


1 **Effect of Interhemispheric Field-Aligned Currents**  
2 **on Region-1 Currents**

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7

8 **Key Points**

9 Region1 can be a sum of current flowing from/to solar wind & interhemispheric current.

10 Interhemispheric currents can be related to "double auroral oval".

11 Accounting for interhemispheric currents can help predict Region1 features.

12

13 **Abstract**

14 An asymmetry in ionospheric conductivity between two hemispheres results in the formation of  
15 additional, interhemispheric field-aligned currents flowing between conjugate ionospheres within  
16 two auroral zones. These interhemispheric currents are especially significant during summer-  
17 winter conditions when there is a significant asymmetry in ionospheric conductivity in two  
18 hemispheres. In such conditions, these currents may be comparable in magnitude with the  
19 Region 1 field-aligned currents. In this case, the R1 current is the sum of two FACs: one is going  
20 from/to the solar wind, and another is flowing between conjugate ionospheres. These  
21 interhemispheric currents can also cause the formation of auroras extended along the nightside  
22 polar cap boundary, which may be related to the so-called “double auroral oval”. In this study,  
23 we present the results of analytical and numerical solutions for the interhemispheric currents and  
24 their effect on the Region 1 currents.

25

26 **Index Terms**

27 2721, 2431, 2736, 2753, 2784

28 **Key Words**

29 Field-aligned currents; Magnetosphere-ionosphere coupling; Region 1 current; High latitude  
30 ionosphere; Interhemispheric currents

31

## 32 1. Introduction

33 There are three major systems of Field-Aligned Currents (FACs), transporting energy into  
34 and out from polar ionospheres: the R1 FACs at the polar cap boundary, the Region 2 (R2) FACs  
35 at the auroral zone equatorward boundary (both were extensively studied from observational data  
36 [e.g., *Iijima and Potemra*, 1976, 1978; *Weimer*, 2001; *Christiansen et al.*, 2002; *Papitashvili et*  
37 *al.*, 2002; *Anderson et al.*, 2005] and theoretically [e.g., *Jaggi and Wolf*, 1973; *Wolf*, 1975; *Harel*  
38 *et al.*, 1981; *Lyatsky and Maltsev*, 1983; *Spiro and Wolf*, 1984; *Richmond*, 1992; *Potemra*,  
39 1994]), and the so-called “substorm current wedge” appearing during substorms [e.g.,  
40 *McPherron et al.*, 1973].

41 More recent studies [*Benkevich et al.*, 2000; *Benkevich and Lyatsky*, 2000; *Ohtani et al.*,  
42 2005a; 2005b; *Østgaard et al.*, 2005; *Lyatskaya et al.*, 2008, 2009; etc.] showed that an  
43 important role in the global 3-D current system can be played by the interhemispheric currents  
44 (IHCs). The IHCs redistribute ionospheric currents between two polar ionospheres in the regions  
45 of closed magnetic field lines in case of asymmetry of ionospheric conductivity between two  
46 polar ionospheres, which may happen during unequal illumination of polar ionospheres and other  
47 effects [e.g., *Richmond and Roble*, 1987; *Kozlovsky et al.*, 2003; *Atkinson and Hutchinson*, 1978;  
48 *Rishbeth*, 1997; *Benkevich et al.*, 2000; *Benkevich and Lyatsky*, 2000; *Yamashita and Iyemori*,  
49 2002; *Lyatskaya et al.*, 2008; 2009; *Ohtani et al.*, 2005a, 2005b; *Østgaard et al.*, 2005, and  
50 references therein]. However, since it is difficult to separate the IHCs from other FACs  
51 (especially when they flow in the same region), despite the important role of the IHCs in  
52 dynamics of the global 3-D current system, they have not been sufficiently investigated.

53 The IHCs can be generated on the gradient of ionospheric conductivity (e.g., at the  
54 terminator separating the sunlit and dark ionospheric regions) and at the boundaries of auroral

55 precipitation regions. *Rishbeth* [1997] suggested that IHCs may be “a significant fraction of the  
56 total current, circulating in the ionosphere”, and the results of numerical modeling by *Benkevich*  
57 *et al.* [2000] showed that the IHCs can reach up to half of the R1 currents.

58 The IHCs can also affect the high-latitude ionosphere and upper atmosphere. The Joule  
59 heating by field-aligned and ionospheric currents are the main factor, which affects the  
60 temperature and expansion of the high-latitude ionosphere and upper atmosphere [e.g., *Chun et*  
61 *al.*, 2002; *Baker et al.*, 2004; *Knipp et al.*, 2005; *McHarg et al.*, 2005]. The present research and  
62 modeling results show that the role of IHCs is even more extensive than we suggested in our  
63 previous works.

64 The main purpose of this study is to examine the effect of interhemispheric FACs (IHCs)  
65 (which are flowing between two conjugate ionospheres) on the R1 FACs, which transport the  
66 electric field and energy from the solar wind to the ionosphere. The IHCs are especially  
67 significant during summer-winter conditions when there is significant asymmetry in ionospheric  
68 conductivity in two hemispheres; in these cases, the IHCs may be comparable in magnitude with  
69 and significantly affect the R1 currents. Another goal is to investigate a possible effect of the  
70 IHCs on the auroral events in the vicinity of the polar cap boundary such as the double auroral  
71 oval. These two problems are not investigated yet due to the necessity to solve this problem  
72 simultaneously in two hemispheres with different distributions of ionospheric conductivity.

73

## 74 **2. Interhemispheric Currents near Polar Cap Boundary**

75 For better understanding of the effect of IHCs on the R1 currents, first we consider a  
76 simple case when the polar cap and auroral zone have the shape of a circle and axisymmetric  
77 ring, respectively. The ionospheric conductivity poleward of the auroral equatorward boundary

78 in each hemisphere is assumed to be uniform. We also assume that the electric potential,  $\varphi_1$ ,  
79 coming from the magnetopause, is the same at both polar cap boundaries and varying as

$$80 \quad \varphi_1 = E_0 r \sin \lambda, \quad (1)$$

81 where  $E_0$  is the electric field (which we assume to be homogeneous and the same in both polar  
82 caps), and the angle,  $\lambda$ , is the longitude ( $\lambda=0$  at the midnight meridian). The potential at the  
83 equatorward boundaries of the auroral zones is assumed to be zero due to the shielding effect on  
84 the plasma sheet inner boundary [e.g., *Jaggi and Wolf, 1973*], which is related to auroral zone  
85 equatorward boundaries. In this case, the potential in the auroral zone,  $\varphi_A$ , is a simple function of  
86 the radius,  $r$ , and the angle,  $\lambda$ ,

$$87 \quad \varphi_A = E_0 \frac{r_{PC}^2}{r} \sin \lambda, \quad (2)$$

88 where  $r_{PC}$  is the polar cap radius. The potential and electric field distribution is the same in both  
89 hemispheres. The FACs are derived as  $\nabla \cdot \mathbf{J}_i$ , where  $\mathbf{J}_i$  are the ionospheric currents,  $\mathbf{J}_i = \Sigma \mathbf{E}$ ,  
90 where  $\Sigma$  is the height-integrated ionospheric conductivity, and  $\mathbf{E}$  is the electric field. The  
91 obtained distribution of the FACs is schematically shown in Figure 1a. If conductivity  
92 distributions in two hemispheres are the same, the FACs distributions are also the same.

93 Then we consider a case when the ionospheric conductivity in two auroral zones is  
94 uniform but different in two hemispheres. In this case, we can expect that a part of ionospheric  
95 currents in one hemisphere can go along the highly-conductive magnetic field lines from one  
96 ionosphere to be closed in the opposite ionosphere, which results in the formation of the IHCs.  
97 The distribution of ionospheric currents and IHCs,  $I_{ih}$ , in this case is shown schematically in  
98 Figure 1b.

99 In the case of symmetric ionospheric conductivity in two hemispheres (as in Fig. 1a), the  
 100 R1 currents ( $I_{R1}$ ) on the polar cap boundaries correspond to the traditional R1 currents, equal to  
 101 the  $I_{sw}$  currents going from and to the solar wind (these currents are generated near the  
 102 magnetopause due to solar wind - magnetosphere dynamo effect). In this case,  $I_{R1} = I_{sw}$ . However,  
 103 in the case of different ionospheric conductivities in two hemispheres (as in Fig.1b), the FACs at  
 104 the polar cap boundaries are the sum of two FACs: the traditional R1 currents ( $I_{sw}$ ) going from/to  
 105 the solar wind, and the IHCs ( $I_{ih}$ )

$$106 \quad I_{R1} = I_{sw} + I_{ih}, \quad (3)$$

107 Since both  $I_{sw}$  and  $I_{ih}$  currents in Figure 1b flow at the polar cap boundary, it is difficult to  
 108 separate the  $I_{ih}$  from  $I_{sw}$ . To separate these currents, in the winter auroral zone we included a  
 109 narrow conductive ring attached to the polar cap boundary (see Figure 2) with ionospheric  
 110 conductivity equal to the conductivity in the conjugate region in the opposite summer  
 111 ionosphere. Due to the small width of the ring, it insignificantly affects the magnitude of the  
 112 currents; however, it relocates the  $I_{ih}$  currents to the equatorward boundaries of this ring, which  
 113 allows separating the  $I_{sw}$  and  $I_{ih}$  currents. Note that a similar meridional displacement of  $I_{ih}$   
 114 currents relatively to  $I_{sw}$  at the night side in reality can be caused by the equatorward  $\mathbf{E} \times \mathbf{B}$   
 115 convection drift of magnetospheric plasma across the polar caps, which results in the  
 116 equatorward displacement of the  $I_{ih}$  FACs while they propagate between two hemispheres; this  
 117 effect is known as the Alfvén wings (e.g., *Lyatsky et al.* [2010a]). The resulting equatorward  
 118 displacement,  $\Delta r$ , of  $I_{ih}$  relatively to the polar cap boundary can be estimated on the ionospheric  
 119 level as  $\Delta r \approx V_d \Delta t$ , where  $V_d$  is the equatorward  $\mathbf{E} \times \mathbf{B}$  convection velocity, and  $\Delta t$  is the  
 120 propagation time of the Alfvén wave, transporting FACs between two hemispheres (e.g.,  
 121 *Kivelson and Ridley* [2007]; *Lyatsky et al.* [2010b]). For reasonable values of  $V_d \approx 0.3$  km/s and



122  $\Delta t \approx 5$  min ( $\Delta t \approx l / V_A$  where  $l$  is the length of the field line and  $V_A$  is an average Alfvén velocity  
123 along this field line), we obtain  $\Delta r \approx 100$  km at the ionosphere level, which is sufficient for  
124 separation of these two currents. For simplicity, we assume that the equatorward displacement of  
125 the  $I_{ih}$  currents is the same for all local times; in this case, the problem is similar to that  
126 (considered above) with a narrow conductive ring attached to the polar cap boundaries in the  
127 winter ionosphere.

128 The resulting model is shown in Figure 2. The conductivity of the summer hemisphere  
129 (which is not shown) is high and uniform; the conductivity of the winter hemisphere is low and  
130 uniform everywhere except the narrow ring with conductivity equal to the conductivity in the  
131 conjugate summer ionosphere, which provides the separation between the  $I_{sw}$  and  $I_{ih}$  currents.  
132 This model also allows us to compare the results of analytical solution with numerical simulation  
133 (we remind that the results obtained in this case are related to the night side only).

134 In each hemisphere, there are three given regions: (1) the polar cap with the radius  $r_1$ , (2)  
135 an adjacent narrow ring (shown in white on Fig.2 with the outer radius  $r_2$ , and (3) the remaining  
136 auroral zone with outer radius  $r_3$ . For simplicity, we suggest the Pedersen conductivity,  $\Sigma_P$ , to be  
137 equal to the Hall conductivity,  $\Sigma_H$ , in each of the regions (which is approximately correct in the  
138 case of relatively-low geomagnetic activity). In the entire Southern auroral zone, the conductivity  
139 is uniform,  $\Sigma_P = \Sigma_H = 3S$ , while in the Northern conductivity in the region 1 and 3:  $\Sigma_{P1} = \Sigma_{H1} = \Sigma_{P3}$   
140  $= \Sigma_{H3} = 1S$ , while in the region 2 (where  $r_1 < r < r_2$ )  $\Sigma_{P2} = \Sigma_{H2} = 3S$ . For calculating the potential  
141 distribution outside the polar caps, we solved the problem accounting for different conductivities  
142 in two auroral zones. Inside the polar caps, where the conductivities are different but uniform in  
143 each polar cap, the potential does not depend on conductivity and is derived by Eq. (1).

144 First, we computed the potential distribution, which is the same in both hemispheres due to  
 145 high conductivity along the field lines. The analytical solution for the potential in three  
 146 consecutive axially symmetric regions with accounting for both Pedersen and Hall conductivities  
 147 can be written in the following form [Lyatsky and Maltsev, 1983; Lyatsky et al., 2006]:

$$148 \quad \varphi_2 = E_0 \left[ r_1 \frac{r/r_2 - r_2/r}{r_1/r_2 - r_2/r_1} \sin \lambda + \alpha r_2 \frac{r_1/r - r/r_1}{r_1/r_2 - r_2/r_1} \sin(\lambda - \lambda') \right] \quad (4)$$

$$149 \quad \varphi_3 = E_0 \alpha r_2 \frac{r/r_3 - r_3/r}{r_2/r_3 - r_3/r_2} \sin(\lambda - \lambda') \quad (5)$$

150 where  $E_0$  is the electric field within the polar cap,  $r_1$  is the radius of a polar cap boundary (region  
 151 1 in the Fig. 2),  $r_2$  and  $r_3$  are the radii of the outer boundaries of the narrow ring (region 2) and  
 152 the auroral zone (region 3), respectively, while  $\varphi_2$  and  $\varphi_3$  are potentials at the boundaries of these  
 153 regions. The potential on the polar cap boundary is given by Eq. (1), the potential at the auroral  
 154 zone equatorward boundary is assumed to be zero.

155 The coefficient  $\alpha$  and the angle  $\lambda'$  in Eqs. (4, 5) are the functions of the radii and the  
 156 Pedersen and Hall ionospheric conductivities of these zones:

$$157 \quad \tan \lambda' = \frac{\Sigma_{H2} - \Sigma_{H3}}{\chi_2 \Sigma_{P2} + \chi_3 \Sigma_{P3}}; \quad -\frac{\pi}{2} < \lambda' < \frac{\pi}{2} \quad (6)$$

$$158 \quad \alpha = \chi_1 \Sigma_{P2} \left[ (\chi_2 \Sigma_{P2} + \chi_3 \Sigma_{P3})^2 + (\Sigma_{H2} - \Sigma_{H3})^2 \right]^{-1/2} \quad (7)$$

$$159 \quad \chi_1 = \frac{2r_1^2}{r_2^2 - r_1^2}; \quad \chi_2 = \frac{r_2^2 + r_1^2}{r_2^2 - r_1^2}; \quad \chi_3 = \frac{r_3^2 + r_2^2}{r_3^2 - r_2^2} \quad (8)$$

160 Note that Lyatsky and Maltsev [1983] considered only the case of symmetric distributions  
 161 of ionospheric conductivity in two hemispheres, and they did not account for IHCs. In the case of  
 162 different ionospheric conductivity in two hemispheres and the existence of IHCs, we should  
 163 assume the conductivities in the regions 2 and 3 to be the sums of the related ionospheric

164 conductivities in Northern and Southern auroral zones. The FACs (including the IHCs) are found  
165 from the computed electric field in each of these regions. Then we used our numerical model that  
166 includes IHCs [Benkevich *et al.*, 2000] for the same conductivity distribution. The obtained  
167 results were compared and found very close. For the potential difference across the polar caps of  
168 100 KV, we obtained the following magnitudes of the currents in Northern hemisphere:  $I_{R1}=0.46$   
169 MA,  $I_{ih}=0.19$  MA,  $I_{R2}=0.14$  MA; and in Southern hemisphere:  $I_{R1}=0.66$  MA,  $I_{ih}=0.19$  MA,  
170  $I_{R2}=0.42$  MA. The computed distributions of the ionospheric and field-aligned currents are  
171 shown in Figure 3.

172

### 173 3. Discussion and Conclusion

174 In this study, we investigated the effect of the interhemispheric currents (IHCs) on the R1  
175 FACs, which transport the electric field and energy from the solar wind into the ionosphere. In  
176 the case of asymmetry in ionospheric conductivity between two hemispheres (particularly,  
177 during summer-winter conditions and specific UT intervals), the R1 currents on the polar cap  
178 boundaries are significantly different from the traditional R1 FACs related to symmetric  
179 ionospheric conductivity in two hemispheres. In the case of interhemispheric asymmetry in  
180 ionospheric conductivity, the FACs on the polar cap boundary include also (additionally to the  
181 traditional R1 currents) the IHCs going along the closed magnetic field lines between two  
182 conjugate ionospheres. The magnitude of these IHCs is proportional to the difference in  
183 ionospheric conductivities in two hemispheres on the polar cap boundaries, and during summer-  
184 winter seasons the IHCs can be comparable in magnitude with the R1 FACs. This shows the  
185 important contributions from the IHCs to the global current system.

186           Accounting for the  $\mathbf{E} \times \mathbf{B}$  convection drift of magnetospheric plasma with the frozen-in  
187 magnetic field results in an equatorward displacement of the IHCs on the night side (while these  
188 currents propagate between two hemispheres). This displacement of the IHCs relatively to the  
189 polar cap boundary results in the formation of double-stream FACs near the nightside polar cap  
190 boundary. As a result, the two FACs, separated along the meridian, in summer hemisphere have  
191 the same direction, whereas in the winter hemisphere these currents flow in opposite directions.  
192 The spatial separation of the FACs near the polar cap boundary can partially explain the  
193 separation of FACs near the polar cap boundary, observed with the ST-5 spacecraft [e.g., *Le et*  
194 *al.*, 2008, 2009].

195           In the winter hemisphere, the spatially-separated double-stream FACs flow in opposite  
196 directions; these FACs can be responsible for the formation of the so-called “double auroral  
197 oval” [e.g., *Elphinstone et al.*, 1995; *Lyatsky et al.*, 2001; *Kornilova et al.*, 2006; *Ohtani et al.*,  
198 2012]. Indeed, it is well known [e.g., *Knight*, 1973; *Janhunen and Olsson*, 1998] that the energy  
199 flux of precipitating electrons depends on the direction of FACs: to provide upward-directed  
200 FACs in heated plasma in the convergent magnetic field, it should be a field-aligned electric field  
201 accelerating these electrons. Thus, the upward FACs are associated with fluxes of accelerated  
202 precipitating electrons, which can result in increasing auroral activity. Since the double-stream  
203 FACs in winter hemisphere flow in opposite directions, one of these FACs (upward-directed) can  
204 be responsible for the generation of the auroras and the formation of auroras along the nightside  
205 polar cap boundary, which is the main feature of the double auroral oval. Note that this  
206 explanation for these events is only one of possible effect contributing to the double auroral oval  
207 configuration; other explanations were proposed, e.g., by *Ohtani et al.* [2012] and recently  
208 *Sandholt et al.* [2014].

209            Thus, in this study we showed that any asymmetry in solar luminosity and, consequently,  
210    ionospheric conductivity in two hemispheres results in the generation of the IHCs flowing  
211    between two hemispheres. These IHCs can significantly affect the global 3-D current system in  
212    winter/summer conditions and some UT intervals.

213            The main results of this study can be summarized are follows:

214            (1) Thus, in the case of asymmetry in ionospheric conductivity between two hemispheres,  
215    the R1 currents are the sum of two FACs: the traditional R1 FACs (the  $I_{sw}$  currents) going  
216    from/to the solar wind, and the interhemispheric currents (IHCs). In a sunlit hemisphere, the  
217    IHCs are going in the same direction as the  $I_{sw}$  currents, which results in increasing R1 currents.  
218    In the winter hemisphere, however, the IHCs are directed oppositely to the  $I_{sw}$  currents; as a  
219    result, the magnitude of the R1 currents in dark winter hemisphere can be less than each of these  
220    currents. In the case considered in this study, the IHCs in the winter hemisphere comprise  
221    approximately 40% of the total R1 currents. The strong contribution from the IHCs to the R1  
222    currents explains an important role played by the IHCs in the dynamics of the total 3-D current  
223    system.

224            (2) Although both  $I_{sw}$  currents and IHCs are placed near the polar cap boundary (the  
225    boundary of open-closed field lines), the locations of these two currents do not totally coincide  
226    (at least at the night side) due to an equatorward displacement of the IHCs while they propagate  
227    to the opposite hemisphere. This equatorward displacement of the IHCs with respect to the  $I_{sw}$   
228    currents results in the formation the double-stream FACs near the nightside polar cap boundaries.

229            (3) The formation of double-stream FACs near the nightside winter polar cap boundary  
230    can lead to some interesting results. Since upward FACs are usually associated with fluxes of  
231    accelerated electrons precipitating into the ionosphere (that is explained as a result of the Knight

232 mechanism [e.g., Knight, 1973]), the double-stream FACs over the nightside polar cap boundary  
233 can create a band of precipitating accelerated electrons and auroras stretched out along the polar  
234 cap boundary. In the evening sector, this band can be associated with the upward  $I_{sw}$  FACs (the  
235 traditional R1 FACs) while in the morning sector the upward-directed IHCs, located somewhat  
236 equatorward of the polar cap boundary, can be observed as part of so-called “double auroral  
237 oval” [e.g., Elphinstone et al., 1995; Lyatsky et al., 2001; Kornilova et al., 2006; Ohtani et al.,  
238 2012].

239

#### 240 **Acknowledgment**

241 This study is supported by the National Science Foundation under Award No. ANT-1204019.

242

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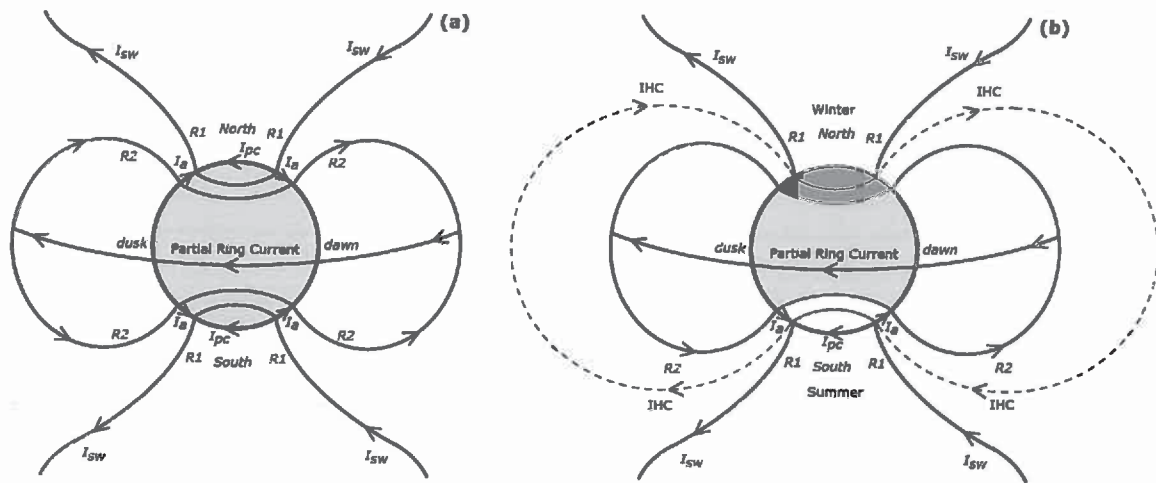
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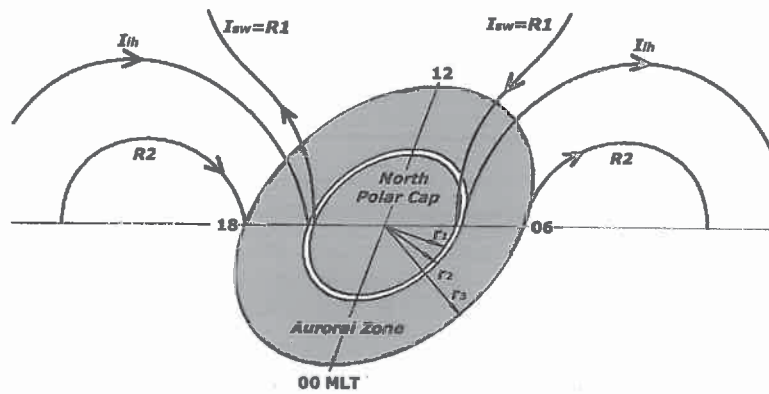
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346 **Figure 1.** A sketch of FACs and ionospheric currents in the dawn-dusk meridional cross-section  
 347 for the cases: (a) when the ionospheric conductivity is the same in both hemispheres and (b)  
 348 when the conductivity in Southern high-latitude ionosphere is higher than that in the Northern  
 349 hemisphere. In the first (a) case, the traditional R1 currents are going on the polar cap boundaries  
 350 from and to the solar wind (these FACs closing through the solar wind we will call the  $I_{sw}$ ), while  
 351 in the case (b) the R1 currents are the sum of the  $I_{sw}$  and IHCs. Shown also are the R2 FACs  
 352 closing the partial Ring Currents in the vicinity of the equatorial plane, and ionospheric currents  
 353 in the polar caps,  $I_{pc}$ , and auroral zones,  $I_a$ . The ionospheric conductivity in Northern auroral  
 354 zone and polar cap in Figure 1b is assumed to be very low.

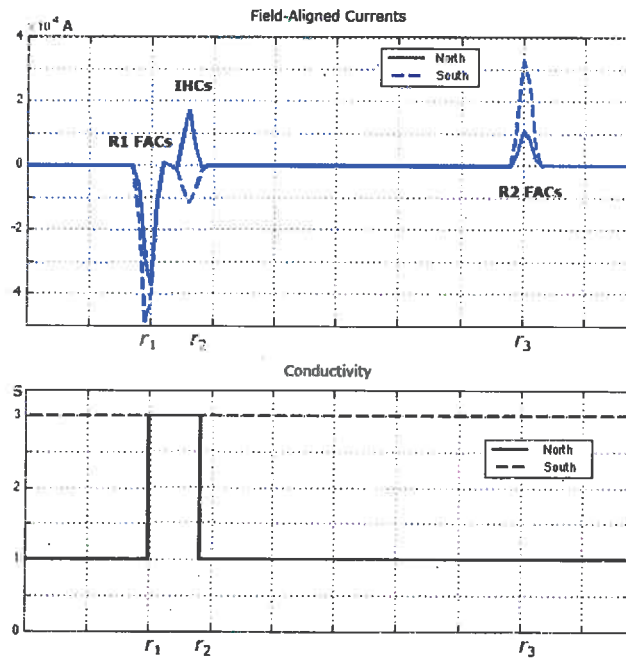
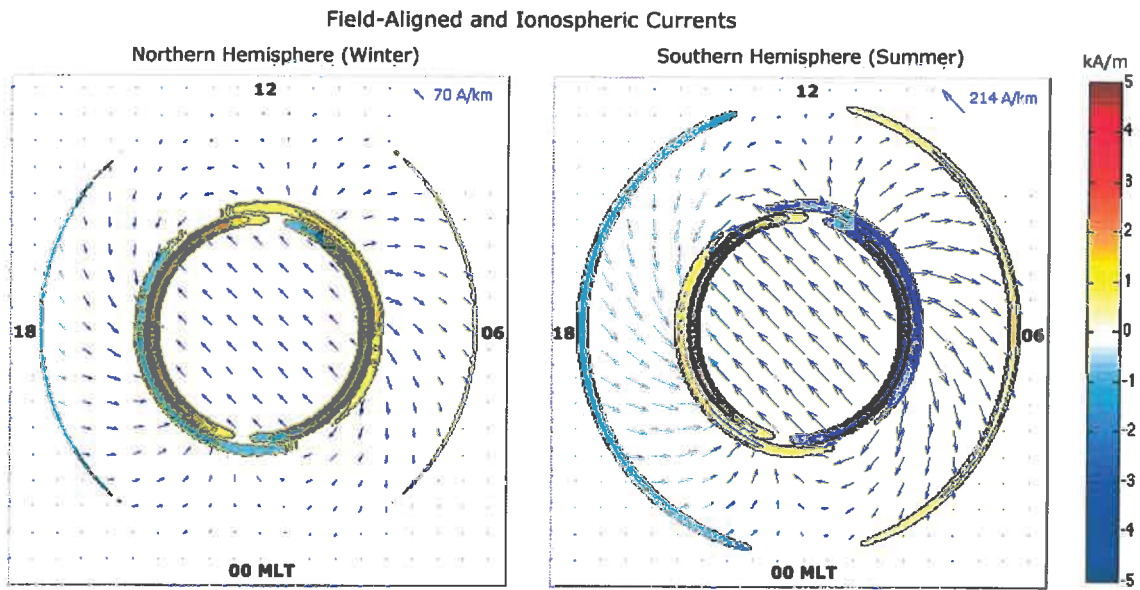
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357 **Figure 2.** A sketch showing the solar wind  $I_{sw}$  currents (going from/to the solar wind) and  
 358 separated interhemispheric  $I_{ih}$  currents (flowing at the outer boundary of the narrow ring of  
 359 enhanced conductivity shown in white) in Northern winter hemisphere. The R2 currents at the  
 360 auroral zone outer boundary are also shown. The ionospheric conductivity in the auroral zone is  
 361 assumed to be much less than that in the opposite summer auroral zone. Note that in the case of  
 362 separated  $I_{sw}$  and  $I_{ih}$  currents, the R1 currents are equal to solar wind currents  $I_{R1}=I_{sw}$ .

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364

365 **Figure 3.** Computed currents in Northern winter hemisphere (top left) and Southern summer  
 366 hemisphere (top right). Ionospheric currents are shown by blue arrows. The magnitude of FACs  
 367 is shown as the contour plots. FACs entering the ionosphere are shown in blue while going out  
 368 from in red and yellow. The FACs currents going from/to the solar wind at the polar cap

369 boundaries are shown as the  $I_{sw}$  currents, the FACs at the outer boundaries of the narrow rings  
370 slightly equatorward of the polar caps are the interhemispheric currents ( $I_{ih}$ ); the FACs at the  
371 outer boundaries of the auroral zones are the R2 currents. Note that  $I_{ih}$  currents have the same  
372 direction as the  $I_{sw}$  currents in summer hemisphere (the top right panel) but the opposite  
373 directions in winter hemisphere (the top left panel). The panels below show the meridional plots  
374 of the relative locations of the FACs, integrated within  $20^\circ$  of longitude along the dawn meridian  
375 (06 MLT), and the conductivity profile (lower panel) in the same meridian. Currents and  
376 conductivity in the Northern hemisphere are shown in solid, while in Southern in dashed lines.