Lithofacies and microbiofacies (foraminifers and radiolarians) of the Permian Sequence in the Shalamzar area, Central Alborz, North Iran

Seyed Hamid Vaziri¹, Akira Yao² and Kiyoko Kuwahara²

Abstract

The Permian sequence in the Shalamzar area occurs along the Gord Calleh Mountain in Central Alborz (NW Tehran), North Iran and attains a thickness of up to 622 meters. This sequence is composed mainly of clastic rocks in the lower part and fossiliferous carbonate rocks in the upper part. The clastic facies were deposited in continental (meandering rivers) and transitional (deltaic and littoral) environments, and the carbonate facies were deposited in a shallow marine environment of the continental margin in the Paleo-Tethys.

This sequence consists of three formations: the Dorud (Asselian-Sakmarian), the Ruteh (Artinskian-Murgabian) and the Nesen (Early Djulfian) formations, and overlies non-conformably on the volcanic rocks (Devonian?) of the basement. This sequence is disconformably overlain by the Elikah Formation (Scythian-Ladinian). The rich foraminifer fauna indicates an Asselian to Early Djulfian age of the succession. Five biozones by foraminifers are established in the Permian System of this region. These biozones include the *Pseudofusulina-Schwagerina* Assemblage Zone (Asselian-Sakmarian), *Schubertella-Mesosubertella* Assemblage Zone (Artinskian), *Dunbarula-Neoschwagerina* and *Neoendothyra-Globivalvulina* Assemblage zones (Murgabian), and *Codonofusiella-Reichelina* Assemblage Zone (Early Djulfian).

Ten species of the Permian radiolarians include the *Latentifistula?* sp. A, *Latentifistula?* sp. B, Latentifistularia A, Latentifistularia B, *Orbiculiforma?* sp. A, *Orbiculiforma?* sp. B, Spherical Radiolaria A, Spherical Radiolaria B, Spherical Radiolaria C and Ellipsoidal Radiolaria A were detected from the Ruteh Formation for the first time in the present study. Index foraminifer's age of the Ruteh Formation assigns the age of detected radiolarians to Artinskian and Murgabian. The detected radiolarians are correlative with those of China, Thailand, Japan, Malaysia, Turkey, Far East of Russia, Kazakhstan, North America, Philippines and Italy. According to time unit, all recorded radiolarians can be correlated with the Permian radiolarian zones in South China, Southwest Japan and Oregon, USA and also are correlated with the *Schubertella-Mesoschubertella* Assemblage Zone (Artinskian), *Dunbarula-Neoschwagerina* and *Neoendothyra-Globivalvulina* Assemblage zones (Murgabian).

Key-words: Permian, Foraminifera, Radiolaria, Shalamzar, Central Alborz, North Iran

¹Department of Geology, Faculty of Science, Islamic Azad University, North Tehran Branch, P. O. Box 19585-851, Tehran, Iran. E-mail: s.h.vaziri@kavosh.net

²Department of Geosciences, Graduate School of Science, Osaka City University, Sugimoto 3-3-138, Sumiyoshi-ku, Osaka 558-8585, Japan. E-mail: yao@sci.osaka-cu.ac.jp, kuwahara@sci.osaka-cu.ac.jp

1. Introduction

Lower to Upper Permian rocks are widely distributed throughout North Iran (Alborz Range). They consist mainly of clastic rocks in the lower part and fossiliferous carbonate rocks in the upper part. The clastic rocks were deposited in continental (meandering rivers) and transitional (deltaic and littoral) environments (Mokhtarpour, 1997a), and carbonate rocks were deposited in a shallow marine environment of the continental margin in the Paleo-Tethys. These have been named as the Dorud and the Ruteh formations by Assereto (1963), and the Nesen Formation by Glaus (1964).

The first research on the Abyek-Hiv area has been carried out by Sieber (1970). Annells *et al.* (1977) prepared geological map of Shakran on a 1:100.000 scale in Central Alborz that Shalamzar area is situated in south part of this map. Vaziri (1993) studied biostratigraphy and lithostratigraphy of the Permian rocks in the study area and prepared a revised geological map of the Abyek-Hiv area on a 1:50.000 scale. Sedimentary petrology and depositional environments of the Lower Permian (Dorud Formation) clastic rocks in the Alborz Range (including study section) were studied by Mokhtarpour (1997a).

In the present paper, lithofacies and microbiofacies (foraminifers) have been studied by the first author (S. H. Vaziri) at the Department of Geology, Islamic Azad University, North Tehran Branch (Iran) with a reviewing at Osaka City University and radiolarian biostratigraphical study have been done by first author with contributions from the second and third authors at the Department of Geosciences, Osaka City University.

For the foraminiferal biostratigraphical study, one hundred and twenty limestone samples were collected from the Dorud, Ruteh and the Nesen formations and for study of radiolarians, twenty-seven samples were collected from chert nodules and beds of the Ruteh and the Nesen formations in the study section. No previous studies have been done on the Permian radiolarians in the Alborz Range. Therefore, the present study is the first record of Permian radiolarians from the Alborz Range, North Iran.

2. Geological setting

The clastic sediments of the Dorud Formation characterize the Lower Permian rocks of the Alborz Range. The type section of the Dorud Formation is located in Central Alborz near the village of Dorud (North Tehran), where it has a thickness of 180 meters (Assereto, 1963). It overlies unconformably the Jeirud Formation (Upper

Devonian) and consists of red sandstones, shales, conglomerate, fossiliferous limestones and quartzite. The Dorud Formation is subdivided into two groups, including clastic and carbonates facies. The clastic rocks of the Dorud Formation were deposited in continental (meandering rivers) and transitional (deltaic and littoral) environments (Mokhtarpour, 1997a).

The Dorud Formation is equivalent to the Vazhnan Formation (Baghbani, 1990b, 1996) in the Abadeh (Central Iran) and Jolfa (northwestern Iran) regions, the Howz-Dora Formation (Partoazar, 1995) in the Shotori Range, Tabas area (eastern Iran) and the Faraghan Formation (Szabo and Kheradpir, 1978; Ghavidel-Syooki, 1984, 1988) in the Zagros Range, southwestern Iran (Fig. 1).

The Ruteh Formation conformably overlies the Dorud Formation. The type section of the Ruteh Formation is located in Central Alborz near the village of Ruteh (North Tehran), where it has a thickness of 230 meters and consists of dark-grey medium-bedded to massive fossiliferous limestones (Assereto, 1963). The Ruteh Formation was deposited in barrier, lagoon-tidal flat and open marine environments (Lasemi and Mokhtarpour, 1993; Mokhtarpour, 1997b). This formation is equivalent to the Surmaq Formation (Taraz, 1969, 1974; Iranian-Japanese Research Group, 1981; Partoazar, 1995; Baghbani, 1996; Kobayashi and Ishii, 2003) in the Abadeh (Central Iran) and the Jolfa (northwestern Iran) regions, middle part of the Jamal Formation (Stocklin et al., 1965; Partoazar, 1995) in the Shotori Range, Tabas area (eastern Iran) and the lower part of the Dalan Formation (Szabo and Kheradpir, 1978; Baghbani, 1990a) in the Zagros Range, southwestern Iran (Fig. 1).

The type section of the Nesen Formation is located in northeast Nesen village (Central Alborz) and has a thickness of 144 meters (Glaus, 1964). This formation consists of black shales (abundant brachiopods and corabearing) and dark-grey cherty fossiliferous limestones. The Nesen Formation was deposited in barrier, lagoon and open marine environments (Mokhtarpour, 1997b). The lower as well as the upper boundary of the Nesen Formation is disconformable. The Elikah Formation (Lower to Middle Triassic) disconformably overlies the Ruteh Formation.

The Nesen Formation in the Alborz Range is equivalent to the Jolfa Formation (Partoazar, 1995) in the Jolfa Region (northwestern Iran), upper part of the Jamal Formation (Stocklin *et al.*, 1965; Partoazar, 1995; Aghanabati, 2004) in the Shotori Range, the Tabas area (eastern Iran), middle part of the Dalan Formation (Szabo and Kheradpir, 1978; Baghbani, 1990a) in the Zagros Range (southwestern Iran), and the Abadeh Formation

Series	International Stages	Tethyan Stages	Alborz Range (North Iran) (Assereto, 1963; Glaus, 1964 and Vaziri, 1993)	Jolfa Region (NW Iran) (Teichert et al., 1973; Partoazar, 1995 and Baghbani, 1996)	Abadeh Region (Central Iran) (Iranian-Japanese Research Group, 1981; Partoazar, 1995; Baghbani, 1990b, 1996; Kobayashi and Ishii, 2003)	Shotori Range (Tabas area, E. Iran) (Stocklin et al., 1965; Partoazar, 1995 and Aghanabati, 2004)	Zagros Range (SW Iran) (Szabo & Keradpir, 1978, Ghavidel-Syooki, 1988, Baghbani, 1990 a and Aghanabati, 2004)
	Wuchia- Chang- pingian hsingian	Djulfian Dorash- amian	Nesen Formation	Ali Bashi Formation	Hambast Formation Abadeh Formation		
Permian	Wor- Capita- dian nian	Murgabian Midian D	Ruteh Formation	Surmaq Formation	Abaden Formation	Jamal	Dalan Formation
Upper	Kungurian Roadian	Bolorian Kuberg- Mur			Surmaq Formation	Formation	
Permian	Artinskian	Artinskian Bol	Ruteh Formation				Faraghan
Lower Per	Asselian Sakmarian	Asselian Sakmarian	Dorud Formation	Vazhnan Formation	Vazhnan Formation	Howz-Dora Formation	Formation

Fig. 1 Stratigraphic correlation chart of the Permian rocks-units of Iran.

(Taraz, 1969, 1974; Partoazar, 1995, 2002) in the Abadeh Region, Central Iran (Fig. 1).

The rich foraminifer fauna from the Permian rocks along Alborz Range indicates an Early to Late Permian (Asselian to Early Djulfian) age. These rocks contain abundant fusulinids, corals, brachiopods, bellerophons and rare radiolarians.

In the Alborz Range, the Dorashamian Stage and at least the upper part of Djulfian are missing. The sedimentation gaps of Upper Djulfian and Dorashamian deposits in the Alborz Range are related to activity of Late Hercynian orogenic phase. This is to say that the *Araxoceras tectum* subzone and the *Vedioceras* subzone as well as the *Physonites*, *Shevyrevites-Iranites* and *Paratirolites* zones are missing in the Alborz Range (Golshani *et al.*, 1986). But Early Djulfian Stage is well represented by the *Codonofusiella-Reichelina* Assemblage Zone. Therefore, a larger sedimentary gap exists between the Permian-Triassic strata in the Alborz Range than either

Jolfa Region (northwestern Iran) or Abadeh Region (Central Iran). This is widely traced laterally, which indicates a cessation in the continuous sedimentation (Golshani *et al.*, 1986).

3. Lithostratigraphic description of the section

Well-defined outcrops of the Lower to Upper Permian rocks occur along the Gord Calleh Mountain (Kuh-e-Gord Calleh), northeast Shalamzar village (Abyek-Hiv area) in Central Alborz, northern Iran with the following coordinates: N 36°. 04′, E 50°. 11′ (Figs. 2 & 5a). These rocks in northeast Shalamzar attain a thickness of up to 622 meters. Permian rocks overlie non-conformably the black volcanic rocks (Devonian?) and on the top are covered disconformably by the Elikah Formation (Fig. 3).

The Permian rocks of the Shalamzar area consist of three formations: the Dorud, the Ruteh and the Nesen formations with Asselian to Early Djulfian age. They can

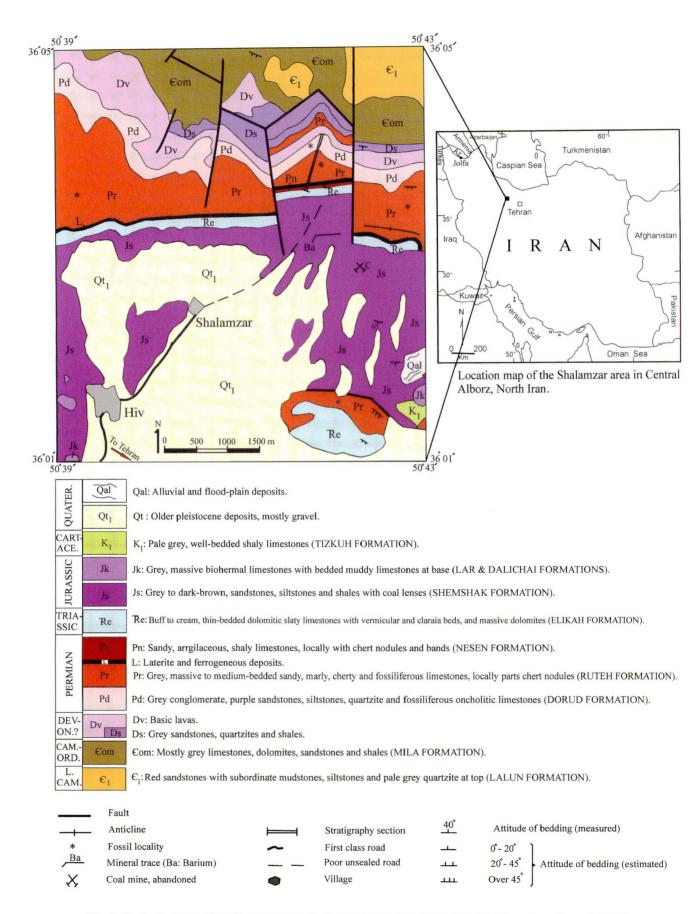


Fig. 2 Geological map of the Shalamzar area in Central Alborz, North Iran (modified from Annells et al., 1977).

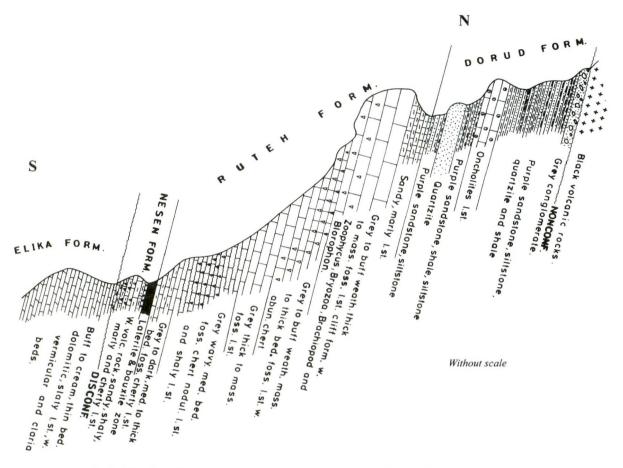


Fig. 3 Stratigraphic profile of the Permian sequence and Permo-Triassic boundary in northeast of Shalamzar, Central Alborz, North Iran

be subdivided into 12 members based on their facies characteristics (Fig. 4).

3.1 The Dorud Formation

This formation with a thickness of 269 m consists of grey congelomerate, red to purple sandstone, silty shale, quartzite and oncolitic limestones with Early Permian (Asselian-Sakmarian) age, which is covered conformably by the Ruteh Formation. This formation in the study section can be subdivided into four members as follows (Fig. 4):

Member 1 (6 m):

Grey conglomerate (composed of limestone, sandstone, quartz and volcanic fragments). Volcanic fragments belong to the volcanic rocks of the basement (Fig. 5b).

Member 2 (198 m):

Alternating purple to reddish, thin-bedded sandstones, siltstones and shales with intercalations of light grey, thin-

bedded quartzite. The bedding shows N 290° W strike, dipping at 45° SW.

Member 3 (8 m):

Grey to buff, oncolitic limestones with fusulinids, brachiopods and corals. The bedding shows N 290° W strike and dip at 45° SW. This member can be compared with the third subunit of the Dorud Formation in type section (Asserto, 1963) in northeast Tehran. This member includes the Girvanella permica Pia (Pl. 3, figs. 1, 2) as a cyanobacteria (Riding, 1991; Flugel, 2004) and the following foraminifers that indicate a Asselian to Sakmarian age (thin sections no. H.V-3 ~ 10): Earlandia sp., Eolasiodiscus cf. donbassicus Reitlinger, Eostaffella sp., Geinitzina postcarbonica Spandel, Globivalvulina cf. scaphoidea Reitlinger, Globivalvulina sp., Globivalvulina mosquensis Reitlinger, Mesoschubertella sp., Monogenerina sp., Nankinella sp., Pseudofusulina sp., Pseudofusulina aff. P. tchernyshevi (Schellwien), Pseudofusulina cf. krotovi candata Rauser, Pseudofusulina cf. diserta Saerovich, Pseudostaffella sp., Pseudoschwagerina sp.,

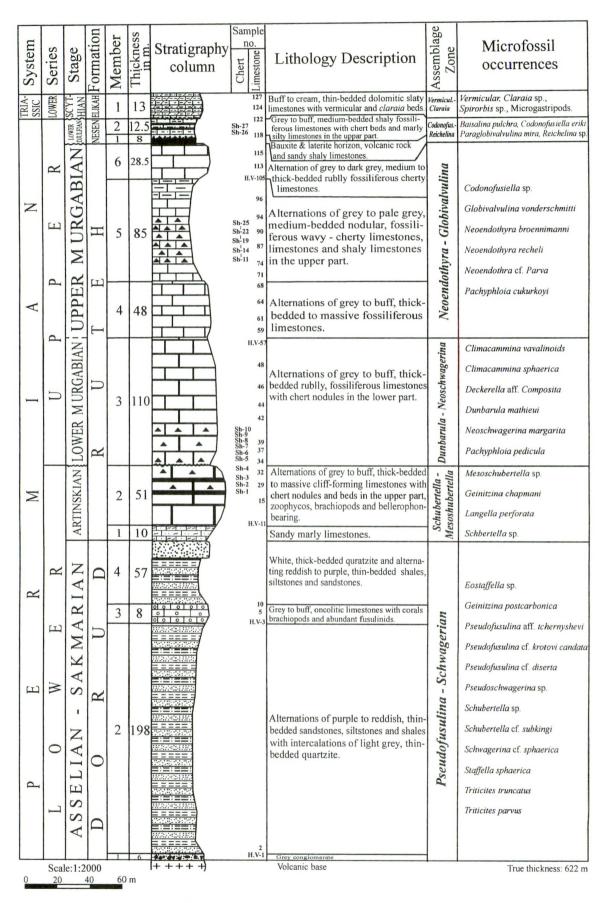


Fig. 4 Lithological characteristics and biozones of the Permian sequence and Permo-Triassic boundary in northeast of Shalamzar, Central Alborz, North Iran.

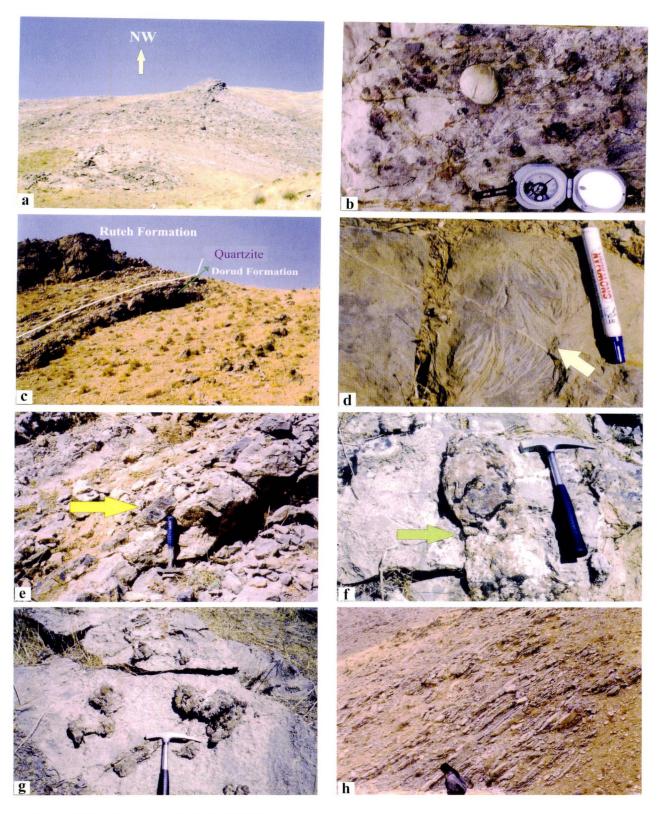


Fig. 5 (a) Permian sequence in the Gord Calleh section, Central Alborz. (b) Basal conglomerate of the Dorud Formation (member 1) (c) The contact between the Dorud and Ruteh formations, and cliff-forming limestones of the Ruteh Formation (member 2). (d) Ichnofossil of the *Zoophycos* in limestone of the Ruteh Formation (member 2). (e, f) Chert bed in alternating limestones of the Ruteh Formation (member 2). (g) Chert nodules in limestone of the Ruteh Formation (member 3). (h) Alternating wavy-cherty limestones of the Ruteh Formation (member 5).

Pseudoschwagerina cf. beedi uralensis Rauser, Schubertella sp., Schubertella cf. subkingi Dunbar et Skinner, Schwagerina sp., Schwagerina cf. sphaerica Sherb, Staffella sp., Staffella sphaerica (Abich), Tetrataxia sp., Tetrataxia cf. planulata Morozova, Triticites parvus Chen. Triticites truncatus Chen, Tuberitina sp. (Pl. 3, figs. 3-9).

Member 4 (57 m):

Consists of white, thick-bedded quartzite (14 m, Fig. 5c) and alternating reddish to purple, thin-bedded shales, siltstones and sandstones. Sandstones have sedimentary structures such as lamination and cross lamination. The bedding shows N 315° W strike and dip at 30° SW.

3.2 The Ruteh Formation

This formation with thickness of 332.5 meters conformably overlies the Dorud Formation (Fig. 5c), and consists of fossiliferous limestones. This formation in the study section can be subdivided into six members as follows (Fig. 4):

Member 1 (10 m):

Consists of grey sandy marly limestones. The bedding is the same as that of the fourth member of the Dorud Formation.

Member 2 (51 m):

Alternating grey to buff, thick-bedded to massive cliff-forming limestones with chert nodules and beds in the upper part (Figs. 5e, f). The bedding shows N 320° W strike and 40° SW dip. This member is considered Artinskian on the basis of microfossils, which indicates the Parafusulina Zone (Schubertella sp.-Mesoschubertella sp.), macrofossils such as Fenestrllina sp., Septopora, Orthotetina, Orthotechia, Bellerophontid and brachiopods, ichnofossil of the Zoophycos (Fig. 5d), and cyanobacteria of the Tubiphytes obscurus Maslov. Tubiphytes is commonly interpreted as a calcified cyanobacterium (Riding, 1991; Pratt, 1995). Senowbari-Daryan and Flugel (1993) asserted that a combination of non-preserved soft organism (central tube) and a cyanobacterial envelope is the original interpretation of Maslov (1956) regarding Tubiphytes. Riding (1993) emphasized that the correct name of Tubiphytes obscurus is Shamovella obscura. The traditional name is consequently used in the article.

This member also includes the following microfossils (thin sections no. $H.V-11 \sim 33$):

Foraminifera: Climacammina sp., Climacammina sphaerica Potieskaya, Deckerella sp., Endothyra sp., Geinitzina sp., Geinitzina postcarbonica Spandel, Geinitzina uralica Suleimnov, Geinitzina chapmani Schubert Var. Longa Sulemanov, Globivalvulina sp., Hemigordius sp., Kahlerina sp., Kahlerina pachytheca Koch. Devide et Ramors, Langella sp., Langella cf. ocarina S.de Civr & Dess., Langella perforata Lange, Lasiodiscus sp., Mesoschubertella sp., Pachyphloia sp., Pachyphloia pedicula Lange, Palaeosiroplectammina sobhaniansis Partoazar & Vaziri, Palaeobigenerina sp., Pseudolangella fragils Decivr. & Dess., Pseudostaffella sp., Schubertella sp., Tuberitina sp., Tuberitina collosa Reitlinger (Pl. 3, figs. 11-15); Algae: Gymnocodium sp., Gymnocodium bellerphontis (Rotheletz), Permocalculus sp., Pseudovermiporella sp., Vermiporella sp., Vermiporella nipponica Endo (Pl. 3, fig. 10).

Member 3 (110 m):

Alternating grey to buff, thick-bedded rubbly, fossiliferous limestones with chert nodules in the lower part (Fig. 5g). The bedding shows N 270° W strike and 40° SW dip.

This member includes the following microfossils that indicate an Early Murgabian age (thin sections no. H.V-34~ 57): Foraminifera: Climacammina sp., Climacammina vavalinoids Lange, Climacammina sphaerica Potieskaya, Climacammina major Morozova, Codonofusiella sp., Codonofusiella erki Rauser, Cribrogenerina sumatrana (Volz), Deckerella sp., Deckerella aff. composita Reitlinger, Dunbarula mathieui Ciry, Endothyra sp., Geinitzina sp., Geinitzina postcarbonica Spandel, Geinitzina uralica Suleimnov, Geinitzina chapmani Schubert Var. Longa Sulemanov, Globivalvalina sp., Globivalvalina mosquensis Reitlinger, Globivalvulina scaphoidea Reitlinger, Glomospira sp., Hemigordius sp., Kahlerina sp., Langella sp., Langella conica De Civer. & Dess., Cryptosetida (Langella) bozorgniansis Partoazar & Vaziri, Langella cf. cukurkoyi De Civer. & Dess., Neoschwagerina margarita Deprat, Nankinella sp., Pachyphloia sp., Pachyphloia pedicula Lange, Pachyphloia cff. schwageri Lange, Palaeotextularia sp., Staffella sp., Tuberitina sp.(Pl. 4, figs. 1-5); Algae: Gymnocodium sp., Permocalculus sp., Pseudovermiporella sp., Vermiporella sp., Vermiporella nipponica Endo; Cyanobacteria: Tubiphytes sp., Tubiphytes obscurus Maslov.

The member 3 overlies the member 2 paraconformably and there are no Bolorian-Kubergandian deposits.

Member 4 (48 m):

Alternating grey to buff, thick-bedded to massive fossiliferous limestones with N 270° W strike and 40° SW dip. This member includes the following microfossils (thin sections no. H.V-58 ~ 69): **Foraminifera:** *Climacammina* sp., *Codonofusiella* sp., *Codonofusiella erki* Rauser,

Dunbarula mathieui Ciry, Earlandia sp., Frondinodosaria sp., Geinitzina sp., Geinitzina sp., Geinitzina chapmani Schubert Var. Longa Sulemanov, Globivalvalina sp., Globivalvalina vonderschmitti Reichel, Hemigordius sp., Langella sp., Cryptosetida (Langella) bozorgniansis Partoazar & Vaziri, Langella cf. acanthi (Lange), Langella cf. cukurkoyi De Civer. & Dess., Neoendothyra sp., Neoendothyra reicheli Reitlinger, Neoendothyra broennimanni Bozorgnia, Pachyphloia sp., Pseudolangella sp., Schubertella sp., Staffella sp., Tetrataxia sp., Tuberitina sp. (Pl. 4, figs. 6, 7, 9-12, 14); Algae: Gymnocodium sp., Permocalculus sp., Pseudovermiporella sp., Vermiporella sp.; Cyanobacteria: Tubiphytes obscurus Maslov.

Member 5 (85 m):

Alternating grey to pale grey, medium-bedded fossiliferous wavy-cherty limestones, limestones and shaly limestones in the upper part (Fig. 5h). The bedding shows N 275° W strike and 42° SW dip. This member includes the following microfossils (thin sections no. H.V-70 ~ 104): Foraminifera: Aghathammina sp., Climacammina sp., Climacammina sphaerica Potieskaya, Climacammina vavalinoids Lange, Cryptosetida (Langella) bozorgniansis Partoazar & Vaziri, Deckerella sp., Deckerella aff. composita Reitlinger, Dunbarula mathieui Ciry, Geinitzina sp., Geinitzina chapmani Schubert Var. Longa Sulemanov, Globivalvalina sp., Globivalvalina vonderschmitti Reichel, Hemigordius sp., Langella sp., Langella cf. cukurkoyi De Civer. & Dess., Lasiodiscus sp., Neoendothyra sp., Neoendothyra reicheli Reitlinger, Neoendothyra broennimanni Bozorgnia, Neoendothyra cf. parva (Lange), Pachyphloia sp., Pachyphloia cukurkoyi De Civer. & Dess., Pachyphloia pedicula Lange, Palaeobigenerina sp., Palaeotextularia sp., Staffella sp., Tetrataxia cf. planulata Morozova, Tuberitina sp.(Pl. 4, fig. 13); Algae: Gymnocodium sp., Permocalculus sp., Pseudovermiporella sp., Vermiporella sp., Vermiporella nipponica Endo (Pl. 4. fig. 15 & Pl. 5, fig. 1); Cyanobacteria: Tubiphytes obscurus Maslov.

Member 6 (28.5 m):

Alternating grey to dark-grey, medium to thick-bedded rublly fossiliferous cherty limestones. The bedding shows N 295° W strike and 40° SW dip.

This member includes the following microfossils (thin sections no. H.V-105 ~ 115): Foraminifera: Climacammina sp., Climacammina sphaerica Potieskaya, Climacammina vavalinoids Lange, Cryptosetida (Langella) bozorgniansis Partoazar & Vaziri, Deckerella sp., Dunbarula mathieui Ciry, Geinitzina sp., Geinitzina cf. uralica Suleimnov, Geinitzina reperta Bikova, Globivalvalina sp.,

Globivalvalina vonderschmitti Reichel, Hemigordius sp., Kahlerina pachytheca Koch, Langella sp., Neoendothyra sp., Neoendothyra broennimanni Bozorgnia, Neoendothyra reicheli Reitlinger, Neoendothyra cf. parva (Lange), Pachyphloia sp., Pachyphloia cukurkoyi De Civer. & Dess., Pachyphloia pedicula Lange, Staffella sp., Tetrataxia sp., Tuberitina sp. (Pl. 4, fig. 8); Algae: Gymnocodium sp., Permocalculus sp., Pseudovermiporella sp., Vermiporella sp.; Cyanobacteria: Tubiphytes sp., Tubiphytes obscurus Maslov.

The followings are microfossils present in the members 4, 5 and 6 of the Ruteh Formation that include the *Neoendothyra* assemblage indicating a Late Murgabian age for these members.

3.3 The Nesen Formation

The Nesen Formation in the study section consists of dark grey, medium-bedded shaly cherty limestones, bauxite-laterite horizon and volcanic rock that can be subdivided into two members with a thickness of 20.5 meters. The Nesen Formation overlies the Ruteh Formation disconformably and there are no Midian deposits. These beds show N 320° W strike and 20° SW dip.

Member 1 (8 m):

Consists of bauxite-laterite horizon with thickness of 6 meters (Fig. 6d), volcanic rock and grey, thin-bedded sandy shaly limestones. This volcanic rock is similar to the volcanic rocks present at the base of the Nesen Formation in the Elikah valley (Glaus, 1964).

Member 2 (12.5 m):

Grey to buff, medium-bedded shaly fossiliferous limestones with bedded cherts (Fig. 6e) and marly silty limestones at the top. This member includes the following microfossils (thin sections no. H.V-116 ~ 122): Foraminifera: Aghathammina sp., Aghathammina pusilla Geinitz, Baisalina pulchra Reitlinger, Climacammina sp., Codonofusiella sp., Codonofusiella eriki Rauser, Cryptosetida (Langella) bozorgniansis Partoazar & Vaziri, Dagmarita sp., Deckerella sp., Dunbarula mathieui Ciry, Earlandia sp., Geinitzina sp., Geinitzina reperta Bikova, Geinitzina postcarbonica Spendel, Globivalvalina sp., Globivalvulina vonderschmitti Reichel, Hemigordius sp., Monogenerina sp., Nankinella sp., Neoendothyra reicheli Reitlinger, Pachyphloia sp., Pachyphloia cf. iranica Bozorgnia, Palaeotextularia sp., Paraglobivalvulina sp., Paraglobivalvulina mira Reitlinger, Reichelina sp., Staffella sp. (Pl. 5, figs. 3-14); Algae: Gymnocodium bellerphontis (Rotheletz), Permocalculus sp., Vermiporella

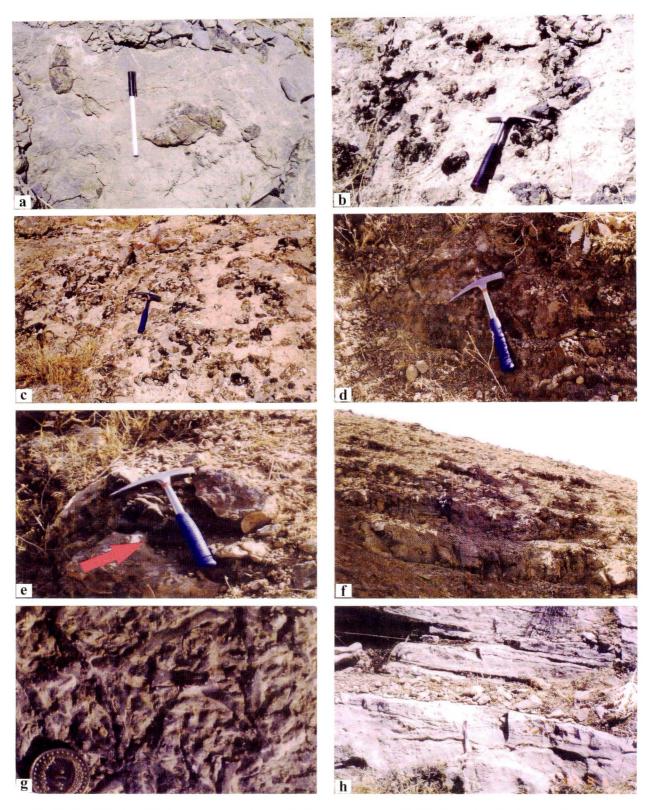


Fig. 6 (a-c) Chert nodules in limestones of the Ruteh Formation (member 5). (d) Bauxite-laterite horizon of the Nesen Formation (member 1). (e) Chert bed in alternating limestones of the Nesen Formation (member 2). (f) Alternating dolomitic slaty limestones with vermicular and *Claraia* beds of the Elikah Formation (member 1). (g) "Calcareous vermicular" in member 1 of the Elikah Formation. (h) Buff to cream dolomites of the Elikah Formation (member 2).

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ЩЩ	1	1	Ц	\perp	1	Ц	1	Ц	\perp	Ц	1	Ц	\perp	\perp	Ц	1	Ц	1	•	1	Ш	1	Ц	Ц	1	Ц	\perp	Ц	Ш	1	Щ	1	Ц	Ш	Ц	\perp	1	Ш		Ш			Ц			Ц		Ц		Ш	Ц		Ш					L	Ц				Cribrogenerina sumatrana
\coprod	1	Ш	\coprod	Ш	1	Ц	1	Ц	1	Ц	1	Ц	1	\perp	Ц	1	Ц	1	Ц	Ц	Ш	1	Ц	Ц	1	Ц	\perp	Ц	Ц	1	Ц	L	Ц	Ш	Ц	Ш	1	Ц	\perp	Ш	Ц	1	Ц	\perp	1	Ц	1	Ц	1	П	Ц		Ш									•	Dagmarita sp.
\coprod	1	Щ	Н	Ш	1	Ц	1	Ц	\perp	9	1	Ц	1		•	9	0	1	Ц	Ц	Ц	1	Ц	Ц	1	Ц		Ц	Ц	1	Ц	\perp	Ц	Ц	Ц		1	Ц		Ц	Ц		Ц	\perp			9	Ц						•	•							•	Deckerella sp.
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		9	L	9	\perp	Ц	\perp	Ц	\perp	Ц	\perp	Ц	1	L	Ц	\perp	Ц	\perp	Ц		Ш			Ц		Ц	\perp	Ц	Ц	\perp	Ц		Ц	•	Ц		\perp	Ш		Ш	Ц	\perp	Ш					Ш													•		Earlandia sp.
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\coprod	\perp	Ш	11	Ш	\perp	Ц	\perp	Ц	╀	Ц	1	Ц	4	┸	Ц	1	Ц	1	Ц	Ц	Ц	\perp	L	Ц	1	Ц	\perp	Ц	Ц	\perp	Ц	L	Ц	Ц	Ц	\perp	\perp	Ц	Ц	Ц	Ц	\perp	Ц	Ц		Ц		Ц		Ш			Ш				•		•				Geinitzina reperta
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ЩЩ	9	1	Ш	Ш	1	Ц	1	Ц	L	Ц	1	Ц	1	\perp	Ц	1	Ш	1	\sqcup			9		Ц	1	Ц		Ц	Ш	1	Ц		Ц	Ш	Ц		1	П			Ц		П			П									I		I			I			Globivalvulina scaphoidea
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\coprod	1		Ц	\perp	1	Ц	1	Ц	L	Ц	1	Ц		1	Ц		П	1						П	1	Ц		Ц	П		П		•	\prod		9		П					П	•		П	Γ	П	I		\coprod	I			I		I		I	9	Γ	П	Globivalvulina voderschchmitti
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				9 9	•					1	•	П	I		•	•	0	•)			I		-	9	9		4		I	\prod	Γ	•	•	•		•		9	•	9	•	•	4	9 0					П	•	T	П	T	T	П	T			T		П	Langella sp.
		Ш	П		I		•	4		•	I	П				I	П	I				I			T	П	I			I	\prod	I		\prod			\int	\prod				T					T		I	П		I	П		T	П	T			T	T	П	Langella perforata
\coprod						Ц		Ц		Ц		П														0									•				•			I					I	•	Ι			Ι	•		I		T				·		Langella bozorgniansis
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		Ш	П	Ш		Ц				Ц					Ц			\perp								\prod		\prod			\prod	I			\prod			П	\prod			I	\prod		I		T	\prod	I			T	П	T	T	П	T		1	T			Langella cf. ocarina
	T	IT	П	П	T	П	T	IT	T	ΙT	T	П	T		IT	T	П	T	П	T		T		I	I	\prod	I	\prod	\prod	T	П	T	IT	П	IT	П		П		П	П	T	П	T	Т	П	T	П	T	П		T	П	T	T	П	T	П		T	T	П	Langella cf. acantha
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Fig. 7 List of the foraminifera species at the Gord Calleh section in Shalamzar area, Central Alborz, North Iran

Fig. 8 List of the foraminifera species at the Gord Calleh section in Shalamzar area, Central Alborz, North Iran

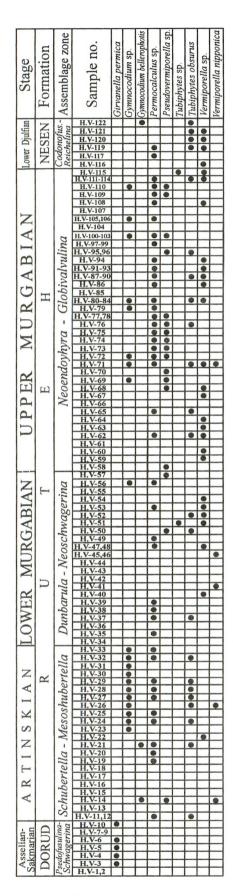


Fig. 9 List of the algae and cyanobacteria species at the Gord Calleh section in Shalamzar area, Central Alborz, North Iran.

sp.; Cyanobacteria: *Tubiphytes obscurus* Maslov. (Pl. 5, fig. 2).

The above-mentioned microfossil assemblages, especially the *Baisalina pulchra*, *Codonofusiella eriki*, *Paraglobivalvulina mira* and *Reichelina* sp. indicate an Early Djulfian age for the formation.

The Nesen Formation overlies disconformably, via a lateritic-bauxitic horizon (6 m), the Ruteh Formation and is covered disconformably by the Elikah Formation (Scythian-Ladinian). In the Alborz Range, the Permian-Triassic boundary is marked by somewhat larger sedimentary gap, resulting in the non-deposition of Upper Djulfian-Dorashamian sediments. This gap is represented by the development of a few meters of lateritic-bauxitic layers (Golshani *et al.*, 1986).

List of the foraminifera species in the study section are shown in Figs. 7, 8 and list of the algae and cyanobacteria species are shown in Fig. 9.

3.4 The Elikah Formation

The Elikah Formation in the study section consists of platy, buff to cream, thin-bedded dolomitic slaty limestones with vermicular and besides bivalves (Claraia), serpulids (Spiorbis) and thick-bedded to massive dolomites with Early to Middle Triassic (Scythian-Ladinian) age (Figs. 6f-h & Pl. 5, fig. 15). The Elikah Formation was deposited on a vast platform along the shelves of Paleo-Tethys and Neo-Tethys (Seyed-Emami, 2003). The rather poor fossil content can deduce the age of the Elikah Formation, mainly concentrated in the lower part. The main fossil taxa are foraminifers, pelecypods, ammonoids and conodonts, indicating an Early and Middle Triassic age for the formation (Allenbach, 1966; Stepanov et al., 1969; Nakazawa et al., 1981; Hirsch and Sussli, 1973; Kozur et al., 1975; Sweet, 1979; Vaziri, 2004). In few localities such as Aruh area (west of Firuzkuh) and Veresk area (northeastern Firuzkuh) along Central Alborz and Shahrud area in eastern Alborz, where the pre-Shemshak erosional phase (Early Cimmerian tectonic event) has not been deep enough, a light grey limestone up to 100 meters (Aruh Limestone Member) builds up of the top of the Elikah Formation (Zaninetti et al., 1972; Stampfli et al., 1976; Nabavi, 1987; Jahani, 2000; Vaziri and Majidifard, 2001; Seyed-Emami, 2003; Vaziri, 2004) and age of the Elikah Formation spans up to Late Triassic (Carnian). In the Shahrud area in eastern Alborz, this limestone contains serpulids (Spiorbis) and foraminifers of the Ammobaculites sp., Involutina sp., Ophthalmidium sp., Reophax sp., Sigmolina sp. and Trochammina sp. that indicate a Carnian age (Vaziri and Majidifard, 2001).

Except in the Jolfa area in northwestern Alborz, where

the Permo-Triassic boundary is rather continuous (Stepanov *et al.*, 1969; Golshani *et al.*, 1986), elsewhere along the Alborz Range, the Elikah Formation overlies with a distinct disconformity Upper Permian or even older strata (Seyed-Emami, 1971, 2003) and may have thickness up to 1000 meters.

4. Foraminiferal biozones of Permian system in the study section

Five biozones by foraminifers have been established in the study section as the following (Fig. 4):

4.1 Pseudofusulina - Schwagerina Assemblage Zone

This biozone includes four members of the Dorud Formation with thickness of 269 meters. It indicates an Asselian-Sakmarian age and includes foraminifers of the Eostaffella sp., Geinitzina postcarbonica Spandel, Pseudofusulina aff. tchernyshevi (Schellwien), Pseudofusulina cf. krotovi candata Rauser, Pseudofusulina cf. diserta Saerovich, Pseudoschwagerina sp., Schubertella sp., Schubertella cf. subkingi Dunbar et Skinner, Schwagerina cf. sphaerica Sherb, Staffella sphaerica (Abich), Triticites truncatus Chen and Triticites parvus Chen (Figs. 7, 8).

4.2 Schubertella - Mesoschubertella Assemblage Zone

This biozone includes members 1 and 2 of the Ruteh Formation with thickness of 61 meters. It indicates an Artinskian age and includes foraminifers of the Mesoschubertella sp., Geinitzina chapmani Schubert Var. Longa Sulemanov, Langella perforata Lange and Schubertella sp. (Figs. 7, 8).

4.3 Dunbarula - Neoschwagerina Assemblage Zone

This biozone includes member 3 of the Ruteh Formation with thickness of 110 meters. It indicates an Early Murgabian age and includes foraminifers of the Climacammina vavalinoids Lange, Climacammina sphaerica Potieskaya, Deckerella aff. composita Reitlinger, Dunbarula mathieui Ciry, Neoschwagerina margarita Deprat and Pachyphloia pedicula Lange (Figs. 7, 8).

4.4 Neoendothyra - Globivalvulina Assemblage Zone

This biozone includes members 4, 5 and 6 of the Ruteh Formation with thickness of 134.5 meters. It indicates a Late Murgabian age and includes foraminifers of the *Codonofusiella* sp., *Globivalvalina vonderschmitti* Reichel, *Neoendothyra broennimanni* Bozorgnia, *Neoendothyra reicheli* Reitlinger, *Neoendothyra* cf. *parva* (Lange) and *Pachyphloia cukurkoyi* De Civer. & Dess.

(Figs. 7, 8).

4.5 Codonofusiella - Reichelina Assemblage Zone

This biozone includes members 1 and 2 of the Nesen Formation with thickness of 20.5 meters. It indicates an Early Djulfian age and includes foraminifers of the *Baisalina pulchra* Reitlinger, *Codonofusiella eriki* Rauser, *Paraglobivalvulina mira* Reitlinger and *Reichelina* sp. (Figs. 7, 8).

5. Materials and Methods

The Ruteh and the Nesen formations in the study section consist mainly of limestones with chert nodules and beds. For study of radiolarian faunas, twenty-five samples (numbered from Sh-1 to Sh-25) from nodular and bedded cherts of the Ruteh Formation were collected and two samples (numbered Sh-26 and Sh-27) from bedded cherts of the Nesen Formation. All chert samples were used for the extraction of radiolarian fossils. These were about 200 g to 300 g per sample.

The chert samples were immersed in a 5 % hydrofluoric acid (HF) solution for 24 hours. After removing the HF solution, the residues were collected using 35 and 200 mesh sieves. In some samples, the same treatment was conducted several times in order to obtain sufficient radiolarian fossils. The residues, in which radiolarian fossils were concentrated by panning, were mounted on a glass slide, a medium (Entellan new) added, and covered with a cover glass.

Radiolarian specimens on slides were observed under a transmitted light microscope. In case of necessity, a scanning electron microscope (JSM-5500) was used to observe and to take photographs of radiolarian specimens.

6. Results

From Twenty-five chert samples that were collected from the members 2, 3 and 5 of the Ruteh Formation, ten samples yielded radiolarian faunas. A total of ten species of the Permian (Artinskian and Murgabian) radiolarians were recorded that include six species and four unidentified species are systematically described in the present paper.

Occurrences of radiolarian species characteristic are described below. Photomicrographs of characteristic radiolarian species are shown in Plates 1 and 2. Their horizons are shown in the columnar section (Fig. 10).

(a) Samples Sh-1 to Sh-4 of the member 2 include foraminifers, ostracods, sponge spicules, bryozoans (Pl. 2, fig. 12) and radiolarians of the *Latentifistula*? sp. A,

Orbiculiforma? sp. B and Spherical Radiolaria B. (Pl. 2, figs. 1, 4).

- **(b)** Samples Sh-5 to Sh-10 of the member 3 include foraminifers, sponge spicules, bryozoans (Pl. 2, fig. 13) and radiolarian of the *Latentifistula*? sp. B (Pl. 2, fig. 9).
- (c) Samples Sh-11 to Sh-25 of the member 5 include foraminifers, ostracods, sponge spicules, bryozoans (Pl. 2, figs. 14, 15) and radiolarians of the *Latentifistula*? sp. A,
- Latentifistula? sp. B, Latentifistularia A, Latentifistularia B, Orbiculiforma? sp. A, Spherical Radiolaria A, Spherical Radiolaria B, Spherical Radiolaria C and Ellipsoidal Radiolaria A (Pl. 1, figs. 1-15 & Pl. 2, figs. 2, 3, 5, 6-8, 10, 11).
- (d) In two samples Sh-26 and Sh-27 from bedded cherts of the Nesen Formation no radiolarians were found. These samples only include foraminifers and abundant

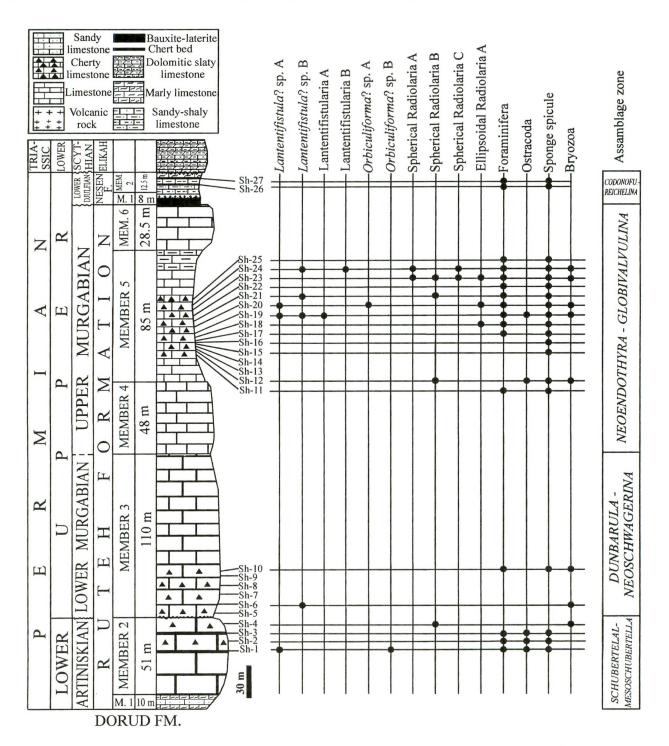


Fig. 10 Columnar section showing the horizons of radiolarian occurrence at the Gord Calleh section in northeast of Shalamzar, Central Alborz, North Iran.

sponge spicules.

According to time unit, the detected Permian (Artinskian and Murgabian) radiolarians in the Shalamzar area (Ruteh Formation) can be correlated with the Permian radiolarian zones in South China, Southwest Japan and Oregon, USA as follow:

- 1. Early Permian (Artinskian) radiolarians include the Latentifistula? sp. A, Orbiculiforma? sp. B and Spherical Radiolaria B in the member 2 of the Ruteh Formation can be correlated with the Albaillella xiaodongensis and Albaillella sinuate zones in southeast Guangxi, South China (Wang et al., 1998), Albaillella sinuate Zone in Southwest Japan (Ishiga 1990) and the Albaillella sinuate Zone in Oregon, USA (Blome and Reed, 1992), which are shown in Fig. 11. These radiolarian faunas also are correlated with the Schubertella-Mesoschubertella Assemblage Zone and the age is assigned to Artinskian.
- **2.** Late Permian (Murgabian) radiolarians include the *Latentifistula*? sp. A, *Latentifistula*? sp. B, Latentifistularia A, Latentifistularia B, Orbiculiforma? sp. A, Spherical

Radiolaria A, Spherical Radiolaria B, Spherical Radiolaria C and Ellipsoidal Radiolaria A in the members 3 and 5 of the Ruteh Formation can be correlated with the Pseudoalbaillella globosa, Follicucullus monacanthus and Follicucullus scholasticus - Follicucullus ventricosus zones in southeast Guangxi, South China (Wang et al., 1998), Pseudoalbaillella globosa, Follicucullus monacanthus and Follicucullus scholasticus zones in Dachongling section, South China (Sun et al., 2002), Pseudoalbaillella globosa, Follicucullus monacanthus and Follicucullus scholasticus - Follicucullus ventricosus zones in Southwest Japan (Ishiga 1990; Kuwahara et al., 1998) and the Pseudoalbaillella globosa, Follicucullus monacanthus and Follicucullus scholasticus zones in Oregon, USA (modified from Blome and Reed, 1992), which are shown in Fig. 11. The detected Late Permian radiolarians also are correlated with the Dunbarula-Neoschwagerina and Neoendothyra- Globivalvalina Assemblage zones, and the age is assigned to Murgabian.

On the basis of the above-mentioned correlation and

	TIME	South China SE. Guangxi (Wang et al., 1998)	South China Dachongling (Sun et al., 2002)	Southwest Japan (Kuwahara et al., 1998 & Ishiga, 1990)	Oregon, USA (modified from Blome & Reed, 1992)		North Iran Central Alborz Detected radiolarians (Present study)
z	Changhsingian			Neoalbaillella optima Neoalbaillella ornithoformis		Dorash-	
IA					Neoalbaillella optima	er Djulfian	
PERM	Wuchiapingian	Follicucullus bipartitus- Follicucullus charveti	Neoalbaillella ornithoformis	Follicucullus charveti- Albaillella yamakitai	Guadalupian-Djullian Neoalbaillella omithoformis	Lower Djulfian Upper	
2		Follicucullus scholasticus-	Follicucullus charveti	Follicucullus scholasticus-		Midian	
E	Maokouan	Follicucullus ventricosus	Follicucullus scholasticus	Follicucullus ventricosus	Follicucullus scholasticus		Latentifistula? sp. A
P P	(Guadalupian)	Follicucullus monacanthus	Follicucullus monacanthus	Follicucullus monacanthus	ugidn Pollicucullus monacanthus	rgabia	Latentifistula? sp. B Latentifistularia A Latentifistularia B Orbiculiforma? sp. A Spherical Radiolaria A
D		Pseudoalbaillella globosa	Pseudoalbaillella globosa	Pseudoalbaillella globosa	Pseudoalbaillella globosa	rg. Mu	Spherical Radiolaria B Spherical Radiolaria C Ellipsoidal Radiolaria A
PERMIAN		Pseudoalbaillella ishigai	Pseudoalbaillella longtanensis	Pseudoalbaillella longtanensis	Obber Pseudoalbaillella sp. C	Bolorian-Kuberg.	
	Qixianian (Leonardian)	Albaillella sinuata		Albaillella sinuata	vennardian Albaillella sinuata	Artinskian	Latentifistula? sp. A Orbiculiforma? sp. B
LOWER		Abaillella xiaodongensis			Геоп	Artin	Spherical Radiolaria B

Fig. 11 Time correlation of detected Permian radiolarians in the Shalamzar area, Central Alborz, North Iran with the Permian radiolarian zones in South China, Southwest Japan and Oregon, USA.

index foraminifer's age of the Ruteh Formation, the age of detected radiolarian faunas is assigned to Artinskian and Murgabian.

7. Conclusions

The Permian sequence in the Shalamzar area consists mainly of clastic rocks in the lower part (Dorud Formation) and fossiliferous carbonate rocks in the upper part (Ruteh and Nesen formations), and is bounded between two unconformities. The sedimentation gaps in lower and upper boundaries are related to the Hercynian orogenic phase. The rich foraminifer fauna indicate an Early to Late Permian (Asselian to Early Djulfian) age of the succession.

In the Present study, five biozones by foraminifers are established in the Permian System of the study section that include the *Pseudofusulina-Schwagerina* Assemblage Zone (Asselian-Sakmarian), *Schubertella-Mesosubertella* Assemblage Zone (Artinskian), *Dunbarula-Neoschwagerina* and *Neoendothyra-Globivalvalina* Assemblage zones (Murgabian), and *Codonofusiella-Reichelina* Assemblage Zone (Early Djulfian).

The Dorud Formation in the study section consists of conglomerate, sandstone, siltstone, shale, quartzite and few fossiliferous limestones that do not have chert. The clastic rocks of the Dorud Formation were deposited in continental (meandering river) and transitional (deltaic and littoral) environments (Mokhtarpour, 1997a). This formation is correlated with the *Pseudofusulina-Schwagerina* Assemblage Zone and the age is assigned to Asselian to Sakmarian.

The Ruteh and the Nesen formations in the study section consist mainly of fossiliferous limestones with chert nodules and beds that were deposited in barrier, lagoon, tidal flat and open marine environments (Mokhtarpour, 1997b). The Ruteh Formation is correlated with the Schubertella-Mesoschubertella, Dunbarula-Neoschwagerina and Neoendothyra-Globivalvalina Assemblage zones, and the age is assigned to Artinskian-Murgabian. The Nesen Formation is correlated with the Codonofusiella-Reichelina Assemblage Zone and the age is assigned to Early Djulfian.

There are remarkable chert nodules and beds in the Ruteh and Nesen formations of northern Iran in order to study radiolarians. The present study examined chert samples from the Ruteh and Nesen formations for radiolarian study and ten species of radiolarian faunas were detected. Detected Permian (Artinskian and Murgabian) radiolarians include the *Latentifistula*? sp. A, *Latentifistula*? sp. B, Latentifistularia A, Latentifistularia B, *Orbiculiforma*? sp. A, *Orbiculiforma*? sp. B, Spherical Radiolaria A,

Spherical Radiolaria B, Spherical Radiolaria C and Ellipsoidal Radiolaria A were recorded from the Ruteh Formation in northeast Shalamzar, Central Alborz for the first time. The detected radiolarians are correlative with those of south and southwestern China, eastern Thailand, central and southwest Japan, Malaysia, Turkey (NW Anatolia), Far East of Russia (Primoryie), Kazakhstan (southern Urals), North America, Philippines and Italy (western Sicily). According to time unit, all recorded radiolarians can be correlated with the Permian radiolarian zones in South China, Southwest Japan and Oregon, USA and these also are correlated with the Schubertella-Mesoschubertella Assemblage Zone (Artinskian), Dunbarula-Neoschwagerina and Neoendothyra-Globivalvalina Assemblage zones (Murgabian). On the basis of the above-mentioned correlation and index foraminifer's age of the Ruteh Formation, the age of detected radiolarian faunas is assigned to Artinskian and Murgabian.

The occurrence of *Orbiculiforma*? sp. A, *Orbiculiforma*? sp. B, Spherical Radiolaria A, Spherical Radiolaria B, Spherical Radiolaria C and Ellipsoidal Radiolaria A were reported also from Jolfa Region in northwestern Iran by Vaziri and Yao (2005) for the first time.

The present study shows that Permian rocks of northern Iran contain radiolarian faunas, but they are not numerous and their diversity is scarce. It is dependent on the depth of the sedimentary site and type of facies. Also this study shows that distribution and diversity species of radiolarian faunas in the upper part of the Ruteh Formation (member 5) is more than the lower part of this formation (Fig. 10) and the relative ratio of radiolarians to sponge spicules of sample changes through the study section. In each sample that ratio of sponge spicules increases together with the amount of lithic fragments, radiolarians decrease or disappear. For example: chert samples of the Nesen Formation (samples Sh-26 and Sh-27) contain abundant sponge spicules but they have not radiolarians. This may be related to global environmental changes.

8. Systematic Paleontology of the Radiolarians

All specimens described in this paper are deposited in Earth History laboratory, Department of Geosciences, Graduate School of Science, Osaka City University, Japan.

Spherical Radiolaria A

(Pl.1, figs. 1-4)

Description: Shell is spherical to subspherical in shape with fine pores. Pores are subcircular to subelliptical. Most species have some big pores.

Measurements (in μ m): Diameter of shells is "between 210 to 225". Average of diameter is 217.5.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples Sh-23, Sh-24), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region, Northwestern Iran by Vaziri and Yao (2005).

Occurrence: Worldwide.

Spherical Radiolaria B

(Pl.1, figs. 5-11)

Description: Shell is spherical to subspherical in shape with fine pores. Pores are subcircular to subelliptical. Most species have some big pores.

Measurements (in μ m): Diameter of shells is "between 155 to 220". Mainly "between 190 to 220". Average of diameter is 198.

Age: Early to Late Permian (Artinskian-Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples Sh-4, Sh-12, Sh-21, Sh-23), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region, Northwestern Iran by Vaziri and Yao (2005).

Occurrence: Worldwide.

Spherical Radiolaria C

(Pl. 1, figs. 12, 13)

Description: Shell is spherical to subspherical in shape with fine pores. Pores are subcircular to angular. Some species have few big pores.

Measurements (in μ m): Diameter of shells is 82.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples Sh-23, Sh-24), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region, Northwestern Iran by Vaziri and Yao (2005).

Occurrence: Worldwide.

Ellipsoidal Radiolaria A

(Pl. 1, fig. 15)

Description: Shell is ellipsoidal in shape with clear pores. Pores are angular.

Measurements (in μ m): Length of shells is 245. Diameter of shells is 170.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples Sh-18, Sh-20, Sh-23), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region,

Northwestern Iran by Vaziri and Yao (2005).

Occurrence: Worldwide.

Suborder SPUMELLARIA Ehrenberg, 1875 Family OBICULIFORMIDAE Pessagno 1973 Genus ORBICULIFORMA Pessagno, 1973

Orbiculiforma? sp. A

(Pl. 1, fig. 14)

Description: Shell is circular and flat in outline. Pores are circular or angular.

Measurements (in μ m): Diameter of shells is 450.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (sample Sh-20), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region, Northwestern Iran by Vaziri and Yao (2005).

Occurrence: North and Northwestern Iran, South China, Southwest Japan.

Orbiculiforma? sp. B

(Pl. 2, fig. 1)

Description: Shell is circular and flat in outline with a cavity in center. Shape and depth of the central shell are depression and central cavity flanked by prominent rim. Pores are circular or angular.

Measurements (in μ m): Diameter of shells is 450.

Age: Early Permian (Artinskian).

Locality: Ruteh Formation, Gord Calleh section (Sh-1), Shalamzar area, Central Alborz, North Iran. This species is also reported from the Jolfa Formation (Early Djulfian) in the Jolfa Region, Northwestern Iran by Vaziri and Yao (2005).

Occurrence: North and Northwestern Iran, South China, Southwest Japan.

Superfamily LATENTIFISTULIDEA Nazarov and Ormiston, 1983

Family LATENTIFISTULIDAE Nazarov and Ormiston, 1983

Genus LATENTIFISTULA Nazarov and Ormiston, 1983

Latentifistula? sp. A

(Pl. 2, figs. 2-6)

Description: This species is characterized by having three characteristic arms, each with strong longitudinal beams connected with transverse bars, coarse sponge meshwork at distal part and a small spine at tip. We know latentifistularia with an initial spicule, with one apical and two to three basal spines, and perforate to partly perforate microsphere. Here, species are broken and one arm can be seen. Shell is

knobby and cylindrical in shape with long furrow and small pores. Pores irregularly arranged.

Measurements (in μ m): Length of arms is "between 450 to 520" and average of length is 453. Diameter of arms is "between 70 to 160" and average of diameter is 95. Some species have rounded bulge in up with 138 average diameters.

Age: Early to Late Permian (Artinskian-Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples 1, Sh-19, Sh-20), Shalamzar area, Central Alborz, North Iran.

Occurrence: North Iran, South China, Central and Southwest Japan, Eastern Thailand, Far East Russia (Primoryie), Kazakhstan (southern Urals), Malaysia, Turkey (NW Anatolia), North America, Italy (western Sicily).

Latentifistula? sp. B

(Pl. 2, figs. 8-10)

Description: Here, species are broken. The first species (fig. 4) has two arms. The second species (fig. 5) has one arm and only the third species (fig. 6) shows three arms. Shells are knobby in shape with fine and coarse-grained pores. Pores irregularly arranged.

Measurements (in μ m): Length of arm in second species is 550. Length of arms in first species is 140 and 460, and in third species they are 200, 270 and 300.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (samples Sh-6, Sh-19, Sh-21, Sh-24), Shalamzar area, Central Alborz, North Iran.

Occurrence: North Iran, South China, Central and Southwest Japan, Eastern Thailand, Far East Russia (Primoryie), Kazakhstan (southern Urals), Malaysia, Turkey (NW Anatolia), North America, Italy (western Sicily).

Latentifistularia A

(Pl. 2, fig. 7)

Description: Latentifistularia has three arms. Here, species is broken and only one arm can be seen. Shell is knobby in shape and has rounded bulge in up. Shell is perforate and pores are angular.

Measurements (in μ m): Length of shell is 540 and width of shell is 100. Width of rounded bulge in up is 240. This rounded bulge has a big hole in center.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (sample Sh-19), Shalamzar area, Central Alborz, North Iran.

Occurrence: North Iran, South and Southwestern China, Southwest Japan, Malaysia, Thailand, Turkey (NW

Anatolia), North America, Philippines.

Latentifistularia B

(Pl. 2, fig. 11)

Description: Species is fusiform with big spine. Shell is perforate and pores are angular.

Measurements (in μ m): Length of species is 160 and its diameter is 75. Length of spine is 85.

Age: Late Permian (Murgabian).

Locality: Ruteh Formation, Gord Calleh section (sample Sh-24), Shalamzar area, Central Alborz, North Iran.

Occurrence: North Iran, South and Southwestern China, Southwest Japan, Malaysia, Thailand, Turkey (NW Anatolia), North America, Philippines.

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Scanning photomicrographs of Early and Late Permian (Artinskian and Murgabian) radiolarians from the Rutch Formation, Gord Calleh section, Central Alborz, North Iran

Plate 1 (Murgabian)

Figs. 1, 2:	Spherical Radiolaria A, sample no. Sh-23.
Figs. 3, 4:	Spherical Radiolaria A, sample no. Sh-24.
Fig. 5:	Spherical Radiolaria B, sample no. Sh-12.
Fig. 6:	Spherical Radiolaria B, sample no. Sh-21.
Figs. 7-11:	Spherical Radiolaria B, sample no. Sh-23
Figs. 12, 13:	Spherical Radiolaria C, sample no. Sh-23.
Fig. 14:	Orbiculiforma? sp. A, sample no. Sh-20.
Fig. 15:	Ellipsoidal Radiolaria A, sample no. Sh-20.

Plate 1 X330 50 m 10 X430 50 mm 20 17 SEI 20 17 SEI

Plate 2 * Samples Sh-1, Sh-4: Artinskian Others: Murgabian

Fig. 1:	Orbiculiforma? sp. B, sample no. Sh-1.
Figs. 2, 5:	Latentifistula? sp. A, sample no. Sh-19.
Figs. 3, 6:	Latentifistula? sp. A, sample no. Sh-20.
Fig. 4:	Latentifistula? sp. A, sample no. Sh-1.
Fig. 7:	Latentifistularia A, sample no. Sh-19.
Fig. 8:	Latentifistula? sp. B, sample no. Sh-19.
Fig. 9:	Latentifistula? sp. B, sample no. Sh-6.
Fig. 10:	Latentifistula? sp. B, sample no. Sh-21.
Fig. 11:	Latentifistularia B, sample no. Sh-24.
Fig. 12:	Bryozoa, sample no. Sh-4.
Fig. 13:	Bryozoa, sample no. Sh-6.
Fig. 14:	Bryozoa, sample no. Sh-19.
Fig. 15:	Bryozoa, sample no. Sh-24.

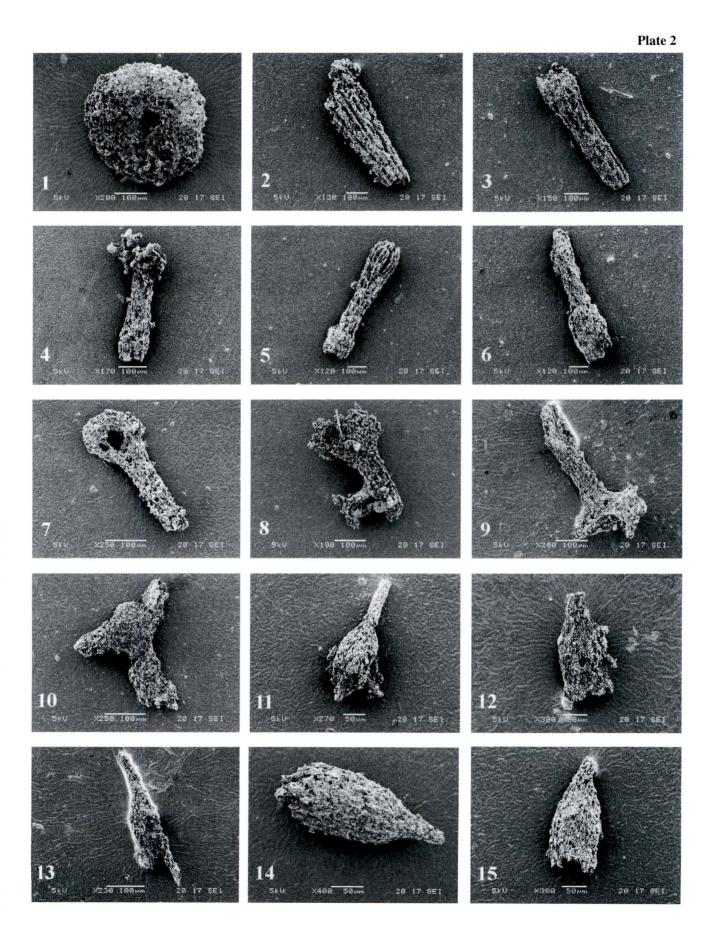


Plate 3

Dorud Formation, Gord Calleh section, Early Permian (Asselian- Sakmarian), Shalamzar area, Central Alborz, North Iran

- Figs. 1, 2: Girvanella permica Pia, X20, thin section no. H.V-6.
- Fig. 3: Schwagerina sp., X20, thin section no. H.V-6.
- Fig. 4: Pseudoschwagerina sp., X20, thin section no. H.V-6.
- Fig. 5: Pseudoschwagerina sp. and Schubertella sp., X29, thin section no. H.V-6.
- Fig. 6: Pseudofusulina sp., X20, thin section no. H.V-6.
- Fig. 7: Pseudofusulina aff. p. tchernyshevi (Schellwien), X20, thin section no. H.V-6.
- Fig. 8: Pseudofusulina cf. diserta Sherbovich, X12.6, thin section no. H.V-10.
- Fig. 9: Tuberitina sp., X100, thin section no. H.V-8.

Ruteh Formation, Gord Calleh section, Early Permian (Artinskian), Shalamzar area, Central Alborz, North Iran

- Fig. 10: Permocalculus sp., X20, thin section no. H.V-11.
- Fig. 11: Hemigordius sp., X40, thin section no. H.V-27.
- Fig. 12: Palaeospiroplectammina sobhaniansis Partoazar and Vaziri, X100, thin section no. H.V-26.
- Fig. 13: Palaeobigenerina sp., X40, sample no. H.V-18.
- Fig. 14: Climacammina sphaerica Potieskaya, X40, thin section no. H.V-21.
- Fig. 15: Kahlerina packytheca Koch. Devide et Ramors, X100, thin section no. H.V-21.

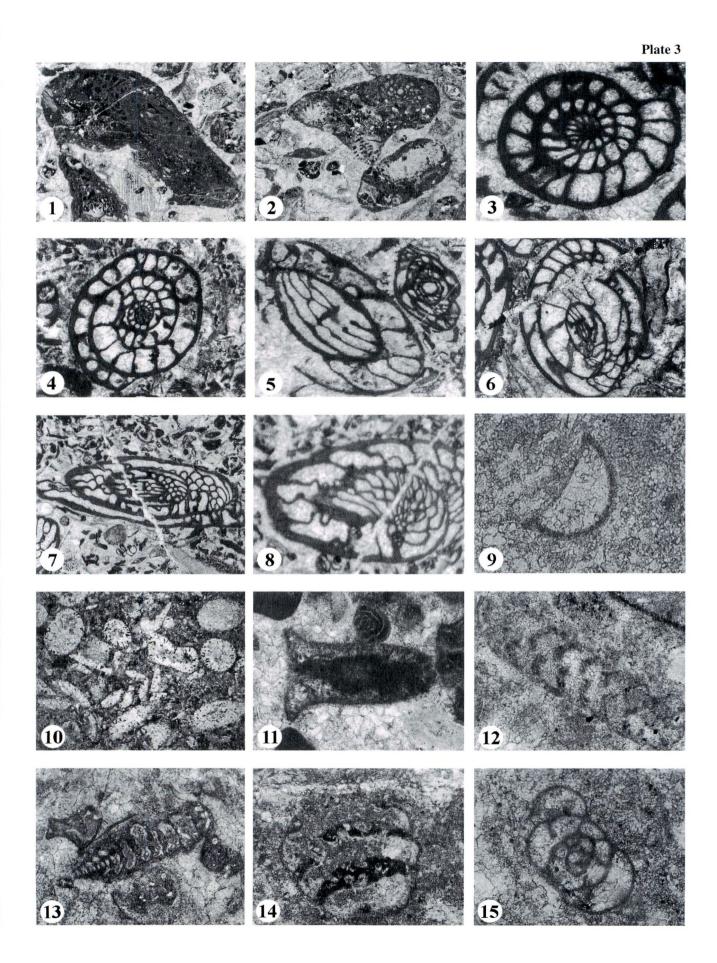


Plate 4

Ruteh Formation, Gord Calleh section, Late Permian (Murgabian), Shalamzar area, Central Alborz, North Iran

- Fig. 1: Neoschwagerina margarita Deprat, X20, thin section no. H.V-37.
- Fig. 2: Cryptosetida (Langella) bozorgniansis Partoazar and Vaziri, X40, thin section no. H.V-49.
- Fig. 3: Langella perforata (Lange), X100, thin section no. H.V-34.
- Fig. 4: Deckerella composita Reitlinger, X40, thin section no. H.V-37.
- Fig. 5: Climacammina sphaerica Potieskaya, X20, thin section no. H.V-37.
- Fig. 6: Climacammina valvulinoides Lange, X20, thin section no. H.V-66.
- Fig. 7: Langella cf. acanthi (Lange), X20, thin section no. H.V-60.
- Fig. 8: Pachyphloia cukurkoyi De Civer. & Dess., X100, thin section no. H.V-105.
- Fig. 9: Geinitzina chapmani Schubert Var. Longa Sulemanov, X40, thin section no. H.V-66.
- Fig. 10: Schubertella sp., X40, thin section no. H.V-65.
- Fig. 11: Globivalvulina sp., X100, thin section no. H.V-64.
- Fig. 12: Globivalvulina cf. vonderschmitti Reichel, X100, thin section no. H.V-60.
- Fig. 13: Neoendothyra reicheli Reitlinger, X63, thin section no. H.V-95.
- Fig. 14: Neoendothyra broennimanni Bozorgnia, X100, thin section no. H.V-64.
- Fig. 15: Pseudovermiporella sp., X12.6, thin section no. H.V-96.

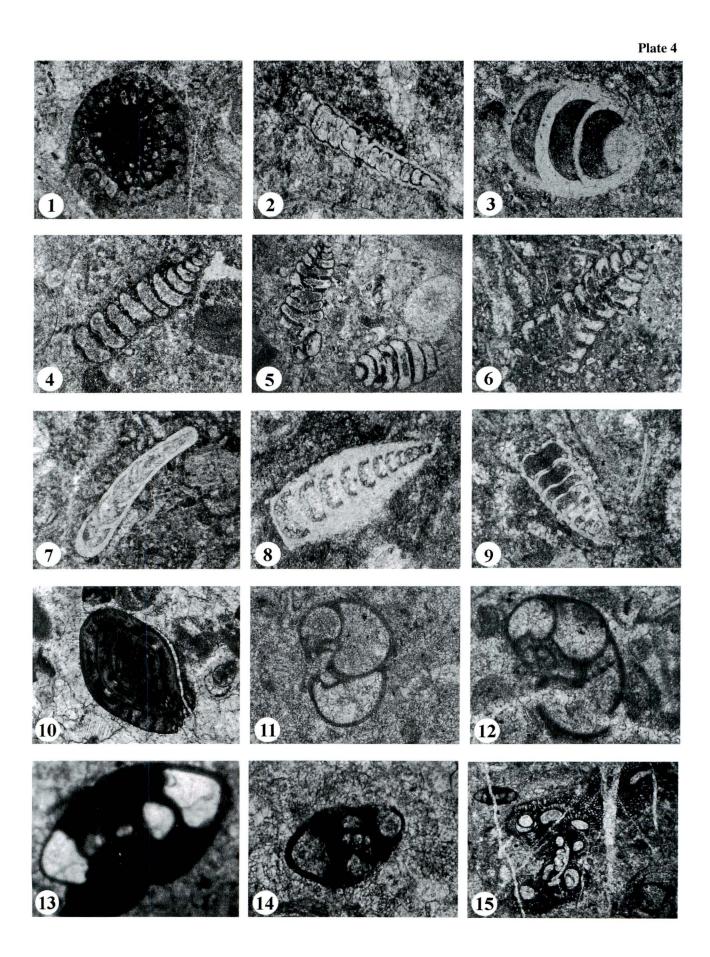


Plate 5

Ruteh Formation, Gord Calleh section, Late Permian (Murgabian), Shalamzar area, Central Alborz, North Iran

Fig. 1: Vermiporella nipponica Endo, X40, thin section no. H.V-71.

Nesen Formation, Gord Calleh section, Late Permian (Early Djulfian), Shalamzar area, Central Alborz, North Iran

- Fig. 2: Tubiphytes obscurus Maslov, X40, thin section no. H.V-121.
- Fig. 3: Geinitzina postcarbonica Spendel, X100, thin section no. H.V-120.
- Fig. 4: Globivalvulina vonderschmitti Reichel, X40, thin section no. H.V-122.
- Fig. 5: Paraglobivalvulina sp., X40, thin section no. H.V-120.
- Fig. 6: Paraglobivalvulina mira Reitlinger, X40, thin section no. H.V-120.
- Fig. 7: Hemigordius sp., X40, thin section no. H.V-121.
- Fig. 8: Codonofusiella sp., X100, thin section no. H.V-121.
- Fig. 9: Dunbarula mathieui Ciry, X100, thin section no. H.V-120.
- Fig. 10: Staffella sp., X40, thin section no. H.V-121.
- Fig. 11: Reichelina sp. and Dagmarita sp., X40, thin section no. H.V-122.
- Fig. 12: Reichelina sp., X100, thin section no. H.V-121.
- Figs. 13, 14: Baisalina pulchra Reitlinger, X100, thin section no. H.V-122.

Elikah Formation, Gord Calleh section, Early Triassic (Scythian), Shalamzar area, Central Alborz, North Iran

Fig. 15: Spirorbis sp., X40, thin section no. H.V-127.

