

Progress and challenges in short- to medium-range coupled prediction

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Synopsis

The availability of GODAE Oceanview-type ocean forecast systems provides the opportunity to develop high-resolution, short- to medium-range coupled prediction systems. Several groups have undertaken the first experiments based on relatively unsophisticated approaches. Progress is being driven at the institutional level targeting a range of applications that represent their respective national interests with clear overlaps and opportunities for information exchange and collaboration. These include general circulation, hurricanes, extra-tropical storms, high-latitude weather and sea-ice forecasting as well as coastal air-sea interaction. In some cases, research has moved beyond case and sensitivity studies to controlled experiments to obtain statistically significant metrics.

Lead author's biography

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Introduction

The Global Ocean Data Assimilation Experiment (GODAE)¹ (Bell et al., 2010) succeeded in demonstrating the feasibility of constraining a mesoscale ocean model to perform routine analyses and forecasts through the data assimilation of the Global Ocean Observing System (GOOS). Development of ocean forecasting has since been consolidated and extended under the GODAE OceanView (GOV)² (Schiller and Dombrowsky, 2014). There are now several agencies and centres supporting first- or second-generation global and basin-scale pre-operational and operational ocean prediction systems as described in this special issue. These systems provide routine estimates of the ocean state for both nowcasts and short-range forecasts. The performance has been shown to have sufficient skill in the upper ocean to positively impact a wide range of ocean specific applications (e.g., defence³, search and rescue⁴ etc). Unlike waves where there is a very tight relationship between the skill of the winds and the skill of the waves, the oceans inertia and heat capacity leads to a circulation that has unique time and space scales that is related more to the integrated (time history) of surface fluxes of mass, heat and momentum rather than an immediate response to the atmospheric weather. Important exceptions apply, however, for example over the continental shelf and in the turbulent surface layer where the time and space scales are a blend between the atmosphere, waves, sea-ice and ocean systems. These regions also correspond to the highest biological and human activity and the majority of applications for ocean prediction. Therefore, minimising errors in the applied stress and fluxes will have a high yield for the benefit of ocean prediction.

61 The availability of GOOS and GOV-type forecast systems provides the opportunity to develop
62 high-resolution, short- to medium-range coupled prediction systems (SMRCP) for the earth
63 system. Making progress in this field is a significant challenge due to the added complexity in
64 all areas of development, coupled frameworks, coupled modelling, coupled initialisation,
65 observational requirements (including experimental campaigns) and large and more diverse
66 teams of scientific experts. There have been several vision papers^{5,6} (Brassington, 2009;
67 Brunet et al., 2010) and workshops relevant to this area driven predominantly by the needs of
68 Numerical Weather Prediction (NWP) at ECMWF⁷ and followed on by the UK Met Office⁸.
69 The GOV science team recognised the need to explore the potential benefit to both oceanic
70 and atmospheric prediction through the use of GOV-type system in coupled prediction
71 research. The Short- to Medium-Range Coupled Prediction Task Team (SMRCP-TT) was
72 set-up at the beginning of GOV in 2009 to coordinate an information exchange for the new
73 developments beginning at some centres in the area of coupled prediction on the medium-
74 range. The scope and objectives of the TT were defined to focus on issues of direct relevance
75 to GOV activities and expertise, while recognising that the area of coupled prediction requires
76 inputs from a number of other disciplines coordinated by other international bodies. The
77 scope of the TT was therefore defined as covering: SMRCP of the ocean, marine boundary
78 layer, surface waves and sea-ice; on global and regional scales; to pursue the development
79 of coupled prediction systems for improving and extending ocean/wave/sea-ice state
80 estimation and forecast skill; with specific coupling focii: ocean-wave-atmosphere and ocean-
81 sea-ice-atmosphere. A key achievement of this group was to initiate a linkage with the
82 Working Group for Numerical Experimentation and to convene a Joint GOV-WGNE workshop
83 was held March 2013, Washington DC, USA ([https://www.godae-
84 oceanview.org/outreach/meetings-workshops/task-team-meetings/coupled-prediction-
85 workshop-gov-wgne-2013/](https://www.godae-oceanview.org/outreach/meetings-workshops/task-team-meetings/coupled-prediction-workshop-gov-wgne-2013/)).

86
87 Land surface modelling for atmospheric forecasting has a longer history^{9,10,11} (de Rasnay et al
88 2014, Ek et al 2003, Pitman 2003) than atmosphere-ocean forecasting and predates the
89 development of earth modelling frameworks. Land-surface schemes were first introduced as
90 a sub-model and embedded within the atmospheric model software. As land-surface models
91 have increased in sophistication these have matured into stand alone models. This
92 component of the earth system is beyond the scope of this paper.

93
94 Earth system modelling has evolved through specialist communities for each of the major
95 components. The requirement to develop coupled earth system models, initially for climate
96 applications, has seen the development of computational frameworks to permit component
97 models to be coupled through the synchronous and efficient exchange of fluxes for high
98 performance computational environments. The US government agencies have adopted the
99 Earth System Modeling Framework (ESMF; <http://www.earthsystemmodeling.org>) as the
100 basic architecture for coupling models. ESMF allows for the passing of variables among the
101 models in memory and organises horizontal interpolation between the fields in the different
102 model components via an exchange grid. On top of ESMF, the National Unified Operational
103 Prediction Capability (NUOPC; <http://www.weather.gov/nuopc>) standardises ESMF interfaces
104 further to promote plug-compatibility of models in couplers and passes information through
105 separate flux computation modules. NUOPC is a consortium of the Navy, NOAA, and Air
106 Force modelers and their research partners. Similar efforts have been undertaken within
107 Europe such as the Ocean Atmosphere Sea Ice Soil coupler version 4 (OASIS4)¹² (Redler *et*
108 *al.*, 2010). Achieving all of the requirements for earth system frameworks including platform
109 independence, interoperability, scalability and others has been elusive but major progress
110 has been achieved in the past decade of development. Availability of these frameworks has
111 aided and accelerated research and development for SMR applications.

112
113 In this paper we summarise some of the progress being made within national/international
114 centres in section 2, identify a selection of applications that demonstrate the impact of
115 coupling in section 3; provide a brief overview of some of the known challenges in section 4
116 and conclude with a discussion on the future outlook for this area.

117 118 **Progress by national programs**

119

120 Coupling of the ocean, atmosphere and sea-ice has been developed over a number of years
121 for seasonal and longer-range prediction, but it has been a relatively new area for the
122 development of SMRCP forecasts. During the past 5 years research programs have emerged
123 within the leading centres: Bureau of Meteorology, Australia; Met Office, United Kingdom
124 (UK); National Oceanic and Atmospheric Administration (NOAA)/National Centers for
125 Environmental Prediction(NCEP), United States of America (USA); European Centre for
126 Medium-range Weather Forecasting (ECMWF); Naval Research Laboratory, USA;
127 Environment Canada, Canada; Mercator-Océan/Météo France, France; and NASA, USA. The
128 present systems being applied to study the impacts of coupling are summarised in Tab 1 and
129 outlined below in more detail. The modelling systems range from regional to global and are
130 relatively sophisticated given the availability of earth-system frameworks from the climate
131 community, an example of which is shown in Fig 1. These systems however use relatively
132 unsophisticated approaches to data assimilation where the Background error covariances are
133 uncoupled or weakly coupled and a variety of approaches are adopted to initialise the
134 coupled model.

135 **Bureau of Meteorology, Australia**

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137
138 The Australian Bureau of Meteorology has pursued research into the impact of coupling
139 between the OceanMAPS forecast system and operational NWP systems using a regional
140 nested framework referred to as CLAM (Coupled Limited Area Model). CLAM is based on the
141 UK Met Office Unified Model (UM) version 6.4¹³ (Davies *et al.*, 2005), the Ocean Atmosphere
142 Sea Ice Soil coupler version 4 (OASIS4)¹² (Redler *et al.*, 2010) and MOM4p1¹⁴ (Griffies,
143 2009). The NWP system known as the Australian Community Climate Earth System
144 Simulator (ACCESS), comprises a suite of atmospheric model configurations from global to
145 regional using four-dimensional variational data assimilation (4DVAR), which was developed
146 for the UM¹⁵ by Rawlins *et al.* (2007). The ocean forecast system is known as the Ocean
147 Model, Analysis and Prediction System (OceanMAPS; Brassington *et al.*, 2012)¹⁶, which uses
148 an eddy-resolving ocean model and an ensemble optimal interpolation scheme called the
149 Bluelink Ocean Data Assimilation System (BODAS; Oke *et al.*, 2008)¹⁷.

150
151 The CLAM infrastructure has been used both in Tropical Cyclone (TC) forecasting research¹⁸
152 (Sandery *et al.*, 2010) and in ACCESS-RC (RC stands for the operational regional
153 atmospheric model (ACCESS-R) coupled to a matching nested regional ocean model), an
154 application of CLAM designed to study the impact of coupling on regional ocean and weather
155 prediction. CLAM was recently used to develop an ensemble coupled initialisation method
156 using cyclic bred vectors¹⁹ (Sandery and O’Kane, 2014). Results using ACCESS-RC have
157 found that ocean-atmosphere coupling offers improvements in the atmospheric model sea
158 surface temperature (SST) boundary condition in the tropics and is significant to severe
159 weather events at three day lead time compared to persisting an SST analysis initial
160 condition. CLAM offered a significant improvement in the forecast of rainfall for the Brisbane
161 flooding event of 2011²⁰ (Barras and Sandery, 2012). Whilst ACCESS-RC is nested inside
162 data assimilating component systems, until recently it has not explicitly had its own data
163 assimilation.

164
165 A collaborative project between the Bureau of Meteorology and the University of Melbourne
166 funded by the Lloyd’s Register Foundation is examining the impact of coupling on the
167 prediction of marine extremes. This research makes use of a multiply nested Weather
168 Research and Forecasting model (WRF)²¹ with resolution to resolve convective storm
169 development and ocean surface conditions from OceanMAPS¹⁶ and regional/nested ocean
170 model simulations based on MOM4p1. Initial focus has been on the sensitivity to the
171 mesoscale SST gradients of storm development²² to justify further research into the coupled
172 response.

173 **Met Office, UK**

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176 The development of coupled predictions for short-range forecasting at the UK Met Office is
177 being undertaken through a number of projects, all using versions of the Hadley Centre
178 Global Environment Model version 3 (HadGEM3). HadGEM3 combines the Met Office Unified
179 Model (UM) atmosphere^{23,24} (Walters *et al.*, 2011; Brown *et al.*, 2012) and JULES land

180 surface model coupled using the OASIS coupler to the Nucleus for European Modelling of the
181 Ocean (NEMO)²⁵ (Madec 2008) and the CICE sea-ice model²⁶ (Hunke and Lipscombe 2010).
182 The assessment of the impact of coupled predictions over atmosphere- and ocean-only
183 predictions demonstrated a positive impact on 1-15 day atmosphere forecasts from coupling
184 most notably in the Tropics²⁷. The HadGEM3 model is running operationally on a daily basis
185 to produce seasonal forecasts in the GloSea5 system²⁸ (MacLachlan et al. 2014). The ocean
186 component of these operational coupled forecasts have been compared with the operational
187 Forecast Ocean Assimilation Model (FOAM)²⁹ (Blockley et al. 2014) ocean forecasts for the
188 first 7-days of the forecast, and shown to be of comparable accuracy. The ocean fields from
189 these coupled forecasts are now being provided operationally to users through the MyOcean
190 project (www.myocean.eu.org).

191 The assessment, development and operational running of the coupled forecasts described
192 above have all been carried out using initial conditions generated separately for the
193 atmosphere and land from the Met Office NWP analysis, ocean and sea-ice from the FOAM
194 analysis. A “weakly” coupled data assimilation (DA) system is being developed in parallel with
195 the above work in order to provide improved initial conditions for the coupled forecasts (see
196 Tab 2). For this work, and the work described above, the UM is run at 60km horizontal
197 resolution on 85 vertical levels, NEMO is at 25km horizontal resolution on 75 vertical model
198 levels, and CICE is run with 5 thickness categories. The coupled model is corrected using two
199 separate 6-hour window DA systems: a 4DVAR system for the atmosphere assimilating the
200 standard set of atmosphere data¹⁵ (Rawlins et al. 2007) with associated soil moisture content
201 nudging and snow analysis schemes on the one hand, and a 3DVAR First Guess at Analysis
202 Time (FGAT) system NEMOVAR³⁰ (Waters et al 2013) for the ocean and sea-ice (using in
203 situ SST, temperature and salinity profile, satellite SST, satellite altimeter, and sea ice
204 concentration data). The background information in the DA systems comes from a previous 6-
205 hour forecast of the coupled model. Given the short time window the coupling frequency was
206 increased from the default 3 hours to 1 hour. This also has a particular benefit in improving
207 the model representation of the diurnal cycle.

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NOAA/NCEP, USA

211 Whereas coupled modelling has been part of the operational model suite at NCEP (and in a
212 broader scale within NOAA) for almost a decade, efforts of systematic model coupling have
213 been taking off only in the last few years.

214
215 Historically, coupled modelling has been used in tropical cyclone (hurricane in the US)
216 modelling and in seasonal modelling. In hurricane modelling, the impact of ocean temperature
217 and heat content on intensification has been long recognised, and operational GFDL and
218 HWRF models have included an active ocean component for more than a
219 decade^{31,32,33,34,35,36,37} (e.g., Bender et al., 1993, 2007, Bender and Ginis, 2000, Yablonsky
220 and Ginis, 2008, 2009, Tallapragada et al., 2013, Kim et al., 2014). Similar approaches have
221 been used by the US Navy³⁸ (e.g., Hodur, 1997). Experimental coupled hurricane modelling
222 has also focused on the air-sea interactions including explicit modelling of wind waves in a
223 coupled system^{39,40,41,42} (e.g., Moon et al., 2004, 2007, Fan et al, 2009, and academia (e.g.,
224 Chen et al. 2007). The wave coupling has not (yet) made its way into operations at NCEP, but
225 the results of the coupling experiments have contributed to much improved surface flux
226 parameterisations in the coupled ocean-atmosphere models for hurricanes.

227
228 Coupled modelling has also been the staple of reanalysis and seasonal forecasting at NCEP.
229 The most recent reanalysis⁴³ (Saha et al. 2010) and the presently operational Climate
230 Forecast System (CFS-v2, Saha et al., 2014)⁴⁴, represents a coupled atmosphere – ocean –
231 land – ice system, albeit with uncoupled data assimilation efforts for all sub-systems. Land
232 surface models within atmospheric models, has a fairly long history at NCEP for mesoscale
233 models¹⁰ (e.g., Ek et al., 2003), and is in operations in the global and seasonal models^{45,46}
234 (e.g., Wei et al., 2012, Meng et al., 2012). Since the underlying land model is a full model that
235 has been used as a standalone model, this is affectively an example of coupled modelling,
236 although historically this modelling has not been labeled as such.

237

238 Within NOAA, ESMF and the NUOPC layer are used in NOAA's Environmental Modeling
239 System (NEMS). NEMS now incorporates, and is the model driver for, most weather models
240 at NCEP. Ocean, ice and wave models such as HYCOM, MOM5, CICE, GFDL ice model
241 and WAVEWATCH III are now available in NEMS, or will be available in late 2014. This
242 provides NOAA with a set of well-defined building blocks for coupling in general.

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ECMWF, Europe

246 Developments of coupled forecasting systems at ECMWF follow three lines: improvement in
247 the modelling of air-sea interaction processes, use of coupled ocean-wave-seaice-
248 atmosphere models in forecasts at all time ranges (medium range, monthly and seasonal),
249 and the development of ocean-atmosphere coupled data assimilation systems.

250

251 Growing ocean waves play a role in the air-sea momentum and heat transfer while breaking
252 ocean waves affect the upper ocean mixing. Ocean waves also provide an additional force
253 on the mean circulation, the so-called Stokes-Coriolis force. Furthermore, the surface stress
254 felt by the mean circulation is the total surface stress applied by the atmosphere minus the
255 net stress going into the waves. Finally, momentum transfer and the sea state are affected by
256 surface currents. These effects have been introduced in the ECMWF coupled forecasting
257 system, and are currently being assessed. The impact of breaking waves in the upper ocean
258 mixing has been shown to have a large impact on the prediction of SST. Janssen et al 2013⁴⁷
259 provide a detailed description on the representation of these effects, and illustrate their impact
260 on ocean-only simulation and on coupled forecasts.

261

262 Since the thermodynamical coupling is thought to be important in the modeling of tropical
263 convection the coupled ocean-atmosphere-wave model, traditionally used only for the
264 monthly and seasonal forecasts ranges, is also used in the medium range weather prediction,
265 since November 2013. Results show that the coupled model provides better forecasts of the
266 tropical atmosphere, improved forecasts of the MJO, and has impacts on the representation
267 of slow-moving tropical cyclones⁴⁷ (Janssen et al 2013).

268

269 ECMWF has implemented a coupled ocean-wave-atmosphere data assimilation system
270 called CERA (Coupled ECMWF ReAnalysis). This system uses the ECMWF coupled model
271 with an incremental variational approach to assimilate simultaneously ocean and atmospheric
272 observations. The ultimate purpose is to generate better and self-consistent coupled states
273 for atmosphere-ocean reanalysis. The CERA system is based on an incremental variational
274 approach where the ECMWF coupled system is used to compute the misfits with ocean and
275 atmospheric observations in the outer loop. The ocean and the atmosphere share a common
276 24-hour assimilation window but still run separate inner loops. The ocean increment is
277 computed using a 3DVAR method based only on the first misfit computation, while the
278 computation of the atmospheric increment is based on a 4DVAR approach with two outer
279 iterations. An SST nudging scheme has been developed in the ocean model to avoid the
280 rapidly-growing bias of the coupled model.

281

Naval Research Laboratory, USA

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284 The US Navy is actively operating and developing coupled forecasting systems on global and
285 regional scales. For regional scales the air-ocean version of the Coupled Ocean Atmosphere
286 Mesoscale Prediction System (COAMPS)⁴⁸ (Holt et al., 2011) was declared operational in
287 2011. Air-ocean coupled model runs are routinely performed at the Navy operational
288 production centres. The COAMPS system is being updated to include coupling of a wave
289 model⁴⁹ (Allard et al. 2012). Operational implementation of a regional, air-ocean-wave
290 coupled system is planned for 2015. Fig 1 shows the coupling interfaces for the fully coupled
291 COAMPS. The various components of the coupled system are integrated through ESMF.

292 A coupled global ocean/ice model will be operational in 2014. At the present time, the
293 coupled ocean/ice model is restricted to the Arctic Ocean (Arctic Cap Nowcast Forecast
294 System). The new global ocean/ice system will produce nowcasts and 120-hour coupled
295 model forecasts of ice fields from CICE and ocean fields from HYCOM at 1/12 degree
296 resolution.

297 A coupled global atmosphere/ocean/ice/wave/land prediction system providing daily
298 predictions out to 10-days and weekly predictions out to 30-days is being developed as a
299 Navy contribution to the Earth System Prediction Capability (ESPC). A schematic of the
300 system is shown in Fig 2. Initial Operational Capability (IOC) is targeted for 2018. ESPC is a
301 national partnership among federal agencies and the research community in the U.S. to
302 develop the future capability to meet the grand challenge of environmental predictions in the
303 rapidly changing environment. The system will be based on NUOPC and use analysis fields
304 of each component as initial conditions and make daily forecasts out to 10-days. Throughout
305 each weekly cycle, predictions out to 30-days will be constructed.

306 Data assimilation in coupled COAMPS currently consists of independent 3DVAR analyses in
307 the ocean and atmosphere. The first-guess fields (6- or 12-hour forecasts) for each fluid are
308 obtained from the coupled model state. This assimilation configuration is referred to as
309 weakly-coupled. A strongly coupled 4DVAR assimilation system for both the ocean and
310 atmospheric components of COAMPS is under development. In this scheme separate
311 4DVAR assimilation systems of the atmosphere and ocean models will be linked through the
312 existing coupling terms and ESMF coupling infrastructure in COAMPS. The tangent linear
313 and adjoint components of these coupling terms will be developed and used to minimise the
314 cost function of the coupled system. The state and observation vectors in the assimilation will
315 be extended to include both ocean and atmosphere variables.

316 For the global ESPC coupled model a hybrid version of the Navy Coupled Ocean Data
317 Assimilation (NCODA) 3DVAR⁵⁰ (Cummings and Smedstad, 2013) has been developed. The
318 hybrid covariances are a weighted average of the static multivariate correlations already in
319 use and a set of coupled covariances derived from a coupled model ensemble. The coupled
320 model ensemble is created using the Ensemble Transform (ET) technique in both the ocean
321 and atmosphere. One idea being explored is to form a combined ocean/atmospheric
322 innovation vector that is assimilated in independent hybrid 3DVAR-ocean and 4DVAR-
323 atmosphere assimilation systems using ensemble-based coupled covariances.

324 An observation operator has been developed for direct assimilation of satellite SST radiances
325 using radiative transfer modeling⁵¹ (Cummings and Peak, 2014). The radiance assimilation
326 operator has been integrated into NCODA 3DVAR. The operator takes as input prior
327 estimates of SST from the ocean forecast model and profiles of atmospheric state variables
328 (specific humidity and air temperature) known to affect satellite SST radiances from the NWP
329 model. Observed radiances are simulated using a fast radiative transfer model, and
330 differences between observed and simulated radiances are used to force a SST inverse
331 model. The inverse model outputs the change in SST that takes into account the variable
332 temperature and water vapour content of the atmosphere at the time and location of the
333 satellite radiance measurement. Direct assimilation of satellite SST radiances is an example
334 of coupled data assimilation. An observation in one fluid (atmospheric radiances) creates an
335 innovation in a different fluid (ocean surface temperature). The observed radiance variables
336 depend on both ocean and atmosphere physics. The radiance assimilation operator is ideally
337 suited for coupled ocean/atmosphere forecasting systems where the atmosphere and ocean
338 states have evolved consistently over time.

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340
341

Environment Canada

342 The Canadian Operational Network of Coupled Environmental Prediction Systems
343 (CONCEPTS) including Mercator-Océan participation (France) is providing a framework for
344 research and operations on coupled atmosphere-ice-ocean (AIO) prediction. Operational
345 activity is based on coupling the Canadian atmospheric Global Environmental Multi-scale
346 (GEM) model with the Mercator system based on the NEMO, together with the CICE sea ice
347 model. Within CONCEPTS two main systems are under development: a short-range regional
348 coupled prediction system and a global coupled prediction system for medium- to long-range
349 applications⁵² (Smith et al., 2013).

350
351 A fully coupled AIO forecasting system for the Gulf of St. Lawrence (GSL) has been
352 developed⁵³ (Faucher et al., 2010) and has been running operationally at the Canadian
353 Meteorological Centre (CMC) since June 2011. The original ocean-ice component of this
354 system⁵⁴ (Saucier et al., 2003) is currently being replaced by NEMO and CICE. This system

355 is also the basis for the development of an integrated marine Arctic prediction system in
356 support of Canadian METAREA monitoring and warnings. Specifically, a multi-component
357 (atmosphere, land, snow, ice, ocean, wave) regional high resolution marine data assimilation
358 and forecast system is being developed for short-term predictions of near surface
359 atmospheric conditions, sea ice (concentration, pressure, drift, ice edge), freezing spray,
360 waves and ocean conditions (temperature and currents).
361

362 More recently a coupled global AIO system is under development. The first step was the
363 development of the Global Ice-Ocean Prediction System (GIOPS)⁵⁵ Smith et al. (2014).
364 GIOPS is now producing daily 10-day forecasts in real-time at CMC. A 33km resolution global
365 version of the GEM model has been interactively coupled with GIOPS. The models are
366 coupled via a TCP/IP socket server called GOSSIP and exchange fluxes at every timestep.
367 Fluxes are calculated on the higher resolution $\frac{1}{4}^\circ$ NEMO grid. Coupled and uncoupled
368 medium-range (16-day) forecasts have been made and evaluated over the summer and
369 winter of 2011. These forecast trials show statistically significant improvements with the
370 coupled model.
371

372 ***Mercator-Océan/Météo France***

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374 Mercator Océan is developing and operating global and regional ocean analysis and forecast
375 systems. In a closer and long term collaboration with Météo France, Mercator Océan provides
376 ocean initial states for the seasonal forecast systems. More recently, new developments were
377 conducted to investigate high resolution ocean and atmosphere coupling. Meteo-France La
378 Réunion is one of the six Tropical Cyclone Regional Specialized Meteorological Centers
379 handled by the World Meteorological Organization. It is responsible for the issuing advisories
380 and tracking of tropical cyclones (TC) in the South-West Indian Ocean (SWIO). In order to
381 provide better guidance to TC forecasters, Meteo-France has developed ALADIN-Reunion⁵⁶
382 (Faure et al., 2008), a regional adaptation of ALADIN-France⁵⁷ (Fischer et al. 2005). This
383 model has been run operationally since 2006 at 10 km resolution with a specific assimilation
384 scheme, which provides better TC analysis.
385

386 Since 2008, Meteo-France has run a new operational limited-area model AROME-France⁵⁸
387 (Seity et al., 2011) at 2.5km-resolution. This system is designed for very short range forecast
388 in order to improve the representation of mesoscale phenomena and extreme weather
389 events. AROME has its own mesoscale data assimilation system that enable to take benefits
390 from mesoscale data such as radar data. Meteo-France is planning to operate an SWIO
391 regional AROME configuration in the near future.
392

393 Meteo-France and Mercator-Ocean are also exploring the potential benefit of developing an
394 operational coupled version of AROME with a 1/12 degree regional configuration of the NEMO
395 ocean model²⁵ (Madec, 2008). This technological demonstrator has been developed in 2013
396 to explore its feasibility and the impact of air-sea coupling on TC prediction. The ocean
397 surface can cool by several degrees during the passage of a tropical cyclone (TC) due to the
398 associated extreme winds. This cooling decreases the ocean-to-atmosphere heat and
399 moisture supply, which can modulate the TC intensity. Hence, atmospheric models need an
400 accurate description of the sea surface temperature (SST) under TCs to correctly predict their
401 intensities. This SST evolution and its feedback on the TC evolution can only be captured by
402 ocean-atmosphere coupled models.
403

404 ***NASA, USA***

405 In the framework of the Goddard Earth Observing System (GEOS) Data Assimilation
406 System⁵⁹ (Rienecker et al., 2011) of the NASA Global Modelling and Assimilation Office,
407 coupling of the atmosphere-ocean assimilation systems with focus on SST is ready for an
408 operational atmospheric assimilation system. Full coupling with integrated Ocean DAS
409 (iODAS)⁶⁰, Vernieres et al., (2012), is currently being explored. The atmospheric analysis is
410 carried out by Gridpoint Statistical Interpolation (GSI)⁶¹, Kleist et al., (2009), with the GEOS⁶²
411 (Molod et al., 2012) atmospheric model. The iODAS is based on MOM4-(ocean) and CICE
412 (sea-ice) and is coupled to GEOS through the ESMF.
413

414 Using atmospheric surface fields and fluxes, an atmosphere-ocean interface layer models
415 diurnal warming⁶³ (Takaya et al., 2010) and cool-skin⁶⁴ (Fairall et al., 1996) effects upon the
416 SST boundary condition, the skin SST thus computed is then used by the atmospheric DAS
417 to directly assimilate (infrared and microwave) radiance observations using the CRTM
418 (<http://www.star.nesdis.noaa.gov/smcd/spb/CRTM/>) and GSI. Emphasis is on surface
419 temperature sensitive channels of the AVHRR (IR), followed by MW instruments such as TMI-
420 TRMM, AMSR-2, GMI-GPM. In addition, a plan to assimilate in-situ observations within the
421 interface layer is being considered. Other experiments are in-progress to evaluate the impact
422 of the two-way feedback of interactive aerosols at 1/4 degree resolution configuration. The
423 current and near-future plan is to use a simplified version of CICE to provide sea-ice
424 temperature and WavewatchIII so that wave effects can also be included in the interface
425 layer.

426 **Demonstrated benefits**

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429 As noted in the introduction, despite the relatively simple approaches to SMRCP there are
430 many examples that demonstrate quantifiable benefits. At this early stage of research and
431 development it is important to highlight where these benefits are being realised relative to
432 applications to identify leading centres, encourage other institutions to undertake similar
433 research, encourage collaboration between centres for common applications and attract
434 additional funding. Importantly, the list of applications and the examples described represent
435 those of the groups participating in the GOV TT-SMRCP and identified through the Joint
436 GOV-WGNE workshop and represent is not an exhaustive review of all the activities being
437 undertaken by the international community.

438 ***General atmospheric circulation***

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440
441 An example of the impact of the coupling on the ocean forecast skill from the UK Met Office
442 system out to 15 days is shown in Fig 3 for the Tropical Pacific region, the area with the
443 largest positive impact. The coupling clearly benefits ocean forecast skill compared with
444 running the same ocean model in forced mode, with lower RMS and mean errors throughout
445 the 15-day forecasts. To assess the benefit of the weakly-coupled data assimilation, one-
446 month experiments have been carried out, including 1) a full atmosphere/land/ocean/sea-ice
447 coupled DA run, 2) an atmosphere-only run forced by OSTIA⁶⁵ (Donlon et al. 2012) SSTs and
448 sea-ice with atmosphere and land DA, and 3) an ocean-only run forced by atmospheric fields
449 from run 2 with ocean and sea-ice DA. In addition, 5-day coupled forecast runs, started twice
450 a day, have been produced from initial conditions generated by either run 1 or a combination
451 of runs 2 and 3.

452
453 Fig 4 shows the monthly average surface air temperature increments and sea surface
454 temperature increments from the Met Office weakly-coupled and un-coupled analysis runs
455 over December 2011. The ocean and atmosphere increments from the coupled runs are a
456 little smaller in large parts of the globe suggesting a better balance of the fluxes in these runs.
457 There are some locations where this is not the case, but this may be useful to suggest
458 improvements to coupled DA system and also to highlight coupled model biases. In particular,
459 improvements to the lake assimilation may be needed. There are also clearly some issues at
460 high latitudes which merit further investigation. Atmospheric forecasts assessments (not
461 shown) indicate the coupled DA system to be producing improved forecast skill in some
462 variables and regions near the surface such as temperature and relative humidity in the
463 tropics. Ocean forecast skill is similar in coupled runs starting from both coupled and un-
464 coupled analyses at least for the first 5-days, and the impact on longer lead-time forecasts will
465 be investigated in the future.

466 THE ECMWF CERA system produces a coupled 10-day forecast where ocean and
467 atmosphere evolve freely. These coupled forecasts have been compared with the ones
468 produced by an atmospheric operational-like system using the ECMWF atmospheric model at
469 the same resolution (T159L91) as the CERA system. The operational-like system is forced by
470 observed SST during the assimilation and the corresponding atmospheric-only 10-day
471 forecasts are forced by persisted SST anomalies. Fig 5 shows the root mean square error

472 (RMSE) of the SST from the 10-day forecasts in the Tropics for September 2010 with respect
473 to the OSTIA SST analysis. The CERA system provides an initial SST state that is farther
474 from the reference than the operational-like system. But, as the RMSE in the operational-like
475 system increases faster, the CERA system shows better forecast skill for SST by day 4 of the
476 forecast.

477
478 Experiments undertaken by NRL have been performed where the local ensemble transform
479 (ET) analysis perturbation scheme is adapted to generate perturbations to both atmospheric
480 variables and sea surface temperature (SST). The adapted local ET scheme is used in
481 conjunction with a prognostic model of SST diurnal variation and the Navy Operational Global
482 Atmospheric Prediction System (NOGAPS) global spectral model to generate a medium-
483 range forecast ensemble. When compared to a control ensemble, the new forecast
484 ensemble with SST variation exhibits notable differences in various physical properties
485 including the spatial patterns of surface fluxes, outgoing long-wave radiation (OLR), cloud
486 radiative forcing, near-surface air temperature and wind speed, and 24-hour accumulated
487 precipitation. The structure of the daily cycle of precipitation also is substantially changed,
488 generally exhibiting a more realistic midday peak of precipitation. Diagnostics of ensemble
489 performance indicate that the inclusion of SST variation is very favorable to forecasts in the
490 Tropics. The forecast ensemble with SST variation outperforms the control ensemble in the
491 Tropics across a broad set of metrics and variables. The SST variation has much less impact
492 in the Mid-latitudes. Further comparison shows that SST diurnal variation and the SST
493 analysis perturbations are each individually beneficial to the forecast from an overall
494 standpoint. The SST analysis perturbations have broader benefit in the tropics than the SST
495 diurnal variation, and inclusion of the SST analysis perturbations together with the SST
496 diurnal variation is essential to realise the greatest gains in forecast performance⁶⁶ (McLay *et*
497 *al.* 2012).

498
499 The Environment Canada global coupled model based on GIOPS⁵⁵ (Smith *et al.*, 2014)
500 shows robust performance in the tropical atmosphere compared to both tropical moored
501 buoys and analyses produced by the European Centre for Medium Range Weather
502 Forecasts. Evaluation against CMC ice analyses in the northern hemisphere marginal ice
503 zone shows the strong impact that a changing ice cover can have on coupled forecasts. In
504 particular, the coupled system is very sensitive to the ice lead fraction in pack ice and the
505 formation of coastal polynyas. As the ice model does not explicitly model landfast ice there is
506 a tendency to overpredict the opening of the ice cover along coastal regions, which has a
507 strong impact on heat and moisture fluxes to the atmosphere. This sensitivity is under further
508 investigation.

509 ***Madden Julian Oscillation***

510
511
512 The impact of representing the SST in monthly forecasts of the Madden Julian Oscillation
513 (MJO) has been explored at ECMWF. The ECMWF monthly forecasting system has been
514 used to conduct sets of monthly hindcasts where the SSTs have been modified in a controlled
515 manner. The impact of temporal and spatial resolution of SST products has been assessed,
516 as well as the impact of coupling with an active ocean. It is found that while the temporal
517 resolution of the SST matters, the temporal coherence between ocean and atmosphere
518 seems important to simulate tropical convection and propagation of the MJO. By increasing
519 the temporal resolution from weekly to daily the hindcasts of the MJO do not improve,
520 probably because in this experimental setting, the high frequency is uncorrelated between
521 ocean and atmosphere. However, MJO hindcasts improved by coupling to an ocean model
522 instead of using an uncoupled atmosphere model forced by observed SST. In the past it had
523 been shown that ocean-atmosphere coupling produced better MJO hindcasts than
524 prescribing *persistence of SST anomalies* as lower boundary conditions for the atmosphere.
525 However, this was the first time that we have obtained results indicating that ocean-
526 atmosphere coupling produced better MJO forecasts than prescribing *observed SST*⁶⁷
527 Boisseson *et al* 2012. See also Janssen *et al* 2013⁴⁷ for the impact of coupling in the medium
528 range weather forecasts and MJO, using a more recent model version.

529
530 CFSv2 increased useful prediction skills for MJO from 10-15 days for CFSv1 to around 20
531 days⁶⁸ (Wang, W. *et al.*, 2013). This improvement was mostly realized by having better model

532 physics and more accurate initializations. But it did not eliminate all biases for weaker
533 amplitudes and slower propagation of MJO events as compared to observations. While, the
534 weak amplitude could be due to the slower response of the convection to the large-scale
535 dynamical fields, the slow eastward movement is related to lower skill in predicting the
536 propagation across the Maritime Continent, a common problem for several statistical and
537 dynamical models^{69,70,71} (Seo et al., 2009; Matsuedo and Endo, 2011; Rashid et al., 2010).

538 ***Hurricane/Tropical Cyclone prediction***

540
541 In order to evaluate the potential benefit of the ocean atmosphere coupling on TC forecasts in
542 the South West Indian Ocean, Mercator-Ocean has developed a new coupled regional model
543 based on the Meteo-France operational atmospheric model AROME and the NEMO ocean
544 model. As the AROME assimilation system is not available yet for the SWIO region, the
545 atmospheric model is initialised from ALADIN-Réunion 10km analyses, which are generated
546 every 6 hours. The TC specific assimilation scheme allows representing accurately the TC
547 structure, intensity and position in the analysis based on the best estimates provided by TC
548 forecasters. ALADIN-Réunion is also used for lateral boundary conditions. Experiments have
549 been conducted with TCs from the last 6-years using NEMO, which is initialised from the
550 global ¼ degree reanalysis GLORYS⁷² (Ferry et al., 2012). Because of the resolution
551 difference between GLORYS and the NEMO regional configuration, an adjustment period is
552 needed for the model to reach its new equilibrium state. This step is achieved by using a
553 digital filtering initialisation procedure during a 3-days integration period. During this period,
554 the ocean model is also forced with 6-hours ALADIN analysis, which allows equilibrating the
555 ocean surface and mixed layer with the high resolution atmospheric forcing. The coupled
556 system is then integrated during 96-hours with a coupling frequency of 15-minutes via the
557 OASIS3 coupler⁷³ (Valcke et al., 2013).

558
559 The coupled model performances have been evaluated against AROME forecasts forced with
560 the Meteo-France SST analysis over an ensemble of 23 intensifying TC simulations (5
561 different TCs from the 2008-2012 seasons). Sea surface temperature (SST) forecast errors
562 are then calculated by comparing the averaged SST within a 150 km radius centered on the
563 TC with the SSMI TMI-AMSRE product⁷⁴ (Gentemann et al., 2003). TC forecasts are
564 evaluated against TC best-tracks provided by Meteo-France La Réunion. The ensemble
565 averaged SST and minimum pressure errors are presented in Fig 8 as a function of the
566 forecast time for the coupled and the forced simulations.

567
568 Concerning SST (Fig 8a), an important improvement is achieved with the coupled model
569 when compared to the forced model. Averaged SST forecast error never exceed $\pm 0.4^{\circ}\text{C}$
570 in the coupled model, while it can reach $+1.2^{\circ}\text{C}$ with Meteo-France SST analysis. The initial
571 SST error ($+0.8^{\circ}\text{C}$) is mainly due to the lower spatial resolution and the temporal smoothing
572 of the operational SST analysis. The initial oceanic state generated from GLORYS with the
573 DFI procedure is really close to the observations. In the forced ensemble, the SST error
574 slowly increases with the forecast time while it stays close to zero in the coupled ensemble.
575 Hence, the coupling limits effectively SST error growth during the forecast.

576
577 The SST improvements lead to a better TC intensity forecast in the coupled ensemble as
578 shown in Fig 8b. While both coupled and forced ensembles show good skills in predicting TC
579 intensity during the first 30-hour (error $< 10\text{hPa}$), models behaviours differ quickly at longer
580 ranges. Coupled forecasts tends to slightly underestimate TC intensity at all forecast times,
581 but with error $< 10\text{hPa}$ even at 96-hour range. In forced simulations, intensity error quickly
582 increases with time and reaches up to 35hPa at 96-hour range. Consequently, the coupling
583 with NEMO greatly improves AROME TC intensity forecast for ranges greater than 30 hours
584 through a more realistic SST representation.

585
586 These encouraging preliminary results achieved with AROME-NEMO will lead to the
587 development of a real-time operational version to assist TC forecasters in La Reunion. New
588 regional configurations will also be developed for the other French overseas territories where
589 Meteo-France provides weather forecast (South-West Pacific Ocean New Caledonia and
590 Polynesia, Atlantic Ocean French Guinea and Caribbean). NEMO will also benefit of the new

591 operational Mercator-Ocean global 1/12 degree daily forecasts which should improve oceanic
592 initial and boundary conditions.

593
594 The NOAA-GFDL coupled hurricane prediction system that has been run operationally for
595 many years, was designed to account for the effects of upper ocean heat content and the role
596 of the ocean response on TC forecasts. This system has demonstrated significant
597 improvements in TC forecasting skill in the Gulf of Mexico³² (Bender et al, 2007).

598
599 Experiments using a coupled limited area modelling system for tropical cyclones (CLAM-TC)
600 for a number of cases in the Australian region have shown that the representation of the
601 ocean cooling response to the passage of a Tropical Cyclone improves in the coupled system
602 both because surface fluxes are more realistically represented with a high resolution regional
603 atmospheric model compared to a global model and that the negative feedback provided by
604 the ocean response tends to limit over-estimates of the storm intensity¹⁸ (Sandery et al ,
605 2010). The ocean component of this system initialises from the data assimilating OceanMAPS
606 providing an improved representation of sub-surface heat content, which is also an additional
607 benefit of running such a system. The CLAM-TC system was extended to study coupled
608 initialisation and in turn an ensemble method was developed that provided further
609 improvements in forecasting the ocean response to TC-Yasi for both SST and sea-level
610 anomalies¹⁹ (Sandery and O’Kane, 2014). Prediction of SST resulting from the ocean
611 response to tropical cyclone Yasi in the Coral Sea on the 2nd of February 2012 was improved
612 using a coupled ocean-atmosphere ensemble initialisation method as shown in Fig 9.

613 614 ***Extra-tropical cyclones – East Coast Lows***

615
616 East Coast Lows are subtropical low pressure weather systems that can rapidly intensify as
617 they propagate over the marine boundary of Australia’s east coast producing strong localised
618 convection, lightning and heavy precipitation. Several storms have produced severe impacts
619 in terms of coastal flooding, damage from hailstones, and in some cases the grounding of
620 ships and losses of life. Adjacent to the east coast is the so-called East Australian Current, a
621 western boundary current of the South Pacific sub-tropical gyre transporting warm/fresh
622 seawater poleward from the Coral Sea to the Tasman Sea. The EAC is frequently unstable
623 producing several anticyclonic eddies per year from the separation point and along the
624 northern New South Wales coast which can persist for months⁷⁵ (Brassington et al., 2010)
625 providing sources of heat into the Austral winter. A specific case on the 7-9 June 2007 that
626 occurred off Newcastle, NSW has been studied using downscaled Weather Research and
627 Forecast model (WRF) simulations. A simulation is initialised with highly resolved SST
628 (BLUElink) and then compared to a second simulation initialised with coarse resolution (Ctrl)
629 SST boundary conditions to examine the impact of the gradients in SST arising from the large
630 scale warm ocean eddies that persist into the Austral winter²² (Chambers et al). Simulations
631 based on the highly resolved SST produced higher values of 48-hour total precipitation along
632 an SST front (see Fig 10) resulting in more localised convection consistent with observations
633 from coastal rain gauges and with lightning strike locations. It is concluded that the SST
634 gradient along the southern flank of a large warm eddy significantly increased the severity of
635 the coastal weather impacts that occurred during this storm.

636 637 ***High latitude weather and sea-ice forecasting – Gulf St Lawrence***

638
639 Sea-ice acts as a barrier between the atmosphere and the ocean, modulating the fluxes of
640 heat and moisture across an interface often with temperature differences of greater than
641 20°C. As such, rapidly evolving changes in the ice cover can have important impacts for polar
642 weather prediction. This can result from a variety of processes such as ice formation and
643 break-up, coastal polynyas and leads in pack ice. Differences between coupled and
644 uncoupled model forecasts after 12-hours from the Canadian Gulf of St. Lawrence coupled
645 forecasting system are shown in Fig 11. This system has shown the strong impacts that a
646 dynamic sea-ice cover⁷⁶ (Smith et al., 2012) can have on 48-hour atmospheric forecasts
647 leading to large changes in surface air temperature (up to 10°C), low-level cloud cover, and
648 precipitation. The top panel is for a winter case (Mar. 10, 2012) with sea-ice concentration on
649 the left and 2m temperature on the right showing that rapid ice changes can cause surface
650 temperature changes of up to 7-8°C over the open water. Due to the presence of a relatively

651 thin seasonal thermocline (~20m) with cold (<0°C) winter surface waters below, upwelling
652 events in summer can also lead to important impacts on weather predictions. For example,
653 the bottom panel in Fig 11 shows a summer case (Jul. 10, 2012) with 10m winds on the left
654 and 2m temperature on the right showing that coastal upwelling in the coupled forecasts can
655 produce surface temperature changes of several degrees Celcius locally.

656

657

Nearshore coastal weather – Adriatic Sea

658

659 A coupled COAMPS⁴⁸ model was executed in the Adriatic Sea from 25 January to 21
660 February, 2003. The atmospheric model configuration was triply nested (36, 12, 4 km
661 horizontal resolution), while the ocean model consisted of two nests (6 and 2 km), with the
662 inner-most nests of both models centered over the northern Adriatic. Both coupled and
663 uncoupled model runs were performed. In the coupled model run the winds, wind stresses,
664 and heat fluxes were interchanged between the atmosphere and ocean (i.e., the ocean feeds
665 back to the atmosphere and the atmosphere feeds back to the ocean) every 12 minutes using
666 grid exchange processors based on the Earth System Modeling Framework (ESMF). In the
667 uncoupled run, wind forcing from the atmospheric model was passed to the ocean model, but
668 the ocean *did not* feedback to the atmosphere, i.e., the heat fluxes calculated by the
669 atmospheric model were computed using daily averaged analysis-quality SST rather than the
670 time-dependent ocean model forecast SST used in the coupled run. Couple and uncoupled
671 statistics are presented for the Acqua Alta platform near Venice, Italy in Fig 12. Inspection
672 of the wind stress time series shows good agreement, with the RMSE slightly larger in the
673 coupled run (0.112) versus the uncoupled run (0.108). The overall smaller mean stresses in
674 the COAMPS runs (0.118 coupled, 0.135 uncoupled) compared to the observations (0.151)
675 are attributed to intensity and positional differences of the Trieste bora jet during the time
676 period of the experiment. The sensible and latent heat flux comparisons, however, showed a
677 clear improvement in the coupled model run. These results illustrate how the coupled model
678 can more accurately predict surface heat fluxes in near-shore regions where a complex SST
679 field is subject to intense atmospheric events and turbulent heat fluxes have high spatial
680 inhomogeneity and large gradients.

681

682

Data assimilation of brightness temperatures

683

684 The NASA, coupled GEOS-DAS have explored the data assimilation of brightness
685 temperature using a surface sensitive (10.35µm) channel of the AIRS instrument on AQUA
686 satellite. The comparison of an experiment that had an active interface-layer with a control
687 experiment with no interface layer (the SST boundary condition was skin SST) was used to
688 diagnose the benefit.

689

690 Preliminary results, at 1 degree resolution, show improved assimilation of all 10-12micron IR
691 observations and decreased bias in precipitation with respect to GPCP data. Fig 7 shows
692 three panels with time series of total number of observations assimilated (top panel), global
693 mean of observation-minus-background (OMB), middle panel, and standard deviation of the
694 OMB (bottom panel). The use of an improved skin temperature estimate reduced the number
695 of observations rejected by the analysis quality control, corresponding also to a reduced
696 standard deviation in OMB. Similar results were obtained for other 10-12 µm IR channels of
697 AIRS-AQUA, IASI-METOP-A, HIRS4-METOP-A, N19 (not shown).

698

699

Known challenges

700

701 Based on the current sophistication of the coupled modelling systems and the range of
702 applications under active investigation many challenges toward coupled prediction have
703 already been addressed. Sufficient progress has been made in observing, modelling and
704 initialisation to forecast waves, the ocean state and sea-ice to suggest that coupled modelling
705 of the marine environment is feasible. The pursuit of seasonal and climate modelling has
706 introduced several software frameworks that facilitate the coupling of component model
707 software that is scalable for super-computing environments. In practice there are several
708 short-comings in their design for GOV-type forecasting and eventual operational applications
709 which require more frequent restarting and data exchanges. This is not impeding progress in

710 basic research but is impacting the efficiency and size of the problems being undertaken and
711 will require further optimisation in design before implementation into operational applications.
712

713 The pursuit of coupled modelling specific to applications for hurricanes has yielded several
714 advances in air-wave-sea coupled parameterisations for high-wind conditions in the tropics.
715 Significant effort will be required to generalise the coupled parameterisations across all
716 applications. However, less sophisticated parameterisations from existing models are
717 demonstrating positive impacts for a wide range of environments.
718

719 The initialisation of coupled models is currently based on uncoupled or weakly coupled data
720 assimilation for each component model and an inefficient coupled initialisation procedure to
721 produce balanced fields in the coupled model. Some promising results are evident from
722 research focusing on the coupled assimilation of brightness temperatures. Coupled data
723 assimilation is required to provide the optimum dynamically balanced coupled fields but there
724 are several challenges to realising this goal.
725

- 726 1. Proper handling of different time scales in the ocean and atmosphere. These scales
727 may be similar enough in the atmosphere boundary layer and ocean mixed layer to
728 allow coupled modelling and coupled data assimilation to succeed. This aspect of
729 the problem needs to be thoroughly studied.
- 730 2. A goal of coupling is to reduce some of the biases in interfacial fluxes that occur in
731 each component model in their uncoupled form. However, any residual biases in a
732 coupled model will distribute throughout the coupled model state requiring more
733 sophisticated analyses to diagnose, attribute and develop bias correction schemes.
- 734 3. It is still a remaining challenge to decide the best way to weight coupled covariances
735 from ensembles in the hybrid schemes. Similarly, to find appropriate methods for
736 coupled initialisation and maintaining coupled model ensemble spread given the
737 disparate temporal and spatial scales of the ocean and atmosphere. It is also
738 unclear how large an ensemble is needed.
- 739 4. Progress would benefit from community-established benchmarks, test cases, or
740 metrics to establish beneficial impact of fully coupled analyses
741

742 In the near-surface ocean, the diurnal cycle imposes time-scales of a few hours⁶⁴ (Fairall et
743 al., 1996). Modelling of the diurnal warming layer is important for computation of the skin
744 temperature. For coupled data assimilation, it is essential to incorporate observational
745 information *directly* from satellite brightness temperature observations and near-surface
746 buoys so that the modeled skin and near-surface temperature profile is estimated accurately,
747 and thus temporally evolved by the model at the *correct* time-scale. It is also relevant to note
748 the different vertical length-scales observed by the observations: IR observations measure
749 “closest” to the skin or air-sea interface (few microns deep); MW observations penetrate
750 slightly deeper (to few mm); and further down to centimeter – and meter scale – we have in-
751 situ measurements e.g., ships and buoys.
752

753 For coupled prediction in polar environments a significant uncertainty lies in the extent to
754 which we can accurately predict small scale ice features and the evolution of the ice cover.
755 Coupled forecasts are strongly sensitive to variations in the ice cover in the marginal ice zone
756 as well as due to coastal polynya formation and leads in the pack ice. As most sea ice
757 observational data are of fairly low resolution, the evaluation of small scale features like leads
758 remains a challenge. The use of ever finer resolution models demands the development of
759 new sea ice rheologies suitable for resolving kilometre scale features. Currently it is not clear
760 how significant these errors are for coupled polar prediction and further study is required⁵²
761 Smith et al., (2013).
762

763 Notably the majority of the applications presented have focused on atmospheric phenomena
764 reflecting the maturity of this community and the extensive range of peer-reviewed
765 benchmarks for uncoupled systems from which the impact of coupling can more readily be
766 assessed. Coupled prediction is expected to also have a significant impact on several ocean
767 applications e.g., sonar prediction, search and rescue and hazardous chemical spills. In
768 addition to the fact that the ocean community is less mature it also reflects the paucity of

769 observations available to establish benchmarks for the leading parameters for these
770 applications such as the sonic layer depth and surface currents.

771

772 **Future/outlook and conclusion**

773

774 All groups contributing to this paper have developed research programs specifically targeting
775 a subset of applications that represent their national interest. The modelling systems range
776 from regional to global and the initialisation and data assimilation is uncoupled or weakly
777 coupled. In many cases the research challenges identified are common across these
778 programs indicating significant benefit from a community-based approach to share advances
779 in coupled science and promote international experiments and observation campaigns.
780 Despite the challenges of achieving skilful forecasts from such complex systems the results to
781 date using relatively unsophisticated techniques have already yielded positive results. Most
782 groups are optimistic that coupled prediction will deliver further improvements with
783 continued research and development.

784

785 The Bureau of Meteorology plan to extend the research into East Coast Lows focusing on
786 diagnosing the dynamical response of the atmospheric boundary layer and the impact of
787 coupled modeling. Ensemble Kalman Filter data assimilation has been extensively
788 investigated for regional ocean prediction and preliminary work is being pursued into their
789 extension to coupled DA. With the implementation of a near-global 1/10 degree BLUElink
790 OceanMAPS the impact of these boundary conditions will be assessed for the ACCESS-G
791 NWP system.

792

793 Work at the UK Met Office on coupled prediction at short time-scales is targeted at three main
794 areas: coupled model development; coupled data assimilation development; and UK
795 environmental prediction. Assessment and development of the coupled model HadGEM3 at
796 these time-scales is an on-going area of work; current developments include improvements to
797 the representation of the diurnal cycle of SST, and implementation of a wave model within the
798 coupled model framework. A higher-resolution version of this global system (12km
799 atmosphere and 1/12° ocean) is also being developed in order to assess its performance
800 compared to the uncoupled NWP system. The weakly coupled data assimilation system
801 described in section 2 is being further assessed and developed, and is planned to be
802 implemented as a demonstration operational system in the Met Office's operational suite in
803 2015. Work to develop a coupled modeling framework around the UK to provide
804 environmental predictions is also underway.

805

806 NOAA/NCEP have established a wide range of coupling projects that are underway or
807 planned using the ESMF - NEMS environment, including: Completing ESMF and NUOPC
808 versions of all component models mentioned in section 2; Converting the coupled HWRF
809 hurricane weather model to the NNMB core in NEMS by 2016, transitioning this coupled
810 model from a custom coupling environment to ESMF – NEMS. In this time frame, the HWRF
811 model will be coupled to a full HYCOM ocean model, and coupling with the wave model will
812 begin; A NEMS based prototype for Arctic modelling is intended to be delivered by 2016,
813 tentatively providing a coupled ocean – sea-ice – atmosphere system, possibly also with a
814 wind wave component added; global model coupling using an atmosphere – ocean – sea-ice
815 coupling will be extended for the CFS-v3, and considered for inclusion in the Global
816 Ensemble System (GEFS) and the deterministic Global Forecast System (GFS); a Nearshore
817 Wave Prediction System (NWPS) will be rolled out to the NWS field offices in the coming
818 year⁷⁷ (Van der Westhuysen et al., 2013). Initially this will consist of a wind wave model with
819 input from weather, ocean and coastal circulation (inundation) models. In future upgrades,
820 this model is intended to become a coupled wave-surge model; NOAA has also funded a
821 project to develop the next generation forecast system for the Great Lakes, consisting of a 3D
822 unstructured grid circulation model, an ice model and a wave model. In operations, this
823 coupled lake model is likely to be fully coupled to a regional mesoscale weather model.

824

825 ECMWF will continue developments on coupled forecasting systems. It is planned to include
826 a dynamical sea-ice model in the medium-range, monthly and seasonal forecasting systems,
827 as well as increasing the resolution of the ocean and atmospheric models. For the time
828 being, the initial conditions for the coupled forecasting system will continue being produced by

829 separate atmospheric and ocean/sea-ice assimilation systems. The developments of the
830 coupled assimilation will continue under the CERA system, targeting a fully coupled
831 assimilation system. The computation of several outer iterations in the incremental variational
832 approach of the CERA system has already allowed the observations in one media to impact
833 the analysis of the other media within the same assimilation cycle. It is expected that the
834 combination of variational and ensemble data assimilation methods will improve the
835 formulation of the background error covariances. In the next few years, ECMWF has planned
836 to produce with the CERA system several extended climate coupled reanalyses spanning the
837 20th century and the satellite era in the context of the ERA-CLIM2 project funded by the
838 European Commission.

839
840 Within CONCEPTS the future activities include research and development to address the
841 challenges outlined above, particularly for polar prediction. This will include evaluating and
842 improving the representation of leads, incorporating wave-ice interactions, atmosphere-ice-
843 ocean momentum transfer, constraining sea-ice thickness and sea-ice forecast verification.
844 Regional coupled systems will also be further developed and applied to the Great Lakes and
845 the North Pacific to support high resolution modelling of the Canadian west coast. The
846 development of global coupled modelling systems will continue for applications of medium-
847 and long-range forecasts. In this context there will be an expansion to coupled models for
848 probabilistic forecasting through the Global Ensemble Prediction System at the Canadian
849 Meteorological Centre.

850
851 Based on Mercator's encouraging results, Meteo-France will develop operational systems
852 covering overseas territories using same modelling tools as described in section 2. The
853 Global Mercator-Ocean operational system will be used to initialise the coupled forecast and
854 dedicated ocean configurations could be developed to improve the consistency between the
855 initialisation phase and the forecast one.

856
857 NASA plans have some commonalities with the Canadian CONCEPTS in terms of
858 constraining sea-ice thickness. Along the same lines, they plan to improve near-surface heat
859 transfer over sea-ice by modelling ice skin temperature using CICE thermodynamics. Plans
860 have been outlined to couple the GMAO ocean analysis (iODAS) to its atmospheric analysis
861 system so that the *foundation temperature* (currently OSTIA SST, used by the atmospheric
862 analysis) is replaced with the corresponding temperature analyzed in the ocean model.

863
864 Following the initial concept papers^{5,6} and early workshops in 2008⁷ and 2009⁸ research and
865 development in this field has made significant advances in terms of the sophistication of the
866 modeling systems being implemented as outlined in **Tab 1**, the rigor of the experiments to
867 quantify impacts and the range of applications. The GOV science team initiated the SMRCP
868 task team to promote the use of coupling based on GOV-type ocean prediction systems and
869 to establish a linkage with the atmospheric community. Outlined in this paper there are many
870 examples where coupled systems are now being based on GOV-type ocean prediction
871 systems for short- to medium-range forecasting with demonstrated impacts. The next steps
872 for the SMRCP-TT are to continue to develop linkages with WGNE and other communities
873 involved in coupled forecasting and to jointly develop and promote international initiatives to
874 address the known challenges.

875
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880
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 1125 TABLES
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System	Ocean (Model DA)	Atmos (Model DA)	Wave (Model DA)	Sea-ice (Model DA)	Coupler	Interfacial flux param.	Global/Regional	Target app(s)
BLUElink	OFAM (MOM4p1) BODAS	ACCESS 4DVAR WRF	High-wind param. roughness ,WW3		OASIS4	-	Regional	Tropical Cyclones, Rainfall, East Coast Lows
UK Met Office	NEMO vn3.4, NEMOVAR 3DVar	UM, Hybrid 4DVar	WWIII, no DA	CICE, NEMOVAR 3DVar	OASIS		Global, Local	Global for seamless forecasting: NWP out to seasonal. Local for environmental prediction around UK.
NOAA/NCEP	HYCOM, MOM5	NCEP	WW3	CICE, GFDL sea ice	ESMF plus NUOPC	-	Global/Regional	NWP, Monthly, Seasonal forecast Hurricane prediction
ECWMF	NEMO	IFS	WAM	LIM2	Single Executable	-	Global	NWP, Monthly, Seasonal forecast and climate reanalyses
GOFS COAMPS	HYCOM NCOM NCODA 3DVAR NCOM 4DVAR	NAVGEN COAMPS NAVDAS 4DVAR	WW3 NCODA 2DVAR	CICE NCODA 3DVAR	ESMF plus NUOPC on global scale	(see Figure 3)	Global and Regional	High Impact Weather, Extended Forecasts
CONCEPTS	NEMO	GEM	WW3	CICE	GOSSIP	Coupling by GEM fluxes	Global/regional	Global and regional Canadian NWP, Operational marine support in ice infested waters
Mercator	NEMO. GLORYS 1/4° reanalysis. Forecast : Coupled regional 1/12° configuration	ALADIN 10km. Forecast : AROME 2.5km	NO	NA	OASIS3	ECUME	Regional, Indian Ocean (46E-68E/9°S-22°S)	Tropical cyclones forecast
GEOS-DAS	iODAS (MOM4p1)	GEOS	-	CICE	ESMF	-	Global	Tropical Cyclones; Reanalysis

1128 Tab 1 Overview of the types of systems being employed to examine the impact of coupled
 1129 modelling together with the type of target applications.

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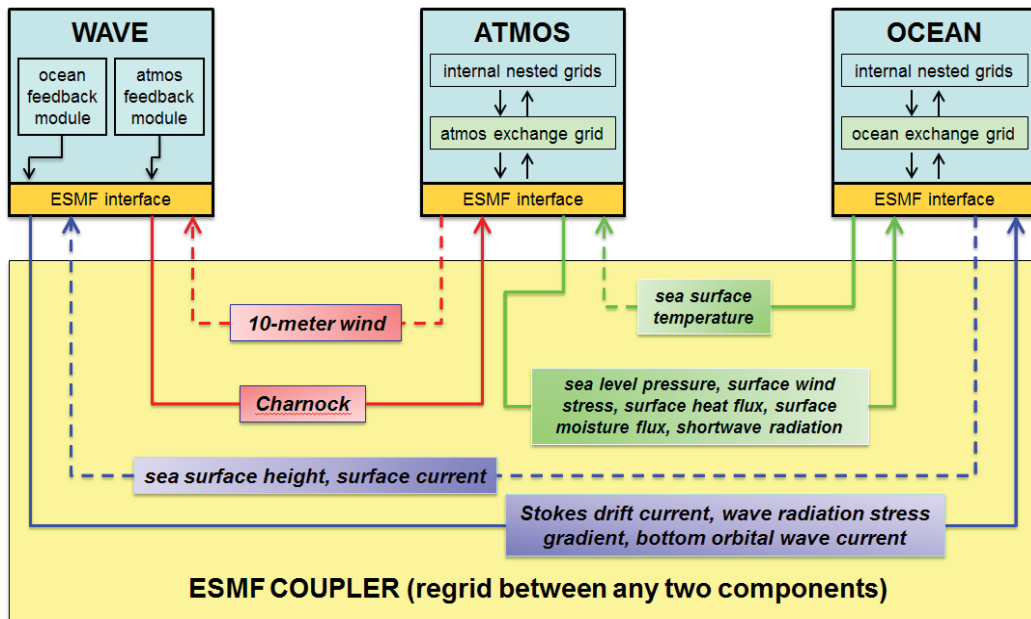
	Model	Observations	DA	Initialisation
Atmosphere	UM ~60km/L85	AIRS, IASI, ATOVS, GPSRO, SSMI, Aircraft, Sondes, Surf-Scat	4D-Var ~120km	Direct
Land	JULES ~60km/4 layers	3D-Var Screen, ASCAT, NESDIS	Nudging Analysis	T/2 Direct
Ocean	NEMO ~25km/L75	In situ SST, T/S profiles, AATSR, AVHRR, AMSRE, Jason 1+2, ENVISAT	3D-Var FGAT	IAU
Sea Ice	CICE ~25km 5 categories	SSMI	3D-Var FGAT	IAU

1132 Tab 2 Model, observations, data assimilation and initialisation methods used in the UK Met
 1133 Office's weakly coupled data assimilation system.

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FIGURES

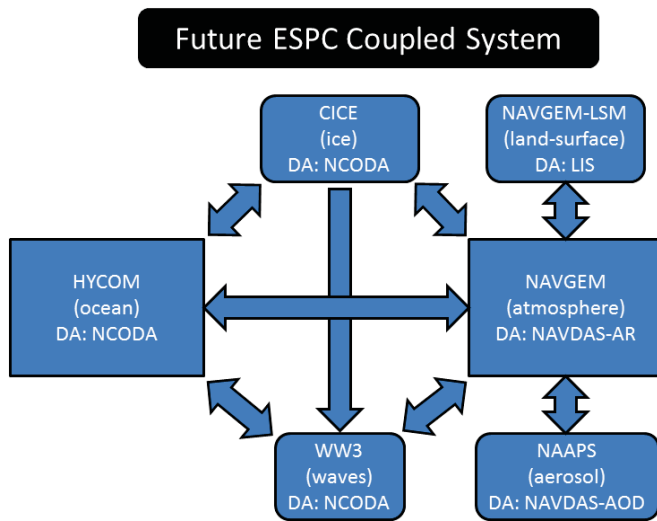


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Fig 1 ESMF coupling framework for the COAMPS air/ocean/wave system showing the variables and exchange parameters passed among the coupled models.

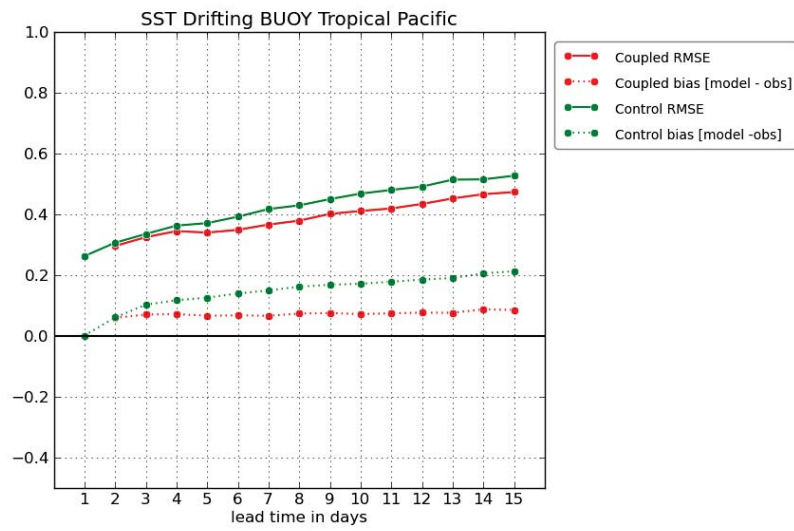
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1146 Fig 2 A schematic of the future ESPC coupled system.

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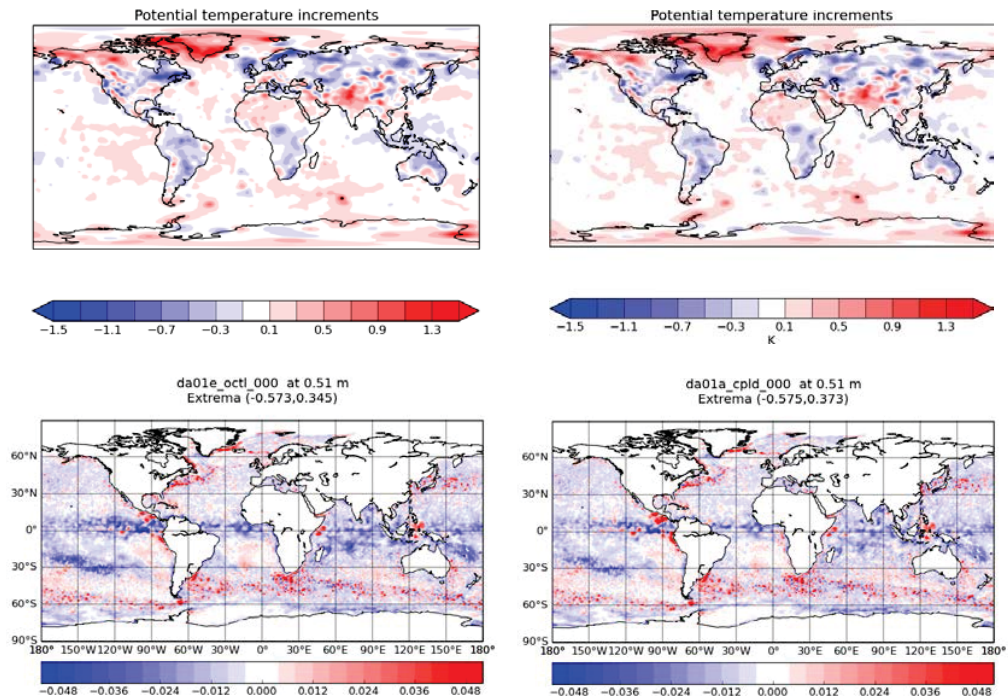
1151 Fig 3 SST (K) observation-minus-forecast RMS (solid) and mean differences (dotted) for a set
1152 of coupled forecasts (red) and ocean-only forecasts (green) in the Tropical Pacific region. The
1153 observations used in this assessment are the drifting buoys.

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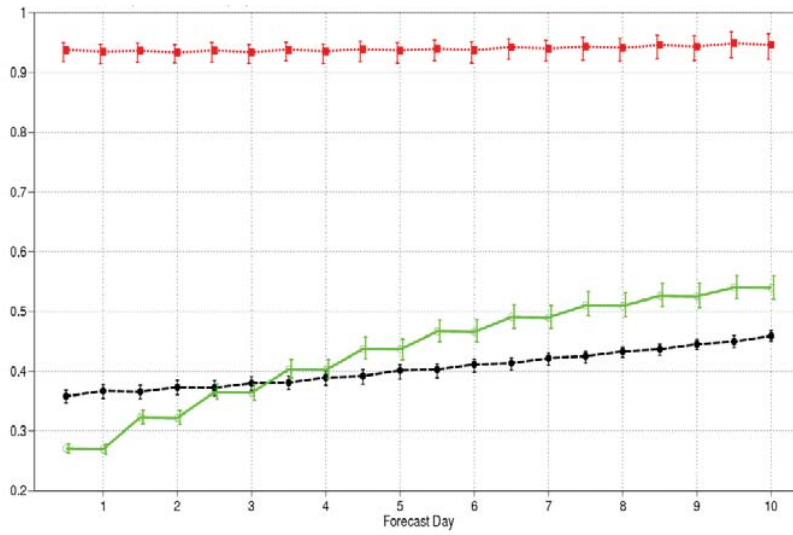
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Fig 4 Monthly average assimilation increments for Dec 2011 for surface air temperature (top row) and sea surface temperature (bottom row) for the un-coupled systems (left column) and the weakly coupled system (right column).

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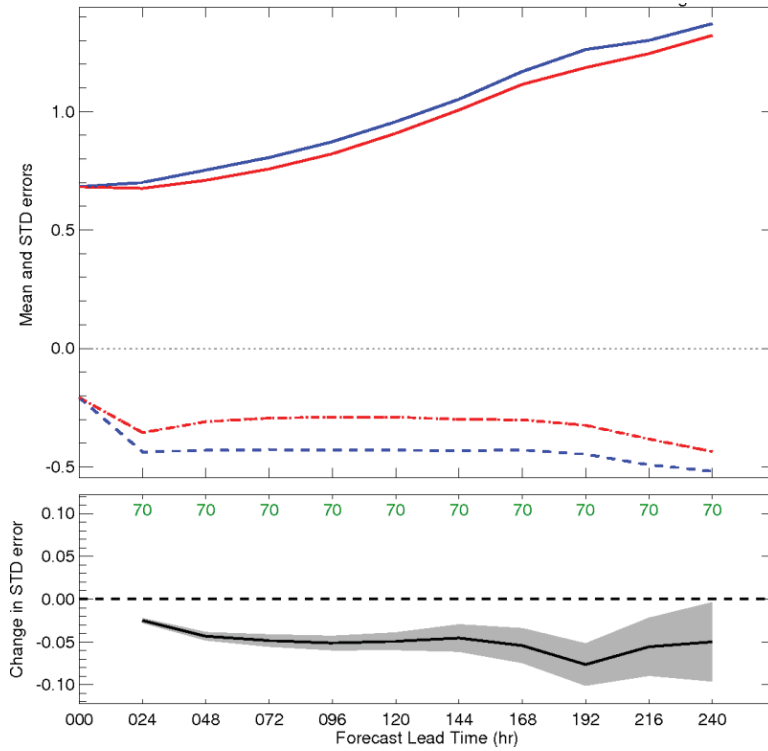
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1166 Fig 5 Root mean square error (RMSE) of the SST forecast in the Tropics from the CERA
1167 system (black) and from the operational-like system (green) for September 2010. The OSTIA
1168 SST analysis is used as reference. The red curve is the RMSE of the SST climatology used
1169 to create the SST anomalies persisted in the forecasts from the operational-like system.

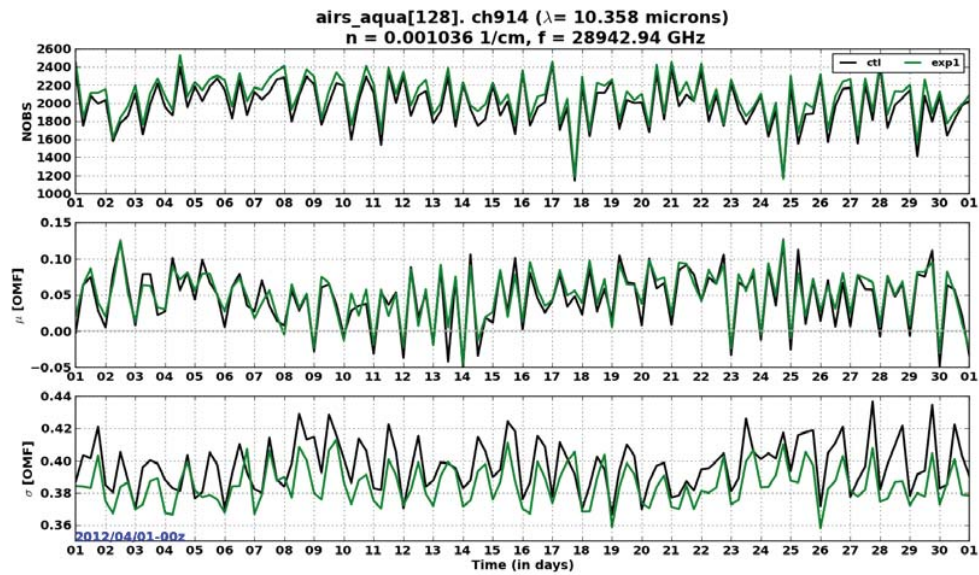
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1174 Fig 6 Evaluation of global coupled forecasts over the tropical Indian Ocean from CMC over
1175 the winter 2011 period. Mean (dashed) and standard deviation (solid) differences between
1176 925 hPa temperature forecasts and ECMWF analyses are shown for uncoupled (blue) and
1177 coupled forecasts (red). The bottom panel indicates the statistical significance of standard
1178 deviation.

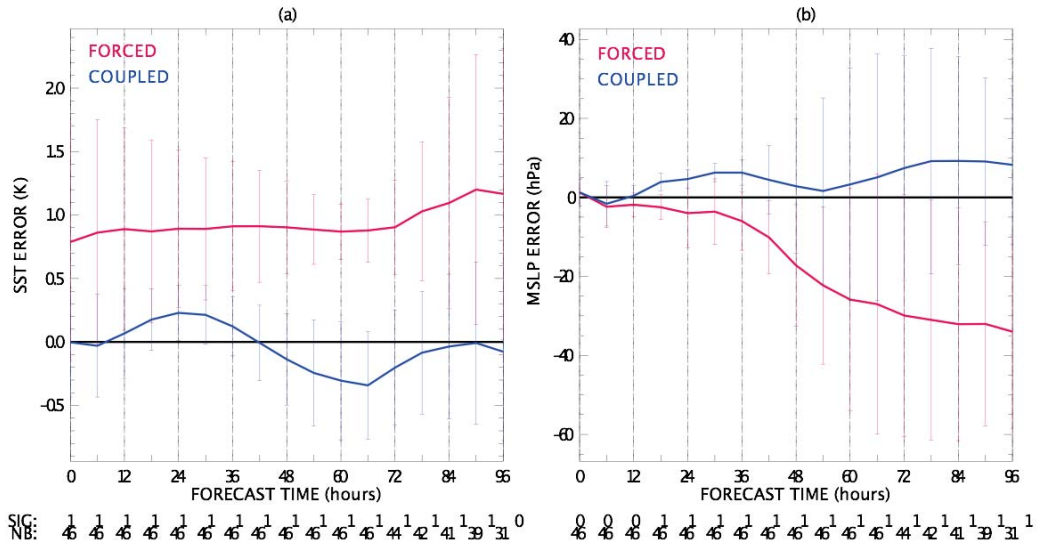
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 1182 Fig 7 Improved assimilation of Brightness Temperature from a surface sensitive (10.35 μm)
 1183 channel of the AIRS instrument on AQUA satellite. We compare an experiment (exp1) that
 1184 had an active interface-layer with a control (ctl) with no interface layer and hence used SST
 1185 boundary condition as skin SST. The three panels plot time series of total number of
 1186 observations assimilated (top panel), global mean of observation-minus-forecast (OMB),
 1187 middle panel, and standard deviation of the OMB (bottom panel). Notice that the analysis
 1188 quality control accepts more observations in exp1, with lesser standard deviation in OMB.
 1189 Similar results are obtained for other 10-12 μm IR channels of AIRS-AQUA, IASI-METOP-
 1190 A, HIRS4-METOP-A, N19 (not shown).

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 1197 Fig 8 (a) SST error ensemble mean evolution (K) as a function of forecast time. (b) Central
 1198 Pressure error ensemble mean evolution (hPa) as a function of forecast time. The total
 1199 number of forecasts and the statistical significance of the difference between the forced and
 1200 coupled ensembles are given for each forecast time below the figure.

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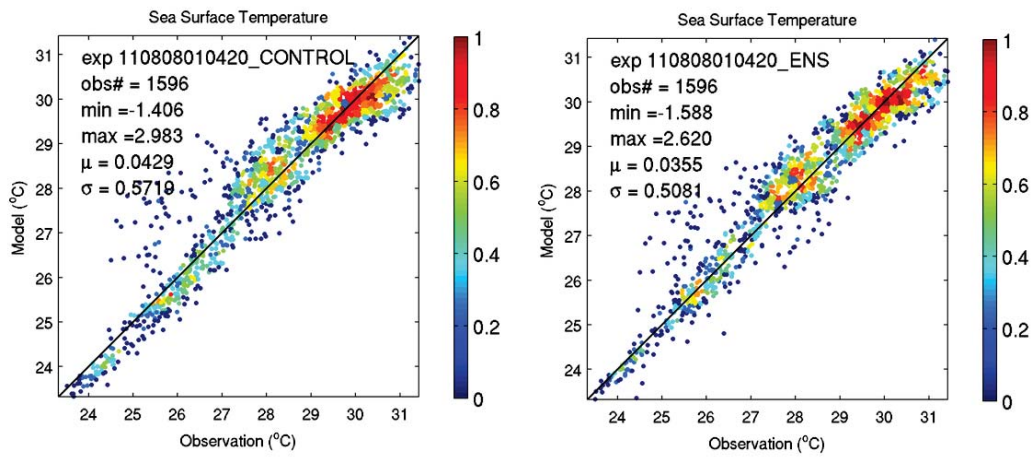
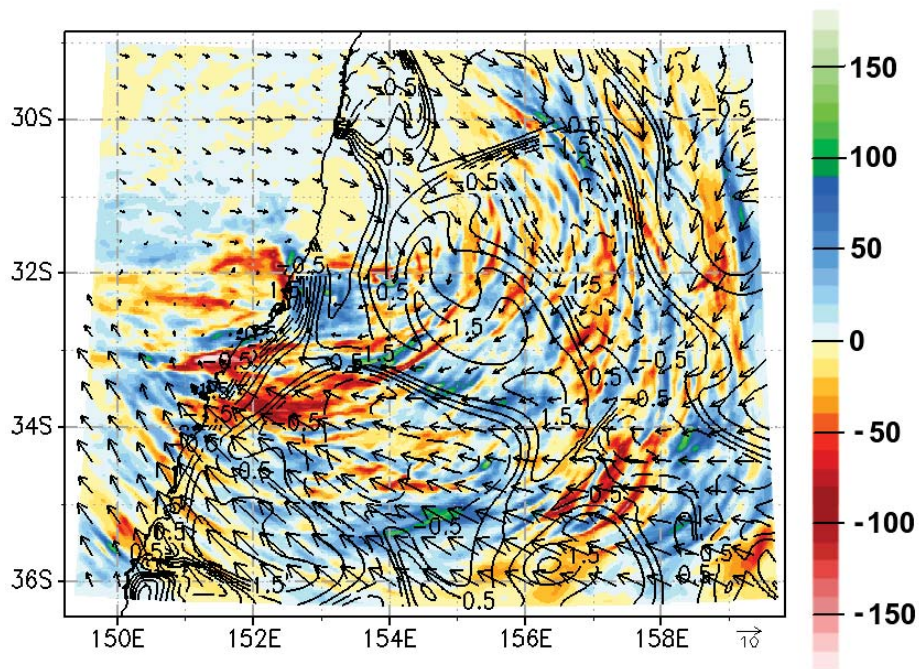


Fig 9 Prediction of SST resulting from the ocean response to tropical cyclone Yasi in the Coral Sea on the 2nd of February 2012 was improved using a coupled ocean-atmosphere ensemble initialisation method. The colours represent a normalised 2D histogram.

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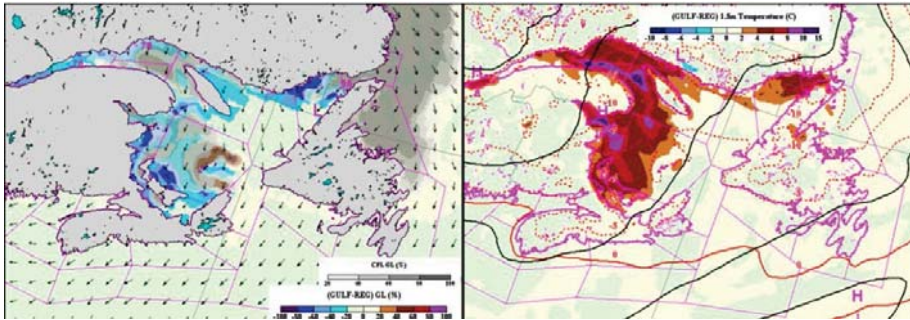
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Fig 10 48 hour (0000 UTC 7 June to 0000 UTC 9 June, 2007) total rainfall differences (colours, mm) (BLUElink - Ctrl). SST differences ($^{\circ}\text{C}$) between the simulations overlaid as contours. In addition the BLUElink simulation average 10 metre wind vectors are overlaid as arrows to indicate the surface flow (m s^{-1} , representative vector in bottom right).

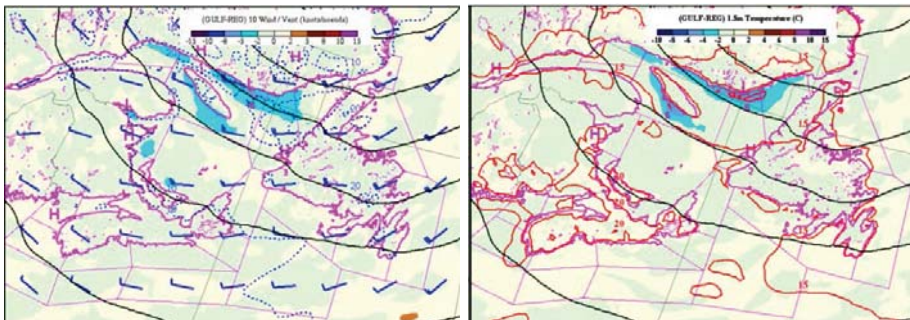
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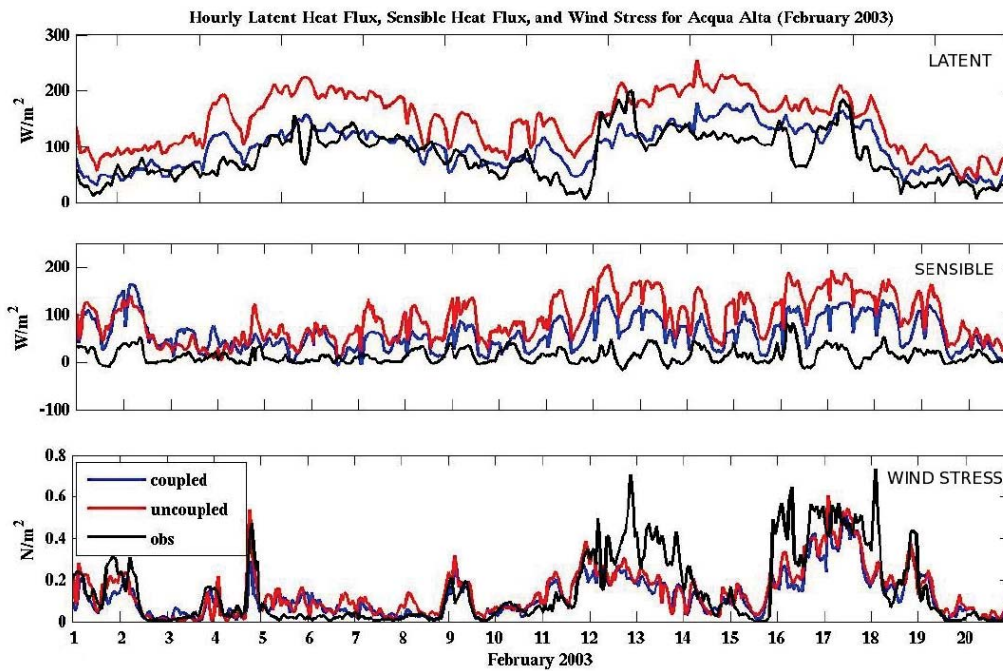


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Fig 11 Differences between coupled and uncoupled model forecasts after 12 hours in the Canadian Gulf of St. Lawrence forecasting system. Top panel is for a winter case (Mar. 10, 2012) with sea ice concentration on the left and 2m temperature on the right showing that rapid ice change can cause surface temperature changes of up to 7-8 °C over the open water. The grey colour shows the ice concentration and the colour scale shows the Coupled minus uncoupled model differences in ice concentration. The bottom panel shows a summer case (Jul. 10, 2012) with 10m winds on the left and 2m temperature on the right showing that coupling induced coastal upwelling can produce surface temperature of several degrees C locally.



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1232 Fig 12 Hourly latent and sensible heat fluxes (W/m^2) and wind stress (N/m^2) for the fully-
1233 coupled COAMPS run and observations at Acqua Alta (Venice). From Allard *et al.* 2010.
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