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Abstract The development of the new Italian strong-motion database ITACA (ITalian AC-celerometric Archive, <http://itaca.mi.ingv.it>) is in progress under the sponsorship of the National Department of Civil Protection (DPC) within Project S4, in the framework of DPC-INGV 2007–2009 research agreement. This work started from the alpha version of ITACA [8], where 2,182 3-component records from 1,004 earthquakes, mainly recorded by the National Accelerometric Network, RAN, operated by DPC, were processed and included in the database. Earthquake metadata, recording station information and reports on the available geological-geophysical information of 452 recording sites, corresponding to about 70% of the total, were also included. Subsequently, ITACA has been updated and will reach its final stage by the end of Project S4, around mid-2010, with additional features, improved information about recording stations, and updated records, including the Mw6.3 L'Aquila earthquake. All records were re-processed with respect to the alpha version [9], with a special care to preserve information about late-triggered events and to ensure compatibility of corrected records, i.e., velocity and displacement traces obtained by the first and second integral of the corrected acceleration should not be affected by unrealistic trends. After a short introduction of ITACA and its most relevant features and statistics, this paper mainly deals with the newly adopted processing scheme, with reference to the problems encountered and the solutions that have been devised.

Chapter 8

Record Processing in ITACA, the New Italian Strong-Motion Database

R. Paolucci, F. Pacor, R. Puglia, G. Ameri, C. Cauzzi, and M. Massa

Abstract The development of the new Italian strong-motion database ITACA (Italian AC-celerometric Archive, <http://itaca.mi.ingv.it>) is in progress under the sponsorship of the National Department of Civil Protection (DPC) within Project S4, in the framework of DPC-INGV 2007–2009 research agreement. This work started from the alpha version of ITACA [8], where 2,182 3-component records from 1,004 earthquakes, mainly recorded by the National Accelerometric Network, RAN, operated by DPC, were processed and included in the database. Earthquake metadata, recording station information and reports on the available geological-geophysical information of 452 recording sites, corresponding to about 70% of the total, were also included. Subsequently, ITACA has been updated and will reach its final stage by the end of Project S4, around mid-2010, with additional features, improved information about recording stations, and updated records, including the Mw6.3 L’Aquila earthquake. All records were re-processed with respect to the alpha version [9], with a special care to preserve information about late-triggered events and to ensure compatibility of corrected records, i.e., velocity and displacement traces obtained by the first and second integral of the corrected acceleration should not be affected by unrealistic trends. After a short introduction of ITACA and its most relevant features and statistics, this paper mainly deals with the newly adopted processing scheme, with reference to the problems encountered and the solutions that have been devised.

8.1 Introduction

The development of the new Italian strong-motion database ITACA (Italian AC-celerometric Archive, <http://itaca.mi.ingv.it>) is in progress under the sponsorship of the Italian Department of Civil Protection (DPC) within Project S4, in the framework of DPC-INGV 2007–2009 research agreement. This Project has

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46 continued the activity originally developed by Project S6, within the previous
47 2004–2006 DPC-INGV agreement, in which the alpha version of ITACA was
48 originally developed [8].

49 The main goal of the S6 and S4 Projects has been to organize into a compre-
50 hensive, informative and reliable database (and related webtools) the wealth of
51 strong-motion records, obtained in Italy during the seismic events occurred starting
52 from the Ancona earthquake sequence in 1972, up to the L’Aquila 2009 sequence.

53 The beta version of ITACA, which will reach its final stage by the end of
54 the project, around mid-2010, will include several improvements and additional
55 features, namely: – strong motion records from other local and/or temporary net-
56 works, and from recent seismic events, in primis the April 6 2009 M_w 6.3 L’Aquila
57 earthquake and its main aftershocks; – updated reports, with an improved for-
58 mat, on the available geological/geophysical information of recording stations,
59 including average HVSR from microtremors and earthquakes where available; –
60 identification of stations and records showing distinctive features, either due to geo-
61 logical/topographic irregularities or due to seismic source effects; – online tools for
62 selection of spectrum-compatible records.

63 To date, ITACA contains 2,550 three-component waveforms: 2,293 of them were
64 recorded during 1,002 earthquakes with a maximum moment magnitude of 6.9
65 (1980 Irpinia earthquake) in the period range 1972–2004, while the rest comes from
66 the M_w 5.4 2008 Parma (Northern Italy) and from the M_w 6.3 2009 L’Aquila (Central
67 Italy) earthquakes and related $M_w > 4$ aftershocks. Records obtained in 2005–2007
68 will be included by the end of the Project.

69 The recordings mainly come from the National Accelerometric Network (RAN,
70 Rete Accelerometrica Nazionale), now operated by DPC. RAN presently consists
71 of 334 free-field digital stations and 84 analogue stations, the replacement of which
72 with digital instruments is currently in progress. The goal is to achieve a final
73 configuration of more than 500 digital stations installed throughout the Italian
74 territory, with an average inter-station spatial distance of about 20–30 km in the
75 most seismically active regions of Italy (A. Gorini, personal communication, 2010).
76 Further records are provided by the Strong Motion Network of Northern Italy (*Rete*
77 *Accelerometrica dell’Italia Settentrionale*, RAIS <http://rais.mi.ingv.it>, [2]), consist-
78 ing of digital instruments, installed around the Garda lake area, and by sparse
79 stations (analogue and digital) operated by ENEA (Ente per le Nuove tecnologie,
80 l’Energia e l’Ambiente (*Italian energy and environment organization*)), over the
81 time span 1972–2004. In addition to these, waveforms recorded during the L’Aquila
82 seismic sequence by the accelerometer installed on the very broad band AQU station
83 (<http://mednet.rm.ingv.it>) are also present.

84 All ITACA records were re-processed with respect to the alpha version [9], with a
85 special care to preserve information about late-triggered events and to ensure com-
86 patibility of corrected records, i.e., velocity and displacement traces obtained by
87 the first and second integral of the corrected acceleration should not be affected by
88 unrealistic trends.

89 This paper mainly deals with the newly adopted processing scheme, with
90 reference to the problems encountered and the solutions that have been devised.

8.2 Characteristics of the ITACA Dataset

Figures 8.1 and 8.2 summarize the main characteristics of the ITACA dataset in terms of focal parameters and distance ranges. As shown in Fig. 8.1, magnitude (either M_w or M_L) ranges from 2 to 6.9 with the best sampled distance interval from 5 to 100 km. The epicentral distance (R_{epi}), for $M < 5.5$ events, and the Joyner-Boore distance (R_{jb}) for stronger earthquakes are considered, based on the fault geometry data available in the DISS database [3, 6]. Nine 3-component records with epicentral distance $R_{epi} \leq 10$ km are available in the range $5.9 \leq M_w \leq 6.3$ (5 from the L'Aquila earthquake, 3 from the Friuli aftershocks of September 1976, and 1 from the Umbria-Marche September 1997 mainshock). The strongest events in ITACA, i.e., the M_w 6.4 May 6 1976 Friuli and the M_w 6.9 November 23 1980 Irpinia earthquakes, were recorded at $R_{epi} > 10$ km.

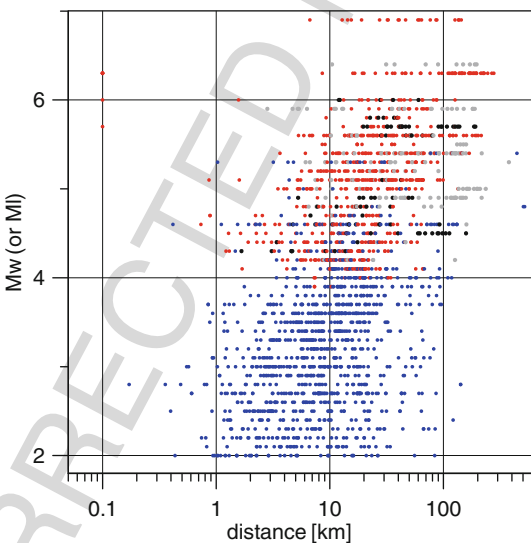


Fig. 8.1 Magnitude vs. distance (either Joyner-Boore for $M > 5.5$ or epicentral distance otherwise) distributions for the ITACA dataset. The records are grouped by focal mechanism (blue: normal; gray: reverse, black: strike, light blue: unknown)

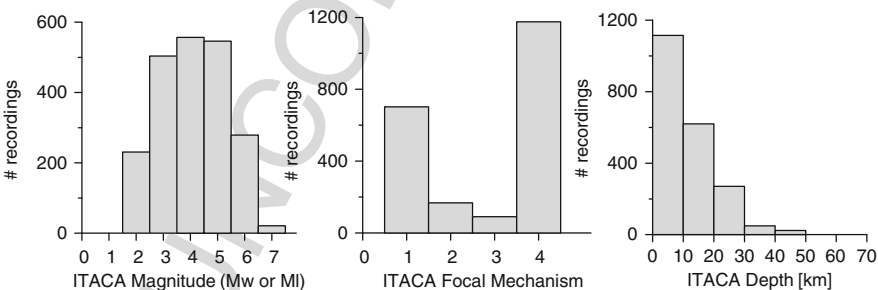


Fig. 8.2 Distribution of ITACA records plotted as a function (a) Magnitude; (b) focal mechanism (1: normal; 2: reverse, 3: strike, 4: unknown); (c) focal depth

136 Distributions of records as a function of magnitude, focal depth and focal
137 mechanism are plotted in Fig. 8.2. Most events with magnitude less than 4 have
138 unknown focal mechanisms. For the strongest earthquakes, the focal mechanisms
139 were assigned following the classification of Zoback [11], as described in Luzi et al.
140 [8]. Among the strongest earthquakes, most of them were caused by normal faults in
141 Central and Southern Apennines (namely, the Irpinia, Umbria-Marche and L'Aquila
142 earthquakes), with focal depths less than 10 km. Earthquakes in NE Italy, includ-
143 ing the Friuli seismic sequence, and in the Northern Apennines, are deeper and
144 mainly characterized by a compressional tectonic regime. Finally, strike-slip events
145 mainly occurred in Southern Italy, including the M_w 5.7 October 31 2002 Molise
146 earthquake, at focal depths generally between 20 and 30 km.

147 As a whole, waveforms collected in ITACA were recorded by 665 strong-motion
148 stations. Among these stations, 287 are presently not in operation, since they were
149 either part of temporary networks or equipped with old analogue instruments, which
150 were removed.

151 Station metadata were included in ITACA after collection of pre-existing data
152 and field investigations performed during the S6 project and the ongoing S4 project.
153 Geophysical and geotechnical information at the ITACA recording stations is avail-
154 able at different levels: from the simple geological description up to a complete
155 geotechnical site characterization, including stratigraphic logs, V_S (S-wave) and
156 V_P (P-wave) velocity profiles, dispersion curves, fundamental frequencies, site
157 response functions, noise measurements etc. For most sites, based on strong and
158 weak motions and noise measurements, it was possible to apply spectral ratio tech-
159 niques, mainly Horizontal to Vertical (HVSr) and, in few cases, Standard Spectral
160 Ratio (SSR), when a nearby reference station was available.

161 All ITACA stations are classified according to the EC8 [5] site classes, i.e., class
162 A: V_{S30} 800 m/s, class B: $V_{S30}=360-800$ m/s, class C: $V_{S30} = 180-360$ m/s, class
163 D: $V_{S30} < 180$ m/s and class E: 5–20 m of C – or D-type alluvium underlain by stiffer
164 material with $V_S > 800$ m/s. However, since V_{S30} will be available only for about
165 100 stations at the end of Project S4, it was decided to denote by a star (*) those
166 stations that were classified only based on the geological/geophysical information
167 available (S4 project – <http://esse4.mi.ingv.it> – Deliverable D4, 2009), but not on
168 a direct measurement of V_{S30} . Among stations with V_{S30} available at present, 8%
169 were classified as A, 42% B, 27% C, 2% D and 21% E.

170 The distributions of peak ground acceleration (*PGA*) and velocity (*PGV*) values
171 reflect the event-distance distribution (Fig. 8.3). With the exception of the L'Aquila
172 seismic sequence, most records with largest peaks are from analogue instruments.
173 A total of 360 waveforms (about 20% of the total) have $PGA > 50$ cm/s² while
174 160 recordings (about 10% of the total) have $PGV > 5$ cm/s. In both cases the
175 maximum of two horizontal components was considered. *PGA* values exceeding
176 400 cm/s² were recorded at stations in the epicentral area, during the L'Aquila,
177 Umbria Marche and Friuli earthquakes. The 1980 Irpinia earthquake generated the
178 largest *PGV* (70 cm/s) at Sturmo station (STR).

8 Record Processing in ITACA, the New Italian Strong-Motion Database

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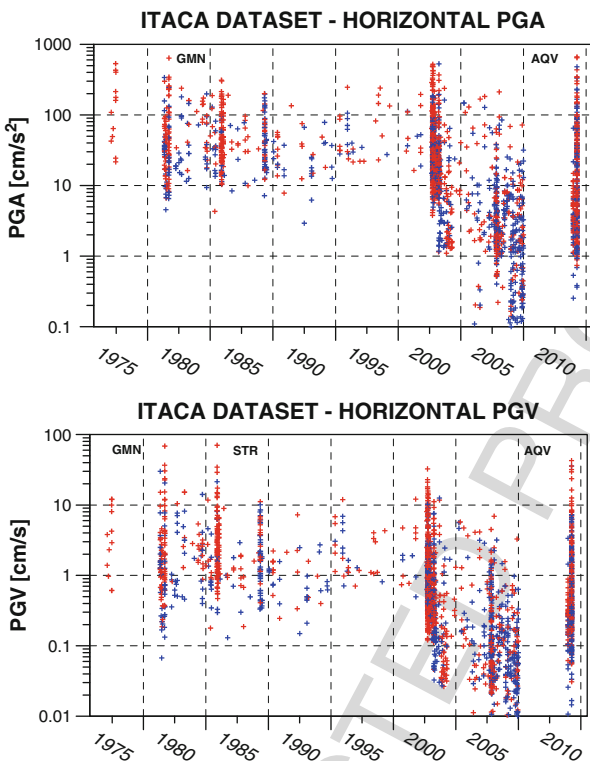


Fig. 8.3 Distribution with time of maximum horizontal PGA (top) and PGV (bottom) of ITACA records. The *blue* symbols represent values recorded at rock site (class A); the *red* ones at all the other site classes

8.3 Record Processing

The problem of defining a procedure to process acceleration time series recorded by analogue and digital instruments has been tackled since the first appearance of ITACA database. The proposed correction scheme involves the processing of analogue and digital records in different ways, with particular attention to the treatment of analogue data, as most of the strongest Italian events were recorded by analogue instruments.

The main steps of the processing procedure are described in Massa et al. [9] and involve: mean removal, baseline correction, instrument correction (for analogue data), band-pass filtering (with acausal filters) and integration of the processed acceleration in order to obtain velocity and displacement waveforms. This scheme was applied to each individual record, with the aim of preserving the low frequency content of the signals. Although the ITACA waveforms were treated by following the worldwide accepted techniques that aim to remove low and high frequency noise, the compatibility among acceleration, velocity and displacement was not

226 guaranteed in the alpha version of ITACA. Within the revision activities to pub-
 227 lish the beta version of the database, several points have been addressed, dealing
 228 with the quality and reliability of corrected records, namely:

- 229
- 230 – to check the accuracy and reliability of the frequency range of the corrected
- 231 records and compare them with the corresponding records available in other
- 232 international databases, such as PEER and European Strong Motion Database
- 233 (ESMDB);
- 234 – to ensure the compatibility of corrected accelerograms, so that no further correc-
- 235 tion is required to obtain the velocity and displacement traces by single and double
- 236 integration, respectively;
- 237 – to identify the late-triggered records, typically on the S-phase, that form a large
- 238 portion of analogue records from small-to-medium magnitude earthquakes.

239 Based on the above discussions a novel procedure for processing the ITACA
 240 strong-motion records has been devised, with the objectives of providing a rational
 241 solution to the previous problems and of being robust as well as reliable enough to be
 242 effectively used for reprocessing of all the ITACA records, including the most recent
 243 ones from the Parma (December 2008) and L'Aquila (April 2009) earthquakes.
 244

245 **8.3.1 ITACA Processing Scheme**

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247 The diagram block of the new procedure is illustrated in Fig. 8.4. Its basic steps are
 248 the followings:
 249

- 250
- 251 – baseline correction (constant de-trending);
- 252 – application of a cosine taper, based on the visual inspection of the record (typically
- 253 between 2 and 5% of the total record length); records identified as late-triggered
- 254 are not tapered;
- 255 – visual inspection of the Fourier spectrum to select the band-pass frequency range;
- 256 whenever feasible, the same range is selected for the 3-components;
- 257 – application of a 2nd order acausal frequency-domain Butterworth filter to the
- 258 acceleration time-series;
- 259 – double-integration to obtain displacement time series;
- 260 – linear de-trending of displacement;
- 261 – double-differentiation to get the corrected acceleration.

262 Note that zero-pads are added at the beginning and end of the signal before the
 263 acausal filter is applied [4]. However, this may pose several problems when using
 264 the corrected accelerograms, especially for engineering applications. As a matter of
 265 fact, very long initial zero-pads would most likely be removed by those end-users
 266 who are interested in using the waveforms for time-consuming non-linear time his-
 267 tory analyses of dynamic response of soils and structures. As a consequence, the
 268 numerical simulations may start from non-zero initial conditions and present spuri-
 269 ous trends in terms of input velocity and displacement, with the risk to compromise
 270 the reliability of results. To overcome this problem, it was decided to re-establish
 after filtering the original initial time-scale, whenever feasible. This is done by

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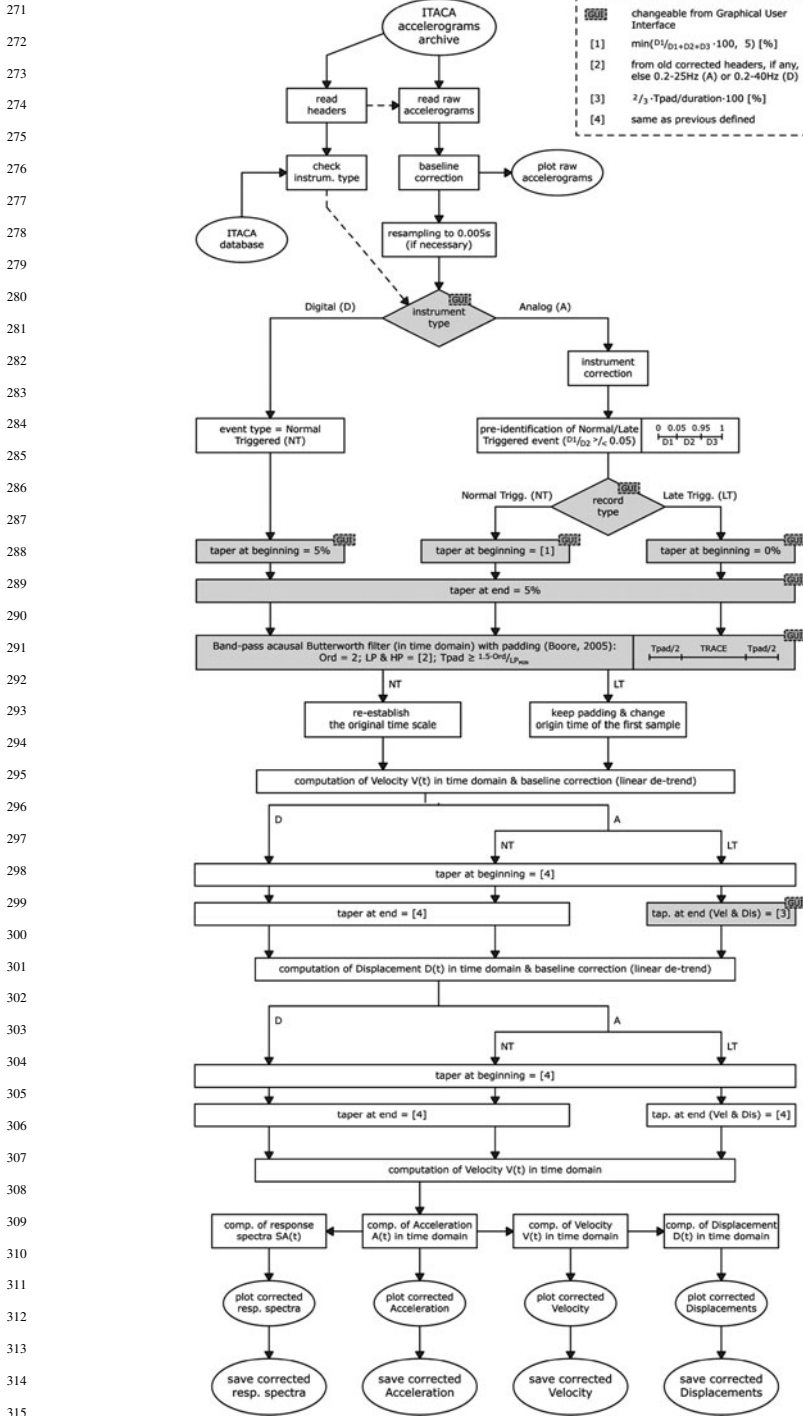


Fig. 8.4 ITACA data processing scheme

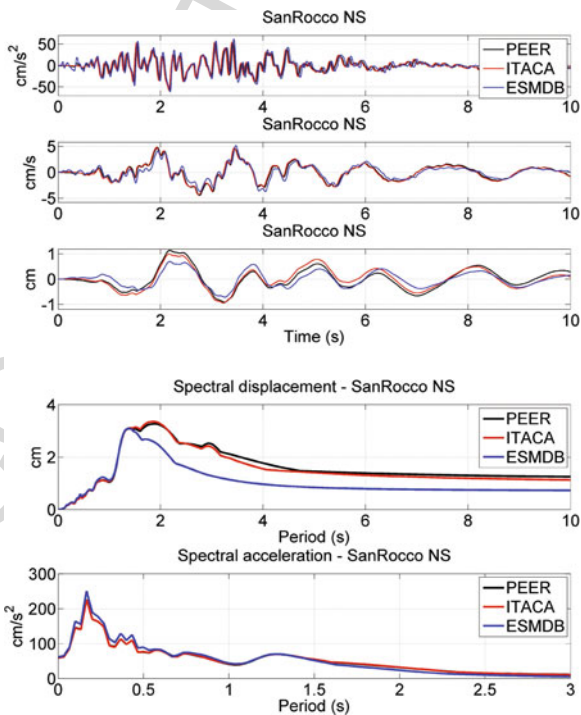
316 removing the zero-pads and by ensuring that the subsequent tapering of velocity
 317 and displacement will produce time histories starting from zero initial conditions.
 318 Otherwise, if tapering is not sufficient for this purpose, the initial zero-pads are
 319 retained. For late-triggered records, no taper is applied and zero-pads are kept.

320 The linear de-trending of displacement traces, and subsequent differentiation to
 321 obtain the corrected accelerations, ensures the compatibility of all corrected records,
 322 in the sense that the integration and double integration of the corrected accelero-
 323 grams produce velocity and displacement time series with zero initial conditions
 324 and without unrealistic trends.

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 326 **8.3.2 Comparison with Records from Other Sources**
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328 Three sources have been considered that contain the most important records from
 329 Italy, namely ITACA itself, the European Strong Motion Database (ESMDB,
 330 <http://www.isesd.cv.ic.ac.uk/ESD/frameset.htm>) and the PEER Strong motion
 331 database (PEER, <http://peer.berkeley.edu/smcat>). Only for L’Aquila 2009 earth-
 332 quake the source external to ITACA was the CESMD (Center for Engineering
 333 Strong Motion Data, <http://www.strongmotioncenter.org>).

334 To clarify the major reasons of difference among records from various sources,
 335 Fig. 8.5 shows a comparison for the San Rocco record, NS component, of the
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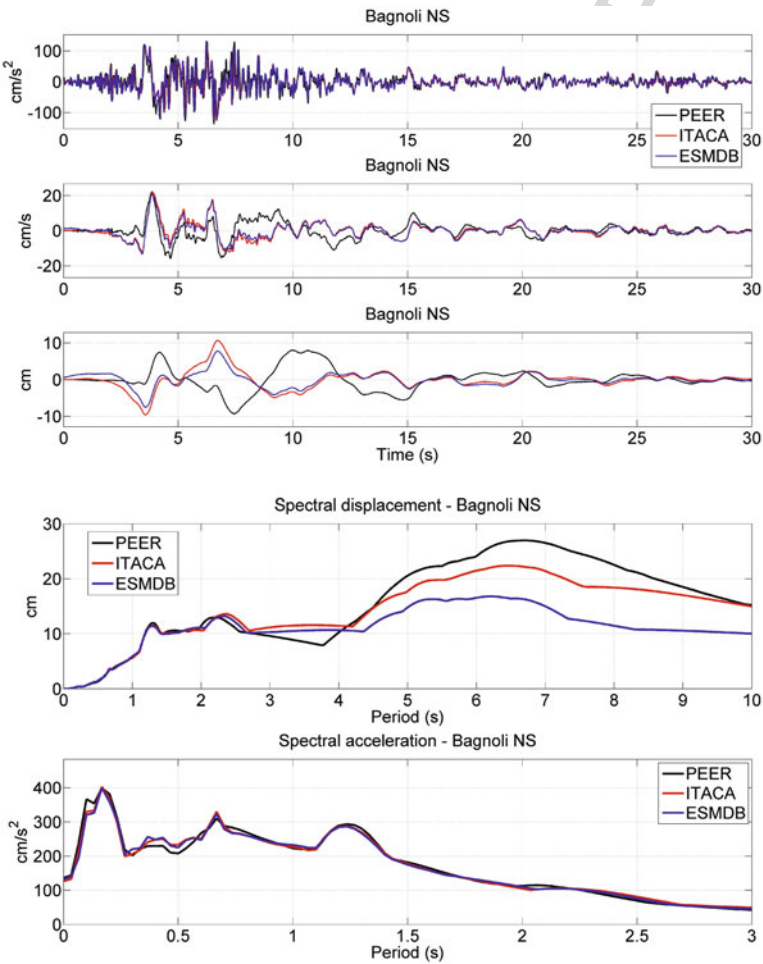
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 350 **Fig. 8.5** Comparison of San
 351 Rocco corrected record, NS
 352 component, from the M_w 6.1
 353 Friuli aftershock of Sep 15
 354 1976, 03:15 GMT, as
 355 available from ITACA,
 356 ESMDB and PEER
 357 databases. From *top* to
 358 *bottom*: corrected
 359 acceleration, velocity,
 360 displacement, spectral
 displacement and spectral
 acceleration

8 Record Processing in ITACA, the New Italian Strong-Motion Database

361 M_w 6.1 Friuli aftershock of September 15 1976 (03:15 GMT). In this case, PEER
 362 and ITACA records are similar, with similar high-pass (HP) filter corners (0.1 and
 363 0.15 Hz, respectively). None of these records have zero-pads at the beginning, but
 364 the tapering allows one to obtain compatible velocity and displacement time series.

365 On the other hand, the ESMDB record is not tapered, it is HP filtered at 0.45 Hz
 366 and keeps zero-pads at the beginning (not shown in the plot). If zero-pads were
 367 removed to re-establish the original time scale, the displacement would be affected
 368 by a trend.

369 As a second example, Fig. 8.6 illustrates the comparison of the corrected Bagnoli
 370 NS record of the M_w 6.9 Irpinia earthquake in 1980. In this case, the HP corner
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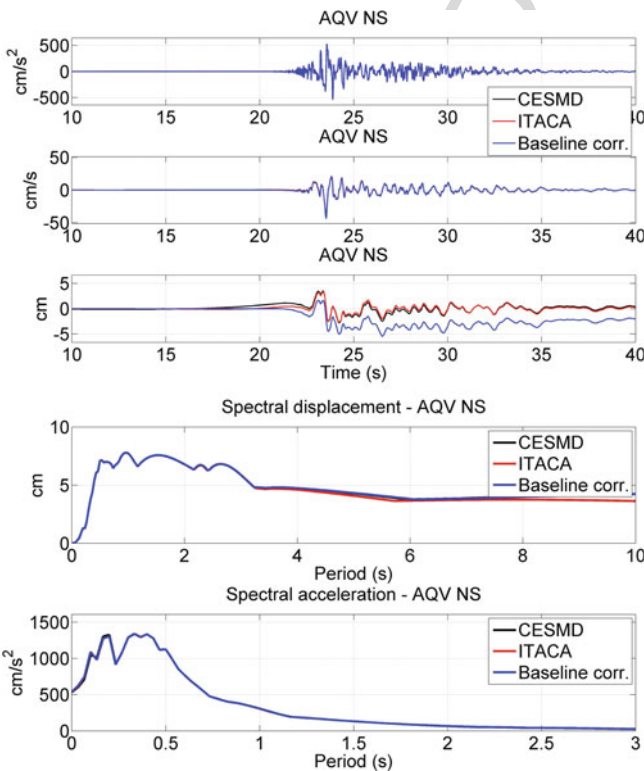


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 404 **Fig. 8.6** As Fig. 8.5, for the NS component of Bagnoli corrected record, from the M_w 6.9 Irpinia
 405 earthquake, 1980

406 frequency of corrected records are similar (0.1 Hz for both ITACA and PEER
 407 and 0.15 Hz for the ESMDB), but the PEER velocity and displacement traces are
 408 different from the other two.

409 Such a difference could be due to causal filtering of the record, affecting the phase
 410 of the signal. ITACA and ESMDB time series, both processed by acausal filter, are
 411 quite similar in this case, although the ESMDB record has zero pads at beginning
 412 that are not shown in the plot.

413 As a further example, Fig. 8.7 illustrates a case of corrected ground motion from
 414 digital records. Reference is made to the NS component of the AQV record of the
 415 M_w 6.3 L'Aquila earthquake and the alternative source is the CESMD. In this case
 416 the HP frequency is 0.1 Hz for ITACA and 0.05 Hz for CESMD. The difference in
 417 the HP frequency is the reason of the clearer evidence of the acausal transient
 418 in the CESMD displacement trace. To avoid the onset of such spurious transients
 419 in the displacement waveforms from acausal high-pass filtering and to recover reli-
 420 able permanent displacements from double integration of accelerations, records of
 421 L'Aquila were also processed using a baseline correction technique that consists of
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 448 **Fig. 8.7** As Fig. 8.5, for the AQV corrected record, NS component, from the M_w 6.3 L'Aquila
 449 earthquake, 2009. Superimposed is the record corrected with a piecewise baseline on velocity to
 450 retrieve permanent displacements

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451 least-squares fitting the velocity time histories by three consecutive line segments,
 452 and subsequently removing these trends from the velocity time histories [1]. The
 453 resulting permanent displacements were found to be consistent with the GPS and
 454 INSAR findings (Anzidei et al., 2009; Atzori et al., 2009). Note that long period
 455 response spectral ordinates are practically unchanged using the three different pro-
 456 cessing techniques, confirming the findings by Paolucci et al. [10] regarding the
 457 reliability of long period response spectral ordinates from digital accelerograms.

458 Due to the space limitations of the paper, instead of documenting similar compar-
 459 isons on a much larger set of records, we summarize here the most significant
 460 outcomes of such comparisons:

- 461 – for digital records, results of ITACA, PEER, ESMDB and CESMD processing are
 462 similar;
- 463 – for analogue records, ITACA and ESMDB provide similar results except for (i) a
 464 more conservative selection of the ESMDB band-pass frequency range in several
 465 cases, (ii) tapering on a longer portion of records in ITACA and (iii) the retention
 466 of zero-pads in the ESMDB records;
- 467 – ITACA and PEER analogue records practically coincide whenever the PEER
 468 records are processed by acausal filters.

470 8.3.3 Processing of Late-Triggered Records

471 A significant portion of analogue strong-motion records of ITACA consists of
 472 accelerograms triggered by the S-phase arrival (*late-triggered records*). Processing
 473 such records faces several major difficulties, especially because tapering of the
 474 initial part of the signal would inevitably cancel out some important portions of the
 475 signal itself. In the new version of ITACA, late-triggered records are identified
 476 by a specific field, so that the end-user may decide to query the database without
 477 considering such records.

478 To support the identification of late-triggered (LT) records in the processing
 479 stage, a criterion was introduced based on the cumulated Arias intensity function,
 480 $I(t)$. For this purpose, each record is subdivided into three portions, as shown in
 481 Fig. 8.8, where D_1 is the time between the starting of the record and the time t_{05}
 482 for which $I(t_{05})=0.05$, and $D_2 = t_{95}-t_{05}$, where $I(t_{95})=0.95$. It was found that most
 483 of the LT records in ITACA could be identified by the condition $D_1/D_2 < 0.05$,
 484 although visual inspection of the records is always required.

485 Once the LT record has been identified, the procedure for correction is similar to
 486 the one for NT records, except for the following:

- 487 – the initial part of the record is not tapered;
- 488 – the zero-pads are always retained.

489 We can gain an interesting insight about the quality of LT records, by considering
 490 two co-located stations in Nocera Umbra, an analogue one (denoted by NCR in
 491 ITACA) and a digital one (denoted by NCR2). Table 8.1 lists the events for which
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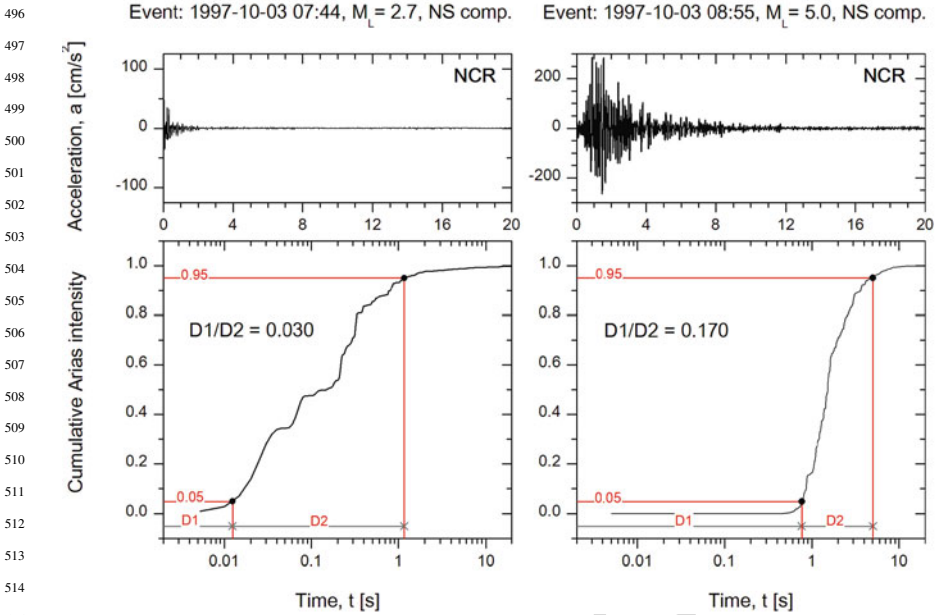


Fig. 8.8 Two analogue records from the same station NCR, identified as *late triggered* (LT, left) and *normally triggered* (NT, right)

both digital and analogue records are available, as well as the corresponding $D1/D2$ ratios and the $N_d(0-0.5\text{ s})$ parameter between the NCR and NCR2 response spectra normalized by NCR2. The latter parameter (N_d) measures the average difference of the response spectral ordinates in the 0–0.5 s period range. Therefore, $N_d=0$ means

Table 8.1 List of events and parameters associated to analogue records at NCR station

ID	M_w	M_L	$D1/D2$		$N_d(0-0.5\text{ s})$		Class. rec.
			NS	EW	NS	EW	
19971003_074404		2.7	0.030	0.019	0.55	0.25	LT
19971003_121624		2.9	0.003	0.006	0.81	0.42	LT
19971003_124844		3.1	0.011	0.017	0.56	0.39	LT
19971007_012434	4.2	4.1	0.005	0.108	0.02	0.17	LT
19971007_050956	4.5	4.3	0.036	0.028	0.68	0.49	LT
19971012_110836	5.2	5.1	0.024	0.031	0.16	0.10	LT
19971014_075405		3.3	0.008	0.011	0.42	0.01	LT
19971014_152309	5.6	5.5	0.059	0.099	0.10	0.06	LT
19971108_153153		4.1	0.014	0.034	0.85	1.32	LT
19980405_155221	4.8	4.5	0.059	0.050	0.05	0.07	LT
19971002_105956	4.7	4.1	0.073	0.082	0.12	0.07	NT
19971003_085522	5.2	5.0	0.170	0.164	0.09	0.10	NT
19971006_232453	5.4	5.4	0.346	0.398	0.07	0.09	NT
19971011_032057		3.7	1.163	0.526	0.12	0.04	NT

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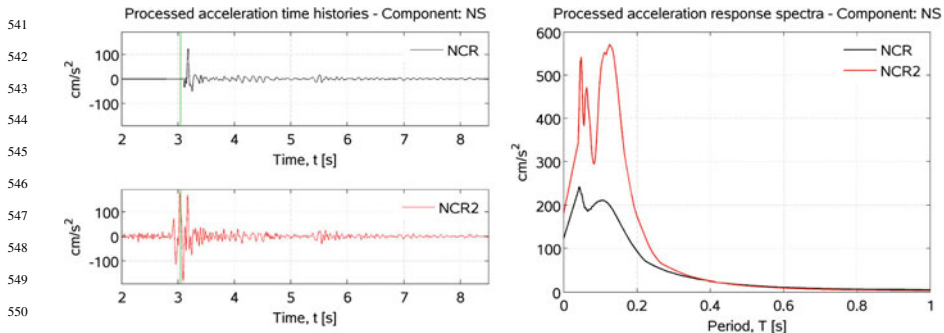


Fig. 8.9 Analogue (NCR) and digital (NCR2) corrected accelerograms of event 19971014_075405 (NS component in Table 8.1)

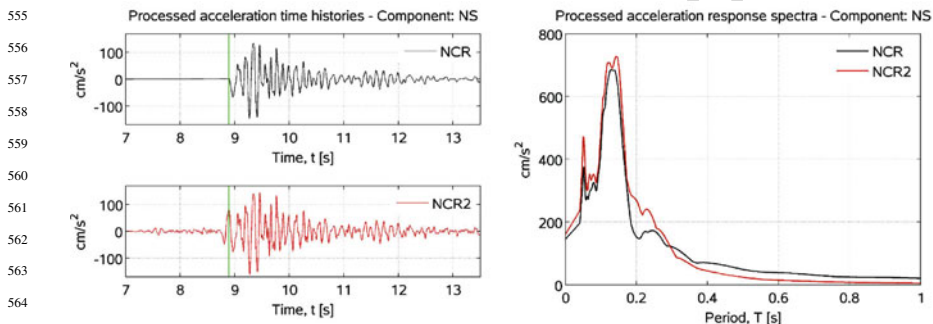


Fig. 8.10 Same as Fig. 8.9 for the NS component of event 19980405_155221

that the analogue and digital spectra coincide, while $N_d=1$ means that the average difference is 100%.

Examples of the corrected LT records at NCR, with the corresponding digital co-located records of NCR2 and the corresponding 5% damped response spectra of acceleration are shown in Fig. 8.9 and Fig. 8.10. It is clear that the case plotted in Fig. 8.9 illustrates a very poor quality record ($N_d = 0.47$, according to Table 8.1), while the corrected analogue accelerogram in Fig. 8.10 ($N_d = 0.05$) approaches the spectral ordinates of the digital record and can be considered usable for engineering applications.

Another interesting illustration about the quality of the LT records and their relationship with the proposed parameter D_1/D_2 is shown in Fig. 8.11 that shows the plot of N_d vs. D_1/D_2 . This plot suggests that the proposed rule-of-thumb $D_1/D_2 < 0.05$ to identify LT records is rather satisfactory, but it is difficult to use the same parameter D_1/D_2 to discriminate between “good” and “poor” quality LT records. A similar conclusion was drawn by Douglas [7], when considering a similar criterion to check the quality of LT records, based on the bracketed duration for acceleration values larger than 0.005 g.

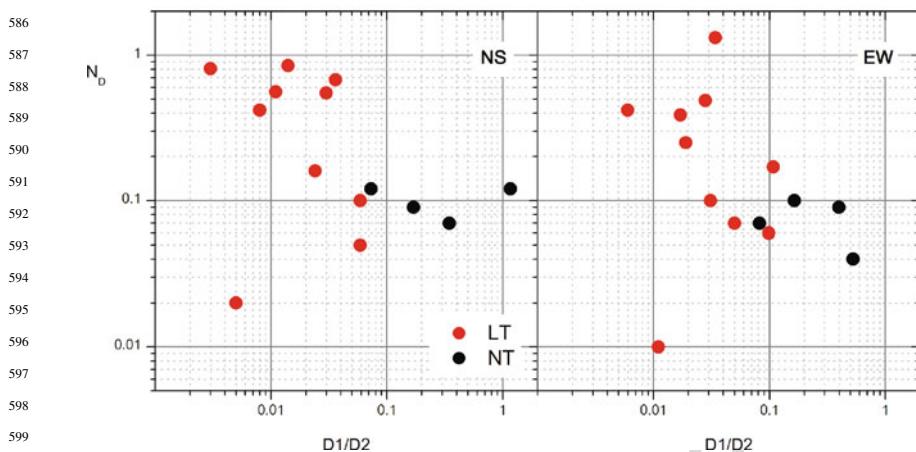


Fig. 8.11 Variation of the index $N_d(0-0.5\text{ s})$ as a function of the ratio D_1/D_2 for the records of NCR station

8.4 Conclusions

A notable effort has been made in the recent years to collect and organize in a single, informative and reliable Italian strong-motion database by the joint cooperation of the Italian Department of Civil Protection, the Istituto Nazionale di Geofisica e Vulcanologia, and several University research groups. The Italian ACelerometric Archive ITACA contains most of the strong motion accelerograms recorded in Italy since 1972. The final release of ITACA from Project S4 will be available by June 2010.

The quality and the level of station and event metadata were appreciated by many researchers and professionals who accessed ITACA after the L'Aquila earthquake. The rapid response of ITACA for collecting, processing and disseminating the data of this earthquake from Italian networks was also appreciated by the professional community.

Among different topics addressed in Project S4 to improve ITACA, this paper illustrated the main issues that were faced to provide reliable corrected accelerograms from a large set of records with a wide variation in quality and amplitude that are usable both for the engineering and research communities.

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Chapter 8

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