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# Impact of human intervention and climate change on natural flow regime

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## Abstract

According to the ‘natural flow paradigm’, any departure from the natural flow condition will alter the river ecosystem. River flow regimes have been modified by anthropogenic interventions and climate change is further expected to affect the biotic interactions and the distribution of stream biota by altering streamflow. This study aims to evaluate the hydrologic alteration caused by dam construction and climatic changes in a mesoscale river basin, which is prone to both droughts and monsoonal floods. To analyse the natural flow regime, 15 years of observed streamflow (1950-1965) prior to dam construction is used. Future flow regime is simulated by a calibrated hydrological model Soil and Water Assessment Tool (SWAT), using ensemble of four high resolution (~25 km) Regional Climate Model (RCM) simulations for the near future (2021-2050) based on the SRES A1B scenario. Finally, to quantify the hydrological alterations of different flow characteristics, the Indicators of Hydrological Alteration (IHA) program based on the Range of Variability Approach (RVA) is used. This approach enables the assessment of ecologically sensitive streamflow parameters for the pre- and post-impact periods in the regions where availability of long-term ecological data is a limiting factor. Results indicate that flow variability has been significantly reduced due to dam construction with high flows being absorbed and pre-monsoon low flows being enhanced by the reservoir. Climate change alone

29 may reduce high peak flows while a combination of dam and climate change may significantly  
30 reduce variability by affecting both high and low flows, thereby further disrupting the  
31 functioning of riverine ecosystems. We find that, in the Kangsabati River basin, influence of dam  
32 is greater than that of the climate change, thereby emphasizing the significance of direct human  
33 intervention.

34

35 *Keywords: Anthropogenic impact, climate change, flow alteration, IHA, RCM, SWAT*

## 37 **1 Introduction**

38 Flow regime alteration of important seasonal flow components, such as high flows and low flows,  
39 by anthropogenic activities, especially large dams, has generated immense scientific interest with  
40 regards to implications for riverine ecosystems, biodiversity conservation and invasion by non-  
41 native species (Bunn and Arthington, 2002; Lytle and Poff, 2004; Meijer et al. 2014).  
42 Degradation of ecological health is now associated with the downstream section of dams (Poff  
43 and Zimmermann 2010; Suen 2011). Carlisle et al. (2010) reported that across regions and  
44 anthropogenic conditions, biological impairment is directly related to the magnitude of  
45 streamflow reduction. Moreover, regulation of river flow and alteration of flood and drought  
46 timing is expected to favour species that spawn during certain times (Freeman et al. 2001).

47 Along with direct anthropogenic impacts, human-induced climate change is also expected to  
48 affect the hydrologic cycles and thereby alter natural flow characteristics. Increasing  
49 temperatures will directly increase evaporation and alter plant transpiration rates, thereby  
50 reducing runoff (Bates et al. 2008). Doll and Zhang (2010) have shown that by mid-21<sup>st</sup> century,  
51 climate change effect on flow regimes may be greater than that caused by dams and water  
52 withdrawals. Global analysis of potential changes in runoff regimes shows that by the year 2050,  
53 most regions will experience significant changes in hydrological regime (Arnell and Gosling  
54 2013). Changes brought about by climate change will interact with existing anthropogenic  
55 factors and thus cause additional stress to riverine ecosystems (Fung et al. 2013; Ravazzani et al.  
56 2015).

57 Much needed interaction between scientists from hydrological, ecological and geomorphological  
58 foci over the past 20 years has increased our understanding of riverine dynamics, which is an  
59 essential prerequisite for gauging future implications of human actions. Such studies typically  
60 require long-term monitoring and assessment of baseline conditions to benchmark the effect of  
61 changes (Wagener et al. 2010). However, increasingly, resources for developing such  
62 quantitative understanding and data are declining (Mishra and Coulibaly 2009). Shifting baseline  
63 conditions due to human intervention has added to the existing issue of insufficient ecological  
64 information (Wagener et al. 2010). Historically, insufficient resources for regular survey and  
65 assessment of ecological conditions of riverine systems have been a significant limitation for  
66 carrying out change detection studies in developing countries.

67

68 In this study of the Kangsabati River basin, we address two important research gaps related to  
69 natural flow regime alteration; (i) effect of anthropogenic activity (damming) and future climate  
70 change for a mesoscale river basin with a strong monsoonal influence on hydrology and (ii)  
71 usage of sparse and scattered ecological data to derive inferences regarding potential impacts of  
72 damming and climate change on riverine ecosystem. We quantify observed alterations in the  
73 flow regime due to damming and then model the ramifications of climate change using the  
74 conventional ‘top-down’ hydrological modelling approach forced by (Regional Climate Model)  
75 RCM simulations for the mid-21<sup>st</sup> century period. The study approach makes three novel  
76 contributions to the existing body of knowledge.

- 77 • Few gauging stations in the developing world have long-term and accessible observed  
78 discharge data which can be used for determining impact of a dam constructed 50 years  
79 ago. This study is valuable because it extends our understanding of observed changes in  
80 river flow regime in a developing country context.
- 81 • A methodological innovation in the modelling approach is that we examine the potential  
82 impact of climate change alone by isolating the climate change signal. We also compare  
83 potential future climate change impacts with combined impact of dam and climate change.
- 84 • For a mesoscale river basin, GCM outputs are not useful because they do not provide the  
85 necessary spatial variability, which RCM simulations provide. The four RCM simulations  
86 used here represent the most comprehensive set of high resolution future climate  
87 simulations available for this region, which make them useful for assessing potential  
88 scenarios of future climate change impact on the river flow regime.

89

### 90 **1.1 Description of the study area**

91 The Kangsabati River (basin area: 5,796 km<sup>2</sup>) originates in the Chotanagpur plateau of central  
92 India, flows in a southeasterly direction to merge with the Ganges River in India, as its last  
93 contributing river (Figure 1). Upper reaches have hardpan sub-surface geology while the middle  
94 reaches consist of transitional undulating terrain, which levels out into the alluvial plains of the  
95 lower reaches. The geology of this lateritic region and the excessively drained topography cause  
96 high monsoon runoff coupled with low flow conditions during the dry months. Therefore,  
97 despite a high average annual rainfall (western part, 1300 mm and eastern part, 1600 mm), the

98 basin has been traditionally considered drought prone due to low water holding capacity of the  
99 lateritic soil, high summer temperatures and high evapotranspiration rates (Mishra and Desai,  
100 2005; Saxena, 2012).

101 The Kangsabati reservoir is located at the confluence of the Kangsabati River and a major  
102 tributary, Kumari. A dam constructed in 1965 on the Kangsabati River was followed by a second  
103 connected dam over Kumari River in 1973. In the intermediate period, partial regulation of the  
104 total flow took place. Since 1974 inflow to the Kangsabati reservoir comprises of the combined  
105 streamflow of Kangsabati and Kumari sub-basins. The diverted water is primarily used for  
106 irrigation in the reservoir command, the area of which is approximately 5,568 km<sup>2</sup>. The dam also  
107 provides flood water storage to mitigate the flooding problems in the lower reaches. High water  
108 demand in the command area has also led to over-exploitation of groundwater resources and  
109 consequently affected the river flow.

110

111 Figure 1

112

113 This river sustains the natural ecosystems which provides locals with their staple food; fish. It  
114 also has the most diverse macrophytic riverine vegetation in the region with up to 80 species  
115 found across the pre-monsoon, monsoon and post-monsoon seasons (Pradhan et al. 2005). Most  
116 siluroid fishes in the region are commercially important and the lower reaches of the Kangsabati  
117 River possess the greatest variety of fishes in the region. However, these fishes are highly  
118 vulnerable to environmental degradation, particularly habitat destruction (Giri et al. 2008). The  
119 studies performed in this region, being sporadic and short term, do not allow for a coherent long-  
120 term ecosystem analysis of river discharge and ecological health.

121 Figure 2 presents the observed discharge at Mohanpur gauging station for the period 1950-2010,  
122 where the 1950-1965 represents the natural flow regime, 1965-1973 represents partial effect of  
123 dam, while dam altered flow regime prevails from 1974-2010. Barring the 1978 floods, the dam  
124 has effectively kept peak flood levels below the 4000 m<sup>3</sup>/s mark. The dampening effect of the  
125 dam is also clearly visible with larger bases of the flood peaks after 1985. Beyond existing  
126 anthropogenic interventions, impending climate change is expected to alter the hydrological  
127 characteristics of the region by reducing the frequency of extreme precipitation events and  
128 lengthening dry spells (Mittal et al. 2013).

129

130 Figure 2

131

## 132 **1.2 Study design**

133 Assessment of ecologically important natural flow regime characteristics necessitates long-term  
134 data, especially for the period prior to the onset of an impact event or change. The gauging  
135 station at Mohanpur, about 80 km downstream of the reservoir has pre-dam discharge data for  
136 the period 1950-1965, which may be considered enough for a bias-free and appropriate  
137 assessment (Kennard et al. 2010). After the intermediate period of 9 years (1966-1973), where  
138 the influence of damming is partial and therefore difficult to understand in terms of impact, a  
139 total of 37 years of post-dam discharge information is available (1974-2010). This constitutes the  
140 observed data and forms the basis for the pre- and post-dam analysis at Mohanpur. Variability in  
141 regulated rivers is highly influenced by water use, while climatic forcing at different time scales  
142 also brings about hydrological changes. Therefore, it is crucial to separate flow regime changes  
143 caused by climate change from dam effects, so that a better knowledge of ecosystem impacts and  
144 potential restoration may be developed (Zolezzi et al. 2009). Based on this understanding,  
145 analysis of impact of dam and climate change on streamflow has been carried out in three parts;  
146 (i) effect of dam (ii) impact of future climate change (climate change signal) and (iii) impact of  
147 both dam and climate change in the future. Hydrologic alteration of biologically relevant flow  
148 regimes expected to be caused by dam construction and climate change are assessed using  
149 Indicators of Hydrologic Alteration (IHA) (The Nature Conservancy 2009).

150

## 151 **2. Methods and Data**

### 152 **2.1 SWAT hydrologic model**

153 SWAT 2009 (Neitsch et al. 2009) is used to simulate river discharges for observed and future  
154 period. SWAT typically operates on a daily time step and accounts for spatial heterogeneities of  
155 soil, land cover and elevation, by subdividing basin into multiple hydrological response units  
156 (HRUs). The rainfall-runoff model simulates the discharge from each sub basin and routes the  
157 streamflow to the watershed outlet (Neitsch et al. 2009). Preprocessing and model setup were  
158 performed using the Arc-SWAT extension for ArcGIS 9.3. The Sequential Uncertainty Fitting  
159 algorithm (SUFI-2) (Abbaspour et al. 2007) is used to calibrate SWAT and quantifies uncertainty

160 using  $P$  factor and  $R$  factor statistics. The  $P$  factor, which varies from 0 to 1, represents the  
161 fraction of observed discharge which falls within the 95PPU band, while the  $R$  factor is derived  
162 by taking the ratio of the average width of the 95PPU and the standard deviation of the observed  
163 discharge. While a value of less than 1 is considered desirable for  $R$  factor, the ideal value for  $P$   
164 factor is 1 (100% values within the band) (Vaghefi et al. 2013). 95PPU is 95 Percent Prediction  
165 Uncertainty, calculated at the 2.5% and 97.5% levels of an output variable, disallowing 5% of  
166 the bad simulations. Three evaluation criteria are used to assess model performance: Percent bias  
167 (PBIAS), Nash-Sutcliffe efficiency (NSE) and coefficient of determination ( $R^2$ ). PBIAS, NSE  
168 and  $R^2$  describe the goodness-of-fit between simulated and observed flow; and the model  
169 simulation would be considered satisfactory when PBIAS values are  $< 25\%$  and best when their  
170 values approach one in case of NSE and  $R^2$  (Moriasi et al. 2007).

171

## 172 **2.2 Assessment of hydrologic alteration**

173 IHA methodology based on Range of Variability Approach (RVA) is applied to assess the degree  
174 of departure from natural flow regime that has already occurred due to dam construction and is  
175 expected in the future due to climate change (Richter et al. 1997). RVA is the most widely used  
176 approach for quantifying hydrologic alterations in order to set appropriate environmental flow  
177 targets (Zolezzi et al. 2009). To analyse the degree of hydrologic alteration in ecologically  
178 relevant statistics, a subset of indices is used, as there exists redundancy among the indices  
179 representing different flow components (Olden and Poff 2003).

180 For RVA analysis, the pre-impact streamflow data is divided into three different categories;  
181 values upto 33<sup>rd</sup> percentile (lower category), 34<sup>th</sup> to 67<sup>th</sup> percentile (middle category) and values  
182 greater than 67<sup>th</sup> percentile (high category). A Hydrologic Alteration factor is calculated for each  
183 of the three categories as: (observed frequency – expected frequency) / expected frequency. A  
184 positive Hydrologic Alteration (HA) value indicates an increase in frequency of values in the  
185 category while negative indicates a reduction. In the absence of specific ecological information,  
186 the range between the 34<sup>th</sup> and the 67<sup>th</sup> percentile, i.e. the middle category is identified as the  
187 targeted range of variability for the post-impact period.

188

## 189 **2.3 Observed input data**



190 SWAT model required input for topography, soil and land use/land cover which are compiled  
191 from Global Land Cover Facility (GLCF) website, National Bureau of Soil Survey and Land Use  
192 Planning (NBSS&LUP), unsupervised classification of digital remote sensing images of LandSat  
193 5 Thematic Mapper (TM) for year 1990 (dated 07/11/1990 and 21/11/1990) and Landsat 7  
194 Enhanced Thematic Mapper (ETM+) for year 2001 (dated 26/10/2001 and 02/11/2001)  
195 respectively. Observed climate data including precipitation, maximum air temperature and  
196 minimum air temperature from 1991 to 2010 for five weather stations (Figure 1) are gathered  
197 from India Meteorological Department (IMD) and Agro-Meteorology Department, Government  
198 of West Bengal. Observed discharge data from river gauging stations, Simulia, Tusuma,  
199 Rangagora, Kharidwar and Mohanpur are collected from the Central Water Commission (CWC)  
200 and Irrigation and Water Ways Department (IWWD), Government of West Bengal. The  
201 Kangsabati reservoir is included with reservoir operational information starting from 1974, when  
202 the second phase of Kangsabati dam completed. Reservoir management information includes  
203 measured monthly outflow to calculate reservoir outflow, reservoir surface area when reservoir  
204 is filled to emergency (12498 ha) and principal spillway (11101 ha), volume of water needed to  
205 fill the reservoir to the emergency ( $123500 \text{ m}^3$ ) and principal spillway ( $98186 \text{ m}^3$ ).

206

#### 207 **2.4. Future climate data**

208 Daily precipitation, maximum and minimum temperature from four RCM simulations and their  
209 ensemble mean are used to drive calibrated SWAT. The historical (control) simulations for the  
210 period 1970-1999 and A1B SRES emission scenario based future climate simulations for the  
211 period 2021-2050 from four RCM simulations, REMO-ECHAM5, REMO-HadCM3, HadRM3-  
212 ECHAM5 and HadRM3-HadCM3; are obtained by the forcing from two CMIP3 GCMs namely  
213 ECHAM5-MPIOM and HadCM3 and two RCMs; REMO and HadRM3. The performance of  
214 these RCMs for the Kangsabati basin has been validated by comparing 20 year model  
215 simulations for the period 1989–2008, driven by lateral boundary forcings from ERAInterim  
216 reanalysis data (Simmons et al. 2007), with the observational datasets; Climate Research Unit  
217 (CRU) for temperature and Asian Precipitation Highly Resolved Observational Data  
218 (APHRODITE) for precipitation. Both the RCMs have demonstrated an adequate ability to  
219 capture the seasonal characteristics and interannual variability (IAV) of temperature and  
220 precipitation (Mittal et al. 2013). The ensemble mean of four RCM simulations are used to

221 simulate future streamflow, due to which the use of bias correction is considered unnecessary  
222 (Maurer and Pierce 2014). The use of ensemble reduces the uncertainties in climate projection  
223 and provides more quantitative information for subsequent hydrologic impacts research (Jung et  
224 al. 2012).

225

### 226 **3. Results and Discussion**

#### 227 **3.1 SWAT model parameter sensitivity analysis**

228 SWAT model was calibrated for the Kangsabati river basin using monthly observed streamflow  
229 at the five gauging station, during for the period 1991 to 2000. Due to the unavailability of  
230 observed weather data for the pre-dam period from 1950 to 1965, SWAT calibration was carried  
231 out using the post-dam period data. Initially, wide but meaningful ranges are assigned to  
232 sensitive parameters and with further simulations final ranges of model parameters were  
233 determined. The parameters with highest sensitivity are used to calibrate and validate the model.  
234 Table 1 shows the sensitive parameters included in the final calibration, their initial ranges,  
235 initial and final values and their t and p values. Eleven parameters representing the surface runoff,  
236 groundwater and soil properties are found to be sensitive in the estimation of streamflow. t-  
237 statistics provides a measure of sensitivity (larger in absolute values are more sensitive) and p-  
238 values determined the significance of the sensitivity with a values close to zero having more  
239 significance. Having high t-statistics and low p-value; Curve Number (CN2), alpha baseflow  
240 (ALPHA BF) and groundwater delay (GW DELAY) parameters are found to be the most  
241 sensitive to streamflow.

242

243 Table 1

244

#### 245 **3.2 SWAT model calibration, validation and uncertainty analysis**

246 The statistical comparison between observed and SWAT simulated streamflow at different  
247 gauging stations during calibration period from 1991 to 2000 shows PBIAS values ranging from  
248 -12.4 to 7.9%, higher values of  $R^2$  (ranging from 0.66–0.87) and NSE (ranging from 0.63–0.74)  
249 for all the gauging stations (Table 2). This suggests that model simulation can be judged as  
250 satisfactory as PBIAS values range between the  $\pm 25\%$  limits,  $R^2$  is greater than 0.6 and NSE is  
251 greater than 0.5 (Moriassi et al. 2007), although NSE values for Simulia and Kharidwar are below

252 0.65, considered to be an acceptable value (Ritter et al. 2013). The  $P$  factor indicates that for all  
253 stations, more than 72% of the data are bracketed in the prediction uncertainty of the model,  
254 whereas the  $R$  factors are mostly around 1 except Mohanpur gauging station where the  $P$  factor is  
255 40% and  $R$  factor is 0.59.

256 For validation for the period 2001 to 2010, PBIAS,  $R^2$  and NSE validation values ranges from -  
257 4.8 to 11.8%, 0.66 to 0.85 and 0.53 to 0.76 respectively, indicating a good relationship between  
258 observed and simulated streamflow values except for Simulia, Kharidwar and Mohanpur station  
259 with low NSE values of 0.53, 0.64 and 0.49 which are unsatisfactory according to Ritter et al.  
260 (2013). The  $P$  factor indicates that for all stations, more than 62% of the data are bracketed in the  
261 prediction uncertainty of the model, whereas the  $R$  factors are mostly around or below 1 except  
262 Tusuma gauging station where the  $R$  factor is 0.65. In general, in the downstream of Kangsabati  
263 dam, the model prediction has larger uncertainties. Poor calibration and validation results in case  
264 of managed streamflow have also been observed before (Faramarzi et al. 2010; Vaghefi et al.  
265 2013).

266

267 Table 2

268

### 269 **3.3 SWAT model simulation**

270 The calibrated model is used to simulate streamflow for two time periods, 1970-1999 (control)  
271 and 2021-2050 (future), based on ensemble mean of four RCM simulations for the SRES A1B  
272 scenario. To analyse the impact of climate change and the combined effect of dam and climate  
273 change, two separate simulations are carried out.

274 *Simulation 1* (impact of climate change) - SWAT model streamflow simulations for these control  
275 period simulations without the inclusion of the Kangsabati dam represent the natural flow regime  
276 of the basin. Comparison of this flow regime with SWAT simulated flow regime for the future  
277 period (2021-2050) is used to isolate the impact of climate change.

278 *Simulation 2* (impact of dam and climate change) - SWAT model is run for the future period  
279 (2021-2050) based on the RCM simulations and their ensemble. Kangsabati dam is included in  
280 this simulation to analyse the streamflow conditions due to both, dam and climate change.  
281 Comparison of future period simulations with observed streamflow for pre-dam period (natural

282 flow regime - 1950-1965) is used to assess the combined impact of dam and climate change on  
283 the natural flow regime.

284

### 285 **3.4 Impact of dam on flow regime**

286 The primary function of the Kangsabati dam is to divert water for irrigation and to mitigate the  
287 impacts of monsoon floods. The IHA based analysis is described from the perspective of pre-  
288 monsoon, monsoon and post-monsoon periods. Seasonal variations in flow of the river after dam  
289 construction are much relevant to the physiological and life cycle stages of various freshwater  
290 fishes. In this case, *Bagarius bagarius*, the largest freshwater migratory siluroid fish (catfish),  
291 which is abundantly present in rivers flowing through West Bengal, is found to be absent in the  
292 downstream section of the Kangsabati dam ((Hamilton, 1822, Mishra and Coulibaly 2009). It is  
293 categorized threatened by the International Union for Conservation of Nature (IUCN 2013),  
294 primarily due to its decline as a result of dam construction which prevents their upstream  
295 migration for spawning (Lakra et al. 2011). The effect of Kangsabati dam on the observed flow  
296 regime is depicted in Figure 3, through monthly average flows, monthly low flows and Flow  
297 Duration Curves (FDCs) for the representative months of April (pre-monsoon), July (monsoon)  
298 and November (post-monsoon). During pre-monsoon, the middle value for monthly average and  
299 monthly low flows is higher in the post dam period, largely due to periodic dam releases during  
300 the otherwise dry period characterized by natural minimum flows. Post-impact period is also  
301 characterized by greater flow variability, with more frequent high flow events. The  
302 corresponding FDC clearly corroborates this assessment, by depicting persistent higher flow  
303 rates for more than 80% of the time period as well as significantly lower flow rates for the  
304 remaining 20% of the time. Whereas in the monsoon season, the dam dampens the monthly  
305 average and low flows by absorbing high flow pulses and maintaining a more consistent flow  
306 rate. The FDC clearly demonstrates the overall effect, where the difference between the area  
307 under the curves for the pre-dam and post-dam periods corresponds to the amount of water  
308 diverted for irrigation purpose. Irrigation requirements for the Rabi (winter season) crop further  
309 reduce the discharge downstream of the dam during the post-monsoon month of November. A  
310 fraction of the high flows is diverted for this purpose, thereby reducing the monthly average  
311 flows.

312

313 Figure 3

314

### 315 **3.5 Effect of climate change on flow regime**

316 Figure 4 (a) show the effect of climate change on simulated flow regime through FDC, EFCs and  
317 hydrologic alteration graphs based on the output of SWAT “Simulation 1- impact of climate  
318 change” for the time period 2021-2050. The comparative analysis of FDCs in Figure 4 reveals  
319 the effect of the climate change vis-à-vis the combined effect of dam and climate change. In this  
320 case, as the FDC demonstrates, climate change reduces flows, but the area under the curve is  
321 affected to a lesser degree than for the impact of dam alone (Section 3.4).

322 Figure 4 (b) demonstrates the deviation factor of coefficients of dispersion (CD), which  
323 represents the change in flow variability as represented by EFCs during the mid-21st century  
324 compared to control period. Climate change causes deviation in both extreme low flow and high  
325 flow components. Deviation for high flow peak and frequency is higher ( $>0.6$ ), but the deviation  
326 for extreme low flow peak, duration, timing and frequency is lower ( $<0.6$ ). A slight change in the  
327 timing of high flow pulses affects the benthic siluroid fishes which are very good indicators of  
328 habitat degradation (Wootton et al. 1996). The change in flow timings affects their life cycle by  
329 disrupting various stages such as spawning, egg hatching, rearing, movement onto the  
330 floodplains for feeding and reproduction or migration upstream and downstream (Poff et al.  
331 1997).

332 Figure 4 (c) shows the extent of Hydrologic Alteration (HA) in monthly flows during the mid-  
333 21st century. A high HA ( $>0.5$ ) for the high and low category is projected for the monsoon  
334 months (JJAS) while rest of the months show less hydrologic alteration in both the categories.  
335 Months of May and September shows high alteration in both the middle and high category,  
336 whereas less alteration is observed in all three categories during post-monsoon (ON) and winter  
337 months (DJF). This reduction may significantly affect the connectivity with the flood plains by  
338 potentially reducing the magnitude and areal spread of floods. Such drastic changes will affect  
339 the yolk-sac-larva of threatened species of siluroid fish *Mystus gulio*, which develops in  
340 floodplain freshwater (IUCN, 2013; Termvidchakorn and Hortle, 2013). Along with decrease in  
341 the number of high flow events, reduction of high flow duration and changes in their timing will  
342 add additional physiological stress to the fish species (Sharma and Shrestha, 2001). As of now,  
343 Mishra and Coulibaly (2009) reported a decline of 27.8% in *Mystus gulio* catch across

344 southwestern Bengal. Observed reduction may be due to a combination of stressors such as  
345 overfishing, flow alteration and habitat loss, but in the absence of biological information and  
346 temporal monitoring of stressors, the influence of individual factors cannot be determined  
347 (Sarkar and Bain, 2007).

348

349 Figure 4

350

### 351 **3.6 Effect of dam and climate change on flow regime**

352 The output of SWAT “simulation 2 – dam and climate change” is used to analyse the combined  
353 impact of dam and climate change on hydrologic indicators in the Kangsabati basin during the  
354 mid-21st century (2021-2050) compared to pre-dam period (1950-1965). The combined effect of  
355 dam and climatic changes, depicted in Figures 4 indicates significant reduction in the magnitude,  
356 frequency and duration of extreme high and medium flow rates in the simulated flow, whereas, a  
357 small increase in low flows is observed in comparison with sole effect of climate change.

358 The FDC shows how extreme high flows above 2000 m<sup>3</sup>/s are eliminated in this scenario. The  
359 natural flow regime shows a consistent temporal distribution with ~ 75% flows lying in the range  
360 from 200 m<sup>3</sup>/s to 10 m<sup>3</sup>/s (Figure 4 (a)). However, in the altered future condition, this percentage  
361 reduces significantly to ~10%. Similarly, the combined effect of dam and climate change shows  
362 significantly greater alteration in EFCs compared to only climate change.

363 Figure 4 (b) represents the deviation factor in CD of EFCs during future (2021-2050) period in  
364 comparison with natural flow (1950-1965). In case of combined impact of dam and climate  
365 change, deviation of >0.5 is observed for CD of all extreme low flow and high flow  
366 characteristics, as compared to the individual impact of climate change. There is also a moderate  
367 increase in deviation in high flow frequency and duration due to the combined effect of dam and  
368 climate change. Significant changes in timing, frequency and duration of extreme low and high  
369 flows implies that the life cycle of many aquatic species may get disrupted during various stages  
370 such as spawning, egg hatching, rearing and their movement onto the floodplains for feeding and  
371 reproduction (Suren and Riis 2010). Benthic siluroid fishes found commonly in the Kangsabati  
372 River, which have declined since the 1960’s (Mishra and Coulibaly 2009), are highly sensitive to  
373 reduction in high flows which cause habitat degradation through channel bed sedimentation  
374 (Lisle 1989).

375 Figure 4 (c) shows the extent of HA in monthly flows due to the combined effect of dam and  
376 climate change. Unlike previous two scenarios, individual impact of dam and climate change;  
377 high hydrologic alteration in either of high, middle or low alteration category is distributed  
378 throughout the year except for the month of October, where HA is comparatively less. Positive  
379 HA ( $>0.5$ ) in the high category is projected for June, August, September and November months  
380 while negative HA ( $> -0.5$ ) is observed in January, February, March, April, July and December.  
381 Low flows in the month of January, February, March and April are projected to increase due to  
382 the increase in HA in the low category in these months whereas conversely shows reduction  
383 during May, June and November months.

384 Previous global analysis has indicated that the impact of climate change on flow regime is larger  
385 than the effect of dams and water withdrawals (Doll and Zhang 2010). However, this may be on  
386 account of two factors; the likely underestimation of dam impacts (Doll and Zhang 2010) and the  
387 high degree of spatial variation and basin specific impacts. An important factor which needs  
388 consideration is that the Kangsabati reservoir storage represents about one-third of the total  
389 annual discharge and is, therefore, a major factor in altering the flow regime of this basin. Future  
390 climate change will, therefore, put additional stress leading to greater risk of ecological change in  
391 a riverine ecosystem already affected by anthropogenic interference.

392

393

#### 394 **4. Conclusions**

395 This study provides a detailed basin scale assessment of ecologically relevant flow alterations in  
396 a monsoon dominated, drought prone river basin. IHA and EFC parameters describing changes  
397 in long-term monthly average, timing, duration and frequency of extreme flow conditions show a  
398 significant change from the natural flow regime during the observed period. Dampening effect of  
399 dams on hydrological variability and the extreme seasonality of river flows is highly pronounced.  
400 Significant overall flow reduction by the dam for provision of irrigation and domestic water  
401 demands will be exacerbated by climate change. The combined effect of dam and climate change  
402 is found to be significantly greater than the individual impact of dam or climate change.  
403 However, lack of sufficient long-term ecological data is a limitation in the assessment of habitat  
404 changes in the Kangsabati basin. We find that the ecologically sensitive IHA parameters and the  
405 associated inferences that may be drawn regarding the impacts on aquatic species are useful in

406 cases where availability of long-term ecological data is a drawback. There is an urgent need to  
407 correlate real time ecological, bio-geochemical and morphological characteristics with observed  
408 hydrological changes to better assess the vulnerability of aquatic ecosystems to future changes. A  
409 better understanding of ecosystem impacts will be useful to inform the method of river  
410 restoration and ecosystem management programmes in the future.



411

412 **List of Figures**

413 Figure 1 Map of the Kangsabati basin, with land cover land use classes, location of hydro-  
414 meteorological stations and Kangsabati reservoir.

415

416 Figure 2 Hydrograph of the Kangsabati River from 1950 to 2010.

417

418 Figure 3 Effect of dam on monthly average flows, monthly low flows and flow duration curves  
419 during pre-monsoon, monsoon and post-monsoon periods.

420

421 Figure 4 Impact of climate change and combined impact of dam and climate change on various  
422 hydrologic indicators, (a) annual flow duration curve, (b) magnitude, duration, timing and  
423 frequency of extreme low flows and high flow pulses, (c) Hydrologic Alteration

424

425

426 **List of Tables**

427 Table 1 Sensitive SWAT parameters included in the final calibration, their initial ranges, final  
428 values and their t and p values.

429

430 Table 2 SWAT Model performance of five calibrated subbasins in the Kangsabati basin.

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## Tables

Table 1 Sensitive SWAT parameters included in the final calibration, their initial ranges, final values and their t and p values

| Parameter name <sup>1</sup> | Definition  | t-statistics <sup>2</sup> | p-value <sup>2</sup> | Initial value | Range of values in SWAT-CUP             | Final value                   |
|-----------------------------|---|---------------------------|----------------------|---------------|---|-------------------------------|
| v__ALPHA_BF.gw              | Base-flow alpha factor (days)   | 4.29                      | 0.00                 | 0.048         | 0.0-0.7                                 | 0.54                          |
| v__GW_DELAY.gw              | Groundwater delay (days)  | -7.65                     | 0.00                 | 31            | 0.0-250.0                               | 80.75                         |
| v__GWQMN.gw                 | Threshold depth of water for return flow  | -1.53                     | 0.13                 | 0             | 0.0-1.2                                 | 0.81                          |
| v__GW_REVAP.gw              | Groundwater revap (water in the shallow aquifer returning to root zone) coefficient | 1.43                      | 0.15                 | 0.02          | 0.0-0.2                                 | 0.10                          |
| v__ESCO.hru                 | Soil evaporation compensation factor  | 1.06                      | 0.29                 | 0.95          | 0.75-0.95                               | 0.78                          |
| v__CH_N2.rte                | Manning's N value for the main channels   | -0.11                     | 0.91                 | 0.014         | 0.12-0.4                                | 0.35                          |
| v__CH_K2.rte                | Effective hydraulic conductivity in main channel                                    | -1.42                     | 0.16                 | 0             | 0.0-74.0                                | 33.37                         |
| Parameter name <sup>1</sup> | Definition  | t-statistics <sup>2</sup> | p-value <sup>2</sup> | Initial value | Initial range of multiplier in SWAT-CUP | Final value of the multiplier |
| r__CN2.mgt                  | SCS runoff curve number for moisture condition II                                   | 11.31                     | 0.00                 | 75-98         | -0.4-0.004                              | -0.18                         |
| r__SOL_AWC.sol              | Available water capacity of first soil layer (mm/mm)                                | 0.35                      | 0.72                 | 0.06          | 0.0-0.4                                 | 0.39                          |
| r__SOL_K (1).sol            | Saturated hydraulic conductivity of first soil layer (mm/h)                         | 1.42                      | 0.16                 | 500           | 0.0-1.6                                 | 0.22                          |
| r__SOL_BD (1).sol           | Moist bulk density of first soil layer (mg/m <sup>3</sup> )                         | 0.98                      | 0.33                 | 1             | 0.0-0.7                                 | 0.61                          |

<sup>1</sup>The qualifier (v\_) refers to the substitution of a parameter by a value from the given range, while (r\_\_) refers to a relative change in the parameter where the current values is multiplied by 1 plus a factor in the given range.

<sup>2</sup>The t-statistics and p-values are results from 500 runs of SUFI2 simulations; the larger t-statistics and smaller p-value, shows more sensitive parameter.

Table 2 SWAT Model performance of five calibrated subbasins in the Kangsabati basin.

| River discharge station | Calibration (1991-2000) |                 |       |      |                | Validation (2001-2010) |                 |       |      |                |
|-------------------------|-------------------------|-----------------|-------|------|----------------|------------------------|-----------------|-------|------|----------------|
|                         | <i>P</i> factor         | <i>R</i> factor | PBIAS | NSE  | R <sup>2</sup> | <i>P</i> factor        | <i>R</i> factor | PBIAS | NSE  | R <sup>2</sup> |
| Simulia                 | 0.74                    | 1.08            | 7.9   | 0.63 | 0.69           | 0.71                   | 0.94            | 11.8  | 0.53 | 0.69           |
| Tusuma                  | 0.78                    | 0.75            | -6.1  | 0.72 | 0.86           | 0.76                   | 0.65            | 5.3   | 0.76 | 0.79           |
| Rangagora               | 0.68                    | 1.10            | -12.4 | 0.74 | 0.66           | 0.62                   | 1.02            | -4.8  | 0.67 | 0.66           |
| Kharidwar               | 0.72                    | 0.84            | 5.8   | 0.64 | 0.75           | 0.79                   | 0.90            | 8.3   | 0.64 | 0.75           |
| Mohanpur                | 0.40                    | 0.59            | 5.4   | 0.68 | 0.87           | 0.63                   | 0.72            | 6.2   | 0.49 | 0.85           |

## Figures

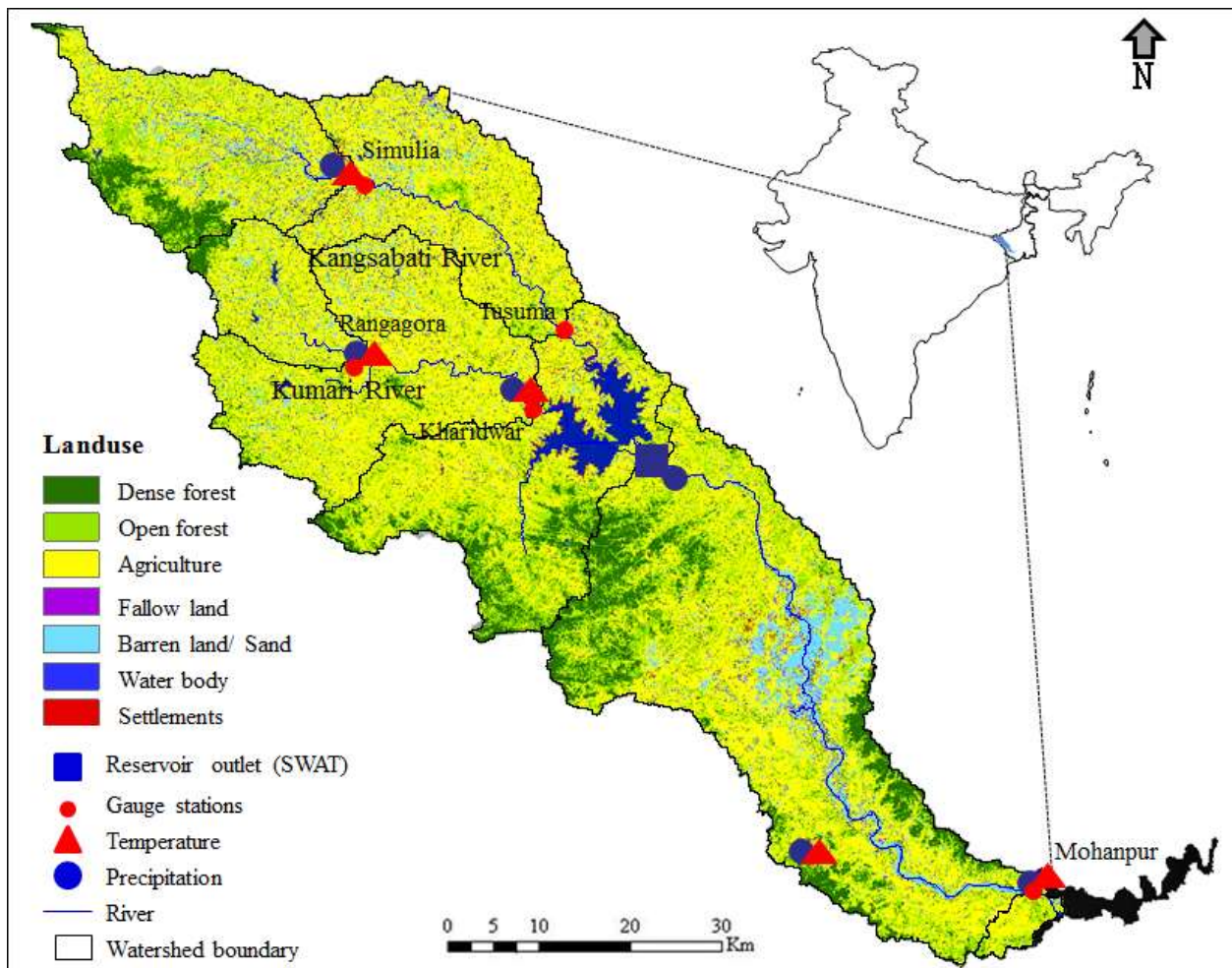


Figure 1

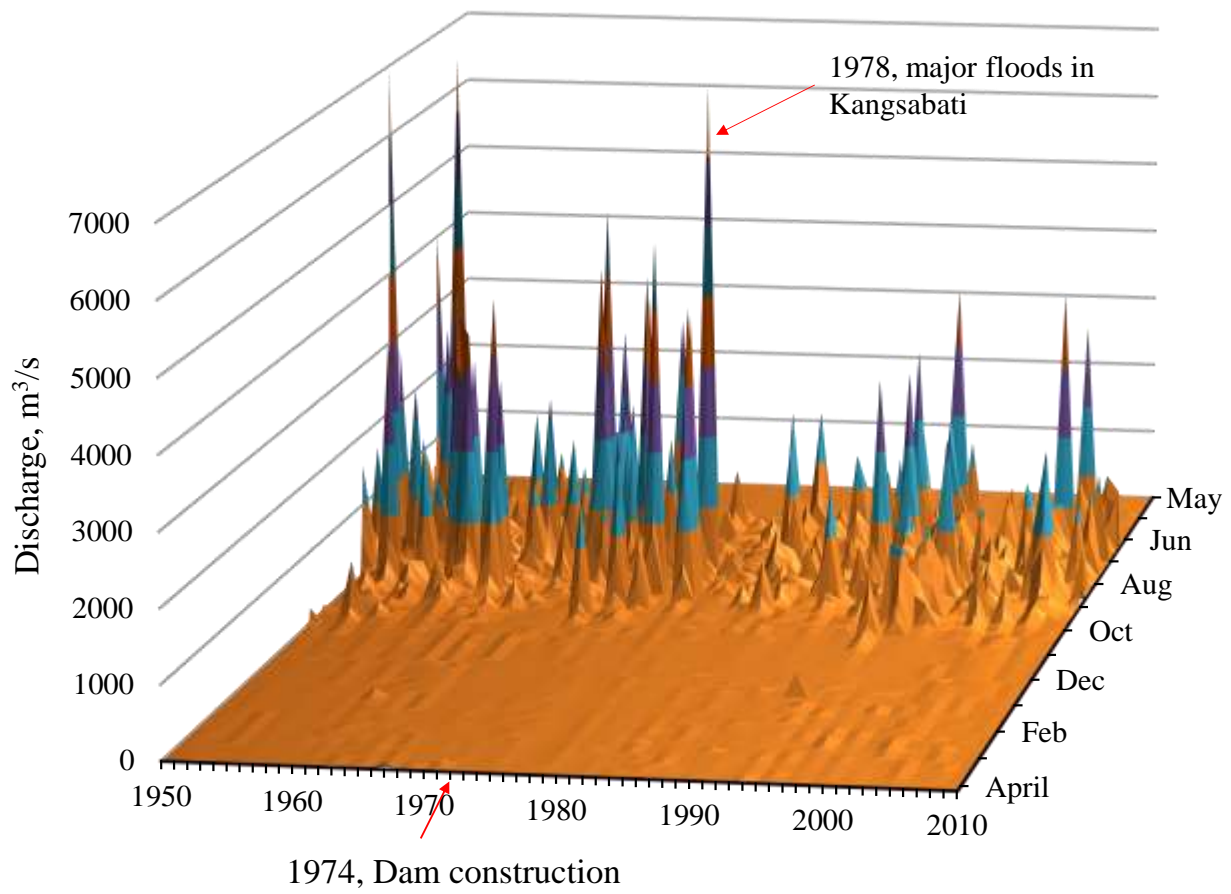


Figure 2



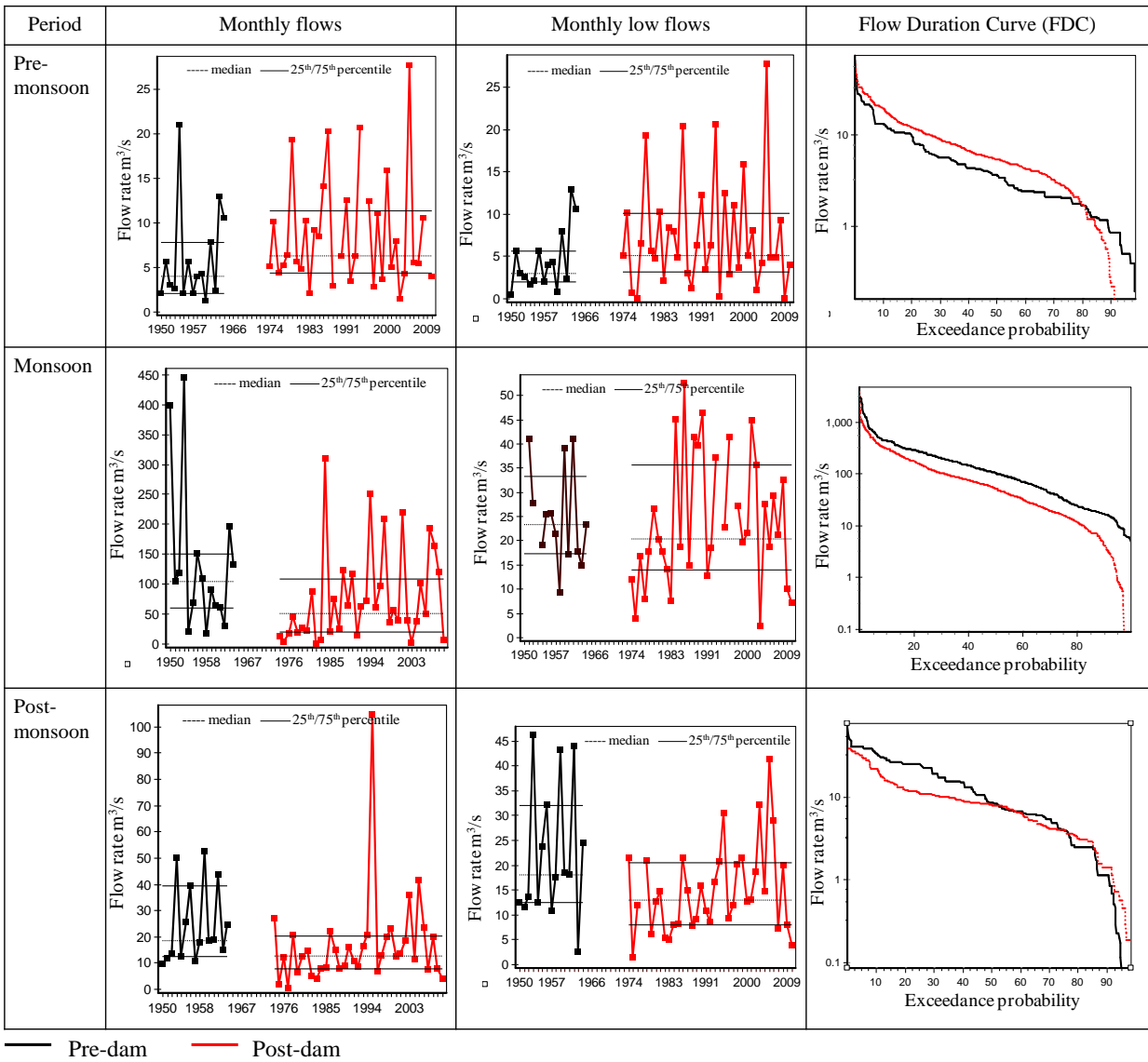


Figure 3

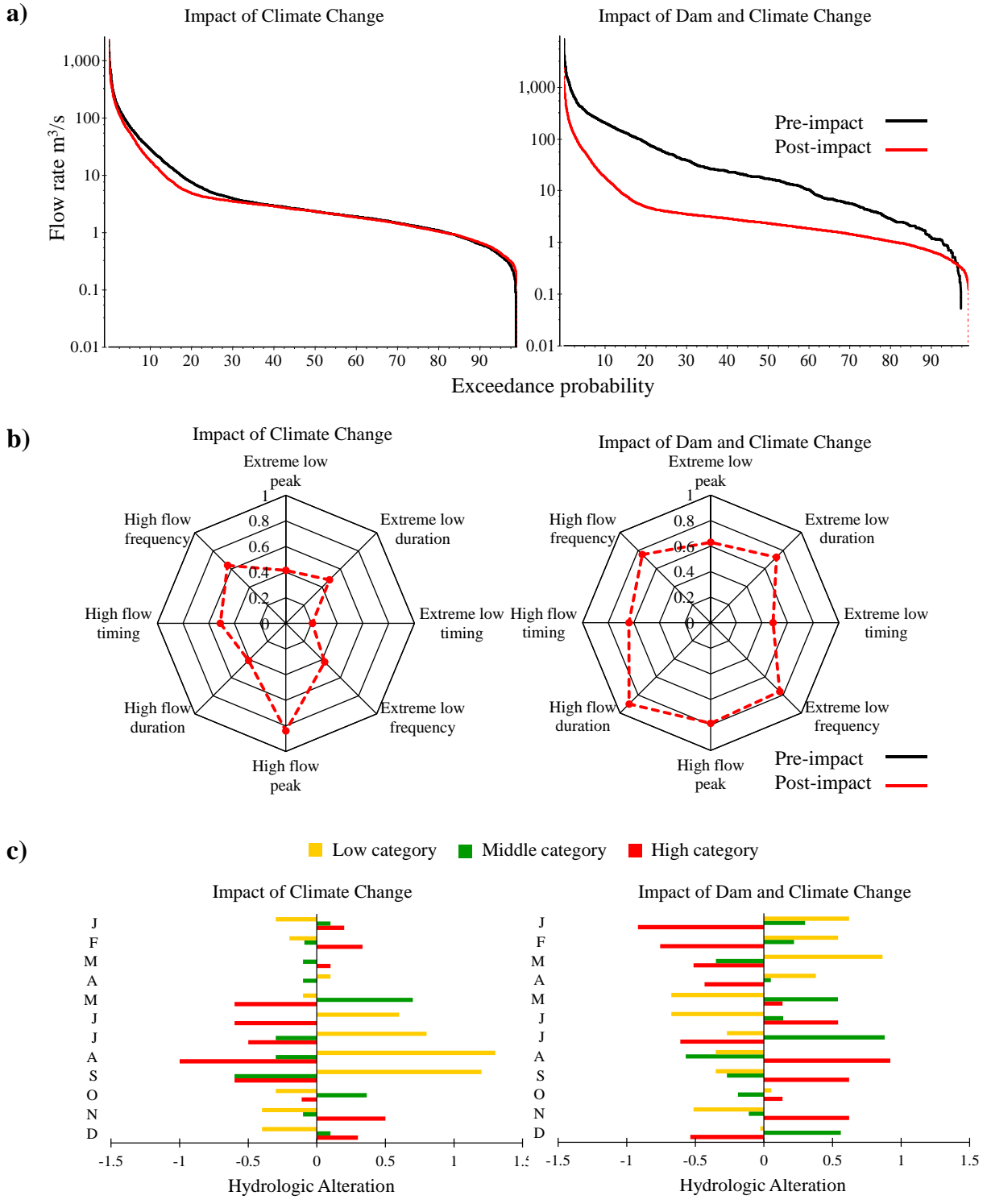


Figure 4