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Modeling the inter-annual variability of salinity in the lagoon of Venice in relation to the Water Framework Directive typologies}

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Keywords

Water Framework Directive (WFD) \sep transitional water \sep salinity \sep Venice Lagoon \sep numerical modeling \sep typologies \sep inter-annual variability

1 1. Introduction

2 Coastal transitional ecosystems are defined by Tagliapietra et al. (2009)
3 as “coastal water bodies with limited seawater supply”. Alternatively, if we
4 follow the definition proposed by the Water Framework Directive, transi-
5 tional waters can be identified as “bodies of surface water in the vicinity of
6 river mouths which are partly saline in character as a result of their prox-
7 imity to coastal waters but which are substantially influenced by freshwater
8 flows” ((European Parliament, 2000), art. 2(6)). Depending on freshwater
9 influence, coastal lagoons are assigned by the Directive to either “transi-
10 tional waters” or “coastal waters” (Tagliapietra and Volpi Ghirardini, 2006).
11 Both definitions recognize the importance of salinity and implicitly admit
12 the presence of spatial variation of salinity in the water bodies.

13 Transitional environments, especially lagoons, are characterised by strong
14 spatial heterogeneity, extreme values and broad fluctuations of several envi-
15 ronmental variables (Rosselli et al., 2009). Chemico-physical processes de-
16 termine gradients and patchiness (Attrill, 2002) for each variable, which in
17 turn leads to patchy or gradient-based distribution of biological components
18 (Levin et al., 2001; McLusky, 2004; Pèrez-Ruzafa et al., 2010).

19 The identification of environmental gradients and their interaction with
20 the biota in transitional ecosystems is key to the development of a frame-
21 work for the assessment of environmental quality. The Water Framework
22 Directive itself (henceforth, WFD) states that chemico-physical and hydro-
23 morphological elements, together with biological communities, should be con-
24 sidered when assessing the ecological status of water bodies ((European Par-
25 liament, 2000), Annex II). However, the biological community responds more

26 strongly to some of these parameters than others. Salinity and residence time,
27 the latter a measure of seawater renewal or confinement, are recognized as the
28 main factors and as proxies of the overall gradient (McLusky, 2004; Franco
29 et al., 2008; Pèrez-Ruzafa et al., 2007).

30 The spatial biological variation recognized in all lagoons (particularly in
31 micro-mesotidal lagoons, Barnes (1994)), with substitution of species along
32 environmental gradients, was related to seawater renewal by Guèlorge and
33 Perthuisot (1983). They defined the main factor controlling the distribution
34 of organisms and the features of populations as “the time of renewal of the
35 elements of marine origin at any given point”. They called it “confinement”
36 since it is strictly related to the degree of separation (seclusion) from the sea
37 and the distance from seaward inlets. Since a widely accepted mathematical
38 definition of confinement is still lacking, hydrodynamic parameters such as
39 residence time could be used as a proxy.

40 The literature on the effect of salinity variation on the biota is extensive.
41 At the community level, a model of benthic invertebrate species richness
42 along a marine-freshwater salinity gradient, based on studies performed on
43 the Baltic Sea and associated systems, was initially proposed by Remane
44 (1934), who described the overall reduction in the number of species in the
45 presence of progressively decreasing salinity levels. Various authors have
46 discussed different aspects of the model, and proposed modifications (Barnes,
47 1989; Hedgpeth, 1967; Odum, 1988). De Jonge (1974) underlines the need
48 to correlate organism distribution with average salinity and its fluctuation,
49 and to consider not only the number of species but also the composition of
50 the fauna. Telesh and Khlebovich (2010) discussed the concept of “critical

51 salinity” as a physiological and evolutionary barrier for marine and freshwater
52 fauna. Several studies have identified salinity as one of the most influential
53 environmental variables for the composition and abundance of invertebrate
54 communities in transitional waters (Williams, 1998, 2001; Pinder et al., 2005;
55 Piscart et al., 2005). Salinity is also a major factor in the distribution of
56 individuals and species among fish (Maci and Basset, 2009; Marshall and
57 Elliott, 1998) and submerged aquatic vegetation (Howard and Mendelssohn,
58 1999; Biber and Irlandi, 2006; Lirman et al., 2008).

59 Assuming salinity and residence time as the main proxies of the “compos-
60 ite gradient” in transitional waters, the effect on organisms of their spatial
61 and temporal variability is remarkable. The spatial and temporal variability
62 of salinity in transitional waters depends on freshwater inputs, precipitation
63 and evaporation rates, exchange with the sea and hydrodynamic transport.
64 The spatial and temporal variability of hydrodynamic transport (residence
65 time or renewal time) depends on freshwater inputs, precipitation and winds,
66 and exchange with the sea, a key role being played by the morphology of the
67 basin, which in turn is modified by the hydrodynamics.

68 Organisms of transitional ecosystems react in similar ways to pollution,
69 salinity change (Wilson, 1994), and more generally to the extreme and vari-
70 able conditions of transitional environments, making it difficult to separate
71 responses to anthropogenic stress from responses to natural variation. Transi-
72 tional ecosystems can be viewed as naturally stressed environments, particu-
73 larly if compared to marine conditions (Elliott and McLusky, 2002; McLusky
74 and Elliott, 2007). The term ”Estuarine Quality Paradox” has been intro-
75 duced by Dauvin et al. (2007) and by Elliott and Quintino (2007) to refer to

76 this concept. In transitional environments, where natural and anthropogenic
77 stresses are often associated, one way to approach the problem is to quantify
78 the natural variability and the resulting stress and then subtract this from
79 the anthropogenic stress.

80 The classification of transitional waters on the basis of salinity is an open
81 question. The Remane model (Remane, 1934) and subsequent studies of the
82 role of salinity gradients in structuring benthic communities form the ba-
83 sis of the “Venice system” (Venice System, 1959; Segerstraale, 1959). Given
84 the complexity of the relationship between community structure and salinity,
85 some authors have proposed overlapping limits between classes in their clas-
86 sification systems (Greenwood, 2007; Bulger et al., 1993; Wolf et al., 2009).
87 Attrill (2002) preferred salinity range to absolute salinity values, as variation
88 in salinity (and in environmental factors generally) may be more important in
89 structuring communities than extreme values. He also explicitly used salinity
90 range as a proxy for a set of variable conditions.

91 A well-known classification of lagoons according to water exchange with
92 the sea was developed by Kjerfve and Magill (1989), who considers leaky,
93 restricted and choked lagoons with gradually decreasing seawater exchange
94 and thus increasing seawater renewal time.

95 The difficulty of constructing a single classification system valid for all
96 transitional environments lies in the heterogeneity within and among these
97 systems and in their high temporal variability. The complex response by
98 the community to variation in environmental factors further complicates the
99 establishment of a common system of classification.

100 European Directive 2000/60/EC establishes a framework for water policy

101 and includes strategies to safeguard the ecological and chemical status of
102 water resources. To achieve these aims it requires the characterisation of
103 water bodies by the identification of “types” at appropriate spatial scales
104 (European Commission, 2003).

105 The classification of water bodies in terms of quality, which takes account
106 of abiotic and biotic elements, environmental pressures and resulting impacts,
107 is based on these types.

108 This entails identifying areas with well-defined physical characteristics
109 and serves to ensure common reference conditions. A water body thus classi-
110 fied as belonging to a specific type is considered homogeneous and represents
111 the unit that will be used for assessing compliance with the Directive’s envi-
112 ronmental objectives.

113 The WFD describes two systems for specifying types in transitional wa-
114 ters. System B, which is the most common, makes reference to obligatory
115 descriptors (Latitude, Longitude, tidal amplitude and salinity) and to op-
116 tional descriptors, of which residence time is one.

117 However a complete typology for transitional waters has not yet been
118 defined (Hering et al., 2010). The Common Implementation Strategy (CIS)
119 working groups are seeking to develop commonly agreed typologies at the
120 European level. Other European groups are working on the issue of intercal-
121 ibration between member states (Vincent et al., 2003; Hering et al., 2010).

122 Although the implementation guidance of the Directive recognises the
123 natural temporal variability of biological quality elements (European Com-
124 mission (2003), section 4.2 and 4.7), little is said about temporal variations
125 in the abiotic parameters on which the typologies are based. In this re-

126 gard it is merely suggested that the characteristics of a water body should
127 be determined by considering mean annual values (European Commission
128 (2003), section 3.2.3) without reference to the length of the time series. As
129 a consequence, different temporal scales could be considered.

130 Numerical models can be used to simulate the hydrodynamic and trans-
131 port process in a basin, and can also represent the spatial and temporal
132 variability of salinity and evaluate hydrodynamic transport scales in several
133 points of the basin.

134 The WFD does not refer to the use of numerical models. It explicitly
135 mentions modelling as a suitable method only to extrapolate reference con-
136 ditions ((European Parliament, 2000), Annex II art. 1.3) when a reference
137 site is not available. Hojberg et al. (2007) points out that monitoring and
138 modelling are inter-dependent (Holt et al., 2000; Parr et al., 2003; Irvine,
139 2004; Moschella et al., 2005; Dabrowski and Berry, 2009), but when imple-
140 menting the monitoring obligations of the WFD, models are rarely used in
141 practice. It is important to note that the acceptable level of monitoring
142 precision and confidence in the WFD is not well described. Rather, it is a
143 subjective issue that depends on socio-economic interests and the risk strat-
144 egy of the decision-makers. Hattermann and Kundzewicz (2010) analyzes
145 how numerical models could be used at various stages in the application of
146 the WFD.

147 While the WFD treats the use of numerical models only marginally, the
148 literature contains extensive references to their application to the study of
149 several aspects of lagoon dynamics and lagoon management. Numerical mod-
150 els can be used to calculate hydrodynamic transport in transitional environ-

151 ments on the scale of the whole basin and to calculate its spatial variability
152 within basins (Wang et al., 2004; Cucco et al., 2006; Gourgue et al., 2007;
153 Cucco et al., 2009; Jouon et al., 2006). The results can be used to distin-
154 guish the circulation in different parts of the basin, to identify areas that
155 are at higher risk of accumulating substances (Cucco and Umgiesser, 2006;
156 Luick et al., 2007; Wang et al., 2009; Rapaglia et al., 2010) and to determine
157 the main forcing factors and/or processes conditioning residence time itself
158 (Tartinville et al., 1997; Wijeratne and Rydberg, 2007; Plus et al., 2009; Mal-
159 hadas et al., 2010; Huang et al., 2010; Cavalcante et al., 2011) Salinity can
160 also be successfully simulated in transitional waters (Solidoro et al., 2004;
161 Huang, 2007; Huang et al., 2002) and numerical models can be used to study
162 the spatial and temporal variability of coastal lagoons (Obrador et al., 2008;
163 Lopes et al., 2010; Faure et al., 2010). In addition, numerical models have
164 been used to advance proposals for the zoning of shallow basins (Ferrarin
165 et al., 2008, 2010), and to evaluate the consequences of different manage-
166 ment strategies (Tsihrintzis et al., 2007; Gong et al., 2008; Hakanson and
167 Duarte, 2008). The adoption as normal practice of the calibration and val-
168 idation of every module of the model, together with the modelling quality
169 assurance procedures, allows the associated error to be accurately estimated
170 and ensures the reliability of numerical models.

171 We use a hydrodynamic numerical model to simulate the circulation of
172 water masses and the dispersion of a passive tracer, in order to develop an
173 objective, transparent, and cheap method for typing lagoons, classified as
174 transitional waters by the WFD. This method can be applied to different
175 years to explore the interannual variability of the descriptors and its effect

176 on the typing process. It may represent a first step in the evaluation of
177 natural variability and could be adapted to identify the natural stresses on
178 organisms in future studies. Finally, the results do not purport to offer a
179 conclusive solution to the typing of lagoons, but they can be employed to
180 suggest management approaches for the lagoon of Venice.

181 The present study takes account of a limited number of variables, in
182 agreement with the Directive's suggestions (European Commission (2003),
183 section 3). Working within the System B framework, we considered annual
184 mean salinity (an obligatory factor) and mean residence time (an optional
185 factor). Following Tagliapietra and Volpi Ghirardini (2006), our approach to
186 the typing process takes account only of abiotic parameters. The final reso-
187 lution of the Symposium of the Venice System (Venice System, 1959), which
188 established a classification system for Marine Waters based on salinity, rec-
189 ommended the use of additional details in addition to the average values,
190 including the salinity range over different timescales. The words "poikilo-
191 halinity" and "homoiohalinity" indicate unstable (variable) and stable (con-
192 stant) salinity respectively; other studies have proposed several statistical
193 measurements of the variability of salinity (De Jonge, 1974). From these
194 considerations, we decided to introduce a new factor: the annual standard
195 deviation of salinity, in order to take account of the variability around the
196 mean value.

197 The following sections illustrate the criteria used to select the sites, fac-
198 tors and methods, and then the results obtained. Section 2 sets out the
199 reasons for choosing the Lagoon of Venice as a case study, describes the
200 lagoon's main characteristics and justifies the three descriptors adopted in

201 the present study. Section 3 presents a short overview of the methods em-
202 ployed in identifying water body types, explains the advantages of using a
203 numerical model combined with datasets to perform the typing process and
204 describes in detail the method adopted. Section 4 illustrates the results ob-
205 tained and section 5 presents our conclusions and considerations on water
206 body management.

207 **2. Selection and description of the Case Study**

208 The Lagoon of Venice is a complex system, characterized by a number of
209 gradients and a mosaic of environments and morphologies that are the re-
210 sult of complex environmental and anthropic drivers. It is one of the biggest
211 in the Mediterranean and the biggest in Italy. This unique natural envi-
212 ronment, of high ecological value, is subject to a difficult coexistence with
213 human activities, such as industry, tourism, fisheries and pressures from the
214 drainage basin. An appropriate management system is thus fundamental.
215 Several studies, including monitoring activities and previous applications of
216 numerical models, provide sufficient expertise to apply a numerical model
217 and a sufficiently broad dataset to calibrate it and validate it.

218 The Venice lagoon is located in the northwest Adriatic Sea ($45^{\circ} 24' 47''$
219 N, $12^{\circ} 17' 50''$ E), it has a surface area of about 550 km^2 , with a north-south
220 length of 50 km and a mean horizontal width of 15 km. Approximately 436
221 km^2 are subject to tidal excursion, while the remainder has been closed off
222 to create fish-farms with limited and artificially regulated water exchange
223 (Guerzoni and Tagliapietra, 2006). Three inlets on the western side of the
224 lagoon allow water exchange with the sea. From north to south, these are

225 named Lido, Malamocco and Chioggia (mean depth 14, 17 and 8 m respec-
226 tively) and are shown in Fig. 1. The bathymetry of the lagoon is variable,
227 since it includes navigable channels, subtidal flats and intertidal features such
228 as saltmarshes. The latter are alternately submerged and exposed for vary-
229 ing periods of time with a frequency that depends on tidal cycles. In terms
230 of depth distribution 5% of the lagoon is deeper than 5 m and 75% is less
231 than 2 m. The mean depth is 1.2 m, but there are some areas with depths
232 greater than 30 m (Molinaroli et al., 2007).

233 The mean water volume of the lagoon is around $590 * 10^6 \text{m}^3$ and the ex-
234 change of water through the inlets in each tidal cycle represents about a third
235 of the total volume of the lagoon (Gačić et al., 2004). The tidal exchange
236 of seawater and the inflow of freshwater from several rivers determine the
237 lagoon's brackish character and the seasonal spatial gradients in the distri-
238 bution of abiotic and biotic variables.

239 The DRAIN project (1999-2000) estimated that inputs of freshwater to
240 the lagoon from the drainage basin (surface area 1850 km^2) amount to an
241 annual mean flux of around $35.5 \text{ m}^3 \text{ s}^{-1}$ (Zonta et al., 2005). The main rivers
242 with natural discharge regimes are the Silone (accounting for 23% of the total
243 flux) and the Dese (21%) together with the navigable channels called Naviglio
244 Brenta (14%) and Taglio Nuovissimo (13%). The most important rivers are
245 located in the northern part of the lagoon, which receives more than 50%
246 of the annual discharge from the drainage basin (Zuliani et al., 2005). Most
247 stretches of the rivers entering the southern part of the lagoon are artificially
248 regulated.

249 The Venice lagoon can be classified as a microtidal environment (mean

250 tidal range less than 1 m), with a mean tidal range of 61 cm, which decreases
251 to 35 cm during neap tide and increases to 79 cm during spring tide (M.
252 Sigovini and D. Tagliapietra, unpublished data). It is defined as a polyhaline
253 lagoon, with salinity varying along a gradient from the landward side to the
254 sea (Guerzoni and Tagliapietra, 2006; ICRAM, 2007; Solidoro et al., 2004).
255 Following Kjerfve and Magill (1989), it could be defined as “restricted” la-
256 goon, where tide and wind are the main forcing factors of circulation. Salinity
257 and residence time may be considered the main variables characterizing the
258 system’s conditions, and are also related to its trophic state (Solidoro et al.,
259 2004; Bianchi et al., 1999).

260 For the purposes of the WFD, the lagoon falls into the Transitional Waters
261 category for the Mediterranean Ecoregion. Applying system B to the Venice
262 Lagoon, we made the following considerations: Latitude and Longitude are
263 not relevant in this case due to the limited variability of both (the lagoon can
264 be enclosed within a square whose sides are around half a degree in length,
265 corresponding to 50 km). Therefore, salinity (both annual mean and range)
266 was the only obligatory factor adopted for the definition of types.

267 Several systems for classifying water bodies, based on various approaches
268 (ICRAM, 2007; CVN, 2004a,b; Zanon, 2006) are available in the local liter-
269 ature. Solidoro et al. (2004) applied the same numerical model used in this
270 study, with lower spatial and temporal resolution, and divided the lagoon
271 into 3 areas with respect to salinity and 11 areas with respect to internal
272 exchanges.

273 3. Selection and description of the Method

274 Several European studies have applied the requirements of the WFD to
275 case studies of coastal and transitional waters (Schernewski and Wielgat,
276 2004; Bulger et al., 1993). Their methods include the combined use of
277 GIS and numerical modelling techniques, as well as statistical approaches
278 based on water quality databases (Urbanski et al., 2008; Basset et al., 2006).
279 Some studies have adopted transitional water typologies based on hydro-
280 morphological characteristics such as morphology, tidal range and salinity
281 (Carstens et al., 2004; Tagliapietra and Volpi Ghirardini, 2006; Kagalou and
282 Leonardos, 2008). Others studies have also included human activities, pres-
283 sures and nutrient loads (Boix et al., 2005; Ferreira et al., 2006). The pub-
284 lished papers based on the implementation of the WFD to transitional waters
285 in the Mediterranean ecoregion do not include reference sites or reference cri-
286 teria but identify “a priori” typologies based on WFD system B descriptors.
287 One way to approach the typing process is to define broad types (e.g.. Moss
288 et al. (2003)) but these have yet to be determined for transitional waters
289 (Borja et al., 2009). Another approach is to draw up a detailed typology
290 reflecting ecological gradients and community structures, moving towards a
291 site-specific assessment (Hering et al., 2010).

292 Some studies consider the possible consequences of inter-annual varia-
293 tion. Lucena-Moya et al. (2009) includes the effect of intra-annual salinity
294 variation on phytoplankton and invertebrate communities by introducing a
295 classification into subtypes. Wolf et al. (2009) approached the longitudinal
296 zoning of tidal marshland streams by combining the abiotic salinity clas-
297 sification proposed by the WFD with a biotic classification based on the

298 salinity preference scores of benthic macroinvertebrate fauna. Galvan et al.
299 (2010) approached the heterogeneity within and between transitional waters
300 by adopting a hierarchical classification system. This study combined hy-
301 drological and morphological indicators and applied a circulation model to
302 estimate some parameters.

303 Mathematical models have been applied to several aspects of the WFD,
304 from the estimation of indexes for the biological community (Ponti et al.,
305 2008; Mistri et al., 2008) to the assessment of chemico-physical status (Garcia
306 et al., 2010; Bald et al., 2005) and ecological status (Nielsen et al., 2003).
307 Yang and Wang (2010) suggested introducing a model for managing diffuse
308 source pollution into the Programme of Measures associated with River Basin
309 Management Plans. Martins et al. (2009) combines classical monitoring of
310 water status with modelling of hydrodynamics, water quality and ecological
311 aspects. Nobre et al. (2010) presents an example of ecosystem modelling
312 as a tool for Integrated Coastal Zone Management and the adoption of an
313 ecosystem-oriented approach to marine resource management. The use of
314 numerical models to simulate ecological aspects as required by the WFD
315 and the establishment of reference situations by modelling are discussed by
316 (Nielsen et al., 2003; Wasson et al., 2003).

317 In section 1 we discussed how salinity and residence time can be consid-
318 ered as the main environmental proxies in complex transitional waters, and
319 how the temporal variability of the parameters can be a useful descriptor
320 itself. Often the temporal and spatial coverage of salinity data is too limited
321 to provide an adequate picture of its variability (Wolf et al., 2009). The
322 costs of a sampling grid able to reflect the spatial and temporal variability

323 of the main parameters, or even just salinity, are sometimes too high (Irvine,
324 2004). To solve this problem and to evaluate the implications of the variabil-
325 ity of this parameter for the typing process, we developed a numerical salinity
326 model with high spatial and temporal resolution, comparing the result with
327 a limited number of continuous, strategically located sampling points. This
328 method has the advantage of being less expensive than high-frequency mon-
329 itoring with high spatial resolution; the model makes it possible to estimate
330 residence time in every element of the grid and to obtain a map showing
331 annually averaged values. To represent interannual variability, we applied
332 the model to two years, 2003 and 2005, which were very different from the
333 climatological and hydrological point of view.

334 This study adopted the SHYFEM model ([https://sites.google.com/
335 site/shyfem/](https://sites.google.com/site/shyfem/)), which was developed expressly for coastal lagoons (Umgiesser
336 and Bergamasco, 1995). It has already been applied successfully to the Venice
337 Lagoon (Umgiesser et al., 2004; Bellafiore et al., 2008; Ferrarin et al., 2008)
338 where it has been used to simulate residence time and salinity Cucco et al.
339 (2006); Solidoro et al. (2004). A full description of the model can be found
340 in Umgiesser et al. (2004).

341 *3.1. Grid and model set-up*

342 With respect to the grids used in previous studies, (Solidoro et al., 2004;
343 Umgiesser et al., 2004) the spatial resolution and the detail of the contours
344 have been improved in order to better represent the bathymetric gradient
345 at reduced computational cost. The main channels crossing the islands have
346 been introduced and the spatial resolution of the shoals and some saltmarshes
347 has been increased in order to improve the simulation of the currents in

348 shallow water and the wet/dry behaviour of the saltmarshes. The grid itself
349 consists of 8029 nodes and 14021 elements (compared to 4367 nodes and
350 7858 elements in the previous grid) and the bathymetric data adopted were
351 collected in the year 2000 (Molinaroli et al., 2009).

352 Simulations start on January 1st and are 1 year long. They represent
353 the years 2003 and 2005, for which the salinity measurements have good
354 spatial and temporal coverage respectively. The model was applied in its
355 two-dimensional version to the lagoon only. The set-up adopted and the
356 method applied to calibrate the modelled water levels are the same as in
357 Umgiesser et al. (2004), where equations and the details of the numerical
358 treatment can be found. In all simulations, realistic forcing factors with a
359 maximum admissible time-step of 300 s and a spin-up time of 5 days were
360 adopted. The initial water level and velocity values were set to 0 and the
361 initial salinity was assigned spatially interpolated values from experimental
362 data corresponding to the start time of the simulation.

363 The timeseries for precipitation and wind (speed and direction) were con-
364 sidered in this application to be spatially homogeneous in the domain. The
365 same principle was adopted for air temperature, solar radiation, relative hu-
366 midity and cloud cover, which were used to calculate the effect of evaporation
367 on water level and salinity. To consider the effect of freshwater inputs, the
368 daily discharges of 11 rivers were included. Their location is shown in Fig. 2.

369 *3.2. The data*

370 The real forcing data used for the model and the comparison data for
371 salinity were processed for both simulated years (2003 and 2005). The tide
372 level data used to force the open boundary levels were collected at each

373 of the seaward inlets every 5 minutes by the Venice Tide Forecasting Cen-
374 tre, which manages a network of automatic weather and tide gauges in the
375 lagoon (<http://www.comune.venezia.it/>). The meteorological data were
376 collected every hour in 2003 and 2005 by the Italian National Research Coun-
377 cils Institute of Marine Sciences (ISMAR-CNR, Venice city). Missing data
378 were retrieved with reference to the corresponding meteorological data mea-
379 sured in Venice city by the Cavanis Institute (www.cavanis.org).

380 Comparison of meteorological characteristics in 2003 and 2005 with the
381 long-term average (1959-2004) shows that 2003 had lower annual precipita-
382 tion and higher air temperature (544 mm and 14.8 C), while 2005 (788.6
383 mm and 13.7 C) was similar to long-term trends (1954-2004 annual average:
384 789.5 mm and 13.6 C). Analysis of monthly precipitation (Pennacchi and
385 Benedetti, 2005, 2006) shows that both years had a maximum in April, and
386 from July to October rainfall in 2005 was much higher than in 2003 (the sum
387 of the values for these months is equal to 450.8 mm in 2005 and 172.8 in
388 2003).

389 In both years annual wind intensities and annual wind directions were in
390 agreement with literature data for the region (Gačić et al., 2009; De Biasio
391 et al., 2008), which indicate NE (Bora, close to 29% of the whole examined
392 database in the last cited paper) and SE (Sirocco, close to 3%) as the main
393 wind directions. 2003 had stronger winds than 2005, with more frequent Bora
394 events, particularly in the winter months, and less frequent Sirocco events in
395 the spring and summer months. These differences may have led to shorter
396 residence times in 2003, especially in areas with more extended wind fetch.

397 Sensitivity analysis confirmed that river discharge is the most important

398 factor for improving the accuracy of the models reproduction of salinity val-
399 ues. For this reason, averaged daily discharges were adopted as river inputs
400 for each of the 11 rivers included in the model. For both years, the discharge
401 data were collected from the Drainage Basin Authority. The data differ from
402 those of the DRAIN project in terms of the time and location of the mea-
403 surements. Analysis of monthly discharge data shows that maximum flows
404 generally occurred in February-March and October-November in all rivers,
405 whereas the low-water period was from June to September. Each river shows
406 inter-annual variability in its annual and monthly discharges. It is impor-
407 tant to note three aspects: i) total annual discharge in 2003 was less than in
408 2005 ($21 \text{ m}^3 \text{ s}^{-1}$ and $29 \text{ m}^3 \text{ s}^{-1}$ respectively); ii) for all rivers, annual mean
409 discharges in 2003 and 2005 were different, but not all rivers had lower dis-
410 charges in 2003 than in year 2005; iii) although the total annual discharge
411 for all rivers was lower in 2003 than in 2005, there were cases in which the
412 monthly discharge of the same river in the same month was higher in 2003
413 than in 2005, meaning that the variability of the discharge was higher in 2003
414 with respect to 2005.

415 Finally the correlation between river discharge and precipitation inside
416 the lagoon is low, showing that the freshwater inputs imposed in the model
417 are not redundant.

418 The salinity measurements for 2003 were collected at 28 stations pertain-
419 ing to the MELa project (Fig. 2, red circles) with monthly sampling during
420 ebb tide. The salinity data for 2005 (Fig. 2, green triangles) were collected
421 by the SAMANET automatic network (Ferrari et al., 2004) every 30 minutes
422 at 8 sampling points.

423 Comparison of salinity at the sampling points used for both 2003 and
424 2005 shows that the difference between years in terms of the annual average
425 and the annual maximum salinity is small. The most important differences
426 concern the annual minimum values and therefore the annual salinity range.
427 At some points the standard deviation is greater in 2003 than in 2005 because
428 this depends not only on the total quantity of fresh water but also on the
429 temporal distribution of the inputs.

430 The data were used to initialise the numerical model and to evaluate
431 the model's performance both spatially, at various sites in the lagoon, and
432 temporally, at high temporal resolution. The first of these steps ensures
433 that the model is representative of salinity throughout the lagoon, and the
434 second ensures that the model is able to reproduce the temporal variability
435 of salinity at each point.

436 *3.3. The typing process*

437 We considered the descriptors belonging to System B of the WFD: an-
438 nual mean salinity, annual standard deviation of salinity and annual mean
439 residence time, for the reasons set out in section 1.

440 The typing of the lagoon was carried out by defining classes (ranges) of
441 values for each considered variable and generating the corresponding maps.
442 Subsequently the classified maps of two or more variables were superimposed.
443 The resulting map shows areas characterised by different combinations of
444 classes for each considered variable.

445 Table 1 shows the defined classes and their ranges.

446 The annual mean salinity was divided into 4 classes, as in the Directive,
447 except that the two least saline categories were combined into one. The salin-

448 ity ranges are thus 0-5, 5-18, 18-30 and higher than 30, which coincide with
449 the intervals of the Venice System, and correspond to oligohaline, mesohaline,
450 polyhaline and euhaline respectively.

451 The classes for the annual standard deviation of salinity were defined after
452 analysing the distribution of values. The extreme standard deviation values
453 were excluded because they were not very frequent and most of them were
454 recorded in areas characterised by special conditions (such as salt marshes).
455 Given the distribution of values in the domain, we decided to divide the an-
456 nual standard deviation of salinity into 3 classes with ranges of 0-2 (low), 2-4
457 (medium), and higher than 4 (high). They represent the degree to which the
458 sampling point is characterised by the mixing of waters with differing salin-
459 ity. Thus, low standard deviation may be associated with stability, medium
460 standard deviation with moderate variability and high standard deviation
461 with high variability.

462 The calculation of residence time followed the method described in Cucco
463 and Umgiesser (2006). Residence time in the lagoon with real forcing factors
464 depends on the wind regime and ranges from more than a month to a few
465 days. Specifically, a long, strong Bora event can “clean” the basin very fast,
466 whereas a Sirocco event can slow the water renewal process by restricting
467 the outflow through the inlets. Long, strong Bora events happen frequently,
468 whereas Sirocco events are more isolated and spread out over the year and
469 are of long duration and strong intensity only in November, which is the
470 period characterised by “high water” phenomena. Wind data for both 2003
471 and 2005 followed this pattern. In order to evaluate the mean residence time
472 in an annual simulation under real forcing conditions, the residence time was

473 thus calculated every 2 months, corresponding to different real forcing condi-
474 tions. The average of the 6 replicates represents our assessment of the annual
475 mean residence time under real forcing conditions. The residence time ranges
476 considered are 0-5 days, 5-15 days and higher than 15 days, which may be
477 related to the “open”, “restricted” and “confined” classes respectively. The
478 upper and lower bounds of the ranges were chosen on the basis of geomorpho-
479 logical considerations: in both 2003 and 2005, the isoline of 15-day residence
480 time roughly coincided with the line of the salt marshes in the southern part
481 of the lagoon. In the northern part of the lagoon the lines still coincide,
482 but less precisely. Finally the isoline of a 5-day residence time marks the
483 limit of marine influence in the area of the lagoon around the inlets. The
484 combination of two variables with their respective classes gives rise to either
485 12 theoretical types (annual mean salinity with standard deviation of salin-
486 ity, annual mean salinity with annual mean residence time) or 9 theoretical
487 types (annual mean residence time with standard deviation of salinity). The
488 combination of all the variables gives rise to 36 theoretical types. The next
489 step is the simplification of the superimposed maps in accordance with the
490 size of the areas, followed by the assignment of each area to a specific type.

491 **4. Results and discussion**

492 *4.1. Spatial (MELa , 2003) and temporal (SAMA, 2005) variability*

493 The hydrodynamic model was calibrated in a calm period with reference
494 to water level data collected by ISPRA in the year 2003. Figure 3 shows the
495 comparisons of measured and simulated water levels at different points in
496 the lagoon and Table 2 shows the statistics calculated for each point, during

497 the whole simulation. The minimum error of the model is 2 cm and the
498 maximum error is 5 cm, with the error increasing from the seaward inlets to
499 the landward side of the lagoon.

500 Figure 4 shows a comparison of measured and modeled salinity timeseries
501 data at 6 stations in the year 2003. The statistics calculated for each sampling
502 point in 2003 are shown in Tab. 3. The position of each station is shown in
503 Fig. 2. The correlation coefficient ranges from 0.67 to 0.99 and the error of
504 the modeled salinity varies from a minimum of 0.4 to a maximum of 4.7. The
505 model overestimates values during the summer period, especially in the inner
506 north-central area, which extends beyond the city of Venice (cross-hatched
507 area in Fig. 2). This is probably a consequence of the uncertainty concerning
508 freshwater input, considering that only the main sources are included in the
509 model (without the discharges from less important channels, Venice city or
510 other human settlements on the islands) and that errors in the measured
511 discharges may be significant. The stratification of salinity may be significant
512 in the north-central area because of the interaction between river discharges
513 and the complex morphology of this area.

514 Stations 10B, 16B, 2B and 1B have higher root mean square error (RMSE)
515 values. The first two are behind the southern salt marsh line, where mixing
516 processes are more complex. Station 10B has a low correlation coefficient,
517 whereas 16B has a high correlation coefficient, indicating that in 10B the
518 freshwater inputs are not properly estimated, whereas this effect is less pro-
519 nounced in 16B. Stations 2B and 1B are situated in a complex system of river
520 inputs and salt marshes: station 1B has high variability because of freshwa-
521 ter inputs and station 2B is bordered by salt marshes in a very shallow area,

522 and its low correlation coefficient is the consequence of high evaporation and
523 the modulation of freshwater inputs by salt marshes.

524 Figure 5 and Table 4 show corresponding statistics for the 6 stations in
525 2005. The correlation coefficient ranges from 0.34 to 0.69 and the RMSE
526 varies from 1.7 to 7.7. The model reproduces the main pattern of variation,
527 but the variability of the measurements is greater than the simulated values.
528 Station 5, just off the industrial zone, is slightly underestimated, probably
529 because the model does not consider freshwater inputs from the zone itself.
530 Station 7 shows high RMSE values and is systematically underestimated:
531 this station is located in a channel near the mouth of the Dese river system
532 and there is probably a stratification effect that the model is not able to
533 reproduce in this application.

534 Annually averaged maps were calculated for each variable in each simula-
535 tion (Fig. 6). The main characteristic of each map is a transversal gradient,
536 which reflects the mixing processes of fresh and salt water. The standard
537 deviation of salinity increases from the sea to the land and from the seaward
538 inlets to the river mouths, and the residence time gradient is similar. Annual
539 mean salinity increases from the land to the sea.

540 The differences between the 2003 and 2005 maps are shown in the bottom
541 row of Fig. 6. They indicate the inter-annual spatial variability of each
542 parameter as determined by the model. Annual mean salinity in the year
543 2003 is greater than in the year 2005 (showing a positive difference in most
544 parts of the lagoon and a spatially averaged difference of nearly 3). This
545 result is in agreement with the lower annual rainfall and river discharge of
546 2003. In this case the spatial distribution of the differences is similar to

547 the spatial distribution of residence time, highlighting the role of mixing
548 processes. In most of the lagoon, the difference between 2003 and 2005 in
549 standard deviation of salinity is between -1 and 1, with a spatial average of
550 nearly 1. The difference is positive and higher than 1 in the northern part
551 of the lagoon and in isolated areas along the landward shore: this means
552 that the standard deviation in 2003 is greater than in 2005 in areas where
553 the effect of freshwater discharge is greater. This behaviour can be explained
554 by local freshwater discharges: in 2003 they were generally lower but more
555 erratic. The difference between 2003 and 2005 in terms of residence time
556 is both positive and negative, with spatially averaged values of 1.3 and -1.6
557 respectively. The residence time is longer (3-5 days) in 2003 than in 2005
558 in the northern part of the lagoon, mainly along the landward shore, where
559 the influence of river discharge is important. It is shorter in the central and
560 southern part of the lagoon (where the differences range from -1 to -3 days),
561 perhaps due to the different wind regime in the two years. It is important
562 to note that the difference in residence time indicates a basic division of
563 the lagoon into two parts: a northern basin, with positive differences, and a
564 south-central basin, with negative and less evident differences. The south-
565 central basin can in turn be divided by another strip of zero difference running
566 across the lagoon from the Malamocco inlet along its main channel (the most
567 important artificial channel in the lagoon)

568 *4.2. Proposed typologies and water bodies in the Venice lagoon*

569 A geographical analysis tool was used to superimpose the distribution of
570 two or three variables in 2003 and 2005. Comparison of the resulting maps
571 indicates that the spatial distribution of each type in the lagoon can change

572 noticeably: the surface area of a specific type may change or one type can
573 be replaced by another.

574 For example, Figure 7 shows the combination of annual average salinity
575 with residence time in 2003 (left panel) and in 2005 (right panel), and the
576 histogram of the log-transformed surface area of each possible type in the two
577 years. The numerical matrix under the histogram contains the numerical la-
578 bels of the 12 possible combinations of salinity and residence time classes.
579 The most extensive types correspond to the combination of the “open” class
580 with the “euhaline” class (14), the “restricted” class with the mesohaline,
581 polyhaline or euhaline salinity classes (22, 23, 24), and the “confined” resi-
582 dence time class with the “mesohaline, polyhaline or euhaline” salinity classes
583 (32, 33, 34). The histogram in the picture shows that the restricted meso-
584 haline and restricted polyhaline types (22, 23) and the confined mesohaline
585 and confined polyhaline types (32, 33) are more extensive in 2005, whereas
586 the others are less extensive. This is a consequence of the larger inputs of
587 freshwater in 2005. To simplify the number of combinations we subsequently
588 assimilated types with an area less than 10 km^2 to the most extensive ad-
589 jacent type. In our example this means that the restricted mesohaline type
590 (22) and the confined mesohaline type (32) were included in the restricted
591 polyhaline type (23). The partitioning obtained from the combination of
592 standard deviation of salinity with residence time is similar to the parti-
593 tioning derived from the combination of mean salinity with residence time,
594 indicating that standard deviation of salinity is important not in establishing
595 boundaries but in providing additional information about the stability of the
596 types. Because 2005 represents a typical year in terms of the annual aver-

597 ages of the climatic forcing factors, we assume that the types and the spatial
598 partitioning obtained from the combination of mean salinity with residence
599 time in that year can be taken as the reference situation. The next step is to
600 associate each defined type in the 2005 map with the corresponding standard
601 deviation class, in order to indicate its stability. This led to the identification
602 of 9 types (expressed as a combination of mean salinity, residence time and
603 standard deviation of salinity), spatially partitioning the Venice lagoon into
604 the water bodies schematically shown in Figure 8.

605 In this partitioning, three water bodies correspond to the areas near the
606 three inlets: two are of the “open euhaline stable” type, whereas the less
607 stable water body, corresponding to the area of the southernmost inlet, is
608 of the “open euhaline medium” type (note: Bellafiore and Umgiesser (2010)
609 showed that the Chioggia inlet is influenced by the coastal freshwater dis-
610 charge of the river Brenta, the mouth of which is near the inlet itself). The
611 most extensive water body in the lagoon, which might be divisible on the
612 basis of other factors not considered in our study, is the “restricted euhaline
613 medium” type. The extreme southern and northern parts of the lagoon are
614 divided into water bodies of specific types. The areas on the landward side
615 belong to the same types, although they are spatially separated. Our results
616 shows that it is possible to consider a hierarchical partitioning of the Lagoon
617 of Venice. As an initial approximation based on the broadest partitioning
618 criteria, our results indicate that the lagoon can be divided into an extensive
619 polyhaline subbasin and a reduced northern subbasin with specific charac-
620 teristics. This division reflects the results obtained by M. Sigovini and D.
621 Tagliapietra (unpublished data), which identify most of the Venice lagoon

622 as microtidal, except for the northern part which appears to be nanotidal
623 (mean tidal range less than 0.5 m). From a more detailed point of view, the
624 Venice lagoon can be divided into 14 water bodies. This partitioning reflects
625 some aspects of the study of (Molinaroli et al., 2009), which is based on the
626 division of the Venice lagoon into the classical four sub-basins. The north-
627 ern sub-basin (A), identified as still in a quasi-natural condition, contains
628 water bodies of 7 different types, making it the most complex sub-basin.
629 The northern-central and southern-central sub-basins (B and C) correspond
630 to the most disturbed areas of the lagoon and include water bodies of 5
631 different types. The southernmost sub-basin (D), which is partly still in a
632 semi-natural condition, includes water bodies of 4 types.

633 5. Conclusions

634 We developed a model which is able to reliably reproduce the spatial and
635 temporal evolution of salinity in most parts of the Venice lagoon, and thus
636 to provide a good assessment of its variability. The model is also able to
637 calculate the residence time and takes into account the inter-annual vari-
638 ability of the studied parameters. Most of the data used by the model are
639 available via the usual monitoring programmes and thus, with little eco-
640 nomic effort, this numerical tool offers support for lagoon management on
641 various levels, in terms of both WFD requirements and other applications.
642 The model makes it possible to tackle several open questions concerning the
643 management strategies of transitional environments, such as:

- 644 1. How to sub-divide a basin into water bodies. Local authorities often as-
645 sume a division of a basin into distinct water bodies without explaining

646 the objective criteria adopted for the zoning. The method developed
647 in this study can be applied to different lagoons and provides unbi-
648 ased and objective zoning indications for the basin. A numerical model
649 simulating the abiotic factors can be adopted as a tool for designing
650 monitoring programs, showing the position and the size of the types in
651 different years. Taking figure 7 as an example, it is possible to identify
652 which type accounts for the largest portion of the lagoon, or alterna-
653 tively, which type is most likely to shift from a dry year to a standard
654 year (unstable). On the other hand, the model can be employed to
655 estimate the variation of salinity associated with input of water from
656 the drainage basin, which generally contains a high concentration of
657 nutrients and pollutants derived from human activities. This knowl-
658 edge, together with knowledge of the residence time, can be used as
659 an operational tool to evaluate the response of water quality elements
660 (including biological elements), helping to distinguish natural from an-
661 thropogenic stresses.

- 662 2. How to manage the spatial and temporal variability of descriptors in
663 transitional waters. Interannual variation in the annually averaged val-
664 ues of the parameters is considerable, and depends on the meteorolo-
665 gical and hydrological characteristics of the year in question. The
666 resulting variability of types and their spatial distribution is signifi-
667 cant, and the typology of the system could be regarded as changing
668 from year to year. This means that a given water body can belong to
669 one type in one year and to a different type in another year, in other
670 words that not only the borders of the water bodies are fuzzy, but their

671 types too. This could be a problem for managers, since water bodies are
672 the prescribed unit for management, monitoring and the achievement
673 of quality targets, and are assumed to belong to a fixed type, which
674 is not always true. The model can solve this problem by identifying a
675 variable that indicates the stability of each type, or by detecting when-
676 ever the type itself shifts from one class combination to another. This
677 aspect is important when establishing the reference status of a water
678 body, since the Directive does not consider the inter-annual variability
679 of types in transitional environments.

680 3. Finally, this study demonstrated that the tool can also be used to per-
681 form a hierarchical division of a lagoon. Thus, according to the pur-
682 pose, either approximate or finely detailed typologies can be adopted,
683 for example to select the adequate number of sampling stations for
684 monitoring.

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ACCEPTED MANUSCRIPT

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Salinity	PSU	Std.Dev. S		Residence time		days
Class	range	Class	range	Class	range	
oligohaline	0-5	stable	0-2	open	0-5	
mesohaline	5-18	medium	2-4	restricted	5-15	
polyhaline	18-30	unstable	> 4	confined	> 15	
euhaline	> 30					

Table 1: Classes of salinity, standard deviation of salinity and residence time.

station	1	2	3	4	5	6	7	8	9
name	Le Saline	Torcello	Pagliaga	Punta Salute	Fusina	Torson di Sotto	Vigo	Petta di Bo'	Settemorti
n	2621	2621	2621	873	2621	2621	2621	2621	2621
r	0.99	0.99	0.99	1.00	1.00	0.99	1.00	0.99	0.99
RMSE	0.03	0.03	0.04	0.02	0.03	0.05	0.03	0.04	0.05
BIAS	-0.02	-0.01	0.01	-0.01	0.02	0.04	-0.02	0.01	-0.02
SI	0.14	0.11	0.15	0.09	0.09	0.18	0.11	0.16	0.17

Table 2: Comparison of measured and simulated water level data at various points of domain in 2003. n=number of records, r=linear correlation coefficient, RMSE=root mean square error, BIAS=difference between mean of observations and simulations and SI=scatter index, calculated as the RMSE normalized with observed mean.

station	n	r	RMSE	BIAS	SI	station	n	r	RMSE	BIAS	SI
B01	13	0.95	3.71	2.15	0.15	B15	12	0.93	1.03	-0.23	0.03
B02	13	0.67	3.72	-1.31	0.12	B16	11	0.82	3.70	0.36	0.12
B03	13	0.93	1.24	0.49	0.04	B17	13	0.97	1.52	-1.08	0.05
B04	13	0.88	2.10	-1.04	0.07	B18	13	0.96	0.60	0.04	0.02
B05	12	0.93	1.61	-0.48	0.05	B19	12	0.90	1.11	0.56	0.03
B06	13	0.93	1.49	-0.78	0.05	B20	12	0.93	2.54	-1.77	0.08
B07	13	0.89	2.01	-1.30	0.06	C1	12	0.96	1.57	0.94	0.05
B08	12	0.74	2.60	-1.43	0.08	C2	13	0.95	2.66	1.14	0.09
B09	13	0.91	2.05	-1.42	0.06	C3	12	0.90	1.63	0.00	0.05
B10	13	0.62	4.69	0.14	0.17	C4	13	0.98	0.66	0.38	0.02
B11	13	0.94	0.83	-0.42	0.02	C5	13	0.89	1.12	-0.42	0.03
B12	13	0.98	0.43	-0.17	0.01	C6	13	0.81	2.26	0.68	0.07
B13	13	0.92	0.81	-0.41	0.02	C7	13	0.99	0.83	0.55	0.02
B14	12	0.85	1.05	0.25	0.03	C8	13	0.95	1.18	0.51	0.03
mean		0.90	1.81	-0.14	0.06						

Table 3: Comparison of measured and simulated salinity data at various points of domain in 2003. n=number of records, r=linear correlation coefficient, RMSE=root mean square error, BIAS=difference between mean of observations and simulations and SI=scatter index.

Station	n	r	RMSE	BIAS	SI
1	12804	0.61	2.21	-0.50	0.07
2	16541	0.64	2.53	-0.60	0.09
3	15192	0.69	1.31	0.52	0.04
4	14766	0.64	1.88	0.68	0.06
5	11116	0.34	3.96	1.15	0.13
6	12886	0.64	1.71	0.83	0.05
7	12768	0.57	7.75	6.06	0.35
mean	13724	0.60	3.05	1.17	0.11

Table 4: Comparison of measured and simulated salinity data at various points of domain in 2005. n=number of records, r=linear correlation coefficient, RMSE=root mean square error, BIAS=difference between mean of observations and simulations and SI=scatter index.

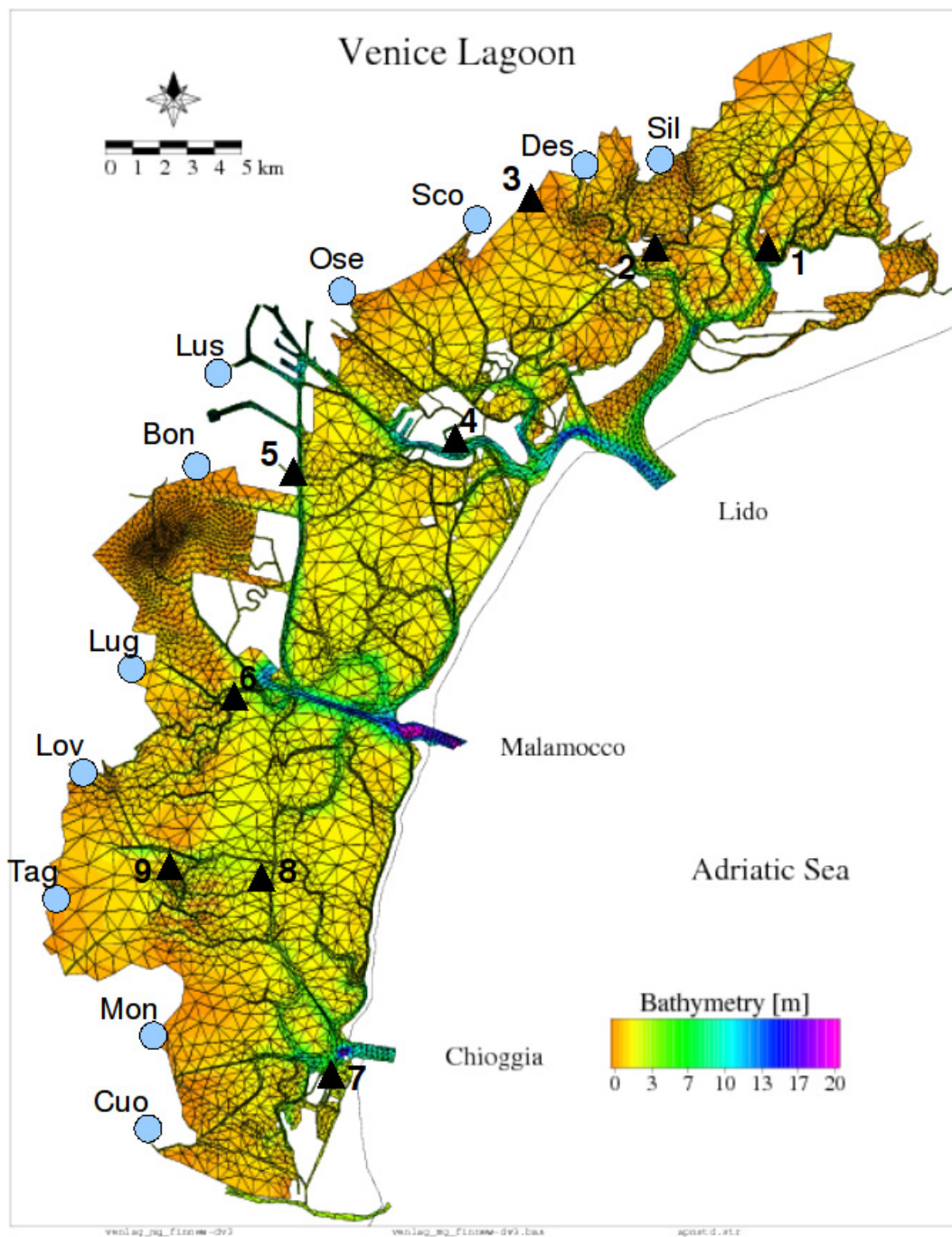


Figure 1: Venice lagoon, numerical grid, bathymetry, rivers and APAT tide gauges.

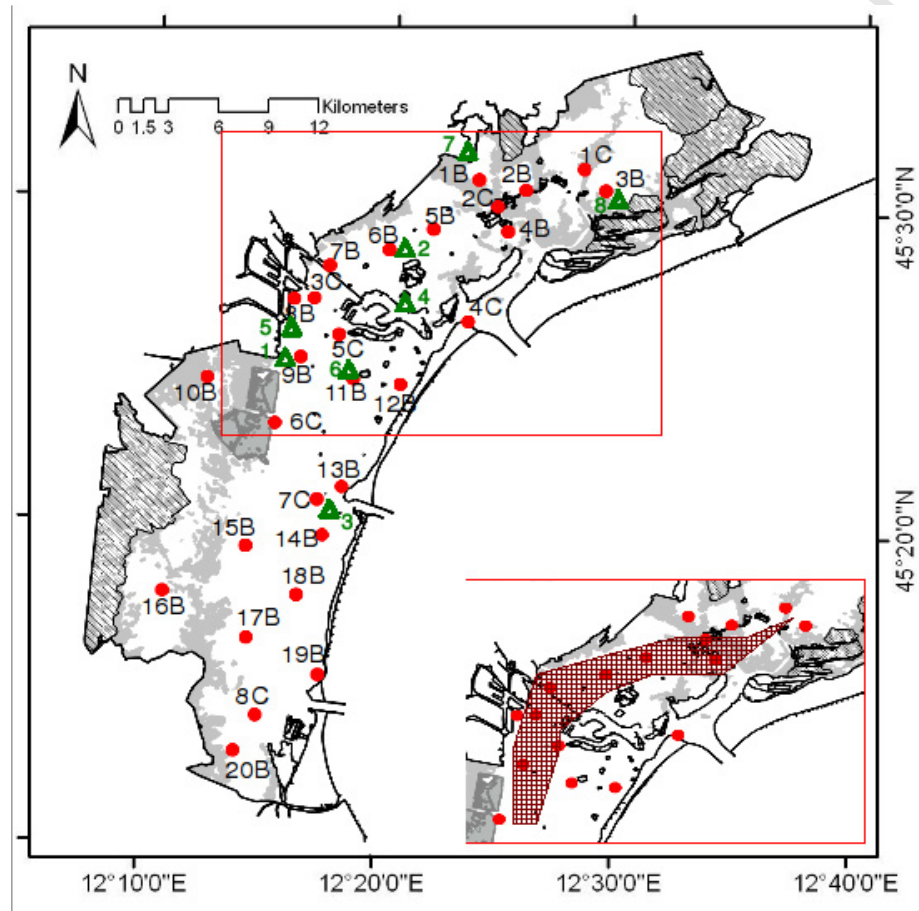


Figure 2: MELa stations (red circles) and SAMA stations (green triangles). Cross-hatching in close-up shows area where model overestimates salinity.

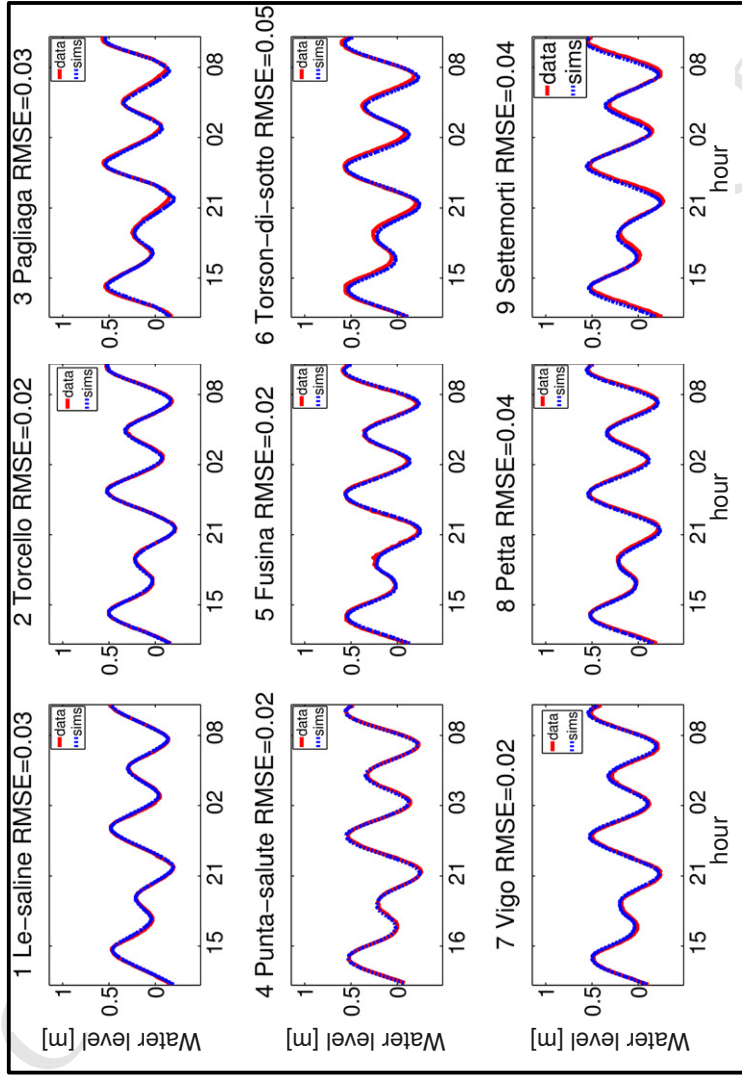


Figure 3: Comparison of measured and simulated water levels in 9 stations shown in Fig. 1

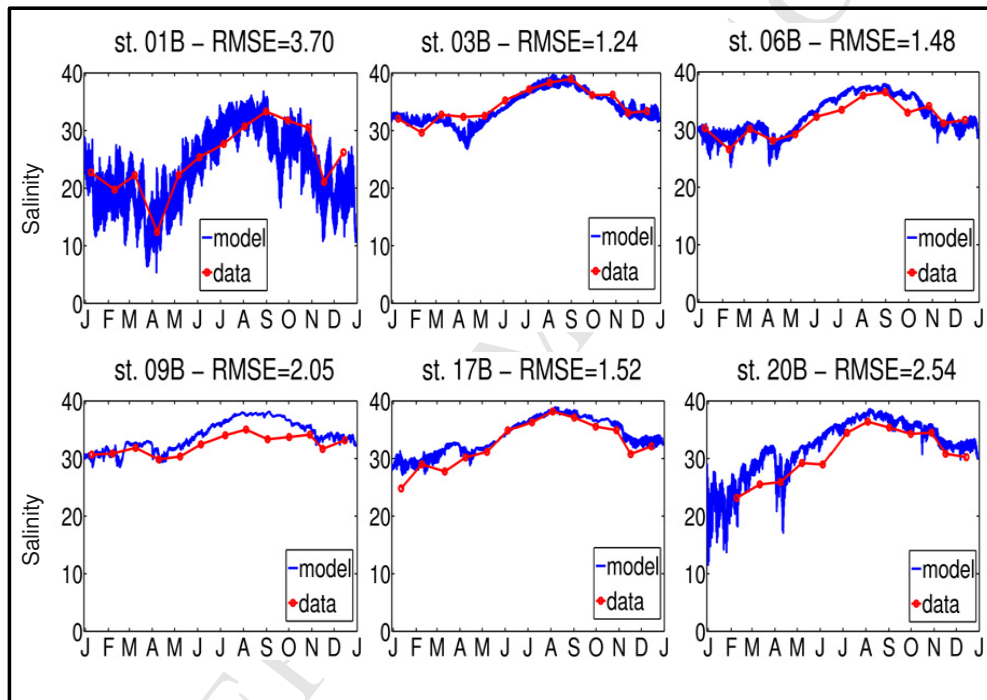


Figure 4: Comparison of measured and modeled salinity in 2003.

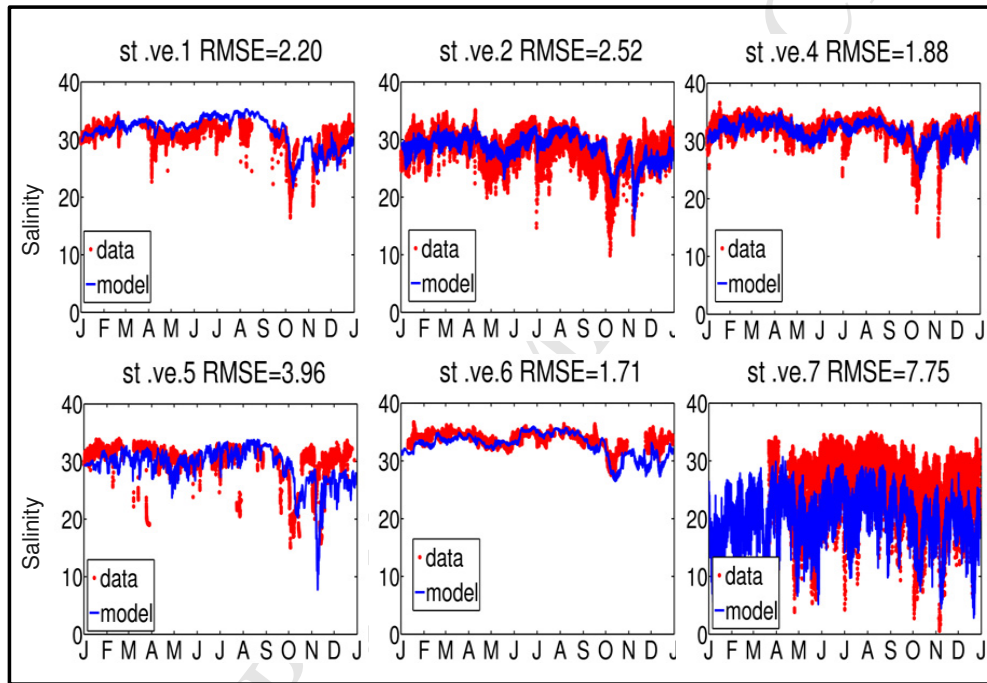


Figure 5: Comparison of measured and modeled salinity in 2005. Stations 1, 2 and 7 of SAMA monitoring network are close to stations B09, B06 and B01 of MELa monitoring project shown in Fig. 4.

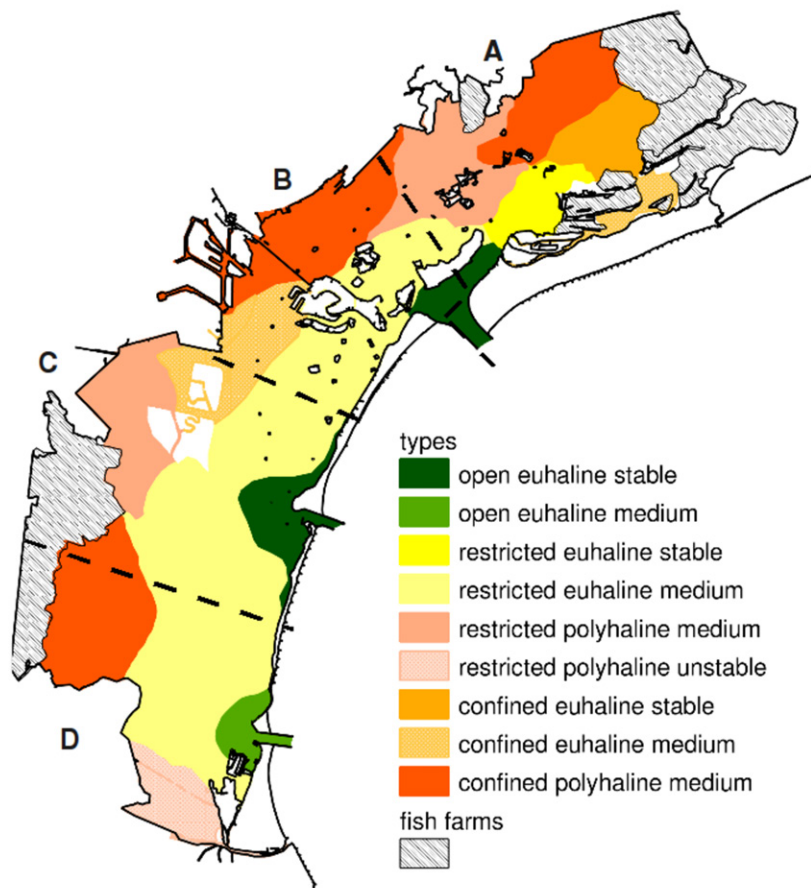


Figure 8: Comparison of types and water bodies identified in this study with the 4 sub-basins as in Molinaroli et al. (2009).