

Challenges in developing better observational constraints and models for aerosols : Emerging ideas for design and use of future observing systems

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Aerosol interaction with meteorology

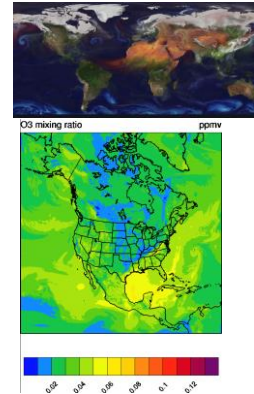
- 1 Low Cloud Feedback
- 2 High Cloud Feedback
- 3 Convective Storm Systems
- 4 Cold Cloud & Precipitation
- 5 Aerosol Attribution and Air Quality
- 6 Aerosol Processing, Removal and Redistribution
- 7 Aerosol Direct Effect and Absorption
- 8 Aerosol Indirect Effect

Aerosol modeling : Status and future



Observations

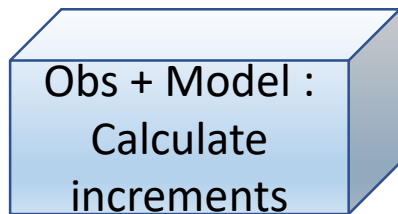
- Multi-platform multi-spectral aerosol retrievals from satellites
- In-situ aerosol mass and optical properties



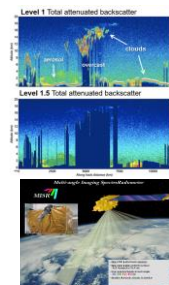
Aerosol modeling

- Best possible meteorology : Multi-model ensemble
- Aerosol mass, size distribution and optical properties

Data Assimilation



- Point by point observation uncertainty
- Choice of aerosol variables
- Clear sky radiance

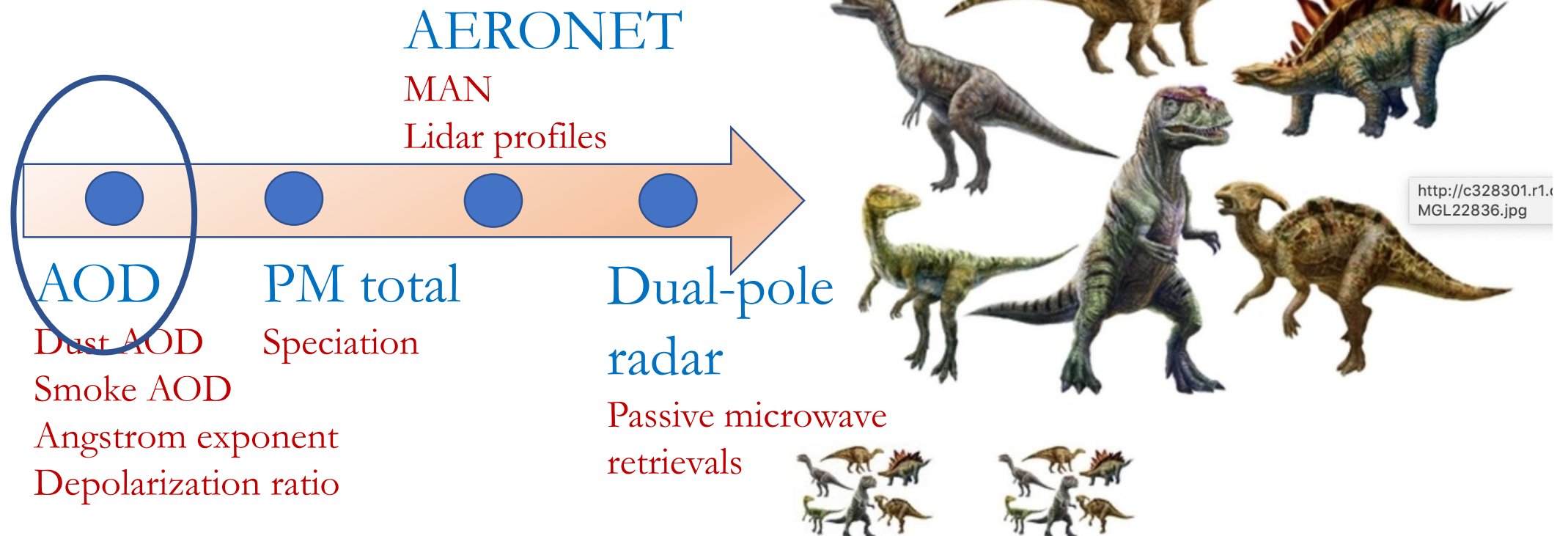


Evaluation and development

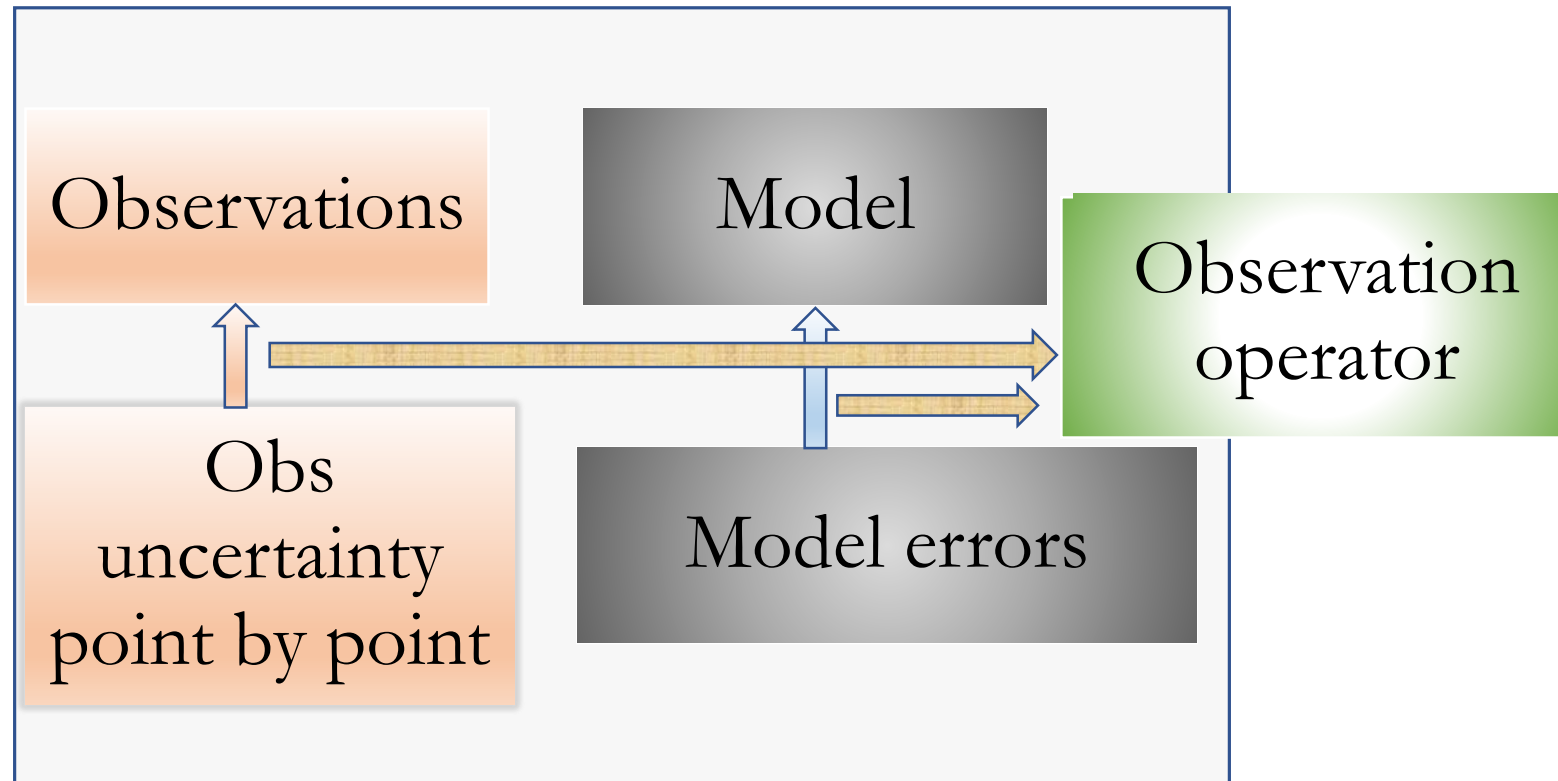
- Use of aerosol variables not used in assimilation – process-based improvements to models

Observation constrained modeling

Good representation of atmospheric aerosol sources and sinks



Observation constrained modeling (Contd)



Improvements to parameters [Johnson et al. 2019]

Direct assimilation (e.g. AOD)

Inversion for emission

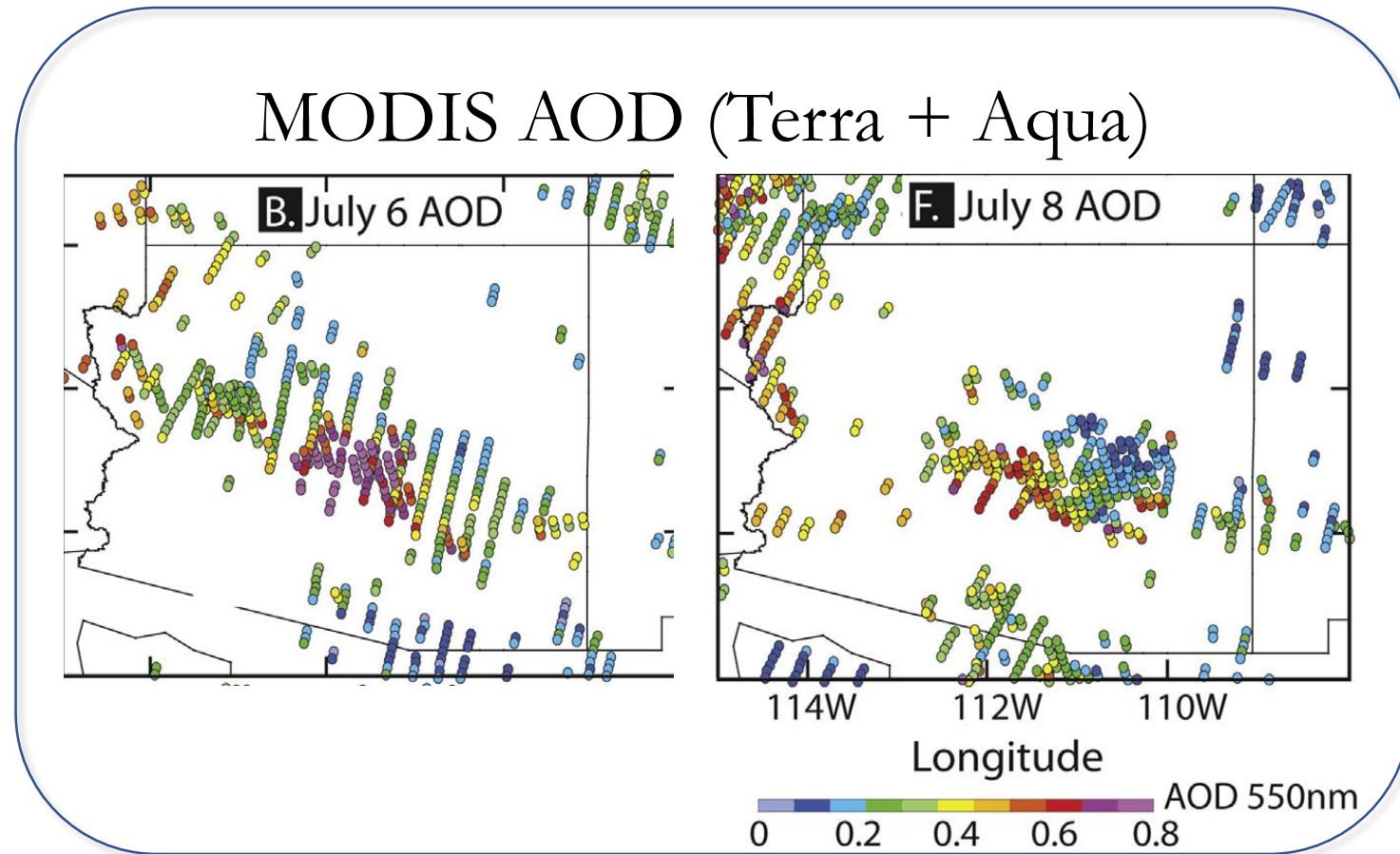
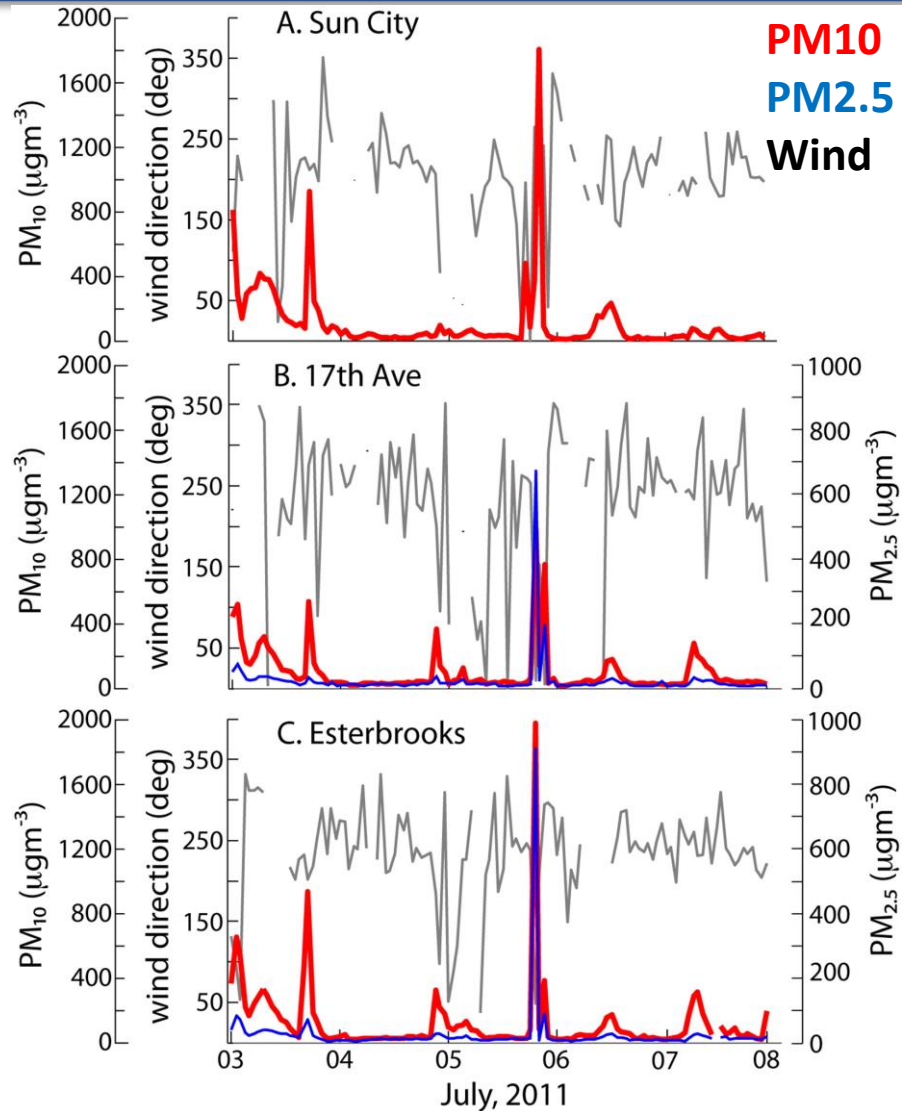
Identify key mechanisms for aerosols

Aerosol regional modeling

**Case study of dust storms in Arizona : Mile
high wall of dust hits Phoenix**

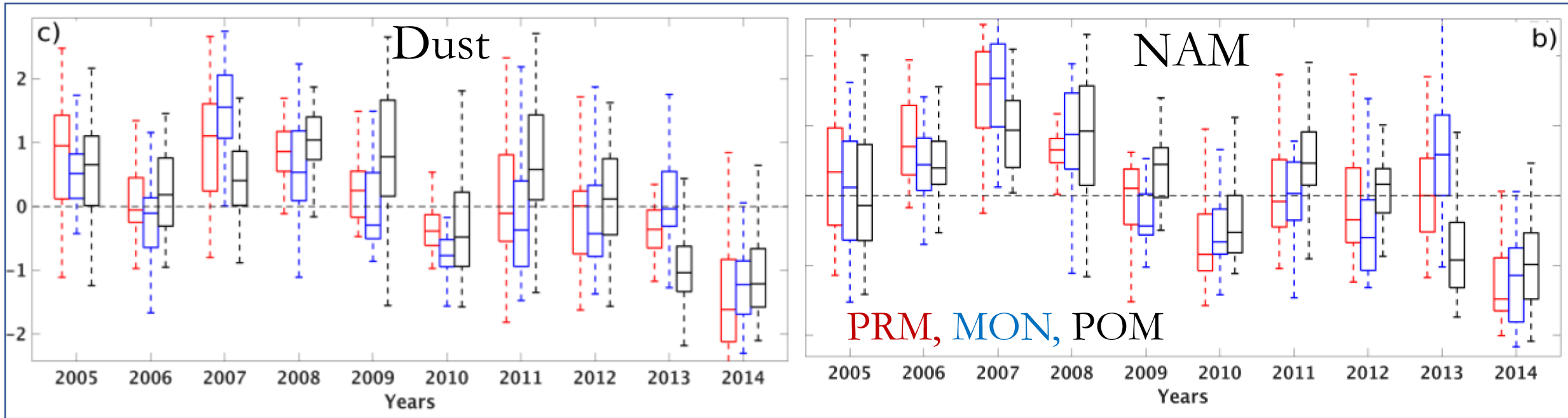


Aerosol optical properties, dust concentrations, and transport



A. Raman et al., 2014

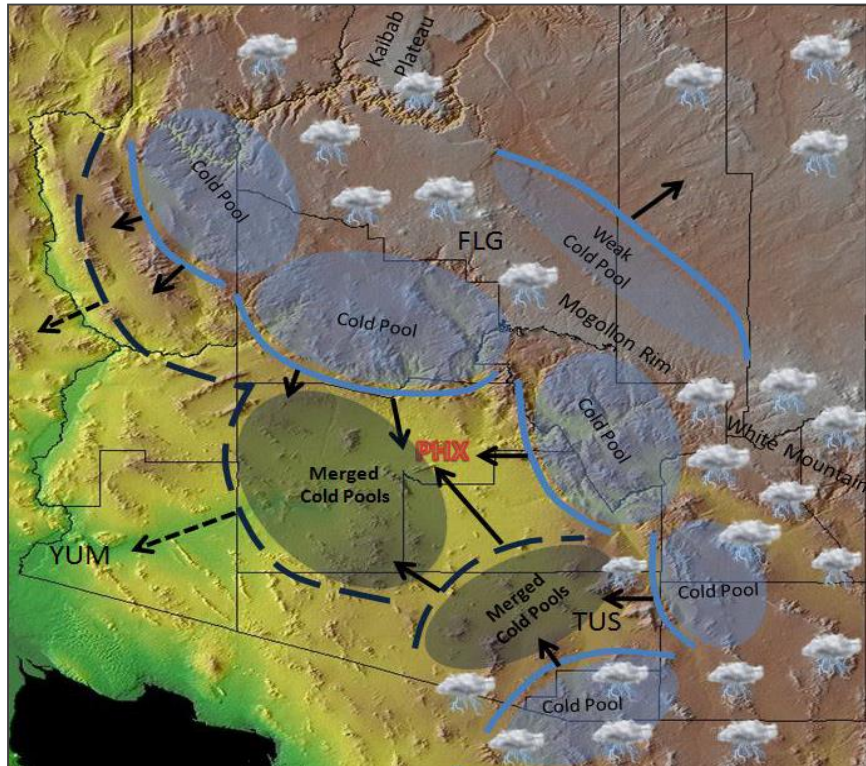
Aerosol trends over the dusty SouthWest



Change in TERRA AOD Standardized anomaly over a decade shows statistically significant **decrease in anomalies** over dusty hotspots and North American Monsoon (NAM) alley **for pre-monsoon and monsoon.**

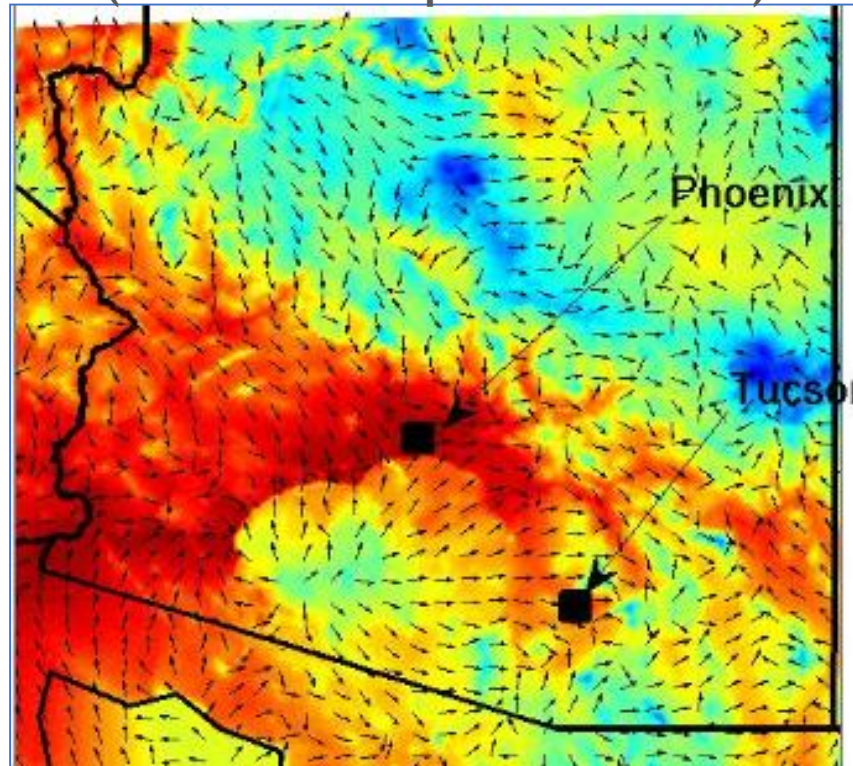
Modeling of Arizona dust storms

Conceptual Diagram of Cold Pool Formation and Movement



**Strong SE winds $23-25\text{ms}^{-1}$
and cold pool
development !**

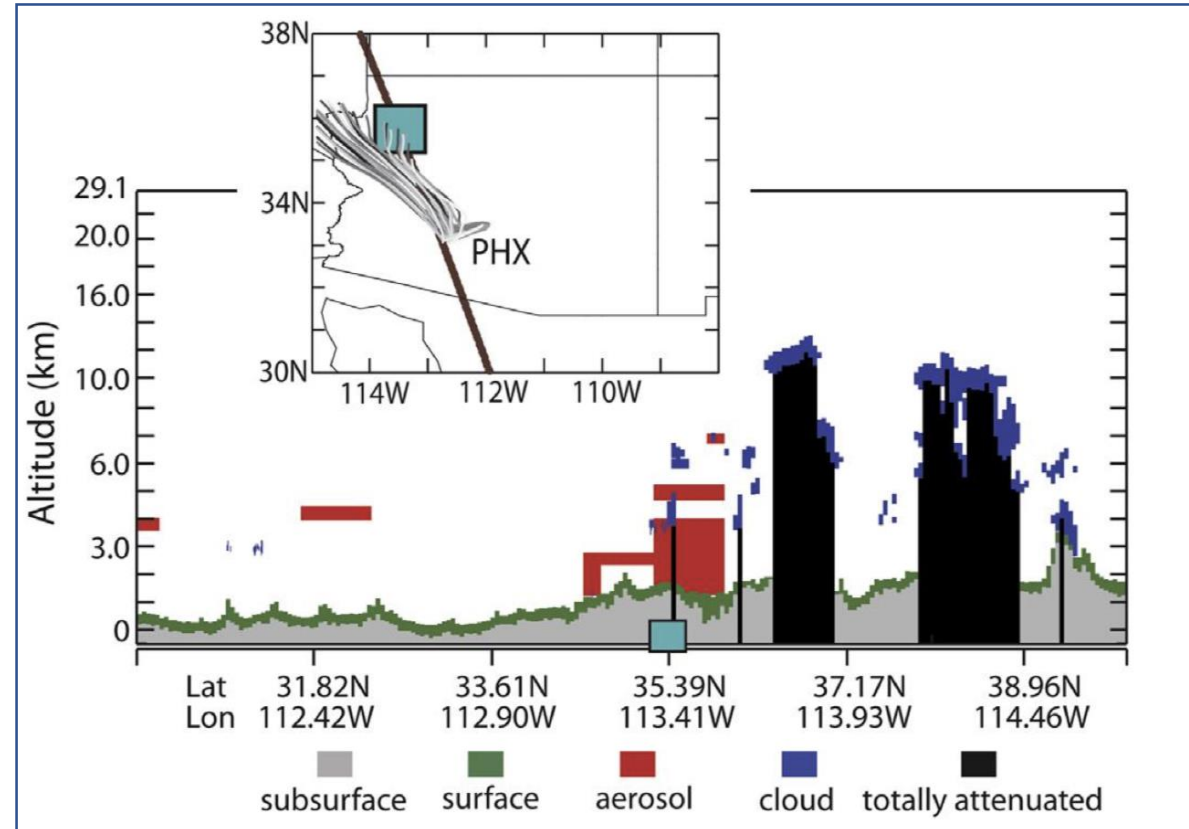
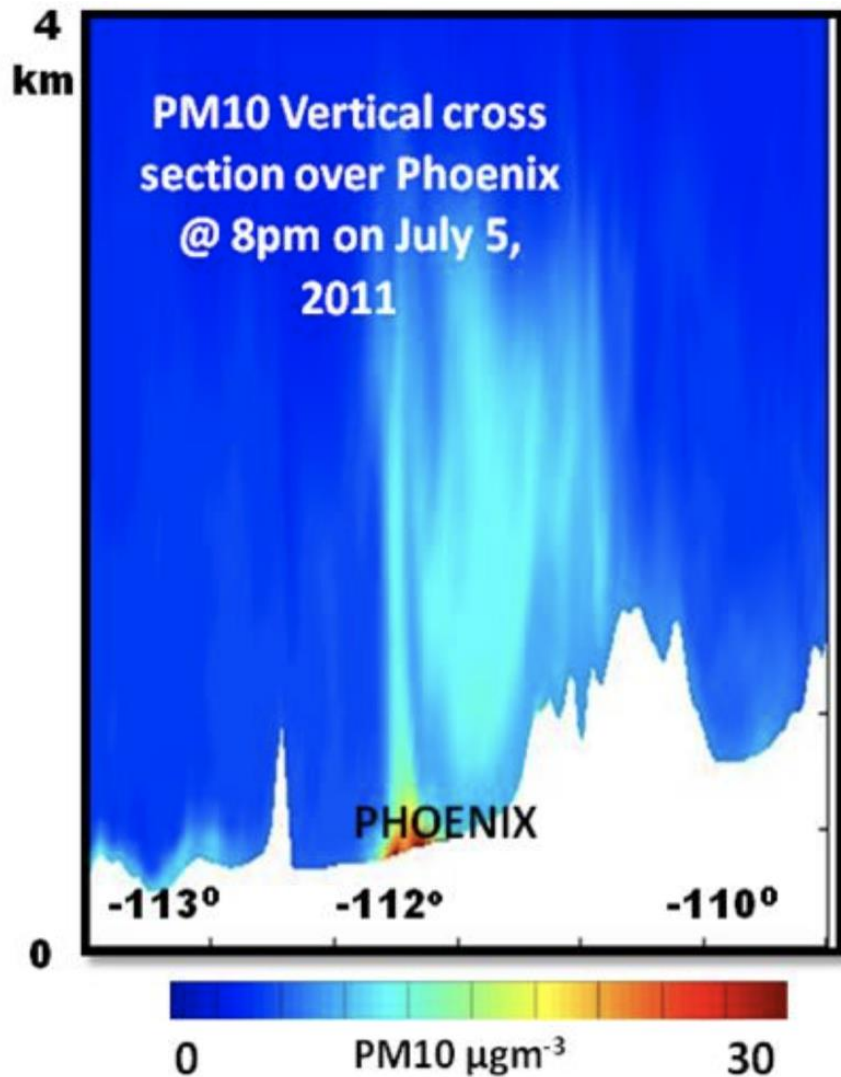
WRF-Chem Simulated Cold Pool
(2011/07/05 8pm Local Time)



WRF-Chem 1.2 km inner domain simulation of July 5, 2011 haboob

- Decrease in 2m temperature
- Downbursts and surface divergence
- Dust uplift

Modeling of Arizona dust storms (Contd)

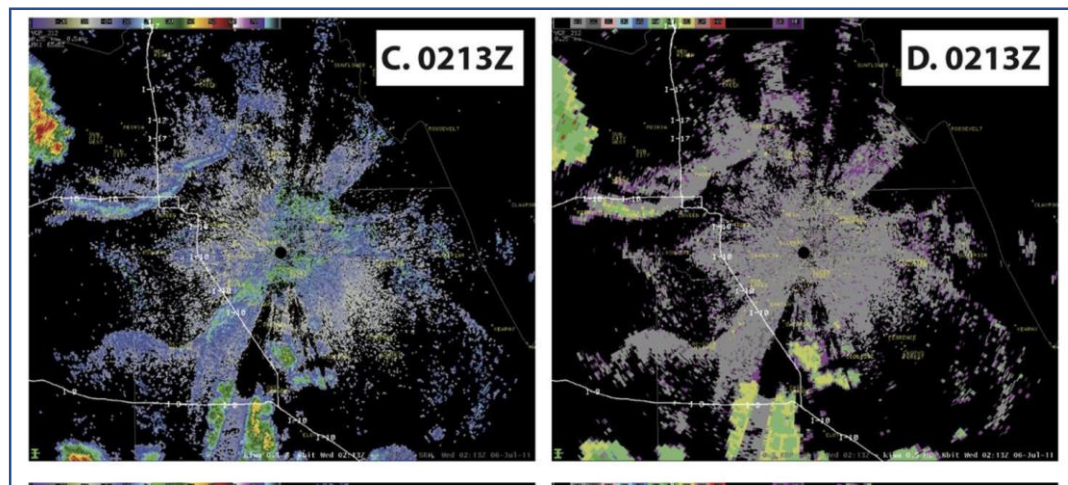


HYSPLIT ensemble trajectories of dust transported from Phoenix to Northwestern AZ. Also seen from CALIPSO aerosol classification

Limitations and future direction

$$F = CSsp (u_{10} - u_{*t}) u_{10}^2 \quad \text{if } u_{10} > u_{*t}$$

S is the erodibility, sp is the fraction of soil composition, u10 is the 10m wind speed, u_{*t} is the threshold wind friction velocity, and C is a tunable constant



Dual-pole radar reflectivities
and hydrometeor classification

Convective scale modeling, using wind
probability instead of wind speed

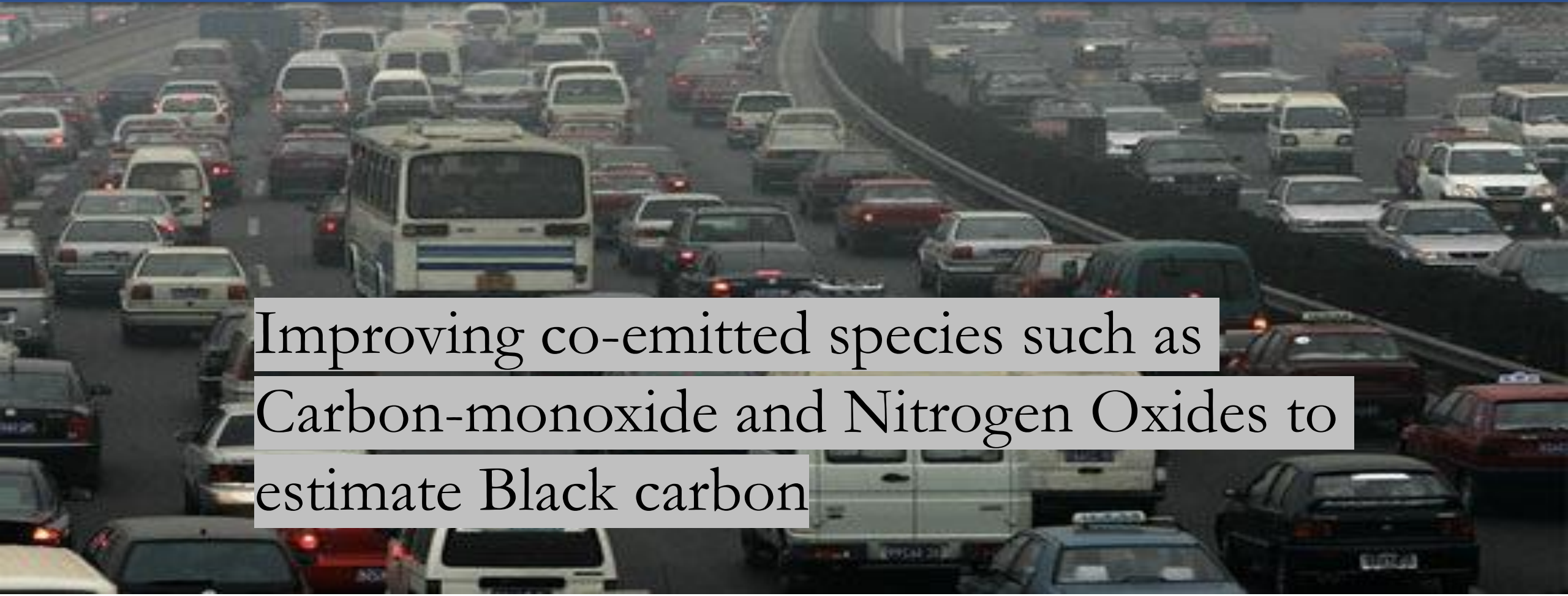
Fine scale meteorology
processes, dust impacts

Lack of
observations

GEO+LEO, MISR
plume height and
spherical, non-spherical
AOD

Missing dust sources
using real time NDVI

Black carbon aerosols and uncertainties due to lack of observations

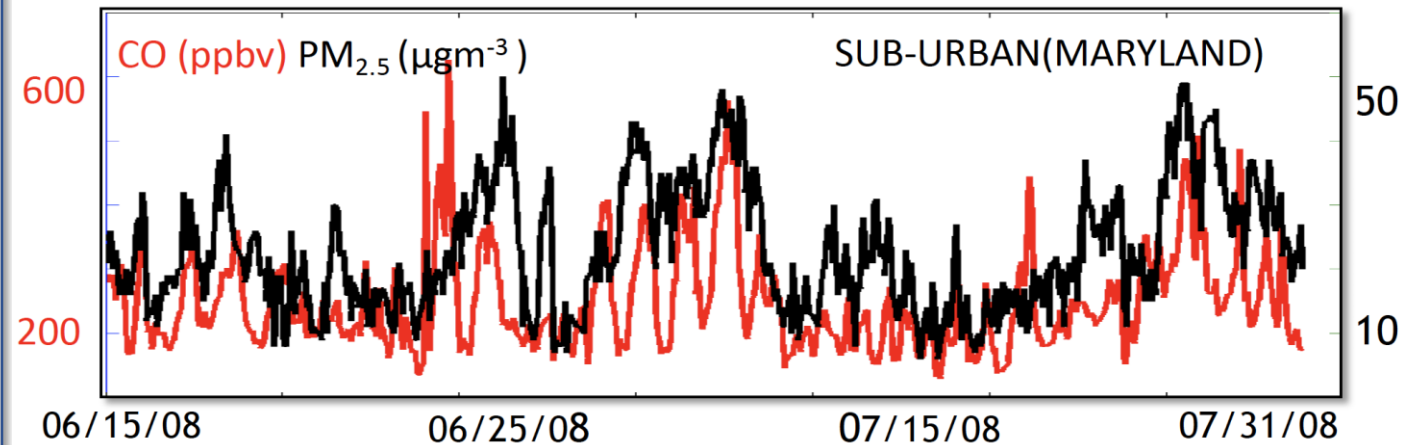


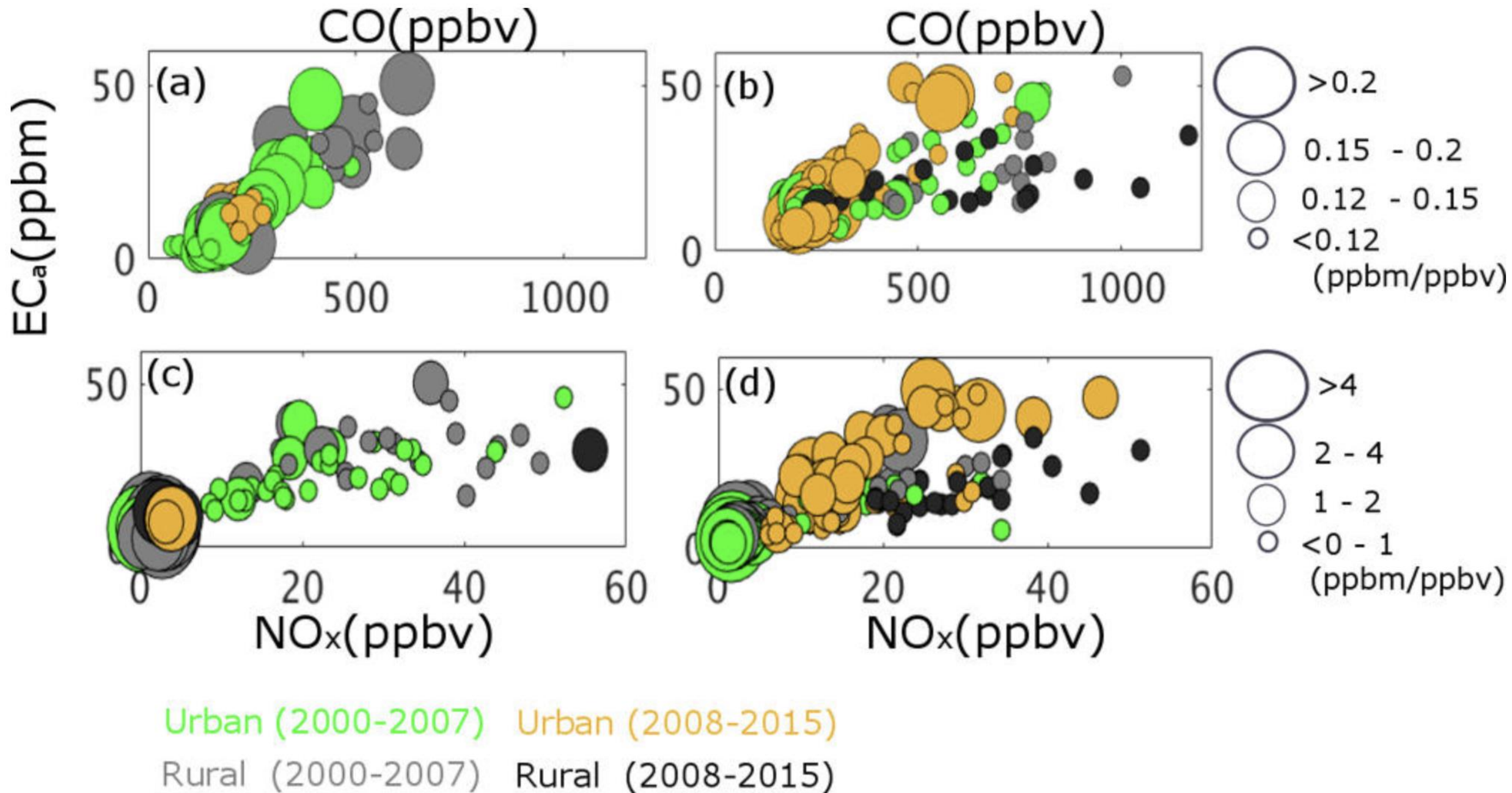
Improving co-emitted species such as Carbon-monoxide and Nitrogen Oxides to estimate Black carbon

Black carbon aerosols : Requirements for observations

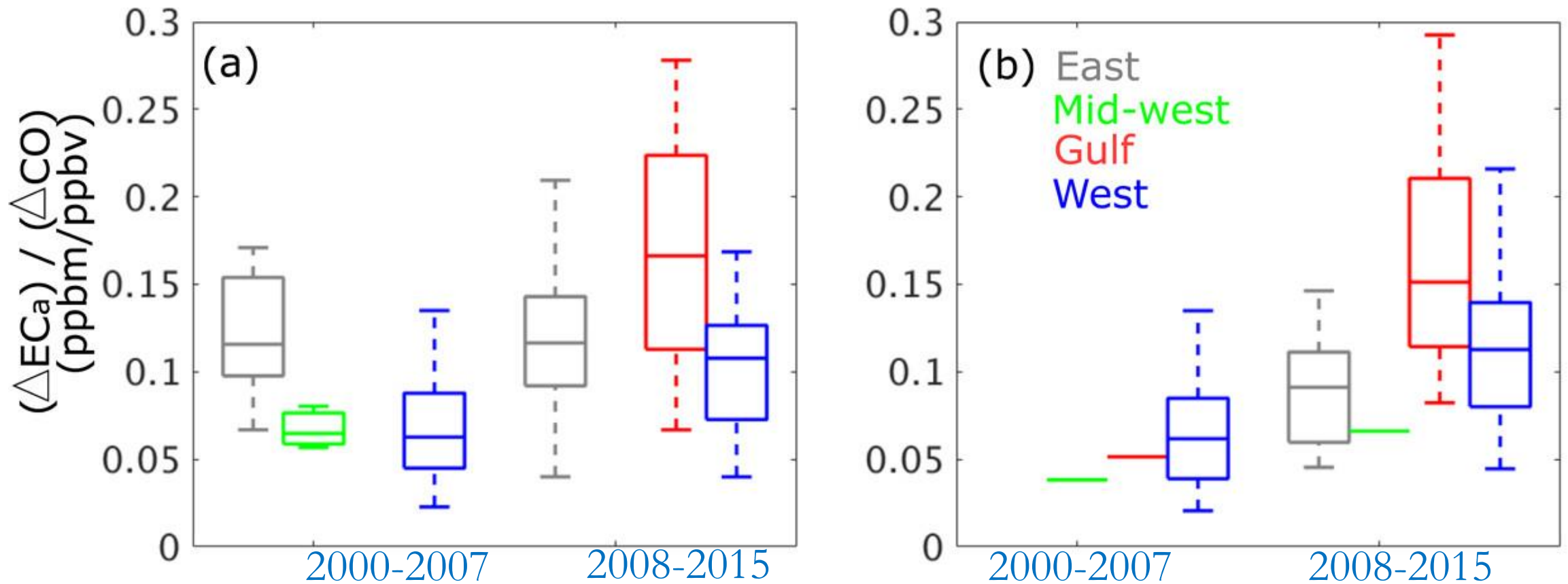
- Reductions in BC, PM_{2.5} (PM 2.5 μm or less in diameter) emissions from on-road diesel engines have not been significant (e.g. Dallmann, T. R. and Harley, R. A, 2010).
- Large uncertainties in BC mixing state in aerosol models.

Using co-emitted species

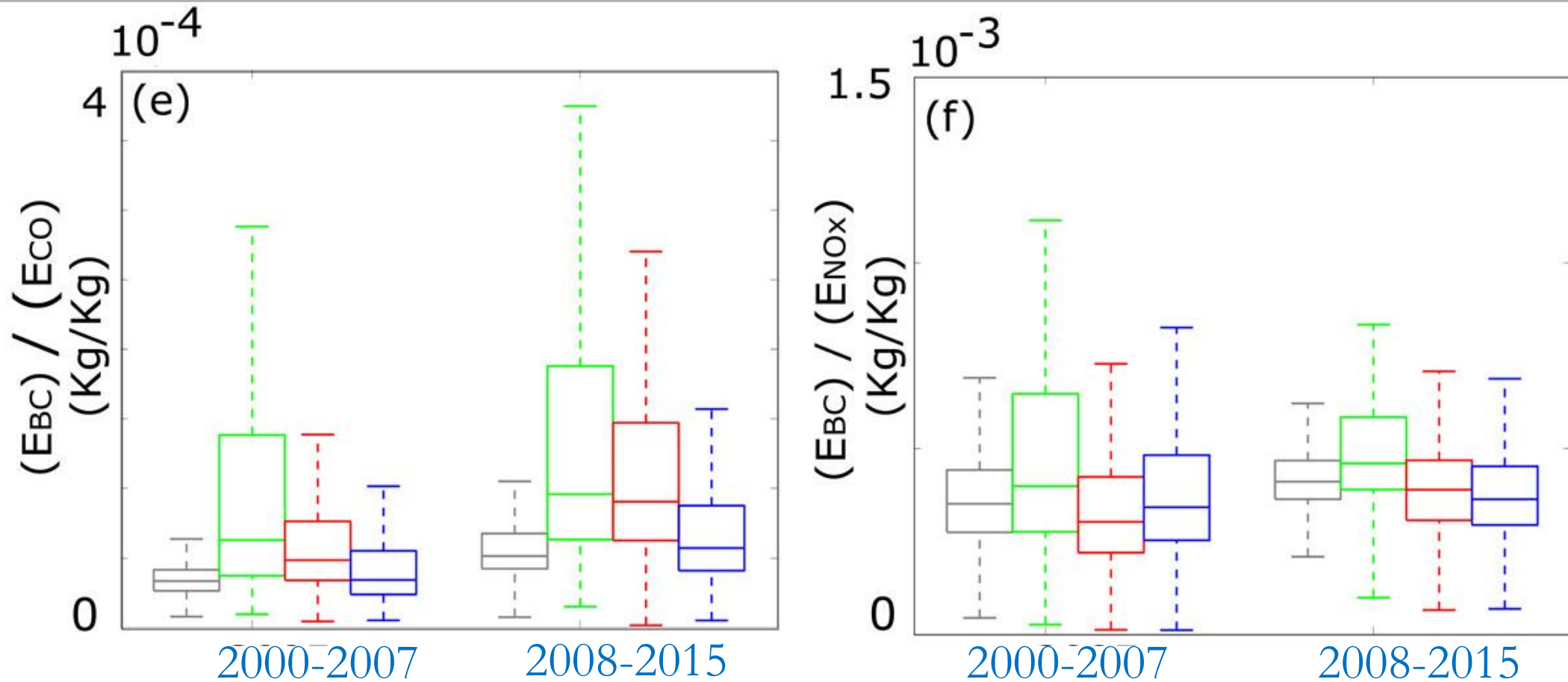




Co-emitted gases : Utility in Elemental carbon enhancement trends



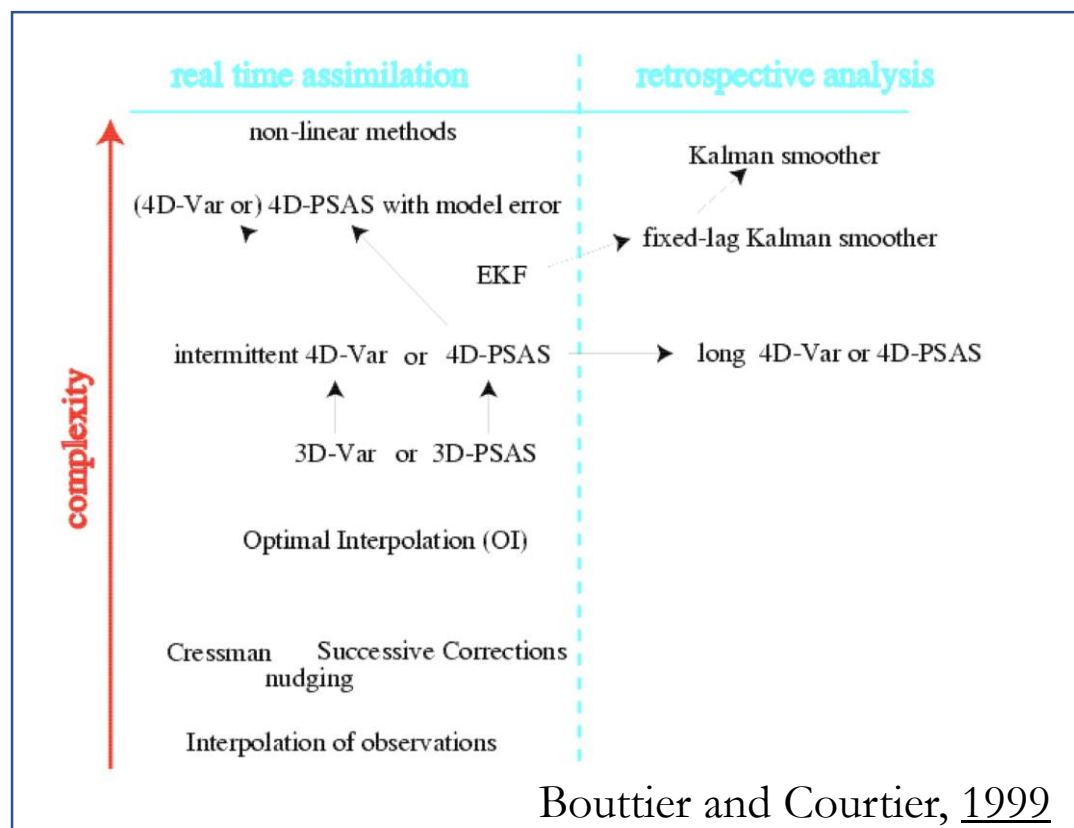
Co-emitted gases : Utility in BC emission trends (Contd)



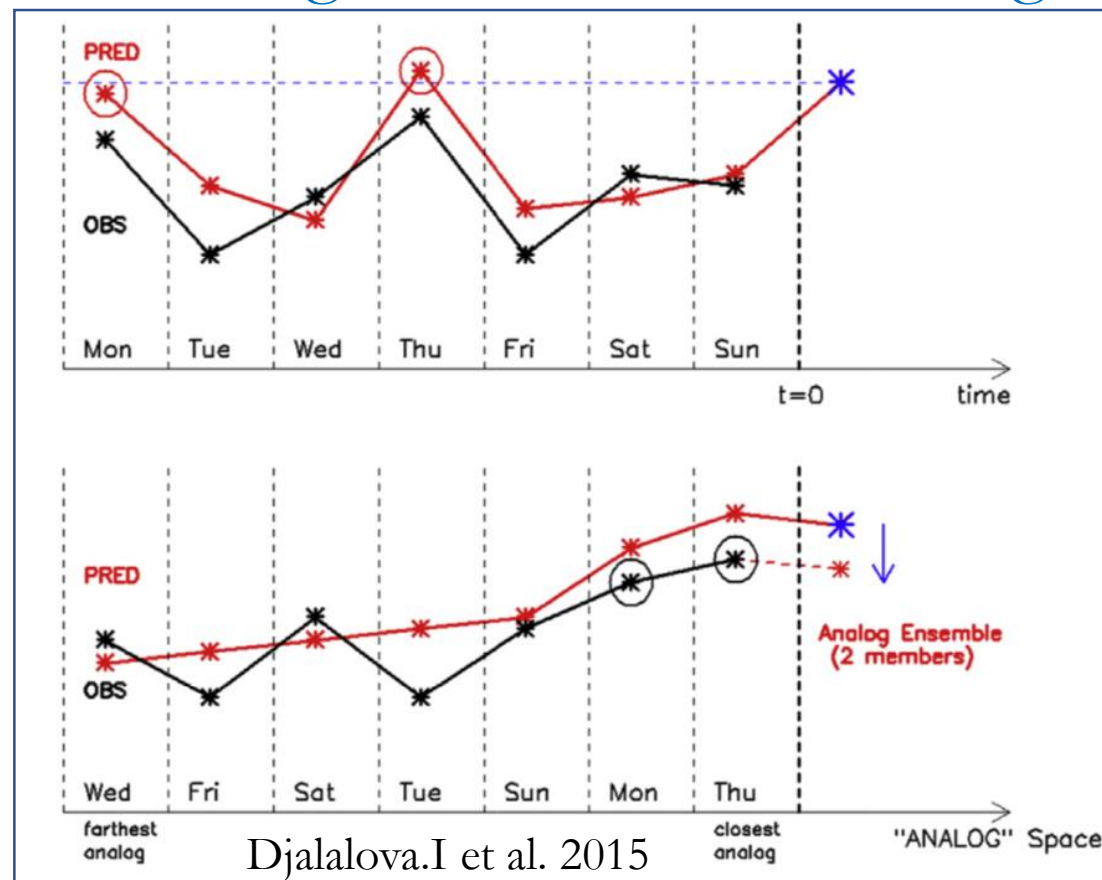
Take home messages

- CO and NO_x provide a novel pathway to improve observational coverage for sparsely measured , co-emitted aerosol species like Black Carbon.
- Such approach can significantly improve characterization of black carbon aerosol sources from satellite retrievals due to ample measurements of CO and NO_x from space.
- Developing tracers for regional and sectoral CO in addition to total CO can improve source attribution and emission fluxes for black carbon aerosol in models.

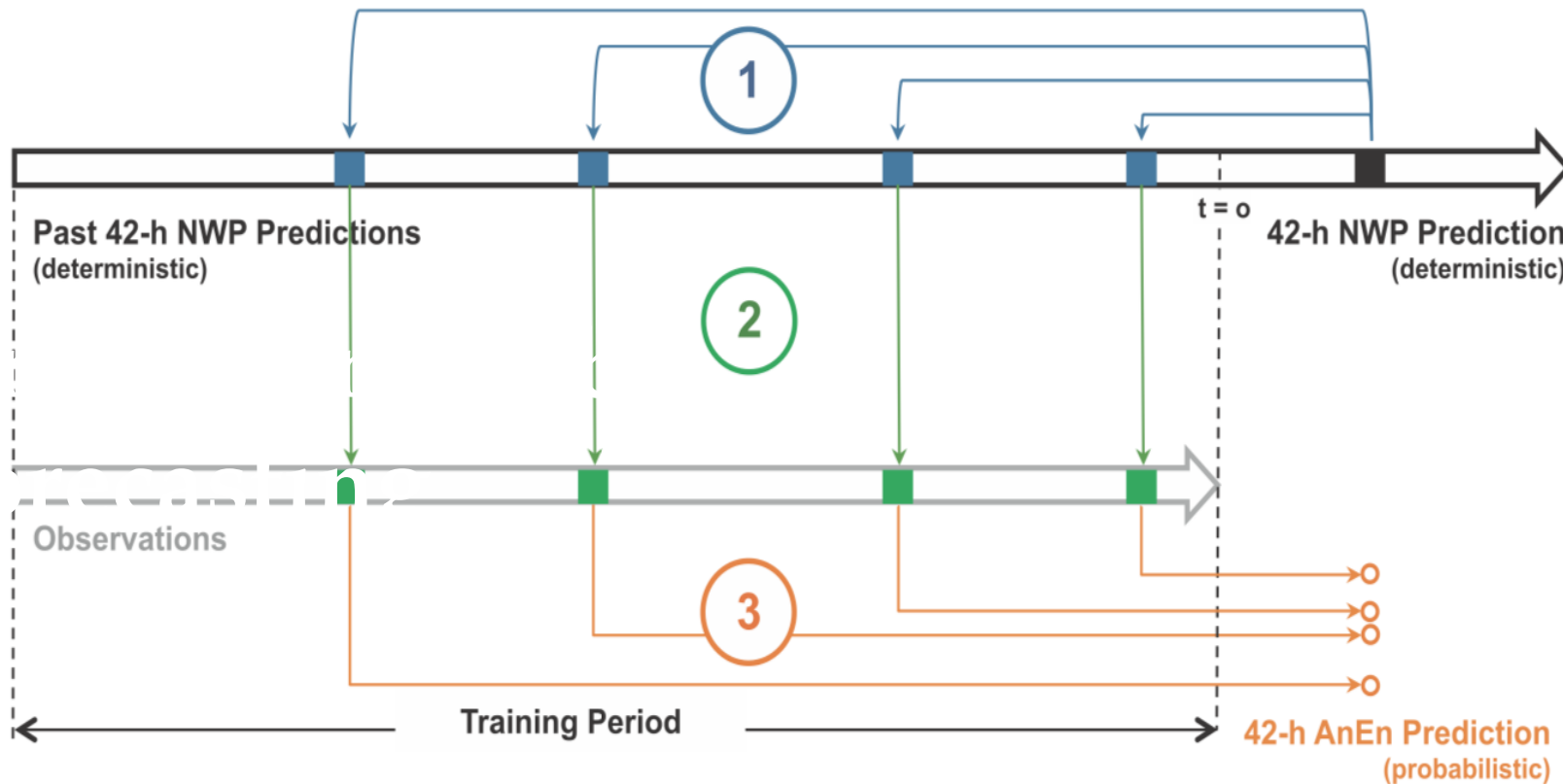
Combining models and observations



Analog Ensemble Forecasting



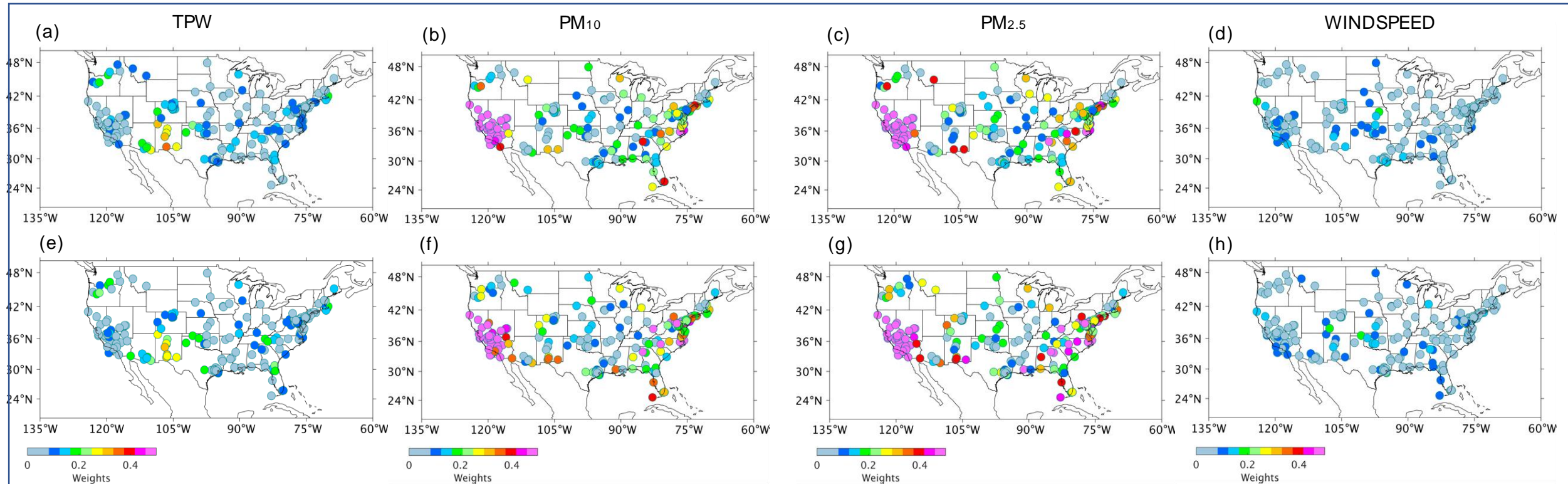
Analog Ensemble Forecasting



Analog Kalman Filter operates on the analog set of forecasts and the predicted errors in the historical forecasts

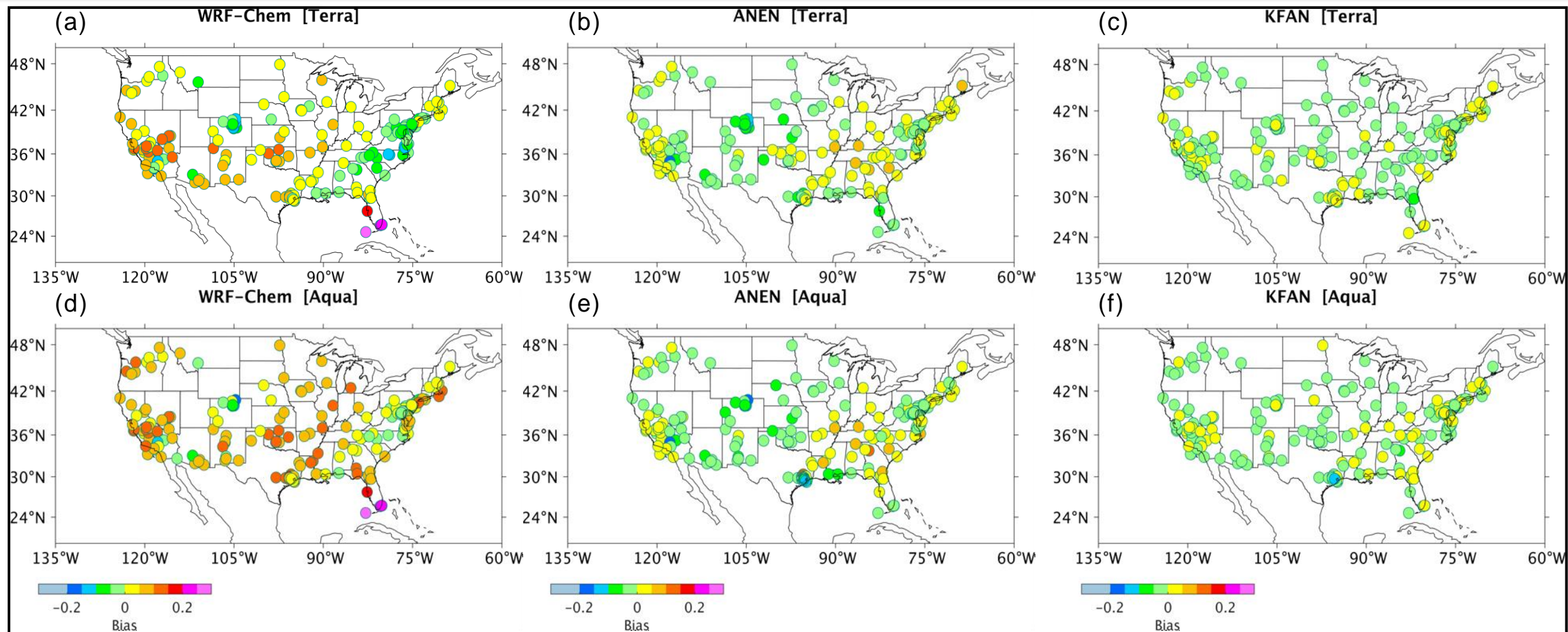
Delle Monache et al. 2011
Delle Monache et al. 2006

Predictor weights correspond to AOD-predictor correlation in the model



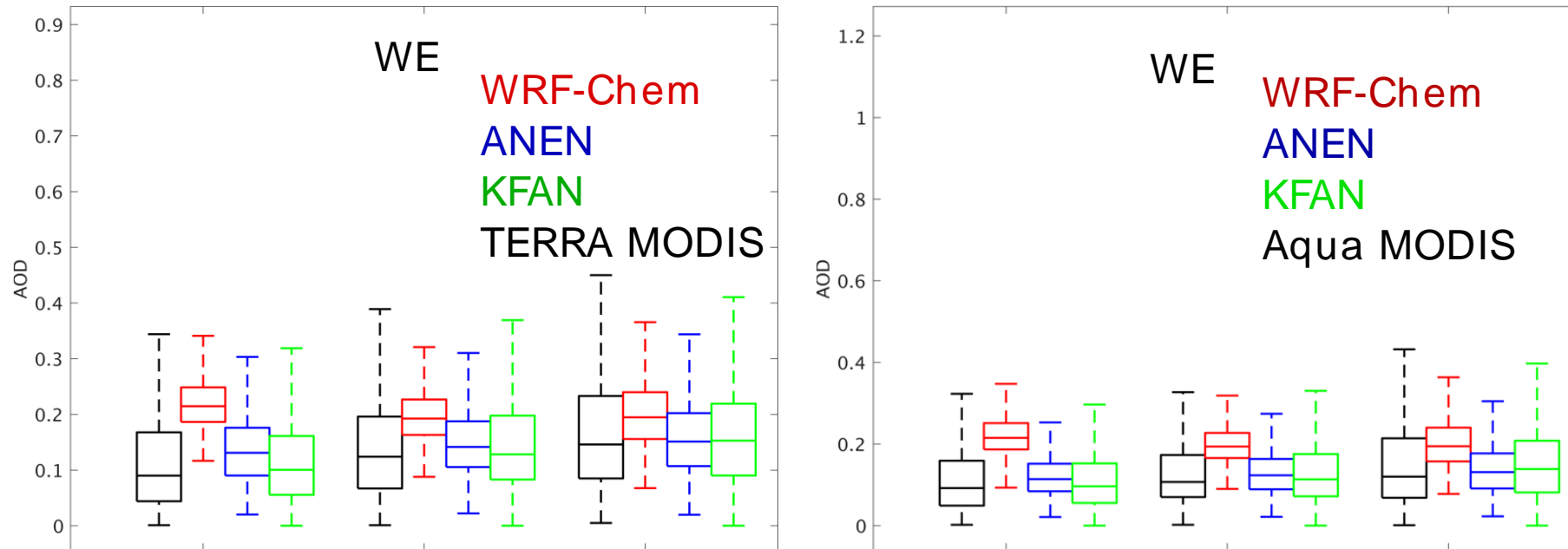
Correlation values of AOD and predictors . (a) AOD Vs Total Precipitable Water, (b) AOD Vs PM₁₀ , (c) AOD Vs PM_{2.5} , and (d) AOD Vs Horizontal windspeed at the surface collocated for Terra overpass time. (e) - (h) Similar to (a) - (d) but for Aqua over pass time

Analog Ensemble Forecasting : Results



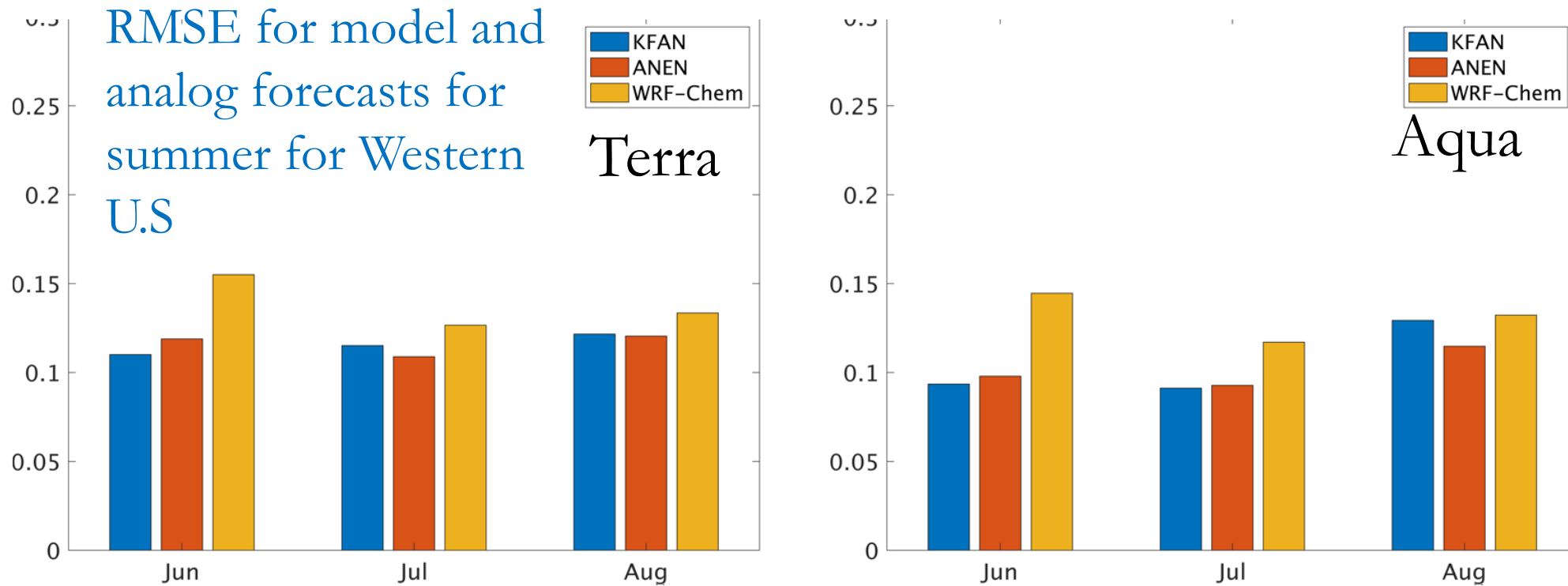
Mean bias (Model - Observation) for June - August, 2012 from WRF-Chem (a,d), ANEN (b,e), and KFAN (c,d) for points collocated with Terra (top) and Aqua (bottom) respectively.

Analog Ensemble Forecasting : Meteorology driven aerosols



Boxplot of AOD from MODIS, WRF-Chem , ANEN, and KFAN for June-August, 2012
Midline in the boxplot represents the median of AOD , top and bottom lines of the box represents 25th and 75th percentile, and top and bottom lines outside the box represents maximum and minimum values of AOD that are not outliers. The boxplots are shown for different EPA regions and outliers are not shown here.

Analog Ensemble Forecasting : Meteorology driven aerosols



Change in RMSE is $\sim 30\%$ for June and July. Smaller reduction for August is due to the higher relative humidity (RH) during this period across U.S and the model errors in accurately predicting the total precipitable water and effects of higher on RH on model AOD.

Take home messages

- Analog Ensemble forecasting with a combination of Kalman Filter provides improvement in AOD in seasons when AOD is mostly driven by wind and aerosol emissions.
- Despite the strong dependence on observations, analogs are heavily driven by the choice and quality of the predictors which is evident from the case of U.S East coast where during the month of August, errors in model predictions of precipitable water and wet scavenging prohibits improvement to AOD.

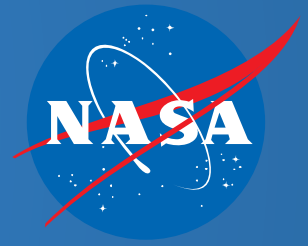
Aerosol global modeling

Implementing a process driven sea salt aerosol emission parameterization in GEOS

Whitecap fraction : Area of the ocean surface covered by active wave breaking (Stage A or active) and mature foam (stage b).



Sea salt emissions =
 f (*Whitecap fraction*)
 f (Size distribution of sea salt particles)



Whitecap models : Wind dependence

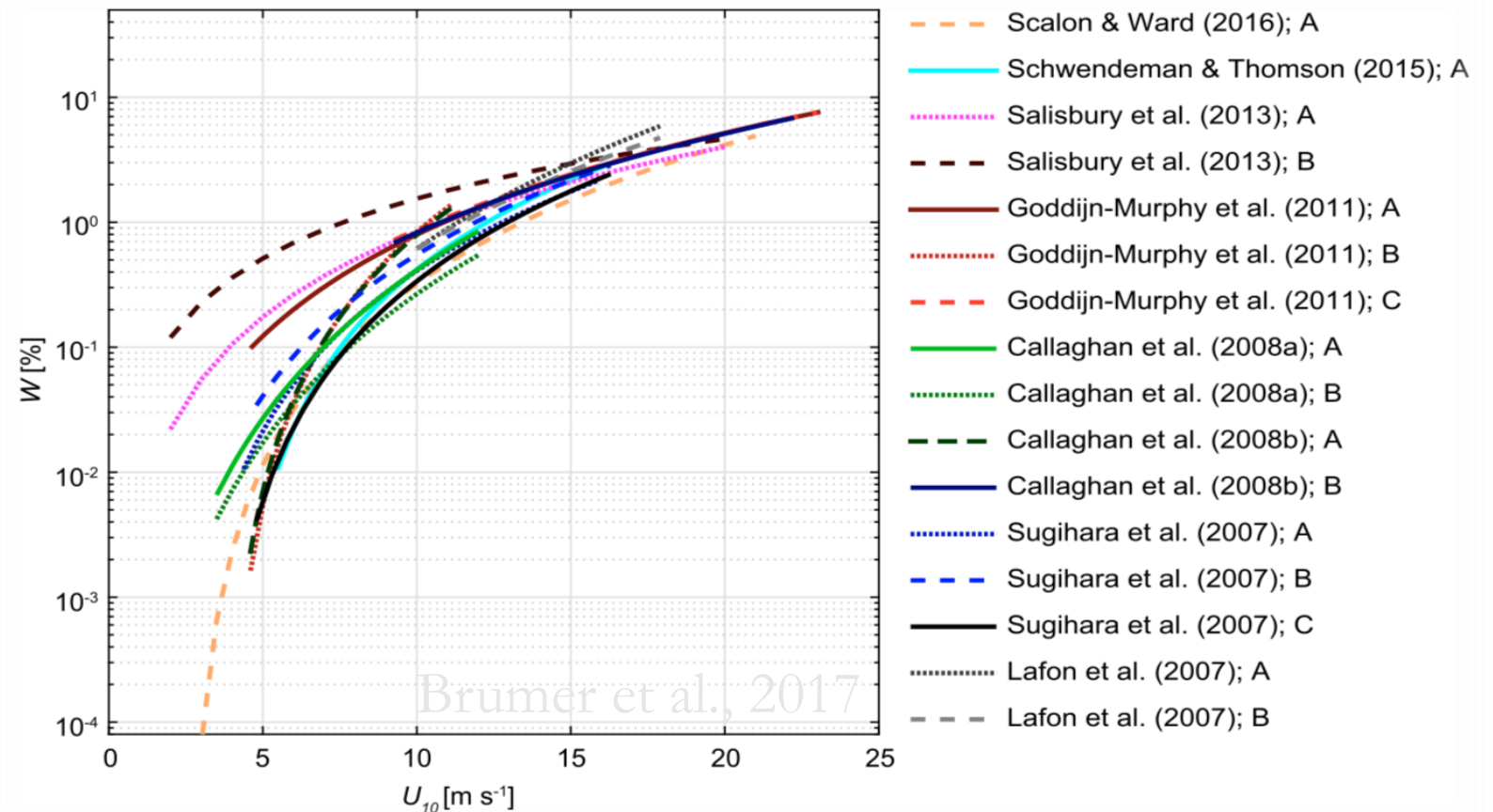
At a given wind speed, W variability is ~ 1 -2 orders of magnitude

(e.g. Monahan, 1971)

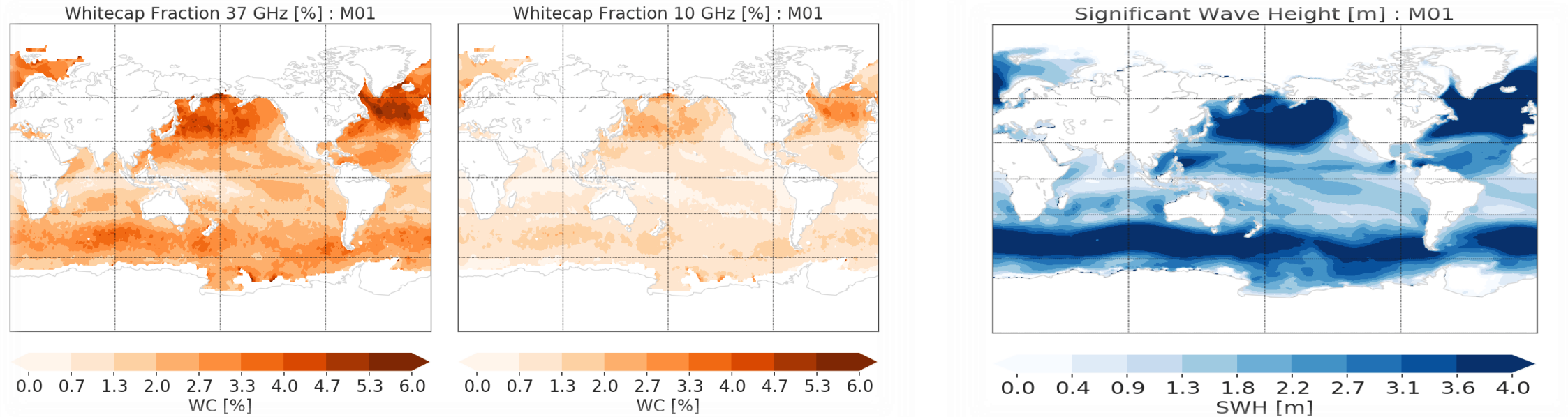
(e.g. Wu, 1988)

(Goddijn-Murphy et al. (2011))

(e.g. Callaghan et al., 2008)

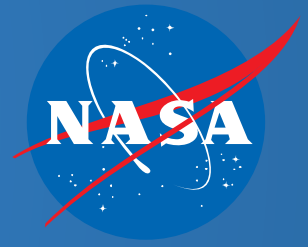


Observation constrained modeling



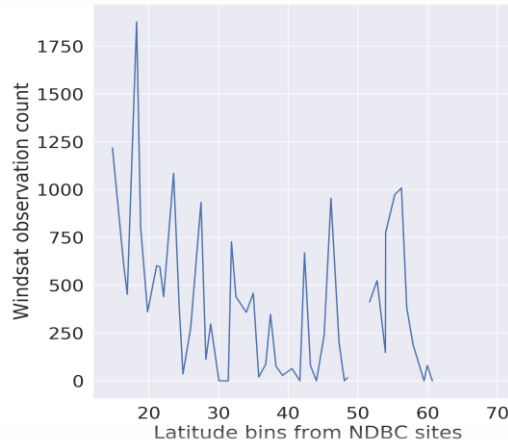
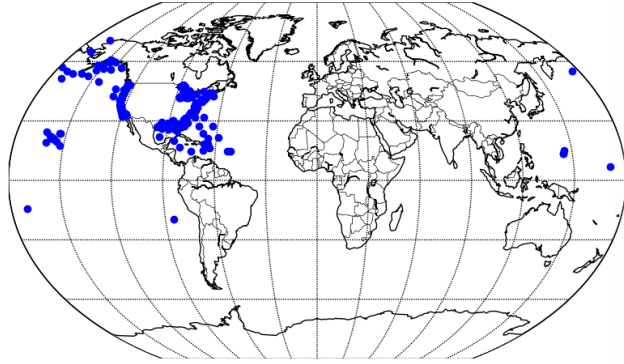
- Windsat W : $1^\circ \times 1^\circ$ multi –frequency retrievals [Anguelova et al., 2019]
- 10 GHz includes more active W and 37 GHz include fresh + mature (foam) W

- GEOS-UMWM
 - $0.5^\circ \times 0.5^\circ$ resolution runs for 2014 replayed to MERRA-2
 - Wind, sea-ice, air density input to UMWM from GEOS



W parameterization development

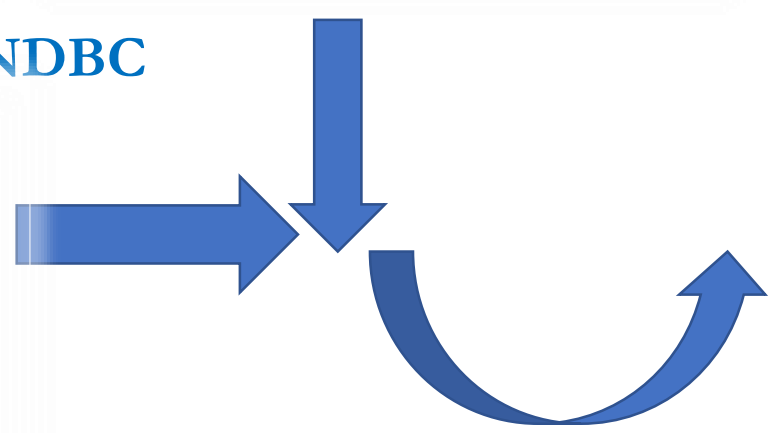
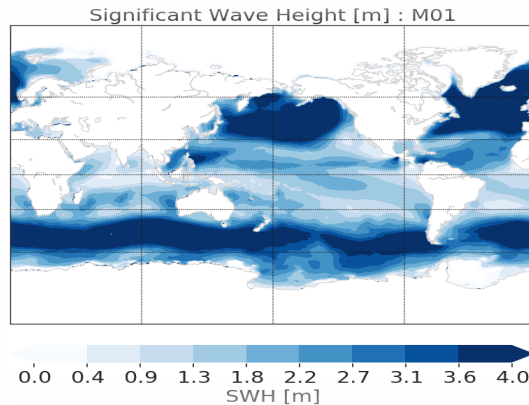
NDBC Wind/wave collocated stations



Model fitting

**Steepness, Re,
dissipation rate,
peak, air-sea
temperature
difference,
peak wave
velocity**

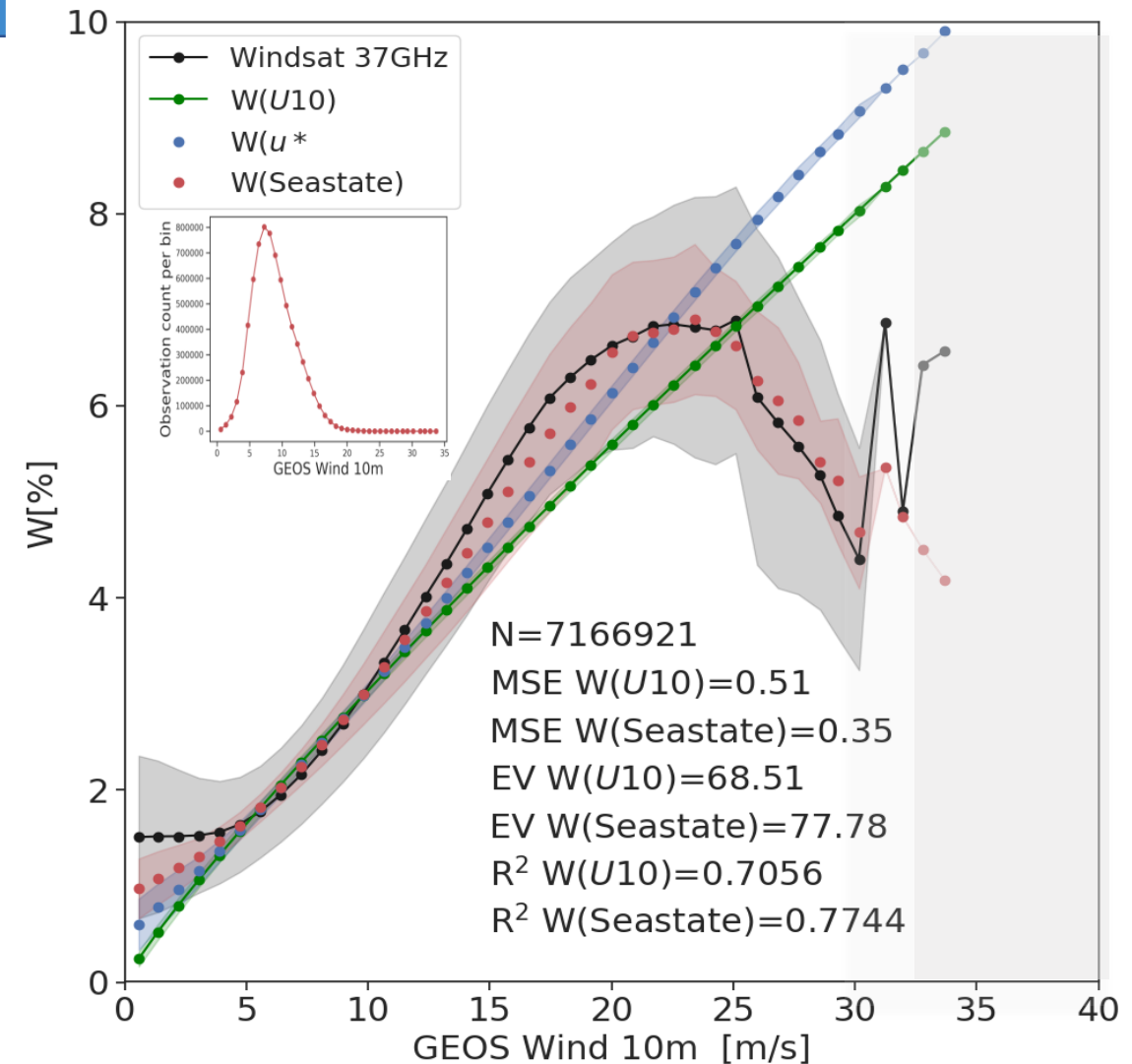
Sample GEOS-UMWM at NDBC



**Test against
independent
Windsat W**

Variability with wind speed

- Whitecap decreases for higher windspeed.
- In order to capture this behavior in models, additional terms based on wind stress were added to the Seastate W model for wind speed > 20 m/s.



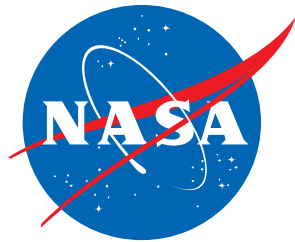
Aerosols as a part of the Earth System : What do we need from models and observations?

- Observations of vertical profiles of aerosol concentrations and number distribution.
- Point by point data uncertainties from satellite retrievals.
- Ensemble Forecasting of aerosols
- Use of aerosol products from satellites to improve model parameterizations in addition to direct assimilation of AOD [e.g. R.Kahn , 2020]

Acknowledgements

Thanks to NASA ACMAP , NASA ROSES, Modeling and Analysis Program, GMAO core, NCAR graduate visitor program for research funding and University of Arizona for computing resources.

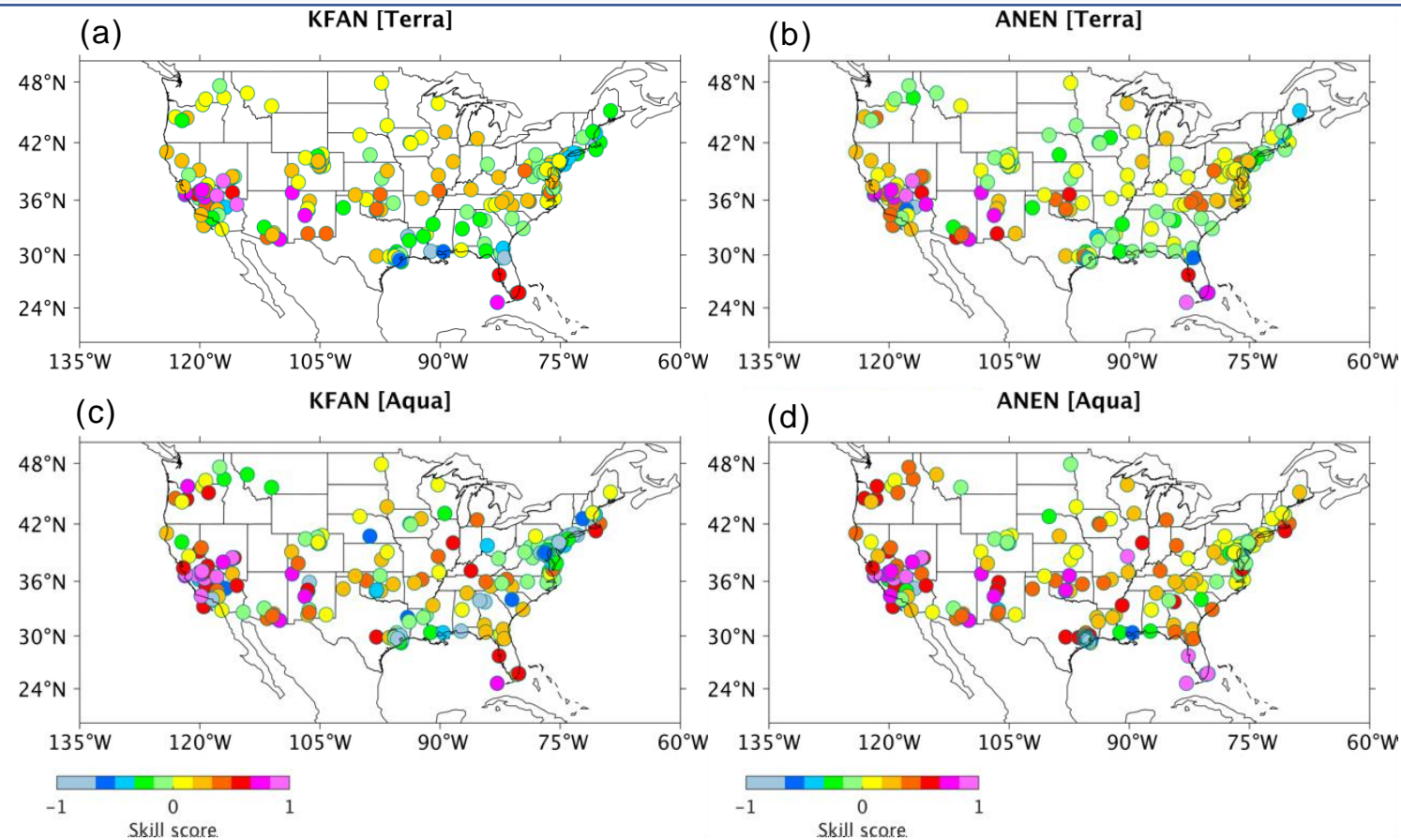
Thanks to Magdalena Anguelova, NRL for Windsat retrievals.



Thank you

Extra slides

Importance of meteorology for aerosol analogs



Skillscore for analog predictions of AOD for June-August, 2012. Skill score is calculated here using Mean Square Errors in comparison with MODIS Terra and Aqua retrievals. (a , b) KFAN and ANEN AOD estimated with reference to WRF-Chem AOD Mean Square Errors for Terra overpass. (c, d) similar to (a, b) for Aqua overpass.

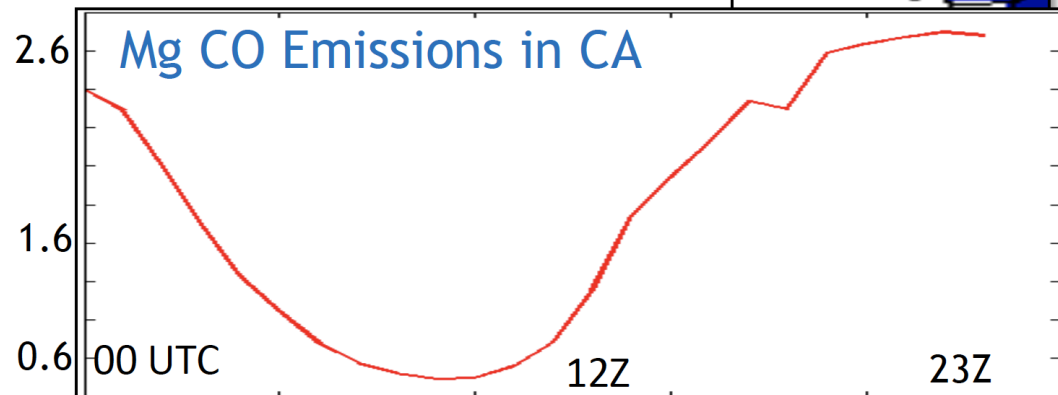
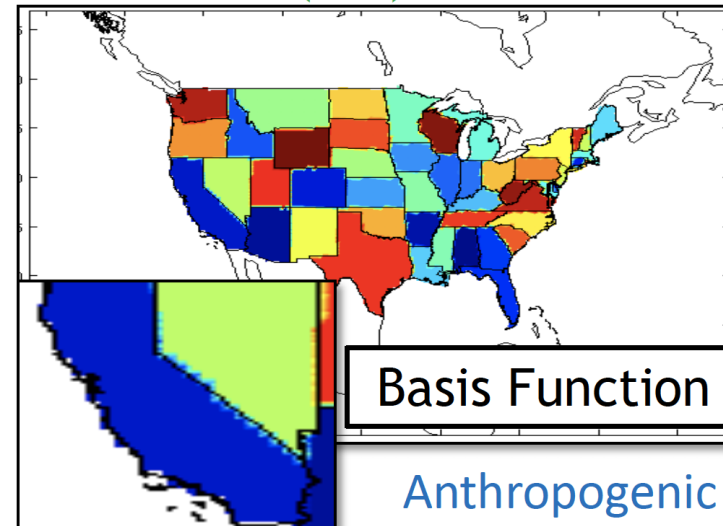
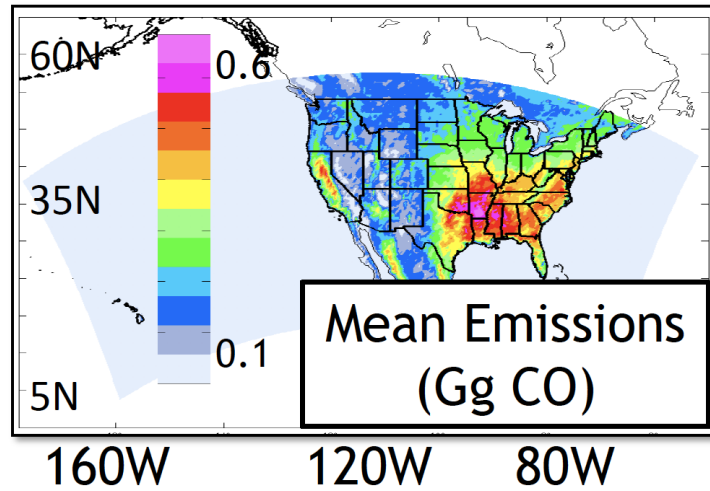
Negative skill scores in the East coast:

- Strong correlation with meteorology, in particular precipitable water

- Impact of wet scavenging processes.

Using $PM_{2.5}/CO$ ratios in WRF-Chem to improve $PM_{2.5}$ concentrations

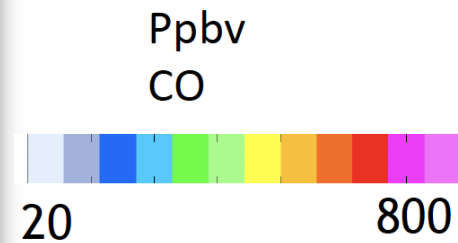
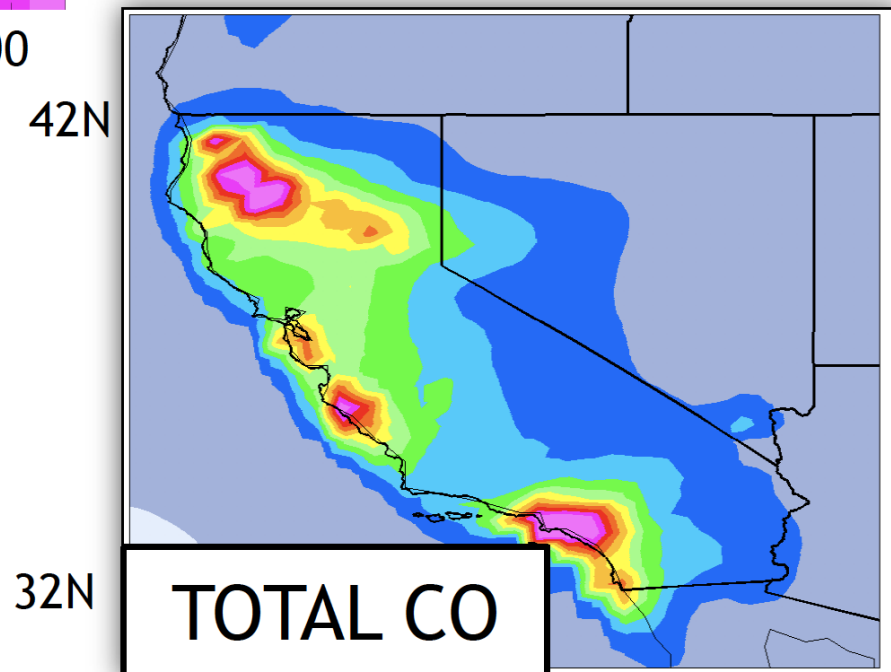
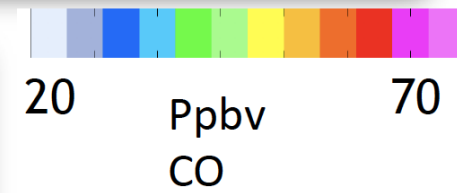
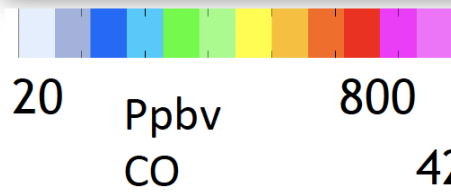
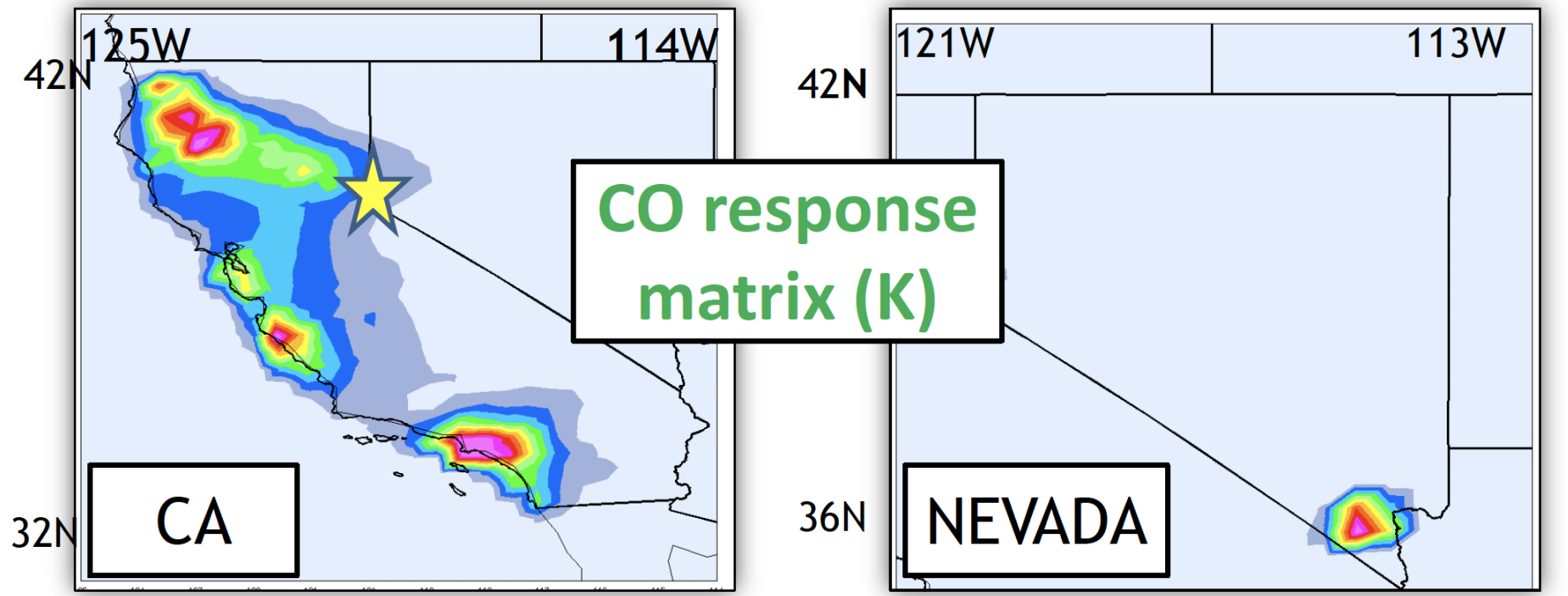
Bayesian Synthesis Inversion : Methodology illustration : **Prior CO emissions (Xa)**



Anthropogenic emissions : NEI 2005

Fire : FINN

Biogenic : Megan



A. Raman et al. IGAC
2014

Optimizing PM given CO observations

- Information from the observed relationship between CO and PM
- Modeled responses K from a set of source basis functions
- An estimate of the change in emissions from a source inversion of CO.

$$p(\mathbf{PM}|\mathbf{EPA CO}) \\ \propto p(\mathbf{PM}|\mathbf{CO}) \times p(\mathbf{CO}|\mathbf{CO}_{emis}) \times p(\mathbf{CO}_{emis}|\mathbf{EPA CO})$$

Mean and covariance of the estimate are given by

$$\hat{X} = X_a + \mathbf{S}_a (\mathbf{K})^T [(\mathbf{K}) \mathbf{S}_a (\mathbf{K})^T + \mathbf{S}_e]^{-1} [\mathbf{Y} - (\mathbf{K})X_a]$$

$$\hat{\mathbf{S}} = [(\mathbf{K})^T \mathbf{S}_e^{-1} (\mathbf{K}) + \mathbf{S}_a^{-1}]^{-1}$$

Where S_a is the prior error covariance and S_e is the observational error covariance.

Results

- Reductions in prior emissions in the New England mostly due to anthropogenic emissions.
- Huge increase in posterior emissions in Western US caused by biomass burning.

