

1 **Supplementary Material**

2 Colloidal mobilization of arsenic from mining-affected soils by surface runoff

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15 *(5 Tables, 8 Figures)*

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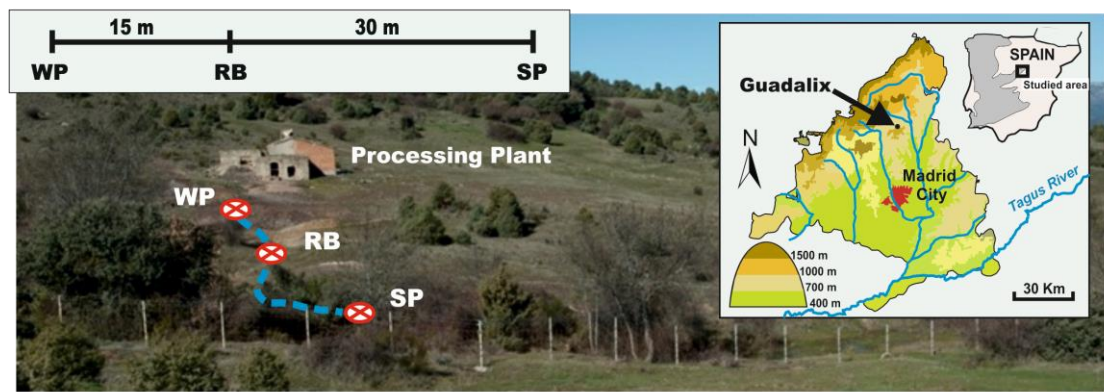
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26 **S1. Experimental locations and characterization procedures**

27 Rainfall simulations tests were performed at the following locations: A) the arsenic-
28 bearing waste pile (WP); B) the river bed (RB) of a small stream (~ 1 m wide) that
29 seasonally collects surface runoff from WP; and C) the sediment that accumulates
30 downstream in an artificial sedimentation pond (SP) (Figure S1) (Recio-Vazquez et al.,
31 2011).

32 Bulk soils from WP, RB and SP were characterized as follows: Texture was determined
33 by the pipette method after removing soil organic matter (Gee and Bauder, 1986). Soil
34 pH and electrical conductivity (EC) were measured in deionized water (1:5 m/m
35 suspensions), using a Thermo Orion model 920 A⁺ pH meter and a Thermo Orion
36 model 125 A conductivity meter, respectively. Total organic carbon (TOC) was
37 determined by wet digestion (Walkley and Black, 1934). Exchangeable bases (Ca, Mg,
38 Na, and K) (Shuman, 1985) were extracted with 1 M NH₄OAc at pH 7 (Thomas, 1982),
39 and exchangeable Al was extracted with 1 M KCl (Barnishel and Bertsch, 1982). Semi-
40 quantitative mineralogical composition of the total (≤ 2 mm) fraction of the WP bulk
41 sample was previously identified by powder X-ray diffraction (XRD) with a Philips
42 PW-1710/00 diffractometer using the CuK α radiation with a Ni filter and a setting of 40
43 kV and 40 mA. Samples were carefully milled over a period of 15 min and pressed to
44 produce pellets of powdered aliquots. XRD analyses were performed using XPOWDER
45 software. Patterns were obtained by step scanning, from 3° to 65° 2 θ , with a count for
46 0.5 s/step exploration speed of 7°/min and 40 kV and 40 mA in the X-ray tube. The
47 qualitative search-matching procedure was based on the ICDDPDF2 and the DIFDATA
48 databases. Those results are presented in Table 1.



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50 Figure S1 – Location of the studied area and relative positions of the experimental
 51 zones: Waste-pile (WP), river-bed (RB) and sedimentation-pond (SP)

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53 S2. Rainfall simulation

54 S2.1. Rainfall simulator design

55 The sprinkling-type portable rainfall simulator to use in rugged terrain conditions and
 56 has been used for decades in a semi-arid environments with good results. The rainfall
 57 simulator is an improvement of the design of Calvo et al. (1988), modified and
 58 optimized by Cerda et al. (1997). The drop distribution generated by this rainfall
 59 simulation presents a drop-size between 2.49-2.53 mm of diameter. In all cases, the
 60 height of the outlet nozzle was set at 175 cm from the soil surface generating a drop
 61 velocity of $3.1 \text{ m}\cdot\text{s}^{-1}$ and a mean drop-size equal to 2.5 mm in diameter as tabulated by
 62 Cerda et al. (1997). The basic components of the rainfall simulator are shown in Figure
 63 S2 and described as follows:

- 64 - Upper structure and nozzle: The structure consists of a square platform of 40x40
 65 cm. A cylindrical pipe which crossed the square structure presents a nozzle in the
 66 middle. The nozzle is a HARDI-1553-10, and has installed a filter with a mesh of
 67 0.3 mm, a piece to increase the size of the drops and a diffuser. Four telescopic
 68 connections of 3.45 m long are connected with the square structure.

- 69 - Pumping system: The water was supplied manually for giving a better stability to
70 the water pressure. The low water flow required (28 mm h^{-1}) allow to use hand air-
71 compressed pumps (Matabi, Kima 12). Distilled water was used to perform the
72 rainfall simulation processes.
- 73 - Wind protector: An important part of the rainfall simulator is the wind protector,
74 which is fitted to the upper square structure. This protection prevents possible
75 evaporation of the smaller drops and the disturbances caused by the wind.
- 76 - Sampling plot: The rainfall area generated by the nozzle is slightly larger than 1 m^2 .
77 In order to avoid border interference, a 0.24 m^2 experimental plot made of
78 galvanized iron (0.55 m of diameter) was used for the measurements.
- 79



80
81 Figure S2 – Rainfall simulator description: View of the rainfall simulation with the
82 pumping system and the wind protector (a), upper structure and nozzle (b), and
83 experimental plot (c).

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87 **S2.2. Rainfall conditions**

88 Distilled water was used for the rainfall simulation experiments. Homogeneous rainfall
89 distribution was achieved at a water pressure of 1.5 kg cm^{-2} (Cerdeira et al., 1997) and a
90 rain intensity of 28 mm h^{-1} . The telescopic legs were height adjusted to overcome slopes
91 of 7%, 5% and 2% at the waste-pile, river-bed and sediment-pond sampling locations,
92 respectively. In all cases, the height of the outlet nozzle was set at 175 cm from the soil
93 surface (Figure S3).

94 Rainfall simulation experiments were carried out during 60 minutes to assure a well
95 representative storm event in agreement with the typical rainfall in the area (according
96 to AEMET data). Runoff suspensions were continuously collected in 5-min integration
97 intervals in the river-bed and sedimentation pond, and in 10-min integration intervals on
98 the waste pile using acid-cleaned polypropylene bottles of 0.5 L of volume. The
99 samples were stored at 4°C until further analyses. The rainfall simulation experiments
100 were performed in October 2011, after at least 15 dry days.

101

102 **S2.3. Experimental locations**

103 The rainfall simulations were performed over the three different experimental zones
104 described previously. A real view of each experiment is presented in Figure S3.



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106 Figure S3 – Rainfall simulation locations:(a) waste-pile, (b) river-bed and (c) the
107 sedimentation pond.

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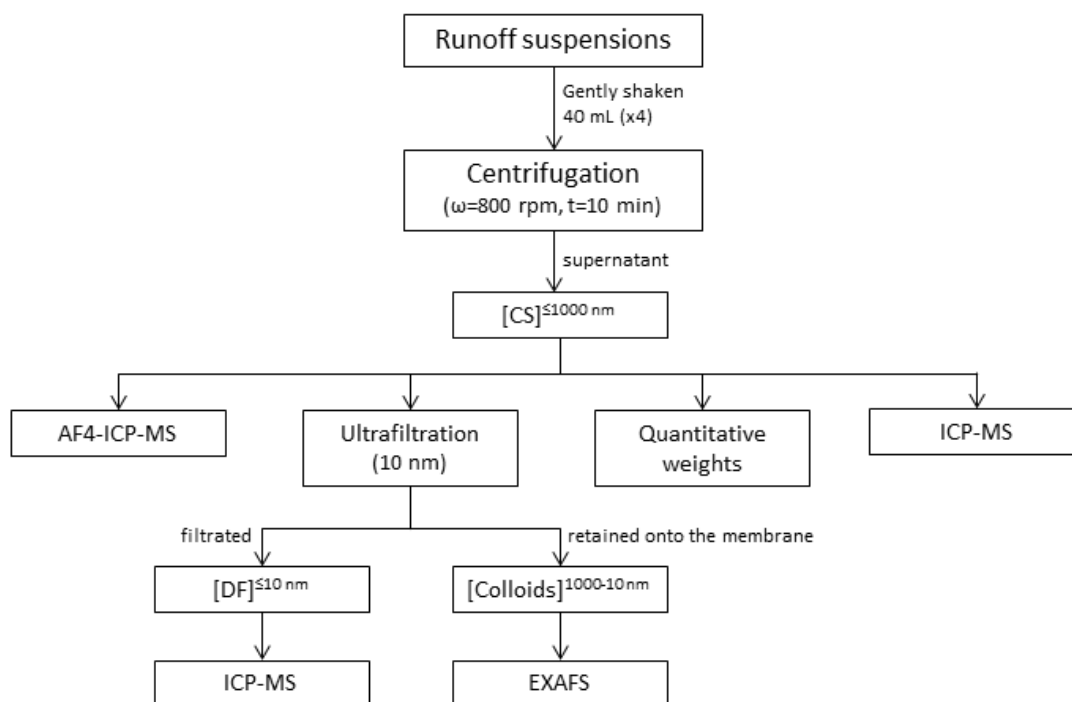
109 **S3. Isolation of colloid-suspensions from runoff**

110 **S3.1. Experimental procedure**

111 In order to separate the colloid-suspension (CS, ≤ 1000 nm), solid colloids (1000-10 nm
112 and the dissolved fraction (DF, < 10 nm), runoff samples were subjected to the protocol
113 described in the Figure S4. These samples were gently shaken and homogenized before
114 four 40 mL aliquots were transferred into four acid-cleaned polypropylene vessels and

115 subsequently centrifuged at 800 rpm during 10 minutes. According to the equations for
116 size-based separation (Bolea et al., 2010), the resulting supernatants (~20 mL)
117 corresponded to their CS. These CS were pipeted to the polypropylene vessels and
118 stored at 4°C. One CS aliquot was used for AF4-ICP-MS analyses within the following
119 48 hours. The second CS aliquot was ultrafiltered through ultrafiltration membranes
120 (*Pall Filtron, Microsep Omega, 10-nm pore size*) to determine the As, Fe, Al, Cu, Zn
121 and Pb concentrations in the dissolved fraction (DF) by ICP-MS or ICP-OES. Colloids
122 with a size range of 1000-10 nm, deposited on the ultrafiltration membranes were stored
123 at 4°C. The colloids from the waste-pile zone (WP) were analyzed by As and Fe K-edge
124 EXAFS and the colloids from sedimentation-pond zone (SP) by XANES spectroscopy.
125 The third CS aliquot was used to quantify the mass of colloids contained, as described
126 by Plathe et al. (2010) but modified as follows: 10 mL of the CS were placed onto pre-
127 weighted glass vessels along with 6 mg of NaCl (*'suprapure', Merck*). The mixtures
128 were shaken for 30 minutes and centrifuged at 4000 rpm for 15 min. The resulting
129 supernatant was discarded and the colloidal mass, coagulated by the effect of the NaCl,
130 were dried at 60°C during 24 hours, and finally weighted.

131 Lastly, ten milliliters of the four CS aliquot were dissolved in two digestion microwave-
132 assisted steps by a microwave program at 200°C during 15 min (Ethos Series 1,
133 *Milestone*): Step I – 11.5 mL of HF/HNO₃/HCl (volume ratio 1.5:0.75:3.5) and Step II –
134 9 mL of H₃BO₃ (5 %). Solutions from the digestion process were filtered and measured
135 by ICP-OES (Iris & Intrepid Radial, *Thermo Fisher Scientific*) for Fe, Al, Cu, Zn and
136 Pb; and by ICP-MS (ELAN DRC-e, *Perkin Elmer*) for As quantification.



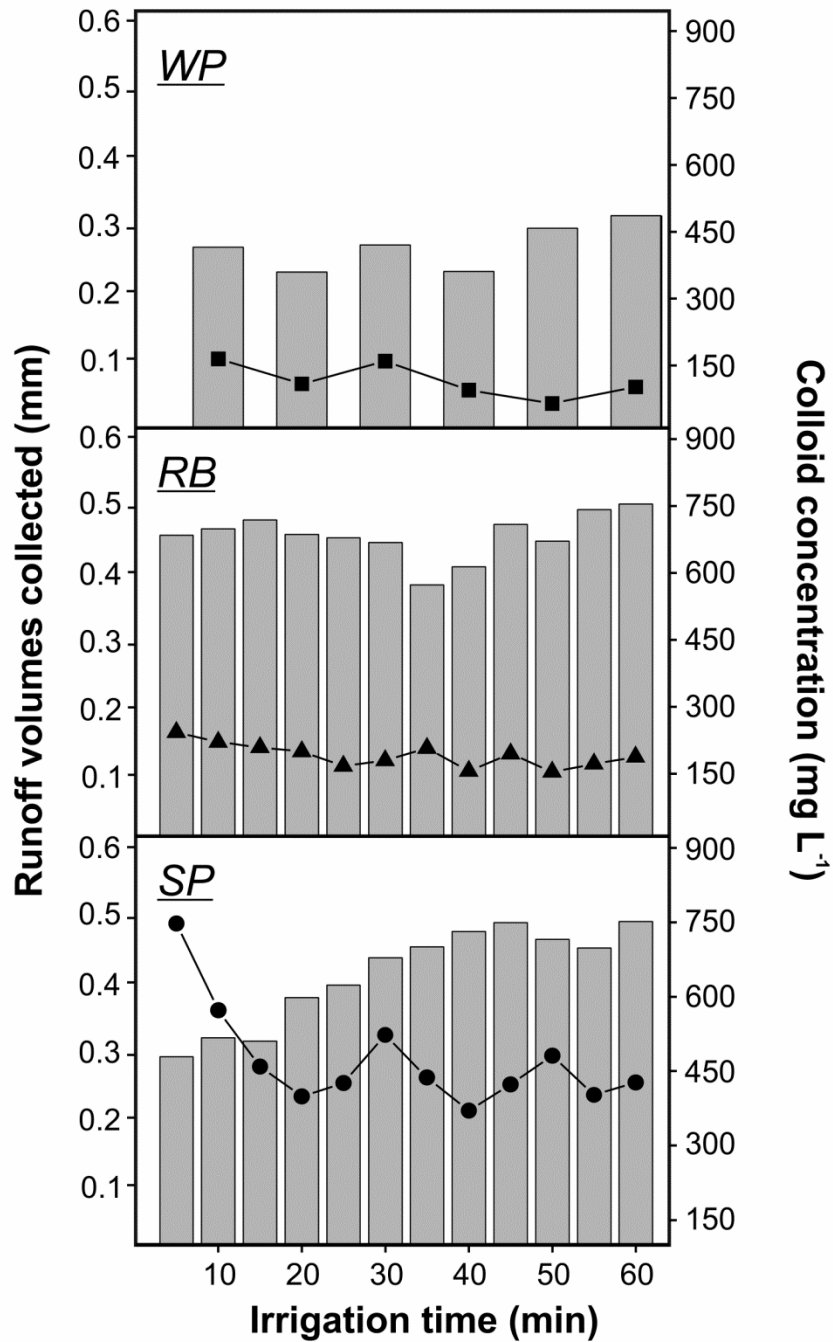
137

138 Figure S4 – Procedure for isolation and characterization of colloid-suspensions (CS, ≤
 139 1000 nm) from original runoff suspensions. Colloids (1000-10 nm) were isolated using
 140 ultrafiltration membranes (10 nm) and the dissolved fraction (DF < 10 nm) was
 141 subsequently measured by ICP-MS

142

143 S3.2. Runoff volumes and colloid concentration

144 All runoff suspensions were collected and weighted. The volumes of runoff suspensions
 145 collected during the rainfall simulations (Table 1) and the colloid concentration of the
 146 corresponding colloid suspensions (Table 2) are shown in Figure S5.



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148 Figure S5 – Runoff volumes collected (bars, left axis) and colloidal concentration of
 149 colloid suspensions (lines and markers, right axis) at waste-pile (WP), river-bed (RB)
 150 and sedimentation-pond (SP).

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154 **S4. ICP-OES and ICP-MS analyses**

155 The pseudo-total concentration of the samples was measured by ICP-OES after aqua
156 regia + microwave digestion (Chen and Ma, 2001). All the chemicals used for sample
157 preservation, analysis and reagent preparation were of analytical quality or higher.
158 Deionized water ($18 \text{ M}\Omega \text{ cm}^{-1}$, MilliQ +, *Millipore Corp.*) was used for all solutions
159 and dilutions. Arsenic determinations were performed by ICP-MS (ELAN DRC-e,
160 *Perkin Elmer*) while Fe, Al, Cu, Zn and Pb quantification were done by ICP-OES (Iris
161 & Intrepid Radial, *Thermo Fisher Scientific*). For ICP analyses, calibration curves were
162 run before and after each sample series (20 samples including matrix-matched blanks
163 and in-between calibration checks). The calibration solutions covered the range of
164 concentration in the samples and were prepared in the same matrix as the extracting
165 reagents from certified stock solutions. Sample blanks were analyzed for correction of
166 background effect on instrument response. Trace metal standards were used to assess
167 instrument precision. We calculated metal concentrations in unknown solutions on the
168 basis of the external calibration averaging the concentrations from two repetitions for
169 each experimental replication. Limits of detection were calculated at three standard
170 deviation of the instrument response from 10 repeated analyses of sample matrix-
171 matched blank solutions. Results from ICP-OES and ICP-MS analyses are presented in
172 Tables 3 and S1.

173 The companion metal concentrations to As and Fe are listed in Table S2. The
174 aluminium concentrations in the colloid suspension and in the dissolved fraction
175 followed a different trend than Fe. At the waste-pile zone, Al concentration was low,
176 averaging $225 \mu\text{g L}^{-1}$ in the colloid suspension and below the quantification limit (QL)
177 in the dissolved fraction at all sampling times. Instead, at the river-bed zone, Al
178 concentrations were higher than those of Fe and decreased with irrigation time from

179 1900 $\mu\text{g L}^{-1}$ to 624 $\mu\text{g L}^{-1}$ and 1320 $\mu\text{g L}^{-1}$ to 456 $\mu\text{g L}^{-1}$ in the colloid suspension and in
 180 the dissolved fraction, respectively. In the sedimentation-pond zone, Al content in the
 181 dissolved fraction was below the QL and that in the colloid suspension increased with
 182 irrigation time from 1030 to 1700 $\mu\text{g L}^{-1}$ unlike the Fe one. Copper, Zn and Pb
 183 concentrations showed different distribution patterns with the irrigation time at each
 184 sampling zone. Essentially, the importance of colloidal versus dissolved mobilization
 185 decreases in the order Pb (~100%) > Cu (~60%) > Zn (~0%) at the waste-pile zone.
 186

187 Table S1 – Companion metal concentrations of selected colloid (≤ 1000 nm)
 188 suspensions (CS) and dissolved fractions (DF) ≤ 10 nm (\pm standard deviation, $n=3$ ^a)

Sample	Time	Fraction	Al	Cu	Zn	Pb
			$\mu\text{g L}^{-1}$ ^{b,c}			
WP	20 min	CS	223±92	116±31	854±50	68±41
		DF	< QL	74±47	922±64	< QL
		Colloids	223 ^d	42	-	68
WP	40 min	CS	255±90	115±48	523±48	75±38
		DF	< QL	70±39	555±59	< QL
		Colloids	255	45	-	75
WP	50 min	CS	198±114	107±51	362±55	63±33
		DF	< QL	70±42	372±69	< QL
		Colloids	198	37	-	63
RB	20 min	CS	1900±105	1760±34	1340±57	< Q
		DF	1320±97	2040±46	1460±67	< QL
		Colloids	580	-	-	< QL
RB	35 min	CS	1470±97	996±46	745±50	< QL
		DF	567±83	1040±38	766±63	< QL
		Colloids	903	-	-	< QL
RB	50 min	CS	624±108	817±40	840±46	< QL
		DF	456±102	858±55	757±60	< QL
		Colloids	168	-	-	< QL
SP	20 min	CS	1030±103	63±52	317±61	< QL
		DF	< QL	< QL	65±62	< QL
		Colloids	1030	63	252	< QL
SP	35 min	CS	1260±86	44±39	155±54	< QL
		DF	< QL	< QL	102±66	< QL
		Colloids	1260	44	53	< QL
SP	50 min	CS	1700±124	20±46	114±58	< QL
		DF	< QL	< QL	63±61	< QL
		Colloids	1700	20	51	< QL

189 ^a Standard deviation was calculated by triplicate determination of the same aliquot

190 ^b Elemental concentration was expressed on ng of metal per liter of colloid suspension (CS)

191 ^c Quantification limits (QL) for ICP-OES measurements were: Al = 100 $\mu\text{g L}^{-1}$, Cu = 12 $\mu\text{g L}^{-1}$, Zn = 23 $\mu\text{g L}^{-1}$,
 192 Pb = 31 $\mu\text{g L}^{-1}$

193 ^d Metal concentrations of colloids (1000-10 nm) were defined as the difference between the concentration of the CS
 194 and the DF

195

196 **S5. Asymmetric flow field-flow fractionation**

197 **S5.1. AF4 characteristics**

198 The channel dimensions were 27.5 cm in length and from 2 to 0.5 cm in width. The
199 spacer used for all the measurements had 350 μm thickness. The accumulation wall
200 consisted of a 1 kDa polyethersulfone (PES) (*Postnova Analytics*). Novachem surfactant
201 (0.01 %, w/v, *Postnova Analytics*) adjusted to pH = 4.5 was used as carrier solution.
202 Carrier was degassed prior to use by an on-line vacuum degasser. The out flow was 0.8
203 mL min^{-1} and a sample loop of 100 μL was used throughout. Colloid-suspensions and
204 calibration standards were detected using a UV-vis diode array detector (*Shimadzu*,
205 wavelength range: 200 to 650 nm). The AF4-UV-vis system was coupled to an ICP-MS
206 in order to perform an on-line multi-element quantification. Samples were introduced
207 into the ICP-MS using a glass concentric slurry nebulizer and a cyclonic spray chamber
208 (*Glass Expansion*). An internal standard solution of 50 ng mL^{-1} Rhodium (*Merck*) was
209 merged with the carrier at 0.3 mL min^{-1} for correction of instrumental drift. The
210 instrumental conditions of the AF4-ICP-MS system are presented in Table S2.

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230 Table S2 – AF4 operating conditions and ICP-MS instrumental and data acquisition
 231 parameters

Crossflow program of AF4			
<i>Carrier</i>	<i>Time (min)</i>	<i>Crossflow type</i>	<i>Crossflow (mL min⁻¹)</i>
Novachem 0.01%, pH = 4.5 ^a	35	Constant	0.1
	3	Linear Decay	0
	2	Constant	0
Instrumental parameters of ICP-MS			
RF power	1200 W		
<u>Argon gas flow rate</u>			
Plasma	15 L min ⁻¹		
Auxiliary	1.2 L min ⁻¹		
Nebulizer	0.9 L min ⁻¹		
Data acquisition parameters of ICP-MS			
Measuring mode	Peak hopping		
Points per spectral peak	1		
Dead time	60 ns		
Sweeps	10		
Dwell time	5 ms		
Integration time per point	50 ms		
Settle time	3 ms		

232 ^a Outflow = 0.8 mL min⁻¹

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235 **S5.2. AF4 principles and channel calibration**

236 According to the AF4 theory, in normal mode it is possible to establish a relationship
 237 between a retention parameter (λ) and the diffusion coefficient (D) of the eluted specie,
 238 which can be related to their hydrodynamic diameter (d) using the Stokes-Einstein
 239 equation (Assemi et al., 2004).

240 Since the colloid-suspension (CS) have a size lower than 1000 nm, the AF4 will be
 241 operated only in normal mode, and the relationship between the retention ratio and size
 242 could be obtained using a set of size standards (Schimpf et al., 2000), assuming a
 243 similar behavior between the size standards and the particles to be characterized.

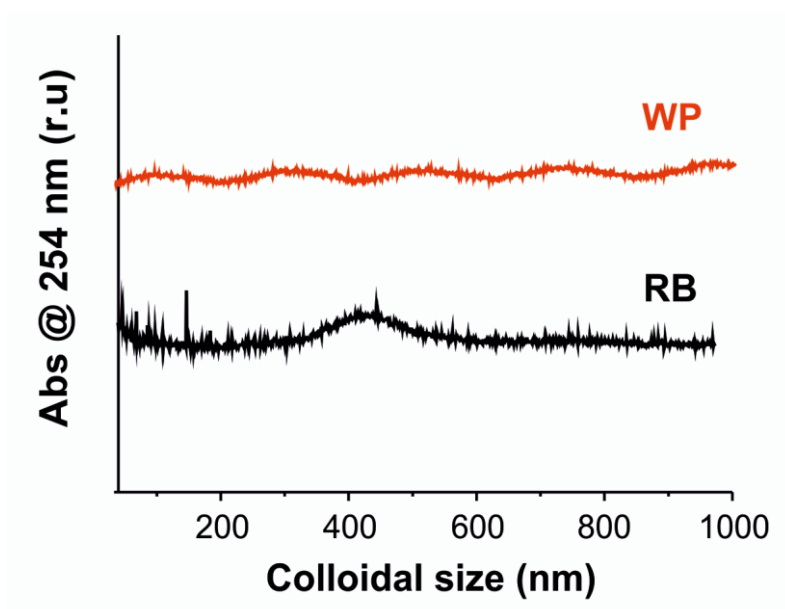
244 Monodispersed size standards of silicon dioxide [SiO₂] of 0.020 ± 0.004 , 0.15 ± 0.03
 245 and 0.50 ± 0.05 μm were used for AF4 calibration. These standards were used because
 246 of its similar nature to the colloid samples. The calibration curve obtained for normal
 247 mode was: $\log R = -2.066 - 0.934 \log d$ ($r = 0.991$), where d is the particle diameter in
 248 μm , and $R = t_0/t_r$, where t_0 is the elution time corresponding to the void peak and t_r is the

249 retention time for a given particle. Diluted solutions (20 mg L^{-1}) of these standards
250 (*Sigma-Aldrich*) were prepared by further dilution with the corresponding carrier.

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252 **S5.3. AF4-ICP-MS measurements**

253 Within the following 48 hours after performing the rainfall simulation tests, all the
254 colloid ($\leq 1000 \text{ nm}$) suspensions were isolated and measured by AF4-ICP-MS. The
255 colloid suspensions from the waste-pile and the river-bed could not be measured by
256 AF4, as presented in the Figure S6.

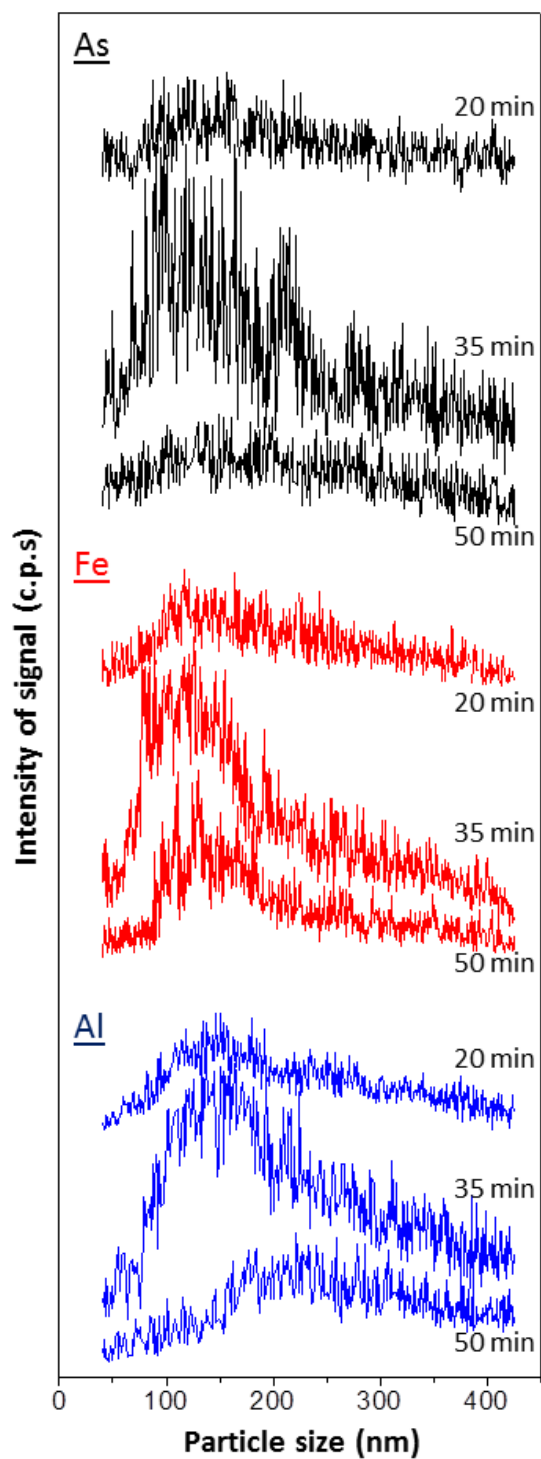


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258 Figure S6 – AF4 fractograms of colloid suspensions from waste-pile (WP) and river-
259 bed (RB) at 20 minutes

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261 The particle size distribution and associated As, Fe and Al concentrations of colloid
262 suspensions at 20, 35 and 50 minutes in sedimentation-pond zone were calculated by
263 AF4-ICP-MS (Figure S7). The maxima concentration peaks were identified by fitting a
264 lognormal function and calculating the maximum and confidence interval at 95 %
265 probability (Figure 1b).



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267 Figure S7 – Arsenic, iron and aluminum AF4-ICP-MS bulk fractograms of the colloid

268 suspensions from sedimentation-pond collected at 20 minutes, 35 minutes and 50

269 minutes.

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272 S4.4. Recovery calculations

273 The recovery of the AF4 determination is a parameter which provides additional
274 information. Reference calculations were made using three injections of sample without
275 applying any crossflow in the AF4 channel (no injection step was applied either). Then,
276 three injections were performed applying the crossflow program described in Table S3.
277 Afterwards, peaks areas were calculated from signal and the recoveries expressed as:
278 $R (\%) = S/S_0 \cdot 100$, where S is the signal area obtained when a crossflow is applied, and
279 S_0 is the signal area obtained with no crossflow. The AF4-ICP-MS recoveries of the
280 colloid suspensions from sedimentation-pond are presented in Table S3.

281
282 Table S3 – Recoveries of colloid suspensions from sedimentation-pond (SP) and target
283 metals obtained by AF4-ICP-MS measurements ^a

DCF	Irrigation time	AF4-ICP-MS recovery (%)		
		<i>As</i>	<i>Fe</i>	<i>Al</i>
<i>SP</i>	<i>20 min</i>	23	48	58
	<i>35 min</i>	25	65	51
	<i>50 min</i>	17	48	39

284 ^a Three AF4 injections were averaged to calculate the recoveries shown

285
286 Taking into account that approximately three quarters of total As found in the
287 sedimentation-pond are presented in the dissolved fraction ($DF \leq 10$ nm), arsenic
288 recoveries ranging 17-25 % are in agreement with the As presented in the colloid-
289 suspension (Table 3). Iron and Al recoveries are comparable to those found in similar
290 studies by AF4-ICP-MS (Serrano et al., 2015), as reported by Neubauer et al. (2013)
291 who explained that lower recoveries than expected are generally associated with
292 adsorption processes of the species injected onto the permeation membrane of the
293 channel, or with losses by filtration through the membrane of those species with
294 molecular weight lower than its pore size (< 1 kDa).

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297 **S6. X-ray absorption spectroscopy**

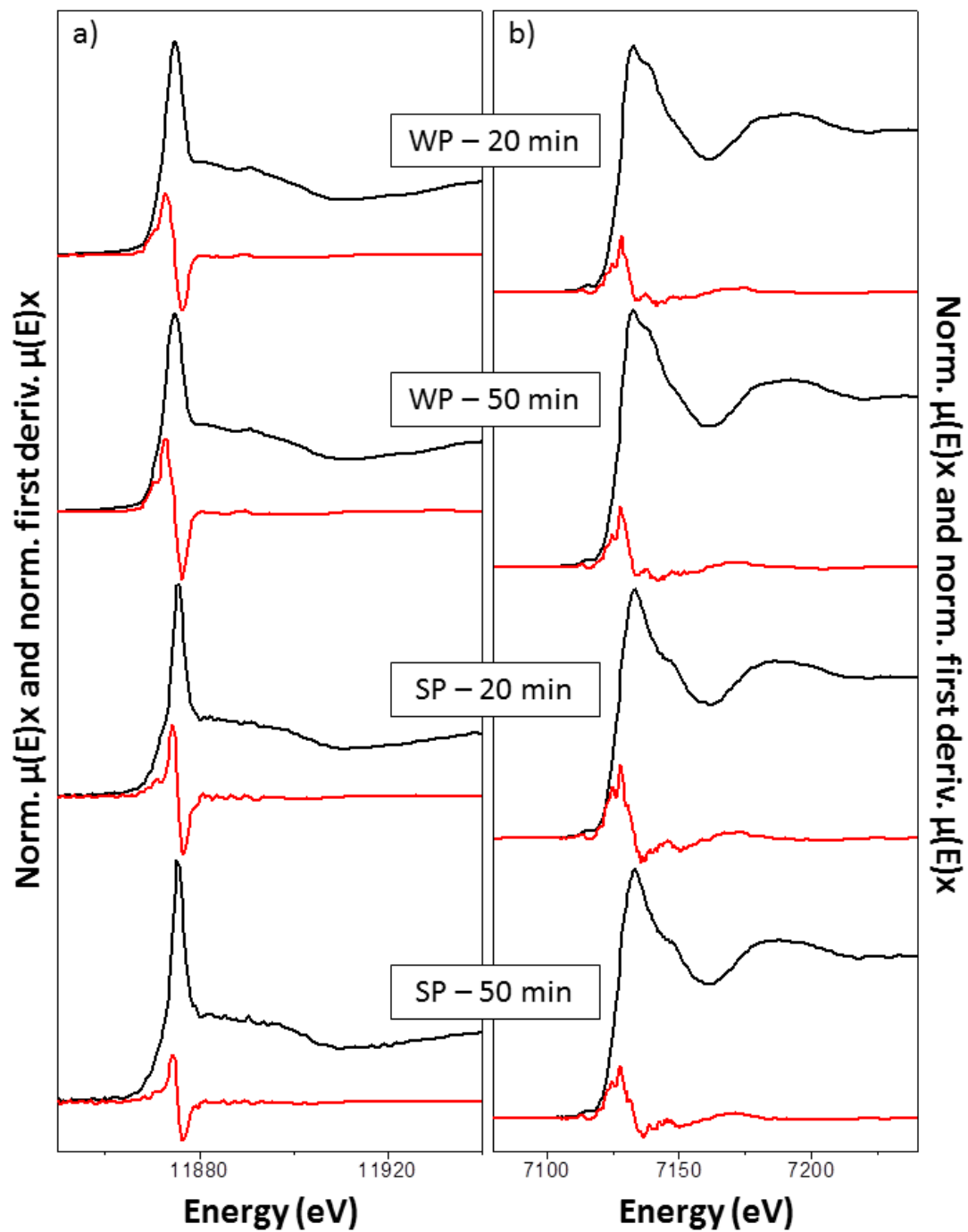
298 **S6.1. Data analysis**

299 The colloid-bearing ultrafiltration membranes were placed on holders made of polyether
300 ether ketone (PEEK) material and sealed with Kapton® tape. Arsenic and Fe K-edge
301 EXAFS spectra were recorded at the bending magnet BM25A beamline at the European
302 Synchrotron Radiation Facility (ESRF, Grenoble, France) (6 GeV, 100 mA, Si(111)
303 monochromator crystals) at room temperature (RT) using a 13-element Ge(Li) solid
304 state detector. Additional As EXAFS spectra from the SP colloids were recorded at
305 beamline 22 of the ALBA Synchrotron Facility (Barcelona, Spain) (3 GeV, 400 mA,
306 Si(311) monochromator) at 80°K using a CdTe solid state detector. The beam energy
307 was calibrated by setting the first inflection point in the K absorption edge of a metallic
308 Fe foil to 7112 eV (Fe measurements) or the first maximum in the K absorption edge of
309 KH_2AsO_4 (*Sigma-Aldrich*) to 11875 eV (As measurements). The sample spectra were
310 acquired in fluorescence mode, starting at 6950 eV (Fe) or 11650 eV (As). The pre-edge
311 step size was set to 5 eV, and the edge step size along the edge was set to 0.5 eV. The
312 EXAFS spectra were collected up to 9 \AA^{-1} for Fe and 10.5 \AA^{-1} for As respectively, using
313 a step size of 0.05 \AA^{-1} in k-space with constant measurement times over the entire
314 EXAFS range. The sample spectra were obtained by averaging several replicate scans
315 (10-12 scans at ESRF; 3-4 at ALBA).

316 Arsenic and iron normalized K-edge XANES spectra were initially compared with the
317 reference compounds and presented in Figure S8 together with their first derivatives.

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321 Figure S8 – Normalized As (a) and Fe (b) K-edge XANES spectra (black line) and their

322 first derivatives (red line) of colloids (1000-10 nm) from waste-pile (WP) and

323 sedimentation-pond (SP) at 20 minutes and 50 minutes of runoff.

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326 Firstly, the entire library (sixteen spectra for Fe, twelve spectra for As) was screened to
327 determine combinations of reference spectra that best matched the data. Subsequently,
328 the k^3 -weighted EXAFS spectra of waste-pile (WP) colloids were analyzed by linear
329 least-squares combination fittings (LCF) over the k-range 2-8 Å⁻¹ for Fe and 2-11 Å⁻¹
330 for As, while the XANES spectra of sedimentation-pond (SP) colloids were also
331 analyzed by LCF over the range - 30 eV and + 60 eV for both As and Fe (Table 4). On
332 the basis of their occurrence and relevance in those preliminary fits, the following
333 reference spectra were included in the final fits which allowed to reproduce all sample
334 EXAFS spectra: An arsenate (scorodite) and As(V) sorbed onto Fe(III)-(hydr)oxides
335 (ferrihydrite) were selected for As LCF analyses; while we used different Fe(III)-
336 (hydr)oxides (e.g., scorodite, ferrihydrite, hematite), different Fe-bearing phyllosilicates
337 (e.g., smectite, nontronite) and a Fe(III)-(oxy)sulfides (schwertmannite) in the Fe
338 fittings.

339 Reference and sample spectra were processed and analyzed by linear combination
340 fitting (LCF) using the software code Athena (Ravel and Newville, 2005). The E_0 was
341 fixed at 7128.5 eV for Fe and 11875 eV for As. The background was subtracted using
342 a linear fit through the pre-edge region and the Autobk routine in Athena for the spline
343 fit through the EXAFS region ($R_{\text{bkg}} = 0.8$ for As and 0.9 for Fe, k -weight = 3, spline k -
344 range = 0.5 – 10.5 Å⁻¹ for As and 0.5 – 8.5 Å⁻¹ for Fe).

345 In the case of As, starting from the best fit with one component, the number of
346 components n was increased as long as the normalized sum of the squared residuals
347 ($\text{NSSR} = \sum(\text{data}_i - \text{fit}_i)^2 / \sum(\text{data}_i)^2$) of best $n + 1$ -component fit was at least 10% lower
348 than the NSSR of the best n -component fit. Otherwise, in the Fe LCF, the best $n + 1$ -
349 component fit was considered to be significantly better than the best n -component fit, if
350 its NSSR was at least 20% lower and if no component account for less than 5% of total

351 Fe. Linear combination fittings were not constrained to sum 100% (Table 3). More
 352 details of reference compound spectra, including their synthesis as well as the reference
 353 publications, are shown in Table S4.

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Table S4 – As and Fe reference compounds used for linear combination fitting

As standards				
Name	Group	Type	Beamline	Source
Scorodite	Arsenate	Natural	BM 4-3, SSRL	Savage et al. (2005)
As(V) sorbed to ferrihydrite	Fe(III) oxide	Synthetic	BM 4-3, SSRL	Root et al. (2009)
Fe standards				
Name	Group	Type	Beamline	Source
Scorodite	Oxide	Natural	BM 4-3, SSRL	Savage et al. (2005)
Ferrihydrite	Oxide	Synthetic	XAS BM, ANKA	Voegelin et al. (2010)
Hematite	Oxide	Natural	BM 7-3, SSRL	O'Day et al. (2004)
Schwertmannite	Oxy-sulfate	Synthetic	BM 4-1, SSRL	measured by S. Hayes ^a
Smectite	Phyllosilicate	Natural	BM 2-3, SSRL	O'Day et al. (2004)
Nontronite	Phyllosilicate	Natural	XAS BM, ANKA	Gorski et al. (2013)

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^a Unpublished reference EXAFS spectrum

359 **S6.2 Shell-fittings**

360 Least-squares fitting of k^3 -weighted Fe and As K-edge spectra were then performed in
 361 Artemis (Ravel and Newville, 2005), only for the waste-pile colloids collected at 20 and
 362 50 minutes (Table S5). Theoretical single and multiple scattering paths used to model
 363 both the Fe and the As K-edge EXAFS spectra were calculated from the structure of
 364 scorodite (Kitahama et al., 1975), using FEFF 8.2 (Ankudinov and Rehr, 2000).
 365 The As K-edge EXAFS spectra were fitted in R -space over a distance $R + \Delta R$ of 0.8 –
 366 3.6 Å (k -weigh = 3, Kaiser-Bessel window sill = 2 Å⁻¹). Two single scattering (SS)
 367 paths were used to model the spectra: As-O and As-Fe. Multiple scattering (MS) within
 368 AsO₄ tetrahedra was accounted for by a three-legged triangular As–O–O MS path
 369 (degeneracy = 12). According to Mikutta et al. (2013), the addition of four-legged MS
 370 paths involving As and O atoms were of minor importance to model the As K-edge
 371 EXAFS of scorodite. In addition, the low data quality of collected EXAFS spectra

372 supported the use of only the triangular As–O–O MS path. While the degeneracies of
373 the MS path were fixed to their theoretical value, their half path lengths were expressed
374 as a function of the SS As–O half path length assuming an ideal tetrahedron. The half
375 path length of the triangular As–O–O MS path were set to 1.8165 ($=1 + \sqrt{(2/3)}$) times
376 the As–O SS half path length. The Debye-Waller parameters for the MS paths were
377 constrained by considering the correlation between the lengths of individual legs
378 (Hudson et al., 1996). Assuming that the lengths of the As–O and the O–As leg of the
379 triangular As–O–O MS path are not correlated, their contribution to the σ^2 of MS half
380 path length equals half the σ^2 of As–O SS path. However, because also the O–O leg
381 contributes to the σ^2 of the MS path, we assumed the σ^2 of the MS path to be equal to
382 the σ^2 of the As–O SS path (Voegelin et al., 2007).

383 The k^3 -weighted Fe K-edge EXAFS spectra were Fourier-transformed over the k-range
384 $2.5 - 10.5 \text{ \AA}^{-1}$ using a Kaiser-Bessel window (sill width = 2.5 \AA^{-1}). The fits were
385 performed in R -space over a distance $R + \Delta R$ of $0.9 - 3.8 \text{ \AA}$. Shell fits include two SS
386 paths for the first and second coordination shells of Fe (Fe–O, Fe–As). In addition, one
387 MS within FeO_6 octahedra was accounted for: a triangular Fe–O–O MS path
388 (degeneracy = 24, $\sigma^2 = \sigma^2$ (Fe–O SS)). The half path length of the triangular Fe–O–O
389 MS path were set to 1.7071 ($=1 + \frac{\sqrt{2}}{2}$) times the Fe–O SS half path length (Voegelin et
390 al., 2010).

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396 Table S5 – Shell-fit results for As and Fe K-edge EXAFS spectra of colloids (1000-10
 397 nm) isolated from colloid suspensions sampled after 20 and 50 minutes of rainfall in
 398 waste-pile (WP) zone. Theoretical simple and multiple scattering paths were calculated
 399 from the structure of scorodite (Kitahama et al. 1975) using FEFF 8.2 (Ankudinov and
 400 Rehr, 2000)

As shell-fit parameters ^a								
Sample	Path	N ^b	σ^{2c} (Å ²)	R ^d (Å)	S ₀ ^{2e}	ΔE^f (eV)	R-factor ^g	red γ^{2h}
WP 20 min	As-O	4	0.0016	1.68	1.06	3.54	0.0405	2.77·10 ⁶
	As-Fe	2	0.0021	3.36				
	MS1 ⁱ	12	0.0064	3.08				

WP 50 min	As-O	4	0.0019	1.69	1.06	3.39	0.0390	3.88·10 ⁶
	As-Fe	2	0.0021	3.36				
	MS1 ⁱ	12	0.0077	3.08				
Fe shell-fit parameters ^a								
Sample	Path	N ^b	σ^{2c} (Å ²)	R ^d (Å)	S ₀ ^{2e}	ΔE^f (eV)	R-factor ^g	red γ^{2h}
WP 20 min	Fe-O	6	0.0074	1.98	1.02	-2.57	0.0545	3.97·10 ⁵
	Fe-As	4	0.0074	3.35				
	MS2 ^j	24	0.0074	3.46				

WP 50 min	Fe-O	6	0.0079	1.98	1.02	-2.84	0.0517	4.24·10 ⁵
	Fe-As	4	0.0068	3.35				
	MS2 ^j	24	0.0079	3.46				

401 ^aThe fit range was set to 2.5 - 10.5 Å

402 ^bDegeneracy (coordination number for single scattering paths), values in bold were fixed during shell fitting

403 ^cDebye-Waller parameter, uncertainty for As-O ± 0.0010-0.0016 Å², for As-Fe ± 0.0032-0.0040 Å², for Fe-O ±
 404 0.0017-0.0022 Å²

405 ^dHalf path length (inter-atomic distance for single scattering paths), uncertainty for As-O ± 0.013-0.015 Å, for As-Fe
 406 ± 0.0033-0.0037 Å, for Fe-O ± 0.013-0.016 Å

407 ^eAmplitude correction factor, constrained to the same value for all paths in a simultaneous fit

408 ^fEnergy shift, constrained to the same value for all paths in a simultaneous fit, uncertainty < ± 1.2 eV

409 ^gNormalized sum of the squared residuals of the fit ($R = \sum(\text{data-fit})^2 / \sum \text{data}^2$)

410 ^hReduced γ^2 (Stern et al. 1995)

411 ⁱMS1 = Triangular As-O-O MS path within As(V) tetrahedron, degeneracy = 12

412 ^jMS2 = Triangular Fe-O-O MS path within Fe(III) octahedron, degeneracy = 24

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414 **S7. References**

- 415 Ankudinov AL, Rehr JJ. Theory of solid-state contributions to the x-ray elastic
416 scattering amplitude. *Physical Review B* 2000; 62: 2437-2445.
- 417 Assemi S, Newcombe G, Hepplewhite C, Beckett R. Characterization of natural organic
418 matter fractions separated by ultrafiltration using flow field-flow fractionation.
419 *Water Research* 2004; 38: 1467-1476.
- 420 Barnishel R, Bertsch PM. *Methods of Soil Analysis. Part 2. Chemical, Microbiological*
421 *Properties*. Aluminum 1982; American Society of Agronomy: 275-300.
- 422 Bolea E, Laborda F, Castillo JR. Metal associations to microparticles, nanocolloids and
423 macromolecules in compost leachates: Size characterization by asymmetrical
424 flow field-flow fractionation coupled to ICP-MS. *Analytica Chimica Acta* 2010;
425 661: 206-214.
- 426 Calvo A, Gisbert J, Palau E, Romero M. Eds. M. Sala and F. Gallart 1988; SEG
427 Monogr.: 6-15, Spain.
- 428 Cerda A, Ibanez S, Calvo A. Design and operation of a small and portable rainfall
429 simulator for rugged terrain. *Soil Technology* 1997; 11: 163-170.
- 430 Chen M, Ma LQ. Comparison of three aqua regia digestion methods for twenty Florida
431 soils. *Soil Science Society of America Journal* 2001; 65: 491-499.
- 432 Gee GW, Bauder JW. *Methods of Soil Analyses. Part 1. Physical and mineralogical*
433 *methods* 1986; Particle-size analyses. Madison, WI.: 383-411.
- 434 Gorski CA, Klüpfel LE, Voegelin A, Sander M, Hofstetter TB. Redox Properties of
435 Structural Fe in Clay Minerals: 3. Relationships between Smectite Redox and
436 Structural Properties. *Environmental Science & Technology* 2013; 47: 13477-
437 13485.

438 Hudson EA, Allen PG, Terminello LJ, Denecke MA, Reich T. Polarized x-ray-
439 absorption spectroscopy of the uranyl ion: Comparison of experiment and
440 theory. *Physical Review B* 1996; 54: 156-165.

441 Kitahama K, Kiriya R, Baba Y. Refinement of crystal-structure of Scorodite. *Acta*
442 *Crystallographica Section B-Structural Science* 1975; 31: 322-324.

443 Mikutta C, Mandaliev PN, Kretzschmar R. New Clues to the Local Atomic Structure of
444 Short-Range Ordered Ferric Arsenate from Extended X-ray Absorption Fine
445 Structure Spectroscopy. *Environmental Science & Technology* 2013; 47: 3122-
446 31.

447 Neubauer E, von der Kammer FD, Hofmann T. Using FLOWFFF and HPSEC to
448 determine trace metal colloid associations in wetland runoff. *Water Research*
449 2013; 47: 2757-2769.

450 O'Day PA, Vlassopoulos D, Root R, Rivera N. The influence of sulfur and iron on
451 dissolved arsenic concentrations in the shallow subsurface under changing redox
452 conditions. *Proceedings of the National Academy of Sciences of the United*
453 *States of America* 2004; 101: 13703-13708.

454 Plathe KL, von der Kammer F, Hasselov M, Moore J, Murayama M, Hofmann T, et al.
455 Using FIFFF and aTEM to determine trace metal-nanoparticle associations in
456 riverbed sediment. *Environmental Chemistry* 2010; 7: 82-93.

457 Ravel B, Newville M. Athena and Artemis: interactive graphical data analysis using
458 IFEFFIT. *Physica Scripta*. 2005; T115: 1007–1010.

459 Root RA, Vlassopoulos D, Rivera NA, Rafferty MT, Andrews C, O'Day PA. Speciation
460 and natural attenuation of arsenic and iron in a tidally influenced shallow
461 aquifer. *Geochimica Et Cosmochimica Acta* 2009; 73: 5528-5553.

462 Savage KS, Bird DK, O'Day PA. Arsenic speciation in synthetic jarosite. *Chemical*
463 *Geology* 2005; 215: 473-498.

464 Schimpf ME, Caldwell K, Giddings JC. *Field-Flow Fractionation Handbook*. In: Wiley
465 & Sons (ed) 2000.

466 Serrano S, Gomez-Gonzalez MA, O'Day PA, Laborda F, Bolea E, Garrido F. Arsenic
467 speciation in the dispersible colloidal fraction of soils from a mine-impacted
468 creek. *Journal of Hazardous Materials* 2015; 286: 30-40.

469 Shuman LM. Fractionation method for soil microelements. *Soil Science* 1985; 140: 11-
470 22.

471 Thomas GW. *Methods of Soil Analysis, Part 2. Chemical and Microbiological*
472 *Properties. Exchangeable cations* 1982; Soil Science Society of America: 159-
473 166.

474 Voegelin A, Kaegi R, Frommer J, Vantelon D, Hug SJ. Effect of phosphate, silicate,
475 and Ca on Fe(III)-precipitates formed in aerated Fe(II)- and As(III)-containing
476 water studied by X-ray absorption spectroscopy. *Geochimica Et Cosmochimica*
477 *Acta* 2010; 74: 164-186.

478 Voegelin A, Weber FA, Kretzschmar R. Distribution and speciation of arsenic around
479 roots in a contaminated riparian floodplain soil: Micro-XRF element mapping
480 and EXAFS spectroscopy. *Geochimica Et Cosmochimica Acta* 2007; 71: 5804-
481 5820.

482 Walkley A, Black IA. An examination of the Degtjareff method for determining soil
483 organic matter, and a proposed modification of the chromic acid titration
484 method. *Soil Science* 1934; 37: 29-38.

485