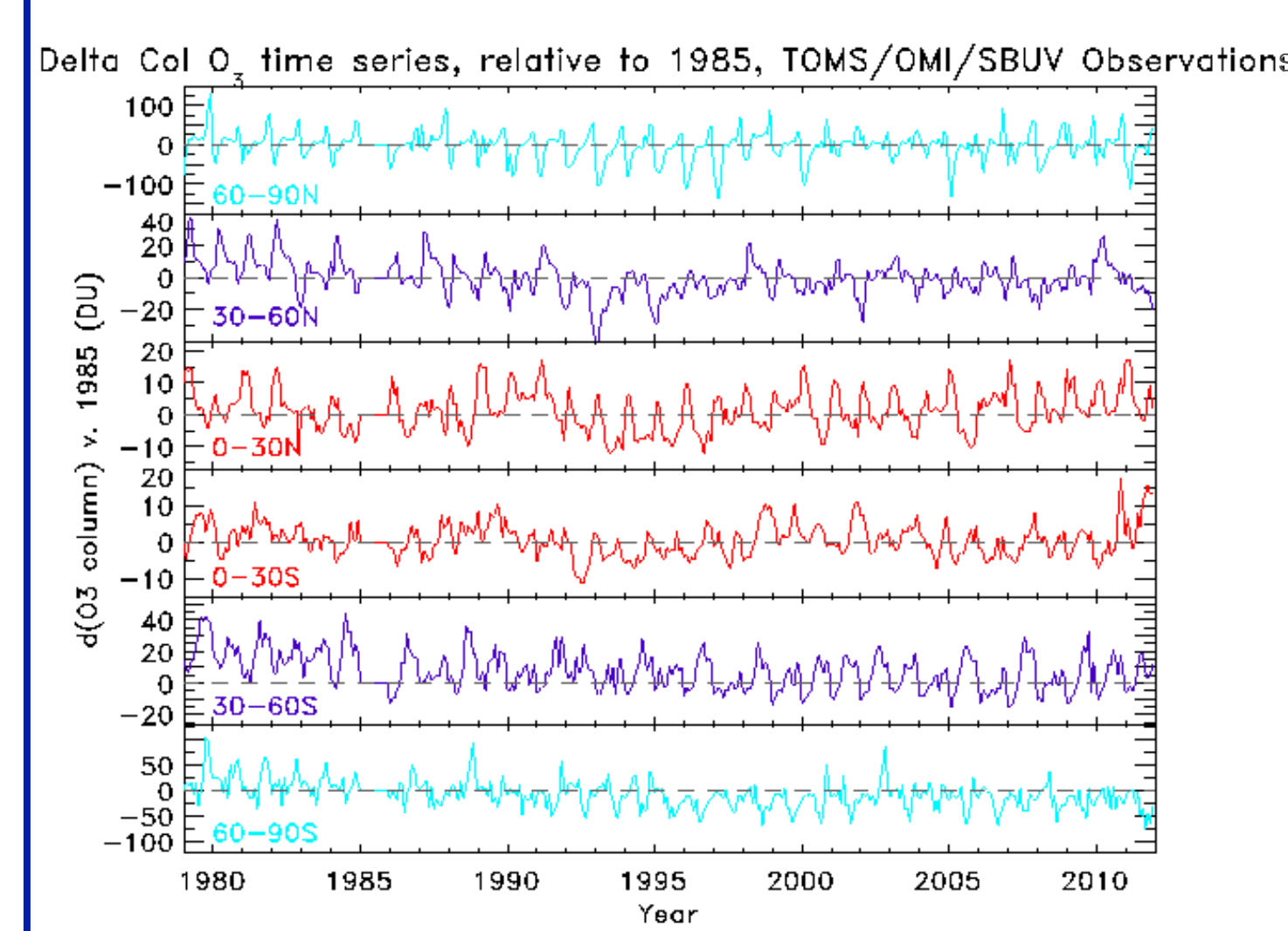
### MERRA & AIRS H<sub>2</sub>O

- Specific humidity files from the NASA Modern-Era Retrospective analysis for Research and Applications (MERRA) starting prior to 1985, and
- H<sub>2</sub>O mixing ratio files from the NASA Atmospheric Infrared Sounder (AIRS) starting in 2002 were used to establish trends in H<sub>2</sub>O by latitude
- OH<sup>TROP</sup> is assumed to follow the square root of the change in H<sub>2</sub>O using a steady-state assumption
- We plan to refine the effect of changing H<sub>2</sub>O on OH<sup>TROP</sup> by examining reaction rates from archived runs of GEOS CCM



### Overhead O<sub>3</sub> Observations

- Total column O<sub>3</sub> trends were obtained from the NASA merged O<sub>3</sub> data set, consisting of measurements from SBUV, TOMS, and Aura OMI instruments
- We then use our photolysis code to estimate the impact on J(O<sub>3</sub>) → O(<sup>1</sup>D) of decreasing initial GEOS CCM overhead O<sub>3</sub> columns by amount suggested by the NASA product
- OH<sup>TROP</sup> is assumed to change by the square root of J(O<sub>3</sub>) → O(<sup>1</sup>D)



## 4. Future plans for refining estimates of ΔOH<sup>TROP</sup>

### Improve Estimate of d(OH<sup>TROP</sup>) / d(H<sub>2</sub>O)

- Reaction rates from recent runs of GEOS CCM are archived for reactions such as:
 
$$\text{H}_2\text{O} + \text{O}(\text{^1D}) \rightarrow 2\text{OH}$$
- Using these reaction rates we will determine the proportion of OH that is produced via reaction with H<sub>2</sub>O
- The determined scaling factor would be used to calculate a new ΔOH<sup>TROP</sup> based on the H<sub>2</sub>O trends
- Estimate time- and pressure-varying values of d(OH<sup>TROP</sup>) / d(H<sub>2</sub>O)
- Evaluate discrepancies between MERRA and AIRS H<sub>2</sub>O trends

### Propagate Uncertainties

- Calculate uncertainties in AIRS and MERRA H<sub>2</sub>O and NASA O<sub>3</sub> product

### Evaluate CH<sub>4</sub>/OH Feedback

- We will use a box model (details below) to probe relationship between CH<sub>4</sub> and OH<sup>TROP</sup> and its dependence on NO<sub>x</sub>

### Box Model

- We will use the GSFC Combined Stratosphere-Troposphere (COMBO) box model provided by Chang Lang (JHU):
  - GMI chemical mechanism
  - 118 species, 321 thermal reactions, and 81 photolysis reactions
  - 5 modules:
    - Aerosol optical depth & surface area
    - Photolysis scheme
    - Thermal reactions scheme
    - Differential eqn solver
    - Input-output
  - Fast-JX photolysis & SMVGear II solver

- Evaluate standard deviation in average fraction of OH production occurring via H<sub>2</sub>O + O(<sup>1</sup>D)
- Estimate uncertainty in the box model evaluation of d(OH<sup>TROP</sup>) / d(CH<sub>4</sub>)

### Assessing OH<sup>TROP</sup> in CCMs

- Through our involvement with the IGAC / SPARC Chemistry-Climate Model Initiative, we have requested:
  1. hourly, instantaneous output from participating CCMs 1 day/season, 1 year/decade
  2. archival of all species, reaction rates, J-values, and physical parameters relevant to OH chemistry
  3. this output for both the REF-C1 (hindcast) and REF-C2 (future) runs
- We plan to assess the causes of differences between OH in the CCMs
- Use of the box model enables us to distinguish between OH differences due to **chemical mechanism** and those due to **differences in OH precursors**
- We can also use this output to predict future trends in OH<sup>TROP</sup>, based on CH<sub>4</sub>, H<sub>2</sub>O, and overhead O<sub>3</sub> from the future CCM runs