Airborne Particles:

What We've Learned About Their Role in Climate From *Remote Sensing*, And Prospects for Future Advances

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Haboob, Khartoum Sudan 2007

Photo Credit: Paul Currion on Flicker

Mt. Etna Eruption 27-30 October 2002



Station Fire near JPL, Pasadena CA August-September 2010

From: http://hometown-pasadena.com

Saharan Dust Plume Tracked Across the Atlantic 05-12 June 1967 ESSA-5 Vidicon Imager



Over Mauritania, Western Sahara, and the eastern Atlantic 7 June 1967

From: *Prospero et al., Earth & Planet. Sci. Lett.* 1970

Global, Over-Ocean Column Aerosol Amount July 1989 - June 1991 NOAA AVHRR



Northern Winter



Northern Summer

From: Husar et al., JGR 1997

Mars Dust Storm – Viking Orbiter 1976



Martian Sky – Viking Lander 1, 1976



SeaWiFS – Sahara Dust over Canary Islands 06 March 1998



MODIS – Fires in Alaska 01 July 2004 21:40 UTC



Even DARF and Anthropogenic DARF are NOT Solved Problems (Yet)





FIGURE SPM-2. Global-average radiative forcing (RF) estimates and ranges in 2005 for anthropogenic carbon dioxide (CO₂), methane (CH₄), nitrous oxide (N₂O) and other important agents and mechanisms, together with the typical geographical extent (spatial scale) of the forcing and the assessed level of scientific understanding (LOSU). The net anthropogenic radiative forcing and its range are also shown. These require summing asymmetric uncertainty estimates from the component terms, and cannot be obtained by simple addition. Additional forcing factors not included here are considered to have a very low LOSU. Volcanic aerosols contribute an additional natural forcing but are not included in this figure due to their episodic nature. Range for linear contrails does not include other possible effects of aviation on cloudiness. $\{2.9, Figure 2.20\}$

IPCC AR3, 2001 (Pre-EOS)

IPCC AR4, 2007 (EOS + ~ 6 years)

Global Energy Flows (W/m²)



Wild et al., BAMS 2012

Aerosol Contribution to Global Climate Forcing

- Cloud-free, global, *Over-ocean*, vis, TOA DARF relative to zero aerosol: -5.5 ± 0.2 W/m² This is a *measurement-based* value, with *uncertainty based on diversity* among estimates (actual uncertainties are probably larger)
- Taking 20% of aerosol to be anthropogenic, the *human-induced component* is: -1.1 \pm 0.4 W/m²
- Global TOA anthropogenic <u>total</u> ARF relative to pre-industrial: -1.3 (-2.2 to -0.5) W/m² This is a model-based value, with uncertainty defined as <u>diversity</u> among estimates; (actual uncertainties are probably much larger)
- The models tend to agree on global AOD (as constrained by satellite & surface obs.), but differ on *regional-scale AOD*, aerosol *SSA*, and *vertical distributions*: *CCSP - SAP 2.3*, 2009

How Good is "Good Enough"??

Climate Sensitivity, Aerosols, and Climate Prediction



- *Models are constrained by* historical global mean surface temperature (*GMST*) *change*
- Forcing by LL greeenhouse gas increase since pre-industrial: ~ 2.6 W/m^2
- Δ GMST *Expected*: ~ 2.1 K; Δ GMST *Observed*: ~ 0.8 K
- Discrepancy dominated by Aerosol Forcing vs. S (disequilibrium, natural variation, etc. are less)
- Model Aerosol Forcing choices compensate for Climate Sensitivity differences (Kiehl, GRL 2007)
 - → Aerosol forcing uncertainty directly impacts confidence in model predictions From a policy perspective, this bears upon the *urgency of mitigation* efforts

AOD Alone is Not Enough – Even for Direct Aerosol Radiative Forcing

Direct Aerosol Radiative *Forcing Efficiency* **per unit AOD**



- Aerosol SSA, Vert. Dist., and Surface Albedo critical, esp. for Surface Forcing
- For Semi-direct Forcing, Aerosol SSA and Vertical Distribution are critical

Constraining DARF – The Next Big Challenge



• Agreement among models is *increasingly good for AOD*, given the combined *AERONET*, *MISR*, and *MODIS* constraints

• The next big observational challenge: Producing *monthly, global maps of Aerosol Type*

How Good is Good Enough?

Instantaneous AOD & **SSA** uncertainty upper bounds for ~1 W/m² TOA DARF accuracy: ~ 0.02 CCSP - SAP 2.3, 2009

Aerosols "Indirect" Forcing of Clouds



Aerosol *Particle Size* Matters

- -- Not easy for remote-sensing techniques to observe the smallest, most numerous CCN
- -- Deducing small-size CCN from larger-particle distribution depends sensitively on ambient RH

• Aerosol Particle Composition Probably Matters Too

-- Remote-sensing not very sensitive to particle chemistry (polarization should help)

• Location, Location, Location

-- Satellite remote-sensing cannot observe aerosol *below* most clouds; difficult observing aerosol near clouds as well

• Clouds, Ambient Meteorology Affect Aerosol Retrievals

The NASA Earth Observing System's Terra Satellite



Terra Project Office / NASA Goddard Space Flight Center

MODerate-resolution Imaging Spectroradiometer [MODIS]

- NASA, Terra & Aqua
 - launches 1999, 2001
 - 705 km polar orbits, descending
 (10:30 a.m.) & ascending (1:30 p.m.)
- Sensor Characteristics
 - 36 spectral bands ranging from 0.41 to 14.385 μm
 - cross-track scan mirror with 2330 km swath width
 - Spatial resolutions:
 - 250 m (bands 1 2)
 - 500 m (bands 3 7)
 - 1000 m (bands 8 36)
 - 2% reflectance calibration accuracy
 - onboard solar diffuser & solar diffuser stability monitor



Improved over AVHRR:

- Calibration
- Spatial Resolution
- Spectral Range & # Bands

Global, Monthly Average MODIS Aerosol Products July 2010



AOD at 550 nm

0.35

0.45

0.55

0.65

-0.05

0.05

0.15

0.25

0.75

Mid-visible Aerosol Optical Depth



Fine-mode Fraction, with AOD encoded as color saturation

Multi-angle Imaging SpectroRaliometer





- <u>Nine</u> CCD push-broom <u>cameras</u>
- <u>Nine view angles</u> at Earth surface: 70.5° forward to 70.5° aft
- Four spectral bands at each angle: 446, 558, 672, 866 nm
- Studies Aerosols, Clouds, & Surface

Ten Years of Seasonally Averaged Mid-visible Aerosol Optical Depth from MISR



... includes bright desert dust source regions

MISR Team, JPL and GSFC

Multi-year Annual Average *Aerosol Optical Depth* from Different Measurements + *Synthesis* (S^{*})



From: Kinne et al. ACP 2006

Aerosol Source Characterization by Combining Measurements and Models

GoCART Inverse-Model-Retrieved Emissions (10⁷ kg/day)

MODIS July 2006 Siberian Smoke Plume Image + AOD, and 5 GoCART Forward-Model Simulations with different source strengths

From: Dubovik et al., ACP 2008

MISR-Derived Ash Plume Aerosol Amount & Properties Eyjafjalljökull Volcano 19 April 2010

MISR Team, JPL and GSFC

MISR Aerosol Type Distribution MISR Version 22, July 2007

Kahn, Gaitley, Garay, et al., JGR 2010

Saharan Dust Source Plume

Bodele Depression Chad June 3, 2005 Orbit 29038

Dust is injected near-surface...

Kahn et al., JGR 2007

Transported Dust Plume

Atlantic, off Mauritania March 4, 2004 Orbit 22399

Transported dust finds elevated layer of relative stability...

Kahn et al., JGR 2007

MISR Stereo-Derived Plume Heights 07 May 2010 Orbit 55238 Path 216 Blk 40 UT 12:39

D. Nelson and the MISR Team, JPL and GSFC

CALIPSO Lidar Aerosol Layer Height "Curtains"

Seasonally aggregated dust & non-dust vertical extinction profiles over Eastern China for 2007

Over-Land Aerosol Short-wave Radiative Forcing w/Consistent Data

The slope of: **MISR AOD MISR SSA TOA albedo vs. AOD** AOD (Green) SSA (Green) 0.96 0.98 1.00 For data stratified by: (d) **Surface BHR MISR Surf. BHR** MISR ANG AE BHR (Green) 1.0 1.5 **Produces**: **Bright surface** Blue Green Red NR + dark aerosol 0.5 = decreasing **Spectral aerosol** albedo w/AOD BHR radiative efficiency $(\mathrm{d}\alpha_{TOA}/\mathrm{d}\tau_{mid-vis})$ 0 0.8 1 0.8 1 0.8 1 0.8 SSA SSA SSA SSA -0.20 -0.10 -0.05 -0.02 -0.01 0.01 0.02 0.05 0.10 0.20

Depends on aerosol microphysical properties relative to surface albedo

Y. Chen et al. JGR 2009

Aerosol Material Fluxes: Atlantic Dust & Asian Pollution

Dust Transport Estimate (Tg) May-October (Top) January-April (Bot) *Kaufman et al., JGR 2005*

2093

20%

MODIS AOD & type, Field Campaign aerosol properties & vertical distribution, GEOS model winds; Compared with GOCART and GMI model Fine-particle mass fluxes

Yu et al., JGR 2008

Current MISR & MODIS Mid-Visible AOD Sensitivities

• MISR: 0.05 or 20% * AOD overall; *better over dark water*

[Kahn et al., 2005; 2010]

• MODIS: 0.05 or 20% * AOD over land 0.03 or 5% * AOD over dark water

[Remer et al. 2005; 2008; Levy et al. 2010]

Based on AERONET coincidences (cloud screened by *both* sensors)

→ Direct Aerosol Radiative Forcing (DARF): Need AOD to <~ 0.02

→ Particle Properties are Categorical rather than continuous Quantities

Kahn, Survy. Geophys. 2012

Comparative Planetology and the Atmosphere of Earth

- Comparative Planetology Discovering how planetary atmospheres are similar, and how they are different, expands our appreciation of Earth itself, by placing specific attributes of our planet into a larger context.
 -- Radiative and Dynamical Scaling Laws
- 2. Subtle Earth Effects Some phenomena in Earth's atmosphere are of much greater physical importance in the atmospheres of other planets.
 -- Venus' Greenhouse; Jupiter's Magnetosphere
- **3. Data Available Only from Other Planets** Data of comparable or higher quality relevant to Earth can sometimes be found in other places. *-- Inner solar system climate record from Mars (and the Moon?)*
- **4.** New Ideas Inspiration leading to a habit of out-of-the-box thinking...

N/S/

