

Illumination and Temperature on Rough Terrain: Fast Methods for Solving the Radiosity Equation

Samuel F. Potter (sfp@umiacs.umd.edu)^{†,\$}, Norbert Schörghofer^{\$}, and Erwan Mazarico* †: University of Maryland Department of Computer Science, \$: Planetary Science Institute, *: NASA Goddard Space Flight Center Acknowledgement: this research was supported by the NASA Planetary Science Division Research Program

Thermal modeling on rough surfaces

- Airless bodies have strong horizontal temperature gradients due to shadows cast by rough topography
- Lunar cold traps exist because of terrain shadowing and are defined by surface temperature
- Thermal models must incorporate shadows, but also long- and short-wavelength radiation between surface elements [3]
- Element-to-element radiation dominates runtime

Contribution

- We develop a fast algorithm for solving the equations governing scattering of long- and short-wavelength radiation
- Solve two discretized radiosity equations to compute temp.
- Offline, we precompute a compressed low-rank version of the discretized kernel matrix by compressing low-rank off-diagonal blocks using sparse SVDs
- Online, multiplication requires nearly O(N) time, where Nis the number of triangular elements
- The matrix only depends on the geometry of the planet so can be used for simulations spanning a long time

Physical model [5]

Energy balance on the surface:

$$\epsilon \sigma T^4 = (1 - \rho)(E + Q) + \epsilon Q_{\rm IR} \tag{1}$$

where:

 $\rho := albedo$

 $E := \text{incoming solar radiation (insolation)} [W m^{-2}]$

 $Q := \text{reflected sunlight } [\text{W m}^{-2}]$

 $Q_{\rm IR} := \text{thermal emission } [{\rm W \, m^{-2}}]$

T := temperature [K]

 $\epsilon := \text{emissivity}$

 $\sigma := \text{Stefan-Boltzmann constant} [\text{W}\,\text{m}^{-2}\,\text{K}^{-4}]$

Governing equations for scattering:

$$Q(\vec{x}) = \frac{1}{\pi} \int_{S} \rho(\vec{y}) (E(\vec{y}) + Q(\vec{y})) F(\vec{x}, \vec{y}) dA(\vec{y})$$
 (2)

$$Q_{\rm IR}(\vec{x}) = \frac{1}{\pi} \int_{S} \left(\epsilon \sigma T(\vec{y})^4 + (1 - \epsilon) Q_{\rm IR}(\vec{y}) \right) F(\vec{x}, \vec{y}) dA(\vec{y})$$
(3)

where:

S :=surface of the planet or crater

dA :=surface area element [m²]

$$F(\vec{x}, \vec{y}) := \frac{\left[\vec{n}(\vec{x}) \cdot (\vec{y} - \vec{x})\right]_{+} \left[\vec{n}(\vec{y}) \cdot (\vec{x} - \vec{y})\right]_{+}}{\pi ||\vec{x} - \vec{y}||^{4}} V(\vec{x}, \vec{y})$$

 $\vec{n} := \text{surface normal on } S$

 $[x]_+ := \max(0, x) = \text{positive part}$

1 if \vec{y} is visible from \vec{x} 0 otherwise

The radiosity method

• Equations (2) and (3) are radiosity integral equations:

$$B(\vec{x}) = E(\vec{x}) + \rho(\vec{x}) \int_{S} G(\vec{x}, \vec{y}) B(\vec{y}) dA(\vec{y})$$
 (4)

where:

 $B := \text{radiosity} [\text{W m}^{-2}]$

 $E := \text{self-emitted radiosity } [\text{W m}^{-2}]$

 $\rho := albedo$

F := geometric kernel

• Midpoint collocation discretization of (4) gives the system:

$$KB = (I - \rho F)B = E \tag{5}$$

$$\boldsymbol{F}_{ij} = \frac{\left[\boldsymbol{n}_i^{\top}(\boldsymbol{p}_j - \boldsymbol{p}_i)\right]_{+} \left[\boldsymbol{n}_j^{\top}(\boldsymbol{p}_i - \boldsymbol{p}_j)\right]_{+}}{\pi \|\boldsymbol{p}_i - \boldsymbol{p}_j\|^4} V_{ij} A_j$$
(6)

where \boldsymbol{F} is the **view factor matrix** and:

 $\boldsymbol{p}_i := ext{centroid of } i^{ ext{th}} ext{ triangle}$

 $oldsymbol{n}_i := ext{surface normal at } oldsymbol{p}_i$

 $A_i :=$ area of *i*th triangle

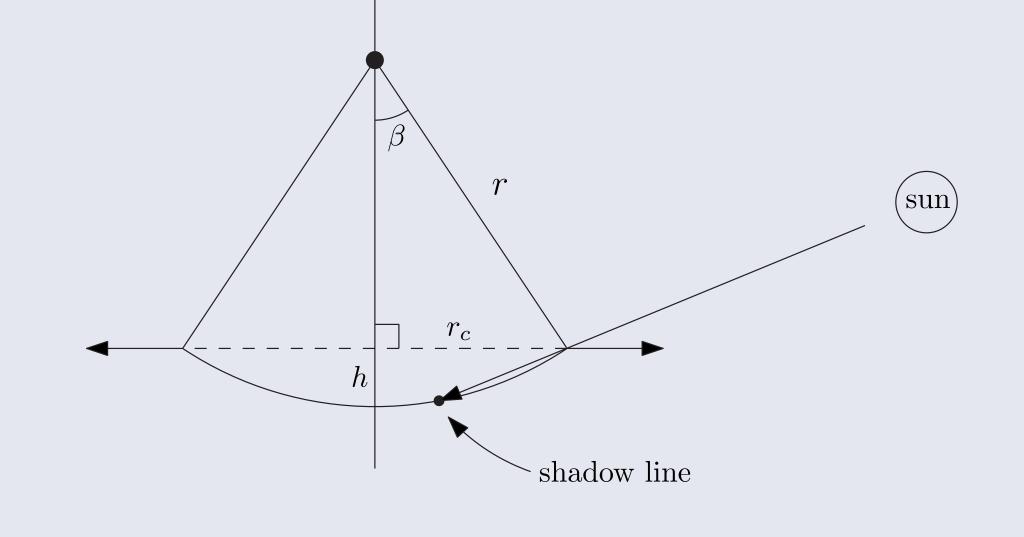
 $V_{ij}:=$ visibility between $oldsymbol{p}_i$ and $oldsymbol{p}_i$

Solving the discrete radiosity system:

- The system (5) can be solved in a small number of Neumann or Jacobi iterations (typically 2 to 5)
- Nearly perfectly conditioned since it's a discretized BIE
- Main challenge: fast multiplication by F

The Ingersoll crater test problem [2]

A crater formed from a spherical cap:



• The steady state temperature for the Ingersoll crater has an exact solution for Lambertian reflectance inside the shadowed portion of the crater:

$$T_{\text{Ingersoll}} = \frac{F_0 \sin(e_0)}{\sigma} \frac{1 - A}{1 - Af} \left(1 + \frac{A(1 - f)}{\epsilon} \right)^{1/4} \tag{7}$$

where σ , A, and ϵ are as before, and:

 $f := S_c/(4\pi r^2)$

 $S_c := \text{crater surface area } [\text{m}^2]$

 $F_0 := \text{solar constant } [\text{W m}^{-2}]$

 $e_0 :=$ solar elevation relative to horizon

We use this test problem to validate our numerical method

Low rank compression of form factor matrix \boldsymbol{F}

- Spatial partitioning: use quadtree or octree to recursively partition triangular elements
- Low-rank interactions: blocks of F that correspond to interactions between nonoverlapping cells in quadtree or octree are typically sparse with a dense low-rank subblock
- SVD compression: compute SVD to find dense subblock and compress it within a given tolerance ϵ
- Best low-rank approximation by SVD:

$$\min_{\operatorname{rank}(\boldsymbol{F}) \le k} \|\boldsymbol{F} - \boldsymbol{F}_k\|_2 = \sigma_{k+1} \tag{8}$$

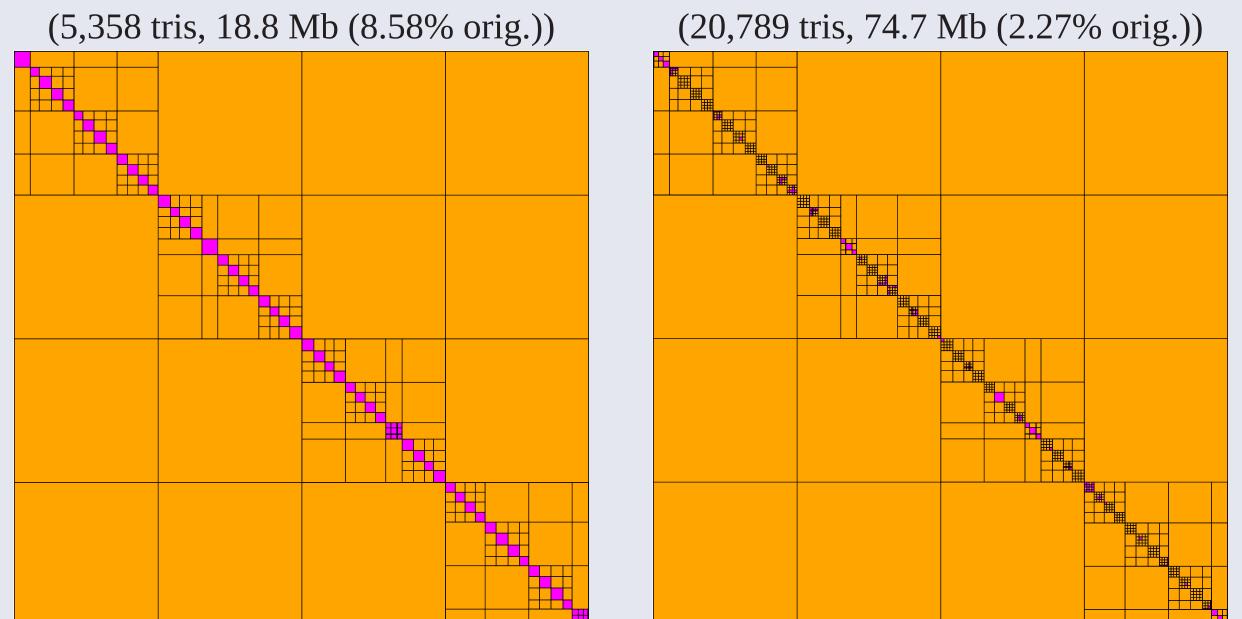
where k is a fixed rank, $m{F}_k = m{U}_k m{\Sigma}_k m{V}_k^ op$ is computed from the rank k truncated SVD of \boldsymbol{F} , and σ_i is the ith singular value of \boldsymbol{F}

• This approach is similar to the \mathcal{H} -matrix format [1]

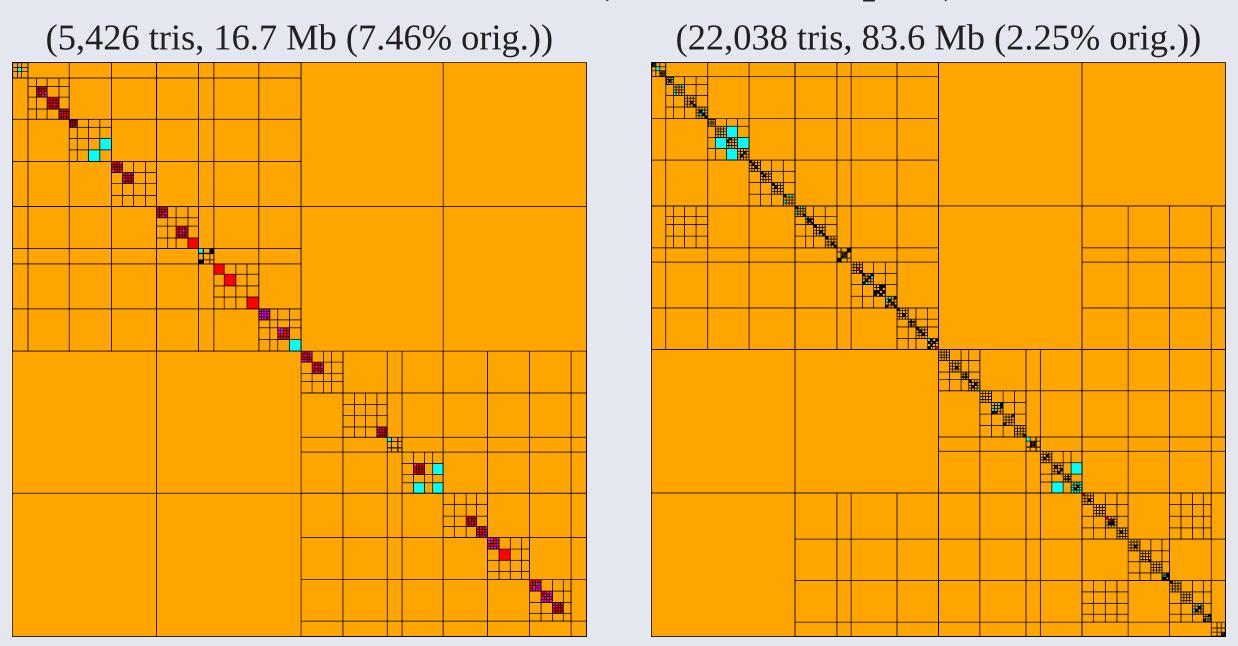
Examples of compressed F matrices

,■*: SVD block, ■: sparse block, ■: dense block, ■: zero block (*: the cyan SVD blocks can be thought of as "especially sparse" SVD blocks)

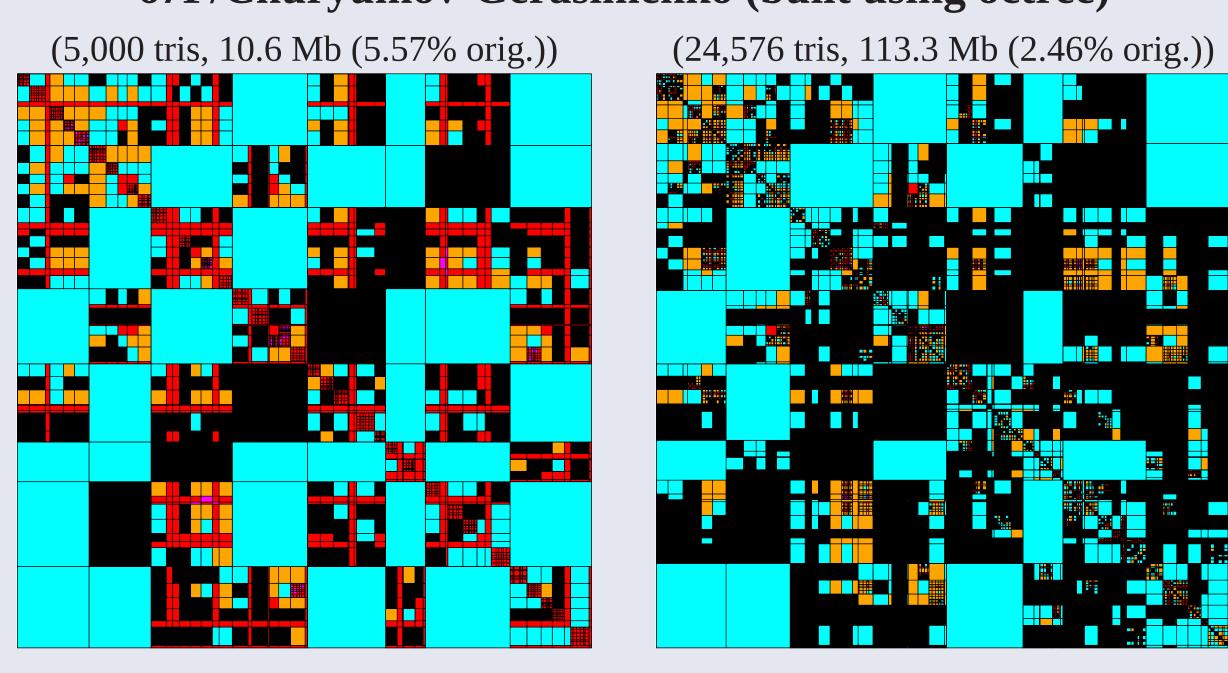
Ingersoll crater



Haworth crater (Lunar south pole)

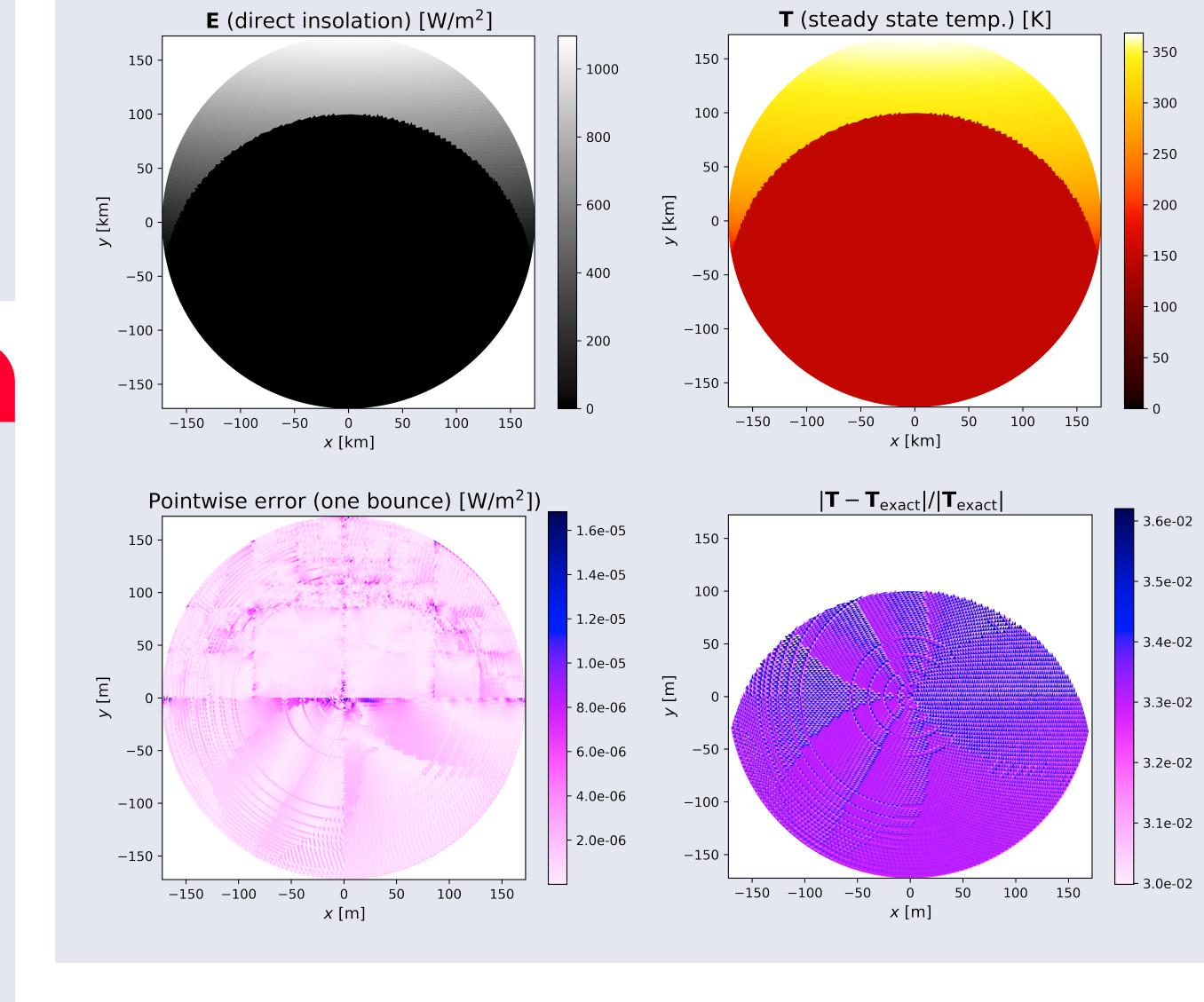


67P/Churyumov-Gerasimenko (built using octree)



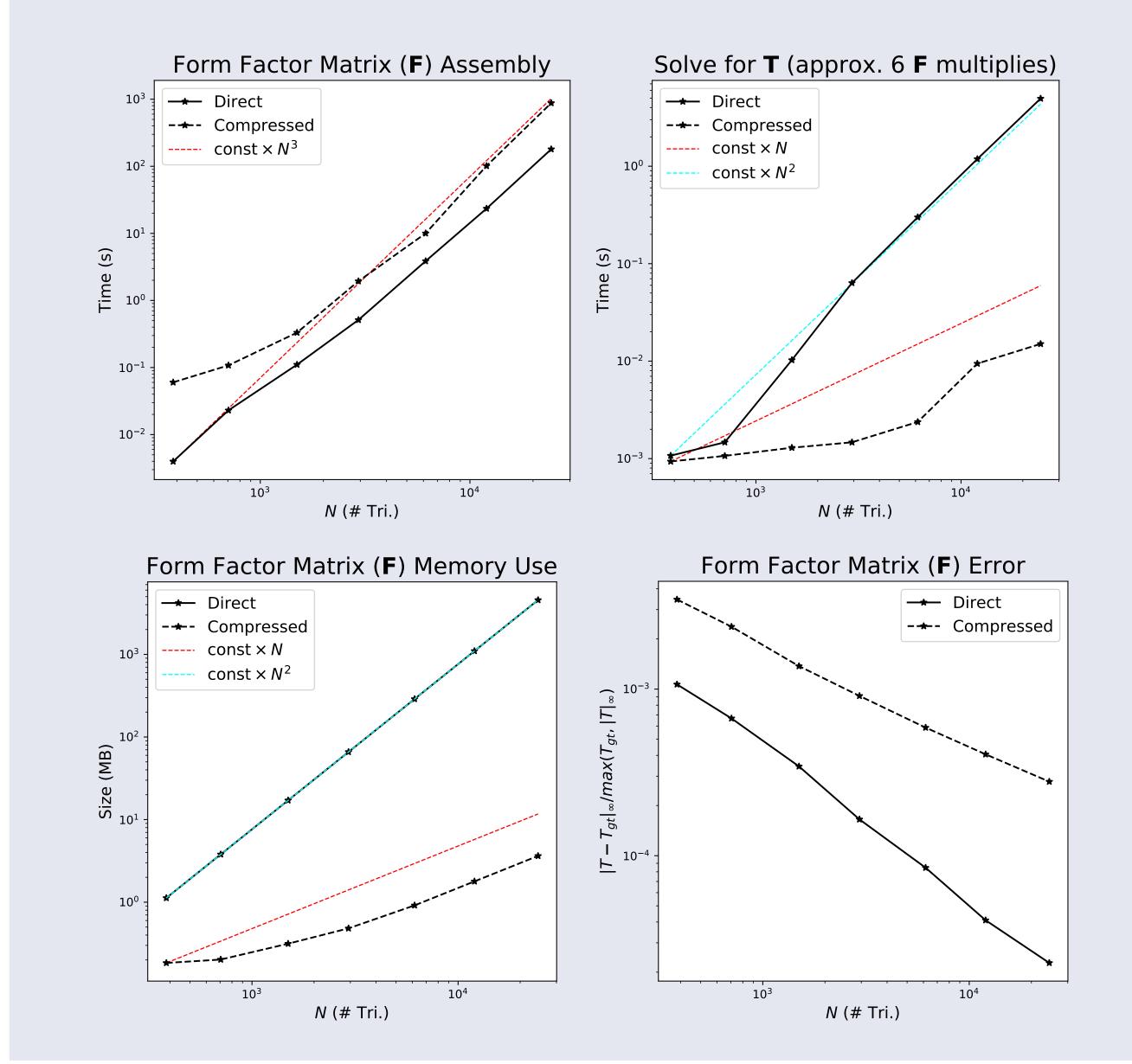
Ingersoll test problem: numerical results

- Only part of the crater is illuminated (*top-left*)
- The steady state temperature in the shadow is constant (*top-right*)
- Pointwise error after one bounce is below the $\epsilon = 10^{-4}$ tolerance (bottom-left)
- ullet Error between T and exact Ingersoll temperature is uniform (bottom-right)
- Errors are sensitive to the triangulation (bottom-left/right)



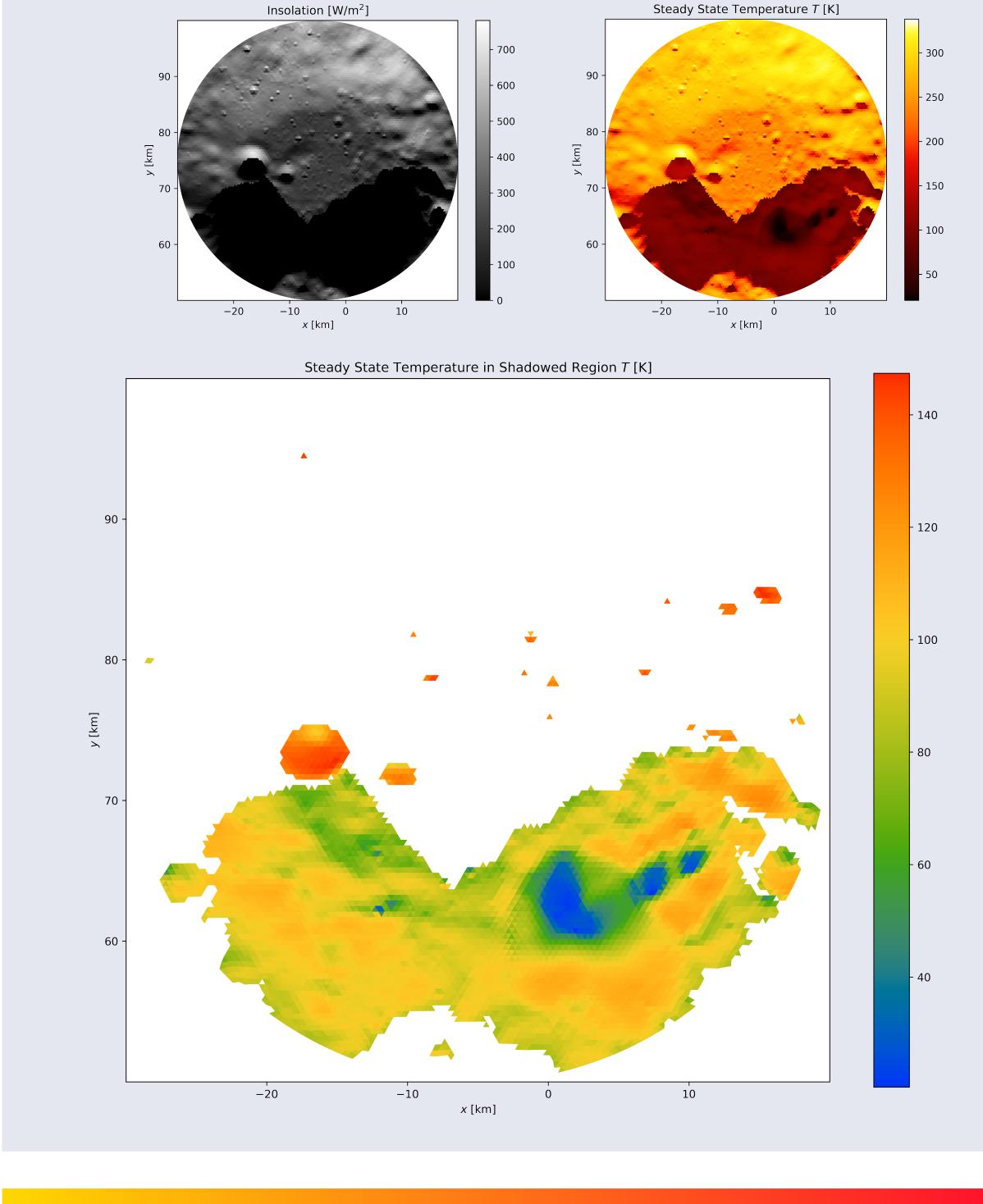
Ingersoll test problem: performance results

- We build the compressed F matrix with a tolerance of $\epsilon = 10^{-4}$ and a maxmium SVD rank of k = 60
- We compare the exact steady state temperature in the shadow region given by (7) with the numerical steady temperature vs. problem size (bottom-right)



Lunar south pole: Haworth crater test problem

- Lunar south pole terrain from DEM obtained by LOLA [6]
- High-quality triangle mesh constructed using Python version of distmesh [4]
- Example here is physically unrealistic because of steady state assumption, but solving the steady state system is the necessary step for solving the time-dependent problem efficiently
- The temperature is computed using the compressed form factor
- The temperature plotted below approximates the temperature using the exact, dense form factor matrix *pointwise*, **validating the** fast method



References

- [1] Wolfgang Hackbusch. Hierarchical matrices: algorithms and analysis, volume 49. Springer, 2015.
- [2] Andrew P Ingersoll, Tomas Svitek, and Bruce C Murray. Stability of polar frosts in spherical bowl-shaped craters on the Moon, Mercury, and Mars. *Icarus*, 100(1):40–47, 1992.
- [3] David A Paige, Matthew A Siegler, Jo Ann Zhang, Paul O Hayne, Emily J Foote, Kristen A Bennett, Ashwin R Vasavada, Benjamin T Greenhagen, John T Schofield, Daniel J McCleese, et al. Diviner lunar radiometer observations of cold traps in the moon's south polar region. science, 330(6003):479–482, 2010.
- [4] Per-Olof Persson and Gilbert Strang. A simple mesh generator in MATLAB. SIAM review, 46(2):329–345, 2004.
- [5] Norbert Schörghofer. Planetary Code Collection. https://github.com/nschorgh/Planetary-Code-Collection, 2018.
- [6] David E Smith, Maria T Zuber, Gregory A Neumann, Erwan Mazarico, Frank G Lemoine, James W Head III, Paul G Lucey, Oded Aharonson, Mark S Robinson, Xiaoli Sun, et al. Summary of the results from the lunar orbiter laser altimeter after seven years in lunar orbit. *Icarus*, 283:70–91, 2017.