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Geochemical and isotopical variations within the Campanian Comagmatic Province: implications on magma source composition

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Abstract

A spatial variation in chemical and isotopical composition is observed between the volcanoes belonging to the Campanian Comagmatic Province. At a given MgO content, magmas from volcanic islands (Procida and Ischia) are enriched in Ti, Na, depleted in La, Ba, Rb, Sr, Th, K contents, and shows lower LREE/HFSE (*e.g.*, La/Nb = 1-2), lower Sr-Pb isotopic ratios and higher Nd isotopic ratios with respect to magmas from volcanoes located inland (Campi Flegrei and Somma-Vesuvius). The observed compositional variations are explained involving two different mantle sources in the genesis of the magmas erupted in this region: a deeper asthenospheric mantle source, from which the Tyrrhenian magmas also derived and a lithospheric mantle source enriched by slab-derived fluids. The contribution of the enriched-lithospheric mantle became more pronounced moving from the Tyrrhenian abyssal plain through the Italian Peninsula where it dominates, likely in response to the thickening of the lithosphere observed under the Peninsula.

Key words *lithosphere – asthenosphere – Campanian Province – Tyrrhenian Basin – geochemistry*

1. Introduction

The Campanian Plain is located between the western margin of the Apennine chain and the eastern border of the Tyrrhenian abyssal plain (fig. 1a,b). Since the Pliocene the Plain has been affected by intense extensive tectonic and magmatism (*e.g.*, Scandone, 1979; Peccerillo and Manetti, 1985; Di Girolamo, 1987; Doglioni, 1991). At the present time the Campanian Plain

is bordered by conjugate NE-SW and NW-SE fault systems formed during the opening of the Tyrrhenian Sea. The volcanic districts of Campi Flegrei, Procida and Ischia, and the strato-volcano of Somma-Vesuvius developed on NE-SW structure lines (see Beccaluva *et al.*, 1991, for a review).

Campanian volcanoes are among the most studied volcanoes in the world. The many papers published in the last decade (*e.g.*, Civetta *et al.*, 1991a,b, 1997; Civetta and Santacroce, 1992; Belkin *et al.*, 1993; Santacroce *et al.*, 1993; D'Antonio and Di Girolamo, 1994; Cioni *et al.*, 1995; Orsi *et al.*, 1995; Ayuso *et al.*, 1998; D'Antonio *et al.*, 1999a,b; de Vita *et al.*, 1999; Marianelli *et al.*, 1999; Pappalardo *et al.*, 1999, 2002; Piochi *et al.*, 1999; Del Moro *et al.*, 2001; Somma *et al.*, 2001) provide a wide data base of chemical and isotopic analyses. However the meaning of petrochemical variations

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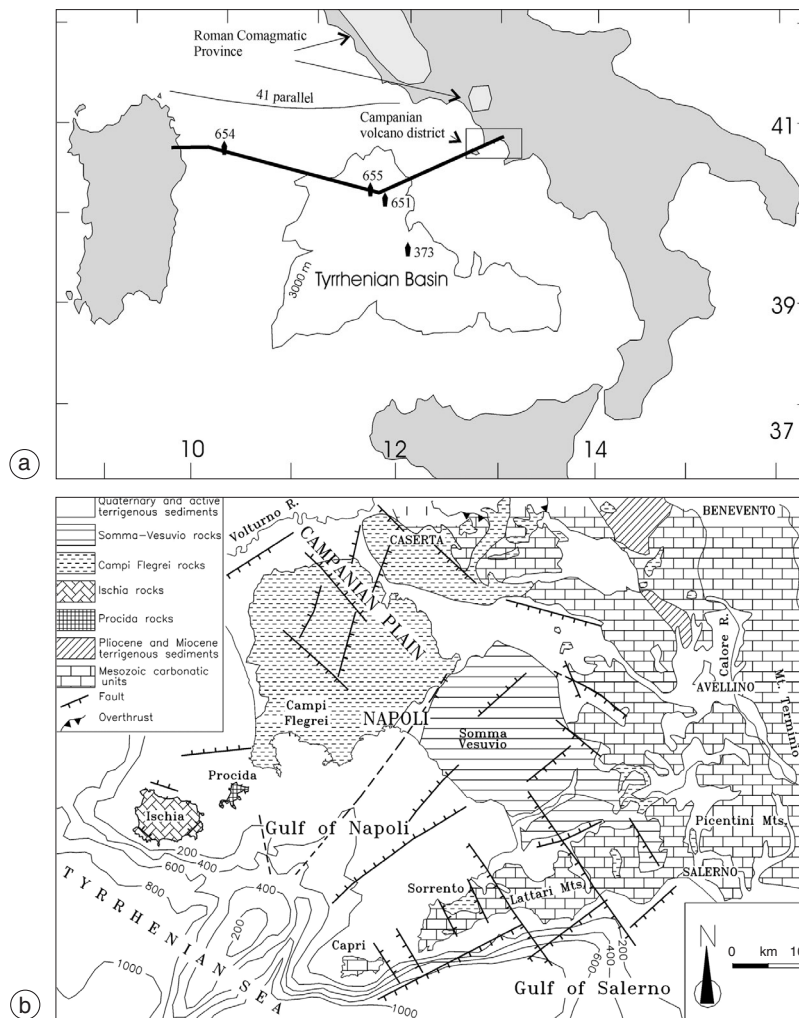


Fig. 1a,b. a) Southern Italy map showing location of the Campanian volcanic area and Tyrrhenian abyssal Plain. Bold line is the cross section in fig. 6; b) geological and tectonic sketch map of the Campanian Plain showing distribution of sedimentary sequences and Campi Flegrei, Ischia, Procida and Somma-Vesuvius volcanic rocks (modified from Orsi *et al.*, 1996).

shown by the erupted products is not completely clear, and a debate exists about the genesis and evolution of Campanian magmas and its geodynamic setting (see Turi and Taylor, 1976; Vollmer, 1976; Hawkesworth and Vollmer, 1979; Cortini and Hermes, 1981; Vollmer *et al.*, 1981; Peccerillo and Manetti, 1985; Ellam *et al.*, 1989; Beccaluva *et al.*, 1991; D'Antonio *et al.*, 1996; Ayuso *et al.*, 1998; Peccerillo, 1999, 2001; Somma *et al.*, 2001).

An open question about the nature of the most mafic magmas is whether they represent deep asthenospheric mantle magmas or reflects shallower imprinting acquired in the lithospheric mantle or in the crust. To shed light on this problem, our paper compares the available com-

positional data (major-, trace-elemental, and isotopes) from rocks erupted from Campanian volcanoes both situated inland (Somma-Vesuvius and Campi Flegrei) and offshore (Procida and Ischia), with that from magmas erupted on the Tyrrhenian abyssal plain. In this way, the regional chemical and isotopic variations of magmas going from east to west can be investigated in relation to the deepening of the crust - mantle boundary (Ferrucci *et al.*, 1989) and the increase in lithospheric thickness (Cella *et al.*, 1998) observed moving from Tyrrhenian Basin towards the Apennine Chain. We assume that if significant modification of the deep mantle magmas occurs *en route* to the surface and/or by the crust, then such a process should be revealed by comparing the geochemical and isotopic signatures of offshore and inland volcanic complexes.

2. Geological setting

The Campanian volcanoes are part of the potassic-rich Italian belt which developed on the eastern side of the Tyrrhenian Sea whose abyssal plain is composed of MORB-like basaltic rocks (Beccaluva *et al.*, 1990). They constitute the Campanian Comagmatic Province that is distinguished from the Roman one on the basis of some important geochemical differences in the outcropping volcanics (Duschenes *et al.*, 1986; Beccaluva *et al.*, 1991). The two Provinces are separated by a lithospheric discontinuity corresponding to the 41° parallel (fig. 1a,b) dividing two different mantle domains (Savelli and Wezel, 1979; Lavecchia, 1988; Spadini and Wezel, 1994; Bruno *et al.*, 2000). The Crust-Mantle boundary beneath the Campi Flegrei and Somma-Vesuvius has been identified at 25-30 km, whereas it rises moving westward, being at less than 25 km beneath Ischia (Corrado and Rapolla, 1981; Ferrucci *et al.*, 1989) and less than 10 km in the Central Tyrrhenian Basin (Duschenes *et al.*, 1986; Gueguen *et al.*, 1997). The nature of the lower crust is not known although the correlation with the Southern Apennine structures (Finetti and Morelli, 1973; Schutte, 1978; Scandone, 1982) and petrological studies on xenoliths (Pappalardo *et al.*, 2002) suggest that it should be a Hercynian basement.

The upper crust consists of a succession of Triassic to Cretaceous limestones and dolomites overlain by Miocene arenaceous and/or flysch sediments, or by pyroclastic rocks (D'Argenio *et al.*, 1973; Ippolito *et al.*, 1975) cropping out in the Apennine Chain. This succession, which is displaced down to 4000 m depth in the Campanian Plain (D'Argenio *et al.*, 1973; Ippolito *et al.*, 1975; D'Argenio *et al.*, 1987), has been drilled at depths of about 2.4 km beneath the Somma-Vesuvius (Brocchini *et al.*, 2001) and has been identified by seismic profiles at depths of more than 3-4 km in the Gulf of Napoli and Pozzuoli (Finetti and Morelli, 1974; Bruno *et al.*, 1998).

The Campi Flegrei, and the Ischia and Procida islands are volcanic fields of the Phlegraean Volcanic District characterized by several monogenic vents that produced both low-energy explosive and effusive eruptions. High-volume ignimbrite eruptions also occurred in the area at least from 200 kyr BP, some of these eruptions generated large caldera structures at Campi Flegrei and Ischia (*e.g.*, Rosi and Sbrana, 1987; Vezzoli, 1988; Orsi *et al.*, 1996; De Vivo *et al.*, 2001). Volcanism is broadly coeval at Campi Flegrei and Ischia where the onset of volcanic activity is dated before 80 kyr and 150 kyr and lasted up to 1532 A.D. and 1302 A.D., respectively (Alessio *et al.*, 1971; Gillot *et al.*, 1982). Volcanic activity occurred from about 80 kyr up to 14 kyr at Procida (see Rosi *et al.*, 1988; Beccaluva *et al.*, 1991). Somma-Vesuvius is a strato-volcano that experienced lava flows and small-to-high scale energy eruptions; the oldest Somma-Vesuvius volcanic products are dated at 0.4 Myr (Brocchini *et al.*, 2001) and the last eruption occurred in 1944 A.D.

At the present time, Campi Flegrei, Ischia and Somma-Vesuvius are the site of widespread fumaroles and thermal springs, as well as seismic activity. Furthermore, bradyseismic phenomena, characterized by ground movements and seismic activity affected Campi Flegrei in 1972, 1984 and 2000 A.D.

3. Petrological setting

The Campanian volcanic rocks belong to the KS series of Appleton (1972) and range in composition from shoshonites to trachy-phonolites.

The latter are the most widespread products. At Somma-Vesuvius mildly to highly undersaturated rocks belonging to the HKS of Appleton (1972), ranging in composition from alkali-basalt to phonolite, also occur. The Campanian rocks have nearly aphyric to strongly porphyritic texture: generally the porphyritic texture is higher in Somma-Vesuvius with respect to Phlegraean rocks, and in Ischia with respect to the Campi Flegrei and Procida rocks. In the KS evolved rocks K-feldspar, plagioclase, Fe-rich diopside, magnetite and biotite represent typical minerals. In the least evolved ones, K-feldspar and biotite are absent, olivine and Mg-rich diopside join the other phases reported. Apatite is the most common accessory crystal. The least evolved HKS rocks from Somma-Vesuvius contain olivine, plagioclase and Fe-rich and Mg-poor diopside. Leucite is the most common mineral in the most evolved rocks. Garnet and phlogopite are very common accessory phases.

Major and trace element and Sr-Nd-Pb isotope variations within the Campanian rocks have been attributed to complex evolutionary processes involving magmas chamber refilling, magmas mixing and crustal assimilation by mantle-derived magmas in a multi-depth magmatic system (*e.g.*, Cioni *et al.*, 1995; Ayuso *et al.*, 1998; D'Antonio *et al.*, 1999a; Pappalardo *et al.*, 1999, 2002; Piochi *et al.*, 1999; Del Moro *et al.*, 2001). In particular, the existence of both deep and shallower crustal reservoirs has been proposed in several studies of the Campanian volcanoes (Belkin *et al.*, 1985; Cortini *et al.*, 1985; Piochi *et al.*, 1999; Marianelli *et al.*, 1999; Pappalardo *et al.*, 2002; Lima *et al.*, 2003; De Astis *et al.*, 2004). Inside the Phlegraean Volcanic District, the deeper reservoir was tapped by a regional fault system during eruptions extruding the least evolved magmas that mingled during ascent with magmas evolving at shallower depth (De Astis *et al.*, 2004; and references therein). Compositional differences have been detected in volcanic rocks from Ischia island with respect to those from Campi Flegrei and Somma-Vesuvius (*e.g.*, Turi and Taylor, 1976; Cortini and Hermes, 1981; Paterne *et al.*, 1988; Vezzoli, 1988; Pappalardo *et al.*, 1999), although their meaning has not been thoroughly investigated.

The source of magmatism has been located in a mantle variably enriched in incompatible elements, radiogenic Sr and unradiogenic Nd (*e.g.*, Hawkesworth and Vollmer, 1979; Peccerillo and Manetti, 1985; Beccaluva *et al.*, 1991; D'Antonio *et al.*, 1996; Ayuso *et al.*, 1998; D'Antonio *et al.*, 1999b; Peccerillo, 1999, 2001). Debate exists about the agent of enrichment that should be i) mantle-derived fluids in an intra-plate tectonic setting (*e.g.*, Cundari, 1980; Vollmer, 1989) or ii) fluids or melts released by an undergoing oceanic slab (*e.g.*, Di Girolamo, 1978, 1987; Hawkesworth and Vollmer, 1979; Peccerillo and Manetti, 1985; Peccerillo, 1990, 2001; Beccaluva *et al.*, 1991; Serri *et al.*, 1993; D'Antonio *et al.*, 1996) which modify a OIB-type (*e.g.*, Beccaluva *et al.*, 1991) or a MORB-type mantle (*e.g.*, D'Antonio *et al.*, 1996).

Recently, the discovery of xenoliths of crustal origin in the least evolved ⁸⁷Sr-enriched rocks from the Campi Flegrei strongly contributed to highlight both the role of crustal contamination in magmas composition and the existence of a homogeneous mantle source beneath the Phlegraean area (Pappalardo *et al.*, 2002). The nature of xenoliths allows recognition of at least two crustal levels of contamination the deeper of which located in the Hercynian basement. Following Cecchetti *et al.* (2001), this depth is about 10-15 km and is similar to that proposed by Belkin *et al.* (1985), Belkin and De Vivo (1993), Marianelli *et al.* (1999), Lima *et al.* (2003) and by Zollo *et al.* (1996) for Somma-Vesuvius. Crustal contamination has been proposed also at Somma-Vesuvius on the basis of mineral chemistry, fluid inclusion and isotope data (*e.g.*, Savelli, 1968; Fulignati *et al.*, 1995, 1998; Gilg *et al.*, 1999, 2001; Del Moro *et al.*, 2001; Pappalardo *et al.*, 2004).

4. Data presentation

4.1. Geochemistry

Major and trace elements have been published in previous papers (Civetta *et al.*, 1991a,b, 1997; Civetta and Santacroce, 1992; Orsi *et al.*, 1992, 1995, 1996; Caprarello *et al.*,

1993; Belkin *et al.*, 1993; Santacroce *et al.*, 1993; D'Antonio and Di Girolamo, 1994; Cioni *et al.*, 1995; Ayuso *et al.*, 1998; D'Antonio *et al.*, 1999a,b; de Vita *et al.*, 1999; Marianelli *et al.*, 1999; Pappalardo *et al.*, 1999, 2002; Piochi *et al.*, 1999; Del Moro *et al.*,

2001; Somma *et al.*, 2001); the database is available on request from the authors. Figure 2 shows selected major and trace element variation diagrams *versus* SiO₂. These diagrams, although sometimes scattering mainly as consequence of different analytical procedures

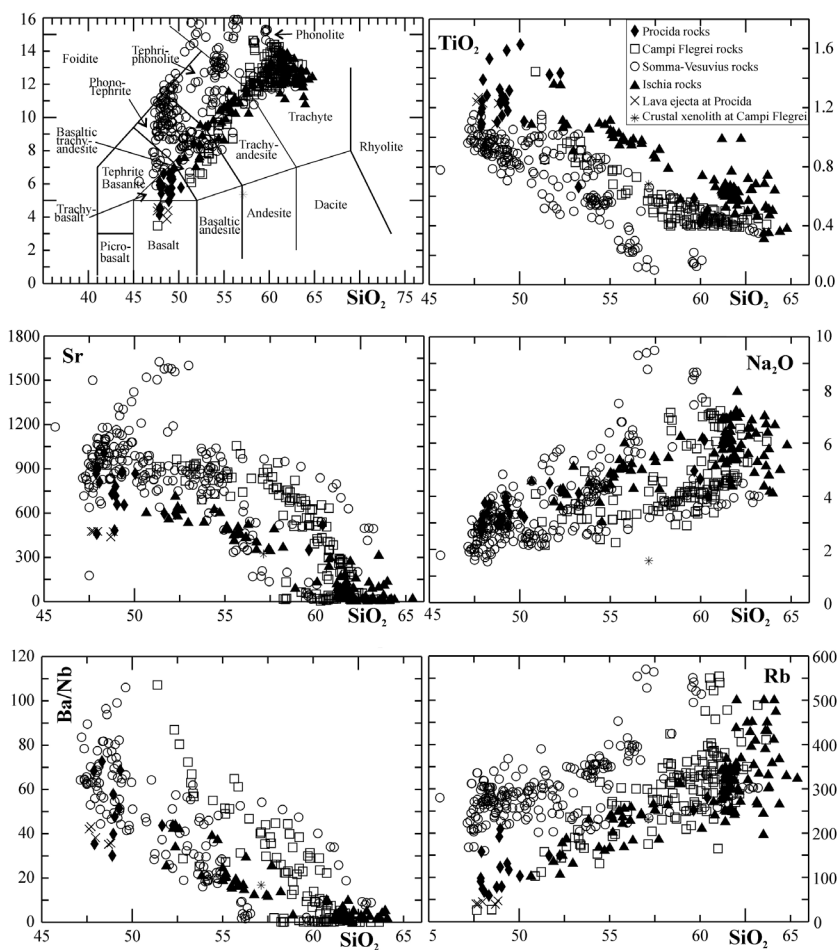


Fig. 2. Selected major and trace elements variation diagrams *versus* SiO₂. Source of data: Civetta *et al.* (1991a,b, 1997); Civetta and Santacroce (1992); Orsi *et al.* (1992, 1995, 1996); Caprarello *et al.* (1993); Belkin *et al.* (1993); Santacroce *et al.* (1993); D'Antonio and Di Girolamo (1994); Cioni *et al.* (1995); Ayuso *et al.* (1998); D'Antonio *et al.* (1999a,b); de Vita *et al.* (1999); Marianelli *et al.* (1999); Pappalardo *et al.* (1999, 2002); Piochi *et al.* (1999); Del Moro *et al.* (2001); Somma *et al.* (2001). Symbols: × = lava ejecta in Solchiaro deposits (Procida); full diamonds = Vesuvius samples; open squares = Campi Flegrei samples; open circles = Vesuvius samples; * = lithic clasts in Campi Flegrei deposits.

(ICP or XRF) and crystal contents (see previous section), display increasing Na₂O content and decreasing P₂O₅, TiO₂, MgO, and Fe₂O₃ contents. K₂O increases up to about 60 wt% of SiO₂ and then decreases; MnO decreases up to 60 wt% of SiO₂ then became strongly scattered. The most mafic compositions (MgO > 9wt%) are represented just by some rocks from Procida. La, Ce, Rb, Y, Zr, Nb and Th increase by factors up to 2 from mafic to silic compositions; Sc, Cr, Co Ni, Eu, V show an opposite trend. Ba is firstly strongly scattered in the range 45-55 wt% of SiO₂ content and then decreases, whereas in the same range, Sr

is initially constant and then decreases. Some samples from Somma-Vesuvius, related to a single eruption, define a separate array. The most evolved rocks from Campi Flegrei and Ischia show a steep enrichment in some incompatible elements (*e.g.*, Th, REE).

Magmas erupted at Somma-Vesuvius, define a trend depleted in Ti, Mg, Fe, P, Ca, V and enriched in Al, Nb, Rb, Th, K, Na with respect to the Campanian trends (fig. 2). Procida, Campi Flegrei and Ischia magmas show chemical similarity, although, at a given SiO₂ content, the Ischia magmas are enriched in Ti, Na, Nb, Yb and depleted in Ba, Sr, Th, Eu (fig. 2). However, in the least evolved rocks (MgO > 6 wt%) from each volcanic district display reasonably constant ratios between elements with similar degree of incompatibility. For example (fig. 2), La/Ce, Th/Zr and Nb/Zr are 0.5 ± 0.1 , 0.06 ± 0.04 , 0.14 ± 0.06 , respectively. Some variations could be related to different analytical procedures and analytical error.

REE and other trace element distributions are reported in fig. 3a,b, where Campanian magmas of similar degrees of evolution ($3.8 < \text{MgO} < 5 \text{ wt}\%$) are compared. The only exception in these spiderdiagrams are the Campi Flegrei rocks with MgO contents less than 3.8 wt% that Pappalardo *et al.* (2002) suggest are representative of the least-contaminated magmas. REE and trace element patterns are roughly sub-parallel to each other with no difference between products from each district. Generally, all selected samples show enrichment in LREEs and other incompatible trace elements (Rb, K, Th, Sr, Ba), and minor enrichment in more compatible elements such as HREE and Y. Negative Nb and Eu anomalies characterize the Campanian rocks, although their amplitude decreases toward the less evolved samples. Mafic rocks from Procida define a less enriched pattern, have lower REE and incompatible trace element contents coherently with their more primitive nature, and no detectable Eu anomaly.

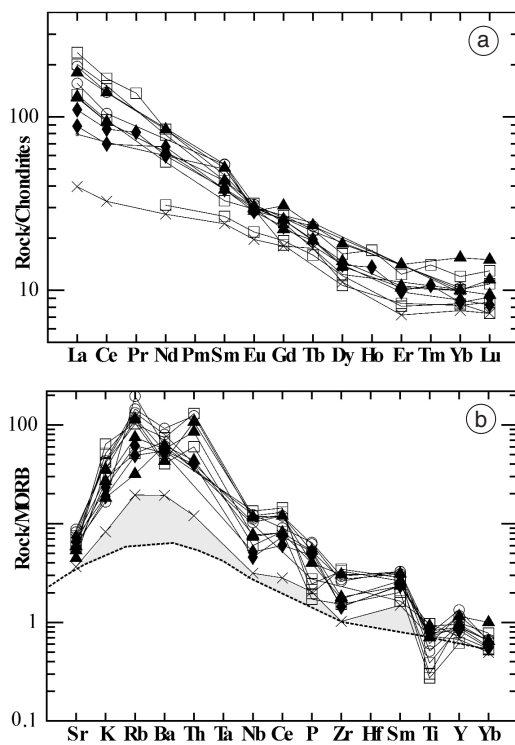


Fig. 3a,b. a) REE patterns for selected Campanian magmas. Normalization on the basis of Nakamura (1974); b) trace elements distribution for selected Campanian magmas. Normalization on the basis of Pearce (1983). Symbols as in fig. 2. Dashed area indicates the intra-plate component following Thorpe *et al.* (1984).

4.2. Isotope geochemistry

⁸⁷Sr/⁸⁶Sr and ¹⁴³Nd/¹⁴⁴Nd ratios are variable and depict a trend from Procida towards the

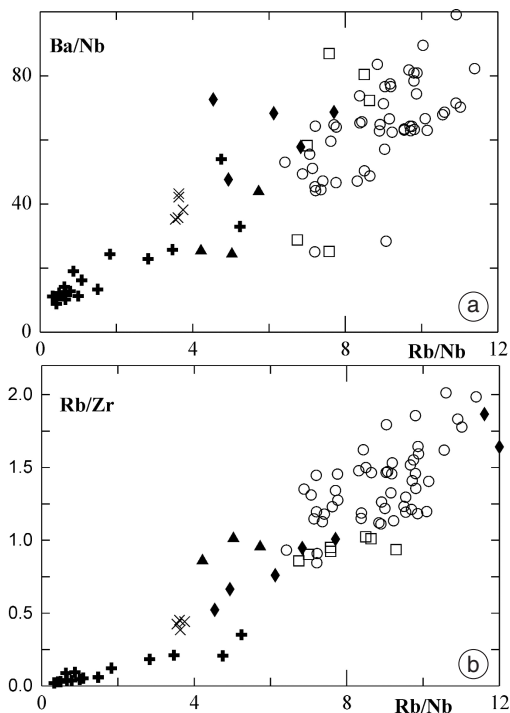


Fig. 4a,b. Selected LILE/HFSE diagrams for Campanian rocks with MgO > 3.8%. Symbols: full crosses = Tyrrhenian samples (Beccaluva *et al.*, 1990; Gasperini *et al.*, 2002). Other symbols as in fig. 2.

crustal-derived xenoliths recovered in Campi Flegrei rocks (fig. 4a). Magmas erupted at Ischia are isotopically distinct from that of Somma-Vesuvius and Campi Flegrei: Sr-isotope composition ranges from 0.7060 to 0.7068 at Ischia and from 0.7068 to 0.7089 at Somma-Vesuvius and Campi Flegrei. $^{87}\text{Sr}/^{86}\text{Sr}$ ratios also vary from 0.7051 to 0.7060 in the Procida volcanic rocks. Nd-isotope composition ranges from 0.51245 to 0.51265 in the Ischia rocks and from 0.51230 to 0.51265 at Campi Flegrei and Somma-Vesuvius. $^{143}\text{Nd}/^{144}\text{Nd}$ values range from 0.51255 to 0.5127 in rocks from Procida.

5. Discussion

The absence of «primary» magmas compositions at Campi Flegrei, Ischia, and Somma-

Vesuvius prevents knowing directly the nature of the mantle source.

Most of the mafic rocks produced in the Campanian region have much lower Cr, Ni and MgO contents with respect to those indicated by Perfit *et al.* (1980) (500-600 ppm, 300 ppm and 6 wt%, respectively), to have been in equilibrium with a peridotite magmas source. Some «primitive» compositions have been recognized in the lava ejecta on Procida island and on the basis of their high MgO, Cr and Ni contents (10-11 wt%, 426-610 ppm, 134-233 ppm, respectively) they were considered nearly primary magmas (D'Antonio *et al.*, 1996, 1999b). The features of the majority of the Campanian volcanic rocks are characterized by high SiO₂ contents and high porphyritic indicis, and the absence of mantle-derived xenoliths. These features suggest that the primary magmas underwent shallow-level differentiation processes after segregation from their mantle source. The isotopic variations reported above, not associated with Fractional Crystallization (FC) process in closed reservoirs, have to be produced by interactions of parental magmas with crustal rocks *en route* to the surface or by addition of crustal materials to the source. Therefore, as we are interested in identifying the nature of the source, we will discuss the rock features inherited by shallow-level processes and then deal with the source related features.

5.1. Shallow level magma evolution

Most of the chemical variations in Campanian rocks are consistent with a differentiation mechanism involving the fractionation of the observed mineral phases (FC) within closed magmatic systems. The trend defined by Sr and Eu respect to the silica content (fig. 2) and the existence of Eu anomaly in the most evolved rocks (fig. 3a,b) are consistent with the role of feldspar (the main crystal phase) crystallization during magmas evolution. The FC process is supported by the constancy of the Sr-isotopic ratios in rocks from some eruptive events at Ischia (Civetta *et al.*, 1991b; Piochi *et al.*, 1999), Campi Flegrei (Pappalardo *et al.*, 1999) and Somma-Vesuvius (Ayuso *et al.*, 1998 and refer-

ences therein). On the other hand, variations of the isotopic compositions of magmas with time have generally been related to mixing processes between distinct parental magmas batches reaching a shallow reservoir (Cioni *et al.*, 1995; D'Antonio *et al.*, 1999b; Pappalardo *et al.*, 1999; Piochi *et al.*, 1999). New studies on Campi Flegrei (Pappalardo *et al.*, 2002), on Somma-Vesuvius (Del Moro *et al.*, 2001; Pappalardo *et al.*, 2004) and Ischia (Piochi *et al.*, 1999) eruptions have also highlighted that the variability of the Sr-isotope ratio can be explained by crustal contamination in a multi-depth magmatic system. In these cases, magmas show an interaction with the wall-rocks because they contain crustal xenoliths (Campi Flegrei) and hornfelsed clasts (Somma-Vesuvius) with Sr isotope ratios > 0.708 and sometimes show chemical and mineralogical disequilibria. Moreover, Pappalardo *et al.* (2002) showed that the Sr-isotopic ratios increase in the younger Campi Flegrei products as their residence time in the magmas chamber increases. Furthermore, crustal interaction is indicated by the negative correlation described by Campanian volcanics and crustal xenoliths on the classical isotope diagram of Sr and Nd (fig. 5a). Finally, crustal contamination for the Campanian magmas is suggested by the La/Nb ratio which is always > 1 (Thompson *et al.*, 1984) and positively correlated to the Sr isotope ratio. Calculation based on the quantitative EC-AFC (Energy Conservation-Assimilation Fractional Crystallisation) approach that accounts for mass and energy conservation (Bohrson and Spera, 2001; Spera and Bohrson, 2001) indicates that the isotopic variations relative to the magmas feeding eruptions at Campi Flegrei and Somma-Vesuvius are explained by contamination with less than 40% of crustal rocks (Pappalardo *et al.*, 2002, 2004).

Despite the evidence of crustal interaction, there is a wealth of papers which suggest the existence of a «crucial» Sr-isotopic value below which we can assume to deal with «mafic rocks». A correlation is observed between the Sr-isotopic ratios of rocks and their crystallization depth at Somma-Vesuvius (Mastrolorenzo *et al.*, 2003; Pappalardo *et al.*, 2004). Olivines and pyroxenes from Somma-Vesuvius rocks crystallized at

depths > 10 km (Belkin *et al.*, 1985; Cortini *et al.*, 1985; Belkin and De Vivo, 1993; Marianelli *et al.*, 1999; Lima *et al.*, 2003) have Sr-isotopic ratios lower than 0.7074 (Civetta and Santacroce, 1992; Ayuso *et al.*, 1998, and references therein). These values are significantly different from those of the late-crystallized feldspars (crystallization depth < 5 km; Belkin *et al.*, 1985; Cortini *et al.*, 1985; Belkin and De Vivo, 1993; Lima *et al.*, 2003), with $^{87}\text{Sr}/^{86}\text{Sr}$ values around 0.7075-7 (Civetta *et al.*, 1991a; Civetta and Santacroce, 1992; Cioni *et al.*, 1995; Ayuso *et al.*, 1998, and references therein). Moreover, clinopyroxene cumulitic rocks from Somma-Vesuvius and Campi Flegrei (Ayuso *et al.*, 1998 and references therein), display $^{87}\text{Sr}/^{86}\text{Sr}$ below 0.7065. Pappalardo *et al.* (2002) exclude that the isotopic variations at Campi Flegrei are related to *in situ* radioactive growth, because a much longer time than that of the magmatic system (*ca.* 60 kyr) would be required to produce the observed isotopic variability, and indicate that petrological variation not reflect mantle melting processes and/or mantle source(s) heterogeneity because the constancy of ratios between elements with similar degrees of incompatibility of the products. The authors, on the basis of a correlation between Sr-isotopic ratios and age, and of Sr-isotopic similarity between the least radiogenic rocks and the mantle xenoliths from Vulture, have linked the least contaminated magmas erupted at Campi Flegrei with those erupted before the Campanian Ignimbrite eruptions characterized by a Sr-isotopic ratios in the range 0.7067-0.7073.

On the basis of these studies, we can choose as the least contaminated Campanian rocks those showing Sr-isotopic ratios lower than 0.70735. We can compare them with the mafic rocks from Procida and Ischia, which have $^{87}\text{Sr}/^{86}\text{Sr}$ widely below this limit and geochemical features akin to more primitive rocks.

5.2. Primary magma compositions and the nature of mantle sources

The possibility to discriminate the role of the shallow evolutionary processes (in particular crustal contamination) allow some representative mafic or near-mafic rocks to be investigated fur-

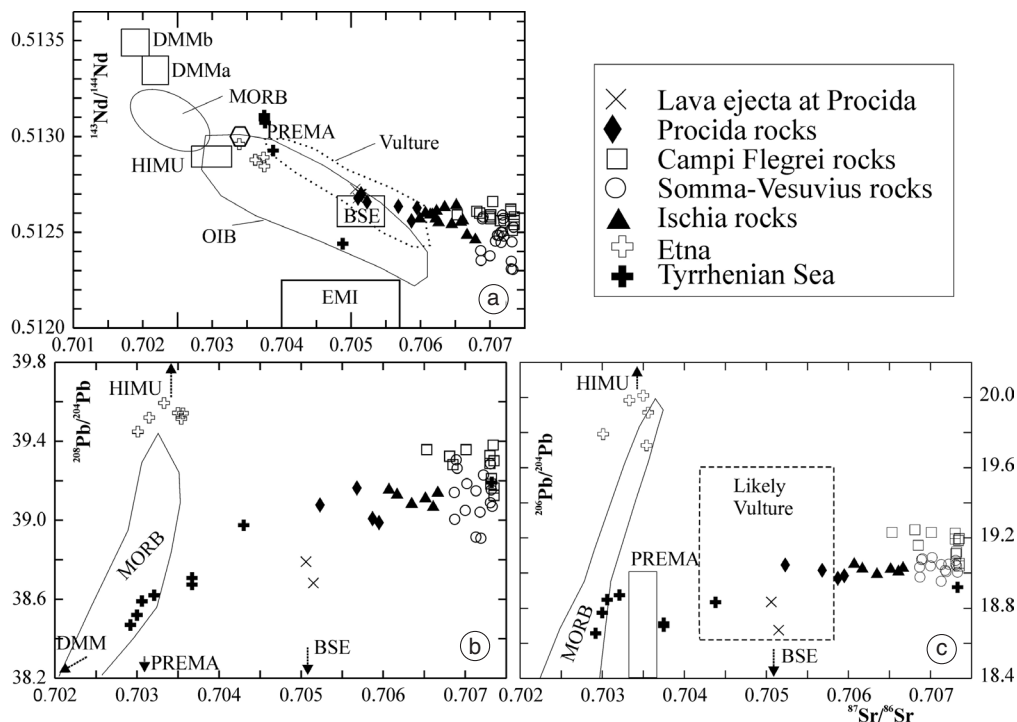


Fig. 5a-c. a) $^{143}\text{Nd}/^{144}\text{Nd}$, b) $^{206}\text{Pb}/^{204}\text{Pb}$, c) $^{206}\text{Pb}/^{204}\text{Pb}$ versus $^{87}\text{Sr}/^{86}\text{Sr}$ diagrams for selected Campanian magmas. Symbols: open crosses = Etna samples (Tanguy *et al.*, 1997); full crosses = Tyrrhenian samples (Beccaluva *et al.*, 1990; Gasperini *et al.*, 2002); field for the Vulture mantle xenoliths (a) from Downes (2001) and hypothesized (c) on the basis of considerations from Downes (2001) and Dunai and Baur (1995). Other symbols as in fig. 2.

ther although they have variable MgO contents (3.8-10-11 wt%). These rocks that approach or mirror the «original» geochemical and isotopic composition are those reported in figs. 3a,b, 4a,b and 5a-c. In the following, the Campanian rocks we have chosen will be compared with the transitional MORB-type basalts erupted in the contiguous sector of Tyrrhenian Basin (Vavilov Basin, sites 655 and 651 of ODP Leg. 107 and DSDP site 373, Leg. 42; Beccaluva *et al.*, 1990; Gasperini *et al.*, 2002) and with the intraplate basalts erupted on the opposite margin of the sea basin on the Sardinia coast (site 654).

The comparison among most mafic rocks from volcanic islands and from volcanoes located inland, as regards the major, REE and several trace elements pattern (figs. 2 and 3a,b) reveal some significant differences. At a given SiO_2

content, Ischia and Procida magmas are enriched in Ti, Na, depleted in La, Ba, Rb, Sr, Th, K contents, and show lower LREE/HFSE (*e.g.*, $\text{La}/\text{Nb} = 1-2$) with respect to the Campi Flegrei and Somma-Vesuvius magmas. Furthermore, LILE/HFSE ratios for rocks characterized by MgO exceeding 3.8% increase moving west to east through the Campanian volcanic area (fig. 4a,b). This trend is highlighted if we consider data (Beccaluva *et al.*, 1990; Gasperini *et al.*, 2002) for MORB-like magmas erupted in the contiguous sector of Tyrrhenian Basin (Vavilov Basin, sites 655 of ODP Leg. 107 and DSDP site 373, Leg. 42) that is also reported in fig. 4a,b.

The least contaminated Campanian samples define a unique trend when plotted in the $^{87}\text{Sr}/^{86}\text{Sr}$ versus $^{143}\text{Nd}/^{144}\text{Nd}$ diagrams (fig. 5a). Procida shoshonitic basalts have the lowest Sr

ratios (0.7050) and highest Nd (0.5128) ratios in the Campanian magmas, whereas the Campi Flegrei and Somma-Vesuvius rocks have the highest Sr ratios (0.7073) and the lowest Nd ratios (0.5126). Ischia magmas plot between these two compositions (fig. 5a). The Sr-isotopic differences are highlighted in $^{87}\text{Sr}/^{86}\text{Sr}$ versus $^{206}\text{Pb}/^{204}\text{Pb}$ diagrams (fig. 5b,c).

As in fig. 4a,b, in these isotope-isotope diagrams (fig. 5a-c) we have also drawn the field for MORB-like magmas erupted in the contiguous sector of Tyrrhenian Basin (Vavilov Basin, sites 655 of ODP Leg. 107 and DSDP site 373, Leg. 42) interpreted as derived from an upwelling asthenospheric mantle (Beccaluva *et al.*, 1990) probably having a plume-like imprint (Gasperini *et al.*, 2002). In the Sr-Nd isotope diagram (fig. 5a), the primitive magmas erupted at Procida plot between the Campi Flegrei and Somma-Vesuvius magmas and the Tyrrhenian magma field. This is supported also by Sr-Pb trends (fig. 5b,c) where the relationship between the Tyrrhenian and the Campanian magmas is highlighted. Therefore, we consider the Tyrrhenian basalts as end-members for the isotopic trend of fig. 5a-c.

As we have excluded from the discussion the crustally contaminated rocks, we assume that the other end member is also located in the mantle domain. The trace element patterns of the Campanian magmas (fig. 3a,b) are similar to intraplate-magmas with respect to immobile elements Zr, Hf, Ti, Y, Yb and, to a lesser extent, Nb but they are enriched in low ionic potential elements. Recent studies (Kostoula *et al.*, 1999; Downes, 2001) suggest a lithospheric origin for the mantle xenoliths recovered in the volcanic products from Vulture Volcano located to the east of the Campanian Plain. These xenolith samples are geochemically and isotopically similar to the Campanian magmas (fig. 5a-c). On this basis we suggest that the other end member could be sub-continental enriched lithospheric mantle. In particular, the enrichment in the low ionic potential elements, especially Rb and Ba, the small negative anomaly at Nb, and the ratio between incompatible elements (*e.g.*, Th/Yb > 2) should be generated by the interaction with subduction related fluid as also suggested for the Vulture xenoliths

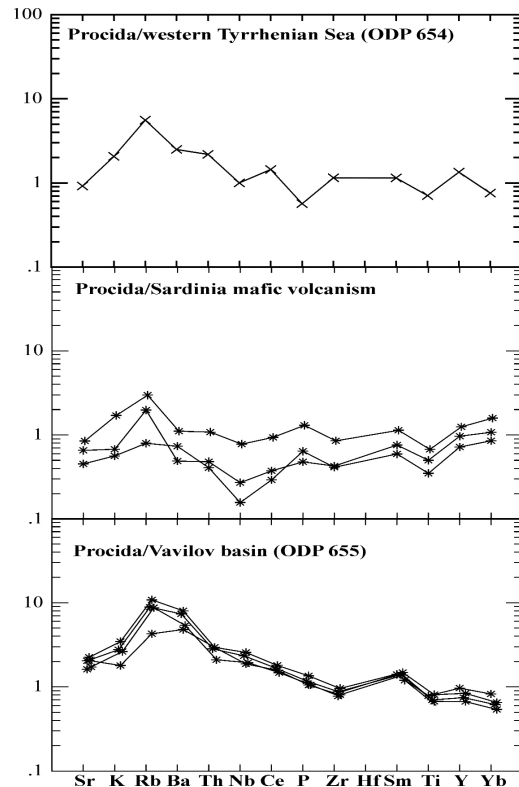


Fig. 6. Trace elements distribution for Procida mafic magmas normalized with respect to sample 654 from western margin of Tyrrhenian Sea (data from Beccaluva *et al.*, 1990), to samples 655 from Vavilov Basin (data from Gasperini *et al.*, 2002) and to samples from Sardinia mafic volcanism (data from Lustrino *et al.*, 2000).

(Downes, 2001). Following Peccerillo and Manetti (1985) the higher K-enrichment of the most Somma-Vesuvius rocks with respect to the other Campanian rocks should be the consequence of melting at greater depths.

Trace element distributions and isotope compositions of the volcanic rocks from the islands (higher Ti and P content and lowest Sr-Pb isotope ratios) compared with those from the mainland volcanoes, indicate that the asthenospheric contribution is more evident in the offshore volcanoes (Procida and Ischia) and it tends to vanish moving inland (Campi Flegrei and Somma-

Vesuvius) where the lithospheric contribution becomes dominant. This can be a consequence of the progressive drop of the lithosphere-as-thenosphere boundary moving from the Tyrrhenian Basin towards the Apennine Chain. Its depth varies from 25 km under the Tyrrhenian Basin, to 65 km beneath the Tyrrhenian coast of Italy and finally up to a depth of 100 km under the Apennine Chain (Cella *et al.*, 1998).

To test this hypothesis we compared the primitive Procida magmas (MgO = 11 wt%) with the anorogenic intraplate tholeiitic magmas erupted on the other border of Tyrrhenian Sea

(MgO = 7 wt%), near the Sardinia coast (site 654), in the Vavilov Basin (site 655) and on the Sardinia Island (fig. 6). The similarity between Procida and Sardinia mafic rocks, also in term of Sr, Nd and Pb isotope ratios (see Lustrino *et al.*, 2000), as well as the similarity in immobile elements (*i.e.* Sr, Nb, Zr, Sm, Yb) contents shown by Procida magmas and Tyrrhenian magmas, evidence that moving radially away from the basin toward its eastern (Italian Peninsula) and western (Sardinia Island) borders, the contribution of the asthenospheric MORB-type mantle decreases in favour of that of the «in-

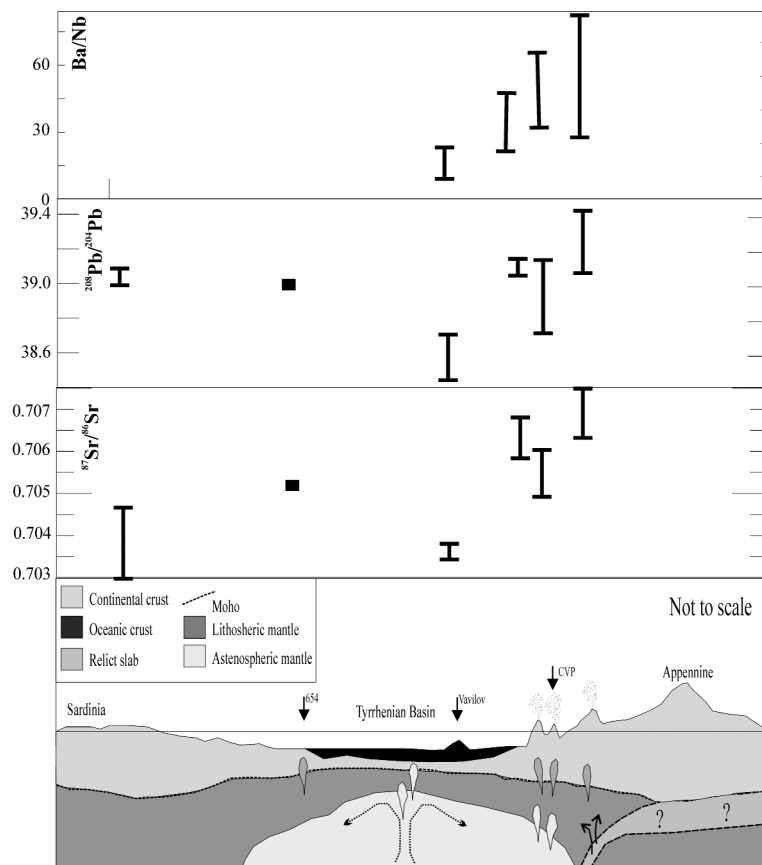


Fig. 7. Interpretative cross section from Tyrrhenian Sea towards the Campanian margin as indicated in fig. 1a,b. Chemical and isotopic variations moving from the Tyrrhenian Basin to the Campanian Comagmatic Province and to Sardinia are shown.

traplate» lithospheric mantle. Furthermore, moving away from the Tyrrhenian Basin the contribution of subduction increases.

6. Conclusions

A spatial variation in chemical composition has been recognised among the volcanoes belonging to the Campanian Comagmatic Province (fig. 7). When the role of low-P evolution processes has been evaluated, the most primitive (not contaminated by crustal assimilation) rocks erupted from Procida and Ischia islands have: i) lower LILE/HFSE ratios with respect to those erupted at Campi Flegrei and Somma-Vesuvius. These ratios are insensitive to partial melting and fractionation processes and can be considered to be representative of the magmas source; ii) lower Sr, Pb isotopic ratios and higher Nd isotopic ratios with respect to that of Campi Flegrei and Somma-Vesuvius.

These compositional variations can be explained assuming that both asthenospheric and «intra-plate» lithospheric mantle sources are involved in the genesis of the «primitive» magmas erupted in the Campania Comagmatic Region. The asthenospheric mantle, from which Tyrrhenian magmas were extracted, represents the deeper source. Moving from the Tyrrhenian abyssal plain to the Italian peninsula the contribution of enriched-lithospheric mantle became more pronounced in response to the increasing lithospheric thickness that magmas must cross *en route* to the surface (fig. 7).

The involvement of the lithosphere may be explained by geotherms uprising related to the asthenospheric upwelling. This hypothesis is supported by the presence of a low-velocity body underneath Somma-Vesuvius (De Natale *et al.*, 2001) which extends from 15 km down to the upper mantle (35 km) and represents the presence of a partially molten body into the lithosphere.

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