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Fate of pesticides in field ditches: the TOXSWA simulation model

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ABSTRACT

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The TOXSWA model describes the fate of pesticides entering field ditches by spray drift, atmospheric deposition, surface runoff, drainage or leaching. It considers four processes: transport, transformation, sorption and volatilisation. A sample simulation showed that sorption to macrophytes can considerably reduce the concentration in the water phase. However, sorption to macrophytes leads to a slower transport and thus to longer residence times for the pesticide in the water layer. Sedimentation and resuspension of suspended solids were not considered, so that the model can only be applied for periods shorter than one month.

Keywords: environmental protection, pesticide concentration, surface water

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Preface

Predicting concentrations at which levels negative effects on aquatic organisms occur requires a methodology to predict exposure concentrations for these organisms. In 1991 the Directorate of Science and Transfer of Knowledge of the Dutch Ministry of Agriculture, Nature Management and Fisheries asked the DLO Winand Staring Centre (SC-DLO) to develop such a methodology. The project 'Modelling transport routes and fate of pesticides in soil and field ditches' started at SC-DLO at the end of 1991. The general administrative order on Environmental Admission Requirements for Pesticides ('AMvB-3a') (Ministry of Housing, Spatial Planning and Environment, 1995) mentions that the TOXSWA model is to be included in the future risk assessment procedure in the Netherlands.

In the first phase of this project P. Groenendijk and J.W.H. van der Kolk made an inventory of all relevant processes in field ditches. In August 1992 the project team was joined by P.I. Adriaanse, who focused specifically on the development of the TOXSWA simulation model (TOXic substances in Surface Waters), which describes the fate of pesticides in field ditches. On November 8, 1994 an international workshop was organised at the DLO Winand Staring Centre, Wageningen, the Netherlands, where aquatic fate modelling, the TOXSWA model and related pesticide regulations in the Netherlands were presented and discussed (Crum and Deneer (eds), 1995). About eighty specialists from eight different countries attended seven lectures and contributed to the discussion. By the end of 1995, W.H.J. Beltman joined the project team who wrote the user's manual for TOXSWA 1.0 and is in charge of defining 'standard scenarios' for the Netherlands.

From the end of 1991 up to mid-1995 the project work took place within the framework of project 7242 of the DLO programme 147 'Ecotoxicological Risks of Pesticides in Aquatic and Terrestrial Ecosystems', financed by the Dutch Ministry of Agriculture, Nature Management and Fisheries. From mid-1995 onwards the project was financed both by the SC-DLO (SEO-project 597) and the Ministry mentioned above (project 592 of DLO programme 276 'Emissions and Ecotoxicological Hazards of Pesticides').

Several persons outside the project team have contributed substantially to the development of the TOXSWA model. The author is most grateful to J.J.T.I. Boesten of the department Fate and Effects of Pesticides at SC-DLO, who provided many important ideas, reflections and criticisms all along the way. Special thanks are due to R.H. Aalderink of the department of Water Quality Management and Aquatic Ecology of the Wageningen Agricultural University; he provided ideas for the numerical solution and commented on it. J.A.P. Heesterbeek of the DLO Agricultural Mathematical Group contributed substantially to the mathematical background of Chapter 7. The 'walk through team', headed by M.J. van der Velden, the Software Quality Manager of SC-DLO, gave useful comments and ideas for improving the TOXSWA computer program.

Model development was critically monitored by a consultation group, consisting of representatives of the Department of Plant Protection (H. de Heer and A.C.P. van Montfort), of the Plant Protection Service (W.W.M. Brouwer), all from the Ministry of Agriculture, Nature Management and Fisheries, of the Toxicology Advisory Centre (J.B.H.J. Linders) and the Laboratory of Soil and Groundwater Research (A.M.A. van der Linden), both at the National Institute of Public Health and Environmental Protection, of the Institute for Inland Water Management and Waste Water Treatment (F. Wagemaker/P.C.M. van Noort) and of SC-DLO (P.E. Rijtema/P. Leeuwangh and P. Groenendijk).

The TOXSWA model development (and its verification) is described in this report. It treats the mathematical model, the numerical solution and the resulting computer model. Some first calculation results are presented as well.

Research on the TOXSWA model will continue at SC-DLO in the coming years. Sorption of a range of pesticides to three macrophyte species is being studied, as little information is available on this sorption parameter. In 1996 and 1997, the sensitivity of the model to input parameters and the initial conditions will be analysed. Model improvements will be implemented, initially in two areas: (i) inclusion of multiple or continuous pesticide applications to the water layer and (ii) inclusion of time series with varying water flow rates and water depths. Furthermore, standard scenarios for the Netherlands will be defined to facilitate the use of the model for risk assessments in the Dutch registration procedure. Model validation will start, using data sets from four experiments carried out in microcosms (1 m^3) and in outdoor ditches (60 m^3) by the DLO Winand Staring Centre (Crum and Brock, 1994; Aalderink and Crum, 1994; Crum et al., in prep.). In addition, data sets from experiments performed elsewhere will be assessed for their potential to validate the TOXSWA model. Field experiments in other EU countries are needed to assess whether the TOXSWA model is able to simulate fate in watercourses outside the Netherlands. The TOXSWA model is being assessed by the surface water fate group of the EU working party called FOCUS. Process descriptions need to be worked out in greater detail, including transformation rates and the effect of environmental parameters on them, or sorption to macrophytes. Finally, more experiments are needed to estimate the dispersion coefficient in various ditches under various conditions. Very few data exist for small watercourses, whereas the dispersion coefficient is very important for estimating the concentration curve. Results of these studies will be reported in future publications.

Summary

The TOXSWA model (TOXic substances in Surface WAters) has been developed to estimate exposure concentrations to pesticides of aquatic organisms in the field. The exposure concentrations are needed during the admission procedure of pesticides on the Dutch market, they are compared with laboratory toxicity data for selected standard organisms to evaluate the risks of agricultural use of pesticides for aquatic ecosystems.

The TOXSWA model describes the behaviour of pesticides in field ditches. The modelled field ditch system is two-dimensional and consists of two types of subsystem, water layer and sediment. In the water layer, concentrations vary only in horizontal direction, while in the sediment, concentrations vary both in horizontal and vertical directions. It can handle a variety of situations as regards hydrological conditions and entry routes of pesticides into surface water.

TOXSWA can be coupled to other models describing pesticides entering ditches via (i) drift or atmospheric deposition, (ii) surface runoff, or (iii) drainage or leaching through the soil.

In principle, TOXSWA can handle various hydrological conditions, so it simulates ditches with varying water depths h and rates of discharge Q .¹ The cross section of the ditch is trapezium-shaped; upward or downward seepage takes place through the ditch bottom and walls. Upward or downward seepage limits and enhances, respectively, the penetration depth of the substance into the sediment. No resuspension or sedimentation of suspended solids occurs in the water layer.

TOXSWA considers four processes: (i) transport, (ii) transformation, (iii) sorption and (iv) volatilisation. In the water layer, pesticides are transported by advection and dispersion, including transport of pesticides sorbed to suspended solids. In the sediment pesticides are transported by diffusion as well. The transformation rate covers the combined effects of hydrolysis, photolysis and biodegradation; metabolites are not considered. Sorption to suspended solids and to sediment is described using the non-linear Freundlich equation. Sorption to macrophytes is described using a linear isotherm. Pesticides are transported across the water-sediment interface by advection (upward or downward seepage) and by diffusion.

A pesticide mass balance is set up for an elemental volume in the water and in the sediment subsystem. The mass balances for the water and sediment subsystems account for incoming and outgoing mass fluxes. They result in two partial differential equations: the conservation equations for the water layer and the sediment.

¹ TOXSWA version 1.0 implements a constant water depth and discharge; these can be freely chosen.

Different types of input of pesticide into the water subsystem are possible: (i) distributed pulse input, e.g. spray drift or a momentary runoff, (ii) point-type pulse input, e.g. spillage of pesticide or a brief release from an individual drain, (iii) continuous point release, e.g. discharging tributary or (iv) continuous distributed release e.g. continuous release from many nearby drains. The water and sediment subsystems are coupled by assuming that the concentration in the liquid phase at the sediment surface equals the water phase concentration of the overlying water column. At the bottom of the sediment there is an inflow or outflow of water with pesticide.

The conservation equations for the water layer and the sediment are solved using a variable weight finite-difference method. This implies that spatial derivatives of the conservation equations may be approximated by e.g. a forward, central or backward difference. Temporal derivatives may be approximated in an implicit or explicit way, or as a weighted average of these two extremes. As processes in one subsystem may be more dynamic than in the other, different sizes of time steps are implemented for the water layer and the sediment.² The approach of the variable weight finite-difference method allows for an optimal computation time by adapting selected time and space steps to acceptable numerical dispersion. The numerical solution introduces numerical dispersion. In TOXSWA, the physical dispersion is corrected for this numerical dispersion. The approximations of the two conservation equations result in two matrix equations, which are solved for specified initial and boundary conditions. This allows the solution vector, consisting of the pesticide concentration, to be found. The TOXSWA model does not solve the conservation equations for the water and sediment subsystems simultaneously. The subsystems are linked by assuming that the concentration in the liquid phase calculated at time t can be used for calculating the flux at the sediment - water interface one time step later (at time $t+\Delta t$).

Verification is defined as the examination of the numerical technique in the computer model to ascertain that it truly represents the mathematical model and that there are no inherent numerical problems with obtaining a solution. This concept is worked out using the notions of convergence, stability and consistency. Convergence states that when the finite-difference grid is refined the truncation errors go to zero. Stability concerns the unstable growth or stable decay of errors in the arithmetic operations needed to solve the finite-difference equations. This condition is translated into a positivity condition, yielding conditions for the time and space steps resulting in stable and positive solutions. Consistency is the requirement that when the finite-difference grid is refined the truncation errors go to zero, but moreover that the finite-difference model approximates the partial differential equation desired and not some other partial differential equation. This is explained for the two conservation equations. It is also checked that the introduced mass of pesticide (100%) can be traced at every moment, so that the mass balances tally and that TOXSWA conserves well the introduced mass in the system. Furthermore, a model simulation is found to correspond very well with an analytical solution for a water subsystem, which

² In TOXSWA version 1.0 the numerical weight factors are fixed on values yielding an explicit central difference calculation scheme. The time steps for the water layer and the sediment also need to be similar.

responds to a Dirac delta function-type input. Correspondence between a model simulation and an analytical solution for the sediment subsystem responding to the same type of input was also excellent.

To illustrate the possibilities of the TOXSWA model, a computation was carried out for the insecticide chlorpyrifos. In total 2 g chlorpyrifos was deposited on a ditch of 200 m long and 3.65 m wide at its water surface. The ditch contained suspended solids and macrophytes and had a water flow rate of 100 m/d. Chlorpyrifos sorbs strongly to organic matter and to macrophytes. Simulation was run for four days. Initially, 53% of the 2 g chlorpyrifos was dissolved in the water phase, 42% was adsorbed to the macrophytes and 4% to the suspended solids. After four days, most of the pesticide had flowed out of the ditch, but of the remaining mass (0.1 g) 72% had penetrated into the sediment, 15% was dissolved in the water phase, 12% was adsorbed to the macrophytes and 1.3% to the suspended solids. The mass balance for the water layer showed that, besides outflow, volatilisation, penetration in sediment and transformation were the most important factors (in order of priority) contributing to the pesticide decrease in the water layer.

The TOXSWA model calculates the average concentration to which aquatic organisms are exposed at 3, 21 and 28 days after application, as well as immediately after application (at 0 d). They represent the average of the concentration course for pesticide dissolved in the water phase with time. The exposure concentration at time $t = 0$ d corresponds to the total pesticide concentration, immediately after application. The exposure concentration has been defined as the concentration at that position in the ditch where the longest exposure duration is expected, i.e. at the downstream end of the section of the ditch where the pesticide input took place.

The TOXSWA model checks whether the mass balances tally during the calculations. Results for the example simulation for chlorpyrifos showed that after four days the missing quantity in the mass balances for both the water layer and the sediment was less than 0.005% of the initial mass (plus, for the sediment, incoming mass of the water layer). So, this shows that TOXSWA 1.0 conserves well the mass of pesticide applied.

In the TOXSWA model the assumptions are made that the pesticide mixes instantaneously across the cross section and that sorption equilibrium is also instantaneous. Measurements for chlorpyrifos performed during earlier experiments of SC-DLO showed, however, that both assumptions do not correspond with reality. This may result in pesticide concentrations simulated in the water phase being too low, especially during the first one or two days.

The model has been developed to simulate periods of up to about one month; moreover, as no resuspension and sedimentation of suspended solids is included, this version of the model is not suitable for calculating long term exposure concentrations in the sediment or accumulation of pesticide in the sediment.

1 Introduction

The Dutch Pesticide Act of 1975 requires an evaluation of the hazards of pesticides, with regard to public health as well as the environment, before pesticides can be registered. In 1991, the Dutch Government began to implement the Multi-Year Crop Protection Plan, which aims to produce a considerable reduction in the dependence on and the use of pesticides in the Netherlands, as well as to reduce the emission of pesticides to the environment. Stricter rules for admission of pesticides are among the measures mentioned in this plan to alleviate side-effects of pesticides on the environment. In 1995 the Dutch Government issued a so-called general administrative order ('AMvB-3a') in which stricter rules concerning leaching to groundwater, persistence in the soil and toxicity for aquatic organisms were specified. Risk evaluation for aquatic organisms is based upon comparison of the estimated exposure concentration in the field with laboratory toxicity data.

At present, a relatively simple model, called SLOOT.BOX (Linders et al, 1990) is used to predict exposure concentrations in field ditches. This model is now being used by the Dutch Board for the Authorization of Pesticides to predict the exposure concentration for short term effects on aquatic organisms. Differentiation of admission of pesticides according to regional characteristics, or realistic prediction of chronic exposure of aquatic organisms (i.e. up to 28 days) requires a more detailed model. This should take all relevant processes into account and should consider not only spray drift deposition but also other entry routes of pesticides to ditches (e.g. runoff, drainage, leaching and atmospheric deposition) (Fig. 1).

The objective for the development of the TOXSWA model has been to provide the Board for the Authorization of Pesticides with an improved tool for the estimation of pesticide concentrations in field ditches. Acute (up to 4 days) as well as chronic (up to 28 days) exposure of aquatic organisms need to be predicted in a realistic way if one is to account for regional differences.

The present report provides a detailed description of the TOXSWA model development with a limited verification of the implemented computer program. The TOXSWA model includes all processes relevant for a description of the fate of pesticides in field ditches; it considers various entry routes of pesticides like spray drift deposition, seepage or point-type inputs. This makes the TOXSWA model a suitable tool for producing realistic estimations of the exposure concentrations, differentiated in time and space, for aquatic organisms in field ditches.

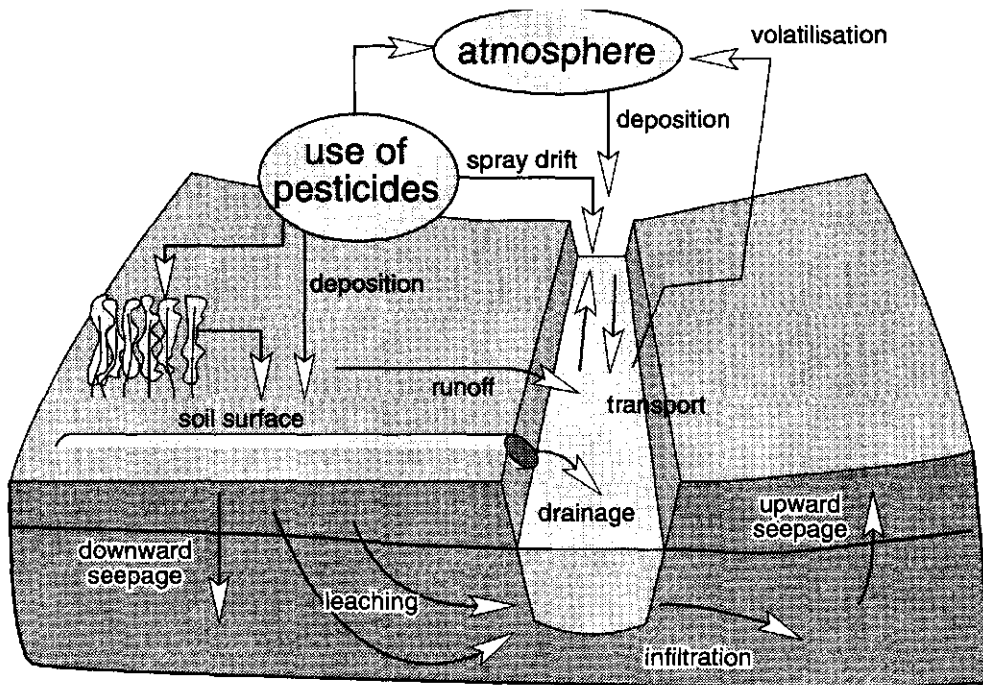


Fig. 1 Possible entry routes of pesticides to a field ditch

In Chapter 2 the modelled system and model assumptions are described, while Chapter 3 draws up mass balances for water layer and sediment. Chapter 4 discusses the processes that appear in TOXSWA and Chapter 5 specifies boundary and initial conditions. In Chapter 6 the differential equations expressing the mass balances are solved with a variable weight finite-difference method, while Chapter 7 verifies the implementation of the resulting computer program. Chapter 8 gives some early computation results and discussions. Finally the conclusions, discussion and recommendations are formulated in Chapter 9.

2 System description

2.1 Introduction and system definition

The behaviour of pesticides in field ditches, including their sediments, is described in this section. Pesticides can end up in field ditches because of spray drift, runoff from neighbouring field lots or by releases. In a ditch with a draining function, pesticides can also enter the ditch water by upward seepage through the ditch bottom. In the water, pesticides sorb to dissolved particles, to suspended solids and to macrophytes. In the sediment, pesticides sorb to the solid bottom material. In the course of time, pesticides degrade.

The field ditch system is divided into two subsystems, water layer and sediment. Mass balances are drawn up for both subsystems. These result in two partial differential equations, which are coupled to each other and in principle need to be solved simultaneously. These equations are solved with the aid of numerical calculation techniques and for specified initial and boundary conditions. These solutions form the basis for the resulting computer program.

The field ditch system is characterised in the following way (Fig. 2).

- The cross section of the ditch is trapezium-shaped and the wetted perimeter separates the water layer from the neighbouring soil.
- Discharge and water level in the ditch vary in time and space.
- Upward and downward seepage occur through the bottom and walls of the ditch.

The coordinate system is defined in Figure 3.

- x axis: positive in the most frequent direction of flow in the ditch;
- z axis: positive with depth, zero at the sediment surface;
- y axis: perpendicular to the x axis and the z axis.

Other characteristics of the system follow below.

- There is a concentration gradient of the substance at issue in the x direction of the water layer; there are no concentration gradients in the y direction or the z direction.

This means that the **assumption** has been made that the substance is ideally mixed, vertically and horizontally, in the water layer.

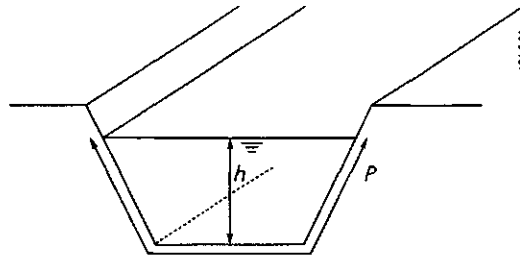


Fig. 2 Outline of model system

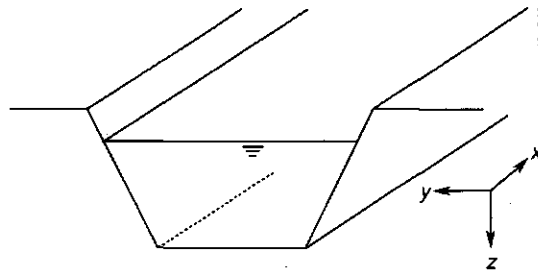


Fig. 3 Definition of coordinate system

- There is a concentration gradient of the substance in the z direction of the sediment as well as in the x direction. This implies that there is no concentration gradient in the y direction.

This means that it has been **assumed** that the substance is ideally mixed horizontally in the sediment.

- Transport of the substance occurs in the x direction in the water layer and in the z direction in the sediment.

It has been **assumed** that transport of the substance in the x direction in the sediment is so small that it can be neglected, so this flux is assumed to be zero.

(Hence, the concentration gradient in the x direction in the sediment originates in the concentration gradient in the x direction of the overlying water.)

The system thus described has been divided into two subsystems to model the substance behaviour. These are the water layer and the sediment, which together form the entire system. This has been done for the following reasons.

1. The substance behaviour in both subsystems can be described by a one-dimensional model. In the water layer this is the x direction, the direction of flow. In the sediment this is the z direction, i.e. downwards.
2. Partly different processes play a role in both subsystems.

The subsystems communicate with each other by exchanging water and substance through the wetted perimeter of the ditch.

2.2 Processes and suspended solids behaviour

In the water subsystem, the substance is subject to the following processes.

- Advection in the x direction.
- Dispersion in the x direction.
- Exchange with the atmosphere, through a diffusive flux across the water-air interface.
- Exchange with the sediment, through a combination of an advective and a diffusive flux across the wetted perimeter.
- Transformation, described as overall transformation without distinction between dissolved substance and substance sorbed to suspended solids or to macrophytes.
It has been **assumed** that the transformation rate of the substance dissolved in the water phase equals the transformation rate of the substance sorbed to suspended solids or to macrophytes.
- Sorption to suspended solids and to macrophytes.
It has been **assumed** that sorption to dissolved particles can be neglected in describing the fate of pesticides, so this process has not been included in the model.

Diffusion in the x direction has not been included, as dispersion prevails; the dispersion process exists even in stagnant waters, due to e.g. wind effects and inversion of the water column caused by air temperature changes between day and night.

In the sediment subsystem the substance behaviour has been described by the following processes.

- Advection in the z direction.
- Dispersion in the z direction.
- Diffusion in the z direction.
- Exchange with the water layer, being a combination of an advective and a dispersive flux across the wetted perimeter.
- Transformation, described as overall transformation without distinction between substance dissolved and substance sorbed.
The **assumption** is that the transformation rate of the substance dissolved in pore water equals that of the substance sorbed to the solid phase of the sediment.
- Sorption to the solid phase of the sediment.

Figure 4 shows a diagram of these processes.

The system description assumes the suspended solids concentration to be constant. The suspended solids flow along with the water. This simplification, i.e. a constant concentration of suspended solids, implies that when additional water is supplied to the ditch (e.g. by drains or trenches, by seepage or by precipitation) the mass of suspended solids will increase as well. These so to say artificially added suspended solids do not contain pesticides at the moment they are added. But immediately afterwards, the total mass of pesticides in the water layer will be redistributed and the pesticides will also sorb to the additional suspended solids. Consequently, the

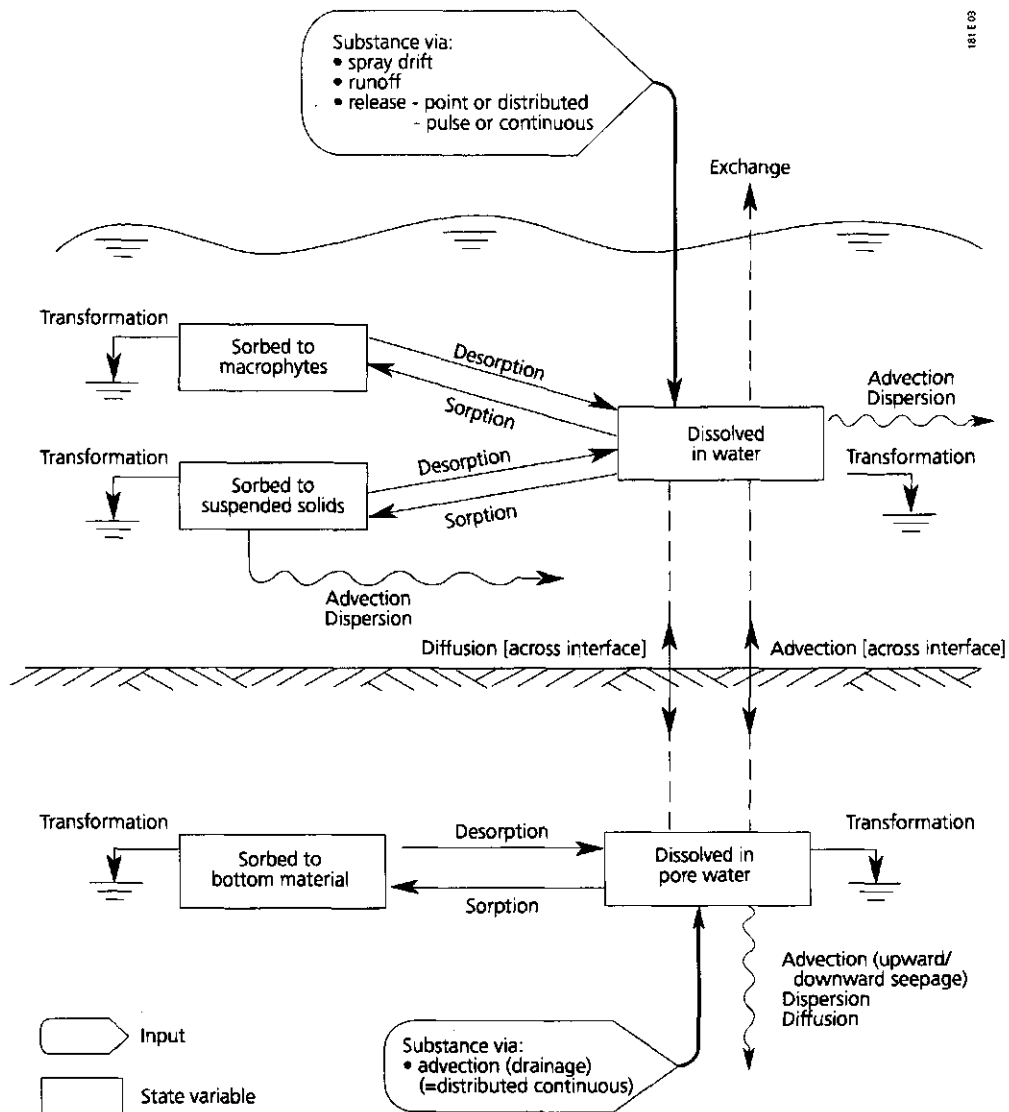


Fig. 4 Diagram of modelled processes

concentration of pesticides dissolved in the water phase will decrease (that is, not only by dilution, but also by sorption to the additional suspended solids).

The **assumption** is therefore that the concentration of suspended solids in the water layer is constant.

Pesticides sorbed to suspended solids are subject to the same advection and dispersion as pesticides dissolved in the water phase.

It has also been **assumed** that pesticides sorbed to suspended solids undergo the same advection and dispersion as dissolved pesticides.

No sedimentation or resuspension of suspended solids occurs. Pesticides generally sorb strongly to suspended solids. Therefore, neglecting the sedimentation flux is only acceptable if this is indeed negligibly small. This is the case e.g. over relatively short periods, or over somewhat longer periods in clear, moderately eutrophic ditches. This means that the development of turbid hypertrophic ditches for a succession of

years will not be described realistically, because an essential process is disregarded. The same argumentation holds for the resuspension flux. The description is only realistic for those ditches in which the flow velocity is so low and the wind influence so small that these provoke a negligible resuspension of suspended solids. This approach is acceptable for sheltered, very slowly flowing field ditches, but whether it also is acceptable for large, full ditches in open polders or for inclined ditches is less clear.

Hence, the **assumption** has been made that sedimentation and resuspension of suspended solids are negligible.

3 Mass balances for water layer and sediment

3.1 Water layer

The water layer is assumed to be ideally mixed laterally and vertically and the substance shows a concentration gradient only in the direction of flow. This means that a one-dimensional mass balance can be drawn up for the elemental volume $A\Delta x$, in which A represents the wetted surface perpendicular to the direction of flow.

The mass balance for the substance in the water layer now becomes as follows (Fig. 5).

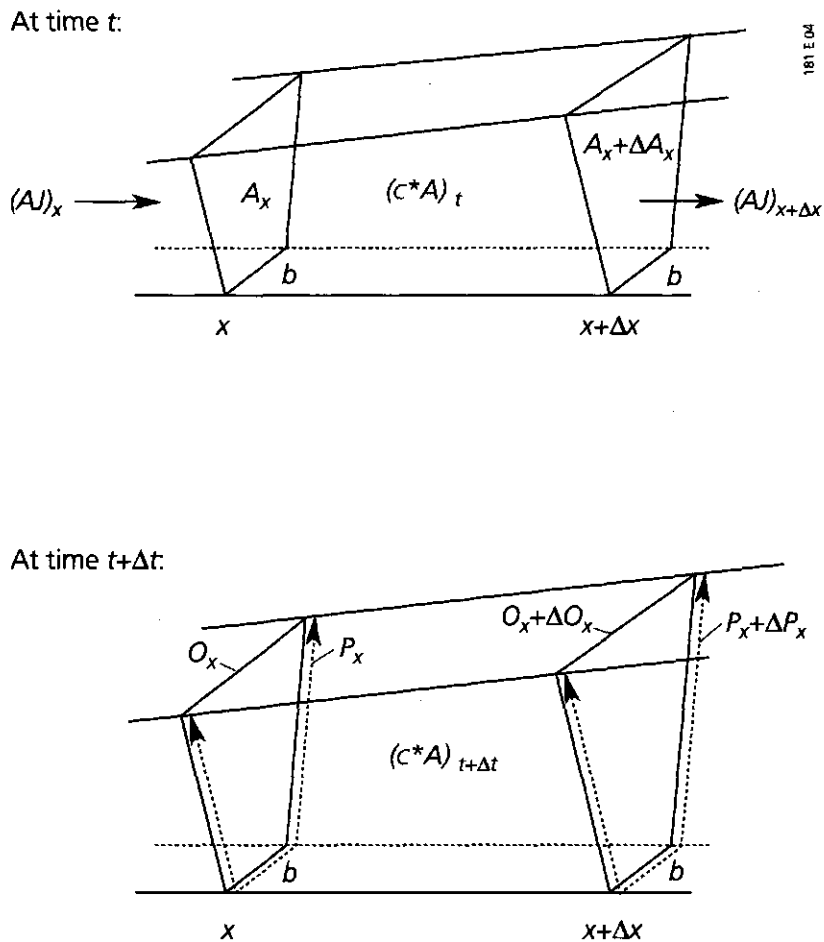


Fig. 5 Mass balance at times t and $t + \Delta t$

- * (c^*A) is the mass above a unit of length, here Δx
- * A and h at $t+\Delta t$ are changed compared to those at time t
- * P_x en O_x at $t+\Delta t$ are also changed compared to those at time t

In the above,

A_x = cross sectional area of flow at location x (L^2) (time, space)³

b = width of ditch bottom (L)

c^* = mass concentration of substance in water layer (this includes substance sorbed to suspended solids and to macrophytes ($M.L^{-3}$) (time, space)

h = water level above ditch bottom (L) (time, space)

J = areic⁴ mass flux of substance in water layer by advection and dispersion ($M.L^{-2}.T^{-1}$) (time, space)

O_x = width of water surface at location x (L) (time, space)

P_x = wetted perimeter at location x (L) (time, space).

The width of the water surface O_x equals:

$$O_x = b + 2hs_1 \quad (3.1)$$

with

s_1 = side slope, horizontal/vertical (1).

The wetted perimeter is described as:

$$P_x = b + 2h\sqrt{s_1^2+1} \quad (3.2)$$

The change in the mass of the substance (i.e. the term c^*A) over length Δx is important for the mass balance of this elemental volume; this change is due to mass transport through the cross sectional area (i.e. the term AJ) at x and $x+\Delta x$.

The mass balance reads:

accumulation = input - output.

Notated in differentials:

$$\frac{[(c^*A)_{t+\Delta t} - (c^*A)_t]\Delta x}{\Delta t} = \frac{[(AJ)_x - (AJ)_{x+\Delta x}]\Delta x}{\Delta x} \Leftrightarrow \quad (3.3)$$

³ The dimension of the symbol introduced is given between the first pair of brackets. L stands for length, T for time, M for mass, N for mole and θ for temperature (Schurer and Rigg, 1980). The quantities on which the introduced variable is dependent are given between the second pair of brackets.

⁴ Areic means that it is divided by the area concerned.

The limit notation for Δx and Δt reaching zero:

$$\frac{\partial(c^*A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} \quad (3.4)$$

Accounting for sources and sink terms, as well as the exchange of mass across the boundaries of the subsystem, the mass balance reads as follows.

In words:

accumulation = input - output + sources - sinks \pm exchange.

In differentials:

$$\begin{aligned} [(c^*A)_{t+\Delta t} - (c^*A)_t] \Delta x &= \Delta t [(AJ)_x - (AJ)_{x+\Delta x}] && \text{flux} \\ &- \Delta t \cdot k [(c^*A)_x + \frac{1}{2} \Delta(c^*A)_x] \Delta x && \text{transformation} \\ &+ \Delta t \cdot J_{wa} [O_x + \frac{1}{2} \Delta O_x] \Delta x && \text{removal to atmosphere} \\ &- \Delta t \cdot J_{wb} [P_x + \frac{1}{2} \Delta P_x] \Delta x && \text{exchange with sediment} \end{aligned}$$

$$\begin{aligned} \Leftrightarrow \frac{[(c^*A)_{t+\Delta t} - (c^*A)_t]}{\Delta t} &= \frac{[(AJ)_x - (AJ)_{x+\Delta x}]}{\Delta x} \\ &- k [(c^*A)_x + \frac{1}{2} \Delta(c^*A)_x] \\ &+ J_{wa} (O_x + \frac{1}{2} \Delta O_x) \\ &- J_{wb} (P_x + \frac{1}{2} \Delta P_x) \end{aligned} \quad (3.5)$$

In the above,

J_{wa} = areic mass flux of substance across the water-air interface; the flux is negative in the upward direction ($M.L^{-2}.T^{-1}$) (time, space)

J_{wb} = areic mass flux of substance across the water-sediment interface; the flux is positive in the downward direction ($M.L^{-2}.T^{-1}$) (time, space)

k = transformation rate coefficient⁵ for substance in the water column (T^{-1}) (-).

The limit notation (Δx , Δt , ΔO_x , ΔP_x , and $\Delta(c^*A)$ reach zero):

$$\frac{\partial(c^*A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} - k(c^*A) + J_{wa} \cdot O_x - J_{wb} \cdot P_x \quad (3.6)$$

This is the conservation equation for the water layer.

⁵ In this report lowercase k is used to indicate rate coefficients and uppercase K to indicate equilibrium coefficients.

3.2 Sediment

Diffusion across the water-sediment interface and sorption to sediment are important processes in describing the behaviour of pesticides in surface waters. In small watercourses, the walls account for an important part of the total exchange area between water and sediment. Therefore, it is more realistic to use the wetted perimeter, instead of the bottom width, to calculate the exchange area (De Heer, 1979).

Perpendicular to the wetted perimeter, the substance is transported from the water layer to the sediment. Per unit of length in the direction of flow, this transport takes place through an area P_0 (L^2).

$$P_0 = b + 2h\sqrt{s_1^2 + 1} \quad (3.7)$$

with

P_0 = wetted perimeter (L)

h = water level above bottom of ditch (L)

At a distance d from the water-sediment interface transport in the sediment takes place through the area P_d . (Fig. 6.)

$$P_d = b + 2d \cdot \tan\left(\frac{1}{2}\beta\right) + 2(h+d)\sqrt{s_1^2 + 1} \quad (3.8)$$

with

P_d = length of wetted perimeter at distance d from the water-sediment interface (L)

β = $\arctan(1/s_1)$ (1)

d = distance of water-sediment interface to area concerned (L).

Hence, the area through which transport takes place increases with increasing distance from the water-sediment interface.

Transport perpendicular to the wetted perimeter occurs in two dimensions, the z direction and the y direction. This two-dimensional transport has been simplified to a one-dimensional transport in the z direction. This implies that transport around corners is neglected; the ditch walls are, as it were, straightened and the flow pattern has been simplified to one-dimensional transport in a widening sediment column. (Fig. 7).

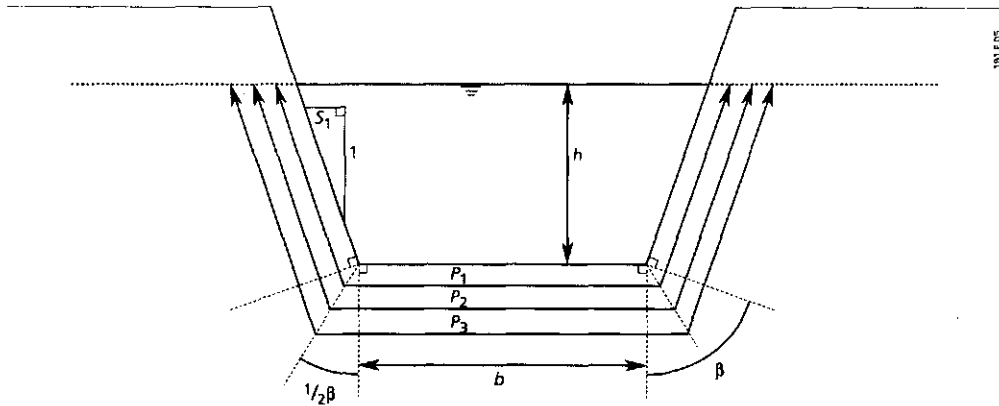


Fig. 6 Cross section of a ditch with the shape of stacked trapezia

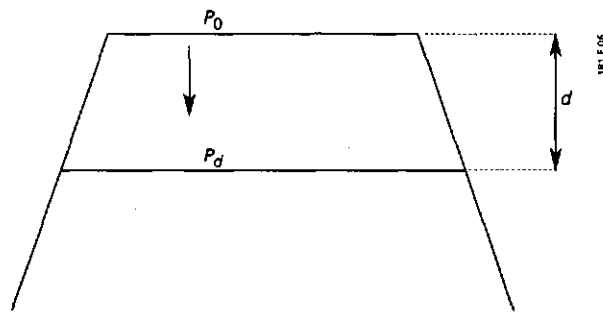


Fig. 7 The sediment subsystem simplified to one dimension

Hence, the **assumption** is that the two-dimensional transport can be calculated by means of a one-dimensional description of mass transport in a sediment column widening with depth.

This one-dimensional description has been applied to the sediment subsystem situated below the exchange area between water layer and sediment; this area has a constant size.

The **assumption** has been made that the size of the exchange area is constant, and does not depend on the varying water level in the ditch.

The size of the exchange area per unit of length in the x direction equals the wetted perimeter of the ditch corresponding to a particular water level h_w (Fig. 8). This can be the minimal water level occurring in the ditch. The mass transport thus takes place in a trapezium-shaped sediment column underneath the area of constant size $P_{z=0}$, per unit of length.

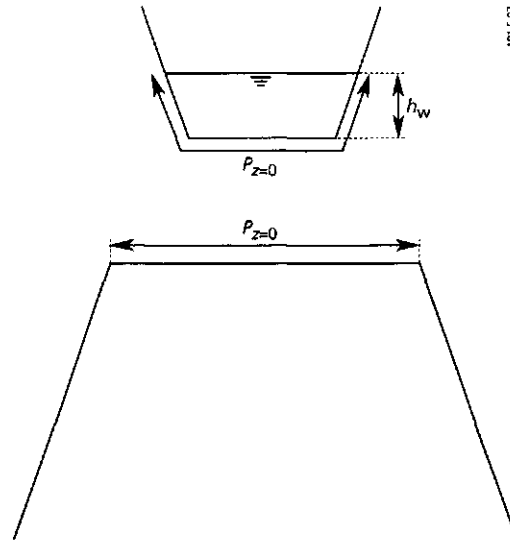


Fig. 8 The sediment subsystem situated below an area of constant size, through which exchange of pesticide between water layer and sediment occurs

$$P_{z=0} = b + 2h_w \sqrt{s_1^2 + 1} \quad (3.9)$$

The mass balance reads:
accumulation = input - output.

In Figure 9:

- c_b^* = mass concentration of substance in sediment ($M.L^{-3}$) (time, space)
- J_{lb} = areic mass flux of substance in the liquid phase of the sediment by advection, dispersion and diffusion ($M.L^{-2}.T^{-1}$) (time, space)
- $P_{z=0}$ = length of wetted perimeter at depth $z = 0$ (L) (space).

In differentials, the mass balance reads:

$$\begin{aligned} (P_z + \frac{1}{2}\Delta P_z) \cdot 1 \cdot \Delta z \cdot \Delta c_b^* &= \Delta t \cdot 1 \cdot (P_z \cdot J_{lb,z} - (P_z + \Delta P_z) \cdot J_{lb,z+\Delta z}) \\ \Leftrightarrow \frac{\Delta c_b^*}{\Delta t} &= \frac{P_z \cdot J_{lb,z} - P_z \cdot J_{lb,z+\Delta z}}{(P_z + \frac{1}{2}\Delta P_z) \Delta z} - \frac{\Delta P_z \cdot J_{lb,z+\Delta z}}{\Delta z \cdot (P_z + \frac{1}{2}\Delta P_z)} \end{aligned} \quad (3.10)$$

The limit notation (Δz , Δt and ΔP_z approach zero):

$$\frac{\partial c_b^*}{\partial t} = - \frac{\partial J_{lb}}{\partial z} - \frac{\partial P}{\partial z} \cdot \frac{J_{lb}}{P} \quad (3.11)$$

$$\Leftrightarrow P \frac{\partial c_b^*}{\partial t} = - \frac{\partial(PJ_{lb})}{\partial z}$$

This is the conservation equation for the sediment subsystem below the exchange area of constant size.

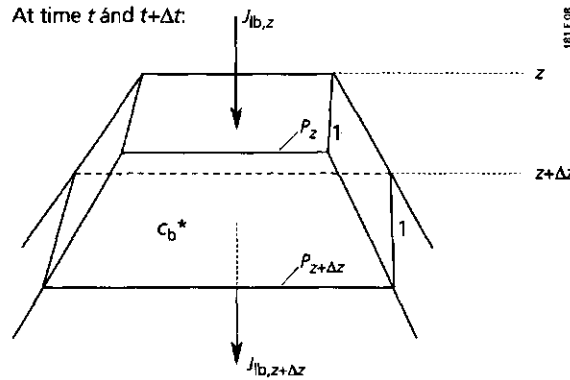


Fig. 9 Mass balance for the sediment subsystem
 * area per unit of length in the x direction is considered
 * wetted areas P_z and $P_{z + \Delta z}$ are constant in time

Analogously to the conservation equation for the water layer (Eq. 3.6), a sink term - $k_b c_b^* P$ has been added to the conservation equation for the sediment (Eq. 3.11). There is no source term for the pesticide. The assumption of negligible fluxes in the x direction in the sediment implies that there is no exchange of mass across the boundaries of the subsystem.

Note that the exchange term between water layer and sediment in the conservation equation for the water layer (Eq. (3.6)) becomes a mathematical boundary condition here. Hence, this term is not explicitly visible in the conservation equation for the sediment.

Taking the above-mentioned sink term into account, the complete conservation equation for the sediment now reads:

$$P \frac{\partial c_b^*}{\partial t} = - \frac{\partial(PJ_{lb})}{\partial z} - k_b c_b^* P \quad (3.12)$$

with

k_b = transformation rate coefficient for substance in the sediment (T^{-1}).

This conservation equation is coupled to the conservation equation for the water layer at the wetted perimeter P (at $z=0$) by the boundary condition for the flux J_{lb} at $z=0$. In principle, the conservation equations should be solved simultaneously.

4 Basic processes in water layer and sediment

4.1 Water layer

The conservation equation for the water layer reads:

$$\frac{\partial(c^*A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} - k(c^*A) + J_{wa} \cdot O_x - J_{wb} \cdot P_x \quad (3.6)$$

The substance concentration in the water layer, c^* , is described as follows.

$$c^* = c + \frac{DW \cdot P_{z=0}}{A} \cdot X_{mp} + ss \cdot X_{ss} \quad (4.1)$$

with

- c^* = mass concentration of substance in the water layer (M.L⁻³)
- c = mass concentration of substance in the water phase (M.L⁻³)
- DW = dry weight of macrophytes per area of sediment (M.L⁻²)
- X_{mp} = content of substance sorbed to macrophytes, i.e. the ratio of the mass of substance sorbed divided by the mass of dry macrophytes (M.M⁻¹)
- ss = mass concentration of suspended solids in the water layer, i.e. the ratio of the mass of dry suspended solids divided by the volume of water (M.L⁻³)
- X_{ss} = content of substance sorbed to suspended solids, i.e. the ratio of the mass of substance sorbed divided by the mass of dry suspended solids (M.M⁻¹)

By analogy to sorption to soil and sorption to the solid phase of sediment, the content of substance sorbed to the suspended solids equals:

$$X_{ss} = K_{F,ss} \cdot c_{e,ss} \cdot \left(\frac{c}{c_{e,ss}} \right)^{n_{ss}} \quad (4.2)$$

with

- $K_{F,ss}$ = Freundlich coefficient for sorption to suspended solids (L³.M⁻¹)
- $c_{e,ss}$ = concentration c at which $K_{F,ss}$ has been estimated (M.L⁻³)
- n_{ss} = Freundlich exponent for sorption to suspended solids (1)

The Freundlich coefficient for sorption to suspended solids, $K_{F,ss}$, is related to the $K_{om,ss}$ in the following way:

$$K_{F,ss} = m_{om,ss} \cdot K_{om,ss} \quad (4.3)$$

with

$m_{om,ss}$ = mass fraction of organic matter in the suspended solids ($M.M^{-1}$)

Hence, it has been **assumed** that sorption to suspended solids and sorption to sediment are analogous processes to sorption to soil and that both can be described with the aid of the Freundlich equation.

$K_{om,ss}$ = slope of sorption isotherm, based on the organic matter content ($L^3.M^{-1}$)

The content sorbed to macrophytes can be described as follows:

$$X_{mp} = K_{mp} \cdot c \quad (4.4)$$

with

K_{mp} = distribution coefficient for substance between macrophytes and water, i.e. the slope of the sorption isotherm based on the mass of dry macrophytes ($L^3.M^{-1}$).

The **assumption** is that sorption of the substance to suspended solids, to macrophytes and to the solid phase of the sediment (a.o. Wauchope and Myers, 1985) is a relatively rapid process compared to other processes, so one can assume an instantaneous equilibrium.

The wetted area A of the conservation equation equals:

$$A = bh + h^2 s_1 \quad (4.5)$$

The flux J of the conservation equation describes the transport of substance, both dissolved in water and sorbed to the suspended solids. It consists of an advective and a dispersive component (which are assumed identical for the dissolved and sorbed substance).

$$J = u(c + ss.X_{ss}) - E_x \frac{\partial(c + ss.X_{ss})}{\partial x} \quad (4.6)$$

with

J = areic mass flux of substance (both dissolved in water and sorbed to suspended solids) in the water layer ($M.L^{-2}.T^{-1}$)

u = flow velocity of the water ($L.T^{-1}$)

E_x = dispersion coefficient in the direction of flow ($L^2.T^{-1}$).

The term $\partial(AJ)/\partial x$ equals

$$\begin{aligned} & \frac{\partial(A[u(c + ss.X_{ss}) - E_x \frac{\partial(c + ss.X_{ss})}{\partial x}])}{\partial x} \\ & = \frac{\partial(Q(c + ss.X_{ss}) - AE_x \frac{\partial(c + ss.X_{ss})}{\partial x})}{\partial x} \end{aligned} \quad (4.7)$$

in which

$$Q = A.u$$

with

$$Q = \text{rate of discharge in the water layer (L}^3\text{.T}^{-1}\text{)}$$

The transformation rate coefficient k is a measure of the (entire) transformation of the substance. The substance may be dissolved in water or it may be sorbed to suspended solids or to macrophytes. The three most important transformation processes in the water layer are photolysis, hydrolysis and biodegradation (Thomann and Mueller, 1987). Photolysis depends mainly on the light intensity in the water column, hydrolysis depends especially on the pH and biodegradation is mainly determined by the extent and type of the bacterial community and the temperature.

The exchange flux of the substance between water body and atmosphere, J_{wa} , is described by the film model of two laminar layers at an interface (Liss and Slater, 1974, Mackay and Leinonen, 1975, review: Mackay, 1981).

$$J_{wa} = -k_1(c - c_1) = +k_g(c_a - c_{a,l}) \quad (4.8)$$

with

$$J_{wa} = \text{areic mass flux at the air-water interface (M.L}^{-2}\text{.T}^{-1}\text{)}$$

$$k_1 = \text{exchange coefficient of substance in the liquid phase (L.T}^{-1}\text{)}$$

$$k_g = \text{exchange coefficient of substance in the gas phase (L.T}^{-1}\text{)}$$

$$c_1 = \text{equilibrium mass concentration of substance at the water-gas interface in the water phase (M.L}^{-3}\text{)}$$

$$c_a = \text{mass concentration of substance in the air (M.L}^{-3}\text{)}$$

$$c_{a,l} = \text{equilibrium mass concentration of substance at the water-gas interface in the gas phase (M.L}^{-3}\text{)}$$

If the exchanging substance obeys Henry's law, then the equilibrium situation at the interface may be described by:

$$c_{a,l} = K_H \cdot c_1 \quad (4.9)$$

with

$$K_H = \text{dimensionless Henry coefficient (1)}.$$

Henry's coefficient is estimated from the quotient of mass concentration of saturated vapour of the substance (via vapour pressure) and the solubility of the substance in water.

$$K_H = \frac{P \cdot M_m}{R \cdot T} \cdot \frac{1}{c_{sol}} \quad (4.10)$$

with

$$P = \text{saturated vapour pressure of substance (L}^{-1}\text{.M.T}^{-2}\text{)}$$

$$M_m = \text{molecular mass (M.N}^{-1}\text{)}$$

$$R = \text{universal gas constant (L}^2\text{.M.T}^{-2}\text{.N}^{-1}\text{.}\theta^{-1}\text{)}$$

T = temperature at which the saturated vapour pressure, the solubility and the exchange coefficients in the liquid and gas phases are defined (θ)
 c_{sol} = solubility of substance in water (M.L^{-3}).

Eliminating c_l and $c_{a,l}$ from Eq. (4.8) with the aid of Eq. (4.9) flux J_{wa} becomes as follows:

$$J_{\text{wa}} = -k_{t,l} \left(c - \frac{c_a}{K_H} \right) \quad (4.11)$$

in which

$$\frac{1}{k_{t,l}} = \frac{1}{k_l} + \frac{1}{K_H \cdot k_g} \quad (4.12)$$

with

$k_{t,l}$ = overall transfer coefficient for the air-water interface, based on the liquid phase (L.T^{-1}).

The exchange flux between water layer and sediment consists of an advective and a diffusive component.

$$J_{\text{wb}} = J_{\text{wb,adv}} + J_{\text{wb,dif}} \quad (4.13)$$

with

J_{wb} = areic mass flux at the water-sediment interface ($\text{M.L}^{-2}.\text{T}^{-1}$)
 $J_{\text{wb,adv}}$ = areic mass flux by advection at the water-sediment interface ($\text{M.L}^{-2}.\text{T}^{-1}$)
 $J_{\text{wb,dif}}$ = areic mass flux by diffusion at the water-sediment interface ($\text{M.L}^{-2}.\text{T}^{-1}$).

The description of this flux given below assumes that diffusion across the water-sediment interface is rapid compared to diffusion in sediment. This implies that the interface resistance is negligible and that the mass concentration of the substance in the water layer equals that in the outer sediment pores.

The **assumption** is thus that the resistance to transport of the substance across the water-sediment interface can be neglected.

The advective component of the exchange flux consists either of a supply of water with the substance towards the ditch because of drainage from neighbouring field lots, or of infiltrating water from the ditch into the field lots. Because of the assumption of no transport resistance across the water-sediment interface ($c = c_{\text{lb},z=0}$), this yields:

$$J_{\text{wb,adv}} = \frac{\ell}{P_{z=0}} \cdot q \cdot c \quad (4.14)$$

with

ℓ = length of drained or infiltrated lot, oriented perpendicular to the ditch and extending on one or two sides of the ditch (L)

q = areic volume flux, i.e. volume of drained or supplied water divided by lot area and time ($L.T^{-1}$). The flux is positive for infiltration and negative for upward flow (drainage from the field lot).

The diffusive component of the exchange flux is described as the areic mass flux by diffusion in the sediment at the location $z=0$.

$$J_{wb,dif} = -(\epsilon \cdot D_{lb} \cdot \frac{\partial c_{lb}}{\partial z})_{z=0} \quad (4.15)$$

with

D_{lb} = diffusion coefficient of substance in the liquid phase of the sediment ($L^2.T^{-1}$)

ϵ = volume fraction of pore water, i.e. volume of liquid divided by volume of bottom material (1).

c_{lb} = mass concentration of substance in the liquid phase of the sediment ($M.L^{-3}$).

The diffusion coefficient of the substance in the liquid phase of the sediment, D_{lb} , is calculated as:

$$D_{lb} = \lambda \cdot D_w \quad (4.16)$$

with

λ = tortuosity factor, i.e. ratio of surface area of bottom material to liquid phase (1)

D_w = diffusion coefficient of substance in water ($L^2.T^{-1}$).

The conservation equation for the water layer (Eq. 3.6), is now entirely defined. All the terms of the equation have been unambiguously described. In fact, this results in one equation with only one unknown state variable (e.g. c). After solution, the other state variables (e.g. X_{ss} , X_{mp}) can be calculated. These are also determined by system parameters (like DW , $m_{om,ss}$, k , ϵ , D_w) and by inputs (like q and c_a). The state variables Q and h form an exception. They are not defined by the conservation equation, but partly define this themselves. They have been calculated in a water flow model and they need to be known at every chosen location and time if one is to calculate the water quality variables.

4.2 Sediment

The conservation equation for the sediment reads:

$$P \frac{\partial c_b^*}{\partial t} = - \frac{\partial(PJ_{lb})}{\partial z} - k_b c_b^* P \quad (3.12)$$

The concentration of substance in the sediment, c_b^* , is defined as:

$$c_b^* = \epsilon c_{lb} + \rho_b X_b \quad (4.17)$$

with

c_b^* = mass concentration of substance in sediment (M.L⁻³)

ρ_b = bulk density of dry bottom material, i.e. volumic mass of dry bottom material (M.L⁻³)

X_b = content of substance sorbed, i.e. the ratio of the mass of substance sorbed divided by the mass of dry bottom material (M.M⁻¹)

The content of substance sorbed to sediment, X_b , is (Jafvert, 1990):

$$X_b = K_{F,wb} \cdot c_{e,wb} \cdot \left(\frac{c_{lb}}{c_{e,wb}} \right)^{n_{wb}} \quad (4.18)$$

with

$K_{F,wb}$ = Freundlich coefficient for sorption to bottom material (L³.M⁻¹)

$c_{e,wb}$ = Concentration c at which $K_{F,wb}$ has been estimated (M.L⁻³)

n_{wb} = Freundlich exponent for sorption to bottom material (1)

$$K_{F,wb} = m_{om,wb} K_{om,wb} \quad (4.19)$$

with

$m_{om,wb}$ = mass fraction of organic matter of the suspended solids (M.M⁻¹)

$K_{om,wb}$ = slope of sorption isotherm, based on the organic matter content (L³.M⁻¹).

The transformation rate coefficient for the substance in the sediment, k_b , is a measure of the entire transformation of the substance in pore water as well as in sorbed form. The main transformation processes in the sediment are hydrolysis and biodegradation.

The perimeter below the water-sediment interface, P , depends on the distance to this interface and the chosen water level h_w in the ditch. For the simplified case of one-dimensional transport in the z direction, P is defined as follows (Fig. 10).

$$P = b + 2z \cdot \tan\left(\frac{1}{2}\beta\right) + 2(h_w + z) \sqrt{s_1^2 + 1} \quad (4.20)$$

with

β = arctan(1/ s_1) (1)

z = depth below the water-sediment interface (L)

In the pore water, the substance is transported by advection, dispersion and diffusion. This flux, J_{lb} , is defined as follows⁶.

$$J_{lb} = \frac{\ell}{P} \cdot q \cdot c_{lb} - \epsilon (E_{lb} + D_{lb}) \frac{\partial c_{lb}}{\partial z} \quad (4.21)$$

with

E_{lb} = dispersion coefficient in pore water ($L^2 \cdot T^{-1}$).

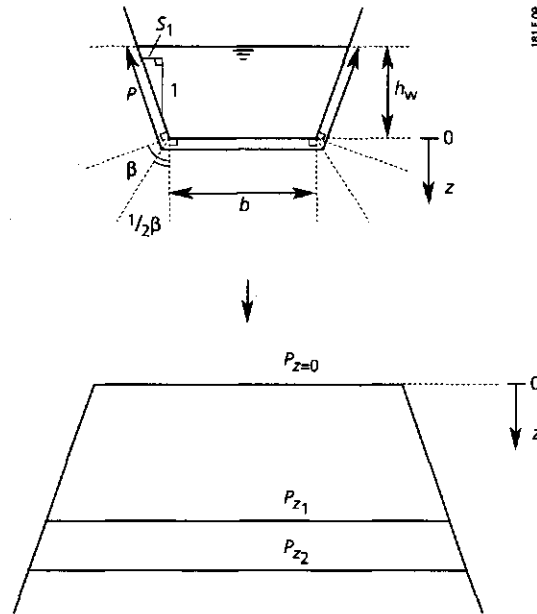


Fig. 10 Calculation of wetted perimeter in the sediment subsystem

The dispersion coefficient is defined as:

$$E_{lb} = L_{dis} \cdot |w| \quad (4.22)$$

with

L_{dis} = dispersion length (L)

w = average flow velocity of pore water (i.e. $\ell q / P \epsilon$) ($L \cdot T^{-1}$).

A special problem may occur in the sediment. Very sharp concentration gradients may exist here. The magnitude (and direction) of the diffusive and dispersive fluxes are calculated with the aid of concentration gradients. (See e.g. Eq. (4.21).) Dispersion of the substance originates in the unequal flow velocities of pore water (i.e. in the velocity distribution) and is thus caused by advection. It then logically follows that the combined material flux resulting from advection and its ensuing

⁶ The first term represents the advection flux, which equals: flow velocity of pore water * mass concentration in pore water * volume fraction of pore water = $(\ell q) / (P \epsilon) * c_{lb} * \epsilon = (\ell / P) * q * c_{lb}$

dispersion, should have the same direction as the advection. (Otherwise the effect would cancel out its cause) (Bolt, 1979, pp. 301 and 346).

Figure 11 shows the four possible situations, occurring in the sediment. In Situations 1 and 3, the advective and dispersive fluxes have the same direction, so the above-mentioned problem does not arise. In Situation 2 as well as in Situation 4, the advective and dispersive fluxes have opposite directions, so it is possible that the dispersive flux cancels out or even exceeds the advective flux. This is prevented as long as the following condition is fulfilled.

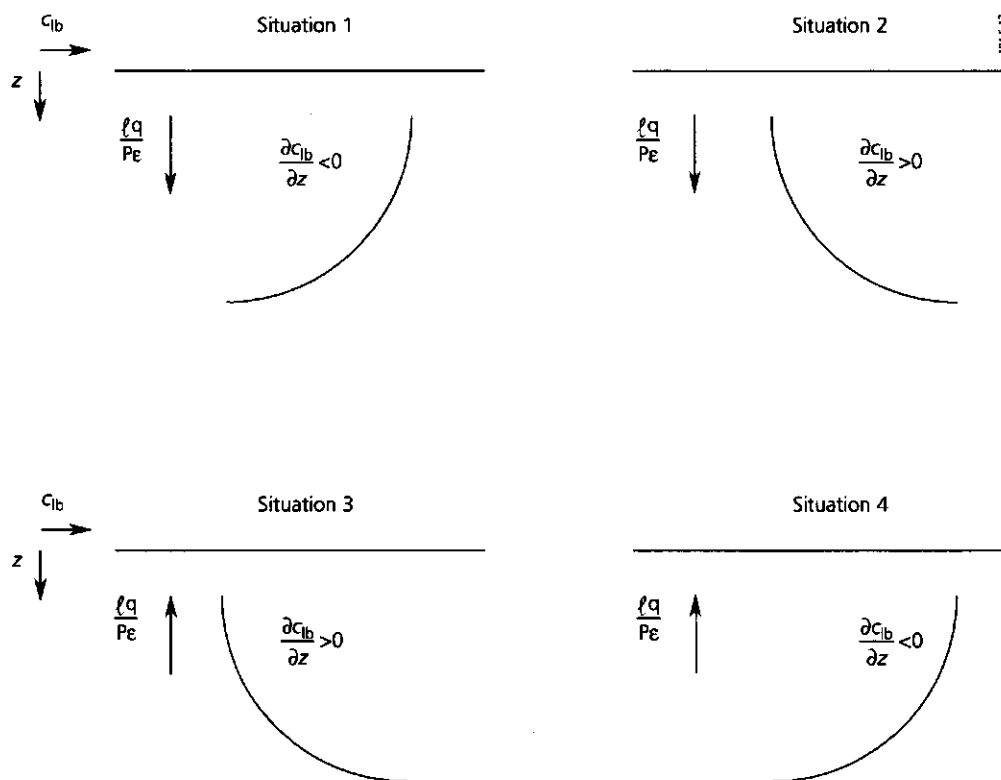


Fig. 11 Four situations in the sediment, depicting the possible combinations of the directions of the advection flow (lq/Pe) and the concentration gradient $\partial c_{lb}/\partial z$

For Situation 2:

$$q > 0 \text{ and } \partial(c_{lb})/\partial z > 0.$$

According to Eq. (4.22), with $q > 0$:

$$E_{lb} = L_{dis} \cdot \frac{lq}{Pe} \quad (4.22a)$$

The dispersive flux should not exceed the advective flux:

$$\begin{aligned}
\left| \frac{\ell}{P} q c_{lb} \right| &> \left| \epsilon E_{lb} \frac{\partial c_{lb}}{\partial z} \right| \Leftrightarrow \\
q c_{lb} &> L_{dis} q \frac{\partial c_{lb}}{\partial z} \Leftrightarrow \\
q \left(c_{lb} - L_{dis} \frac{\partial c_{lb}}{\partial z} \right) &> 0 \Leftrightarrow (q \text{ is positive}) \\
c_{lb} - L_{dis} \frac{\partial c_{lb}}{\partial z} &> 0
\end{aligned} \tag{4.23}$$

For Situation 4:

$q < 0$ and
 $\partial(c_{lb})/\partial z < 0$.

Combining Eq. (4.22) with $q < 0$ results in:

$$E_{lb} = L_{dis} \cdot \frac{-\ell q}{P\epsilon} \tag{4.22b}$$

The condition that the dispersive flux should not exceed the advective flux leads to:

$$\begin{aligned}
\left| \frac{\ell}{P} q c_{lb} \right| &> \left| \epsilon E_{lb} \frac{\partial c_{lb}}{\partial z} \right| \Leftrightarrow \\
- q c_{lb} &> L_{dis} \cdot -q \cdot \frac{\partial c_{lb}}{\partial z} \Leftrightarrow \\
- q \left(c_{lb} + L_{dis} \frac{\partial c_{lb}}{\partial z} \right) &> 0 \Leftrightarrow (-q \text{ is positive}) \\
c_{lb} + L_{dis} \frac{\partial c_{lb}}{\partial z} &> 0
\end{aligned} \tag{4.24}$$

This means that the dispersive flux is smaller than the advective flux if the following conditions are fulfilled:

$$\text{for } q > 0 \text{ and } \frac{\partial c_{lb}}{\partial z} > 0: \quad c_{lb} - L_{dis} \frac{\partial c_{lb}}{\partial z} > 0 \tag{4.23}$$

and

$$\text{for } q < 0 \text{ and } \frac{\partial c_{lb}}{\partial z} < 0: \quad c_{lb} + L_{dis} \frac{\partial c_{lb}}{\partial z} > 0 \tag{4.24}$$

In the case the dispersion flux would exceed the advective flux, it has been assumed that the dispersive flux will only cancel out the advective flux. Hence, only the diffusive flux is left in these cases. (See also De Heer, 1979, p.122.)

The conservation equation for the sediment is now entirely specified. All the terms of this equation have been unambiguously defined in this section. Just as for the water layer, the result is again one equation with one unknown variable (e.g. c_{1b}). This is coupled to the equation for the water layer via the boundary condition $c_{1b,z=0} = c$. In principle, the two conservation equations, for the water layer and for the sediment, need to be solved simultaneously.

5 Boundary conditions and initial condition, other input

5.1 Water layer

The conservation equation for the water layer reads:

$$\frac{\partial(c^*A)}{\partial t} = -\frac{\partial(AJ)}{\partial x} - k(c^*A) + J_{wa} \cdot O_x - J_{wb} \cdot P_x \quad (3.6)$$

The boundary condition at the upper end of the subsystem states that water without pesticide flows into the ditch.

For $t \geq 0$ and $x = 0$:

$$J = 0 \quad (5.1)$$

The boundary condition at the end of the ditch is:

For $t \geq 0$ and $x = \text{end value of ditch (e.g. 100 m)}$:

$$J = u(c + ss \cdot X_{ss}) - E_x \cdot \frac{\partial(c + ss \cdot X_{ss})}{\partial x} \quad (5.2)$$

The initial condition is that there is no pesticide in the water layer.

For $t = 0$ and $x > 0$:

$$c^* = 0 \quad (5.3)$$

Applications of pesticides into the ditch are not, strictly speaking, included in the boundary conditions (at $x = 0$ or $x = \text{end value}$) or the initial condition (at $t = 0$). These inputs to the subsystem are included as options in the conservation equation. A distinction is made between four situations.

In the first and second situations different (point-type or distributed) pulse inputs occur. In an infinitesimally small period of time, a mass of substance is released into the ditch water. The mass of substance applied per unit of volume during this period, $p(t,x)$, is described below. In both cases, a term $p(t,x)$ is added to the conservation equation.

In the first situation a distributed pulse input of the substance occurs, e.g. spray drift, a momentary runoff or a brief release from a number of nearby drains or trenches. It is assumed that the supplied lineic mass of substance M_L (dimension $M \cdot L^{-1}$) is instantaneously and ideally mixed over the water column at the time of release. Such an pulse input may occur repeatedly.

$$p_d(t,x) = \sum_{s=1}^m \frac{M_s}{A} \cdot \delta(t-t_s) \quad (5.4)$$

in which

$p_d(t,x)$ = distributed pulse input ($M.L^{-3}.T^{-1}$)

s = number of pulse input (total m inputs) (1)

and δ is the Dirac delta function (dimension T^{-1}) defined by (Fig. 12):

$\delta(t-t_s) = 0$ for $t \neq t_s$, and

$$\int_{-\infty}^{\infty} \delta(t-t_s) dt = 1 \quad (5.5)$$

with

t_s = time of pulse input (T).

In the second situation there is a point-type pulse input, e.g. spillage of the substance, rinsing application equipment, or a brief release from individual drains or trenches. The mass of substance, released in an infinitesimally short time, is M (dimension M) and this is assumed to be in an infinitesimally thin, vertical slice of the ditch. An pulse input, occurring repeatedly, can be described by:

$$p_p(t,x) = \sum_{s_0=1}^{m_0} \frac{M \cdot \delta(x-x_0) \cdot \delta(t-t_{s_0})}{A} \quad (5.6)$$

in which

$p_p(t,x)$ = point-type pulse input ($M.L^{-3}.T^{-1}$)

s_0 = number of pulse input at location x_0 (total m_0 inputs)

and δ is the Dirac delta function (dimension L^{-1}) defined by (Fig. 13):

$\delta(x-x_0) = 0$ for $x \neq x_0$ and

$$\int_{-\infty}^{\infty} \delta(x-x_0) dx = 1 \quad (5.7)$$

with

x_0 = location of pulse input (L).

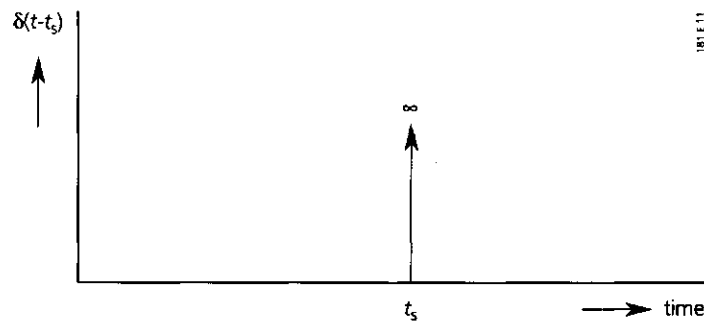


Fig. 12 Dirac delta function for pulse input

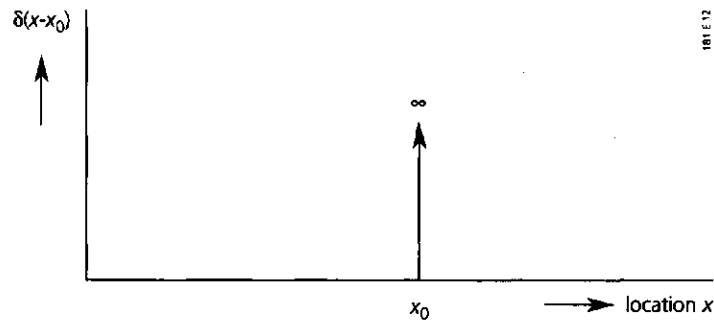


Fig. 13 Dirac delta function for point-type input

If such a pulse input occurs at a second location, the term $p_p(t,x)$ becomes as follows. (Etc.: for every subsequent location another term is added.)

$$p_p(t,x) = \sum_{s_0=1}^{m_0} \frac{M \cdot \delta(x-x_0) \cdot \delta(t-t_{s_0})}{A} + \sum_{s_{00}=1}^{m_1} \frac{M \cdot \delta(x-x_{00}) \cdot \delta(t-t_{s_{00}})}{A} \quad (5.8)$$

in which

s_{00} = number of pulse input at location x_{00} (total m_{00} inputs).

In the third and fourth situations, there are one or more point-type or distributed sources, which start emitting from a certain moment and continue doing so for a certain period. For both situations, the mass of substance applied per unit of volume per unit of time, $b(t,x)$, is described below. A term $b(t,x)$ is added to the conservation equation.

In the third situation, a continuous point release takes place, e.g. a discharging tributary or an individual trench adding water with a constant concentration of the substance. The mass of substance released per unit of time is M_T (dimension $M \cdot T^{-1}$, e.g. calculated as $Q_{\text{tributary}} c_{\text{tributary}}$) and this is ideally mixed over the cross sectional area A .

$$b_p(t,x) = \sum_{r=1}^n \frac{M_T}{A} \cdot \zeta(t) \cdot \delta(x-x_r) \quad (5.9)$$

in which

$b_p(t,x)$ = point source with continuous input ($M.L^{-3}.T^{-1}$)

r = number of point-type input (total n inputs)

and $\zeta(t)$ is the release time function (dimension (1)) defined by:

$\zeta(t)$ = 0 for $t < t_{rs}$ or $t > t_{re}$

= 1 for $t_{rs} \leq t \leq t_{re}$

with

t_{rs} = time at which the continuous release at location x_r starts (T)

t_{re} = time at which the continuous release at location x_e ends (T).

In the fourth situation, there is a continuous, distributed release, e.g. continuing release from many nearby drains or trenches. The lineic mass of substance released per unit of time is M_{LT} (dimension $M.L^{-1}.T^{-1}$, e.g. calculated as discharge per unit of length in the flow direction x times the concentration in trench/drain water: $(q\ell).c_{dg}$, with c_{dg} = mass concentration of substance in drain/trench water).

$$b_d(t,x) = \sum_{r=1}^n \frac{M_{LT}}{A} \cdot \zeta(t) \quad (5.10)$$

with

$b_d(t,x)$ = distributed source with continuous input ($M.L^{-3}.T^{-1}$).

Depending on the type of application of the substance to the subsystem, one or more terms are added to the conservation equation, Eq. (3.6). These optional terms have been defined by Eqs. (5.4), (5.6), (5.8), (5.9) and (5.10). Hence, the completed conservation equation for the water layer reads:

$$\frac{\partial(c \cdot A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} - k(c \cdot A) + J_{wa} \cdot O_x - J_{wb} \cdot P_x$$

(3.6a)

option: $+p_d \cdot A$

option: $+p_p \cdot A$

option: $+b_p \cdot A$

option: $+b_d \cdot A$

5.2 Sediment

The conservation equation for the sediment reads:

$$P \frac{\partial c_b^*}{\partial t} = - \frac{\partial(PJ_{lb})}{\partial z} - k_b c_b^* P \quad (3.12)$$

The mass flux at the water-sediment interface is composed of an advective and a diffusive component (there is no dispersive component across the interface):

$$J_{lb} = \frac{\ell}{P_0} q \cdot c_{lb} - \varepsilon D_{lb} \cdot \frac{\partial c_{lb}}{\partial z} \quad (5.11)$$

The boundary condition at the sediment surface is:

For $t \geq 0$ and $z = 0$:

$$c_{lb} = c \quad (5.12)$$

The boundary condition at the lower end of the sediment subsystem consists of the mass fluxes by advection, dispersion and diffusion.

For $t \geq 0$ and $z = \text{end value of sediment (e.g. 0.10 m)}$:

$$J_{lb} = \frac{\ell}{P} q \cdot c_{lb} - \varepsilon (E_{lb} + D_{lb}) \frac{\partial c_{lb}}{\partial z} \quad (5.13)$$

There are two options for the initial condition. One option assumes the sediment to be free of pesticide.

For $t = 0$ and $z > 0$:

$$c_b^* = 0 \quad (5.14)$$

In the other option there is sediment with a certain mass of pesticide, specified as a function of depth.

For $t = 0$ and $z > 0$:

$$c_b^* = f(z) \quad (5.15)$$

in which

$f(z)$ = mass concentration of substance in sediment as a function of depth at starting time $t = 0$ ($M.L^{-3}$).

6 Numerical solution of the mass conservation equations

6.1 Introduction

The conservation equation for the water subsystem Eq. (3.6) and that for the sediment subsystem Eq. (3.12) have been solved numerically with the aid of the finite-difference method. For this purpose a rectangular grid of points in the water subsystem was defined in the (x,t) plane, numbered $i = 1, 2, 3, \dots$ along the x axis and $j = 1, 2, 3, \dots$ along the t axis. The x axis was assumed to be positive in the direction of dominant flow. Δx_i was defined as the length of a segment around point i , while Δt was defined as the time step (Fig. 14).

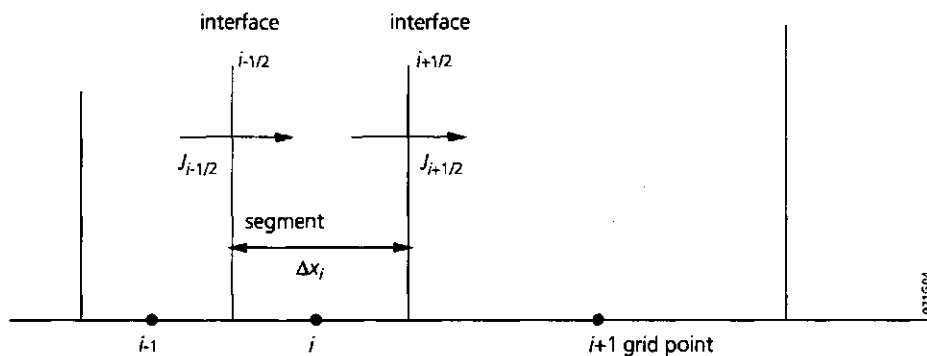


Fig. 14 Outline defining the discretisation of the x axis (water layer)

A rectangular grid of points in the sediment subsystem was defined in the (z,t) plane, numbered $k = 1, 2, 3, \dots$ along the z axis and $j = 1, 2, 3, \dots$ along the t axis. The z axis was assumed to be positive in the downward direction. Δz_k is the thickness of a segment around point k and Δt is again the time step (Fig. 15).⁷

Concentrations in the grid were defined at the grid points, while fluxes were defined at the interfaces, as well as the corresponding areas through which the fluxes occur (e.g. Bella and Dobbins, 1968). In the water subsystem, water flow was described with the aid of the water depth at a grid point and the flow velocity or rate of discharge through an interface.

⁷ As both subsystems, water layer and sediment, are coupled to each other, it was assumed that the same time step Δt was employed in both water layer and sediment. It is, however, possible that one of the subsystems permits a much larger time step than the other and that one could make use of this to shorten the calculation time. In that case a Δt_w and a Δt_s may be defined, in which one is a multiple of the other. TOXSWA 1.0 allows only identical time steps to be applied.

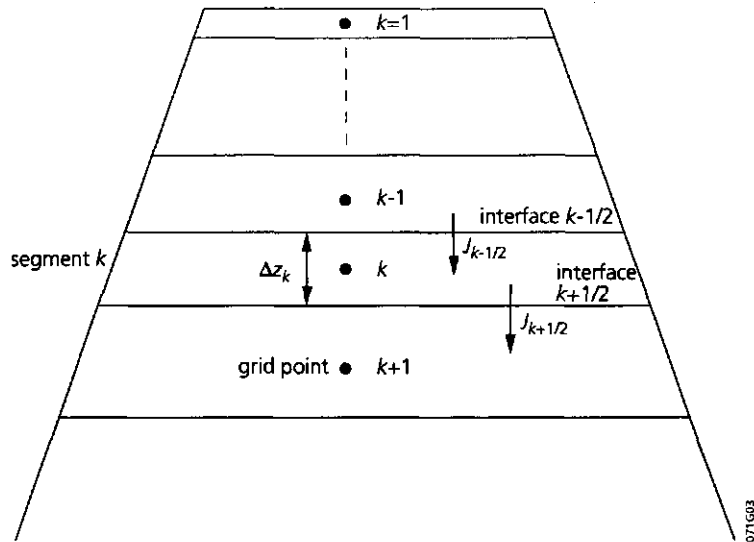


Fig. 15 Outline defining the discretisation of the z axis (sediment)

The terms upper and lower boundary have been defined here in relation to the variable of space: the upper boundary is located at $x = 0$ or $z = 0$ and the lower boundary at the end values of x and z . Table 1 shows an overview of the variables and the locations at which these were defined, for both the water and sediment subsystems.

The finite-difference method offers several possibilities. Spatial derivatives may be approximated with the aid of forward, central or backward differences. Temporal derivatives of the differential equation considered may be approximated in an explicit or an implicit manner. Selecting the calculation scheme depends on two factors, calculation time and acceptable numerical dispersion. The calculation time is mainly determined by the size of the space steps, Δx or Δz , or of the time steps, Δt . The step sizes are prescribed by the conditions which have been imposed to obtain stability, or in fact positivity of the solution. An unstable solution of the differential equations oscillates and diverges; a stable solution converges, but it can still produce negative solutions (i.e. concentrations), which is physically impossible. Therefore, one is interested in stable and positive solutions of the differential equations considered. The imposed positivity conditions lead to conditions for the allowed size of time and space steps. Hence, it is not only the dispersion coefficient and the flow velocity, but also the selected calculation scheme which determine the step sizes allowed.

Table 1 Overview of variables in the conservation equations for the water and sediment subsystems, demonstrating at which location in the grid these variables are defined

	Water layer	Sediment
Water flow		
grid point	h	-
interface	Q, u	$\frac{\ell q}{P \varepsilon}$
Water quality		
grid point	c, c^* A, b (to calculate volumes) k O_x, P_x DW X_{mp}, X_{ss} J_{wa}, J_{wb}	c_{1b}, c_b^* P (to calculate volumes) k X_b ε
interface	uA	$\left(\frac{\ell}{P} \frac{q}{\varepsilon} \right) P$
	E_x, A	D_{1b}, P, E_{1b}, P

Which size of numerical dispersion is acceptable depends on the types of issue studied. If one is studying point sources resulting in sharp concentration gradients, one does not want the selected calculation scheme to smoothen these gradients rapidly (i.e. in an artificial way). This would not agree with the actual situation. In such cases one should select a calculation scheme resulting in a small numerical dispersion. In the case of releases from diffuse origins no steep concentration gradients are present in the surface water, which makes it possible to select a calculation scheme with a relatively large numerical dispersion. Small numerical dispersions are often linked to small possible time and space steps, which results in long calculation times on the computer. If the numerical dispersion is allowed to be larger, it is generally possible to use larger steps, involving less calculation time.

Selecting the calculation scheme thus involves finding a balance between calculation time and acceptable numerical dispersion. Which size of numerical dispersion can be accepted, depends on the types of issue studied (steep concentration gradients or not).

6.2 Numerical weight factors

The finite-difference method in its most generalized form makes use of two weight factors in its calculation scheme. A weight factor θ for time indicates whether the solution scheme is implicit or explicit, or a weighted average of these two extremes. A weight factor α for space indicates whether the solution method is backward,

forward or somewhere in between backward and forward. In the case of a constant segment size, the concentration used to estimate convective flow between the grid points i and $i+1$, $c_{i+1/2}$, can be expressed as:

$$c_{i+1/2} = (1 - \alpha) c_i + \alpha c_{i+1} \quad (6.1)$$

with

$$0 \leq \alpha \leq 1.$$

By analogy:

$$c_{i-1/2} = (1 - \alpha) c_{i-1} + \alpha c_i \quad (6.1a)$$

A backward difference is used for $\alpha = 0$, a forward one for $\alpha = 1$.

Similarly, the concentration used to approximate the derivative with respect to time between time $j\Delta t$ and $(j+1)\Delta t$ (defined as $c_{j+1/2}$) reads:

$$c_i^{j+1/2} = \theta c_i^j + (1 - \theta) c_i^{j+1} \quad (6.2)$$

with

$$0 \leq \theta \leq 1.$$

An explicit calculation scheme is applied for $\theta = 1$, while an implicit calculation scheme is used for $\theta = 0$.

If the segment size is not constant but differs per grid point, the situation becomes more complicated (Fig. 16). It remains possible, however, to obtain the same type of approximation for all segments for only one value of a weight factor β (comparable to the weight factor α discussed above). We will define β in such a way that:

- $\beta = 0$ corresponds with a backward difference,
- $\beta = 1/2$ corresponds with a central difference and
- $\beta = 1$ corresponds with a forward difference.

We did so, because it is logical to have a value 1 for a forward difference.

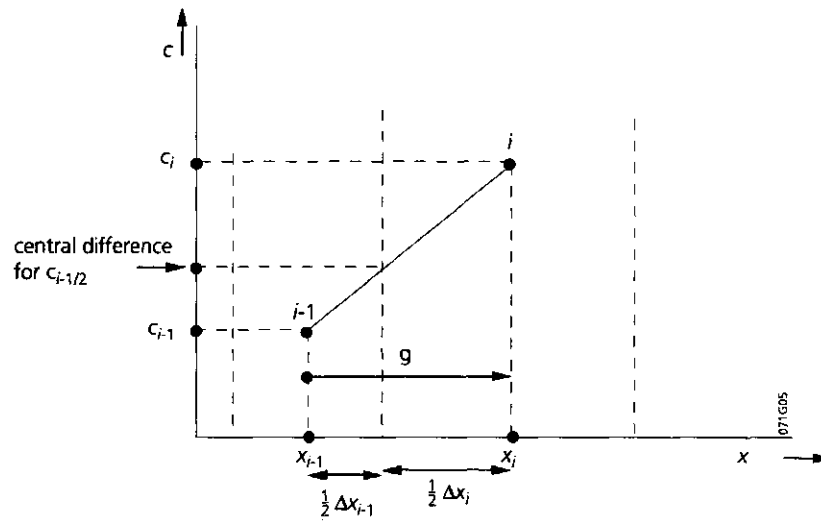


Fig. 16 Linear interpolation to determine the concentration at the interface, $c_{i-1/2}$ with the aid of the concentrations at grid points $i-1$ and i , c_{i-1} and c_i respectively

The curve between the grid points $i-1$ and i in Figure 16 is described by:

$$c_{i-1/2} = c_{i-1} + \frac{c_i - c_{i-1}}{x_i - x_{i-1}} g_{i-1/2} \quad (6.3)$$

in which g is the location at which $c_{i-1/2}$ is calculated. We want to express g as a function of Δx_{i-1} , Δx and β with the restriction that $g = \frac{1}{2}\Delta x_{i-1}$ if $\beta = \frac{1}{2}$, because this is the central difference case. This results in the following function g (Fig. 17):

$$0 \leq \beta \leq \frac{1}{2} : \quad g_{i-1/2} = \beta \Delta x_{i-1}$$

$$\frac{1}{2} < \beta \leq 1 : \quad g_{i-1/2} = \frac{1}{2} \Delta x_{i-1} + (\beta - \frac{1}{2}) \Delta x_i \quad (6.4)$$

Increasing β from 0 to 1 corresponds to going from a calculation of the concentration, $c_{i-1/2}$, only with the aid of c_{i-1} (fully backward) up to a calculation of $c_{i-1/2}$ only with the aid of c_i (fully forward). Linear interpolation takes place between these two extremes.

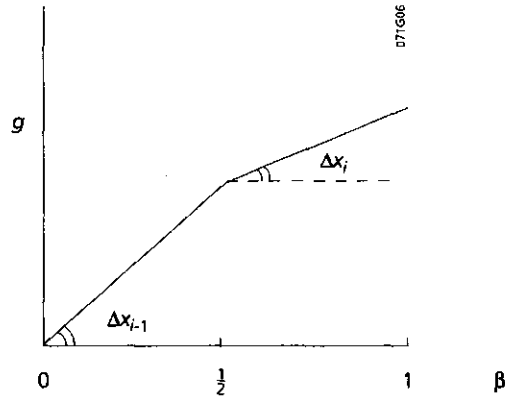


Fig. 17 The function $g_{i-1/2}$ with B representing a weight factor of space and Δx_{i-1} and Δx_i the slope parameters

Eq. (6.3) can be rewritten as:

$$c_{i-1/2} = \left(1 - \frac{g_{i-1/2}}{x_i - x_{i-1}} \right) c_{i-1} + \frac{g_{i-1/2}}{x_i - x_{i-1}} c_i \quad (6.3a)$$

This means that the weight factor $(1-\alpha)$ from Eq. (6.1a) has been replaced by:

$$1 - \frac{g_{i-1/2}}{x_i - x_{i-1}}$$

and the weight factor α by:

$$\frac{g_{i-1/2}}{x_i - x_{i-1}}$$

Below, Eq. (6.3a) will be shown as:

$$c_{i-1/2} = (1 - NWW_{i-1/2}) c_{i-1} + NWW_{i-1/2} c_i \quad (6.3b)$$

NWW , the Numerical Weight factor for the Water layer, stands for:

$$NWW_{i-1/2} = \frac{g_{i-1/2}}{x_i - x_{i-1}} \quad (6.5)$$

Similarly:

$$c_{i+\frac{1}{2}} = (1 - NWW_{i+\frac{1}{2}}) c_i + NWW_{i+\frac{1}{2}} c_{i+1} \quad (6.6)$$

with

$$NWW_{i+\frac{1}{2}} = \frac{g_{i+\frac{1}{2}}}{x_{i+1} - x_i} \quad (6.5a)$$

The original weight factor α from Eqs. (6.1) and (6.1a) has now been replaced by the factor $NWW_{i-\frac{1}{2}}$ or $NWW_{i+\frac{1}{2}}$. A calculation scheme with fully forward differences according to space is described by a value of 1 for all three weight factors introduced. Table 2 shows the outline.

Table 2 Outline of weight factors introduced according to space, using the finite-difference method in TOXSWA

Difference approximation	Weight factor α	Weight factor β	NWW
Forward	1	1	1
Central	$\frac{1}{2}$	$\frac{1}{2}$	depending on the ratio of the segment sizes
Backward	0	0	0

By analogy:

$$c_{lb\ k-\frac{1}{2}} = (1 - NWB_{k-\frac{1}{2}}) c_{lb\ k-1} + NWB_{k-\frac{1}{2}} c_{lb\ k} \quad (6.7)$$

in which the NWB , the Numerical Weight factor for the sediment [water Bottom], equals:

$$NWB_{k-\frac{1}{2}} = \frac{g_{i-\frac{1}{2}}}{z_k - z_{k-1}} \quad (6.8)$$

Similarly:

$$c_{lb\ k+\frac{1}{2}} = (1 - NWB_{k+\frac{1}{2}}) c_{lb\ k} + NWB_{k+\frac{1}{2}} c_{lb\ k+1} \quad (6.7a)$$

with

$$NWB_{k+\frac{1}{2}} = \frac{g_{i+\frac{1}{2}}}{z_{k+1} - z_k} \quad (6.8a)$$

In the TOXSWA model, the calculation scheme can be determined by the model user by selecting the weight factors β for space and θ for time. The selected calculation scheme for the sediment does not necessarily correspond with that for the water layer, so the value of β or θ for the sediment may differ from those selected for the water layer.

6.3 Numerical dispersion

For the classical advection-diffusion equation

$$\frac{\partial c}{\partial t} = -u \frac{\partial c}{\partial x} + D \frac{\partial^2 c}{\partial x^2}, \quad (6.9)$$

so applying to a system without transformation and without external exchanges, the coefficient for numerical dispersion can be approximated by:

$$E_{\text{num}} = \frac{u}{2} [(1 - 2\alpha) \Delta x + (1 - 2\theta) u \Delta t] \quad (6.10)$$

(Van Genuchten and Wieringa, 1974; Aalderink, 1993)

When differences in segment sizes are taken into account, the equation reads:

$$E_{\text{num } i-1/2} = \frac{u}{2} [(1 - 2 NWW_{i-1/2}) (x_i - x_{i-1}) + (1 - 2\theta) u \Delta t] \quad (6.11)$$

or

$$E_{\text{num } i+1/2} = \frac{u}{2} [(1 - 2 NWW_{i+1/2}) (x_{i+1} - x_i) + (1 - 2\theta) u \Delta t] \quad (6.11a)$$

For an explicit calculation scheme with backward differences, this results e.g. in ($\alpha = 0$ or $NWW = 0$ and $\theta = 1$):

$$E_{\text{num } i-1/2} = \frac{u}{2} [(x_i - x_{i-1}) - u \Delta t] \quad (6.12)$$

In such schemes the numerical dispersion can be suppressed by appropriately choosing $(x_i - x_{i-1})$ and Δt .

The numerical dispersion in implicit schemes ($\theta = 0$) is always larger than that in explicit schemes. Note that the expression 'numerical dispersion' is somewhat misleading, because E_{num} can also be negative (e.g. if $\alpha = 1$ and $\theta = 1$), so then 'numerical steepening' occurs. An example of an implicit calculation scheme with backward differences would be ($\alpha = 0$ or $NWW = 0$ and $\theta = 0$):

$$E_{\text{num } i-\frac{1}{2}} = \frac{u}{2} [(x_i - x_{i-1}) + u \Delta t] \quad (6.13)$$

The approximated numerical dispersion for a Crank-Nicholson scheme ($\alpha = \frac{1}{2}$, $\theta = \frac{1}{2}$) with equal segment sizes results in:

$$E_{\text{num}} = 0$$

For unequal segment sizes, the factor NWW will not be exactly $\frac{1}{2}$, which means that the approximated numerical dispersion will not equal zero. Generally, the Crank-Nicholson scheme will lead to small time and space steps, so to long calculation times.

If the numerical dispersion caused by the selected calculation scheme is approximately known, a more accurate approximation of Eq. (6.9) may be obtained, provided the appropriate difference equation uses a 'calculation' dispersion of the form:

$$E_{\text{calculation}} = E_{\text{physical}} - E_{\text{numerical}} \quad (6.14)$$

This is a better method for simulating real, measured concentration courses (Van Genuchten and Wieringa, 1974).

In the TOXSWA model, the numerical dispersion coefficients have been approximated for the water and sediment subsystems. This has been done by rewriting the advection-diffusion equations for both subsystems, expressing c^* and c_b^* as a factor multiplied by the concentration in the water and liquid phase, respectively. Next Eq. (6.10) has been applied (Cf. Wierenga and van Genuchten, 1974 and Boesten, 1986). For the water subsystem, this results in:

$$E_{\text{num}, i-\frac{1}{2}} = \frac{u}{2} \left[(1 - 2NWW_{i-\frac{1}{2}}) (x_i - x_{i-1}) + (1 - 2\theta) \frac{u \left(1 + ssK_F n_{ss} \left(\frac{c}{c_e} \right)^{n_{ss}-1} \right)}{\left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ssK_F n_{ss} \left(\frac{c}{c_e} \right)^{n_{ss}-1} \right)} \Delta t \right] \quad (6.15)$$

$$E_{\text{num}, i+\frac{1}{2}} = \frac{u}{2} \left[(1 - 2NWW_{i+\frac{1}{2}}) (x_{i+1} - x_i) + (1 - 2\theta) \frac{u \left(1 + ssK_F n_{ss} \left(\frac{c}{c_e} \right)^{n_{ss}-1} \right)}{\left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ssK_F n_{ss} \left(\frac{c}{c_e} \right)^{n_{ss}-1} \right)} \Delta t \right] \quad (6.15a)$$

(Assumptions made include that ss , K_F , K_{mp} and A are constant in time.)

The numerical dispersion coefficients for the sediment subsystem are given by:

$$E_{\text{num},k-\frac{1}{2}} = \frac{\ell q}{2P} \left[(1 - 2NWB_{k-\frac{1}{2}}) (z_k - z_{k-1}) + (1 - 2\theta) \frac{\ell q}{P \left(\epsilon + \rho_b K_F n_{wb} \left(\frac{c_{lb}}{c_e} \right)^{n_{wb}-1} \right)} \Delta t \right] \quad (6.16)$$

$$E_{\text{num},k+\frac{1}{2}} = \frac{\ell q}{2P} \left[(1 - 2NWB_{k+\frac{1}{2}}) (z_{k+1} - z_k) + (1 - 2\theta) \frac{\ell q}{P \left(\epsilon + \rho_b K_F n_{wb} \left(\frac{c_{lb}}{c_e} \right)^{n_{wb}-1} \right)} \Delta t \right] \quad (6.16a)$$

(Assumptions made include that ϵ , ρ_b , K_F and P are constant in time.)

6.4 Water layer

The right-hand term of the conservation equation for the water layer, Eq. (3.6), was approximated by applying the generalised finite-difference equation. In this equation, spatial derivatives and concentrations are evaluated at time $j+\frac{1}{2}$, as defined by Eq. (6.2) which implies that θ is user defined.

$$\left(-\frac{\partial(AJ)}{\partial x} \right)_i^{j+\frac{1}{2}} - k (c^* A)_i^{j+\frac{1}{2}} + (J_{wa} O_x)_i^{j+\frac{1}{2}} - (J_{wb} P_x)_i^{j+\frac{1}{2}} \approx \frac{(AJ)_{i-\frac{1}{2}}^{j+\frac{1}{2}} - (AJ)_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} - k (c^* A)_i^{j+\frac{1}{2}} + (J_{wa} O_x)_i^{j+\frac{1}{2}} - (J_{wb} P_x)_i^{j+\frac{1}{2}} \quad (6.17)$$

The term $(AJ)_{i-\frac{1}{2}}^{j+\frac{1}{2}}$ for values of i from 2 to m is given by:

$$(AJ)_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \left[uA (c + ssX_{ss}) - AE_x \left(\frac{\partial(c + ssX_{ss})}{\partial x} \right) \right]_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \quad (6.18)$$

$$Q_{i-\frac{1}{2}}^{j+\frac{1}{2}} (c_{i-\frac{1}{2}}^{j+\frac{1}{2}} + ssX_{ss\ i-\frac{1}{2}}^{j+\frac{1}{2}}) - A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i-\frac{1}{2}}^{j+\frac{1}{2}} \left(\frac{(c_i^{j+\frac{1}{2}} + ssX_{ss\ i}^{j+\frac{1}{2}}) - (c_{i-1}^{j+\frac{1}{2}} + ssX_{ss\ i-1}^{j+\frac{1}{2}})}{\frac{1}{2}\Delta x_i + \frac{1}{2}\Delta x_{i-1}} \right)$$

in which

$$Q_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \theta Q_{i-\frac{1}{2}}^j + (1 - \theta) Q_{i-\frac{1}{2}}^{j+1}$$

$$c_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \theta c_{i-\frac{1}{2}}^j + (1 - \theta) c_{i-\frac{1}{2}}^{j+1} = \theta (1 - NWW_{i-\frac{1}{2}}) c_{i-1}^j + \theta NWW_{i-\frac{1}{2}} c_i^j +$$

$$(1 - \theta) (1 - NWW_{i-\frac{1}{2}}) c_{i-1}^{j+1} + (1 - \theta) NWW_{i-\frac{1}{2}} c_i^{j+1}$$

ss is constant in time and space

$$X_{ss\ i-\frac{1}{2}}^{j+\frac{1}{2}} = \theta X_{ss\ i-\frac{1}{2}}^j + (1 - \theta) X_{ss\ i-\frac{1}{2}}^{j+1} =$$

$$\theta (1 - NWW_{i-\frac{1}{2}}) X_{ss\ i-1}^j + \theta NWW_{i-\frac{1}{2}} X_{ss\ i}^j +$$

$$(1 - \theta) (1 - NWW_{i-\frac{1}{2}}) X_{ss\ i-1}^{j+1} + (1 - \theta) NWW_{i-\frac{1}{2}} X_{ss\ i}^{j+1}$$

$$A_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \theta A_{i-\frac{1}{2}}^j + (1 - \theta) A_{i-\frac{1}{2}}^{j+1} \quad \text{and}$$

$$A_{i-\frac{1}{2}} = f(h_{i-\frac{1}{2}})$$

$$E_{i-\frac{1}{2}}^{j+\frac{1}{2}} = \theta E_{i-\frac{1}{2}}^j + (1 - \theta) E_{i-\frac{1}{2}}^{j+1}$$

$$c_i^{j+\frac{1}{2}} = \theta c_i^j + (1 - \theta) c_i^{j+1}$$

$$X_{ss\ i}^{j+\frac{1}{2}} = \theta X_{ss\ i}^j + (1 - \theta) X_{ss\ i}^{j+1}$$

$$c_{i-1}^{j+\frac{1}{2}} = \theta c_{i-1}^j + (1 - \theta) c_{i-1}^{j+1}$$

$$X_{ss\ i-1}^{j+\frac{1}{2}} = \theta X_{ss\ i-1}^j + (1 - \theta) X_{ss\ i-1}^{j+1}$$

By analogy the term $(AJ)_{i+\frac{1}{2}}^{j+\frac{1}{2}}$ for values of i from 1 to $(m-1)$ reads:

$$(AJ)_{i+\frac{1}{2}}^{j+\frac{1}{2}} = [uA (c + ssX_{ss}) - AE_x \left(\frac{\partial(c + ssX_{ss})}{\partial x} \right)]_{i+\frac{1}{2}}^{j+\frac{1}{2}} = \quad (6.19)$$

$$Q_{i+\frac{1}{2}}^{j+\frac{1}{2}} (c_{i+\frac{1}{2}}^{j+\frac{1}{2}} + ssX_{ss\ i+\frac{1}{2}}^{j+\frac{1}{2}}) - A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}} \left(\frac{(c_{i+1}^{j+\frac{1}{2}} + ssX_{ss\ i+1}^{j+\frac{1}{2}}) - (c_i^{j+\frac{1}{2}} + ssX_{ss\ i}^{j+\frac{1}{2}})}{\frac{1}{2}\Delta x_{i+1} + \frac{1}{2}\Delta x_i} \right)$$

in which

$$Q_{i+\frac{1}{2}}^{j+\frac{1}{2}} = \theta Q_{i+\frac{1}{2}}^j + (1 - \theta) Q_{i+\frac{1}{2}}^{j+1}$$

$$c_{i+\frac{1}{2}}^{j+\frac{1}{2}} = \theta c_{i+\frac{1}{2}}^j + (1 - \theta) c_{i+\frac{1}{2}}^{j+1} = \theta (1 - NWW_{i+\frac{1}{2}}) c_i^j + \theta NWW_{i+\frac{1}{2}} c_{i+1}^j +$$

$$(1 - \theta) (1 - NWW_{i+\frac{1}{2}}) c_i^{j+1} + (1 - \theta) NWW_{i+\frac{1}{2}} c_{i+1}^{j+1}$$

$$X_{ss\ i+\frac{1}{2}}^{j+\frac{1}{2}} = \theta (1 - NWW_{i+\frac{1}{2}}) X_{ss\ i}^j + \theta NWW_{i+\frac{1}{2}} X_{ss\ i+1}^j +$$

$$(1 - \theta) (1 - NWW_{i+\frac{1}{2}}) X_{ss\ i}^{j+1} + (1 - \theta) NWW_{i+\frac{1}{2}} X_{ss\ i+1}^{j+1}$$

$$A_{i+\frac{1}{2}}^{j+\frac{1}{2}} = \theta A_{i+\frac{1}{2}}^j + (1 - \theta) A_{i+\frac{1}{2}}^{j+1} \quad \text{with}$$

$$A_{i+\frac{1}{2}} = f(h_{i+\frac{1}{2}})$$

$$E_{i+\frac{1}{2}}^{j+\frac{1}{2}} = \theta E_{i+\frac{1}{2}}^j + (1 - \theta) E_{i+\frac{1}{2}}^{j+1}$$

$$c_{i+1}^{j+1/2} = \theta c_{i+1}^j + (1 - \theta) c_{i+1}^{j+1}$$

$$X_{ss\ i+1}^{j+1/2} = \theta X_{ss\ i+1}^j + (1 - \theta) X_{ss\ i+1}^{j+1}$$

The term $-k(c^*A)_i^{j+1/2}$ equals:

$$-k (c^* A)_i^{j+1/2} = -k(c_i^{*j+1/2} A_i^{j+1/2}) \quad (6.20)$$

in which

$$c_i^{*j+1/2} = \theta c_i^{*j} + (1 - \theta) c_i^{*j+1}$$

$$A_i^{j+1/2} = \theta A_i^j + (1 - \theta) A_i^{j+1}$$

The term $(J_{wa}O_x)_i^{j+1/2}$ equals:

$$(J_{wa}O_x)_i^{j+1/2} = (J_{wa\ i}^{j+1/2} O_{x\ i}^{j+1/2}) = -k_{L1} \left(c_i^{j+1/2} - \frac{c_{a\ i}^{j+1/2}}{K_H} \right) O_{x\ i}^{j+1/2} \quad (6.21)$$

in which

$$c_{a\ i}^{j+1/2} = \theta c_{a\ i}^j + (1 - \theta) c_{a\ i}^{j+1}$$

(it has been assumed that c_a , the pesticide background concentration in the air, is constant)

$$O_{x\ i}^{j+1/2} = \theta O_{x\ i}^j + (1 - \theta) O_{x\ i}^{j+1}$$

The term $-(J_{wb}P_x)_i^{j+1/2}$ reads:

$$-(J_{wb}P_x)_i^{j+1/2} = -(J_{wb,adv\ i}^{j+1/2} + J_{wb,dif\ i}^{j+1/2}) P_{x\ i}^{j+1/2}$$

For downward water flow ($q > 0$):

$$\begin{aligned}
 &= - \left(\frac{\ell}{P_{z=0\ i}^{j+1/2}} q_i^{j+1/2} c_i^{j+1/2} - \left(\epsilon D_{lb} \frac{\partial c_{lb}}{\partial z} \right)_{z=0\ i}^{j+1/2} \right) P_{x\ i}^{j+1/2} \\
 &= - \left(\frac{\ell}{P_{z=0\ i}^{j+1/2}} q_i^{j+1/2} c_i^{j+1/2} - \epsilon_i^{j+1/2} D_{lb\ i}^{j+1/2} \frac{(c_{lb\ i, k=1}^{j+1/2} - c_i^{j+1/2})}{1/2 \Delta z_1} \right) P_{x\ i}^{j+1/2} \quad (6.22a)
 \end{aligned}$$

For upward water flow ($q < 0$):

$$\begin{aligned}
 &= - \left(\frac{\ell}{P_{z=0\ i}^{j+1/2}} q_i^{j+1/2} c_{lb\ i, k=1}^{j+1/2} - \left(\epsilon D_{lb} \frac{\partial c_{lb}}{\partial z} \right)_{z=0\ i}^{j+1/2} \right) P_{x\ i}^{j+1/2} \\
 &= - \left(\frac{\ell}{P_{z=0\ i}^{j+1/2}} q_i^{j+1/2} c_{lb\ i, k=1}^{j+1/2} - \epsilon_i^{j+1/2} D_{lb\ i}^{j+1/2} \frac{(c_{lb\ i, k=1}^{j+1/2} - c_i^{j+1/2})}{1/2 \Delta z_1} \right) P_{x, i}^{j+1/2} \quad (6.22b)
 \end{aligned}$$

in which

$$P_{z=0\ i}^{j+1/2} = P_{z=0\ i}^j \quad (\text{constant in time})$$

$$q_i^{j+1/2} = \theta q_i^j + (1 - \theta) q_i^{j+1}$$

$$\epsilon_i^{j+1/2} = \epsilon_i^j \quad (\text{constant in time})$$

$$D_{lb\ i}^{j+1/2} = D_{lb\ i}^j \quad (\text{constant in time})$$

$$P_{x\ i}^{j+1/2} = P_{x\ i}^j (= P_{z=0\ i}^j) \quad (\text{constant in time})$$

$$c_{1b, i, k=1}^{j+1/2} = c_{1b, i, k=1}^j \quad (\text{see also section 6.9.})$$

The left-hand term of the conservation equation for the water layer, Eq. (3.6), was also approximated with the aid of the finite-difference method.

$$\left(\frac{\partial c^* A}{\partial t} \right)_i^j \approx \frac{c_i^{*j+1} A_i^{j+1} - c_i^{*j} A_i^j}{\Delta t} \quad (6.23)$$

The variable c^* , the total mass concentration of the substance in the water layer, may be written as a factor multiplied by c , the mass concentration in the water phase, according to Eq. (6.24):

$$c_i^{*j} = \left\{ 1 + \frac{DW P_{z=0} K_{mp}}{A} + SS K_{F,ss} \left(\frac{c_i^j}{c_{e,ss}} \right)^{n_{ss}-1} \right\} c_i^j \quad (6.24)$$

This factor depends on c , so c is calculated for known values of c^* in an iterative way (see section 6.8). Similarly the variable X_{ss} , the content of the substance sorbed to suspended solids, may be expressed as a function of c :

$$X_{ss, i}^j = \left\{ K_{F,ss} \left(\frac{c_i^j}{c_{e,ss}} \right)^{n_{ss}-1} \right\} c_i^j \quad (6.25)$$

Substituting these Eqs. (6.24) and (6.25) into the numerical approximations of the conservation equation (Eqs. (6.17) up to (6.23) inclusive) and rearrangement of all terms lead to the following equations:

$$\frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1 - \theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_c} \right)^{n_u-1} \right) +$$

$$\frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1 - \theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_c} \right)^{n_u-1} \right) +$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1 - \theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_c} \right)^{n_u-1} \right) +$$

$$\left(1 + \frac{DW P_{z=0}^{j+1}}{A_i^{j+1}} K_{mp} + ss K_F \left(\frac{c_i^{j+1}}{c_c} \right)^{n_u-1} \right) (A_i^{j+1} + kA_i^{j+\frac{1}{2}} (1 - \theta) \Delta t) +$$

$$k_{t,1} O_{x i}^{j+\frac{1}{2}} (1 - \theta) \Delta t + \frac{\ell}{P_{z=0}} q_i^{j+\frac{1}{2}} P_{x i}^j (1 - \theta) \Delta t +$$

$$\frac{2E_i^j D_{lb} P_{x i}^j}{\Delta z_1} (1 - \theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1 - \theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_c} \right)^{n_u-1} \right) -$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1 - \theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_c} \right)^{n_u-1} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{n_u-1} \right) +$$

$$\frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{n_u-1} \right)$$

$$RDD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta NWW_{i-\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{n_u-1} \right) -$$

$$\frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{n_u-1} \right) -$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i+1})} \theta \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{n_u-1} \right) -$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{n_u-1} \right) +$$

$$\left(1 + \frac{DW P_{z=0}}{A_i^j} K_{mp} + ss K_F \left(\frac{c_i^j}{c_e} \right)^{n_u-1} \right) (A_i^j - kA_i^{j+\frac{1}{2}} \theta \Delta t) -$$

$$k_{t,1} O_{x_i}^{j+1/2} \theta \Delta t - \frac{\ell}{P_{z=0}} q_i^{j+1/2} P_{x_i}^j \theta \Delta t -$$

$$\frac{2\varepsilon_i^j D_{lb} P_{x_i}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = - \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta NWW_{i+1/2} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n_s-1} \right) +$$

$$\frac{2A_{i+1/2}^{j+1/2} E_{x_{i+1/2}}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n_s-1} \right)$$

$$RV = k_{t,1} O_{x_i}^{j+1/2} \frac{c_{a_i}^{j+1/2}}{K_H} \Delta t + 2 \frac{\varepsilon_i^j D_{lb} P_{x_i}^j}{\Delta z_1} c_{lb, i, k=1}^j \Delta t$$

In the case of upward water flow ($q < 0$) in the sediment subsystem the right-hand vector RV contains an additional term:

$$- \frac{\ell}{P_{z=0}^j} q_i^{j+1/2} c_{lb, i, k=1}^j P_{x_i}^{j+1/2} \Delta t$$

and the penultimate terms containing $q_i^{j+1/2}$ in LDD and RDD disappear. Eq. (6.26) can be written as:

$$A.c^{j+1} = B.c^j + d \quad (6.26a)$$

in which A and B are tridiagonal matrices and d is a vector.

The solution of Eq. (6.26a) is:

$$c^{j+1} = A^{-1} (B.c^j + d) \quad (6.26b)$$

in which A^{-1} is the inversed tridiagonal matrix A (see also Press et al, 1986). The value of c_i^{j+1} at time $j+1$ can be calculated for known values of c_i^j at time j (cf. also section 6.8).

6.5 Boundary conditions and initial condition for water layer

6.5.1 Lower boundary condition

The boundary condition at the end of the part of the ditch considered reads:

$$J_{m+\frac{1}{2}} = \left(u (c + ssX_{ss}) - E_x \frac{\partial(c + ssX_{ss})}{\partial x} \right)_{m+\frac{1}{2}} \quad (6.27)$$

This was approximated by defining a certain number of additional grid points, *ebt* (end buffer, total number), behind the last grid point m in the ditch. The mass flux leaving the water layer at grid point *ebt* is calculated as:

$$J_{ebt+\frac{1}{2}} = \max[(u (c + ssX_{ss}))_{ebt+\frac{1}{2}}, 0] \quad (6.28)$$

This implies that an outgoing water flow ($u>0$) produces only advective transport leaving the water subsystem, while an ingoing water flow ($u<0$) causes water free of pesticide to enter the subsystem.

Eq. (6.28) implies a discontinuity from a physical point of view; changing the direction of flow at the end of the water subsystem, $u_{ebt+\frac{1}{2}}$, entails a discontinuous change of the flux $J_{ebt+\frac{1}{2}}$. This has been compensated for by extending the calculations for an additional number of grid points, *ebt*, called the end buffer, behind the last grid point in the ditch, m .

This buffer is a fictitious entity and serves only as a solution for the numerical problem created by the discontinuous boundary condition. Therefore, the substance can only be (re)distributed or transported in the buffer. The substance cannot originate or disappear in the buffer.

This means that certain processes can occur in the buffer behind the part of the ditch considered, while others cannot:

- advection and dispersion of substance do occur, but only advection occurs in the last half segment;
- sorption to suspended solids does not occur;
- sorption to macrophytes does not occur;
- there is no volatilisation to the atmosphere;
- there is no exchange with the sediment (so no advective or diffusive flux to or from the sediment) and

— there is no transformation.

In the end buffer the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* change. Annex 1 describes these elements for the grid points $m+1$ up to $ebt-1$ inclusive.

When the water flow in the ditch is positive, the lower boundary condition is defined by assuming an outgoing advective flux and by neglecting the dispersion in the last half segment. In this case the concentration $c_{ebt+1/2}^j$ cannot be calculated with the aid of the concentration c_{ebt+1}^j , so it has been assumed that $c_{ebt+1/2}^j = c_{ebt}^j$. Similarly, $X_{ss\ ebt+1/2}^j = X_{ss\ ebt}^j$. Annex 2 provides the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* for a positive water flow at grid point ebt .

When the water flow is negative, the incoming substance flux is zero. (There is again no dispersion in the last half segment.) This implies that $c_{ebt+1/2}^j = 0$ and $X_{ss\ ebt+1/2}^j = 0$. Annex 3 provides the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* for a positive water flow at grid point ebt .

6.5.2 Upper boundary condition

When the water flow in the ditch is positive, water free of pesticide enters the subsystem at the upper boundary. Hence, $(AJ)_{1/2} = 0$. The elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* are presented in Annex 4.

In situations of low hydraulic gradients, and where wind influence and pumping stations are present, it is quite common to encounter alternating flow directions in watercourses. When the water flow alternates between being positive and negative, the situation becomes similar to that of the lower boundary. This means that the boundary condition at the beginning of the part of the ditch considered equals:

$$J_{fbt+1/2} = \left(u (c + ssX_{ss}) - E_x \frac{\partial(c + ssX_{ss})}{\partial x} \right)_{fbt+1/2} \quad (6.29)$$

This situation was approximated by defining a certain number of additional grid points fbt in front of the first grid point in the ditch (front buffer, total number). The mass flux leaving the water subsystem is now calculated according to:

$$J_{1/2} = \min[(u (c + ssX_{ss}))_{1/2}, 0] \quad (6.30)$$

An outgoing, negative water flow causes only advective transport, while an ingoing, positive water flow means that water free of pesticide enters. The added front buffer with fbt segments compensates for this discontinuous boundary condition, analogously to the lower boundary.

For the grid points 2 up to fbt inclusive in the front buffer this means that:

— sorption to suspended solids does not occur;

- sorption to macrophytes does not occur;
- there is no volatilisation;
- no exchange with sediment occurs and
- there is no transformation.

For the situation of alternating positive and negative water flows, Annex 5 shows the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* for the grid points $fb = 2$ up to $fb = fbt$ inclusive in the front buffer.

When there is a negative water flow, the upper boundary condition consists of an outgoing advective flux; dispersion is neglected in the first half segment. The concentration $c_{fb=1/2}$ and $X_{ss, fb=1/2}$ are estimated by $c_{fb=1/2} = c_{fb=1}$ and $X_{ss, fb=1/2} = X_{ss, fb=1}$. Annex 6 presents the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* for the grid point $fb = 1$ in the front buffer in the case of a negative water flow.

When the water flow is positive, the incoming substance flux is zero. (Again there is no dispersion in the first half segment.) Hence, $c_{fb=1/2} = 0$ and $X_{ss, fb=1/2} = 0$. Annex 7 provides the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* at grid point $fb = 1$ for a positive water flow if a front buffer is present (i.e. for alternating flow directions).

The upper and lower boundary conditions are defined in such a way that dispersion in the outer half segments is neglected. This implies that the substance cannot enter the subsystem by dispersion across the boundaries if there is a higher concentration outside the buffers than inside. Therefore, the mass balance will continue to represent the full 100% of the substance originally added to the subsystem.

The final system of numerical equations can now be composed with the aid of section 6.5 and the upper and lower boundary conditions described above. This results in a matrix equation comparable to Eq. (6.26a) for each situation. This matrix equation can be solved for a specified initial condition.

6.5.3 Initial condition

Two examples of initial conditions are worked out here: an input from diffuse origins and a point-type input, both taking place at time $t = 0$.

The first example may be air drift deposition, e.g of y kg of active ingredient per m^2 at time $t = 0$. It is assumed that the pesticide mixes instantaneously with the water layer and that sorption equilibrium with suspended solids and macrophytes is instantaneous. At time $t = 0$ none of the substance is yet sorbed to the sediment. The load per running metre of ditch is $O_x \cdot y$ kg a.i./m' dissolved in a volume of A m^3 . The substance concentration at time $t = 0$ is:

$$c^* = \frac{O_x \cdot y}{A} \quad (6.31)$$

The initial total mass concentration in the buffer(s) equals zero.

In the second example, there is a point-type input in one segment of the water subsystem, represented by e.g. y kg a.i. The total mass concentration c^* in this segment at time $t = 0$ equals:

$$c^* = \frac{y}{A \cdot \Delta x} \quad (6.32)$$

The initial mass concentration in all other segments equals zero.

Given these initial concentrations, the system of numerical equations representing the mass conservation equation of the water subsystem, Eq. (6.26), can now be solved.

6.6 Sediment

The right-hand terms of the mass conservation equation of the sediment, Eq. (3.12), are approximated by the following generalised finite-difference equation, which evaluates the spatial derivatives and the concentrations at time $j+1/2$, as defined again by Eq.(6.2).

$$\left(- \frac{\partial(PJ_{lb})}{\partial z} \right)_k^{j+1/2} - k_b (c_b^* P)_k^{j+1/2} \approx \frac{(PJ_{lb})_{k-1/2}^{j+1/2} - (PJ_{lb})_{k+1/2}^{j+1/2}}{\Delta z_k} - k_b \cdot c_b^* P_k^{j+1/2} \quad (6.33)$$

The term $(PJ_{lb})_{k-1/2}^{j+1/2}$ for values of k from 2 to n is given by:

$$\begin{aligned} (PJ_{lb})_{k-1/2}^{j+1/2} &= \left[\frac{P \cdot \ell}{P} \cdot q \cdot c_{lb} - P \varepsilon (E_{lb} + D_{lb}) \frac{\partial c_{lb}}{\partial z} \right]_{k-1/2}^{j+1/2} \\ &= \ell \cdot q^{j+1/2} \cdot c_{lb \ k-1/2}^{j+1/2} - P_{k-1/2}^{j+1/2} \varepsilon_{k-1/2}^{j+1/2} (E_{lb \ k-1/2}^{j+1/2} + D_{lb \ k-1/2}^{j+1/2}) \left(\frac{c_{lb \ k}^{j+1/2} - c_{lb \ k-1}^{j+1/2}}{1/2 \Delta z_k + 1/2 \Delta z_{k-1}} \right) \end{aligned}$$

$$= \ell q^{j+1/2} c_{lb\ k-1/2}^{j+1/2} - P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) \left(\frac{c_{lb\ k}^{j+1/2} - c_{lb\ k-1}^{j+1/2}}{1/2\Delta z_k + 1/2\Delta z_{k-1}} \right) \quad (6.34)$$

in which

$$q^{j+1/2} = \theta q^j + (1 - \theta) q^{j+1}$$

$$c_{lb\ k-1/2}^{j+1/2} = \theta c_{lb\ k-1/2}^j + (1 - \theta) c_{lb\ k-1/2}^{j+1} = \theta (1 - NWB_{k-1/2}) c_{lb\ k-1}^j +$$

$$\theta NWB_{k-1/2} c_{lb\ k}^j + (1 - \theta) (1 - NWB_{k-1/2}) c_{lb\ k-1}^{j+1} + (1 - \theta) NWB_{k-1/2} c_{lb\ k}^{j+1}$$

$$P_{k-1/2}^j = b + 2 \sum_{p=1}^{k-1} \Delta z_p \tan(1/2\beta) + 2 (h_w + \sum_{p=1}^{k-1} \Delta z_p) \sqrt{s_1^2 + 1}$$

$\epsilon_{k-1/2}^j$ is given as a function of depth

$$E_{lb\ k-1/2}^{j+1/2} = \theta E_{lb\ k-1/2}^j + (1 - \theta) E_{lb\ k-1/2}^{j+1}$$

$$E_{lb\ k-1/2}^j = (L_{dis} |w|)_{k-1/2}^j$$

$$D_{lb\ k-1/2}^j = (\lambda D_w)_{k-1/2}^j$$

$$c_{lb\ k}^{j+1/2} = \theta c_{lb\ k}^j + (1 - \theta) c_{lb\ k}^{j+1}$$

$$c_{lb\ k-1}^{j+1/2} = \theta c_{lb\ k-1}^j + (1 - \theta) c_{lb\ k-1}^{j+1}$$

P , ϵ and D_{lb} are constant in time, so their values at time $j+1/2$ equal their values at time j .

Similarly, $(PJ_{lb})_{k+1/2}^{j+1/2}$ for values of k from 1 to $n-1$ reads:

$$(PJ_{lb})_{k+1/2}^{j+1/2} = \left[\frac{P \cdot \ell}{P} \cdot q \cdot c_{lb} - P \varepsilon (E_{lb} + D_{lb}) \frac{\partial c_{lb}}{\partial z} \right]_{k+1/2}^{j+1/2} =$$

$$\ell \cdot q^{j+1/2} \cdot c_{lb \ k+1/2}^{j+1/2} - P_{k+1/2}^j \varepsilon_{k+1/2}^j (E_{lb \ k+1/2}^{j+1/2} + D_{lb \ k+1/2}^j) \left(\frac{c_{lb \ k+1}^{j+1/2} - c_{lb \ k}^{j+1/2}}{1/2 \Delta z_{k+1} + 1/2 \Delta z_k} \right) \quad (6.35)$$

in which

$$c_{lb \ k+1/2}^{j+1/2} = \theta c_{lb \ k+1/2}^j + (1 - \theta) c_{lb \ k+1/2}^{j+1} = \theta (1 - NWB_{k+1/2}) c_{lb \ k}^j + \theta NWB_{k+1/2} c_{lb \ k+1}^j +$$

$$(1 - \theta) (1 - NWB_{k+1/2}) c_{lb \ k}^{j+1} + (1 - \theta) NWB_{k+1/2} c_{lb \ k+1}^{j+1}$$

$$P_{k+1/2}^j = b + 2 \sum_{p=1}^k \Delta z_p \tan(1/2\beta) + 2 (h_w + \sum_{p=1}^k \Delta z_p) \sqrt{s_1^2 + 1}$$

$$E_{lb \ k+1/2}^{j+1/2} = \theta E_{lb \ k+1/2}^j + (1 - \theta) E_{lb \ k+1/2}^{j+1}$$

$$c_{lb \ k+1}^{j+1/2} = \theta c_{lb \ k+1}^j + (1 - \theta) c_{lb \ k+1}^{j+1}$$

The term $(-k_b \cdot c_{b \ k}^* \cdot P_k^{j+1/2})$ equals:

$$- k_b c_{b \ k}^{* \ j+1/2} P_k^{j+1/2} = - k_b c_{b \ k}^{* \ j+1/2} P_k^j \quad (6.36)$$

with

$$c_{b \ k}^{* \ j+1/2} = \theta c_{b \ k}^{* \ j} + (1 - \theta) c_{b \ k}^{* \ j+1}$$

The left-hand term of the mass conservation equation for the sediment, Eq. (3.12), is also approximated by means of the finite-difference method.

$$\frac{P_{k-\frac{1}{2}}^j \varepsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) (1 - \theta)\Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$LDD = \frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} (1 - \theta) (1 - NWB_{k+\frac{1}{2}}) \frac{\Delta t}{P_k^j} sofd_{k+\frac{1}{2}} +$$

$$\frac{P_{k+\frac{1}{2}}^j \varepsilon_{k+\frac{1}{2}}^j (E_{lb\ k+\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}^j) (1 - \theta)\Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j} -$$

$$\frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} (1 - \theta) NWB_{k-\frac{1}{2}} \frac{\Delta t}{P_k^j} sofd_{k-\frac{1}{2}} +$$

$$\frac{P_{k-\frac{1}{2}}^j \varepsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) (1 - \theta)\Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j} +$$

$$\left(\varepsilon_k^j + \rho_{b\ k}^j K_F \left(\frac{c_{lb\ k}^{j-1}}{c_e} \right)^{n_{wb}-1} \right) \left(1 + k_b P_k^j (1 - \theta) \frac{\Delta t}{P_k^j} \right)$$

$$LBD = \frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} (1 - \theta) NWB_{k+\frac{1}{2}} \frac{\Delta t}{P_k^j} sofd_{k+\frac{1}{2}} -$$

$$\frac{P_{k+\frac{1}{2}}^j \varepsilon_{k+\frac{1}{2}}^j (E_{lb\ k+\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}^j) (1 - \theta)\Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$ROD = \frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} \theta (1 - NWB_{k-\frac{1}{2}}) \frac{\Delta t}{P_k^j} sofd_{k-\frac{1}{2}} +$$

$$\frac{P_{k-\frac{1}{2}}^j \epsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$RDD = - \frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} \theta (1 - NWB_{k+\frac{1}{2}}) \frac{\Delta t}{P_k^j} sofd_{k+\frac{1}{2}} -$$

$$\frac{P_{k+\frac{1}{2}}^j \epsilon_{k+\frac{1}{2}}^j (E_{lb\ k+\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j} +$$

$$\frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} \theta NWB_{k-\frac{1}{2}} \frac{\Delta t}{P_k^j} sofd_{k-\frac{1}{2}} -$$

$$\frac{P_{k-\frac{1}{2}}^j \epsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j} +$$

$$\left(\epsilon_k^j + \rho_{bk}^j K_F \left(\frac{c_{lb\ k}^j}{c_e} \right)^{n_{wb}-1} \right) \left(1 - k_b P_k^j \theta \frac{\Delta t}{P_k^j} \right)$$

$$RBD = - \frac{\varrho q^{j+\frac{1}{2}}}{\Delta z_k} \theta NWB_{k+\frac{1}{2}} \frac{\Delta t}{P_k^j} sofd_{k+\frac{1}{2}} +$$

$$\frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$RV = 0$, except at the last grid point *ebbt* (see section 6.7).

As was done for the water layer, Eq. (6.39) can be solved for c_{lb}^{j+1} by inverting the left-hand tridiagonal matrix (see Press et al, 1986).

6.7 Boundary conditions and initial condition for sediment

6.7.1 Upper boundary condition

For a positive as well as for a negative advection flow q , $c_{lb} = c_i$ at the location $z = 0$ in the sediment, so the concentration of the substance in the liquid phase of the sediment equals the concentration of the substance dissolved in the water layer in segment number i . The conservation equation for the sediment subsystem, Eq. (3.12), now reads:

$$\begin{aligned} \left(P \frac{\partial c_b^*}{\partial t} \right)_1 &= - \left(\frac{\partial(PJ_{lb})}{\partial z} \right)_1^{j+1/2} - (k_b c_b^* P)_1^{j+1/2} \\ &\approx \frac{(PJ_{lb})_{1/2}^{j+1/2} - (PJ_{lb})_{1/2}^{j+1/2}}{\Delta z_1} - k_b c_b^* P_1^j \end{aligned} \quad (3.12a)$$

The term $(PJ_{lb})_{1/2}^{j+1/2}$ stands for:

$$(PJ_{lb})_{1/2}^{j+1/2} = \varrho q^{j+1/2} c_{lb\ 1/2}^{j+1/2} - P_{1/2}^j \varepsilon_{1/2}^j D_{lb\ 1/2}^j \frac{c_{lb\ 1}^{j+1/2} - c_{lb\ 1/2}^{j+1/2}}{1/2 \Delta z_1} \quad (6.34a)$$

$$\text{if } q \geq 0: \quad = \varrho q^{j+1/2} c_i^{j+1/2} - P_{1/2}^j \varepsilon_{1/2}^j D_{lb\ 1/2}^j \frac{c_{lb\ 1}^{j+1/2} - c_i^{j+1/2}}{1/2 \Delta z_1}$$

$$\text{if } q < 0: \quad = \varrho q^{j+1/2} c_{lb\ 1}^{j+1/2} - P_{1/2}^j \varepsilon_{1/2}^j D_{lb\ 1/2}^j \frac{c_{lb\ 1}^{j+1/2} - c_i^{j+1/2}}{1/2 \Delta z_1}$$

in which the variables are defined by Eq. (6.34). The terms $(PJ_{lb})_{1/2}^{j+1/2}$ and $(k_b c_b^* P)_1^{j+1/2}$ were defined by the Eqs. (6.35) and (6.36). Substitution and rearrangement of the numerical approximations of the conservation equation, Eq. (3.12), for $k = 1$ results in the upper boundary condition (Annex 8).

6.7.2 Lower boundary condition

For a downward advection flow, the boundary condition at the lower end of the sediment subsystem considered is:

$$PJ_{lb} = \ell q c_{lb} - \varepsilon P (E_{lb} + D_{lb}) \frac{\partial c_{lb}}{\partial z} \quad (6.40)$$

In the case of an upward advection flow the boundary condition reads:

$$PJ_{lb} = \ell q c_{lot} \quad (6.41)$$

in which c_{lot} is the concentration in the water flowing upward, i.e. seeping from the neighbouring lot into the ditch bottom. This concentration needs to be prescribed as a function of time. This boundary condition is also discontinuous, which is why an additional number of grid points *ebbt* (end buffer sediment [water bottom], total number) have been defined, constituting a buffer behind the last grid point n in the sediment.

The mass flux flowing out of the buffer at grid point *ebbt* equals:

$$\text{if } q > 0 : \quad PJ_{ebbt+\frac{1}{2}} = \ell q c_{lb}$$

(i.e. diffusion and dispersion are neglected), and

$$\text{if } q < 0 : \quad PJ_{lb} = \ell q c_{lot} \quad (6.42)$$

As in the case of the water layer, the buffer only serves the purpose of solving the numerical problem caused by the discontinuous boundary condition. Hence, the substance is exclusively transported in the buffer; no substance originates or disappears in the buffer. This implies that certain processes can occur in the buffer, while others cannot:

- advection, dispersion and diffusion of the substance do occur, but only advection occurs in the last half segment;
- no sorption (to the solid phase of the sediment) occurs and
- there can be no transformation of the substance.

Annex 9 presents the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD* and *RBD* for the grid points $n+1$ up to *ebbt-1* inclusive.

When the water flow is downward, the lower boundary condition is described by an outgoing advection flux; the diffusion and dispersion in the last half segment are neglected. (This implies that the substance cannot enter by diffusion or dispersion and the mass balance will continue to represent 100% of the mass originally added.) The concentration at the boundary, c_{lb}^i , is determined by:

$$c_{lb\ ebbt+\frac{1}{2}}^j = c_{lb\ ebbt}^j \quad (6.43)$$

Annex 10 defines the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD* and *RBD* at the grid point *ebbt*.

When the water flow is upward, the concentration in the liquid phase at the lower boundary is externally determined. So:

$$\begin{aligned} (PJ_{lb})_{ebbt+\frac{1}{2}}^{j+\frac{1}{2}} &= \ell q^{j+\frac{1}{2}} c_{lb\ lot}^{j+\frac{1}{2}} \\ &= \ell q^{j+\frac{1}{2}} \theta c_{lb\ lot}^j + \ell q^{j+\frac{1}{2}} (1 - \theta) c_{lb\ lot}^{j+1} \end{aligned} \quad (6.44)$$

Annex 11 shows the elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* at the last grid point *ebbt*. The last element of the right-hand vector *RV* is found to be not equal to zero in this case.

6.7.3 Initial condition

There are two options for the initial condition of the sediment subsystem. In the first option, the sediment is free of pesticide at time $t = 0$. This implies:

$$c_b^* = 0 \quad (6.45)$$

In the second option, a certain mass of pesticide is located in the sediment; this can be described as a function of depth. At time $t = 0$:

$$c_b^* = f(z) \quad (6.46)$$

in which

$f(z)$ = mass concentration of the substance in sediment as a function of depth at time $t = 0$ (M.L⁻³).

Given the initial concentration, and selecting the right upper and lower boundary conditions, the system of numerical equations, Eq. (6.39), can now be solved.

6.8 Iterative calculations caused by the Freundlich equation for sorption

At two points, iterative calculations are necessary to solve the equations, because the non-linear Freundlich equation was used to describe sorption to the suspended solids and to the solid sediment material. The first point concerns the calculation

of the initial pesticide concentration in the water phase; the second concerns the solution of the matrix equation for the water layer or the sediment.

The conservation equation for the water layer, Eq. (3.6), reads:

$$\frac{\partial(c^* A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} - k(c^* A) + J_{wa} O_x - J_{wb} P_x \quad (3.6)$$

To approximate the fluxes of the substance in the right-hand term of Eq. (3.6), values of c have to be derived from known values of c^* . Combining Eqs. (4.1) and (4.2) results in:

$$c^* = c + \frac{DW P_{z=0}}{A} K_{mp} c + ss K_{F,ss} c_{e,ss} \left(\frac{c}{c_{e,ss}} \right)^{n_s} \quad (6.47)$$

Eq. (6.47) shows that it is impossible to derive values of c from values of c^* in an explicit way. Rearranging Eq. (6.47) yields an implicit equation in c :

$$c = \frac{c^*}{1 + \frac{DW P_{z=0}}{A} K_{mp} + ss K_{F,ss} \left(\frac{c}{c_{e,ss}} \right)^{n_s-1}} \quad (6.47a)$$

For many pesticides, the value of n_s-1 is small (the value of n in sorption studies for soils is often near 0.9), so the right-hand term changes little with a change in c . Using an initial estimation for c in the denominator (and known values for the other variables of the right-hand term), the final value of c can be rapidly found in an iterative manner. The value of c^* at time $t = 0$ serves as the initial estimation for c , yielding a value for c in the left-hand term. This value is then used in the denominator of the right-hand term, etc. In this way, c can be approximated from known values of c^* .

An analogous calculation leads to c_{lb} from known c_b^* for the sediment subsystem. Combining Eqs. (4.17) and (4.18) yields:

$$c_b^* = \epsilon c_{lb} + \rho_b K_{F,wb} c_{e,wb} \left(\frac{c_{lb}}{c_{e,wb}} \right)^{n_{wb}} \quad (6.48)$$

Rearranging Eq. (6.48) results in:

conservation equations simultaneously; instead, the simpler option of solving one equation after the other has been chosen. The next two sections explain this solution method in more detail.

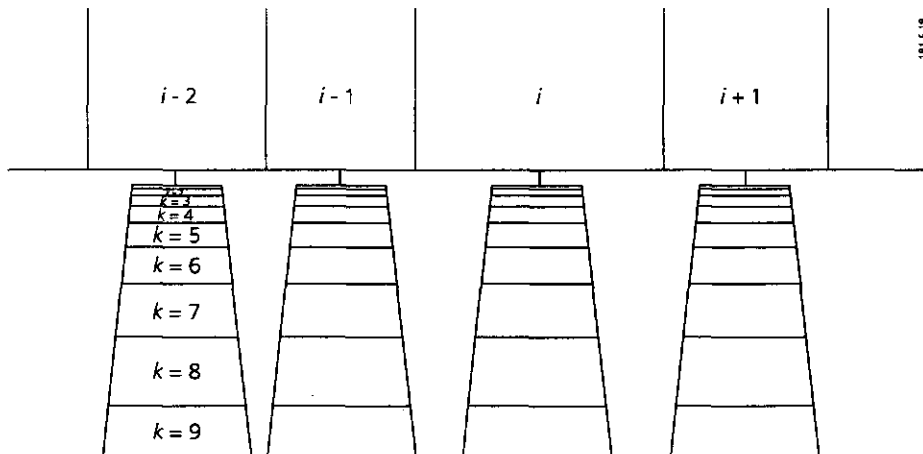


Fig. 18 Detail of the TOXSWA field ditch system, demonstrating the concept of one water subsystem coupled to many sediment subsystems

To solve the conservation equation for the water subsystem, its right-hand terms are evaluated at time $j+1/2$ in section 6.4 (see Eq. (6.17)). For the exchange term with the sediment, $-(J_{wb} \cdot P_x)_i^{j+1/2}$, this yields a term $c_{lb, i, k=1}^{j+1/2}$ in Eq. (6.22). This term is approximated as:

$$c_{lb, i, k=1}^{j+1/2} = c_{lb, i, k=1}^j \quad (6.50)$$

The value of c_{lb} at time $j+1/2$ is thus approximated by its value at time j . This means that the value $c_{lb}^{j+1/2}$ need not yet be known when the conservation equation of the water subsystem is being solved. Hence, it is not necessary to solve simultaneously the conservation equations of the water and sediment subsystems.

Approximating the upper boundary condition of the sediment subsystem yields the terms $c_i^{j+1/2}$ and $c_{lb, 1}^{j+1/2}$ in section 6.7.1 (see Eq. (6.34a)). Hence, the terms c and c_{lb} at times j and $j+1$ appear in the approximation of the upper boundary condition (see also Annex 8). Calculating c_{lb}^{j+1} at time j in the sediment subsystem requires that the value of c_i^{j+1} is known. This would imply that both conservation equations, that for the water and that for the sediment subsystem, need to be solved simultaneously. However, this is a complex procedure. It is simpler to have the solution of the conservation equation of the water subsystem precede that of the sediment subsystem by one time step. This implies that at time j in the sediment subsystem, the concentration c_i in the water layer is known at time j as well as at time $j+1$. The system of equations for the sediment subsystem, Eqs. (6.39), can now easily be solved.

7 Verification

7.1 Introduction

Annexes 12 and 13 show an outline of the computer program of the TOXSWA model. Annex 19 presents the source code of the TOXSWA program, version 1.0, including a guide to the vocabulary used. After a computer program has been written a verification is required. Verification is defined as the examination of the numerical technique in the computer model to ascertain that it truly represents the mathematical model and that there are no inherent numerical problems in obtaining a solution. This also implies a check on errors in the code (programming bugs). (Leaching Modelling Workgroup FOCUS, 1995.) For finite-difference approximations, the concept of verification may be worked out using the three notions of convergence, stability and consistency (Bear and Verruijt, 1987; Lapidus and Pinder, 1982). The condition of convergence states that when the finite-difference grid is refined the truncation errors go to zero. Stability concerns the unstable growth or stable decay of errors in the arithmetic operations needed to solve the finite-difference equations, so it has to do with the boundedness of all perturbations in a computed solution. Consistency is the requirement that when the finite-difference grid is refined, the truncation errors go to zero, but moreover that the finite-difference model approximates the partial differential equation desired and not some other partial differential equation.

The three notions of convergence, stability and consistency lead to the formulation of the following conditions for verification.

Convergence is worked out as follows:

- 1.⁹ A solution obtained with the same numerical solution scheme but time and space steps e.g. 10 times smaller, should differ only slightly from the solution obtained with the original time and space steps. This implies that in this case the truncation errors are so small that the numerical equations provide nearly the same solution.
- 2.¹⁰ The solution obtained should not vary with the combinations of the numerical weight factors selected for space β and time θ ; this means that the solution should be identical whether it has been calculated with the aid of a forward, backward or central difference scheme, or with the aid of an implicit or explicit difference scheme. (The numerical dispersion in TOXSWA, which depends on

⁹ This condition is not yet tested for version 1.0 of the TOXSWA model.

¹⁰ These conditions have not been tested for version 1.0 of the TOXSWA model. Version 1.0 allows only the explicit central difference method to be selected to solve the conservation equations, and calculations are performed for a constant, rather than variable, wetted area A .

the numerical solution scheme selected, has been estimated and the physical dispersion has been corrected for this.)

The stability condition has been translated into the following positivity conditions:

3. The solution of the matrix equations, Eq. (6.26) for the water layer and Eq. (6.39) for the sediment, needs to be stable, i.e. the errors in the arithmetic operations needed to solve the finite-difference equations should be bounded. Moreover, as it concerns concentrations, the solution should be positive. It can be proven that for specific conditions for the *LOD*, *LDD*, *LBD*, *ROD*, *RDD* and *RBD* elements of the matrix equations and due to the specific structure of these elements (described in Annex 14 at the illustration for the three-dimensional case), the solution will be positive. Positivity can be shown to be a stricter condition than the stability condition for the solution.

The following four checks are proposed for the consistency condition:

4. Letting the finite differences approach zero should result in the numerical equations reducing to the original partial differential equations.

The numerical equations in the computer program should truly represent the mathematical model, i.e. the mass balances. This means that the consistency condition also implies the following:

5. The mass of pesticide introduced (100%) should be traceable at every moment with an error of less than 0.1% (this means that the mass balances tally and that no pesticide is lost during the calculation process by any other than those being modelled, e.g. transformation or volatilisation).

Furthermore consistency is usually checked by verifying the numerical procedure against a variety of analytical solutions (Bear and Verruijt, 1987). Two situations are proposed for this check:

6. An analytical solution of the advection-dispersion equation for constant h and A for a system with transformation, volatilisation and convective/dispersive downward seepage responding to a Dirac delta function-type input, as described in Jury and Roth (1990). (Volatilisation and convective/dispersive downward seepage are, from a mathematical point of view, equivalent to transformation.) This solution should correspond with that obtained with the numerical solution scheme. The same type of analytical solution can also be applied to the sediment subsystem. So, the water and sediment subsystem can be checked independently.

- 7.¹¹ A situation like that described under point 6., but now with a wetted surface A varying according to: $A = \varphi \cdot e^{\gamma x + \xi t}$. This situation is relevant because the TOXSWA model can deal with water levels and wetted perimeters varying in time and space. This special description for A again results in a convection-dispersion type of equation, but the parameters of the convection and transformation terms include γ and/or ξ .

7.2 Stability condition

7.2.1 Description of the positivity for the water layer

The stability condition was worked out as stated in criterium 3 of section 7.1.

5. The solution of the matrix equations, Eq. (6.26) for the water layer and Eq. (6.39) for the sediment, needs to be stable, i.e. the errors in the arithmetic operations needed to solve the finite-difference equations should be bounded. Moreover, as it concerns concentrations, the solution should be positive. It can be proven that for specific conditions for the *LOD*, *LDD*, *LBD*, *ROD*, *RDD* and *RBD* elements of the matrix equations and due to the specific structure of these elements (described in Annex 14 at the illustration for the three-dimensional case), the solution will be positive. Positivity can be shown to be a stricter condition than the stability condition for the solution.

The solution of the system of equations, solving the conservation equation for the water subsystem reads:

$$c^{j+1} = A^{-1} (B \cdot c^j + d) \quad (6.26a)$$

The vector d is zero or positive and this implies that c^{j+1} is positive if the product of the inversed matrix A^{-1} and the matrix B is positive, which is true when A^{-1} and B are both positive matrices. A positive matrix is defined as a matrix that has no negative elements. So the tridiagonal matrix B is positive if:

$$\begin{aligned} RBD &\geq 0, \\ RDD &> 0 \text{ and} \\ ROD &\geq 0. \end{aligned}$$

The matrix A^{-1} is positive if:

$$\begin{aligned} LBD &\leq 0, \\ LDD &> 0 \text{ and} \\ LOD &\leq 0. \end{aligned}$$

¹¹ These conditions have not been tested for version 1.0 of the TOXSWA model. Version 1.0 allows only the explicit central difference method to be selected to solve the conservation equations and calculations are performed for a constant, rather than variable, wetted area A .

Annex 14 gives the outline of the proof that $A^{-1} \cdot B$ is positive and stable for the two tridiagonal matrices A and B when the elements RBD , RDD , ROD , LBD , LDD and LOD fulfill the conditions mentioned above. A positive matrix is called stable if the dominant eigenvalue of the matrix is smaller than 1. The outline is given for a system with constant segment size Δx , flow velocity u and water depth h and a linear sorption isotherm describing sorption to suspended solids. For this system the elements LOD , LDD , LBD , ROD , RDD and RBD are exactly defined in section 7.2.2.

For the sediment subsystem the positivity and its proof may be demonstrated analogously.

7.2.2 Restrictions resulting from requirements for positivity for the water layer

The conditions for which the solution of the matrix equation Eq.(6.26) is positive for the water subsystem have been worked out in an advice about the possible time and space steps for which this equation can be solved. The advice is based upon a slightly simplified situation.

When sorption to suspended solids is described with a linear isotherm and when the segment size Δx , the flow velocity u , the water depth h and wetted area A are constant, the elements LOD , LDD , LBD , ROD , RDD , RBD and RV for the water subsystem read:

$$LOD = - (1-\theta) (1-\beta) \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) - (1 - \theta) \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$LDD = (1-\theta) (1 - 2\beta) \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) + 2(1 - \theta) \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$+ (1 + k (1 - \theta) \Delta t) + \left(1 + \frac{DW \cdot P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right)$$

$$+ k_{t,1} \frac{O_x}{A} (1 - \theta) \Delta t + \frac{\ell}{A P_{z=0}} q P_x (1 - \theta) \Delta t + \frac{2\varepsilon D_{lb} P_x}{A \Delta z_1} (1 - \theta) \Delta t$$

$$LBD = (1-\theta) \beta \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) - (1 - \theta) \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$ROD = \theta (1-\beta) \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) + \theta \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$RDD = - \theta (1 - 2\beta) \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) - 2 \theta \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$+ (1 - k \theta \Delta t) + \left(1 + \frac{DW \cdot P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right)$$

$$- \left(k_{t,1} \frac{O_x}{A} \theta \Delta t + \frac{\ell}{A P_{z=0}} q P_x \theta \Delta t + \frac{2\varepsilon D_{lb} P_x}{A \Delta z_1} \theta \Delta t \right)$$

$$RBD = - \theta \beta \frac{u \Delta t}{\Delta x} (1 + ss K_{L,ss}) + \theta \frac{E \Delta t}{\Delta x^2} (1 + ss K_{L,ss})$$

$$RV = k_{t,1} \frac{O_x}{A} \frac{c_a}{K_H} \Delta t + \frac{2\varepsilon D_{lb} P_x}{A \Delta z_1} c_{lb, k=1} \Delta t \quad (7.1)$$

with

$K_{L,ss}$ = slope of (linear) sorption isotherm of suspended solids ($L^3.M^{-1}$).

With

$$M = 1 + \frac{DW \cdot P_{z=0}}{A} K_{mp} + ss K_{L,ss} \quad (M > 0)$$

$$F = k_{t,1} \frac{O_x}{A} + \frac{\ell}{A P_{z=0}} q P_x + \frac{2\varepsilon D_{lb} P_x}{A \Delta z_1} \quad (\text{Assumption}^{12}: F > 0)$$

$$G^{13} = (1 - 2\beta) \frac{u}{\Delta x} (1 + ss K_{L,ss}) + 2 \frac{E}{\Delta x^2} (1 + ss K_{L,ss})$$

$$+ k M + F$$

the element *LDD* may be rewritten as:

$$LDD = G (1 - \theta) \Delta t + M \quad (7.2)$$

and the element *RDD* may be rewritten as:

$$RDD = - G \theta \Delta t + M \quad (7.3)$$

The positivity conditions for the off-diagonal elements of the two tridiagonal matrices lead to an advice about the size of the space step, Δx .

The condition $LBD \leq 0$ (or $RBD \geq 0$) leads to:

$$\Delta x \leq \frac{E}{\beta u} \quad \text{for } u > 0, E > 0 \text{ and } \beta \neq 0 \quad (7.4)$$

The condition $LOD \leq 0$ (or $ROD \geq 0$) leads to:

$$\Delta x \leq - \frac{E}{(1 - \beta) u} \quad \text{for } u < 0, E > 0 \text{ and } \beta \neq 1 \quad (7.5)$$

¹² In fact the term F should be such that $2E/\Delta x^2(1+ss K_{L,ss})+kM+F$ is positive; F can become negative if q , the seepage, is negative, but this only happens for unrealistic high (negative) values of q ($|q|$ in the order of $m.d^{-1}$). It has been assumed that the dispersion E is positive. It may be possible, however, that the E_{cal} with which TOXSWA 1.0 calculates is negative. In that case an additional condition for Δx and Δt should be added.

¹³ The sign of G depends on the values of u and β .

If $u = 0$ the conditions described above do not lead to restrictions for Δx . When a positive flow direction is combined with $\beta = 0$ or when a negative flow direction is combined with $\beta = 1$, there is also no restriction for Δx .

The positivity conditions for the diagonal elements of the two tridiagonal matrices lead to an advice about the size of the time step, Δt .

The condition $LDD > 0$ leads to:

$$G (1 - \theta) \Delta t + M > 0 \quad (7.6)$$

corresponding to:

$$1. \quad G > 0 : \Delta t > - \frac{M}{G (1 - \theta)} \quad (7.7)$$

In this case no restriction for the selection of (an always positive) Δt is found.

$$2. \quad G < 0 : \Delta t < - \frac{M}{G (1 - \theta)} \quad (7.8)$$

and $G < 0$ corresponds to:

$$(i) \quad \beta > \frac{1}{2} \text{ and } u > H > 0 \quad (7.9)$$

or

$$(ii) \quad \beta < \frac{1}{2} \text{ and } u < H < 0 \quad (7.10)$$

with

$$H = \frac{- \{ 2E (1 + ss K_{L,ss}) + k M \Delta x^2 + F \Delta x^2 \}}{(1 - 2\beta) (1 + ss K_{L,ss}) \Delta x}$$

If $\beta = \frac{1}{2}$ or $u = 0$ $G (1 - \theta)$ is always positive, so the condition $LDD > 0$ does not lead to a restriction for Δt . If $\theta = 1$ $G (1 - \theta)$ equals zero, so the condition $LDD > 0$ is always true and only the condition $RDD > 0$ poses restrictions to Δt .

The condition $RDD > 0$ leads to:

$$- G \theta \Delta t + M > 0 \quad (7.11)$$

corresponding to:

$$1. \quad G < 0 : \Delta t > \frac{M}{G \theta} \quad (7.12)$$

In this case no restriction for the selection of Δt is found.

$$2. \quad G > 0 : \Delta t < \frac{M}{G \theta} \quad (7.13)$$

and $G > 0$ corresponds to:

$$(i) \quad \beta < \frac{1}{2} \text{ and } H < u \quad (7.14)$$

H being negative

$$(ii) \quad \beta > \frac{1}{2} \text{ and } H > u \quad (7.15)$$

H being positive

with H as defined before.

When $\beta = \frac{1}{2}$ or $u = 0$ $G \theta$ is always positive, so the condition RDD does lead to a restriction for Δt . Only when $\theta = 0$ $G \theta$ equals zero, the condition $RDD > 0$ is always true but now the condition $LDD > 0$ may lead to restriction for Δt .

Figure 19 summarizes the results for the restrictions imposed on the time step Δt .

It appears that in the case $\theta = 0$ and $|H| > |u|$ (i.e. $-|H| < u < |H|$) the condition $LDD > 0$ is not fulfilled, so the solution to matrix Eq.(6.26) is not positive. So, the matrix Eq. (6.26) has only a positive solution if Δx is selected such that $|H| < |u|$.

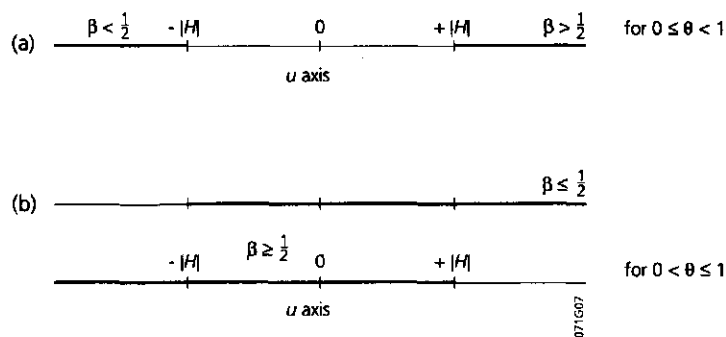


Fig. 19 Outline of the domain (u, β) for which the condition $LDD > 0$ (part(a)) or $RDD > 0$ (part(b)) result in a restriction for the time step, Δt

7.2.3 Restrictions resulting from requirements for positivity for the sediment

For the sediment subsystem an advice about the possible time and space steps is also worked out. The solution of matrix Eq.(6.39) is positive for these time and space steps.

When sorption to the bottom material is described with a linear isotherm and when the segment size Δz and the seepage q are constant, the elements LOD , LDD , LBD , ROD , RDD and RBD for the sediment subsystem read:

$$\begin{aligned}
 LOD &= - \frac{\varrho q}{\Delta z} (1 - \theta) (1 - \beta) \frac{\Delta t}{P_k} sofd \\
 &\quad - \frac{P_{k-1/2} \varepsilon_{k-1/2} (E_{lb\ k-1/2} + D_{lb\ k-1/2}) (1 - \theta) \Delta t}{\Delta z^2 P_k} \\
 LDD &= \frac{\varrho q}{\Delta z} (1 - \theta) (1 - 2\beta) \frac{\Delta t}{P_k} sofd \\
 &\quad + \frac{P_{k+1/2} \varepsilon_{k+1/2} (E_{lb\ k+1/2} + D_{lb\ k+1/2}) (1 - \theta) \Delta t}{\Delta z^2 P_k} \\
 &\quad + \frac{P_{k-1/2} \varepsilon_{k-1/2} (E_{lb\ k-1/2} + D_{lb\ k-1/2}) (1 - \theta) \Delta t}{\Delta z^2 P_k} \\
 &\quad + (\varepsilon_k + \rho_{b,k} K_{L,b}) (1 + k_b (1 - \theta) \Delta t)
 \end{aligned}$$

$$LBD = \frac{\varrho q}{\Delta z} (1 - \theta) \beta \frac{\Delta t}{P_k} sofd$$

$$\begin{aligned}
& - \frac{P_{k+\frac{1}{2}} \epsilon_{k+\frac{1}{2}} (E_{lb\ k+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}) (1 - \theta) \Delta t}{\Delta z^2 P_k} \\
ROD &= \frac{\ell q}{\Delta z} \theta (1 - \beta) \frac{\Delta t}{P_k} sofd + \frac{P_{k-\frac{1}{2}} \epsilon_{k-\frac{1}{2}} (E_{lb\ k-\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}) \theta \Delta t}{\Delta z^2 P_k} \\
RDD &= - \frac{\ell q}{\Delta z} \theta (1 - 2\beta) \frac{\Delta t}{P_k} sofd - \frac{P_{k+\frac{1}{2}} \epsilon_{k+\frac{1}{2}} (E_{lb\ k+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}) \theta \Delta t}{\Delta z^2 P_k} \\
& - \frac{P_{k-\frac{1}{2}} \epsilon_{k-\frac{1}{2}} (E_{lb\ k-\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}) \theta \Delta t}{\Delta z^2 P_k} + (\epsilon_k + \rho_{b,k} K_{L,b}) (1 - k_b \theta \Delta t) \\
RBD &= - \frac{\ell q}{\Delta z} \theta \beta \frac{\Delta t}{P_k} sofd + \frac{P_{k+\frac{1}{2}} \epsilon_{k+\frac{1}{2}} (E_{lb\ k+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}) \theta \Delta t}{\Delta z^2 P_k} \quad (7.16)
\end{aligned}$$

with

$K_{L,b}$ = slope of (linear) sorption isotherm of bottom material ($L^3.M^{-1}$)
 $sofd$ = factor equalling 0 or 1, preventing that the dispersion flux would exceed the advection flux in the sediment (see section 4.2)

With

$$V = \epsilon_k + \rho_{b,k} K_{L,b} \quad (V > 0)$$

and

$$W^{14} = \frac{\ell q}{\Delta z} (1 - 2\beta) \frac{sofd}{P_k} + \frac{P_{k+\frac{1}{2}} \epsilon_{k+\frac{1}{2}} (E_{lb\ k+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}})}{\Delta z^2 P_k}$$

¹⁴ The sign of W depends on the values of q and β .

$$+ \frac{P_{k-1/2} \varepsilon_{k-1/2} (E_{lb, k-1/2} + D_{lb, k-1/2})}{\Delta z^2 P_k} + V k_b$$

the element LDD may be rewritten as:

$$LDD = W (1 - \theta) \Delta t + V \quad (7.17)$$

and the element RDD may be rewritten as:

$$RDD = - W \theta \Delta t + V \quad (7.18)$$

The positivity conditions for the off-diagonal elements of the two tridiagonal matrices lead to an advice about the size of the space steps, Δz .

The condition $LBD \leq 0$ (or $RBD \geq 0$) leads to:

$$\Delta z \leq \frac{P_{k+1/2} \varepsilon_{k+1/2} (E_{lb, k+1/2} + D_{lb, k+1/2})}{\ell q \beta \text{sofd}} \quad (7.19)$$

for $q > 0$, $E + D > 0$ and $\beta \neq 0$ and $\text{sofd} \neq 0$.

The condition $LOD \leq 0$ (or $ROD \geq 0$) leads to:

$$\Delta z \leq - \frac{P_{k-1/2} \varepsilon_{k-1/2} (E_{lb, k-1/2} + D_{lb, k-1/2})}{\ell q (1 - \beta) \text{sofd}} \quad (7.20)$$

for $q < 0$, $E + D > 0$ and $\beta \neq 1$ and $\text{sofd} \neq 0$.

If $q = 0$ the conditions described above do not lead to restrictions for Δz . When a positive (downward) seepage is combined with $\beta = 0$ or when a negative (upward) seepage is combined with $\beta = 1$, there is also no restriction for Δz .

The positivity conditions for the diagonal elements of the two tridiagonal matrices lead to an advice about the size of the time step, Δt . Generally speaking, the water subsystem is more dynamic than the sediment subsystem, so one may expect that the restriction for Δt will be stricter for the water than for the sediment subsystem. In the case of a pond, however, with zero flow, the situation may be different.

The condition $LDD > 0$ leads to:

$$W (1 - \theta) \Delta t + V > 0 \quad (7.21)$$

corresponding to:

$$1. \quad W > 0 : \Delta t > \frac{-V}{W (1 - \theta)} \quad (7.22)$$

In this case no restriction for the selection of Δt is found.

$$2. \quad W < 0 : \Delta t < \frac{-V}{W (1 - \theta)} \quad (7.23)$$

and $W < 0$ corresponds to:

$$(i) \quad \beta > \frac{1}{2} \text{ \acute{a}nd } Z < q \quad (7.24)$$

Z being positive
or

$$(ii) \quad \beta < \frac{1}{2} \text{ \acute{a}nd } Z > q \quad (7.25)$$

Z being negative
with

$$Z = \frac{- \left\{ P_{k+\frac{1}{2}} \epsilon_{k+\frac{1}{2}} (E_{lb, k+\frac{1}{2}} + D_{lb, k+\frac{1}{2}}) + P_{k-\frac{1}{2}} \epsilon_{k-\frac{1}{2}} (E_{lb, k-\frac{1}{2}} + D_{lb, k-\frac{1}{2}}) + V k_b \Delta z^2 P_k \right\}}{\ell (1 - 2\beta) \text{ sofd } \Delta z}$$

When $\beta = \frac{1}{2}$ or $q = 0$ $W (1-\theta)$ is always positive, so the condition $LDD > 0$ does not lead to a restriction for Δt . When $\theta = 1$ $W (1-\theta)$ equals zero, so the condition $LDD > 0$ is always true and the condition $RDD > 0$ may impose restrictions.

The condition $RDD > 0$ leads to:

$$- W \theta \Delta t + V > 0 \quad (7.26)$$

corresponding to:

$$1. \quad W < 0 : \Delta t > \frac{V}{W \theta} \quad (7.27)$$

In this case no restriction for the selection of Δt is found.

$$2. \quad W > 0 : \Delta t < \frac{V}{W \theta} \quad (7.28)$$

and $W > 0$ corresponds to:

$$(i) \quad \beta > \frac{1}{2} \text{ and } Z > q \quad (7.29)$$

Z being positive
or

$$(ii) \quad \beta < \frac{1}{2} \text{ and } Z < q \quad (7.30)$$

Z being negative.

When $\beta = \frac{1}{2}$ or $q = 0$ $W \theta$ is always positive so the condition $RDD > 0$ does lead to a restriction for Δt . When $\theta = 0$, $W \theta$ equals zero, so the condition $RDD > 0$ is always true but now the condition $LDD > 0$ may impose a restriction for Δt .

Figure 20 summarizes the results for the restrictions imposed on the time step, Δt . It appears that when $\theta = 0$ and $|Z| > |q|$ (i.e. $-|Z| < q < |Z|$) the condition $LDD > 0$ is not fulfilled, so the solution to matrix Eq.(6.39) is not positive. So, the matrix Eq. (6.39) has only a positive solution if Δz is selected such that $|Z| < |q|$.

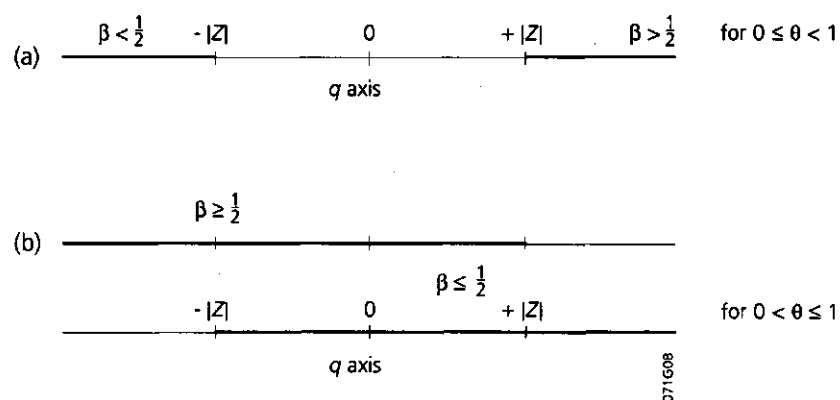


Fig. 20 Outline of the domain (q, β) for which the condition $LDD > 0$ (part (a)) or $RDD > 0$ (part(b)) result in a restriction for the time step, Δt

7.3 Consistency condition

7.3.1 Introduction

The consistency condition was tested according to three of the criteria formulated in section 7.1.

4. Letting the finite differences approach zero should result in the numerical equations reducing to the original partial differential equations.

Inspection of the approximations of the right-hand and left-hand terms of the conservation equations for the water and sediment subsystems makes it clear that all procedures are indeed consistent. (See Eqs. (6.17) and (6.23) for the water subsystem and Eqs. (6.33) and (6.37) for the sediment subsystem.)

5. The mass of pesticide introduced (100%) should be traceable at every moment with an error less than 0.1% (this means that the mass balances tally and that no pesticide is lost during the calculation process by any other process than those being modelled, e.g. transformation or volatilisation).

To check that no mass is lost in the calculation process, separate mass balances are drawn up for the water subsystem and each sediment subsystem. (Each subsystem is treated as a whole entity, meaning that it is not subdivided when the mass balances are calculated.) For each subsystem all incoming and outgoing fluxes, including the transformed pesticide mass, are summated in time and the quantity missing from the mass balance is calculated. The TOXSWA program monitors whether this missing mass does not exceed a certain percentage of the initial total mass present plus incoming mass.¹⁵

The numerical solution of the mass conservation equation is verified against an analytical solution, described below.

6. An analytical solution of the advection-dispersion equation for constant h and A for a system with transformation, volatilisation and convective/dispersive downward seepage responding to a Dirac delta function-type input, as described in Jury and Roth (1990). (Volatilisation and convective/dispersive downward seepage are, from a mathematical point of view, equivalent to transformation.) This solution should correspond with that obtained with the numerical solution scheme. The same type of analytical solution can also be applied to the sediment subsystem. Hence, the water and sediment subsystems can be checked independently.

The analytical solutions are worked out in the next sections.

¹⁵ In the TOXSWA model version 1.0 error messages are written to a message file when the missing mass exceeds 0.1%. (There is a maximum of 20 error messages about mass balances per subsystem, water layer and sediment.)

7.3.2 Analytical solutions of the Convection-Dispersion Equation

In Jury and Roth (1990) transfer functions are used to estimate outflow concentrations from soil columns. Transfer functions are used to model such a complex system in a simple way by characterizing the output flux as a function of the input flux. Outflow concentrations may be described with the aid of a travel time probability distribution or probability density function (pdf) for outflowing solute molecules. The travel time pdf characterizes the distribution of possible travel times that a solute molecule might experience in moving from the inlet end to the outlet end. Example 3.1 of Jury and Roth provides the solution of the Convection-Dispersion Equation in case of the entry of a narrow pulse (delta function) of solute flux concentration through the inlet end $z = 0$ at $t = 0$, assuming that there is no solute present in the system initially (Eq. 3.12, divided by θV !, compare the solution of Problem 3.8 in Jury and Roth (1990)).

$$c_i^r(z, t) = \frac{1}{\sqrt{\pi D t}} \exp\left(-\frac{(z - Vt)^2}{4 D t}\right) - \frac{V}{2D} \exp\left(\frac{Vz}{D}\right) \operatorname{erfc}\left(\frac{z + Vt}{\sqrt{4 D t}}\right) \quad (7.31)$$

in which

c_i^r = probability density of the total resident concentration (here 1, normalised, L^{-1})

D = effective diffusion-dispersion coefficient ($L^2.T^{-1}$)

t = time (T)

z = depth (L)

V = pore water velocity ($L.T^{-1}$)

and where

$\operatorname{erfc}(x) = 1 - \operatorname{erf}(x)$, the complementary error function, defined by

$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x \exp(-y^2) dy \quad (7.32)$$

This applies to non-sorbing, non-decaying substances and normalised situations (so the integral of the total concentration with depth equals always 1.).

For substances undergoing simultaneous linear adsorption and first order decay during transport under steady state water flow conditions, and for an application of mass M the solution reads (Boesten, 1992, combining sections 4.2 and 4.8 of Jury and Roth (1990)):

$$c_i^r = \frac{c_i^r}{\theta R} = \frac{M}{\theta R} \cdot e^{-\mu t} \left\{ \frac{1}{\sqrt{\pi D \frac{t}{R}}} \exp \left(-\frac{\left(z - V \frac{t}{R} \right)^2}{4 D \frac{t}{R}} \right) - \frac{V}{2D} \exp \left(\frac{Vz}{D} \right) \cdot \operatorname{erfc} \left(\frac{z + V \frac{t}{R}}{\sqrt{4 D \frac{t}{R}}} \right) \right\} \quad (7.33)$$

in which

c_i^r = mass of dissolved solute per volume of fluid (M.L⁻³)

c_i^r = total resident concentration (M.L⁻³)

M = applied area-averaged mass (M.L⁻²)

θ = volume fraction of water (1)

R = retardation factor (1)

μ = first-order transformation rate constant (T⁻¹)

So, Eq.(7.33) provides the analytical solution for the generalised form of the Convection-Dispersion Equation in case of the entry of a narrow pulse (delta function) of solute flux concentration through the inlet end $z = 0$ at $t = 0$, assuming that there is no solute present in the system initially and for a substance undergoing simultaneous linear sorption and first-order decay during transport with a steady state water flow. The generalised form of the Convection-Dispersion Equation reads:

$$\begin{aligned} (\theta + \rho_b K_{L,b}) \frac{\partial c_i^r}{\partial t} &= -\theta V \frac{\partial c_i^r}{\partial x} + \theta D \frac{\partial^2 c_i^r}{\partial x^2} - \mu (\theta + \rho_b K_{L,b}) c_i^r \\ \Leftrightarrow R \frac{\partial c_i^r}{\partial t} &= -V \frac{\partial c_i^r}{\partial x} + D \frac{\partial^2 c_i^r}{\partial x^2} - \mu R c_i^r \end{aligned} \quad (7.34)$$

with

$K_{L,b}$ = slope of (linear) sorption isotherm of bottom material (L³.M⁻¹)

and

$$R = 1 + \frac{\rho_b K_{L,b}}{\theta} \quad (1)$$

where R is called the retardation factor (Bolt, 1979). The retardation factor R is defined as the ratio of the mean flow velocity of a non-sorbing substance divided by the mean flow velocity of a sorbing substance in a system with exclusively convective transport.

For the water subsystem the more simple analytical solution of the Convection-Dispersion Equation in case of an instantaneous source, e.g. a sudden drain discharge

has been chosen. This solution applies for an infinite and not a semi-finite system as described here above. This solution is given in Leistra (1973) and reads for a soil system (in case of decay combined with adsorption):

$$c_l^r = \frac{M}{2\theta R} \frac{1}{\sqrt{\pi D \frac{t}{R}}} e^{-\mu t} \exp\left(-\frac{(x - V \frac{t}{R})^2}{4 D \frac{t}{R}}\right) \quad (7.35)$$

7.3.3 Application of the analytical solutions of the Convection-Dispersion Equation to the water and sediment subsystems of TOXSWA

We now apply the solutions described above to the two subsystems of the TOXSWA model.

The conservation equation for the sediment, Eq. (3.12) reads:

$$P \frac{\partial c_b^*}{\partial t} = - \frac{\partial(PJ_{lb})}{\partial z} - k_b c_b^* P \quad (3.12)$$

Assuming sorption to sediment is a linear process, so

$$X_b = K_{L,b} \cdot c_{lb} \quad (7.36)$$

with

$K_{L,b}$ = slope of (linear) sorption isotherm of bottom material ($L^3 \cdot M^{-1}$)

and using Eq. (4.17):

$$c_b^* = \varepsilon c_{lb} + \rho_b K_{L,b} c_{lb} \quad (7.37)$$

for constant perimeters P (so for a rectangular-shaped cross section), and substituting Eq. (4.21), Eq. (3.12) now reads:

$$(\varepsilon + \rho_b K_{L,b}) \frac{\partial c_{lb}}{\partial t} = - \frac{dq}{P} \frac{\partial c_{lb}}{\partial z} + \varepsilon (E_{lb} + D_{lb}) \frac{\partial^2 c_{lb}}{\partial z^2} - k (\varepsilon + \rho_b K_{L,b}) c_{lb}$$

$$\Leftrightarrow R_{wb} \frac{\partial c_{lb}}{\partial t} = -w \frac{\partial c_{lb}}{\partial z} + (E_{lb} + D_{lb}) \frac{\partial^2 c_{lb}}{\partial z^2} - k R_{wb} c_{lb} \quad (7.38)$$

in which

R_{wb} = retardation factor for the sediment subsystem (1)

$$R_{wb} = 1 + \frac{\rho_b K_{L,b}}{\varepsilon}$$

This is the generalised form of the Convection-Dispersion Equation model for a substance undergoing simultaneous transport, linear equilibrium sorption and first order decay. For a semi-infinite system its solution for a pulse input at $z = 0$ and $t = 0$ reads (in the notation used for the TOXSWA model description) (compare Eq. (7.33)):

$$c_b^* = M e^{-kt} \left\{ \frac{1}{\sqrt{\pi (E_{lb} + D_{lb}) \frac{t}{R_{wb}}}} \exp \left(- \frac{\left(z - w \frac{t}{R_{wb}} \right)^2}{4 (E_{lb} + D_{lb}) \frac{t}{R_{wb}}} \right) - \frac{w}{2(E_{lb} + D_{lb})} \exp \left(\frac{wz}{(E_{lb} + D_{lb})} \right) \operatorname{erfc} \left(\frac{z + w \frac{t}{R}}{\sqrt{4 (E_{lb} + D_{lb}) \frac{t}{R}}} \right) \right\} \quad (7.39)$$

The conservation equation for the water layer, Eq.(3.6) reads:

$$\frac{\partial(c^*A)}{\partial t} = - \frac{\partial(AJ)}{\partial x} - k(c^*A) + J_{wa} \cdot O_x - J_{wb} \cdot P_x \quad (3.6)$$

Assuming sorption to suspended solids is a linear process, so

$$X_{ss} = K_{L,ss} \cdot c \quad (7.40)$$

with

$K_{L,ss}$ = slope of (linear) sorption isotherm of suspended solids ($L^3.M^{-1}$)

and combining Eqs. (4.1) and (4.4):

$$c^* = \left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right) c \quad (7.41)$$

For constant wetted areas A , assuming exclusively convective, downward seepage, and substituting Eqs. (4.6), (4.11) and (4.14), the conservation equation, Eq. (3.6) now reads:

$$\left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right) \frac{\partial c}{\partial t} = -u (1 + ss K_{L,ss}) \frac{\partial c}{\partial x} + E_x (1 + ss K_{L,ss}) \frac{\partial^2 c}{\partial x^2}$$

$$- k \left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right) c - k_{t,l} \left(c - \frac{c_a}{K_H} \right) \frac{O_x}{A} - \frac{\ell q}{A} c \Leftrightarrow$$

$$R_{wl} \frac{\partial c}{\partial t} = -u \frac{\partial c}{\partial x} + E_x \frac{\partial^2 c}{\partial x^2} - k R_{wl} c - \frac{k_{t,l} O_x}{(1 + ss K_{L,ss}) A} c - \frac{\ell q}{A (1 + ss K_{L,ss})} c \quad (7.42)$$

(assume c_a , mass concentration of substance in the air, equals zero)

in which

R_{wl} = retardation factor for the water subsystem (1)

$$R_{wl} = \frac{\left(1 + \frac{DW P_{z=0}}{A} K_{mp} + ss K_{L,ss} \right)}{1 + ss K_{L,ss}}$$

For a finite system with $\varepsilon = 1.0$ its solution for a pulse input at $t = 0$ reads (in the notation used for the TOXSWA model description) (compare Eq.(7.35)):

$$\begin{aligned}
c = & \frac{M}{2R_{wl} (1 + ss K_{L,ss})} \frac{1}{\sqrt{\pi E_x \frac{t}{R_{wl}}}} \\
& \cdot \exp \left\{ - \left(k + \frac{k_{z,l} O_x}{(1 + ss K_{L,ss}) A R_{wl}} + \frac{qg}{A (1 + ss K_{L,ss}) R_{wl}} \right) t \right\} \\
& \cdot \exp \left\{ - \frac{\left(x - u \frac{t}{R_{wl}} \right)^2}{4E_x \frac{t}{R_{wl}}} \right\}
\end{aligned} \tag{7.43}$$

7.3.4 Comparison of the analytical solutions with the numerical solutions (i.e. TOXSWA)

To be able to compare the analytical solutions of section 7.4.2 with TOXSWA model calculations the source code of TOXSWA has been changed in such a way that there was no diffusion across the water-sediment interface (See Annex 15). So, only advective transport took place across the water-sediment interface.

In the sediment subsystem a pulse input was applied at $t = 0$ in the first node, i.e. at $z = 0.0005$ m, in case of the TOXSWA calculation. The applied pulse corresponded to an initial total concentration c_b^* of 10 g/m^3 in the first segment of 1 mm thickness. For the analytical solution (Eq. (7.39)) the same mass (corresponding to 10 mg/m^2) was applied at the sediment surface. To be able to compare the TOXSWA results with the analytical solution a rectangular-shaped cross section of the ditch (so constant perimeter P with depth) was selected and bulk density, porosity and organic matter content were kept constant with depth. Sorption to solid bottom material was described with a linear sorption isotherm and there was a constant downward seepage of $2 \text{ mm/(m}^2 \cdot \text{d)}$, expressed as an infiltration flux per m^2 neighbouring field lot. All other input parameters were kept identical to the ones of the example simulation of chlorpyrifos of Chapter 8. Annex 16 summarizes the differences in the input files between the example simulation of Chapter 8 and the simulation carried out here to compare the numerical with the analytical solution for the sediment subsystem.

Figure 21 shows the concentration profiles in the upper centimeters of the sediment, calculated by the TOXSWA model and by the analytical solution, Eq. (7.39). Figure 21 shows an excellent correspondence between the numerical and the analytical solution. For the analytical solution Table 3 details the situation at the first node of 0.0005 m for the period of 0.0 to 0.5 day after application of the pulse load at the sediment surface. Table 3 demonstrates that it takes 0.0625 day for this pulse to arrive at 0.0005 m depth in the sediment. Due to dispersion and diffusion the concentration peak already lowered and therefore this peak is lower than the concentration peak that the TOXSWA model calculated for the first node in the sediment directly after application ($t = 0.001$ d). TOXSWA calculated a concentration peak of 9.97 g/m^3 at $t = 0.001$ d (i.e. after sorption to bottom material and some transport to the second node) (see also Fig. 21).

In the water subsystem a pulse input was applied at $t = 0$ in node 8 at $x = 60$ m in the ditch for the TOXSWA simulation. The pulse corresponded to an initial total concentration c^* of 8.264 g/m^3 in the eighth segment of 8 m long. For the analytical solution, Eq. (7.43), the same total mass (corresponding to $6.6112\text{E-}02 \text{ g/m}^2$, expressed per m^2 wetted cross section area A) was applied at 60 m in the ditch. Total ditch length was 400 m in this TOXSWA simulation, the ditch cross section was again rectangular and there was a constant downward seepage of $2 \text{ mm}/(\text{m}^2 \cdot \text{d})$, expressed as an infiltration flux per m^2 neighbouring field lot. Sorption to suspended solids was described with a linear sorption isotherm and all other input parameters were identical to the ones of the example simulation for chlorpyrifos of Chapter 8. Annex 17 gives an overview of the differences in the input files used here and those used in the example simulation of Chapter 8.

Table 3 Total pesticide concentrations in node 1 ($z = 0.0005$ m) of the sediment as a function of time, calculated with the aid of the analytical solution for the sediment, Eq.(7.39)

Time (d)	Total concentration ($\text{g} \cdot \text{m}^{-3}$)
0.0125	2.96
0.0250	7.30
0.0375	9.03
0.0500	9.62
0.0625	9.73
0.0750	9.64
0.0875	9.46
0.1	9.24
0.2	7.56
0.3	6.45
0.4	5.70
0.5	5.14

Figure 22 shows the concentration profiles in the water layer of the ditch, calculated by the TOXSWA model and by the analytical solution, Eq. (7.43). Correspondence is again excellent. So, TOXSWA 1.0 has been verified successfully both for the water and sediment subsystems.

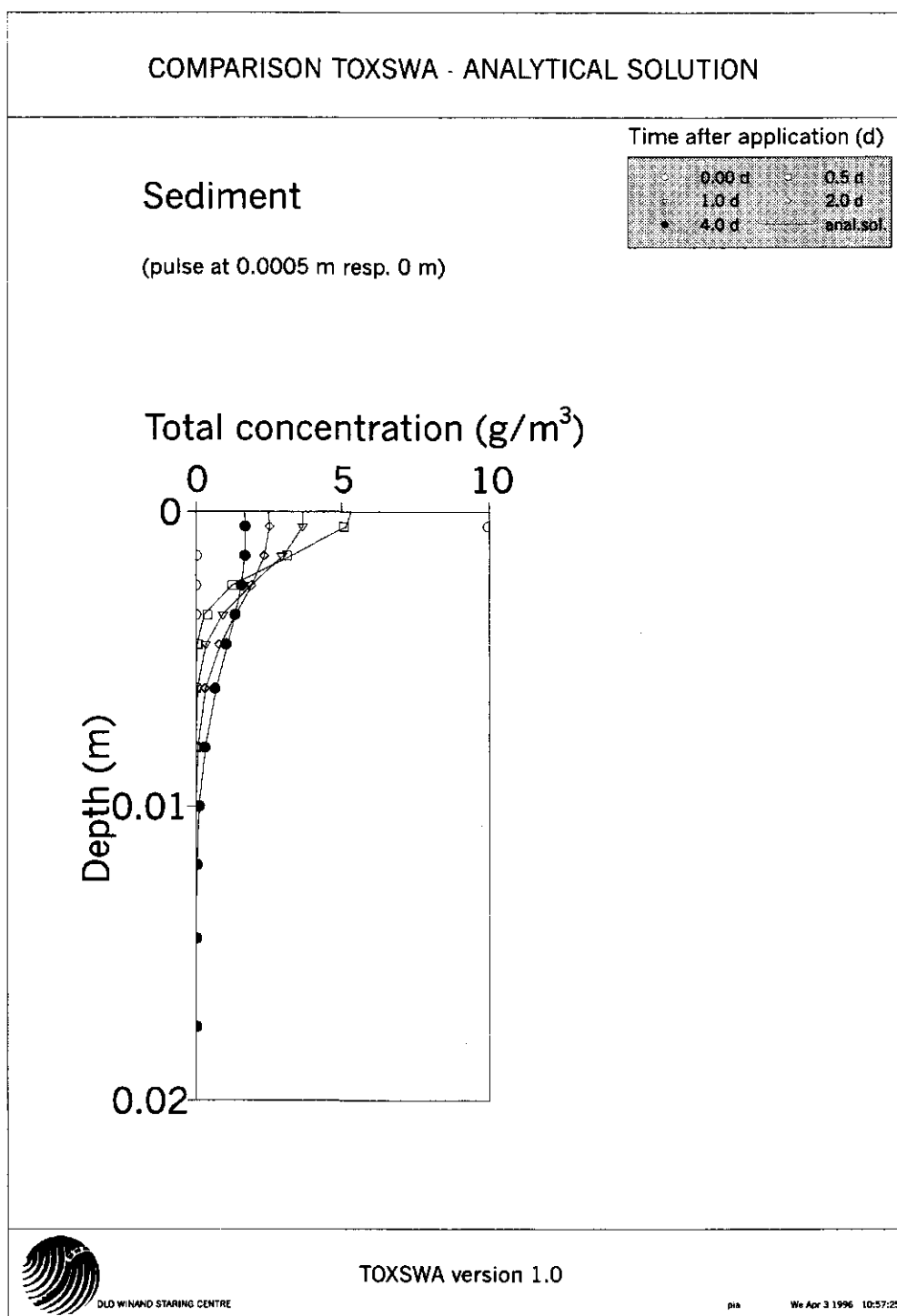


Fig. 21 Comparison of the total chlorpyrifos concentration in the sediment calculated with the aid of the TOXSWA model (indicated by the markers) and with the aid of the analytical solution for the mass conservation equation of the sediment, Eq.(7.39) (indicated by the drawn lines)

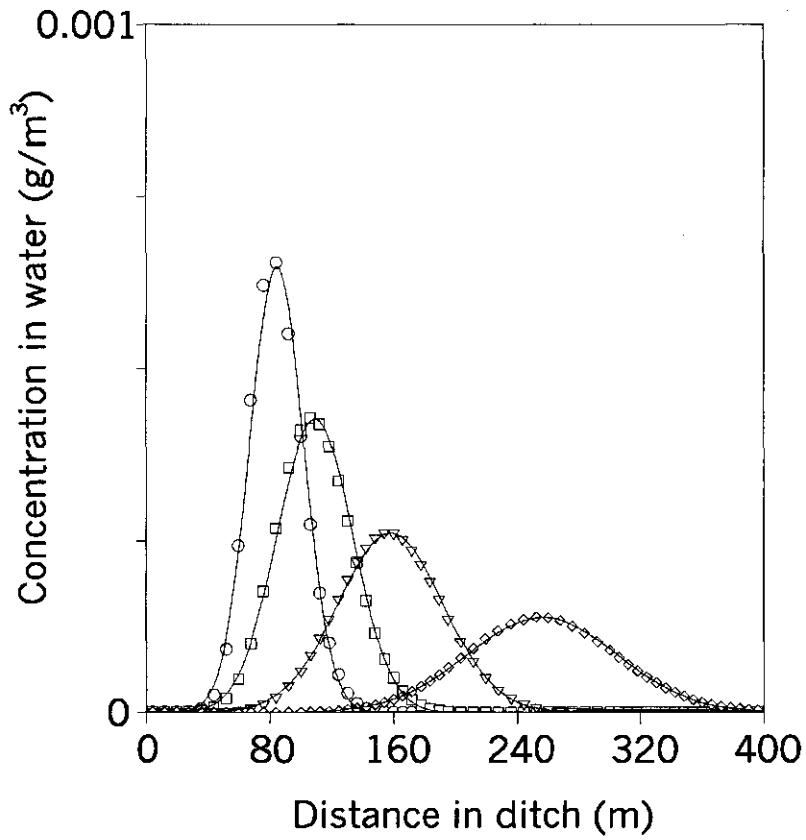
COMPARISON TOXSWA - ANALYTICAL SOLUTION

Water layer

(pulse at 60 m)

Time after application (d)

○	0.5 d	□	1.0 d
△	2.0 d	◇	4.0 d
—	anal. sol.		



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Fig. 22 Comparison of the chlorpyrifos concentration in the water phase of the ditch calculated with the aid of the TOXSWA model (indicated by the markers) and with the aid of the analytical solution for the mass conservation equation of the water layer, Eq.(7.43) (indicated by the drawn lines)

8 Illustration of the potential use of TOXSWA: simulation of chlorpyrifos behaviour after spray drift deposition

8.1 Design of computation and values of parameters

To illustrate the possibilities of the TOXSWA model a computation was carried out for the insecticide chlorpyrifos. The following situation was simulated. Of an application rate of 1 kg a.i. per ha 3% is deposited by spray drift onto a neighbouring ditch. The ditch is 200 m long and deposition of chlorpyrifos occurs over the total length of the ditch except the first 24 m. Water flow rate in the ditch is 100 m/d, the water depth is 50 cm and the cross section of the ditch has a trapezoidal shape. An amount of 250 g of dry macrophyte biomass is present per m² ditch bottom, corresponding to a moderately-grown ditch in summertime; the suspended solids concentration is 50 g/m³. Initially, the water does not contain chlorpyrifos.

The sediment layer in the ditch is 10 cm thick and porosity varies from 80% in the top mm to 40% in the bottom cm. The organic matter content varies from 8% to 0.5%. There is no upward or downward seepage in the sediment. Initially, the sediment is free of pesticide.

The explicit central difference calculation scheme was selected to solve the mass conservation equations. Distances between the nodes in the water and sediment subsystems were 6 to 8 m and 1 to 10 mm, respectively. The time step was 100 s and the run time on a Pentium 90 MHz was 32 minutes.

Chlorpyrifos transformation is characterised by a half-life time of 75 and 175 d in the water layer and the sediment, respectively; these values have been determined at the DLO Winand Staring Centre in earlier experiments. Sorption to sediment is very strong ($K_{om} = 16.4 \text{ m}^3/\text{kg}$ and $n = 0.984$) (Pers. comm. Crum) as is sorption to macrophytes ($K_{mp} = 2.0 \text{ m}^3/\text{kg}$) (Van Huffelen, 1993).

The input files for this example simulation are presented in Annex 18.

8.2 Results and discussion

Figure 23 shows the chlorpyrifos concentration in the water phase as well as the total mass concentration in the sediment as a function of time and depth. The upper graph shows that, due to dispersion, the concentration front flattens as it moves out of the ditch. The spray drift deposition of 0.03 kg/ha on the ditch results in an initial concentration of 4.4 µg/l in the water (after the instantaneous sorption to the macrophytes and suspended solids). Without any sorption the initial concentration would have been 8.3 µg/l. The lower graphs show the penetration of chlorpyrifos into the top 2 cm of the sediment. After 4 days, total concentrations built up vary from negligible (< 0.3 µg/dm³) at 4 m along the ditch up to 415 µg/dm³ at 196 m

along the ditch. After this time, back-diffusion from the sediment into the water layer has already started, a process which is driven by the difference in pesticide concentration between the liquid phase (i.e. not the total mass concentration) of the sediment and the water in the ditch. The graph for the situation at 100 m clearly shows this phenomenon, as the total concentration is already decreasing. It can be derived from Figure 18 that the average rate of displacement of the pesticide in the water layer is about 60 m/d, so about 60% of the water flow rate.

Figure 23 also presents the average concentration to which aquatic organisms are exposed at 3, 21 and 28 days after application, as well as immediately after application (at 0 d). The exposure concentrations have been determined by calculating the average of the concentration course for pesticide dissolved in the water phase with time. (This implies that the exposure concentrations are calculated according to the assumptions of instantaneous mixing and sorption.) The concentration presented at time $t = 0$ d corresponds to the total pesticide concentration, immediately after application, so this means before any adsorption has taken place. The exposure concentration has been defined as the concentration at that position in the ditch where the longest exposure duration is expected, i.e. at the downstream end of the section of the ditch where the pesticide input took place. In this example the exposure concentrations were 8.3 $\mu\text{g/l}$ immediately after application and 3.5 $\mu\text{g/l}$ at $t = 3$ d.

Figure 24 presents the distribution of chlorpyrifos between the different compartments as a function of the time since application. The upper graph indicates that a total of nearly 2 g of the insecticide has been deposited in the ditch. Initially, 53% is dissolved in the water phase, while 42% and 4% are adsorbed to the macrophytes and suspended solids, respectively. After 4 days, the largest part (72%) of the remaining mass (0.1 g) is found in the sediment, 15% is dissolved in the water phase, 12% has been adsorbed to the macrophytes and 1.3% to the suspended solids. The lower graphs present the chlorpyrifos mass per running metre at selected locations in the ditch. Roughly speaking, they present the same pattern as that shown above, albeit with a more rapid decline in pesticide mass, due to the passage of the concentration front through the water layer. The longer chlorpyrifos is present in the overlying water layer, the more it penetrates into the sediment. In the first part of the ditch, a negligible amount ($< 0.6 \mu\text{g/m}^3$) (not visible) has penetrated via diffusion into the sediment, due to 'back dispersion' of the insecticide into this part of the overlying water layer.

Figure 25 shows the mass balances of chlorpyrifos as a function of time for the entire ditch. Separate mass balances are presented for the two subsystems, water layer and sediment. The total mass in the water layer decreases steadily, mainly due to outflow. Other factors contributing to the decrease are, in order of priority, volatilisation, penetration into sediment and transformation. The total mass in the sediment increases up to about 3 days after application. From ca. day 2 onwards, back-diffusion to the water layer becomes significant, as well as transformation, and the total mass in the sediment gradually starts to decrease. Initially, nearly 2 g of chlorpyrifos is present in the water layer, whereas a maximum of about 0.1 g can be found in the sediment later.

The TOXSWA model calculates how much mass is missing in the mass balances for each subsystem, i.e. the (only) water subsystem and, in this case, 29 sediment subsystems. This missing quantity is expressed as a percentage of the initial mass plus the incoming additional mass for each subsystem. In the example simulation for chlorpyrifos the missing quantity in the mass balance for the entire water layer (subdivided into 29 nodes) after 3456 time steps of 100 s (corresponding to 4.00 d) was 0.0037 % ($0.72 \cdot 10^{-4}$ g) of the dose. For the entire sediment, so the summation of the 29 sediment subsystems, the missing mass was 0.0042% ($0.43 \cdot 10^{-5}$ g) of the initial mass plus the incoming mass from the water layer.

The three types of graph in Figures 23, 24 and 25 give an overview of the fate of the pesticide. Concentrations in water layer and sediment are shown and the mass balance graphs explain how these concentrations have developed. The distribution between the different compartments indicates where the insecticide can be found at different points in time. This example of a pesticide with a high sorption capacity shows that considerable amounts of the pesticide can be found in the macrophyte and sediment compartments. Instantaneous sorption equilibrium is assumed in the model. Measurements during past experiments with chlorpyrifos have indicated, however, that it takes about one day before chlorpyrifos deposited onto a 50 cm deep ditch is mixed over the entire depth of the water layer (Crum and Brock, 1994). The sorption equilibrium with macrophytes and with sediment may also take up to about 24 hours (Van Huffelen, 1993). This means that the simulated pesticide concentrations in the water phase may be too low and that the simulated mass concentration in the macrophyte compartment especially may be too high during the first one or two days.

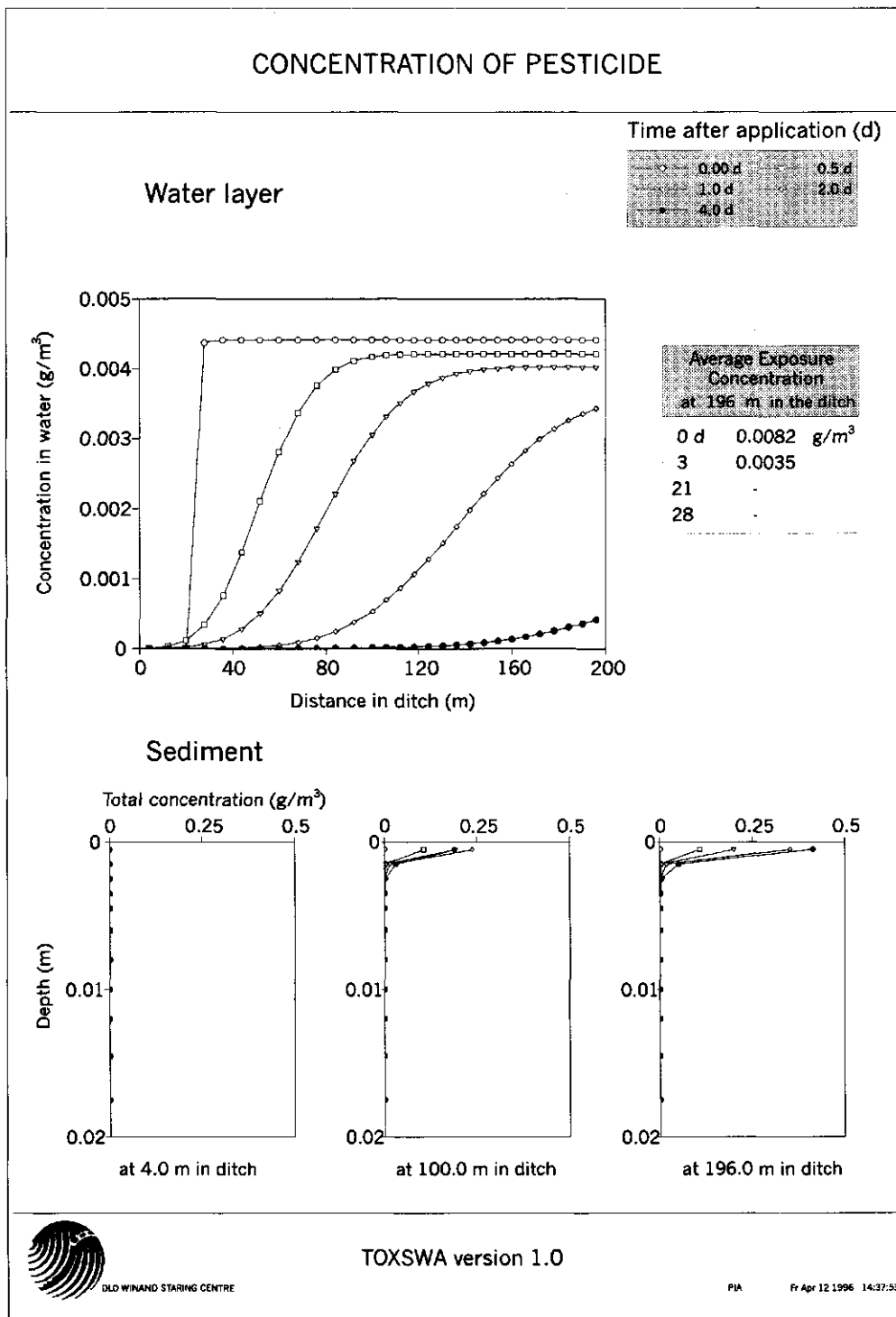


Fig. 23 Chlorpyrifos concentration in the water layer and at selected locations in the sediment after spray drift deposition of 0.03 kg.ha⁻¹ onto the ditch

DISTRIBUTION OF PESTICIDE

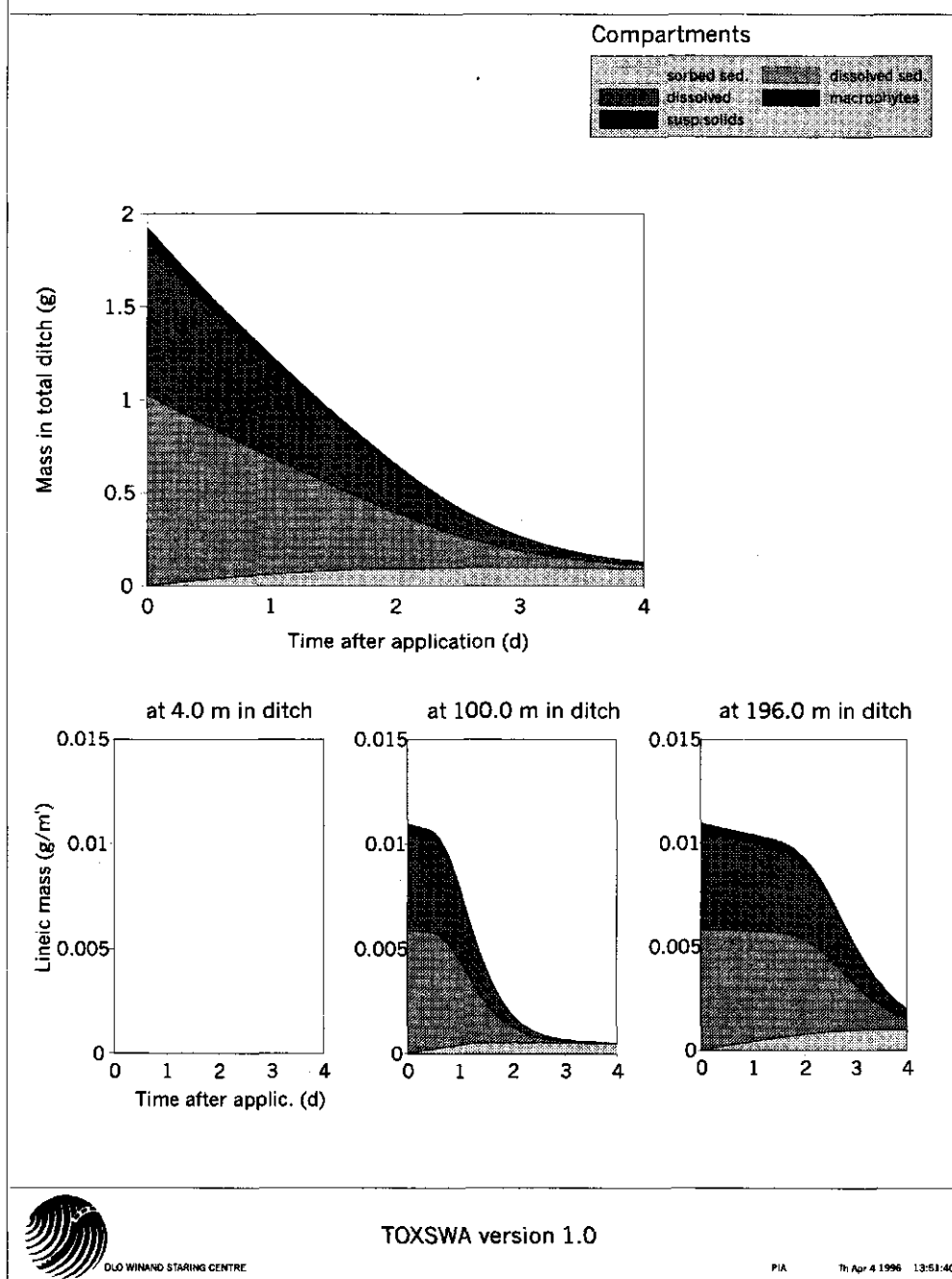
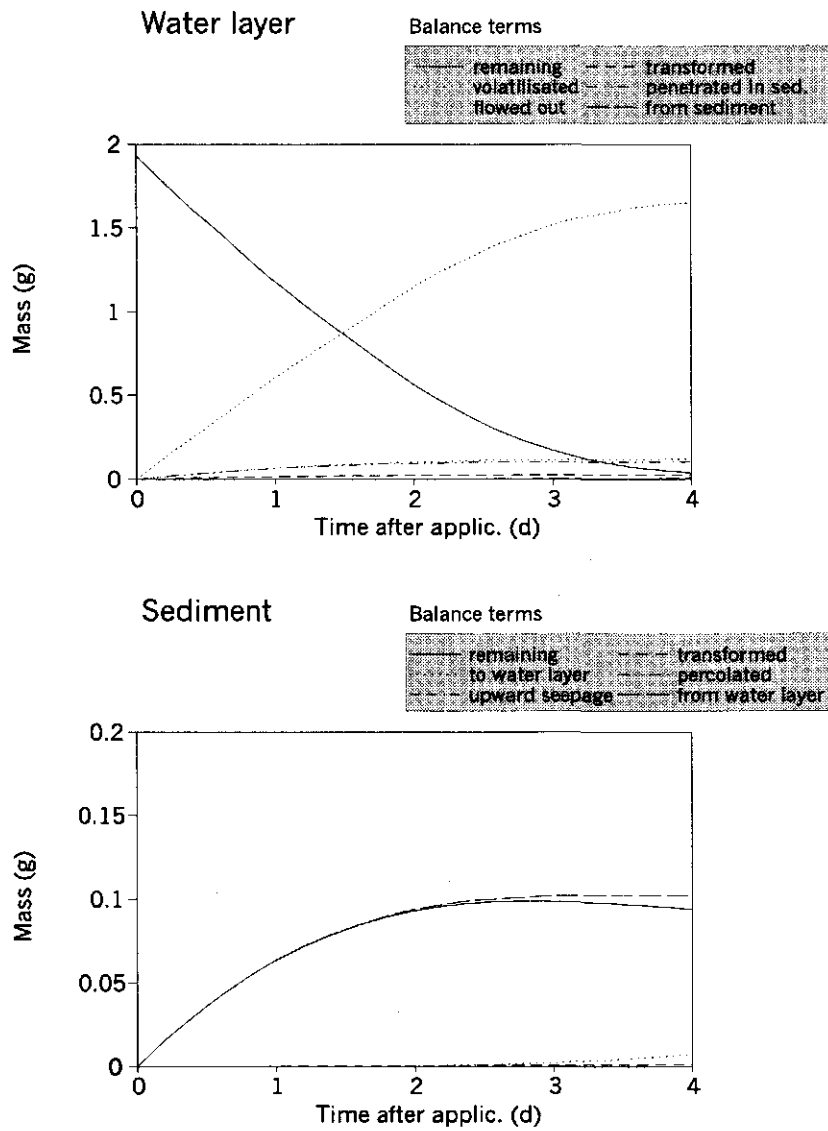


Fig. 24 Distribution of chlorpyrifos between the different compartments as a function of time for the entire ditch (200 m) as well as per running metre at selected locations. (The mass dissolved in the liquid phase of the sediment is too small to be visible)

MASS BALANCE OF PESTICIDE



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Fig. 25 Mass balances of chlorpyrifos as a function of time for the entire ditch (200 m). Separate mass balances are shown for the water and sediment subsystems

9 Conclusions, discussion and recommendations

9.1 Conclusions and discussion

The TOXSWA model has been developed to describe the fate of pesticides in field ditches. It calculates pesticide concentrations in the horizontal direction in the water layer and in the vertical plus horizontal directions in the sediment. It can handle a variety of situations as regards hydrological conditions and entry routes of pesticides into surface water.

A limited verification of the model has taken place by comparing model output with an analytical solution for the sediment subsystem, as well as for the water subsystem. In both cases the concentration profiles calculated by the model showed an excellent correspondence with those calculated according to the analytical solution. So, TOXSWA 1.0 has been verified successfully both for the water and the sediment subsystems.

An example simulation described the behaviour of the insecticide chlorpyrifos after spray drift deposition in a ditch. Results showed that initially, nearly half of the insecticide was adsorbed to the macrophytes and this provokes a retardation of the concentration front of about 60% compared to the water flow rate. After four days about 70% of the remaining pesticide mass was found in the sediment.

The TOXSWA model checks whether the mass balances tally during the calculations. Results for the example simulation for chlorpyrifos showed that after four days the missing quantity in the mass balance for the water layer was less than 0.005% of the dose applied. The missing mass for the entire sediment layer was also less than 0.005% of the initial mass plus incoming mass of the water layer after four days, so this shows that conservation of pesticide mass applied is well done in TOXSWA 1.0.

TOXSWA has been developed to serve as a tool in the pesticide registration procedure in the Netherlands. It has been designed to estimate chronic exposure of aquatic organisms to pesticides, so it simulates periods of up to about one month. Of course, the TOXSWA model also estimates acute exposure, but the simulated acute concentrations might differ from those encountered in the field, due to assumptions made in the model. In reality, pesticides, especially those with low solubilities, may need 24 hours before they are thoroughly mixed over the entire cross section of the ditch; subsequently, it takes time before sorption equilibrium is reached. The TOXSWA model, however, assumes instantaneous mixing over the entire cross section, as well as instantaneous sorption equilibrium with suspended solids, macrophytes and with the solid bottom material (once the pesticide has entered the sediment subsystem).

In accordance with the aim of its development (exposure concentrations for up to about one month) it is assumed that no sedimentation or resuspension occurs. This means that the TOXSWA model cannot be used to estimate long-term exposure concentrations or accumulation of pesticides in the sediment.

9.2 Recommendations

If one is to make full use of all the possibilities the TOXSWA model offers, model improvements are needed:

- inclusion of multiple or continuous pesticide applications to the water layer;
- implementation of varying water depths and rates of discharge in the water layer, coupled to incoming water via entry routes like upward seepage and surface runoff; and
- making the presentation of the model output more user-friendly.

The sensitivity of the model to input parameters, initial and boundary conditions needs to be studied. This can guide further experimental work as well as future use of the TOXSWA model, e.g. when standard scenarios are defined.

Standard scenarios need to be defined to facilitate the use of the model for risk assessments in the Dutch registration procedure. The standard scenarios should comprise the environmental conditions, such as water depth, flow velocity, mass of macrophytes, as well as the pesticide masses entering the ditch by one or several entry routes.

It should be demonstrated that the TOXSWA model simulates correctly real field situations, which means that the model needs to go through a validation process. The validation process is defined as a comparison of the model output with data independently derived from experiments or observations of the environment; this implies that none of the input parameters is obtained via calibration. Validation should be done for a specified range of validity, i.e. that part of reality to which the validation of a model applies (Leaching Modelling Workgroup FOCUS, 1995). Four experiments performed at the DLO Winand Staring Centre provide data sets for stagnant water bodies and one slowly moving water body, and these will be used first. Subsequently, data sets from elsewhere need to be assessed as to their potential use for the validation process. More experiments are probably needed to validate the model, especially in systems with a certain flow velocity.

Preliminary results of a sensitivity analysis for the water subsystem showed that the dispersion coefficient is a very important coefficient to describe well the (longitudinal) mixing of the substance in the watercourse. More experiments need to be performed to estimate the dispersion coefficient in various ditches under various conditions. Very few data exist, especially for small watercourses.

The process descriptions need to be worked out in greater detail. Transformation rates in water and sediment need to be studied more closely, as well as the influence on

these rates of pH, light intensity, bioactivity and temperature conditions. Furthermore, the effect of aerobic and anaerobic conditions on the rate of transformation of the substance in the upper few millimetres below the water-sediment interface needs further study. Very little information is available on sorption to macrophytes, although this process clearly affects the level and duration of the pesticide concentration. More experimental work is needed to determine this sorption coefficient. A study on the sorption of a range of pesticides to three plant species is currently being carried out at the DLO Winand Staring Centre.

In order to assess whether the TOXSWA model is able to simulate the fate of pesticides in watercourses in other countries of the EU, field experiments need to be executed outside the Netherlands. At the moment, the TOXSWA model is being assessed by the surface water fate group of the EU working party called FOCUS, FORum for the Co-ordination of pesticide fate models and their USE.

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List of symbols

A_x	= cross-sectional area of flow at location x	(L ²)
b	= width of ditch bottom	(L)
$b_d(t,x)$	= distributed source with continuous input	(M.L ⁻³ .T ⁻¹)
$b_p(t,x)$	= point source with continuous input	(M.L ⁻³ .T ⁻¹)
c	= mass concentration of substance in the water phase	(M.L ⁻³)
c^*	= mass concentration substance in water layer (this includes substance sorbed to suspended solids and to macrophytes)	(M.L ⁻³)
c_a	= mass concentration of substance in the air	(M.L ⁻³)
$c_{a,l}$	= equilibrium mass concentration of substance at the water-gas interface in the gas phase	(M.L ⁻³)
c_b^*	= mass concentration of substance in sediment	(M.L ⁻³)
$c_{e,ss}$	= concentration c , at which $K_{F,ss}$ has been estimated	(M.L ⁻³)
$c_{e,wb}$	= concentration c , at which $K_{F,wb}$ has been estimated	(M.L ⁻³)
c_l	= equilibrium mass concentration of substance at the water-gas interface in the water phase	(M.L ⁻³)
c_{lb}	= mass concentration of substance in the liquid phase of sediment	(M.L ⁻³)
c_{sol}	= solubility of substance in water	(M.L ⁻³)
d	= distance of water-sediment interface to concerned area	(L)
D_{lb}	= diffusion coefficient of substance in the liquid phase of sediment	(L ² .T ⁻¹)
D_w	= diffusion coefficient of substance in water	(L ² .T ⁻¹)
DW	= dry weight of macrophytes per area of sediment	(M.L ⁻²)
E_{lb}	= dispersion coefficient in pore water	(L ² .T ⁻¹)
E_x	= dispersion coefficient in direction of flow	(L ² .T ⁻¹)
h	= water level above ditch bottom	(L)
h_w	= water level above ditch bottom, defining the exchanging perimeter $P_{z=0}$	(L)
i	= index of grid points (segments) in space in x direction	(1)
j	= index of grid points in time	(1)
J	= areic mass flux of substance in water layer by advection and dispersion	(M.L ⁻² .T ⁻¹)
J_{lb}	= areic mass flux of substance in the liquid phase of the sediment by advection, dispersion and diffusion	(M.L ⁻² .T ⁻¹)
J_{wa}	= areic mass flux of substance across the water-air interface, the flux is negative in upward direction	(M.L ⁻² .T ⁻¹)
J_{wb}	= areic mass flux of substance across the water-sediment interface, the flux is positive in downward direction	(M.L ⁻² .T ⁻¹)
$J_{wb,adv}$	= areic mass flux by advection at the water-sediment interface	(M.L ⁻² .T ⁻¹)
$J_{wb,dif}$	= areic mass flux by diffusion at the water-sediment interface	(M.L ⁻² .T ⁻¹)
k	= index of grid points (segments) in space in z direction	(1)

k	= transformation rate coefficient for substance in the water column	(T ⁻¹)
k_b	= transformation rate coefficient for substance in the sediment	(T ⁻¹)
k_g	= exchange coefficient of substance in the gas phase	(L.T ⁻¹)
k_l	= exchange coefficient of substance in the liquid phase	(L.T ⁻¹)
$k_{t,l}$	= overall transfer coefficient for the air-water interface, based at the liquid phase	(L.T ⁻¹)
K_{mp}	= distribution coefficient for substance between macrophytes and water, i.e. slope of sorption isotherm based at the mass of dry macrophytes	(L ³ .M ⁻¹)
$K_{om,ss}$	= slope of sorption isotherm for suspended solids, based at the organic matter content	(L ³ .M ⁻¹)
$K_{om,wb}$	= slope of sorption isotherm for sediment, based at the organic matter content	(L ³ .M ⁻¹)
$K_{F,ss}$	= Freundlich coefficient for sorption to suspended solids	(L ³ .M ⁻¹)
$K_{F,wb}$	= Freundlich coefficient for sorption to bottom material	(L ³ .M ⁻¹)
K_H	= dimensionless Henry coefficient	(1)
$K_{L,b}$	= slope of (linear) sorption isotherm of bottom material	(L ³ .M ⁻¹)
$K_{L,ss}$	= slope of (linear) sorption isotherm of suspended solids	(L ³ .M ⁻¹)
L_{dis}	= dispersion length	(L)
LBD	= band above the diagonal of the left-hand tridiagonal matrix	(1)
LDD	= diagonal of the left-hand tridiagonal matrix	(1)
LOD	= band under the diagonal of the left-hand tridiagonal matrix	(1)
l	= length of drained or infiltrated lot, oriented perpendicular to the ditch and extended on one or two sides of the ditch	(L)
$m_{om,ss}$	= mass fraction of organic matter of the suspended solids	(M.M ⁻¹)
$m_{om,wb}$	= mass fraction of organic matter of the sediment material	(M.M ⁻¹)
M	= applied area-averaged mass of substance	(M.L ²)
M_m	= molecular mass	(M.N ⁻¹)
n_{ss}	= Freundlich exponent for sorption to suspended solids	(1)
n_{wb}	= Freundlich exponent for sorption to bottom material	(1)
NWB	= Numerical Weight factor sediment [water Bottom]	(1)
NWW	= Numerical Weight factor Water layer	(1)
O_x	= width water surface at location x	(L)
$p_d(t,x)$	= distributed pulse input	(M.L ⁻³ .T ⁻¹)
$p_p(t,x)$	= point-type pulse input	(M.L ⁻³ .T ⁻¹)
P	= saturated vapour pressure of substance	(L ⁻¹ .M.T ⁻²)
P_0	= wetted perimeter	(L)
P_d	= length wetted perimeter at distance d from the water-sediment interface	(L)
P_x	= wetted perimeter at location x	(L)
$P_{z=0}$	= length wetted perimeter at depth $z = 0$	(L)

q	= areic volume flux, i.e. volume drained or supplied water divided by lot area and time. The flux is positive with infiltration and negative with upward flow (drainage from the field lot)	$(L^3.L^{-2}.T^{-1})$
Q	= rate of discharge in water layer	$(L^3.T^{-1})$
r	= number of point-type inputs (total n inputs)	
R	= universal gas constant	$(L^2.M.T^{-2}.N^{-1}.\theta^{-1})$
R_{wb}	= retardation factor for the sediment subsystem	(1)
R_{wl}	= retardation factor for the water subsystem	(1)
RBD	= band above the diagonal of the right-hand tridiagonal matrix	(1)
RDD	= diagonal of the right-hand tridiagonal matrix	(1)
ROD	= band under the diagonal of the right-hand tridiagonal matrix	(1)
RV	= right-hand vector	(1)
s	= number pulse input (total m inputs)	
s_0	= number pulse input at location x_0 (total m_0 inputs)	
s_{00}	= number pulse input at location x_{00} (total m_{00} inputs)	
s_1	= side slope, horizontal/vertical	(1)
ss	= mass concentration of suspended solids in the water layer, i.e. the ratio of the mass of dry suspended solids divided by the volume of water	$(M.L^{-3})$
t	= time	(T)
t_{rs}	= time at which the continuous release at location x_r starts	(T)
t_{re}	= time at which the continuous release at location x_e ends	(T)
t_s	= time of pulse input	(T)
T	= temperature at which the saturated vapour pressure, the solubility and the exchange coefficients in the liquid and gas phases are defined	(θ)
u	= flow velocity water	$(L.T^{-1})$
w	= average flow velocity of pore water (i.e. $lq/P\epsilon$)	$(L.T^{-1})$
x	= downstream distance along ditch axis	(L)
x_0	= location pulse input	(L)
X_b	= content of substance sorbed, i.e. the ratio of the mass of substance sorbed divided by the mass of dry bottom material	$(M.M^{-1})$
X_{mp}	= content of substance sorbed to macrophytes, i.e. the ratio of the mass of substance sorbed divided by the mass of dry macrophytes	$(M.M^{-1})$
X_{ss}	= content of substance sorbed to suspended solids i.e. the ratio of the mass of substance sorbed divided by the mass of dry suspended solids	$(M.M^{-1})$
y	= direction perpendicular to the x and z axis	(L)
z	= depth under the water-sediment interface	(L)

α	= weight factor for space, applied in the numerical solution method	
β	= weight factor for space, applied in the numerical solution method	(1)
β	= $\arctan (1/s_1)$	(1)
δ	= Dirac delta function	(T ⁻¹ or L ⁻¹)
ε	= volume fraction of pore water, i.e. volume of liquid divided by volume of bottom material	(1)
$\zeta(t)$	= release time function defined by: 0 for $t < t_{rs}$ or $t > t_{re}$ 1 for $t_{rs} \leq t \leq t_{re}$	(1)
θ	= weight factor for time, applied in the numerical solution method	(1)
λ	= tortuosity factor, i.e. ratio of surface area of bottom material to liquid phase	(1)
ρ_b	= bulk density of dry bottom material, i.e. volumic mass of dry bottom material	(M.L ⁻³)

Annex 1 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq.(6.26) for the grid points $m + 1$ up to $ebt - 1$ inclusive in the end buffer of the water subsystem

$$LOD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$LDD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \cancel{SS K_F \left(\frac{c_i^{j+1} \tau_{e-1}}{c_e} \right)} \right)$$

$$+ \left(1 + \frac{DW \cdot P_{z=0}}{A_i^{j+1}} K_{mp} + \frac{SS K_F}{c_e} \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_n-1} \right) (A_i^{j+1} + k A_i^{j+1/2} (1-\theta) \Delta t)$$

$$+ k_{t,1} O_{xi}^{j+1/2} (1-\theta) \Delta t + \frac{l}{P_{z=0}} q_i^{j+1/2} P_{xi}^j (1-\theta) \Delta t$$

$$+ \frac{2\varepsilon_i^j D_b P_{xi}^j}{\Delta z_1} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} (1-\theta) NWW_{i+1/2} \Delta t \left(1 + \frac{SS K_F}{c_e} \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_n-1} \right)$$

$$- \frac{2A_{i+1/2}^{j+1/2} E_{x \ i+1/2}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + \frac{SS K_F}{c_e} \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_n-1} \right)$$

$$ROD = \frac{Q_{i-1/2}^{j+1/2}}{\Delta x_i} \theta (1 - NWW_{i-1/2}) \Delta t \left(1 + \frac{SS K_F}{c_e} \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_n-1} \right)$$

$$+ \frac{2A_{i-1/2}^{j+1/2} E_{x \ i-1/2}^{j+1/2}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + \frac{SS K_F}{c_e} \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_n-1} \right)$$

$$\begin{aligned}
RDD &= \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i-\frac{1}{2}} \Delta t \left(1 - \cancel{\frac{SS K_F}{\left(\frac{c_i^j}{c_e}\right)^{\gamma_{n-1}}}} \right) \\
&\quad - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 - \cancel{\frac{SS K_F}{\left(\frac{c_i^j}{c_e}\right)^{\gamma_{n-1}}}} \right) \\
&\quad - \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 - \cancel{\frac{SS K_F}{\left(\frac{c_i^j}{c_e}\right)^{\gamma_{n-1}}}} \right) \\
&\quad - \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - \cancel{\frac{SS K_F}{\left(\frac{c_i^j}{c_e}\right)^{\gamma_{n-1}}}} \right) \\
&\quad + \left(1 - \cancel{\frac{DW P_{z=0}}{A_i^j} K_{mp}} - \cancel{\frac{SS K_F}{\left(\frac{c_i^j}{c_e}\right)^{\gamma_{n-1}}}} \right) (A_i^j - \cancel{k A_i^{j+\frac{1}{2}} \theta \Delta t}) \\
&\quad - \cancel{\frac{k_{t,1} O_{xi}^{j+\frac{1}{2}} \theta \Delta t}{P_{z=0}}} - \cancel{\frac{l}{q_i^{j+\frac{1}{2}} P_{xi}^j \theta \Delta t}} \\
&\quad - \cancel{\frac{2\varepsilon^j D_b P_{xi}^j}{\Delta z_1} \theta \Delta t}
\end{aligned}$$

$$RBD = - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{K_F}{K_H} \left(\frac{c_{i+1}^j}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - \frac{K_F}{K_H} \left(\frac{c_{i+1}^j}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$RV = \frac{c_{a,i}^{j+\frac{1}{2}}}{K_H} \Delta t - \frac{\mathcal{E}_i^j D_{lb,i}^{j+\frac{1}{2}} P_{x,i}^j}{\Delta z_1 c_{lb,i,k=1}^j} \Delta t$$

So, the right-hand vector RV does not appear anymore.

Annex 2 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.26) for the grid point $eb = ebt$ of the end buffer of the water subsystem. The flow velocity u is positive

$$LOD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 - \frac{K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_u-1}}{\gamma_u} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 - \frac{K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_u-1}}{\gamma_u} \right)$$

$$LDD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 - \frac{K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_u-1}}{\gamma_u} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_u-1}}{\gamma_u} \right)$$

$$+ \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 - \frac{K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_u-1}}{\gamma_u} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{r_u-1} \right)$$

$$+ \left(1 + \frac{DW \cdot P}{A_i^{j+1} K_{mp}^{z=0}} + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{r_u-1} \right) (A_i^{j+1} - k A_i^{j+\frac{1}{2}} (1-\theta) \Delta t)$$

$$- k_{t-1} O_{x_i}^{j+\frac{1}{2}} (1-\theta) \Delta t + \frac{\ell}{P_{z=0}} q_i^{j+\frac{1}{2}} P_{x_i}^j (1-\theta) \Delta t$$

$$+ \frac{2\varepsilon_i^j D_w P_{x_i}^j}{\Delta z_1} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{r_u-1} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i-1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{r_u-1} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{r_u-1} \right)$$

$$+ \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 \leftarrow \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right)$$

$$RDD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i-\frac{1}{2}} \Delta t \left(1 \leftarrow \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right)$$

$$- \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot \left(1 \leftarrow NWW_{i+\frac{1}{2}} \right) \Delta t \left(1 \leftarrow \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 \leftarrow \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right)$$

$$+ \left(1 \leftarrow \frac{DW P_{z=0}}{A_i^j} K_{mp} \leftarrow \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}}{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_m-1}} \right) (A_i^j \leftarrow k A_i^{j+\frac{1}{2}} \theta \Delta t)$$

$$\frac{k_{t1} O_{xi}^{j+\frac{1}{2}} \theta \Delta t}{P_{z=0}} \frac{l}{q_i^{j+\frac{1}{2}} P_{xi}^j \theta \Delta t}$$

$$\frac{2\varepsilon_i^j D_{10} P_{x1}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta \cdot NWW_{i+1/2} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$+ \frac{2A_{i+1/2}^{j+1/2} E_{x_{i+1/2}}^{j+1/2}}{\Delta x_i (\Delta x_{i-1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$RV = k_{i,1} O_{x_i}^{j+1/2} \frac{c_{x1}^{j+1/2}}{K_H} \Delta t + 2 \frac{\varepsilon_i^j D_{10} P_{x1}^j}{\Delta z_1} c_{1b,i,k=1}^j \Delta t$$

So, the elements *LBD*, *RBD* and *RV* have disappeared in the segment *ebt* in the case of a positive flow velocity.

Annex 3 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.26) for the grid point $eb = ebt$ of the end buffer of the water subsystem. The flow velocity u is negative

$$LOD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\gamma_{e-1}}} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\gamma_{e-1}}} \right)$$

$$LDD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\gamma_{e-1}}} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i+\frac{1}{2}}) \Delta t \left(1 + \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\gamma_{e-1}}} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\gamma_{e-1}}} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$+ \left(1 + \frac{DW \cdot P_{z=0}}{A_i^{j+1}} K_{mp} \leftarrow ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_{n-1}} \right) \left(A_i^{j+1} \leftarrow k A_i^{j+\frac{1}{2}} (1-\theta) \Delta t \right)$$

$$\leftarrow k_{i,1} O_{x_i}^{j+\frac{1}{2}} (1-\theta) \Delta t \leftarrow \frac{\ell}{P_{z=0}} q_i^{j+\frac{1}{2}} P_{x_i}^j (1-\theta) \Delta t$$

$$\leftarrow \frac{2E_i^j D_{\theta} P_{x_i}^j}{\Delta z_1} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_{n-1}} \right)$$

$$+ \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right)$$

$$RDD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i-\frac{1}{2}} \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right)$$

$$\frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 + \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right)$$

$$+ \left(1 - \frac{DW P_{z=0}}{A_i^j} K_{mp} - \frac{ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}}}{1} \right) (A_i^j - k A_i^{j+\frac{1}{2}} \theta \Delta t)$$

$$\frac{k_{i-1} O_{xi}^{j+\frac{1}{2}} \theta \Delta t}{P_{z=0}} \frac{q_i^{j+\frac{1}{2}} P_{xi}^j \theta \Delta t}{P_{z=0}}$$

$$\frac{2\varepsilon_i^j D_{ib} P_{x_i}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta NWW_{i+\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x_{i+\frac{1}{2}}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$RV = k_{i,1} \frac{O_{xi}^{j+\frac{1}{2}} c_{a_i}^{j+\frac{1}{2}}}{K_H} \Delta t \leftarrow 2 \frac{\varepsilon_i^j D_{ib}^{j+\frac{1}{2}} P_{x_i}^j}{\Delta z_1} c_{ib, i, k=1}^j \Delta t$$

So, the elements *LBD*, *RBD* and *RV* have disappeared in the segment *ebt* in the case of a negative flow velocity.

Annex 4 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.26) for the grid point $i = 1$ in the water subsystem in the case of a positive flow direction

In the case of downward waterflow in the sediment subsystem:

$$LOD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{C_{i-1}^{j+1}}{C_e} \right)^{n-1} \right)$$

$$\frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{C_{i-1}^{j+1}}{C_e} \right)^{n-1} \right)$$

$$LDD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{n-1} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{n-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i+1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{n-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{n-1} \right)$$

$$+ \left(1 + \frac{DW \cdot P_{z=0}}{A_i^{j+1}} K_{mp} + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_m-1} \right) \left(A_i^{j+1} + k A_i^{j+1/2} (1 - \theta) \Delta t \right)$$

$$+ k_{\gamma_1} O_{x_i}^{j+1/2} (1 - \theta) \Delta t + \frac{\ell}{P_{z=0}} q_i^{j+1/2} P_{x_i}^j (1 - \theta) \Delta t$$

$$+ \frac{2\varepsilon_i^j D_{lb} P_{x_i}^j}{\Delta z_1} (1 - \theta) \Delta t$$

$$LBD = \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} (1 - \theta) NWW_{i+1/2} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_m-1} \right)$$

$$- \frac{2A_{i+1/2}^{j+1/2} E_{x_{i+1/2}}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1 - \theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^{j+1}}{c_e} \right)^{\gamma_m-1} \right)$$

$$ROD = \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta (1 - NWW_{i+1/2}) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_m-1} \right)$$

$$- \frac{2A_{i+1/2}^{j+1/2} E_{x_{i+1/2}}^{j+1/2}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_m-1} \right)$$

$$\begin{aligned}
RDD &= \frac{Q_{i-1/2}^{j+1/2}}{\Delta x_i} \theta \cdot NWW_{i-1/2} \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_n-1} \right) \\
&\quad - \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta \cdot (1 - NWW_{i+1/2}) \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_n-1} \right) \\
&\quad - \frac{2A_{i-1/2}^{j+1/2} E_{x \ i-1/2}^{j+1/2}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_n-1} \right) \\
&\quad - \frac{2A_{i+1/2}^{j+1/2} E_{x \ i+1/2}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_n-1} \right) \\
&\quad + \left(1 + \frac{DW P_{z=0}^j}{A_i^j} K_{mp} + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_n-1} \right) (A_i^j - kA_i^{j+1/2} \theta \Delta t) \\
&\quad - k_{t,1} O_{xi}^{j+1/2} \theta \Delta t - \frac{q}{P_{z=0}} q_i^{j+1/2} P_{xi}^j \theta \Delta t \\
&\quad - \frac{2\epsilon_i^j D_{lb} P_{xi}^j}{\Delta z_1} \theta \Delta t
\end{aligned}$$

$$RBD = - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i+\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_{i+1}^j}{c_e} \right)^{n-1} \right)$$

$$RV = k_{t,1} O_{x_i}^{j+\frac{1}{2}} \frac{c_{a,i}^{j+\frac{1}{2}}}{K_H} \Delta t + 2 \frac{\epsilon_i^j D_{lb,i}^{j+\frac{1}{2}} P_{x,i}^j}{\Delta z_1} c_{lb,i,k=1}^j \Delta t$$

In the case of upward water flow in the sediment subsystem the right-hand vector contains an additional term:

$$- \frac{\ell}{P_{z=0}} q_i^{j+\frac{1}{2}} c_{lb,i,k=1}^j P_{x,i}^{j+\frac{1}{2}} \Delta t$$

and the penultimate terms containing $q_i^{j+\frac{1}{2}}$ in the *LDD* and *RDD* elements disappear.

So in all cases ($q < 0$, $q > 0$) the elements *LOD* and *ROD* have disappeared in the first segment; *LBD*, *RBD* and *RV* remain unchanged in this case of a positive flow velocity and an incoming flux of zero in the water subsystem.

Annex 5 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.26) for the grid point $fb = 2$ up to $fb = fbt$ inclusive of the front buffer of the water subsystem. The flow velocity alternates between being positive and negative

$$LOD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 \leftarrow \overline{SS K_F \left(\frac{c_i^{j+1} \gamma_{i-1}}{c_e} \right)} \right)$$

$$- \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 \leftarrow \overline{SS K_F \left(\frac{c_i^{j+1} \gamma_{i-1}}{c_e} \right)} \right)$$

$$LDD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 \leftarrow \overline{SS K_F \left(\frac{c_i^{j+1} \gamma_{i-1}}{c_e} \right)} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 \leftarrow \overline{SS K_F \left(\frac{c_i^{j+1} \gamma_{i-1}}{c_e} \right)} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x \ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 \leftarrow \overline{SS K_F \left(\frac{c_i^{j+1} \gamma_{i-1}}{c_e} \right)} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \frac{SS K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{n-1}}{1} \right)$$

$$+ \left(1 - \frac{DW \cdot P_{z=0}}{A_i^{j+1} K_{mp}} - \frac{SS K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{n-1}}{1} \right) \left(A_i^{j+1} - k A_i^{j+1} (1-\theta) \Delta t \right)$$

$$+ \frac{k_{x1} \theta_{xi}^{j+1} (1-\theta) \Delta t}{P_{z=0}} - \frac{q_i^{j+1} P_{xi}^j (1-\theta) \Delta t}{P_{z=0}}$$

$$+ \frac{2\epsilon_i^j D_w P_{xi}^j}{\Delta z_1} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{SS K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{n-1}}{1} \right)$$

$$- \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \frac{SS K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{n-1}}{1} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 - \frac{SS K_F \left(\frac{c_i^j}{c_e} \right)^{n-1}}{1} \right)$$

$$\begin{aligned}
& + \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
RDD = & \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i-\frac{1}{2}} \Delta t \left(1 \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& + \left(1 \leftarrow \frac{DW P_{z=0}}{A_i^j} K_{mp} \leftarrow \cancel{\text{SS } K_F} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) (A_i^j \leftarrow \cancel{k A_i^{j+\frac{1}{2}}} \theta \Delta t) \\
& \cancel{k_{i-1} O_{xi}^{j+\frac{1}{2}} \theta \Delta t} - \frac{l}{P_{z=0}} q_i^{j+\frac{1}{2}} P_{xi}^j \theta \Delta t
\end{aligned}$$

$$\frac{2\varepsilon_i^j D_{10} P_{x_i}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{\Delta t}{\Delta x_i} K_F \left(\frac{c_{i+1}^j}{c_c} \right)^{\gamma_n-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x_{i+\frac{1}{2}}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - \frac{\Delta t}{\Delta x_i} K_F \left(\frac{c_{i+1}^j}{c_c} \right)^{\gamma_n-1} \right)$$

$$RV \left[\frac{c_{a_i}^{j+\frac{1}{2}}}{K_H} \Delta t \right] \left[\frac{\varepsilon_i^j D_{10} P_{x_i}^j}{\Delta z_1} c_{lb, i, k=1}^j \Delta t \right]$$

So, the vector RV does not appear anymore.

Annex 6 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.26) for the grid point $fb = 1$ in the case of a negative flow velocity in the situation of alternating positive and negative flow velocities in the water layer

$$LOD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_w-1} \right)$$

$$\frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_w-1} \right)$$

$$LDD = - \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 - ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_w-1} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 - ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_w-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\gamma_w-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\nu_e-1}} \right)$$

$$+ \left(1 - \cancel{\frac{DW \cdot P_{z=0}}{A_i^{j+1}} K_{mp} + ss K_F \left(\frac{C_i^{j+1}}{C_e} \right)^{\nu_e-1}} \right) \left(A_i^{j+1} - \cancel{k A_i^{j+1}} (1-\theta) \Delta t \right)$$

$$\cancel{k_{l,1} O_{xi}^{j+1/2}} (1-\theta) \Delta t - \frac{\ell}{P_{z=0}} \cancel{q_i^{j+1/2} P_{xi}^j} (1-\theta) \Delta t$$

$$\cancel{\frac{2\varepsilon_i^j D_w P_{xi}^j}{\Delta z_1}} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_{i+1}^{j+1}}{C_e} \right)^{\nu_e-1}} \right)$$

$$- \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \cancel{ss K_F \left(\frac{C_{i+1}^{j+1}}{C_e} \right)^{\nu_e-1}} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 + \cancel{ss K_F \left(\frac{C_{i-1}^j}{C_e} \right)^{\nu_e-1}} \right)$$

$$\begin{aligned}
& + \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^{j-1}}{c_e} \right)^{\gamma_{n-1}} \right) \\
RDD = & \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot \Delta t \left(1 - NWW_{i-\frac{1}{2}} \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 - ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_x^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& + \left(1 - \frac{DW P_{z=0}}{A_i^j} K_{mp} - ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) (A_i^j - \kappa A_i^{j-\frac{1}{2}} \theta \Delta t) \\
& - \frac{\kappa_{i,1} O_{xi}^{j+\frac{1}{2}} \theta \Delta t}{P_{z=0}} - \frac{\rho}{P_{z=0}} q_i^{j+\frac{1}{2}} P_{xi}^j \theta \Delta t
\end{aligned}$$

$$\frac{2\varepsilon_i^j D_{10} P_{x_i}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = - \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta \cdot NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{K_F}{K_F} \left(\frac{c_{i+1}^j}{c_e} \right)^{r_e-1} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x_{i+\frac{1}{2}}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - \frac{K_F}{K_F} \left(\frac{c_{i+1}^j}{c_e} \right)^{r_e-1} \right)$$

$$RV \left[\frac{k_{t,1} O_{x_i}^{j+\frac{1}{2}} c_{a_i}^{i+\frac{1}{2}}}{K_H} \Delta t - \frac{2 \varepsilon_i^j D_{10} P_{x_i}^j}{\Delta z_1} c_{lb, i, k=1}^j \Delta t \right]$$

The elements *LOD*, *ROD* and *RV* have disappeared in this case.

Annex 7 The elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, *RBD* and *RV* from Eq. (6.26) for the grid point $fb = 1$ in the case of a positive flow velocity in the situation of alternating positive and negative flow velocities in the water layer

$$LOD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1-NWW_{i-\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^{j+\frac{1}{2}}}{c_e} \right)^{\gamma_{e-1}} \right)$$

$$\frac{2A_{i-\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^{j+\frac{1}{2}}}{c_e} \right)^{\gamma_{e-1}} \right)$$

$$LDD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i-\frac{1}{2}} \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^{j+\frac{1}{2}}}{c_e} \right)^{\gamma_{e-1}} \right)$$

$$+ \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) (1 - NWW_{i+\frac{1}{2}}) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+\frac{1}{2}}}{c_e} \right)^{\gamma_{e-1}} \right)$$

$$\frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x, i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} (1-\theta) \Delta t \left(1 + ss K_F \left(\frac{c_i^{j+\frac{1}{2}}}{c_e} \right)^{\gamma_{e-1}} \right)$$

$$+ \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}}{\left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}} \right)$$

$$+ \left(1 + \frac{DW \cdot P}{A_i^{j+1} K_{mp} z=0} + \frac{ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}}{\left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}} \right) \left(A_i^{j+1} - k A_i^{j+\frac{1}{2}} (1-\theta) \Delta t \right)$$

$$- \frac{k_{t-1} O_{xi}^{j+\frac{1}{2}} (1-\theta) \Delta t}{P_{z=0} q_i^{j+\frac{1}{2}} P_{xi}^j (1-\theta) \Delta t}$$

$$- \frac{2\varepsilon_i^j D_w P_{xi}^j}{\Delta z_1} (1-\theta) \Delta t$$

$$LBD = \frac{Q_{i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} (1-\theta) NWW_{i+\frac{1}{2}} \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}}{\left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}} \right)$$

$$- \frac{2A_{i+\frac{1}{2}}^{j+\frac{1}{2}} E_{x\ i+\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} (1-\theta) \Delta t \left(1 - \frac{ss K_F \left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}}{\left(\frac{c_i^{j+1}}{c_e} \right)^{\nu_w-1}} \right)$$

$$ROD = \frac{Q_{i-\frac{1}{2}}^{j+\frac{1}{2}}}{\Delta x_i} \theta (1 - NWW_{i-\frac{1}{2}}) \Delta t \left(1 + \frac{ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{\nu_w-1}}{\left(\frac{c_{i-1}^j}{c_e} \right)^{\nu_w-1}} \right)$$

$$\begin{aligned}
& \frac{2A_{i-1/2}^{j+1/2} E_{x_{i-1/2}}^{j+1/2}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_{i-1}^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
RDD = & \frac{Q_{i-1/2}^{j+1/2}}{\Delta x_i} \theta \cdot NWW_{i-1/2} \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta \cdot (1 - NWW_{i+1/2}) \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& \frac{2A_{i-1/2}^{j+1/2} E_{x_{i-1/2}}^{j+1/2}}{\Delta x_i (\Delta x_i + \Delta x_{i-1})} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& - \frac{2A_{i+1/2}^{j+1/2} E_{x_{i+1/2}}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) \\
& + \left(1 + \frac{DWP_{z=0}}{A_i^j} K_{mp} + ss K_F \left(\frac{c_i^j}{c_e} \right)^{\gamma_{n-1}} \right) (A_i^j - \kappa A_i^{j-n} \theta \Delta t) \\
& - \frac{\kappa_{x,1} O_{x_i}^{j+1/2} \theta \Delta t}{P_{z=0}} - \frac{\rho}{P_{z=0}} q_i^{j-n} P_{x_i}^j \theta \Delta t
\end{aligned}$$

$$\frac{2\varepsilon_i^j D_{10} P_{x1}^j}{\Delta z_1} \theta \Delta t$$

$$RBD = - \frac{Q_{i+1/2}^{j+1/2}}{\Delta x_i} \theta \cdot NWW_{i+1/2} \Delta t \left(1 - \frac{SS K_F}{c_e} \left(\frac{c_{i+1}^j}{c_e} \right)^{\gamma_m - 1} \right)$$

$$+ \frac{2A_{i+1/2}^{j+1/2} E_{x i+1/2}^{j+1/2}}{\Delta x_i (\Delta x_{i+1} + \Delta x_i)} \theta \cdot \Delta t \left(1 - \frac{SS K_F}{c_e} \left(\frac{c_{i+1}^j}{c_e} \right)^{\gamma_m - 1} \right)$$

$$RV = \frac{c_{a1}^{j+1/2}}{K_H} O_{xi}^{j+1/2} \Delta t - \frac{\varepsilon_i^j D_{10} P_{x1}^j}{\Delta z_1} c_{lb i, k=1}^j \Delta t$$

So, the elements *LOD*, *ROD* and *RV* have disappeared in segment $fb = 1$ in the case of a positive flow velocity in the situation of alternating positive and negative flow velocities in the water layer.

Annex 8 The elements LOD , LDD , LBD , ROD , RDD , RBD and RV from Eq. (6.39) for the grid point $k = 1$ of the sediment subsystem. So, this is the upper boundary condition

In the case of downward seepage ($q > 0$):

LOD disappears

$$\begin{aligned}
 LDD &= \frac{q q^{j+1/2}}{\Delta z_1} (1 - \theta) (1 - NWB_{1/2}) \frac{\Delta t}{P_1^j} sofd_{1/2} \\
 &+ \frac{P_{1/2}^j \epsilon_{1/2}^j (E_{lb\ 1/2}^{j+1/2} + D_{lb\ 1/2}^j)}{(\frac{1}{2} \Delta z_2 + \frac{1}{2} \Delta z_1) \Delta z_1} (1 - \theta) \frac{\Delta t}{P_1^j} \\
 &+ \frac{P_{1/2}^j \epsilon_{1/2}^j D_{lb\ 1/2}^j}{\frac{1}{2} (\Delta z_1)^2} (1 - \theta) \frac{\Delta t}{P_1^j} \\
 &+ \left(\epsilon_1^j + \rho_{b1}^j K_F \left(\frac{c_{lb\ 1}^{j+1}}{c_c} \right)^{r_w-1} \right) (1 + k_b (1 - \theta) \Delta t)
 \end{aligned}$$

$$LBD = \frac{q \cdot q^{j+1/2}}{\Delta z_1} (1 - \theta) NWB_{1/2} \frac{\Delta t}{P_1^j} sofd_{1/2} - \frac{P_{1/2}^j \epsilon_{1/2}^j (E_{lb\ 1/2}^{j+1/2} + D_{lb\ 1/2}^j)}{(\frac{1}{2} \Delta z_2 + \frac{1}{2} \Delta z_1) \Delta z_1} (1 - \theta) \frac{\Delta t}{P_1^j}$$

ROD disappears

$$\begin{aligned}
RDD = & - \frac{\varrho \cdot q^{j+1/2}}{\Delta z_1} \theta (1 - NWB_{1/2}) \frac{\Delta t}{P_1^j} \text{sofd}_{1/2} - \frac{P_{1/2}^j \varepsilon_{1/2}^j (E_{1b \ 1/2}^{j+1/2} + D_{1b \ 1/2}^j)}{(\frac{1}{2}\Delta z_2 + \frac{1}{2}\Delta z_1) \Delta z_1} \theta \frac{\Delta t}{P_1^j} \\
& - \frac{P_{1/2}^j \varepsilon_{1/2}^j D_{1b \ 1/2}^j \theta}{\frac{1}{2}(\Delta z_1)^2} \frac{\Delta t}{P_1^j} \\
& + \left(\varepsilon_1^j + \rho_{b \ 1}^j K_F \left(\frac{c_{1b \ 1}^j}{c_e} \right)^{\eta_{\text{nb}} - 1} \right) (1 - k_b \theta \Delta t)
\end{aligned}$$

$$RBD = - \frac{\varrho \cdot q^{j+1/2}}{\Delta z_1} \theta NWB_{1/2} \frac{\Delta t}{P_1^j} \text{sofd}_{1/2} + \frac{P_{1/2}^j \varepsilon_{1/2}^j (E_{1b \ 1/2}^{j+1/2} + D_{1b \ 1/2}^j)}{(\frac{1}{2}\Delta z_2 + \frac{1}{2}\Delta z_1) \Delta z_1} \cdot \theta \cdot \frac{\Delta t}{P_1^j}$$

On the right-hand side the following terms appear:

$$\begin{aligned}
& c_i^j \left(\frac{\varrho \cdot q^{j+1/2}}{\Delta z_1} \theta \cdot \frac{\Delta t}{P_1^j} + \frac{P_{1/2}^j \varepsilon_{1/2}^j D_{1b \ 1/2}^j}{\frac{1}{2}(\Delta z_1)^2} \theta \frac{\Delta t}{P_1^j} \right) \\
& + c_i^{j+1} \left(\frac{\varrho \cdot q^{j+1/2}}{\Delta z_1} (1-\theta) \frac{\Delta t}{P_1^j} + \frac{P_{1/2}^j \varepsilon_{1/2}^j D_{1b \ 1/2}^j}{\frac{1}{2}(\Delta z_1)^2} (1-\theta) \frac{\Delta t}{P_1^j} \right)
\end{aligned}$$

This implies that a right-hand vector is added to Eq. (6.39).

In the case of upward seepage ($q < 0$):

On the right-hand side the following terms appear:

$$c_i^j \left(\frac{P_{1/2}^j \varepsilon_{1/2}^j D_{1b \ 1/2}^j}{\frac{1}{2}(\Delta z_1)^2} \theta \frac{\Delta t}{P_1^j} \right)$$

$$+ c_i^{j+1} \left(\frac{P_{1/2}^j \varepsilon_{1/2}^j D_{1b\ 1/2}^j}{1/2(\Delta z_1)^2} (1-\theta) \frac{\Delta t}{P_1^j} \right)$$

The term

$$\left(\frac{q \cdot q^{j+1/2}}{\Delta z_1} \theta \frac{\Delta t}{P_1^j} \right)$$

is added to the element *RDD*, and the term

$$- \left(\frac{q \cdot q^{j+1/2}}{\Delta z_1} (1-\theta) \frac{\Delta t}{P_1^j} \right)$$

to the element *LDD*.

So, a right-hand vector is added to Eq. (6.39) (analogously to Eq. (6.26) for the water layer). The first element of this vector is composed of four (in case of $q > 0$) or two terms (in case of $q < 0$), which are described above and the other elements are zero.

Annex 9 The elements *LOD*, *LDD*, *LBD*, *ROD*, *RDD*, and *RBD* from Eq. (6.39) for the grid points $n + 1$ up to $ebbt - 1$ inclusive from the end buffer at the lower end of the sediment

$$LOD = - \frac{q q^{j+1/2}}{\Delta z_k} (1 - \theta) (1 - NWB_{k-1/2}) \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$- \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$LDD = \frac{q q^{j+1/2}}{\Delta z_k} (1 - \theta) (1 - NWB_{k+1/2}) \frac{\Delta t}{P_k^j} sofd_{k+1/2}$$

$$+ \frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$- \frac{q q^{j+1/2}}{\Delta z_k} (1 - \theta) NWB_{k-1/2} \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$+ \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\epsilon_k^j \leftarrow P_{b k}^j K_F \left(\frac{C_{lb k}^{j+1} r_{w-1}^{j+1}}{C_c} \right) \right) \left(1 \leftarrow K_b P_k^j (1 - \theta) \frac{\Delta t}{P_k^j} \right)$$

$$LBD = \frac{q q^{j+1/2}}{\Delta z_k} (1 - \theta) NWB_{k+1/2} \frac{\Delta t}{P_k^j} \text{sofd}_{k+1/2}$$

$$- \frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb k+1/2}^{j+1/2} + D_{lb k+1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$ROD = \frac{q q^{j+1/2}}{\Delta z_k} \theta (1 - NWB_{k-1/2}) \frac{\Delta t}{P_k^j} \text{sofd}_{k-1/2}$$

$$+ \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb k-1/2}^{j+1/2} + D_{lb k-1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$RDD = - \frac{q q^{j+1/2}}{\Delta z_k} \theta (1 - NWB_{k+1/2}) \frac{\Delta t}{P_k^j} \text{sofd}_{k+1/2}$$

$$- \frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb k+1/2}^{j+1/2} + D_{lb k+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$+ \frac{q q^{j+1/2}}{\Delta z_k} \theta NWB_{k-1/2} \frac{\Delta t}{P_k^j} \text{sofd}_{k-1/2}$$

$$- \frac{P_{k-1/2}^j \varepsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\varepsilon_k^j \left[\frac{P_{bk}^j K_F}{P_{bk}^j K_F} \left(\frac{C_{bk}^j}{C_e} \right)^{P_{bk}^j - 1} \right] \right) \left(1 - \frac{K_b P_k^j \theta \Delta t}{P_k^j} \right)$$

$$RBD = - \frac{q q^{j+1/2}}{\Delta z_k} \theta NWB_{k+1/2} \frac{\Delta t}{P_k^j} \text{sofd}_{k+1/2}$$

$$+ \frac{P_{k+1/2}^j \varepsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

Annex 10 The lower boundary condition at the grid point *ebbt* in the case of downward water flow in the sediment

$$LOD = - \frac{\ell q^{j+1/2}}{\Delta z_k} (1 - \theta) (1 - NWB_{k-1/2}) \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$- \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$LDD = \frac{\ell q^{j+1/2}}{\Delta z_k} (1 - \theta) (\cancel{1 - NWB_{k+1/2}}) \frac{\Delta t}{P_k^j}$$

$$\frac{\cancel{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) (1 - \theta) \Delta t}}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$- \frac{\ell q^{j+1/2}}{\Delta z_k} (1 - \theta) NWB_{k-1/2} \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$+ \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\epsilon_k^j \cancel{\rho_{bk}^j K_F \left(\frac{c_{lb\ k}^{j+1/2}}{c_e} \right)} \right) \left(1 - \cancel{k_b P_k^j} (1 - \theta) \frac{\Delta t}{P_k^j} \right)$$

$$LBD = \frac{l q^{j+1/2}}{\Delta z_k} (1 - \theta) NWB_{k+1/2} \frac{\Delta t}{P_k^j}$$

$$= \frac{P_{k-1/2}^j \epsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$ROD = \frac{l q^{j+1/2}}{\Delta z_k} \theta (1 - NWB_{k-1/2}) \frac{\Delta t}{P_k^j} \text{sofd}_{k-1/2}$$

$$+ \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$RDD = - \frac{l q^{j+1/2}}{\Delta z_k} \theta (1 - NWB_{k+1/2}) \frac{\Delta t}{P_k^j}$$

$$= \frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$+ \frac{l q^{j+1/2}}{\Delta z_k} \theta NWB_{k-1/2} \frac{\Delta t}{P_k^j} \text{sofd}_{k-1/2}$$

$$- \frac{P_{k-1/2}^j \epsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\cancel{\frac{e_k^j}{P_k^j} \cancel{P_{bk}^j} \cancel{K_F} \left(\frac{\cancel{e_{1bk}^j}}{c_c} \right)^{P_{bk}^j - 1}} \right) \left(1 - \cancel{k_b P_k^j \theta \frac{\Delta t}{P_k^j}} \right)$$

$$RBD = \cancel{\frac{0 q^{j+1/2}}{\Delta z_k} \theta NWB_{k+1/2} \frac{\Delta t}{P_k^j}}$$

$$\cancel{\frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{1bk+1/2}^{j+1/2} D_{1bk+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}}$$

So, the elements *LBD* and *RBD* disappear.

Annex 11 The lower boundary condition at the grid point ebb in the case of upward water flow in the sediment

$$LOD = - \frac{\varrho q^{j+1/2}}{\Delta z_k} (1 - \theta) (1 - NWB_{k-1/2}) \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$- \frac{P_{k-1/2}^j \varepsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$LDD = \frac{\varrho q^{j+1/2}}{\Delta z_k} (1 - \theta) (1 - NWB_{k+1/2}) \frac{\Delta t}{P_k^j}$$

$$\frac{P_{k+1/2}^j \varepsilon_{k+1/2}^j (E_{lb\ k+1/2}^{j+1/2} + D_{lb\ k+1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$- \frac{\varrho q^{j+1/2}}{\Delta z_k} (1 - \theta) NWB_{k-1/2} \frac{\Delta t}{P_k^j} sofd_{k-1/2}$$

$$+ \frac{P_{k-1/2}^j \varepsilon_{k-1/2}^j (E_{lb\ k-1/2}^{j+1/2} + D_{lb\ k-1/2}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\varepsilon_k^j \left[\frac{c_{lb\ k}^{j+1} \gamma_{m-1}}{c_e} \right] \right) \left(1 - \frac{k_b P_k^j (1 - \theta) \Delta t}{P_k^j} \right)$$

$$LBD = \frac{\ell q^{j+\frac{1}{2}}}{\Delta z_k} (1 - \theta) NWB_{k+\frac{1}{2}} \frac{\Delta t}{P_k^j}$$

$$\frac{P_{k+\frac{1}{2}}^j \epsilon_{k+\frac{1}{2}}^j (E_{lb\ k+\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}^j) (1 - \theta) \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$ROD = \frac{\ell q^{j+\frac{1}{2}}}{\Delta z_k} \theta (1 - NWB_{k-\frac{1}{2}}) \frac{\Delta t}{P_k^j} sofd_{k-\frac{1}{2}}$$

$$+ \frac{P_{k-\frac{1}{2}}^j \epsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$RDD = \frac{\ell q^{j+\frac{1}{2}}}{\Delta z_k} \theta (1 - NWB_{k+\frac{1}{2}}) \frac{\Delta t}{P_k^j}$$

$$\frac{P_{k+\frac{1}{2}}^j \epsilon_{k+\frac{1}{2}}^j (E_{lb\ k+\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k+\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

$$+ \frac{\ell q^{j+\frac{1}{2}}}{\Delta z_k} \theta NWB_{k-\frac{1}{2}} \frac{\Delta t}{P_k^j} sofd_{k-\frac{1}{2}}$$

$$- \frac{P_{k-\frac{1}{2}}^j \epsilon_{k-\frac{1}{2}}^j (E_{lb\ k-\frac{1}{2}}^{j+\frac{1}{2}} + D_{lb\ k-\frac{1}{2}}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_k + \frac{1}{2} \Delta z_{k-1}) P_k^j}$$

$$+ \left(\epsilon_k^j \left[\frac{P_{bk}^j K_F \left(\frac{c_{lbk}^j}{c_e} \right)^{\gamma_{wb}-1}}{P_k^j} \right] \right) \left(1 - \frac{K_b P_k^j \theta \Delta t}{P_k^j} \right)$$

$$RBD = \frac{\ell q^{j+1/2}}{\Delta z_k} \theta \frac{NWB_{k+1/2}}{P_k^j} \frac{\Delta t}{P_k^j}$$

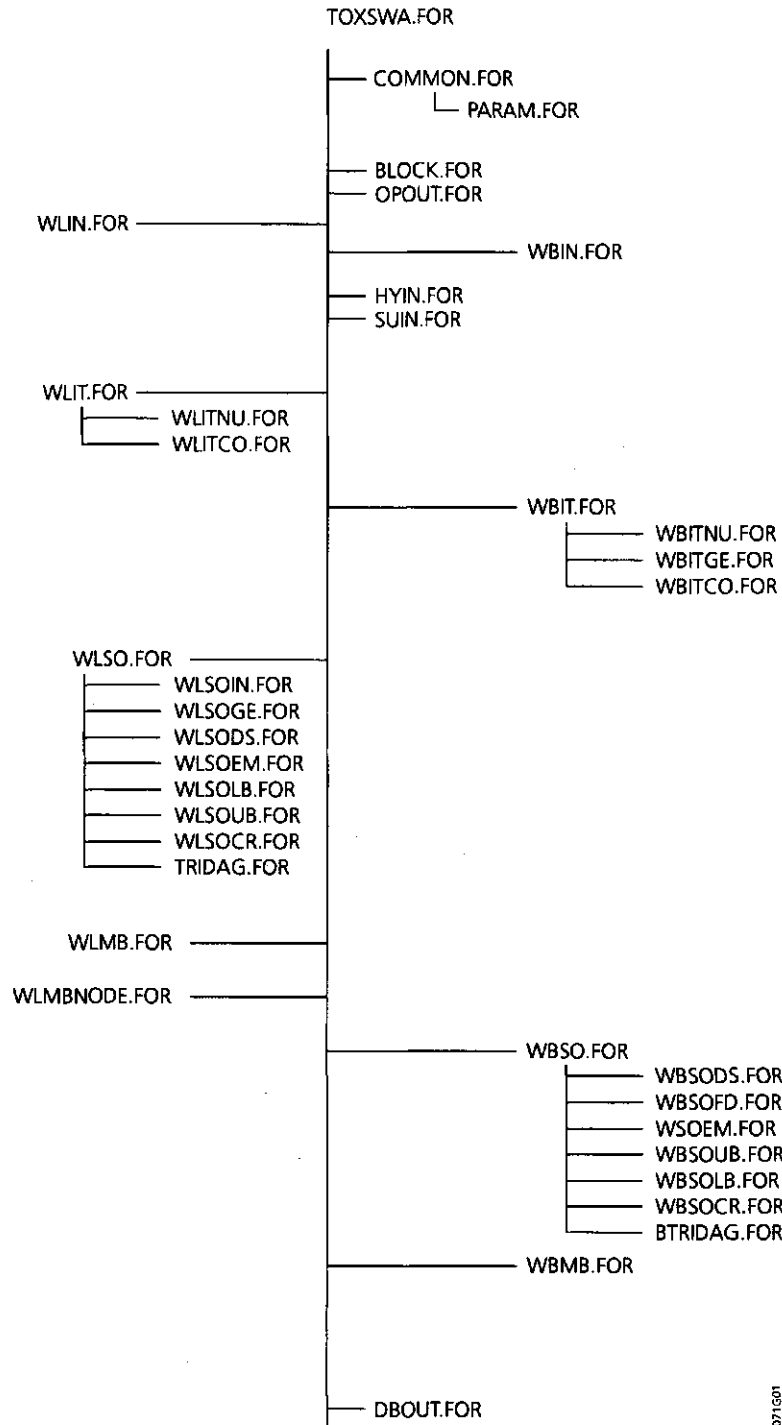
$$\frac{P_{k+1/2}^j \epsilon_{k+1/2}^j (E_{k+1/2}^{j+1/2} - D_{lbk+1/2}^j) \theta \Delta t}{\Delta z_k (\frac{1}{2} \Delta z_{k+1} + \frac{1}{2} \Delta z_k) P_k^j}$$

So, the elements *LBD* and *RBD* disappear. Two additional terms appear on the right-hand side of the numerical equations approximating the conservation equation, namely:

$$- \frac{\ell q^{j+1/2}}{\Delta z_{ebbt}} \theta c_{lb\ lot}^j \frac{\Delta t}{P_{ebbt}^j} - \frac{\ell q^{j+1/2}}{\Delta z_{ebbt}} (1-\theta) c_{lb\ lot}^{j+1} \frac{\Delta t}{P_{ebbt}^j}$$

This implies that, in the case of upward water flow, the last element of the right-hand vector *RV* is composed of the two terms described above.

Annex 12 Structure of the TOXSWA program



071601

Meaning of symbols used:

First pairs:

b bottom
db distribution
hy hydrology
op options
tri tri-diagonal
su substance
wb sediment (water bottom)
wl water layer

Second pairs:

dag diagonal
in input
it initial calculations
mb mass balance
out output
so solution

Third pairs:

co concentration
cr core of matrix
ds dispersion
em elements of matrix
fd flux: dispersion - advection
ge geometry
in input
lb lower boundary
nu numerical aspects
ub upper boundary

Annex 13 Description of the TOXSWA model including subroutines

TOXSWA.FOR

Main program. Initiates and controls time step and node handling. Terminates run.

COMMON.FOR

Declares all the common variables for the program TOXSWA.

PARAM.FOR

Defines parameters that are used for the declaration of the dimensions of certain arrays for the program. These are included in the common block.

BLOCK.FOR

Initialises identifiers for names of output files.

OPOUT.FOR

Reads all selected options concerning the output, such as desired output files, points of time for output or nodes for which detailed output is desired.

WLIN.FOR

Reads all input data for the water subsystem, concerning numerical aspects and length of ditch (from WLNU.INP), geometry and other characteristics of the ditch (from WLPA.INP) and initial concentrations (from WLST.INP). Data are converted to SI-units (m, s, g, Pa, K, J, mol).

WLIT.FOR

Calculates initial conditions for the water layer. Segment lengths in the ditch plus buffers are determined.

WLITNU.FOR

Calculates the numerical weight factors for the ditch and the buffers.

WLITCO.FOR

Calculates initial concentrations c , X_{mp} and X_{ss} for every node.

WBIN.FOR

Reads all input data for the sediment subsystem, concerning numerical aspects and thickness of the sediment (from WBNU.INP), properties of the sediment (from WBPA.INP) and initial concentrations (from WBST.INP). The input data are converted to SI-units (m, s, g, Pa, K, J, mol).

SUIN.FOR

Reads all data concerning substance properties, e.g. transformation rates in water and sediment, sorption parameters, solubility from SU.INP.

HYIN.FOR

Reads all hydrological input data for the water layer and sediment from HY.INP. All input data are converted to SI-units (m, s, g, Pa, K, J, mol).

WBIT.FOR

Calculates initial conditions for the sediment. The thickness of the segments in the sediment plus buffer are determined.

WBITNU.FOR

Calculates the numerical weight factors for the sediment and the end buffer.

WBITGE.FOR

Calculates initial conditions concerning perimeter, porosity, bulk density, Freundlich sorption coefficient and diffusion coefficient.

WBITCO.FOR

Calculates initial concentrations c_{1b} and X_b at every node.

WLSO.FOR

Solves the mass conservation equation for the water layer.

WLSOIN.FOR

Reads concentration for the first node of the sediment, for each node of the water layer, and seepage/infiltration per node.

WLSOGE.FOR

Calculates geometrical time-dependent characteristics.

WLSODS.FOR

Calculates dispersion coefficient; equals physical dispersion minus numerical dispersion.

WLSOEM.FOR

Defines elements of the matrix which are used in subroutines WLSOLB.FOR, WLSOUB.FOR and WLSOCR.FOR.

WLSOLB.FOR

Composes the elements of the matrix for the buffer at the lower boundary.

WLSOUB.FOR

Composes the elements of the matrix for the buffer at the upper boundary.

WLSOCR.FOR

Composes the core matrix, consisting of the numerical equations for solving the mass conservation equation.

TRIDAG.FOR

Inverts the left-hand tridiagonal matrix for the water layer.

WLMB.FOR

Checks the mass balance for the water layer.

WLMBNODE.FOR

Calculates the mass balance for selected segments of the water layer (to allow the provision of output at selected nodes).

WBSO.FOR

Solves the mass conservation equation for the sediment.

WBSODS.FOR

Calculates dispersion coefficient; equals physical dispersion minus numerical dispersion.

WBSOFD.FOR

Prohibits calculated dispersion flux from cancelling out advection flux.

WBSOEM.FOR

Defines elements of the matrix which are used in subroutines WBSOLB.FOR, WBSOUB.FOR and WBSOCR.FOR.

WBSOUB.FOR

Composes the elements of the matrix at the upper node of the sediment (upper boundary).

WBSOLB.FOR

Composes the elements of the matrix for the buffer at the lower boundary.

WBSOCR.FOR

Composes the core matrix, consisting of the numerical equations for solving the mass conservation equation.

BTRIDAG.FOR

Inverts the left-hand tridiagonal matrix for the sediment.

WBMB.FOR

Checks the mass balance for the sediment.

DBOUT.FOR

Calculates output concerning the pesticide mass distribution between the different compartments (water phase, macrophytes, suspended solids, liquid phase sediment and solid phase sediment) for the total ditch as well as per running metre.

Step 1:

Positivity

Condition (c) states that $R \geq 0$ and that for L , $l_{ii} > 0$ (diagonal elements) and $l_{ij} \leq 0$, $i \neq j$ (off-diagonal elements). Define a matrix $B = I - DL$, where I is the identity matrix and where D is a diagonal matrix with $d_{ii} = 1/l_{ii}$. We now assert that $B \geq 0$ and $\rho(B) < 1$. We illustrate this assertion for the three-dimensional case below. Theorem 3.10 of Varga (1962) states that (i) L^{-1} exists and $L^{-1} \geq 0$, corresponds with (ii) $B \geq 0$ and $\rho(B) < 1$ (for B defined as above and for L with $l_{ii} > 0$). It now follows from Theorem 3.10 that L^{-1} exists and $L^{-1} \geq 0$. As R is also positive, $L^{-1}R \geq 0$.

Illustration for the three-dimensional case:

Write $a = LDD > 0$, $-b = LBD \leq 0$, $-c = LOD \leq 0$, so that $a, b, c \geq 0$ under condition (c).

$$L = \begin{pmatrix} a & -b & 0 \\ -c & a & -b \\ 0 & -c & a \end{pmatrix}$$

From the definition for LOD , LDD and LBD in section 7.3 we find easily: $a > b$, $a > c$ and $a^2 > 2bc$. We find

$$B = \begin{pmatrix} 0 & \frac{b}{a} & 0 \\ \frac{c}{a} & 0 & \frac{b}{a} \\ 0 & \frac{c}{a} & 0 \end{pmatrix}$$

for which clearly $B \geq 0$ and for which $\rho(B) = \sqrt[3]{(2bca^2)} < 1$.

Step 2:

Stability

To prove stability we make use of Theorem 3.13 of Varga (1962). Theorem 3.13 states that if $A = L - R$ with $L^{-1} \geq 0$ and $R \geq 0$ and $A^{-1} \geq 0$ then, $\rho(L^{-1}R) < 1$.

If we define $A = L - R$ we know from step 1 that under condition (c) $L^{-1} \geq 0$ and $R \geq 0$. To demonstrate that $A^{-1} \geq 0$ we can use the same arguments as in step 1, because A has a similar structure as L (i.e. $a_{ii} > 0$ and $a_{ij} \leq 0$, $i \neq j$). So, according to theorem 3.13 of Varga $\rho(L^{-1}R) < 1$, so the solution is stable.

So, step 1 and step 2 together outline the proof of statement (s).

Annex 15 Changes in the source code of TOXSWA 1.0 for comparison with the analytical solutions for the water layer and the sediment

```

**      SUMMARY OF THE CHANGES IN THE SOURCE CODE OF THE TOXSWA 1.0
**      PROGRAM, WHICH WERE MADE TO ENABLE THE COMPARISON OF TOXSWA 1.0
**      WITH ANALYTICAL SOLUTIONS FOR THE MASS CONSERVATION EQUATIONS OF
**      THE WATER AND SEDIMENT SUBSYSTEMS
**
**
**      subroutine wlsocem
**
**      SUBROUTINE
**      wlsocem - the elements of the matrix, which are used in the subroutines
**               wlsolb, wlsoub and wlsocr, are defined here
**
**
**      lddwddf      = 0.
**      &            ((2.*por(1)*kdfwbmh(1)*pez0hw)/delz(1))*
**      &            (1.-thetawl)*deltwl
**
**      rddwddf      = 0.
**      &            ((2.*por(1)*kdfwbmh(1)*pez0hw)/delz(1))*thetawl*
**      &            deltwl
**
**      rvwddf       = 0.
**      &            ((2.*por(1)*kdfwbmh(1)*pez0hw)/delz(1)) *
**      &            cowbjkis1(ixnotot) * deltwl
**
**
**
**      subroutine wbsoub
**
**      SUBROUTINE
**      wbsoub - the first row of the two tridiagonal matrices will be
**               composed here
**
**
**      if (qseifjph.ge.zero) then
**
**      bidd(kznotot) = biddadph + biddsdph +
**      &               (pemh(1)*pormh(1)*kdfwbmh(1)*(1.-thetawb)
**      &               *deltwb)
**      &               / (0.5*(delz(1)**2.)*pe(1)) +
**      &               biddca + biddtf
**      blbd(kznotot) = blbdad - blbd sdf
**
**      brdd(kznotot) = -brddaaph - brddsdph
**      &               -(pemh(1)*pormh(1)*kdfwbmh(1)*thetawb*deltwb)/
**      &               (0.5*(delz(1)**2.)*pe(1))
**      &               + brddca - brddtf
**      brbd(kznotot) = -brbdad + brbd sdf
**
**      brv(kznotot) = cowlj(ixnotot) * (
**      &               (leplot*qseifjph*thetawb
**      &               *deltwb)/
**      &               (delz(1)*pe(1))) +
**      &               (pemh(1)*pormh(1)*kdfwbmh(1)*
**      &               thetawb*deltwb) / (0.5*(delz(1)**2.)*pe(1))) +
**      &               cowljpl(ixnotot) * (
**      &               (leplot*qseifjph*
**      &               (1.-thetawb)*
**      &               deltwb) / (delz(1)*pe(1)))

```

```

*      &      (pemh(1)*pormh(1)*
*      &      kdfwbmh(1)*(1.-thetawb)*deltwb) / (0.5*
*      &      (delz(1)**2.)*pe(1))
*      else

**
**
**
      subroutine wimb

** SUBROUTINE
** wimb - the mass balance is checked for the water layer
**

C--- Calculation of incoming substance from sediment and of
C substance penetrating in sediment
C (this is total P*Jwb*delx at time j and location i)
C This takes only place in ditch, not in front- and endbuffer

      totrsinwb = 0.0
      totrsoutwb = 0.0
      do 10 ixnotot = nxnofb+1, nxnofb+nxnodit
        if (qseifj.ge.zero) then
          rsinwb = leplot*qseifj*cowlj(ixnotot)*delx(ixnotot)
*      &      - por(1)*kdfwbmh(1)*delx(ixnotot)*pez0hw*
*      &      ((cowbjkisl(ixnotot)-cowlj(ixnotot))
*      &      /(0.5*delz(1)))
        else
C      qseifj is negative
          rsinwb = leplot*qseifj*cowbjkisl(ixnotot)*delx(ixnotot)
*      &      - por(1)*kdfwbmh(1)*delx(ixnotot)*pez0hw*
*      &      ((cowbjkisl(ixnotot)-cowlj(ixnotot))
*      &      /(0.5*delz(1)))
        end if
        totrsinwb = totrsinwb + amin1(0.0, rsinwb)
        totrsoutwb = totrsoutwb + amax1(0.0, rsinwb)
      10 continue

**
**
**
      subroutine wimbnode

** SUBROUTINE
** wimbnode - the mass balance is made for a selected section of
** the water layer
**

C--- Calculation of incoming substance from sediment and of substance
C penetrating in sediment in selected section
C (this is P*Jwb*delx at time j and location ixnotot)

      if (ixnotot.ge.nxnofb+1 .and. ixnotot.le.nxnofb+nxnodit)
        &      then
          if (qseifj.ge.zero) then
            rsinwb = leplot*qseifj*cowlj(ixnotot)*
            &      delx(ixnotot)
*      &      -por(1)*kdfwbmh(1)*delx(ixnotot)*pez0hw*
*      &      ((cowbjkisl(ixnotot)-cowlj(ixnotot))
*      &      /(0.5*delz(1)))
          else
C      qseifj is negative
            rsinwb = leplot*qseifj*cowbjkisl(ixnotot)*
            &      delx(ixnotot)

```

```

*      &          -por(1)*kdfwbmh(1)*delz(ixnotot)*pez0hw*
*      &          ((cowbjkial(ixnotot)-cowlj(ixnotot))
*      &          /(0.5*delz(1)))
      end if
      rsoutwbn = amax1(0.0, rsinwbn)
      rsinwbn = amin1(0.0, rsinwbn)
    else
      rsoutwbn = zero
      rsinwbn = zero
    end if

**
**
**
      subroutine wbnb

** SUBROUTINE
** wbnb - the mass balance is checked for the water bottom
**

C---- Calculation of incoming substance from water layer and of substance
C lost to water layer
C (this is P*Jlb at time j and k=1/2)

      if (qseifj.ge.zero) then
        rsinwl = leplot*qseifj*cowlj(ixnotot)
*      &          - pemh(1)*pormh(1)*
*      &          kdfwbmh(1)*((cowbj(1)-cowlj(ixnotot))/(.5*delz(1)))
      else
C          qseifj is negative
        rsinwl = leplot*qseifj*cowbj(1)
*      &          - pemh(1)*pormh(1)*
*      &          kdfwbmh(1)*((cowbj(1)-cowlj(ixnotot))/(.5*delz(1)))
      end if
      rsoutwl = amin1(0.0, rsinwl)
      rsinwl = amax1(0.0, rsinwl)

C---- Calculation of substance lost to water layer
C (this is P*Jlb at time j and k=1/2)

      if (qseifj.ge.zero) then
        rsoutwl = leplot*qseifj*cowlj(ixnotot)
*      &          - pemh(1)*pormh(1)*
*      &          kdfwbmh(1)*((cowbj(1)-cowlj(ixnotot))/(.5*delz(1)))
      else
C          qseifj is negative
        rsoutwl = leplot*qseifj*cowbj(1)
*      &          - pemh(1)*pormh(1)*
*      &          kdfwbmh(1)*((cowbj(1)-cowlj(ixnotot))/(.5*delz(1)))
      end if
      rsoutwl = amin1(0.0, rsoutwl)

```


Annex 17 Changes in the input files (compared to those of the example simulation for chlorpyrifos) for calculating TOXSWA output for comparison with the analytical solution for the water layer

```

*   Filename :   WLNJ.INP
*   Content  :   Input data for TOXSWA concerning numerical solution and
*               time and length of ditch considered
*
-----
*
*   Section 3:   Space parameters
*
-----
*   total length of ditch considered
xdlit = 400.          ! unit: m          range: 10. .... 10,000.
*
-----
*
*   Section 4:   Number of space nodes
*
-----
*   number of nodes in ditch
nxndit = 58          ! unit: -          range: 10 .... 200
*
-----
*
*   Section 5:   Location of space nodes
*               (x coordinate starts in front buffer)
*
-----
*   x coordinate of nodes in ditch
xcddit = 4. 12. 20. 28. ! unit: m          range: 0.1 ..... 10,000.
          36. 44. 52. 60.
          68. 76. 84. 92.
          100. 106. 112. 118.
          124. 130. 136. 142.
          148. 154. 160. 166.
          172. 178. 184. 190.
          196.
          204. 212. 220. 228.
          236. 244. 252. 260.
          268. 276. 284. 292.
          300. 306. 312. 318.
          324. 330. 336. 342.
          348. 354. 360. 366.
          372. 378. 384. 390.
          396.
*
-----End of file-----
*
*   Filename :   WLPA.INP
*   Content  :   Input data for TOXSWA concerning geometry and other
*               characteristics of the ditch
*
-----
*
*   Section 1:   Geometry of ditch
*
-----
*   side slope, horizontal/vertical, s_1
sisl = 1.E-05        ! unit: -          range: 1.E-05 ..... 10.0
*
-----End of file-----
*
*   Filename :   WLST.INP
*   Content  :   Input data for TOXSWA concerning initial concentrations
*               in the water layer
*

```


Annex 18 Input files for the example simulation for chlorpyrifos

```
*   Filename :   OPOUT.INP
*   Content  :   Input data for TOXSWA concerning options for obtaining
*                model output
*
*-----
*   Section 1:   Output files selected
*-----
*   output file input.out (echo of all input) desired ? (0 = no, 1 = yes)
* op_input = 1          ! unit: -          range: 0 .... 1
*   output file icwlhy.out (initial calculations for water layer and
* hydrology and exposure concentrations) desired ? (0 = no, 1 = yes)
* op_icwlhy = 1        ! unit: -          range: 0 .... 1
*   output file icwb.out (initial calculations for sediment) desired ?
* (0 = no, 1 = yes)
* op_icwb = 1          ! unit: -          range: 0 .... 1
*   output file wmb.out ( mass balance for water layer) desired ?
* 0 = no, 1 = yes)
* op_wmb = 1           ! unit: -          range: 0 .... 1
*   output file wmbnoder.out ( mass balance for a segment of water
* layer desired ? (0 = no, 1 = yes)
* op_wmbnoder = 1     ! unit: -          range: 0 .... 1
*   output file wbsconoder.out (concentrations in a sediment subsystem
* desired ? (0 = no, 1 = yes)
* op_wbsconoder = 1   ! unit: -          range: 0 .... 1
*   output file wmbnoder.out (mass balance for a sediment subsystem)
* desired ? (0 = no, 1 = yes)
* op_wmbnoder = 1     ! unit: -          range: 0 .... 1
*   output file wmball.out (mass balance of all sediment subsystems)
* desired ? (0 = no, 1 = yes)
* op_wmball = 1       ! unit: -          range: 0 .... 1
*   output file dbnoder.out (distribution of substance in wl+wb at
* noder wl) desired ? (0 = no, 1 = yes)
* op_dbnoder = 1      ! unit: -          range: 0 .... 1
*   output file dbdit.out (distribution of substance in entire ditch)
* desired ? (0 = no, 1 = yes)
* op_dbdit = 1        ! unit: -          range: 0 .... 1
*
*-----
*   Section 2:   Points in time for output
*-----
*   number of points in time at which output is desired (max. of 9)
* nditout = 5          ! unit: -          range: 1 .... 9
*   points in time at which output is desired
* ptditout = 0. 0.5  1.          ! unit: d          range: 0. .... 100.
*                2.  4.          ! (only ptditout(1) may equal 0.)
*
*-----
*   Section 3:   Segments of wl+sediment subsystems selected for output
*-----
*   number of segments wl, coupled to sediment subsystems for which
* output is desired (max. of 9)
* nwbsy = 3            ! unit: -          range: 1 .... 9
*   node number in water layer at/or under which output is desired
* iwbsy = 1 13 29      ! unit: -          range: 1 .... 200
*   number of upper segments forming the top layer for which the
* accumulated pesticide mass will be calculated
* ktop = 11           ! unit: -          range: 1 .... 30
*-----
*   End of file-----
```

```

*   Filename :   WLNU.INP
*   Content  :   Input data for TOXSWA concerning numerical solution and
*               time and length of ditch considered
*
*-----
*   Section 1: Numerical weight factors
*-----
*   numerical weight factor for space
betawl = 0.5          ! unit: -          range: 0.0 ..... 1.0
*   numerical weight factor for time
thetawl = 1.0        ! unit: -          range: 0.0 ..... 1.0
*
*-----
*   Section 2: Time parameters
*-----
*   selected time step for water layer
deltwl = 100.        ! unit: s          range: 1. .... 86,400.
*   total time considered (holds also for sediment)
ttot = 4.0           ! unit: d          range: 0.1 ..... 100.
*
*-----
*   Section 3: Space parameters
*-----
*   total length of ditch considered
xdit = 200.          ! unit: m          range: 10. .... 10,000.
*   length of front buffer (0. if none)
xfb = 0.             ! unit: m          range: 0. .... 1000.
*   length of end buffer (0. if none)
xeb = 0.             ! unit: m          range: 0. .... 1000.
*
*-----
*   Section 4: Number of space nodes
*-----
*   number of nodes in ditch
nxnodit = 29         ! unit: -          range: 10 ..... 200
*   number of nodes in front buffer (0 if none)
nxnofb = 0           ! unit: -          range: 0 ..... 25
*   number of nodes in end buffer (0 if none)
nxnoeb = 0          ! unit: -          range: 0 ..... 25
*
*-----
*   Section 5: Location of space nodes
*               (x coordinate starts in front buffer)
*-----
*   x coordinate of nodes in front buffer (0. if none)
xcdfb = 0.           ! unit: m          range: 0. .... 1000.
*   x coordinate of nodes in ditch
xcdit = 4. 12. 20. 28. ! unit: m          range: 0.1 ..... 10,000.
          36. 44. 52. 60.
          68. 76. 84. 92.
          100. 106. 112. 118.
          124. 130. 136. 142.
          148. 154. 160. 166.
          172. 178. 184. 190.
          196.
*   x coordinate of nodes in end buffer (0. if none)
xcdeb = 0.           ! unit: m          range: 0. .... 1000.
*-----
*-----End of file-----

```



```

* Filename : WLPA.INP
* Content  : Input data for TOXSWA concerning geometry and other
*           characteristics of the ditch
*-----
* Section 1: Geometry of ditch
*-----
* bottom width of ditch, b
wibot = 1.65          ! unit: m          range: 0.1 ..... 10.0
* side slope, horizontal/vertical, s_1
sisl = 2.0           ! unit: -          range: 1.E-05 ..... 10.0
* water depth defining perimeter for exchange water layer - sediment,
* h_w
wdhfl = 0.10         ! unit: m          range: 0.01 ..... 2.0
*-----
* Section 2: Geometry of the neighbouring plot of land
*-----
* length of draining plot, perpendicular to ditch and located at one
* or both sides
leplot = 100.        ! unit: m          range: 1. .... 1000.
*-----
* Section 3: Concentration and organic matter in suspended solids
*-----
* concentration of suspended solids, ss
coss = 50.           ! unit: g/m3       range: 1. .... 100,000.
* mass ratio of organic matter, m_om,ss
raomss = 0.10       ! unit: -          range: 0. .... 1.
*-----
* Section 4: Amount of macrophytes
*-----
* dry weight of macrophyte biomass per m^2 bottom, DW
dwmp = 250.         ! unit: g/m^2      range: 0. .... 1000.
*-----
*-----End of file-----

```



```

* Filename : WBNJ.INP
* Content  : Input data for TOXSWA concerning numerical solution and
*           thickness of the sediment considered
*
*-----
* Section 1: Numerical weight factors
*-----
* numerical weight factor for space
betawb = 0.5 ! unit: - range: 0.0 .... 1.0
* numerical weight factor for time
thetawb = 1.0 ! unit: - range: 0.0 .... 1.0
*
*-----
* Section 2: Time parameters
*-----
* selected time step for sediment
deltwb = 100. ! unit: s range: 1. .... 86,400.
*
*-----
* Section 3: Space parameters
*-----
* length of sediment considered (end buffer excluded)
zwb = 0.10 ! unit: m range: 0.01 .... 0.5
* length of end buffer sediment (0. if none)
zebb = 0. ! unit: m range: 0.0 .... 0.1
*
*-----
* Section 4: Number of space nodes
*-----
* number of nodes in sediment (end buffer excluded)
nznwb = 23 ! unit: - range: 10 .... 200
* number of nodes in end buffer (0 if none)
nznebb = 0 ! unit: - range: 0 .... 25
*
*-----
* Section 5: Location of space nodes
*-----
* z coordinate of nodes in sediment
zcdwb = 0.0005 ! unit: m range: 0.0001 .... 0.6
0.0015
0.0025
0.0035
0.0045
0.0060
0.0080
0.0100
0.0120
0.0145
0.0175
0.0205
0.0240
0.0280
0.0325
0.0375
0.0425
0.0475
0.0550
0.0650
0.0750
0.0850
0.0950
* z coordinate of nodes in end buffer (0. if none)
zcdobb = 0.0 ! unit: m range: 0.0 .... 0.6
*-----
* End of file
*-----

```



```
*-----*
* Section 3: Dispersion in sediment
*-----*
* dispersion length
* ldis = 0.015          ! unit: m          range: 0. .... 1.
*-----*
* End of file
*-----*
```



```

*   Filename :   SU.INP
*   Content  :   input data for TOXSWA concerning the substance
*
*-----
*   Section 1:   Transformation in water layer
*-----
*   rate coefficient for transformation in water layer, k (= ln2/DT_50)
*   krtfwl = 0.0092      ! unit: 1/d      range: 0. .... 50.
*
*-----
*   Section 2:   Sorption to suspended solids
*-----
*   sorption coefficient based at organic matter content, K_om,ss,
*   (distribution coefficient)
*   kdomssdit = 16.4      ! unit: m^3/kg      range: 0. .... 100.
*   concentration of pesticide at which the K_om of the suspended solids
*   has been observed, c_e,ss
*   coobkomss = 1.0E-06    ! unit: kg/m^3      range: 1.0E-08 ..... 0.1
*   Freundlich exponent for sorption to suspended solids, n_ss
*   exfrss = 0.984        ! unit: -          range: 0.1 .... 2.0
*
*-----
*   Section 3:   Sorption to macrophytes
*-----
*   slope of sorption isotherm based at dry weight macrophytes, K_mp,
*   (distribution coefficient)
*   kdmpdit = 2.0         ! unit: m^3/kg      range: 0. .... 100.
*
*-----
*   Section 4:   Volatilisation
*-----
*   transport coefficient of pesticide in liquid phase, k_l
*   klg = 1.70            ! unit: m/d          range: 0.1 .... 10,000.
*   transport coefficient of pesticide in gas phase, k_g
*   kga = 163.1          ! unit: m/d          range: 0.1 .... 10,000.
*   saturated gas pressure of pesticide, P
*   psat = 0.0025        ! unit: Pa           range: 0.00001 ..... 0.1
*   molecular mass of pesticide, M
*   mamol = 350.6        ! unit: g/mol        range: 10. .... 10,000.
*   temperature during observation of saturated vapour pressure and
*   solubility with which Henry coefficient is calculated, T
*   tekhe = 298          ! unit: K            range: 200. .... 350.
*   solubility of pesticide in water, c_sol
*   cosol = 2.0          ! unit: g/m^3        range: 0.001 ..... 100.
*
*-----
*   Section 5:   Exchange water layer - sediment
*-----
*   diffusion coefficient of pesticide in water, D_w
*   kdfw = 40.           ! unit: mm^2/d       range: 1. .... 200.
*
*-----
*   Section 6:   Transformation in sediment
*-----
*   rate coefficient for transformation in sediment, k_b (= ln2/DT_50)
*   krtfwb = 0.0040      ! unit: 1/d          range: 0. .... 50.

```



```

*-----*
*
* Section 7: Sorption to sediment
*-----*
*
* sorption coefficient based at organic matter content of sediment
* material K_om,wb, (distribution coefficient)
* kdomwb1 = 16.4      ! unit: m^3/kg      range: 0. .... 100.
* concentration of pesticide at which the K_om of the sediment
* material has been observed, c_e,wb
* coobkomwb = 1.0E-06 ! unit: kg/m^3      range: 1.0E-08 ..... 0.1
* Freundlich exponent for sorption to sediment material, n_wb
* exfrwb = 0.984     ! unit: -          range: 0.1 .... 2.0
*
*-----End of file-----*

```

Annex 19 Source code of the TOXSWA program, version 1.0, including a guide to the vocabulary used

List of letters and combinations of letters which are used more than once in the source code of TOXSWA 1.0 to form the names of the variables and parameters

a	all
ad	advection
ai(r)	air
all	all sediment subsystems
av	average
b	sediment [water <u>bottom</u>]
bd	bulk density
bot	bottom
c	concentration (in water or liquid phase)
ca	total concentration, c^* or c_p^*
cd	coordinate
cl	calculation
co	concentration
cu	accumulated (over time)
db	distribution
dd	distribution in ditch
del(t)	delta Δ , difference
df	diffusion
dh	depth
dit	ditch
ds	dispersion
dw	dry weight
eb	end buffer (of water layer)
ebb	end buffer of water bottom
en	end-side node
er	error
ex	exponent
f	front
fb	front buffer (of water layer)
fd	flux dispersion exceeds advection (in sediment)
fl	flux
fn	front-side node
fr	Freundlich
ga	gass
he	Henry
hw	water level defining the exchanging perimeter $P_{z=0}$
hy	hydrology
i	at node i
ic	input concentration
imh	at node $i-\frac{1}{2}$
in	input or incoming
iph	at node $i+\frac{1}{2}$
iml	at node $i-1$
ip1	at node $i+1$
j	at time j
jmh	at time $j-\frac{1}{2}$
jph	at time $j+\frac{1}{2}$
jp1	at time $j+1$
k	coefficient
kd	distribution coefficient
kis1	$k = 1$
kr	rate coefficient
l	water layer
lbd	band above the diagonal of the left-hand tridiagonal matrix
ldd	diagonal of the left-hand tridiagonal matrix

le	length
lin	lineic
lm	lineic mass
lo	loop
lod	band under the diagonal of the left-hand tridiagonal matrix
lot	neighbouring field lot
lq	liquid
m	mass
mb	mass balance
mn	minimum
mp	macrophytes
mx	maximum
n	number or numerical
no	node(s)
node	node(s)
nu	numerical
nw	numerical weight factor water layer
nwb	numerical weight factor sediment
nx	number in x direction
nz	number in z direction
ob	observed
old	of past time step
om	organic matter
op	option(s)
out	output
p	parameter
pe	perimeter
per	percentage or percolated
por	porosity
prc	percentage
pt	point of time
qseif	seepage/infiltration flux
qu	quantity
qvo	discharge volume
ra	ratio
rbd	band above the diagonal of the right-hand tridiagonal matrix
rdd	diagonal of the right-hand tridiagonal matrix
rh	right hand
rod	band under the diagonal of the right-hand tridiagonal matrix
rs	rate of substance
rv	right-hand vector
s	start
sisl	side slope
so	sorbed or solution
ss	suspended solids
st	start, beginning
su	substance
sum	sum of
sy	systems
t	time
tf	transformation
top	top layer in sediment or top of water level in ditch
tor	tortuosity
tot	total
ts	time steps
tt	iteration
u	flow velocity in ditch
ui	unit number for input file
uo	unit number for output file
us	upward seepage
var	variable
vol	volatilisation
w	water, wetted
war	wetted area
wb	sediment [water <u>b</u> ottom]

wdh	water depth
wi	width
wl	water layer
x	x direction
xp	exposure
z	z direction

```

** PROGRAM TOXSWA
**
** PROGRAM:
** toxswa.for - fate of pesticides in small surface waters
** author - P.I. Adriaanse
**
** SYNOPSIS:
** run TOXSWA
**
** DESCRIPTION:
** Main program of the TOXSWA simulation model; it calculates
** (i) pesticide concentration profiles in water layer (water phase,
** sorbed to suspended solids and sorbed to macrophytes) and in
** sediment (pore water and sorbed to solid bottom material),
** (ii) mass balances for the water and sediment layer and
** (iii) distribution of pesticide mass between the different
** compartments by calling different subroutines.
**
** References:
** Adriaanse, P.I., 1996. Fate of pesticides in field ditches: the
** TOXSWA simulation model. Report 90, DLO Winand Staring Centre,
** Wageningen.
** Belman, W.H.J., P.I. Adriaanse and M. van Elswijk, 1996. User's
** manual TOXSWA 1.0. Technical Document 33. DLO Winand Staring
** Centre, Wageningen.
**
** Restrictions for use:
** Use is restricted according to the 'General Terms and Conditions
** as to the Making Available of Computer Software', dated 1 November
** 1990 of the DLO Winand Staring Centre, Wageningen.
**
** HISTORY:
** 10 April 1996 P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
**
** C include 'common.for' common variables
** C real cowbjpl_2d(mnxxnnotot, mnznnotot) local variables
** C logical allconczero
**
** C call opout options for model output
** C call wlin general and initial input water layer
** C call wbin general and initial input sediment
** C call whin hydrological input data water layer and
** C call hwin sediment
** C call hwin characteristics substance
** C call suin initial calculations water layer
** C call wlit initial calculations sediment
** C call wbit
**
** C do 90 its = 1, ntswl timestep loop
** C call wlso solution water layer
** C call wlmb mass balance water layer
** C call wlmbnode output mass balance selected node w1
** C call wmbnode sediment subsystems loop (under each water
** C layer segment ixnnotot there is a sediment
** C subsystem)
** C do 50 ixnnotot = nxnofb+1, nxnofb+nxnodit
** C fill array cowbj(1-nznnotot) with right
** C values under this node of water layer
** C do 30 kznotot = 1, nznnotot

```

```

if (its.eq.1) then
  cowbj(kznnotot) = costwb(kznnotot)
else
  cowbjpl_2d(ixnnotot, kznotot)
end if
continue
30
C call wbsol solution sediment
C call wbsob mass balance sediment
C call wmbout output partition mass w1+wb
C call wmbout sediment subsystems loop
C do 40 kznotot = 1, nznnotot
  cowbjpl_2d(ixnnotot, kznotot) = cowbjpl(kznnotot)
  continue
40
C call wbsob
C call wmbout
C do 40 kznotot = 1, nznnotot
  cowbjpl_2d(ixnnotot, kznotot) = cowbjpl(kznnotot)
  continue
40
C call wbsob
C call wmbout
C do 40 kznotot = 1, nznnotot
  cowbjpl_2d(ixnnotot, kznotot) = cowbjpl(kznnotot)
  continue
40
C --- Program stops when all concentrations are zero (below 1E-20)
  allconczero = .true.
  do 60 ixnnotot = 1, nxnnotot
    if (cowbjpl(ixnnotot).gt.1.0E-20) then
      allconczero = .false.
    end if
  end if
  continue
60
C do 80 ixnnotot = nxnofb+1, nxnofb+nxnodit
  do 70 kznotot = 1, nznnotot
    if (cowbjpl_2d(ixnnotot, kznotot).gt.1.0E-20) then
      allconczero = .false.
    end if
  end if
  continue
70
80
  if (allconczero) then
    print*, 'All concentrations in water layer and in'
    print*, 'sediment are zero (less than 1E-20).'
    print*, 'Programs stops.'
    stop
  end if
  continue
  time steps loop
C 90 continue
  print*, 'end of simulation'
  stop
  end
**
** COMMON BLOCK:
** common.for - contains declarations of all common variables for
** the TOXSWA program
** author - P.I. Adriaanse
**
** DESCRIPTION:
** This common block contains all the common variables of TOXSWA; it
** is included in all the subroutines of the TOXSWA program.
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
** implicit none

```



```

C      tiny will be used to prevent dividing by
C      zero or raising a (tiny) negative number
C      to a power
C      parameter (tiny = 1.0E-25)

C      block data

** DATA BLOCK:
** block for - initialises identifiers of the TOXSWA program
** author - P.I. Adriaanse

** DESCRIPTION:
** This data block initialises identifiers of TOXSWA

** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

C      include 'common.for'      common variables
C      include 'local.for'      local variables

C      data u100, u10m, u10p, u10s, u10n, u10p, u10s, u10y, u10z
C      & / 81, 83, 85, 87, 89, 91, 93, 95, 97/
C      data u0er, u0in, u0w1, u0w2, u0c1, u0m1, u0m2, u0d1
C      & / 21, 22, 23, 24, 25, 26, 27, 28/
C      data zero / 0./
C      end

C      subroutine opout

** SUBROUTINE
** opout.for - input data for the TOXSWA program about options for
**              obtaining model output are read
** author      - P.I. Adriaanse

** DESCRIPTION:
** In this subroutine it is read which output the user wants of the
** TOXSWA model.

** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

C      include 'common.for'      common variables
C      integer i,k, helpnditout, helpnwsy
C      statement for Vax Fortran 77 compiler for date and time
C      call date (dmy)
C      call time (hms)

C      open (unit = u0er, file = 'message.out', status = 'unknown')
C      write (u0er, 10) dmy, hms
C      10 format (' TOXSWA simulation:', 2A10, /,

```

```

&      tcoxp, sumct, coxp
&      common/realwlb2/ qvo1jmh, qvo1jph, qvo1jmh, qvo1jplh,
&      wdhlj, wdhlmh, wdhljpl, wdhljpl,
&      wdhljplm, wdhljplm,
&      wazj, wazjmi, wazjpl,
&      wazjpl, wazjplm, wazjplh, wazjph,
&      wazjphmh, wazjphmh, wazjphmh, wazjplmh,
&      qvo1jmh, qvo1jph, witopj, witopjph,
&      zero, delz,
&      wdhitin, wdhitout, wdhitin, wdhitout,
&      qvoditn, qvoditout

&      common/realwb/ betawb, thetabw,
&      deltawb,
&      zwb, zebb,
&      zcwb, zcdebb, zed,
&      kcfwb, bdehwh, bdehwh, por, pormh, porph, tor,
&      kdombw, kcfwb, kcfwbmh, kcfwbmh,
&      coobkombw, kcfwb, raombw,
&      idis,
&      kdwbjkmh, kdwbjkmh, kdwbjkmh, kdwbjkmh,
&      kdwbjkmh, kdwbjkmh, kdwbjkmh

&      common/realwbi/ pe, penn, pepb,
&      castwb, costwb, sowbat,
&      nwmw, nwopb,
&      cowbj, cowbjpl, oldcowbjpl, cowbjkmh, cowbjkph,
&      sofdfactoram, sofdfactorph,
&      sowbjpl, sawbjpl

&      common/realwb2/ b10d, b10d, brod, brdd, brbd, brv, brh,
&      b10dad, b10dad, b10dad, b10dad, b10dad, b10dad, b10dad, b10dad,
&      b10dad, b10dad, b10dad, b10dad, b10dad, b10dad, b10dad, b10dad,
&      brodad, brodad, brodad, brodad, brodad, brodad, brodad, brodad,
&      brddsdadm, brddca, brddca, brddca, brddca, brddca, brddca, brddca,
&      brddsdadm, brddca, brddca, brddca, brddca, brddca, brddca, brddca

```

```

** FILE:
** param.for - include-file, it defines parameters that are used for
**              the declaration of the dimensions of certain arrays
**              for the TOXSWA program
** author      - P.I. Adriaanse

** DESCRIPTION:
** This include-file is included in the common block of the TOXSWA
** program
** version 1.0

** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

integer mxnknofb, mxnknodit, mxnknob, mxnknob, mxnknob,
& mxznznobw, mxznznobw, mxznznobw
real
tiny

parameter (mxnknofb = 20)      maximum number nodes in front buffer w1
parameter (mxnknodit = 100)  maximum number nodes in ditch
parameter (mxnknob = 20)     maximum number nodes in end buffer w1
parameter (mxnknob = 140)    maximum number nodes in ditch and 2 buffers

parameter (mxznznobw = 100)  maximum number nodes in sediment
                              (end buffer excluded)
parameter (mxznznobw = 20)   maximum number nodes in end buffer wb
parameter (mxznznobw = 20)   maximum number nodes in sediment
                              (end buffer included)
parameter (mxznznobw = 120)  maximum number nodes in sediment
                              (end buffer included)

```

```

&
' Messages (error/warning/ttutil):')
C---- input data from file opout.inp
call rdinit (uico, uoer, 'opout.inp')

C
call rdslnr ('op_inpu', 0, 1, op_inpu)
call rdslnr ('op_icwlhy', 0, 1, op_icwlhy)
call rdslnr ('op_icwb', 0, 1, op_icwb)
call rdslnr ('op_wmb', 0, 1, op_wmb)
call rdslnr ('op_wmbnodenr', 0, 1, op_wmbnodenr)
call rdslnr ('op_wbsconodenr', 0, 1, op_wbsconodenr)
call rdslnr ('op_wmbnodenr', 0, 1, op_wmbnodenr)
call rdslnr ('op_wmbball', 0, 1, op_wmbball)
call rdslnr ('op_chnodenr', 0, 1, op_chnodenr)
call rdslnr ('op_dbdnt', 0, 1, op_dbdnt)

if (op_inpu.eq.1) then
open (unit = uoin, file = 'input.out', status = 'unknown')
write (uoin, 20) dmy, hms
20 format ('TOXSWA simulation: ', 2A10, //, ' Output of input: ', //,
& ' All input from subroutine opout:')
end if

if (op_inpu.eq.1) then
write (uoin, 30)
30 format (' Section 1: Selected output files: (no dimensions)')
write (uoin, *) op_inpu
write (uoin, *) op_icwlhy
write (uoin, *) op_icwb
write (uoin, *) op_wmb
write (uoin, *) op_wmbnodenr
write (uoin, *) op_wbsconodenr
write (uoin, *) op_wmbball
write (uoin, *) op_chnodenr
write (uoin, *) op_dbdnt
end if

C
call rdslnr ('nditout', 1, 9, nditout)
call rdslnr ('pdtitout', 0, 100, pdtitout, 9, helpnditout)
if (helpnditout.ne.nditout) then
write (uoer, 40)
40 format (//, ' Error in input opout.inp, //,
& ' actual length array pdtitout is not equal, //,
& ' to declared length, please correct data')
end if

if (op_inpu.eq.1) then
write (uoin, 41)
41 format (' Section 2: Points of time for output: (d)')
write (uoin, *) nditout
write (uoin, *) (pdtitout(i), i = 1, nditout)
end if

C
call rdslnr ('nwbsy', 1, 9, nwbsy)
call rdslnr ('iwsy', 1, 200, iwsy, 9, helpnwbsy)
if (helpnwbsy.ne.nwbsy) then
write (uoer, 50)
50 format (//, ' Error in input opout.inp, //,
& ' actual length array iwsy is not equal,
& ' to declared length, please correct data')
end if

call rdslnr ('ktop', 1, 30, ktop)

if (op_inpu.eq.1) then
write (uoin, 51)
51 format (' Section 3: Selected sediment subsystems for output:')
&
write (uoin, *) nwbsy
write (uoin, *) (iwsy(k), k = 1, nwbsy)
write (uoin, *) ktop
end if

close (uico)
if (op_inpu.eq.1) then
close (unit = uoin)
end if

return
end

subroutine wlin
** SUBROUTINE:
** wlin.for - all input data of the water layer for the TOXSWA
** program are read
** author - P.I. Adriaanse
**
** DESCRIPTION:
** Input data for water layer concerning numerical aspects, the ditch
** geometry and other ditch characteristics and the initial
** concentrations are read; the data are converted to SI-units
** (m, s, g, Pa, K, J, mol).
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
** C include 'common.for' common variables
** C integer i, helpknofb, helpknodit, helpknoeb,
& helpknotot local variables

if (op_inpu.eq.1) then
open (unit = uoin, file = 'input.out', status = 'unknown')
write (uoin, 11) dmy, hms
11 format (' TOXSWA simulation: ', 2A10, // Output of input: //)
write (uoin, 12)
12 format (//, ' All input data from wlin:')
end if

if (op_icwlhy.eq.1) then
open (unit = uowl, file = 'icwlhy.out', status = 'unknown')
write (uowl, 13) dmy, hms
13 format (' TOXSWA simulation: ', 2A10, //
& ' Output of initial calculations water layer and hydrology: ',
& //)
end if

C---- input data (numerical aspects) from file wlinu.inp
call rdinit (uiwn, uoer, 'wlinu.inp')

C
call rdslnr ('betawl', 0, 0, 1, 0, betawl)
call rdslnr ('thetawl', 0, 0, 1, 0, thetawl)
if (op_inpu.eq.1) then
write (uoin, 20)
20 format (//, ' Input from wlinu.inp: //,
& ' Section 1: Numerical weight factors',
& ' (no dimensions)')
& write (uoin, *) betawl, thetawl
end if

C
call rdslnr ('deltwl', 1, 86400, deltwl)
time parameters
end if

```



```

call rdarer ('xtot', 0.1, 100., ttot)
if (op_input.eq.1) then
  write (uoin,30)
30 format (' Section 2: Time parameters (s, resp d)')
  write (uoin,*) deltwl, ttot
end if
ttot = ttot*86400.
ncswl = int (ttot/deltwl) + 1
C
C calculation of timesteps, itsout(i) for
C output with the aid of points of time
C peditout(i) of subroutine epout
if (peditout(1) .le. zero) then
  else itsout(1) = 1
  if itsout(1) = nint (peditout(1)*86400./deltwl)
  end if
do 31 i = 2, nditout
  itsout(i) = nint (peditout(i)*86400./deltwl)
31 continue
if (op_iewlby.eq.1) then
  write (uowl,47)
32 format (' Results of initial calculations wlin:')
33 format (' There are ,16, , timesteps for output, these are: ')
  write (uowl,*) (itsout(i), i = 1, nditout)
end if
C
C space parameters
call rdarer ('xdlit', 10., 10000., xdlit)
call rdarer ('xib', 0., 1000., xib)
call rdarer ('xwbfl', 0., 1000., xwbfl)
if (op_input.eq.1) then
  write (uoin,40)
40 format (' Section 3: Space parameters (m)')
  write (uoin,*) xdlit, xib, xwbfl
end if
C
C number space nodes
call rdinar ('nxndit', 10, 100, nxndit)
call rdinar ('nxnofb', 0, 20, nxnofb)
call rdinar ('nxnoeb', 0, 20, nxnoeb)
if (op_input.eq.1) then
  write (uoin,50)
50 format (' Section 4: Number space nodes (no dimensions)')
  write (uoin,*) nxndit, nxnofb, nxnoeb
end if
C
C location space nodes
C dummy statement
if (nxnofb.eq.0) then
  nxnofb = 0
else
  call rdarer ('xcdfb', 0., 1000., xcdfb, mxnxfbf, helpnxfbf)
end if
C
if (helpnxfbf.ne.nxnofb) then
  write (uover,60)
60 format (' Error in input wlinu.inp, /, /,
& actual length array xcdfb is not equal, /, /,
& to declared length, please correct data', /)
end if
C
call rdarer ('xcedit', 0.1, 10000., xcedit, mxnxcndit, helpnxcndit)
if (helpnxcndit.ne.nxndit) then
  write (uover,61)
61 format (' Error in input wlinu.inp, /, /,
& actual length array xcedit is not equal, /, /,
& to declared length, please correct data', /)
end if
C
if (nxnoeb.eq.0) then
  dummy statement
  nxnoeb = 0
else
  call rdarer ('xdeb', 0., 1000., xdeb, mxnxnoeb, helpnxnoeb)
end if
if (helpnxnoeb.ne.nxnoeb) then
  write (uover,62)
62 format (' Error in input wlinu.inp, /, /,
& actual length array xdeb is not equal, /, /,
& to declared length, please correct data', /)
end if
if (op_input.eq.1) then
  write (uoin,63)
63 format (' Section 5: Location space nodes (m)')
  write (uoin,*) (xcdfb(i), i = 1, nxnxfbf)
  write (uoin,*) (xcedit(i), i = 1, nxcndit)
  write (uoin,*) (xdeb(i), i = 1, nxnoeb)
end if
close (uowm)
C----- input data (geometry and characteristics) from file wipa.inp
call rdinit (uiwp, uover, 'wipa.inp')
C
C geometry ditch
call rdarer ('wibot', 0.1, 10., wibot)
call rdarer ('sisl', 1.5, 0.5, 10., sisl)
call rdarer ('wdhfl', 0.01, 2.0, wdhfl)
if (op_input.eq.1) then
  write (uoin,80)
80 format (' Input from wipa.inp, /, /,
& Section 1: geometry ditch, /, /,
& (m, dimensionless, resp m)')
  write (uoin,*) wibot, sisl, wdhfl
end if
pez0hw = wibot + 2.*wdhfl*sqrt((sisl**2)-1.)
C
C geometry neighbouring plot of land
call rdarer ('leplot', 1., 1000., leplot)
if (op_input.eq.1) then
  write (uoin,81)
81 format (' Section 2: Geometry neighbouring plot (m)')
  write (uoin,*) leplot
end if
C
C concentration and org. matter susp. solids
call rdarer ('coss', 1., 10000., coss)
call rdarer ('raoms', 0., 1., raoms)
if (op_input.eq.1) then
  write (uoin,90)
90 format (' Section 3: Concentration and org. matter susp.solids',
& (g/m**3, dimensionless)')
  write (uoin,*) coss, raoms
end if
C
C amount macrophytes
call rdarer ('dwp', 0., 1000., dwp)
if (op_input.eq.1) then
  write (uoin,100)
100 format (' Section 4: Amount macrophytes (g/m**2)')
  write (uoin,*) dwp
end if
close (uiwp)
C----- input data (initial concentrations w1) from file w1st.inp
call rdinit (uiws, uover, 'w1st.inp')
C
C initial concentrations
call rdarer ('castw1', 0., 100., castw1, mxnxnotot, helpnxnotot)

```

```

nxsnotot = nxsnotot+nxnfb+nxnceb
if (helpnxsnotot.ne.nxsnotot) then
  write (uwer, 110)
110 format (//, ' Error in input wst.inp', //,
  & , actual length array castwl is not equal', //,
  & , to declared length, please correct data')
end if

if (op.inp.eq.1) then
  write (uoin, 111)
111 format (//, ' Input from wst.inp', //,
  & , Section 1: Initial concentrations',
  & , (g/m3'))
  write (uoin, *) (castwl(i), i = 1, nxsnotot)
end if

C
call rdsrer ('coair', 0., 0., coair)
  constant background concentration in air
if (op.inp.eq.1) then
  write (uoin, 120)
120 format (' section 2: Background concentration (g/m3)')
  write (uoin, *) coair
end if
close (uows)

return

end

subroutine wbin

** SUBROUTINE:
** wbin.for - all input data of the sediment layer for the TOXSWA
** program are read
** author - P.I. Adriaanse

** DESCRIPTION:
** Input data for the sediment layer concerning numerical aspects,
** thickness and other characteristics of the sediment layer and the
** initial concentrations are read; the input data are converted to
** SI-units (m, s, g, Pa, X, J, mol).

** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

C
include 'common.for'
common variables
integer k, help4znobw, help3znobw, help2znobw, help1znobw, help4znobw,
& help3znobw, help2znobw, help1znobw, help4znobw, help3znobw,
& help2znobw, help1znobw

C----- the following statements have already been executed in the options
C for output subroutine, opout.for
C----- statement for Vax Fortran 77 compiler for date and time
C
C call date (dmy)
C call time (hms)
C
C open (unit = uwer, file = 'message.out', status = 'unknown')
C write (uwer, 10) dmy, hms
C 10 format (' TOXSWA simulation:', 2A10, /' Error messages:./)

C----- input data (numerical aspects) from file wbinu.inp

```

```

call rdinit (uibn, uoer, 'wbinu.inp')

```

```

if (op.inp.eq.1) then
  write (uoin, 20)
20 format (//, ' All input data from wbin:')
end if

```

```

C
call rdsrer ('betawb', 0., 1., betawb)
call rdsrer ('thetawb', 0., 1., thetabw)

```

```

if (op.inp.eq.1) then
  write (uoin, 30)
30 format (//, ' Input from wbinu.inp', //,
  & , Section 1: Numerical weight factors',
  & , (no dimensions'))
  write (uoin, *) betawb, thetabw
end if

```

```

C
call rdsrer ('deltwb', 1., 86400., deltwb)
if (deltwb.ne.deltwl) then
  write (uwer, 40)
40 format (//, ' Error in wbinu.inp', //,
  & , deltwb should be equal to deltwl in TOXSWA 1.0 !', //,
  & , program stops')
end if

```

```

if (deltwb.ne.deltwl) then
  print*, ' Error in wbinu.inp.'
  print*, ' Deltwb should be equal to deltwl in TOXSWA 1.0.'
  print*, ' Programs stops.'
  stop
end if

```

```

if (op.inp.eq.1) then
  write (uoin, 41)
41 format (' Section 2: Time parameters (s)')
  write (uoin, *) deltwb
end if

```

```

C
call rdsrer ('zwb', 0.01, 0.5, zwb)
call rdsrer ('zebb', 0., 0.5, zebb)

```

```

if (op.inp.eq.1) then
  write (uoin, 50)
50 format (' Section 3: Space parameters (m)')
  write (uoin, *) zwb, zebb
end if

```

```

C
call rdsinr ('nznobw', 10, 100, nznobw)
call rdsinr ('nznobbb', 0, 20, nznobbb)

```

```

if (op.inp.eq.1) then
  write (uoin, 60)
60 format (' Section 4: Number space nodes (no dimensions)')
  write (uoin, *) nznobw, nznobbb
end if

```

```

C
call rdsrer ('zcdwb', 0.0001, 0.6, zcdwb, mxnznobw, helpnznobw)
if (helpnznobw.ne.nznobw) then
  write (uwer, 70)
70 format (//, ' Error in input wbinu.inp', //,
  & , actual length array zcdwb is not equal', //,
  & , to declared length, please correct data', //,
  & , )
end if

```

```

if (nznobbb.eq.0) then
  nznobbb = 0
else
  call rdsrer ('zcdzbb', 0., 0.6, zcdzbb, mxnznobbb, helpnznobbb)
end if

```

```

if (helpznnoebb.ne.nznnoebb) then
  write (uwer, 71)
71 format (//, 'Error in input wbuw.inp', /,
&
, actual length array zcdebb is not equal', /,
&
, to declared length, please correct data')
& end if

if (op_input.eq.1) then
  write (uoin, 72)
72 format (' Section 5: Location space nodes
write (uoin, *) (zcdwb(k), k = 1, nznnoebb)
write (uoin, *) (zcdwb(k), k = 1, nznnoebb)
end if

close (uibn)

C----- input data (sediment characteristics) from file wbp.a.inp
call rdinit (uibp, uwer, 'wbp.a.inp')

C
physical properties sediment
call rdarer ('bdwb', 10., 3000., bdwb, mxznzwb, help4znzwb)
call rdarer ('por', 0., 1., por, mxznzwb, help4znzwb)
call rdarer ('tor', 0., 1., tor, mxznzwb, help4znzwb)
&
if (help4znzwb.ne.nznzwb.or.help4znzwb.ne.nznzwb.or.
help4znzwb.ne.nznzwb) then
  write (uwer, 90)
90 format (//, 'Error in input wbp.a.inp', /,
&
, actual length array bdwb, por or tor is not equal',
&
, , to declared length, please correct data')
& end if

if (op_input.eq.1) then
  write (uoin, 91)
91 format (' Section 1: Physical properties sediment',
&
(kg/m3, (2x) dimensionless), /
write (uoin, *) (bdwb(k), k = 1, nznzwb)
write (uoin, *) (por(k), k = 1, nznzwb)
write (uoin, *) (tor(k), k = 1, nznzwb)
end if

C
compose vectors bdwb(k), por(k) and tor(k)
C
if (nznnoebb.eq.0) then
  'dummy' statement
  nznnotot = nznzwb
else
  do 92 k = 1, nznzwb+nznnoebb
    bdwb(k) = bdwb(nznzwb)
    por(k) = por(nznzwb)
    tor(k) = tor(nznzwb)
  92 continue
  do 93 k = 1, nznzwb+nznnoebb
    bdwb(k) = bdwb(k)*1000.
  93 continue

C
organic matter sediment
call rdarer ('raomwb', 0., 1., raomwb, mxznzwb, help4znzwb)
write (uwer, 100)
100 format (//, 'Error in input wbp.a.inp', /,
&
, actual length array raomwb is not equal', /,
&
, to declared length, please correct data')
& end if

if (op_input.eq.1) then
  write (uoin, 101)
101 format (' Section 2: Organic matter sediment',
&
(dimensionsless)', /
write (uoin, *) (raomwb(k), k = 1, nznzwb)
end if

compose vector raomwb(k) in end buffer

C
if (nznnoebb.eq.0) then
  'dummy' statement
  nznnotot = nznzwb
else
  do 102 k = nznzwb+1, nznzwb+nznnoebb
    raomwb(k) = raomwb(nznzwb)
  102 continue
end if

C
call rdarer ('ldis', 0., 1., ldis)
dispersion in sediment
if (op_input.eq.1) then
  write (uoin, 110)
110 format (' Section 3: Dispersion length
write (uoin, *) ldis
end if

close (uibp)

C----- input data (initial concentrations wb) from file wbst.inp
call rdinit (uibp, uwer, 'wbst.inp')

C
initial concentrations
nznnotot = nznzwb+nznnoebb
call rdarer ('castwb', 0., 1000., castwb, mxznzwb, help4znzwb)
write (uwer, 120)
120 format (//, 'Error in input wbst.inp', /,
&
, actual length array castwb is not equal', /,
&
, to declared length, please correct data')
& end if

if (op_input.eq.1) then
  write (uoin, 121)
121 format (//, 'Input from wbst.inp', /,
&
, Section 1: Initial concentrations',
&
(g/m3), /
write (uoin, *) (castwb(k), k = 1, nznnotot)
end if

close (uibp)
return

end

subroutine hyin
** SUBROUTINE:
** hyin.for - all hydrological input data for the TOMSWA program are
** read, for the water layer as well as for the sediment
** layer.
** author - P.I. Adriaanse
**
** DESCRIPTION:
** Input data for the water layer concern the constant water flow and
** water depth; to facilitate possible later linkage to a hydro-
** dynamic model these constant values are transcribed in variables;
** input data for the sediment layer concern upward or downward
** seepage from the neighbouring lot into the ditch or vice versa
** and the pesticide concentration in the water seeping upwardly into
** the ditch bottom; all input data are converted to SI-units
** (m, s, g, Pa, K, J, mol).
**
** HISTORY:
** 10 April 1986 - P.I. Adriaanse
** version 1.0
**

```



```

30 format (' section 2: Sorption susp. solids',
& write (uoin, *) kdmsedit, coobkoms, exfss)
end if (uoin, *) kdmsedit, coobkoms, exfss
kdmsedit = kdmsedit/1000.
coobkoms = coobkoms*1000.

C --- compose vector kdoms(i), distribution coefficient for sorption to
C suspended solids, for buffers (no sorption, kdoms=0) and in ditch
C in front buffer
if (nxnofb.eq.0) then
else
do 31 i = 1, nxnofb
kdoms(i) = zero
continue
31 end if

C in ditch
do 32 i = nxnofb+1, nxnofb+nxnodit
kdoms(i) = kdmsedit
32 continue

C in end buffer
if (nxnoeb.eq.0) then 'dummy' statement
else
nxnotot = nxnofb+nxnodit
do 33 i = nxnofb+nxnodit+1, nxnofb+nxnodit+nxnoeb
kdoms(i) = zero
continue
33 end if

do 34 i = 1, nxnofb+nxnodit+nxnoeb
kdfw(i) = raoms*kdoms(i)
34 continue

C sorption to macrophytes
call rdsrer ('kdmptit', 0., 100., kdmptit)
if (op_input.eq.1) then
write (uoin, 40)
40 format (' Section 3: Sorption to macrophytes',
& write (uoin, *) kdmptit
end if
kdmptit = kdmptit/1000.

C --- compose vector kdmpt(i), distribution coefficient for sorption to
C macrophytes, for buffers (no sorption, kdmpt=0) and in ditch
C in front buffer
if (nxnofb.eq.0) then
kdmpt(i) = 1.
else
do 41 i = 1, nxnofb
kdmpt(i) = zero
continue
41 end if

C in ditch
do 42 i = nxnofb+1, nxnofb+nxnodit
kdmpt(i) = kdmptit
42 continue

C in end buffer
if (nxnoeb.eq.0) then 'dummy' statement
else
nxnotot = nxnofb+nxnodit
do 43 i = nxnofb+nxnodit+1, nxnofb+nxnodit+nxnoeb
kdmpt(i) = zero
continue
43 end if

C volatilisation
call rdsrer ('klg', 0.1, 10000., klg)
call rdsrer ('kga', 0.1, 10000., kga)
call rdsrer ('psat', 0.0001, 0.1, psat)
call rdsrer ('mamol', 10., 10000., mamol)

call rdsrer ('tekhe', 200., 350., tekhe)
call rdsrer ('cosol', 0.001, 100., cosol)
if (op_input.eq.1) then
write (uoin, 50)
50 format (' Section 4: Volatilisation',
& write (uoin, *) psat, mamol, tekhe, cosol
end if
klg = klg/86400.
kga = kga/86400.
kpsa = 8.3144
khe = (psat*mamol)/(unga*tekhe*cosol)
helpvar = (1./klg)*(1./khe*kga)
kcl = 1./helpvar

C exchange water layer and sediment
call rdsrer ('kdfw', 1., 200., kdfw)
if (op_input.eq.1) then
write (uoin, 60)
60 format (' Section 5: Exchange water layer - sed.',
& write (uoin, *) kdfw
end if
kdfw = kdfw*(1.0E-06)/86400.

C transformation in sediment
call rdsrer ('krtfwb', 0., 50., krtfwb)
if (op_input.eq.1) then
write (uoin, 70)
70 format (' Section 6: Transformation in sediment ',
& write (uoin, *) krtfwb
end if
krtfwb = krtfwb/86400.

C dry sediment material
call rdsrer ('kdomwb1', 0., 100., kdomwb1)
call rdsrer ('coobkomwb', 1.0E-08, 0.1, coobkomwb)
call rdsrer ('exirwb', 0.1, 2.0, exirwb)
if (op_input.eq.1) then
write (uoin, 80)
80 format (' Section 7: Sorption sediment',
& write (uoin, *) kdomwb1, coobkomwb, exirwb
end if
kdomwb1 = kdomwb1*0.001
coobkomwb = coobkomwb*1000.

C compose vector kdomwb(k), distribution
C coefficient for sorption to sediment
C material, for end buffer (no sorption,
C kdomwb(k)=0) and in sediment itself
do 81 k = 1, nznobw
kdomwb(k) = kdomwb1
81 continue

C in end buffer
if (nznobw.eq.0) then 'dummy' statement
else
nznobot = nznobw
do 82 k = nznobw+1, nznobw+nznobeb
kdomwb(k) = zero
continue
82 end if

do 83 k = 1, nznobw+nznobeb
kdfw(k) = raomwb(k)*kdomwb(k)
83 continue

close (uisu)

```

```

C --- Additional output
if (op.icwby.eq.1) then
  write (uowl, 90)
90 format (' ', Results of some initial calculations from',
  & , subroutine suin:')
  write (uowl, 91)
91 format (' kdfirs(1-nxnotot)', ' (m3/kg)')
  write (uowl, *) (kdfirs(i), i = 1, mxnotot)
end if

if (op.icwb.eq.1) then
  open (unit = uowb, file = 'icwb.out', status = 'unknown')
  write (uowb, 92) dmy, hmg
92 format (' TOXSWA simulation: ', 2A10, /)
  & , Results from initial calculations for sediment:', //)
93 format (' Results of some initial calculations from',
  & , subroutine suin:')
  write (uowb, 94)
94 format (' kdfwb(1-nznotot)', ' (m3/kg)')
  write (uowb, *) (kdfwb(k), k = 1, nznotot)
end if

return
end

subroutine wlit
** SUBROUTINE:
** wlit.for - initial calculations for the water layer of the TOXSWA
** program are executed
** author - P.I. Adriaanse
**
** DESCRIPTION:
** Two subroutines are called:
** # wlitru, in which:
** - the numerical weight factors are calculated for the ditch and
** for the needed buffers;
** # wlitco, in which:
** - the initial concentrations c, Xmp and Xes are calculated at
** every node;
** - the exposure concentration at 0 d (total initial concentration
** c*) is determined.
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C local variables
integer i

C---- first calculation of segment length around every node
C
  if (nxnofb.eq.0) then
    delx(1) = zero
  else
    do 10 i = 1, nxnofb
      if (i.eq.1) then
        delx(i) = (xcdfb(1)+xcdfb(2))/2.
      else
        if (i.eq.nxnofb) then
          delx(i) = (xcdfb(i)+xcdfb(i-1))/2.
        else
          delx(i) = (xcdfb(i)+xcdfb(i-1))/2.
        end if
      end if
    end do
  end if

  do 20 i = nxnofb+1, nxnofb+nxnodit
    if (i.eq.nxnofb+1) .and. nxnofb.ne.0) then
      delx(i) = (xcddit(2)+xcddit(1))/2.
    else
      if (i.eq.(nxnofb+1) .and. nxnofb.eq.0) then
        delx(i) = (xcddit(1)+xcdfb(nxnofb))/2.
      else
        if (i.eq.(nxnofb+1) .and. nxnofb.eq.0) then
          delx(i) = (xcddit(2)+xcddit(1))/2.
        else
          if (i.eq.(nxnofb+1) .and. nxnofb.ne.0) then
            delx(i) = (xcddit(2)+xcddit(1))/2.
          else
            delx(i) = (xcddit(i-nxnofb)+xcddit(i-nxnofb-1))/2.
          end if
        end if
      end if
    end do
  end if

  do 30 i = nxnofb+nxnodit+1, nxnofb+nxnodit+nxnoeb
    if (i.eq.(nxnofb+nxnodit+1) then
      delx(i) = (xcdeb(2)+xcdeb(1))/2.
    else
      if (i.eq.(nxnofb+nxnodit+nxnoeb) then
        delx(i) = (xcdeb(1)+xcddit(nxnodit))/2.
      else
        delx(i) = (xcdeb(i-nxnofb+nxnodit+1) +
          xcdeb(i-nxnofb+nxnodit))/2.
        -
        (xcdeb(i-nxnofb+nxnodit) +
          xcdeb(i-nxnofb+nxnodit-1))/2.
      end if
    end if
  end do
30 continue
end if

C--- Compose vector xcd(nxnotot) with x coordinates in ditch included
C the buffers
C
  if (nxnofb.eq.0) then
    in front buffer
    xcd(1) = 1.
  else
    do 40 i = 1, nxnofb
      xcd(i) = xcdfb(i)
    end do
  end if
40 continue

C
  do 50 i = nxnofb+1, nxnofb+nxnodit
    xcd(i) = xcddit(i-nxnofb)
  end do
50 continue

C
  if (nxnoeb.eq.0) then
    in end buffer
    nxnotot = nxnofb+nxnodit
  else
    do 60 i = nxnofb+nxnodit+1, nxnofb+nxnodit+nxnoeb
      xcd(i) = xcdeb(i-nxnofb+nxnodit)
    end do
  end if
60 continue

```

```

do 295 i = 1, nxnnotot
  write (uowl, *) i, nwmnh(i), nwwph(i)
295 continue
end if
return

end

subroutine whitco (ntot, co, somp, soss)

** SUBROUTINE:
** whitco.for - initial concentrations for the water layer of the
** TOXSWA program are calculated
** author - P.I. Adrianse

** DESCRIPTION:
** The initial concentrations in the water phase, c, sorbed at the
** macropores, xmp, and sorbed at the suspended solids, xss, are
** calculated at every node. The exposure concentration at 0 d (total
** initial concentration, c*) is also determined.

** HISTORY:
** 10 April 1996 - P.I. Adrianse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

C include 'common.for' common variables
C integer i, ii, ntot local variables
C real co(ntot), somp(ntot), soss(ntot), warst,
C col(mnmxnnotot), co2(mnmxnnotot)
C do 30 i = 1, ntot node loop

  warst = wibet*wdhet(i)*(wdht(i)**2)*sisl
  if (cswrwl(i).le.0.) then
    co(i) = zero
    somp(i) = zero
    soss(i) = zero
  else
    col(i) = cswrwl(i)
    co2(i) = cswrwl(i)/(1.+ (dwnmp*pez0hw*kdmp(i))/warst +
    & coss*raomss*kdomss(i)*(col(i)/coobkomss)**
    & (exfrs-1.))
    iteration loop
  C 10 if (abs(co2(i)-col(i)).lt.abs(0.001*co2(i))) go to 20
    col(i) = co2(i)
    co2(i) = col(i)/(1.+ (dwnmp*pez0hw*kdmp(i))/warst +
    & coss*raomss*kdomss(i)*
    & (col(i)/coobkomss)**(exfrs-1.))
    go to 10
  continue
  co(i) = co2(i)
  somp(i) = kmpp(i)*co(i)
  soss(i) = raomss*domss(i)*coobkomss*
  & ((co(i)/coobkomss)**exfrs)
  end if
  C 30 continue

C---- Exposure concentration at 0 d
do 35 i = 1, ntot
do 33 ii = i, nwsby
if (i.eq.iwsby(ii)) then
  exp(ii,ii) = castwl(i)
end if
end if
33 continue
35 continue
if (op_icwlyh.eq.1) then

```

```

end if
if (op_icwlyh.eq.1) then
  write (uowl, 70)
70 format (/ , Results of initial calculations from subroutine whit:')
  write (uowl, 71)
71 format (' delx(1-nxnnotot):', ' (m)')
  write (uowl, *) (delx(i), i = 1, nxnnotot)
72 format (' xcd(1-nxnnotot):', ' (m)')
  write (uowl, *) (xcd(i), i = 1, nxnnotot)
end if

C call whitnu numerical weight factors
C call whitco initial concentrations
call whitco (nxnnotot, costwl, sompst, sossst)
return
end

subroutine whitnu

** SUBROUTINE:
** whitnu.for - initial calculations concerning numerical aspects
** for the water layer of the TOXSWA program are
** calculated
** author - P.I. Adrianse

** DESCRIPTION:
** The numerical weight factors are calculated for the water layer
** and the needed buffers.

** HISTORY:
** 10 April 1996 - P.I. Adrianse
** version 1.0

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.

C include 'common.for' common variables
C integer i local variables

C---- numerical weight factors at the segment boundaries
do 230 i = 2, nxnnotot-1
  if (betawl.ge.zero .and. betawl.le.0.5) then
    nwmnh(i) = (betawl*delx(i-1))/(xcd(i)-xcd(i-1))
    nwwph(i) = (betawl*delx(i))/(xcd(i+1)-xcd(i))
  else
    if (betawl.gt.0.5 .and. betawl.le.1.0) then
      nwmnh(i) = (0.5*delx(i-1) + (betawl-0.5)*delx(i)) /
      & (xcd(i)-xcd(i-1))
      nwwph(i) = (0.5*delx(i) + (betawl-0.5)*delx(i-1)) /
      & (xcd(i+1)-xcd(i))
    end if
  end if
230 continue

nwmnh(1) = nwmnh(2)
nwmnh(nxnnotot) = nwwph(nxnnotot-1)

if (op_icwlyh.eq.1) then
  write (uowl, 235)
235 format (/ , Results of calculations from subroutine whitnu:')
  write (uowl, 285)
285 format (' num. weight factors for all nodes:', / , (dimensionless)')
  & , nodenumber nwmnh nwwph

```

```

write (uowl, 40)
40 format (/, 'Results of calculations from subroutine whitco', /,
& write (uowl, *) length exchanging perimeter, pezChw:', (m)')
write (uowl, 41) pezOhw
41 format (' initial wetted area ditch, warst:', (m2)')
write (uowl, *) warst
42 format (' initial water depths for all nodes, whst(1-nxznnotot):',
& write (uowl, *) (whst(i), i = 1, ntot)
write (uowl, 43)
43 format (' initial concentrations for all nodes', /,
& ' nodenr, total, dissolved, at macrophytes and at susp solids', /,
& ' , ixnotot, castwl(i) Co(i) and',
& ' , soss(i)', (g.(m)-3 g.(m)-3 g.(g)-1 ' ,
& /, ' , g.(g)-1)',
& do 44 i = 1, ntot
write (uowl, *) i, castwl(i), co(i), somp(i), soss(i)
44 continue
end if
return
end

```

```

subroutine whbt
** SUBROUTINE:
** wht for - initial calculations for the sediment are executed
** author - P.I. Adrianse
**
** DESCRIPTION:
** The segment lengths in the sediment plus buffer are determined.
** Three subroutines are called:
** # whbtu, in which:
** - the numerical weight factors are calculated for the sediment
** and the buffer;
** # whtgc, in which:
** -the perimeter at a distance d of the water-sediment interface
** is calculated (Pk-1/2, Pk en Pk+1/2);
** -the porosity (epsilon) is calculated for k-1/2 and k+1/2;
** -the bulk density is calculated for k-1/2 and k+1/2;
** -the Freundlich sorption coefficient, Kf, is calculated at k-1/2
** and k+1/2;
** -the diffusion coefficients for the pesticide in the liquid phase
** of the sediment, Dlb,k-1/2 and Dlb,k+1/2, are calculated;
** # whtco, in which:
** - the initial concentrations in the liquid phase, c1b, and sorbed
** to solid sediment material, Xb, are calculated at every node.
** HISTORY:
** 10 April 1996 - P.I. Adrianse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C local variables
integer k
C---- first calculation of segment length around every node
C
do 10 k = 1, nznobw
if (k.eq.1) then
delz(k) = (zcdwb(1)+zcdwb(2))/2.
else
if (k.eq.nznobw) then
delz(k) = zwb - (zcdwb(k-1)+zcdwb(k))/2.
else
delz(k) = (zcdwb(k)+zcdwb(k+1))/2. -
(zcdwb(k-1)+zcdwb(k))/2.
&
end if
end if
10 continue
C
if (nznobw.eq.0) then
in end buffer
C
nznnotot = nznobw+nznobbb
else
do 20 k = nznobw+1, nznobw+nznobbb
if (k.eq.nznobw+1) then
delz(k) = (zcdwb(2)+zcdwb(1))/2. -
(zcdwb(nznobw)+zcdwb(1))/2.
&
else
if (k.eq.nznobw+nznobbb) then
delz(k) = zwb+zebb -
(zcdwb(k-nznobw)+zcdwb(k-nznobw-1))/2.
&
else
delz(k) = (zcdwb(k-nznobw+1)+zcdwb(k-nznobw))/2. -
(zcdwb(k-nznobw)+zcdwb(k-nznobw-1))/2.
&
end if
end if
20 continue
end if
C---- Compose vector zcd(nznnotot) with z coordinates in sediment
C included the buffer
C
do 30 k = 1, nznobw
zcd(k) = zcdwb(k)
30 continue
C
if (nznobbb.eq.0) then
in end buffer
C
nznnotot = nznobw+nznobbb
else
do 40 k = nznobw+1, nznobw+nznobbb
zcd(k) = zcdwb(k-nznobw)
40 continue
end if
if (op.icwb.eq.1) then
write (uowb, 50)
50 format (/, 'Results of initial calculations',
& write (uowb, 51) ' from subroutine whbt:')
51 format (' delz(1-nznnotot) from whbt:', (m)')
write (uowb, *) (delz(k), k = 1, nznnotot)
52 format (' zcd(1-nznnotot) from whbt:', (m)')
write (uowb, *) (zcd(k), k = 1, nznnotot)
end if
C call whbtu - numerical weight factors
C - perimeters
C - porosities
C - diffusion coefficients
C call whtgc - initial concentrations
call whtco (nznnotot, costwb, sowlbst)
return
end
subroutine whbtu
** SUBROUTINE:

```



```

** wbitnu.for - initial calculations concerning numerical aspects for
** the sediment of the TOXSWA program are executed
** author - P.I. Adriaanse
**
** DESCRIPTION:
** The numerical weight factors are calculated for the sediment
** and the end buffer.
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C integer k local variables
C
C----- numerical weight factors at the segment boundaries
do 10 k = 2, nznnotot-1
if (betawb.ge.zero and .betaawb.le.0.5) then
nwbbmh(k) = (betaawb*delz(k-1))/(zcd(k)-zcd(k-1))
nwbbph(k) = (betaawb*delz(k))/(zcd(k+1)-zcd(k))
else
if (betaawb.gt.0.5 and .betaawb.le.1.0) then
nwbbmh(k) = (0.5*delz(k-1))/(betaawb-0.5)*delz(k) /
& (zcd(k)-zcd(k-1))
nwbbph(k) = (0.5*delz(k))/(betaawb-0.5)*delz(k+1) /
& (zcd(k+1)-zcd(k))
&
& end if
&
10 continue
nwbbph(1) = nwbbmh(2)
nwbbmh(nznnotot) = nwbbph(nznnotot-1)
if (op_iwb.eq.1) then
write (uowb, 20)
&
20 format (/)
&
21 format (' num. weight factors for all nodes: ', /,
& nodenumber nwbbmh nwbbph (dimensionless)')
do 22 k = 1, nznnotot
write (uowb, *) k, nwbbmh(k), nwbbph(k)
22 continue
end if
return
end
subroutine wbitge
subroutine wbitge
** SUBROUTINE:
** wbitge.for - initial calculations concerning sediment and sorption
** characteristics for the TOXSWA program are executed
** author - P.I. Adriaanse
**
** DESCRIPTION:
** The following initial calculations are executed here:
** -the perimeter at a distance d of the water-sediment interface
** is calculated (PK-1/2, PK en PK+1/2);
** -the porosity (epsilon) is calculated for k-1/2 and k+1/2;
** -the bulk density is calculated for k-1/2 and k+1/2;
** -the Freundlich sorption coefficient, KF, is calculated for k-1/2
** and k+1/2;
** -the diffusion coefficients for the pesticide in the liquid phase
** of the sediment, Dib,k-1/2 and Dib,k+1/2, are calculated.

```

```

** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C integer k local variables
real betasisl, sumdelz(mxnznnotot), tormb, torph,
& raomwmbmb, raomwbbph
C----- Calculate perimeter at depth d from the water-sediment interface,
C PK-1/2, Pk and Pk+1/2
C P at z=0, pez0hw, has already been calculated in wlin.for
betasisl = atan(1./sisl)
do 10 k = 1, nznnotot
pa(k) = wibot + 2.*zcd(k)*tan(0.5*betasisl) +
& 2.*(wdhfl+zcd(k))*sqrt((sisl**2)+1.)
10 continue
do 15 k = 1, nznnotot
sumdelz(k) = zcd(k) + 0.5*delz(k)
15 continue
do 20 k = 1, nznnotot
if (k.eq.1) then
pemb(k) = pez0hw
else
pemb(k) = wibot + 2.*sumdelz(k-1)*tan(0.5*betasisl) +
& 2.*(wdhfl+sumdelz(k-1))*sqrt((sisl**2)+1.)
&
20 continue
do 25 k = 1, nznnotot
peph(k) = wibot + 2.*sumdelz(k)*tan(0.5*betasisl) +
& 2.*(wdhfl+sumdelz(k))*sqrt((sisl**2)+1.)
25 continue
C----- Calculate porosity, epsilon(k-1/2) and epsilon(k+1/2), in
C the sediment included the buffer
C do 30 k = 1, nznnotot
if (k.eq.1) then
porph(k) = por(1)
porph(k) = (delz(k+1)/(delz(k)+delz(k+1)))*por(k) +
& (delz(k)/(delz(k)+delz(k+1)))*por(k+1)
&
else
if (k.eq.nznnotot) then
porph(k) = (delz(k)/(delz(k-1)+delz(k)))*por(k-1) +
& (delz(k-1)/(delz(k-1)+delz(k)))*por(k)
&
else
porph(k) = (delz(k)/(delz(k-1)+delz(k)))*por(k-1) +
& (delz(k-1)/(delz(k-1)+delz(k)))*por(k)
&
& end if
&
30 continue
C----- Calculate bulk density, bdbw(k+1/2) and bdbw(k-1/2), in
C the sediment included the buffer
C do 40 k = 1, nznnotot
if (k.eq.1) then
bwbmbh(k) = bdbwb(1)
bwboph(k) = (delz(k+1)/(delz(k)+delz(k+1)))*bdbwb(k) +
& (delz(k)/(delz(k)+delz(k+1)))*bdbwb(k+1)
&

```

```

82 continue
   write (uowb, 83)
83 format(' porosity at interfaces for all nodes:', /,
   & , nodenumber , porph', /,
   & , (dimensionless)')
   do 84 k = 1, nznnotot
   write (uowb, *) k, porph(k), porph(k)
84 continue
   write (uowb, 85)
85 format(' bulkdensity at interfaces for all nodes:', /,
   & , nodenumber , bdwbph , bdwbph ,
   & do 86 k = 1, nznnotot
   write (uowb, *) k, bdwbmh(k), bdwbph(k)
86 continue
   write (uowb, 87)
87 format(' Freundlich sorption coefficient at interfaces for',
   & , all nodes:', /,
   & , nodenumber , kdfwbmh , kdfwbph ,
   & do 88 k = 1, nznnotot
   write (uowb, *) k, kdfwbmh(k), kdfwbph(k)
88 continue
   write (uowb, 89)
89 format(' diffusion coefficients at interfaces for all nodes:',
   & , nodenumber , kdfwbmh ,
   & do 90 k = 1, nznnotot
   write (uowb, *) k, kdfwbmh(k), kdfwbph(k)
90 continue
   end if
   return
end

subroutine wbitco (ztot, co, sowb)
** SUBROUTINE:
** wbitco.for - initial concentrations for the sediment layer for
** the TOKSWA program are executed
** author - P.I. Adriaanse
**
** DESCRIPTION:
** The initial concentrations in the liquid phase, clb, and sorbed to
** solid sediment material, Xb, are calculated at every node.
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C integer k, ztot local variables
C real co(ztot), sowb(ztot),
C col{mxnznnotot}, co2{mxnznnotot}
C do 30 k = 1, ztot node loop
   if (castwb(k).le.0.) then
   co(k) = zero
   else
   col(k) = castwb(k)
   co2(k) = castwb(k)/(por(k) + bdwb(k)*kdfwb(k)*
   (col(k)/coobkxmb)**(exfrwb-1.))
   iteration loop
C 10 if (abs(co2(k)-col(k)).lt.abs(0.001*co2(k))) go to 20

```

```

else
   if (k.eq.nznnotot) then
   if (kdfwb(k)/(delz(k-1)+delz(k))*bdwb(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*bdwb(k)
   & bdwbph(k) = bdwb(nznnotot)
   else
   bdwbmh(k) = (delz(k)/(delz(k-1)+delz(k))*bdwb(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*bdwb(k)
   & bdwbph(k) = (delz(k)/(delz(k-1)+delz(k))*bdwb(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*bdwb(k)
   & kdfwb(k-1/2), in the sediment included the buffer
C
C in sediment included the buffer
C do 50 k = 1, nznnotot
   if (k.eq.1) then
   raomwb = raomwb(1)
   raomwbph = (delz(k)/(delz(k)+delz(k+1))*raomwb(k) +
   & (delz(k+1)/(delz(k)+delz(k+1))*raomwb(k+1)
   else
   if (k.eq.nznnotot) then
   raomwbmh = (delz(k)/(delz(k-1)+delz(k))*raomwb(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*raomwb(k)
   & raomwbph = raomwb(nznnotot)
   else
   raomwbmh = (delz(k)/(delz(k-1)+delz(k))*raomwb(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*raomwb(k)
   & raomwbph = (delz(k)/(delz(k)+delz(k+1))*raomwb(k) +
   (delz(k+1)/(delz(k)+delz(k+1))*raomwb(k+1)
   end if
   end if
   kdfwbmh(k) = raomwbmh * kdomwb(k)
   kdfwbph(k) = raomwbph * kdomwb(k)
50 continue
C
C Avoid that two values exist for kdfwb at
C interface between sediment and endbuffer
C kdfwbph(nznwb) = (kdfwb(nznwb)+kdfwb(nznwb+1))/2.
C kdfwbmh(nznwb-1) = (kdfwb(nznwb)+kdfwb(nznwb+1))/2.
C
C---- Calculate diffusion coefficient Dlb,k-1/2 and Dlb,k-1/2 in
C the sediment included the buffer
C
C in sediment included the buffer
C do 70 k = 1, nznnotot
   if (k.eq.1) then
   tormh = tor(1)
   torph = (delz(k+1)/(delz(k)+delz(k+1))*tor(k) +
   & (delz(k)/(delz(k)+delz(k+1))*tor(k+1)
   else
   if (k.eq.nznnotot) then
   tormh = (delz(k)/(delz(k-1)+delz(k))*tor(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*tor(k)
   & torph = tor(nznnotot)
   else
   tormh = (delz(k)/(delz(k-1)+delz(k))*tor(k-1) +
   (delz(k-1)/(delz(k-1)+delz(k))*tor(k)
   & torph = (delz(k)/(delz(k)+delz(k+1))*tor(k) +
   (delz(k+1)/(delz(k)+delz(k+1))*tor(k+1)
   end if
   end if
   kdfwbmh(k) = tormh*kdfw
   kdfwbph(k) = torph*kdfw
70 continue
   if (op-icwb.eq.1) then
   write (uowb, 80)
80 format(' , Results of initial calculations from subroutine',
   & , wbitge:')
   write (uowb, 81)
81 format(' length perimeter at nodes and interfaces for',
   & , all nodes:', /, nodenumber ,
   & , peph ,
   & do 82 k = 1, nznnotot
   write (uowb, *) k, pemh(k), pe(k), pep(k)

```

```

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** Research (SC-DLO), 1996.
**

C include 'common.for' common variables
C integer i, ii, ittle local variables
C real rh(mxnnotot) poscounter
C logical loghelp, convergence_OK

do 10 ixnotot = 1, nxnotot
  if (its.eq.1) then
    else cowlj(ixnotot) = costw1(ixnotot)
  end if
  end if
10 continue

do 20 ixnotot = 1, nxnotot
  read concentration c, lb of first
  node sediment for each node water layer;
  calculate seepage/infiltration plus
  concentration for j and j+1
  calculate wetted areas, water depths, width
  water surface and discharges at nodes and
  segment boundaries at time j, j+1/2 and j+1
  call wisoin
  call wisoge
20 continue

C----- Calculate numerical and calculation dispersion
do 30 ixnotot = 1, nxnotot
  call wisods
30 continue
  if (its.eq.1) then
    poscounter = 1
  end if
  poscounter counts the number of times that
  the positivity conditions are not
  fulfilled

  ittle = 1
  Loop while no final value of cowljpl. (So start of iteration
  procedure due to sorption according to Freundlich equation to
  find c at time j+1)
40 continue

C----- Make estimation for c at time j+1 for calculation of lod, ldd and
  lbd
do 50 ixnotot = 1, nxnotot
  if (ittlo.eq.1) then
    oldcowljpl(ixnotot) = cowlj(ixnotot)
  else
    oldcowljpl(ixnotot) = cowljpl(ixnotot)
  end if
C----- Compose left-hand and right-hand matrices
C
C if (cowlj(ixnotot).le.zero) then
  ldd, lbd and rod, rdd and rbd terms
  cowlj(ixnotot) = tiny
end if
if (oldcowljpl(ixnotot).le.zero) then
  oldcowljpl(ixnotot) = tiny
end if
50 continue

co1(k) = co2(k)
co2(k) = castwb(k)/por(k) + bdeb(k)*kdrwb(k)*
  (co1(k)/coobkomwb)**(exfrwb-1.1)
  go to 10
continue = co2(k)
co1(k)
sowb(k) = kdrfwb(k)*coobkomwb*(co(k)/coobkomwb)**exfrwb
end if
30 continue

if (op.icwb.eq.1) then
  write (uowb,40)
40 format (/, 'Results of initial calculations from subroutine',
  & , whitcov,1)
  write (uowb,41)
41 format (, 'Initial concentrations for all nodes', /,
  & , noder, totat, dissolved and sorbed at solid',
  & , bottom mat. /,
  & , kznnotot castwb(k) co(k) sowb(k) /,
  & , do 42 k = 1, zto (g.(m)-3 g.(g)-1) /,
  & , write (uowb,*) k, castwb(k), co(k), sowb(k)
42 continue
end if
return
end

subroutine wiso
** SUBROUTINE:
** wiso.for - solves the mass conservation equation for the water
** layer of the TOXSWA program
** author - P.I. Adriaanse
**
** DESCRIPTION:
** The mass conservation equation for the sediment subsystem is
** solved by solving the matrix equation, containing two tridiagonal
** matrices describing the pesticide concentration in the liquid
** phase of the sediment (Eq. (6.39) in SC-DLO report 90). The
** equation can be solved in an implicit or in an explicit way and
** due to the sorption description according to the Freundlich
** equation an iteration takes place to calculate the final
** pesticide concentrations.
** Exposure concentrations at 0, 3, 21 and 28 day are also
** calculated in this subroutine.
** The following subroutines are called:
** # wisoin, in which:
** - concentration and seepage rates are read for timestep j;
** # wisoge, in which:
** - some geometrical characteristics for the water layer are
** calculated;
** # wisods, in which:
** - the (physical) dispersion coefficient is corrected for the
** numerical dispersion;
** # wisocm, in which:
** - the elements of the tridiagonal matrices are defined, which are
** used in the subroutines wisoub, wisolb and wisocr;
** # wisolb, in which:
** - the lod, ldd etc terms of the matrices for the lower boundary of
** the water layer are composed;
** # wisoub, in which:
** - the lod, ldd etc terms of the matrices for the upper boundary of
** the water layer are composed;
** # wisocr, in which:
** - the lod, ldd etc terms of the core parts of the matrices for the
** water layer are composed;
** # tridag, in which:
** - the tridiagonal matrix for the water layer is inverted.
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0

```



```

C      if (.not. convergency_OK) then
C          if (ittlo.eq.50) then
C              write (uwer, 91)
C          format (//, 'Message from subroutine wiso: ')
C          write (uwer, 92) ittlo, ixnotot, its
C          format (/, 'More than', /,
C              16, ' iteration loops will be needed to calculate c j+1',
C              /, ' for gridpoint', i6, ' at timestep', i8,
C              /, ' in the water layer')
C          write (uwer, 93) oldcowjpl(ixnotot), cowjpl(ixnotot)
C          format (/, 'Iteration values oldcowjpl and cowjpl are:', E10.4,
C              print*, ' Error: ',
C              print*, ' Never ending loop in wiso; ',
C              print*, ' see message out',
C              print*, ' Program stops',
C              stop
C          end if
C          end if
C          if (.not. convergency_OK) then
C              ittlo = ittlo + 1
C              go to 40
C          end if
C      C---- End of loop while no final value of cowjpl

C---- Now final value of c at time j+1, cowjpl found
C---- Calculate Xes, Xmp and c* at time j+1 at each node
do 100 ixnotot = 1, nmxnotot
if (cowjpl(ixnotot).lt.zero) then
write (uwer, 96)
96 format (//, 'Message from subroutine wiso: ')
write (uwer, 97) its, ixnotot, cowjpl(ixnotot)
97 format (/, ' Negative concentration in ditch', /,
& ' at timestep', i6, ' and node', i3, /,
& ' Concentration is:', E10.4)
cowjpl(ixnotot) = tiny
end if
C---- When the concentration at an individual node becomes smaller than
C IE-25 the concentration will be set at IE-25. (At concentrations
C in the order of IE-34 no convergency will be reached in the
C iteration loop at statement label 90 in this subroutine wiso.for.
C Also other numerical problems due to computer range limitations
C will be prevented in this way.)
if (cowjpl(ixnotot).lt.1.0E-25) then
cowjpl(ixnotot) = 1.0E-25
end if
sossjpl(ixnotot) = raomss*xdomss(ixnotot)*coobkoms*
& ((cowjpl(ixnotot)/ixnotot)/coobkoms)*exfrss
sompjpl(ixnotot) = kdmp(ixnotot)*cowjpl(ixnotot)
& cowjpl(ixnotot) + iixnotot)/warjpl(ixnotot)
& dmp*pezhw*sompjpl(ixnotot)/warjpl(ixnotot)
+ coss*sossjpl(ixnotot)
100 continue

C---- Calculation of exposure concentrations
C      tcoxp(1) = 0.00
C      tcoxp(2) = 3.00
C      tcoxp(3) = 21.00
C      tcoxp(4) = 28.00
do 103 ixnotot = 1, nmxnotot
do 102 i = 1, nwbxy
initialisation of sum of product of
pesticide concentration in water phase
* time step, separately for each node
selected for output
C          if (its.eq.1 .and. ixnotot.eq.iwbxy(i)) then
C              sumct(i) = zero
C              calculation of sumct
C              sumct(i) = deltwl*cowjpl(ixnotot) + sumct(i)
C              if (ixnotot.eq.iwbxy(i)) then
C                  do 101 ii = 2, 4
C                      if (nint(tcexp(ii)).le.nint(ttot-0.5)) then
C                          if (tcexp(ii)*86400.) .ge. (its*deltwl) .and.
C                              (tcexp(ii)*86400.) .lt. ((its+1)*deltwl) then
C                              sumct(ii,ii) = sumct(i)/(its*deltwl)
C                          end if
C                      else
C                          sumct(ii,ii) = -100.
C                      end if
C                  continue
C              end if
C              output of exposure concentrations in output
C              file icwly.out
C              if (op_icwly.eq.1) then
C                  if (its.eq.1) then
C                      write (uowl, 104)
C                  format (//, 'Output from wiso for', /, 3, 21 and 28 days', /,
C                      & ' Exposure concentrations at 0, 3, 21 and 28 days', /,
C                      & ' , node output dist.in ditch time exp concentration', /,
C                      & ' (m) (d)')
C                  end if
C              do 107 ixnotot = 1, nmxnotot
C                  do 106 i = 1, nwbxy
C                      if (ixnotot.eq.iwbxy(i)) then
C                          if (its.eq.1) then
C                              write (uowl, 105) ixnotot, xcd(ixnotot),
C                                  tcoxp(1), tcoxp(2), tcoxp(3), tcoxp(4)
C                              end if
C                          if ((tcoxp(2)*86400.) .ge. (its*deltwl) .and.
C                              (tcoxp(2)*86400.) .lt. ((its+1)*deltwl)) then
C                              write (uowl, 105) ixnotot, xcd(ixnotot),
C                                  tcoxp(2), tcoxp(3), tcoxp(4)
C                              end if
C                          if ((tcoxp(3)*86400.) .ge. (its*deltwl) .and.
C                              (tcoxp(3)*86400.) .lt. ((its+1)*deltwl)) then
C                              write (uowl, 105) ixnotot, xcd(ixnotot),
C                                  tcoxp(3), tcoxp(4)
C                              end if
C                          if ((tcoxp(4)*86400.) .ge. (its*deltwl) .and.
C                              (tcoxp(4)*86400.) .lt. ((its+1)*deltwl)) then
C                              write (uowl, 105) ixnotot, xcd(ixnotot),
C                                  tcoxp(4)
C                              end if
C                          105 format (i8, 5x, f8.2, 7x, f6.2, 2x, E12.5)
C                              end if
C                      end if
C                  continue
C              if (its.eq.nrow) .and. ixnotot.eq.nmxnotot) then
C                  close (unit = uowl)
C              end if
C          end if
C          end if
C      C---- End of calculation of exposure concentrations

```



```

C C      prevent undefined exponentiation in
C C      calculation of Enum
C C      if (cowljimb(ixnotot).le.zero) then
C C      end if
C C      if (cowljiph(ixnotot).le.zero) then
C C      end if
C C      if (ixnotot.gt.1) then
C C      kdswnujimb = (qvoijimb/(2.*warjimb(ixnotot))) *
C C      (1.-2.*nwvmb(ixnotot))
C C      *kcd(ixnotot)-xcd(ixnotot-1)+(1.-2.*thetawl)*
C C      (qvoijimb/warjimb(ixnotot))*deltwl *
C C      (1.+coss*kdfrrs(ixnotot)*exfrss *
C C      (cowljimb(ixnotot)/coobkoms)**(exfrss-1.)) /
C C      (1.-dvwmp*pez0hw*kdmpl(ixnotot)/warjimb(ixnotot) +
C C      coss*kdfrrs(ixnotot)*exfrss *
C C      ((cowljimb(ixnotot)/coobkoms)**(exfrss-1.)))
C C      end if
C C      if (ixnotot.lt.nxnotot) then
C C      kdswnujiph = (qvoijiph/(2.*warjiph(ixnotot))) *
C C      (1.-2.*nwvph(ixnotot))
C C      *kcd(ixnotot+1)-xcd(ixnotot)*deltwl *
C C      (qvoijiph/warjiph(ixnotot))*deltwl *
C C      (1.+coss*kdfrrs(ixnotot)*exfrss *
C C      (cowljiph(ixnotot)/coobkoms)**(exfrss-1.)) /
C C      (1.-dvwmp*pez0hw*kdmpl(ixnotot)/warjiph(ixnotot) +
C C      coss*kdfrrs(ixnotot)*exfrss *
C C      ((cowljiph(ixnotot)/coobkoms)**(exfrss-1.)))
C C      end if
C C      if (ixnotot.gt.1) then
C C      if kdswecljimb(ixnotot) = kds - kdswnujimb
C C      end if
C C      if (ixnotot.lt.nxnotot) then
C C      if kdswecljiph(ixnotot) = kds - kdswnujiph
C C      end if
C C      if (op.icwlvb.eq.1) then
C C      if (fits.eq.1 and ixnotot.eq.1) then
C C      write (uowl,5)
C C      5 format (/,' Some calculation results from subroutine wisods', /,
C C      /,' physical, numerical and calculation dispersion for', /,
C C      /,' first timestep and first node',
C C      /,' , kds kdswnujiph kdswecljiph (m2.(day)-1)',
C C      write (uowl,*) 86400.*kds, 86400.*kdswnujiph,
C C      &
C C      end if
C C      return
C C      end
C C      subroutine wisoom
C C      wisoom.for - the elements of the tridiagonal matrices for the
C C      water layer of the TOXSWA program are defined
C C      here
C C      author - P.I. Adriaanse
C C      ** SUBROUTINE:
C C      ** wisoom.for - the elements of the tridiagonal matrices for the water
C C      ** layer are defined here; these elements are used in the subroutines
C C      ** wisob, wisob and wisoc.
C C      ** HISTORY:
C C      ** 10 April 1996 - P.I. Adriaanse
C C      ** version 1.0
C C      **

```

```

** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C      include 'common.for'      common variables
C      local variables          local variables
C
C----- Elements of the core and buffer matrices will be composed here
C
C      the terms lod., lddadmh, ldddsmb, rod.,
C      rddadmh and rdddsmb will not be defined for
C      ixnotot=1, as well as lddadph, ldddsph,
C      lbd., rddadph, rdddsph and rbd. for
C      ixnotot=ixnotot, because they do not
C      appear in the two tridiagonal matrices
C
C      if (ixnotot.gt.1) then
C      lodad = (qvoijimb(ixnotot)/delx(ixnotot))*(1.-thetawl)*
C      (1.-nwvmb(ixnotot))*deltwl * (1.+coss*
C      kdfrrs(ixnotot-1)*
C      ((oldcowljip(ixnotot-1)/coobkoms)**(exfrss-1.)))
C      lodds = (2.*warjiphmh(ixnotot)*kdswecljimb(ixnotot)/
C      delx(ixnotot))*delx(ixnotot)*deltwl *
C      (1.-thetawl)* deltwl *
C      (1.+coss*kdfrrs(ixnotot-1)*
C      ((oldcowljip(ixnotot-1)/coobkoms)**
C      (exfrss-1.)))
C      lddadmh = (qvoijimbmh(ixnotot)/delx(ixnotot))*(1.-thetawl)*
C      nwvmbmh(ixnotot)*deltwl * (1.+coss*
C      kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.)))
C      end if
C      if (ixnotot.lt.nxnotot) then
C      lddadph = (qvoijiphix(ixnotot)/delx(ixnotot))*(1.-thetawl)*
C      (1.-nwvph(ixnotot))*deltwl * (1.+coss*
C      kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.)))
C      end if
C      if (ixnotot.gt.1) then
C      ldddsmb = (2.*warjiphmh(ixnotot)*kdswecljimb(ixnotot)/
C      delx(ixnotot))*delx(ixnotot)*deltwl *
C      (1.-thetawl)*deltwl *
C      (1.+coss*kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.)))
C      end if
C      if (ixnotot.lt.nxnotot) then
C      ldddsph = (2.*warjiphix(ixnotot)*kdswecljiph(ixnotot)/
C      delx(ixnotot))*delx(ixnotot+1)*delx(ixnotot))
C      (1.-thetawl)*deltwl *
C      (1.+coss*kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.)))
C      end if
C      lddca = (1.+((dvwmp*pez0hw)/warjip(ixnotot))*kdmpl(ixnotot)/
C      +coss*kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.))) *
C      (warjip(ixnotot))
C      lddtf = (1.+((dvwmp*pez0hw)/warjip(ixnotot))*kdmpl(ixnotot)+
C      coss*kdfrrs(ixnotot)*
C      ((oldcowljip(ixnotot)/coobkoms)**(exfrss-1.))) *
C      (ktrfel*warjiph(ixnotot)*(1.-thetawl)*
C      deltwl)
C      lddai = ktl*witopiph(ixnotot)*(1.-thetawl)*deltwl
C      lddwab = (leplot/pez0hw)*gseifjph*pez0hw*
C      lddwbf = ((2.*por(1))*kdfwmbh(1)*pez0hw/delz(1))**
C      (1.-thetawl)*deltwl
C
C      if (ixnotot.lt.nxnotot) then
C      lbdad = (qvoijiphix(ixnotot)/delx(ixnotot))*(1.-thetawl)*
C      nwvphix(ixnotot)*deltwl*(1.+coss*kdfrrs(ixnotot-1)*
C      ((oldcowljip(ixnotot-1)/coobkoms)**(exfrss-1.)))
C      lbdbs = ((2.*warjiphix(ixnotot)*kdswecljiph(ixnotot)/
C      delx(ixnotot))*delx(ixnotot+1)*delx(ixnotot))
C      * (1.-thetawl)*deltwl *

```



```

C
else
  now qseifjph less than zero
  if (ueb.ge.zero) then
    lod(ixnotot) = -lddad - lodds
    ldd(ixnotot) = -lddadmh +
      (qvojpiph(ixnotot)/deltx(ixnotot))*
      (1 - thetau)*deltw
      * (1 + coss*kdfrrs(ixnotot)*
        ((oldcowljpl(ixnotot)/coobkoms)**
          (exfrss-1.)))
      + lddasmh +
      lddca - lddrf + lddai + lddwdbf
    rod(ixnotot) = rodad + rodas
    rdd(ixnotot) = rddadmh -
      (qvojpiph(ixnotot)/deltx(ixnotot))
      * thetau*deltw
      * (1 - coss*kdfrrs(ixnotot)*
        ((cowlj(ixnotot)/coobkoms)**
          (exfrss-1.)))
      - rddasmh +
      rddca - rddrf - rddai - rddwdbf
    rv(ixnotot) = rvai + rvwdbf - rvwbad
  else
    now ueb less than zero
    lod(ixnotot) = -lddad - lodds
    ldd(ixnotot) = -lddadmh + lddasmh +
      lddca - lddrf + lddai + lddwdbf
    rod(ixnotot) = rodad + rodas
    rdd(ixnotot) = rddadmh - rddasmh +
      rddca - rddrf - rddai - rddwdbf
    rv(ixnotot) = rvai + rvwdbf - rvwbad
  end if
end if
end if

if (ixnotot.eq.ixnotot .and. nnoeb.ne.0) then
  last row elements in matrix for lower
  boundary; in case of an end buffer
  first determine direction of flow u
  if (ueb.ge.zero) then
    lod(ixnotot) = -lddad - lodds
    ldd(ixnotot) = -lddadmh +
      (qvojpiph(ixnotot)/deltx(ixnotot))*
      (1 - thetau)*deltw
      * (1 + coss*kdfrrs(ixnotot)*
        ((oldcowljpl(ixnotot)/coobkoms)**
          (exfrss-1.)))
      + lddasmh
      + (oldcowljpl(ixnotot)/coobkoms)**
        (exfrss-1.)))
      * warij(ixnotot)
    rod(ixnotot) = rodad + rodas
    rdd(ixnotot) = rddadmh -
      (qvojpiph(ixnotot)/deltx(ixnotot))*
        thetau*
      * (1 + coss*kdfrrs(ixnotot)*
        ((cowlj(ixnotot)/coobkoms)**
          (exfrss-1.)))
      * warij(ixnotot)
  else
    (qvojpiph(ixnotot)/deltx(ixnotot))*thetaw1*
    deltaw1*
    ((cowlj(ixnotot)/coobkoms)**(exfrss-1.)))
    - rddasmh +
    (1 + coss*kdfrrs(ixnotot)*
      ((cowlj(ixnotot)/coobkoms)**(exfrss-1.)))
      * warij(ixnotot)
  end if
  if (ueb.lt.zero) then
    lod(ixnotot) = -lddad - lodds
    ldd(ixnotot) = -lddadmh + lddasmh +
      (oldcowljpl(ixnotot)/coobkoms)**
        (exfrss-1.)))
      * warij(ixnotot)
    rod(ixnotot) = rodad + rodas
    rdd(ixnotot) = rddadmh - rddasmh +
      (1 + coss*kdfrrs(ixnotot)*
        ((cowlj(ixnotot)/coobkoms)**
          (exfrss-1.)))
      * warij(ixnotot)
  end if
end if

if (ixnotot.ne.nxnofb.nxnodit) then
  terms lod, ldd, lbd, rod, rdd, rbd and rv
  for grid points m+1 until ebt-1 included
  in the end buffer
  lod(ixnotot) = -lddad - lodds
  ldd(ixnotot) = -lddadmh + lddadph + lddasmh + ldddsph +
    ((oldcowljpl(ixnotot)/coobkoms)**
      (exfrss-1.))) * warij1(ixnotot)
  lbd(ixnotot) = lbdad - lbdads
  rod(ixnotot) = rodad + rodas
  rdd(ixnotot) = rddadmh - rddadph - rddasmh - rdddsph +
    (1 + coss*kdfrrs(ixnotot)*
      ((cowlj(ixnotot)/coobkoms)**(exfrss-1.)))
      * warij(ixnotot)
  rbd(ixnotot) = -rbdad + rbdads
  rv(ixnotot) = zero
end if
end if

subroutine wlsoub
** SUAROUTINE:
** wlsoub.for - the lod, ldd etc terms of the matrices for the upper
** boundary of the water layer of the TOXSWA program
** are composed here
** author - P.I. Adriaanse
**
** DESCRIPTION:
** The matrices for the first node or front buffer of the water
** layer will be composed for this timestep. (Only the nonzero terms
** of the left-hand and right-hand triangular matrices will be
** stored, they form two times three vectors; notice that lod(l),
** lbd(nxnorb), rod(l) and rbd(ixnotot) are undefined and not
** referenced in the subroutines.)
**
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for' common variables
C local variables
C
C flow velocity has been positive since start
C calculations, no front buffer needed,
C nxnofb=0 and first row core matrix is
C composed
C if (nxnofb.eq.0 .and. ixnotot.eq.1 .and. ufb.ge.zero) then
  if (qseifjph.ge.zero) then
    ldd(l) = lddadph + ldddsph + lddca + lddrf + lddai +
      lddwad + lddwdbf
    lbd(l) = lbdad - lbdads
    rdd(l) = -rddadph - rdddsph + rddca - rddrf - rddai -
      rbd(l) = -rbdad + rbdads
    rv(l) = rvai + rvwdbf
  else
    if (qseifjph.lt.zero) then
      if ldd(l) = lddadph + ldddsph + lddca + lddrf + lddai +
        lddwad +
      lbd(l) = lbdad - lbdads
    end if
  end if
end if

```



```

** vectors; notice that lod(1), lbd(mxnotot), rod(1) and rbd(mxnotot)
** are undefined and not referenced in the subroutines.)
**
** HISTORY:
** 10 April 1996 - P.I. Adriause
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for'      common variables
C                            local variables
if (gseifiph.ge.zero) then
  lod(ixnotot) = -lodad - lodds
  & ldd(ixnotot) = -lddadmh + lddasph + ldddsph + ldddsph + lddca +
  & ldd(ixnotot) = lddtf + lddai + lddwbad + lddwbdf
  & lbd(ixnotot) = lbdad - lbdbs
  rod(ixnotot) = rodad + rodds
  rdd(ixnotot) = rddasmh - rddasph - rdddsph - rdddsph + rddca -
  & rbd(ixnotot) = -rbdad + rbdbs
else
  rv(ixnotot) = rvai + rvwbdf
  & gseifiph is negative
  lod(ixnotot) = -lodad - lodds
  & ldd(ixnotot) = -lddadmh + lddasph + ldddsph + ldddsph + lddca +
  & lbd(ixnotot) = lbdad - lbdbs
  rod(ixnotot) = rodad + rodds
  rdd(ixnotot) = rddasmh - rddasph - rdddsph - rdddsph + rddca -
  & rbd(ixnotot) = -rbdad + rbdbs
  rv(ixnotot) = rvai + rvwbdf - rvwbad
end if
return
end
subroutine tridag (a, b, c, r, u, n)
** SUBROUTINE:
** tridag.for - solves an equation A.u = r where A is a tridiagonal
** matrix and u and r are vectors (Numerical recipes
** in Fortran, 1992); this algorithm is used to solve
** the mass conservation equation for the water layer
** of the TOXSWA program
** author - P.I. Adriause
**
** DESCRIPTION:
** The equation to be solved is A.u = r;
** A is the tridiagonal matrix of which only the nonzero elements are
** stored as three vectors a(i:n), b(i:n) and c(i:n) representing
** respectively the band under the diagonal, the diagonal itself and
** the band above the diagonal. u and r are vectors of length n. So,
** the subroutine tridag solves for a vector u the tridiagonal
** linear set given by the equation above. a, b, c and r are input
** vectors and are not modified.
** Parameter: nmax is the maximum expected value of n
**
** HISTORY:
** 10 April 1996 - P.I. Adriause
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**

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```

C implicit none
C include 'param.for'      local variables
integer n, j, nmax
real a(n), b(n), c(n), r(n), u(n)
parameter (nmax = mxnxnotot)
real bet, gam(nmax)
C one vector of workspace, gam is needed
if (abs(b(1)).lt.1.0E-06) pause ' tridag: rewrite equations'
C If this happens then you should rewrite
C Your equations as a set of order n-1, with
C u2 trivially eliminated.
bet = b(1)
u(1) = r(1)/bet
do 10 j = 2, n
  gam(j) = c(j-1)/bet
  bet = b(j)-a(j)*gam(j)
  if (abs(bet).lt.1.0E-06) pause ' tridag failed'
  u(j) = (r(j)-a(j)*u(j-1))/bet
10 continue
do 20 j = n-1, 1, -1
  backsubstitution
  u(j) = u(j)-gam(j+1)*u(j+1)
20 continue
return
end
subroutine wimb
** SUBROUTINE:
** wimb.for - the mass balance is checked for the water layer of the
** TOXSWA program
** author - P.I. Adriause
**
** DESCRIPTION:
** All the incoming and outgoing mass terms are calculated for the
** water layer and it is checked that no mass is lost due to errors
** in the calculation process. The user receives an error message
** when more than 0.1% of the originally applied mass plus the mass
** incoming from the sediment is lost.
**
** HISTORY:
** 10 April 1996 - P.I. Adriause
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for'      common variables
C                            local variables
integer i, k, itto, helpi, imbout(21)
real sossj(mxnxnotot), cawlj(mxnxnotot), warst
  & totrainwb, rsainwb, totrsoutwb, totsvol,
  & rsvol, totstf, rsouteb, rsoutfb,
  & totwlist, totmwi,
  & cuinwb, cucutewb, cuvol, cutf, cucuteb, cuoutfb,
  & gumwi, qumwlpbc
  & wantoutpl, wantoutp2
C Logical
C---- Calculation of incoming substance from sediment and of
C substance penetrating in sediment
C (this is total *uwb*delx at time j and location i)
C This takes only place in ditch, not in front and end buffer
  & totrainwb = 0.0

```



```

open (unit = uomw, file = 'wimb.out', status = 'unknown')
write (uomw, 85) deltwl
85 format (' ', 'Output from subroutine wimb', /, ' ', 'F8.1', ' ', s')
&
&
& (timestep length deltwl = ' ', F8.1, ' ', s')
end if
if (wantoutp) then
write (uomw, 95) its, its*deltwl/86400., qumwl, qumwiprc,
&
& totmwt,
& totmwl, abs(cuinwb), -cuoutwb, cuoutfb,
& cutf, -cuouteb, cuoutfb
95 format(/, /, 'timestep', ' ', ' ', ' ', 'F8.4', ' ', 'd'), /, /,
&
& ' mass balance for w1 subsystem', /, /,
& ' qumwl = ', E10.3, ' g for entire ditch', /, /,
& ' ( ', E12.5, '% of initial mass + incoming mass wb)', /, /,
& ' positive mb-terms: ', E10.3, ' g, totmwl = ', E10.3, ' g', /, /,
& ' totmwt = ', E10.3, ' g, cuoutwb = ', E10.3, ' g', /, /,
& ' cuinwb = ', E10.3, ' g, cuoutfb = ', E10.3, ' g', /, /,
& ' 24X, ', 24X, ' ', cutf = ', E10.3, ' g', /, /,
& ' 24X, ', 24X, ' ', cuouteb = ', E10.3, ' g', /, /,
& ' 24X, ', 24X, ' ', cuoutfb = ', E10.3, ' g', /, /,
&
& end if
if (its.eq.ntswl) then
if (its.eq.ntswl) then
close (unit = uomw)
end if
end if
return
end
subroutine wimbnode
**
** SUBROUTINE:
** wimbnode.for - the mass balance is checked for a selected segment
** of the water layer of the TOXSWA program
** author
** - P.I. Adriaense
**
** DESCRIPTION:
** All the incoming and outgoing mass terms are calculated for one
** segment of the water layer and it is checked that no mass is lost
** due to errors in the calculation process. The user receives an
** error message when more than 0.1% of the originally applied mass
** plus the mass incoming from the sediment subsystem located under
** the segment concerned is lost.
** This subroutine was made to enable output of the mass balance for
** selected nodes of the water layer (and corresponding sediment
** subsystems).
**
** HISTORY:
** 10 April 1996 - P.I. Adriaense
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
** C include 'common.for' common variables
** C integer k, l, l1, l2 local variables
** C real sossj(mxknxnotot), cawlj(mxknxnotot), warst
** C real rsinfn, rsinfn, rsouten, rsoutfn, rsinwbn, rsoutwbn,
** C & ravoia, rstfn,
** C & mwlstn, mwlstn,
** C & cuinfn(mxknxnotot), cuinfn(mxknxnotot),
** C & cuoutfn(mxknxnotot), cuoutfn(mxknxnotot),
** C & cuinwbn(mxknxnotot), cuoutwbn(mxknxnotot),
** C & cuvoia(mxknxnotot), cutfa(mxknxnotot),
** C & qumw, qumwpr, bx
** C character filename*12
do 110 iknotot = nxnoEb-1, nxnofb+nxmodit

```

```

filename = 'wimb###.out'
do 100 k = 1, nwsby
if (iknotot.eq.iwbsy(k)) then
if (op.wimbnodeur.eq.1) then
uomw = 49+k
write (filename(5:7), '(I3.3)') iknotot
open (unit = uomw, file = filename, status = 'unknown')
end if
C---- Calculation of substance flowing into and out of selected segment
C (this is JVA at time j at interfaces iknotot-1/2 and iknotot+1/2)
C
C first calculation of concentration at susp.
C solids
if (iknotot.ne.1) then
sossj(iknotot-1) = raoms*kdoms(iknotot-1)*coobkoms*
& sossj(iknotot-1)/coobkoms**
& exfrss)
&
&
&
& sossj(iknotot) = raoms*kdoms(iknotot)*coobkoms*
& ((cowlj(iknotot)/coobkoms)**
& exfrss)
&
& if (iknotot.ne.nxnotot) then
& sossj(iknotot+1) = raoms*kdoms(iknotot+1)*coobkoms*
& ((cowlj(iknotot+1)/coobkoms)**
& exfrss)
&
& end if
if (qvoijmh.ge.zero) then
if (iknotot.eq.1) then
rsinfn = 0.0
else
rsinfn = qvoijmh*
& (cowlj(iknotot-1)+coss*sossj(iknotot-1)) -
& warjmh(iknotot)*kdsowljmh(iknotot)*{
& cowlj(iknotot)+coss*sossj(iknotot) -
& (cowlj(iknotot-1)+coss*sossj(iknotot-1))//
& (0.5*(delx(iknotot)+delx(iknotot-1)))
&
& end if
else
now discharge at interface iknotot-1/2
negative, so there is a front buffer
if (iknotot.eq.1) then
rsinfn = qvoijmh*
& (cowlj(iknotot)+coss*sossj(iknotot))
&
& else
rsinfn = qvoijmh*
& (cowlj(iknotot)+coss*sossj(iknotot)) -
& warjmh(iknotot)*kdsowljmh(iknotot)*{
& cowlj(iknotot)+coss*sossj(iknotot) -
& (cowlj(iknotot-1)+coss*sossj(iknotot-1))//
& (0.5*(delx(iknotot)+delx(iknotot-1)))
&
& end if
end if
rsoutfn = amin1(0.0, rsinfn)
rsinfn = amax1(0.0, rsinfn)
if (qvoijph.ge.zero) then
if (iknotot.eq.nxnotot) then
rsouten = qvoijph*
& (cowlj(iknotot)+coss*sossj(iknotot))
&
& else
rsouten = qvoijph*
& (cowlj(iknotot)+coss*sossj(iknotot)) -
& warjph(iknotot)*kdsowljph(iknotot)*{
& cowlj(iknotot)+coss*sossj(iknotot+1) -
& (cowlj(iknotot)+coss*sossj(iknotot))//
& (0.5*(delx(iknotot+1)+delx(iknotot)))
&
& end if
else
now discharge at interface iknotot+1/2
negative
if (iknotot.eq.nxnotot) then
rsouten = 0.0
else
rsouten = qvoijph*
& (cowlj(iknotot+1)+coss*sossj(iknotot+1)) -
& warjph(iknotot)*kdsowljph(iknotot)*{

```

```

C----- Actual total substance in selected segment at end of timestep j
&
&
&
mwln = cawljpl(ixnotot)*warjpl(ixnotot)*delx(ixnotot)

C----- Integration with respect to time
cuinfn(ixnotot) = cuinfn(ixnotot) + deltwl*rsinfn
cuinen(ixnotot) = cuinen(ixnotot) + deltwl*rsinen
cuouten(ixnotot) = cuouten(ixnotot) + deltwl*rsouten
cuoutfn(ixnotot) = cuoutfn(ixnotot) + deltwl*rsoutfn
cuinwbn(ixnotot) = cuinwbn(ixnotot) + deltwl*rsinwbn
cuoutwbn(ixnotot) = cuoutwbn(ixnotot) + deltwl*rsoutwbn
cuvaln(ixnotot) = cuvaln(ixnotot) + deltwl*rsvaln
cutfn(ixnotot) = cutfn(ixnotot) + deltwl*rstfn

C----- Mass balance checked
C (Quantity (g) to check mass balance selected segment)
&
&
&
qumw = mwlstn + cuinfn(ixnotot) - cuinen(ixnotot) -
cuoutwbn(ixnotot) - cuoutfn(ixnotot) +
cuvaln(ixnotot) + cutfn(ixnotot) -
mwln
&
&
if ((mwlstn+cuinfn(ixnotot)-cuinen(ixnotot)-
cuoutwbn(ixnotot)-cuoutfn(ixnotot)-
cuvaln(ixnotot)-cutfn(ixnotot))*100.
qumper = (qumw/(mwlstn+cuinfn(ixnotot)-
cuinen(ixnotot)-cuinwbn(ixnotot)))*100.
&
&
qumper = 0.01
end if

C----- Output
C output in message.out if error
if (its.eq.1) then
itflo = 0
end if
if (abs(qumper).gt.0.1 .and. itflo.lt.20) then
write (uwer, 63) qumper, ixnotot, its,
its*deltwl/86400.,
qumw, (mwlstn+cuinfn(ixnotot)-
cuinen(ixnotot)-cuinwbn(ixnotot))
&
&
&
63 format (//, ' Warning:', E12.5, ' % of application plus incoming',
, mass, ' /, is missing in the',
, mass balance of node', I8, ' of the water layer', /,
, at timestep', I8, ' (', F6.2, ' day)', /,
, qumw =', E10.3, ' g and', /,
, (mwlstn+cuinfn-cuinen-cuinwbn) =', E10.3, ' g', /,
, See output in wimb##.out', /)
itflo = itflo+1
end if

C
output in file wimb##.out
if (op.wimbo.denr.eq.1) then
if (its.eq.1) then
write (uoma, 60) deltwl
, per running meter ditch', /,
, (timestep length deltwl =', F8.1, ' s')
end if
if (wantoutp) then
hx = delx(ixnotot)
write (uoma, 69) its, its*deltwl/86400.,
ixnotot, qumw/hx,
qumper, mwlstn/hx, mwln/hx,
abs(cuinfn(ixnotot))/hx,
-cuouten(ixnotot)/hx,
abs(cuinen(ixnotot))/hx,
cuoutfn(ixnotot)/hx,
abs(cuinwbn(ixnotot))/hx,
-cuoutwbn(ixnotot)/hx,
cuvaln(ixnotot)/hx,
cutfn(ixnotot)/hx
69 format (//, ' timestep', I8, ' (', F8.4, ' d)', /,
, mass balance for wl subsystem per m', at node', I8, /,

```



```

C      common variables
C      local variables
C      integer icountersofd
C      logical t1, t2, t12, t3, good
C----- The condition that the calculated dispersion flux should not
C cancel out or exceed the advection flux will be implemented at
C time j (and not j+1/2 as clb j+1 is not yet known)
C
C      for the flux across the interface
C      kznotot-1/2
C
C      sofdfactorrmh(kznotot) = 1
C      first determine whether situation 2 plus
C      condition, Eq. (4.23) occurs (see Fig. 11
C      of SC-DLO report 90)
C
C      if ((kznotot.gt.1) .and. good) then
C      t1 = qselj.gt.zero
C      t2 = ((cobjj{kznotot}-cobjj{kznotot-1})/
C      & {zcd{kznotot}-zcd{kznotot-1}}).gt.zero
C      & t3 = (cobjjkmh{kznotot} - cobjj{kznotot-1})/
C      & ldis*(cobjj{kznotot} - cobjj{kznotot-1}).lt.zero
C      & {zcd{kznotot}-zcd{kznotot-1}}.lt.zero
C      t12 = t1 .and. t2
C      good = t12 .and. t3
C      end if
C
C      if ((kznotot.gt.1) .and. good) then
C      now suppose that dispersion flux and
C      advection flux have exactly the same
C      magnitude but they are directed in opposite
C      directions (so their combined effect is
C      zero; Elb J+1/2, k-1/2 = 0 and
C      l.gj+1/2, clb j+1/2, k-1/2 = 0; this last
C      term will be set to zero with the aid of a
C      factor sofdfactorrmh(kznotot) in lod, ldd,
C      rod and rdd
C      kselbjkmh(kznotot) = zero
C      sofdfactorrmh(kznotot) = zero
C      end if
C
C      determine whether situation 4 plus
C      condition, Eq. (4.24) occurs (see Fig. 11
C      of SC-DLO report 90)
C
C      if ((kznotot.gt.1) .and. good) then
C      t1 = qselj.lt.zero
C      t2 = ((cobjj{kznotot}-cobjj{kznotot-1})/
C      & {zcd{kznotot}-zcd{kznotot-1}}).lt.zero
C      & t3 = (cobjjkmh{kznotot} +
C      & ldis*(cobjj{kznotot}-cobjj{kznotot-1})/
C      & {zcd{kznotot}-zcd{kznotot-1}}).lt.zero
C      t12 = t1 .and. t2
C      good = t12 .and. t3
C      end if
C
C      if ((kznotot.gt.1) .and. good) then
C      Elb J+1/2, k-1/2 = 0 and
C      l.gj+1/2, clb j+1/2, k-1/2 = 0; this last
C      term will be set to zero with the aid of a
C      factor sofdfactorrmh(kznotot) in lod, ldd,
C      rod and rdd
C      kselbjkmh(kznotot) = zero
C      sofdfactorrmh(kznotot) = zero
C      end if
C
C      determine whether situation 4 plus
C      condition, Eq. (4.24) occurs
C
C      if (sofdfactorrmh(kznotot).le.zero .or.
C      icountersofd = icountersofd + 1
C      if (icountersofd.le.50) then
C      write (uerr, 20) icountersofd, its, ixnotot, kznotot,
C      20 format(//, 'Warning:', //
C      & ', Phenomenon of dispersion flux cancel out/exceeding', //
C      & ', advection flux would occur in sediment. (this has', //
C      & ', been avoided with aid sofdactors),', //
C      & ', Number of occurrence of this phenomenon:', i6, //,
C      & ', Timestep:', i6, ', under water layer node:', i6, //,
C      & ', sediment node:', i6, //, ' sofdfactorrmh:', f6.2, //
C      & ', sofdfactorph:', f6.2)
C      end if
C      end if
C
C      return
C
C      end
C
C      subroutine wboem
C
C      ** SUBROUTINE:
C      ** wboem.for - the elements of the tridiagonal matrices for the
C      ** sediment layer of the TONSWA program are
C      ** defined here
C      ** author - P.I. Adriaanse
C      **
C      ** DESCRIPTION:
C      ** The elements composing the two tridiagonal matrices for the
C      ** sediment layer are defined here; these elements are used in the
C      ** subroutines wboelb, wboelb and wboecr.
C      ** HISTORY:

```

```

** 10 April 1996 - P.I. Adriaanse
** version 1.0
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for'
C
C common variables
C local variables
C
C----- Elements of the core and buffer matrices will be composed here
C
C the terms blod.., biddadm, brod.. and
C brddadph will not be defined for kznotot=1,
C as well as biddsdafh, blbd...
C brddsdafh and brbd... for kznotot=nznotot.
C because they do not appear in the two
C triangular matrices.
C also biddsdafh and brddsdafh will not be
C defined for kznotot=1, as well as
C biddsdafh, blbdca, blddf, brddadph,
C brddca and brddf for kznotot=nznotot,
C because these terms are modified for these
C nodes
C
C if (kznotot.gt.1) then
C   blodad = ((leplot*gseljph)/delz(kznotot)) * (1.-thetawb) *
C   & (1.-nwph(kznotot)) * (deltwb/pe(kznotot)) *
C   & sofdacformh(kznotot)
C   blodsdaf = (pemb(kznotot)*porph(kznotot))*(kdsbc1jkmh(kznotot))*
C   & kdfwbh(kznotot)/(1.-thetawb)*deltwb /
C   & (delz(kznotot)*0.5*(delz(kznotot)+delz(kznotot-1)))*
C   & pe(kznotot)
C end if
C
C if (kznotot.lt.nznotot) then
C   biddadph = ((leplot*gseljph)/delz(kznotot)) * (1.-thetawb) *
C   & (1.-nwph(kznotot)) * (deltwb/pe(kznotot)) *
C   & sofdacformh(kznotot)
C   brddsdafh = (pemb(kznotot)*porph(kznotot))*(kdsbc1jkmh(kznotot))*
C   & kdfwbh(kznotot)/(1.-thetawb)*deltwb /
C   & (delz(kznotot)*0.5*(delz(kznotot)+delz(kznotot-1)))*
C   & pe(kznotot)
C end if
C
C if (kznotot.lt.nznotot) then
C   brodca = (combj(kznotot)/coobkcomb)**kdfwb(kznotot)*
C   & (combj(kznotot) + bdbw(kznotot))*kdfwb(kznotot) *
C   & (combj(kznotot)/coobkcomb)**(exfrwb-1.)) *
C   & krfwb*thetawb*deltwb
C end if
C
C if (kznotot.lt.nznotot) then
C   brodad = ((leplot*gseljph)/delz(kznotot)) * (1.-thetawb) *
C   & (1.-nwph(kznotot)) * (deltwb/pe(kznotot)) *
C   & sofdacformh(kznotot)
C   brddsdafh = (pemb(kznotot)*porph(kznotot))*(kdsbc1jkmh(kznotot))*
C   & kdfwbh(kznotot)/(1.-thetawb)*deltwb /
C   & (delz(kznotot)*0.5*(delz(kznotot)+delz(kznotot-1)))*
C   & pe(kznotot)
C end if
C
C if (kznotot.lt.nznotot) then
C   brodca = (combj(kznotot)/coobkcomb)**kdfwb(kznotot)*
C   & (combj(kznotot) + bdbw(kznotot))*kdfwb(kznotot) *
C   & (combj(kznotot)/coobkcomb)**(exfrwb-1.)) *
C   & krfwb*thetawb*deltwb
C end if
C
C return
C
C subroutine wbsoub
C
C ** SUBROUTINE.
C ** wbsoub.for - the blod, bidd etc terms of the matrices for the
C ** upper boundary of the sediment layer of the FORSWA
C ** program. It is composed here
C ** author
C ** - P.I. Adriaanse
C
C ** DESCRIPTION:
C ** The first row (kznotot=1) for the lefthand and righthand
C ** triangular matrices will be composed for this timesep.
C ** (Only the nonzero components of the left-hand and right-hand
C ** triangular matrices will be stored, they form two times three
C ** vectors; notice that blod(i), bidd(kznotot), brod(i) and
C ** brbd(kznotot) are undefined and not referenced in the
C ** subroutines).
C
C ** HISTORY:
C ** 10 April 1996 - P.I. Adriaanse
C ** version 1.0
C
C ** COPYRIGHT:
C ** DLO Winand Staring Centre for Integrated Land, Soil and Water
C ** Research (SC-DLO), 1996.
C
C include 'common.for'
C common variables
C local variables

```



```

C
now node kznotot with upward seepage
blod(kznotot) = -blodad - bloddsdf
bldd(kznotot) = -bloddadh + bldddsdfmh +
                bldda + blddtf
brd(kznotot) = brodad + broddsdf
brdd(kznotot) = brddadmh - brdddsdfmh +
                brdda - brddtf
brv(kznotot) = -colotj*(leplot*qsdfjph*thetamb*
                deltwb) / (deiz(kznotot)*pe(kznotot)) -
                colotjpi*(leplot*qsdfjph*
                (1.-thetamb)*deltwb) / (deiz(kznotot)*
                pe(kznotot))
                end if
                end if
return
end

subroutine wbsocr
** SUBROUTINE:
** wbsocr.for - the core parts of the matrices, consisting of the
** numerical equations to solve the mass conservation
** equation for the sediment layer of the TOXSWA
** program are composed here
** author - P.I. Adriaanse
** DESCRIPTION:
** The core parts of the matrices will be composed for this timestep.
** (Only the nonzero components of the left-hand and right-hand
** tridiagonal matrices will be stored, they form two times three
** vectors; notice that blod(j), blbd(kznotot), brodd(i) and
** brbd(mznotot) are undefined and not referenced in the
** subroutines.)
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
C
include 'common.for'
local variables
blod(kznotot) = -blodad - bloddsdf
bldd(kznotot) = bldadph + bldddsdfph - bloddadmh + bldddsdfmh +
                bldda + blddtf
blbd(kznotot) = blbdad - blbdsddf
brd(kznotot) = brodad + broddsdf
brdd(kznotot) = brddadmh - brdddsdfph + brddadmh - brdddsdfmh +
                brdda - brddtf
brbd(kznotot) = -brbdad + brbdsddf
brv(kznotot) = zero
return
end

subroutine bridag (a, b, c, r, u, n)
** SUBROUTINE:
** bridag.for - version of tridag for sediment (only local parameter
** nmax has been changed)

```

```

** tridag.for - solves an equation A.u = r where A is a tri-diagonal
** matrix and u and r are vectors (Numerical recipes
** in Fortran, 1992); this algorithm is used to solve
** the mass conservation equation for the water layer
** of the TOXSWA program
** author - P.I. Adriaanse
** DESCRIPTION:
** The equation to be solved is A.u = r;
** A is the tri-diagonal matrix of which only the nonzero elements are
** stored as three vectors a(i:n), b(i:n) and c(i:n) representing
** respectively the band under the diagonal, the diagonal itself and
** the band above the diagonal. u and r are vectors of length n. So,
** the subroutine bridag solves for a vector u the tri-diagonal
** linear set given by the equation above. a, b, c and r are input
** vectors and are not modified.
** Parameter: nmax is the maximum expected value of n
** HISTORY:
** 10 April 1996 - P.I. Adriaanse
** version 1.0
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
** implicit none
** include 'param.for'
** local variables
** integer n, j, nmax
** real a(n), b(n), c(n), r(n), u(n)
** parameter (nmax = mznznotot)
** real bet, gam(nmax)
**
** one vector of workspace, gam is needed
** if (abs(b(1)).lt.1.0E-06) pause ', tridag: rewrite equations'
** if this happens then you should rewrite
** your equations as a set of order n-1, with
** u2 trivially eliminated.
bet = b(1)
u(1) = r(1)/bet
do 10 j = 2, n
  gam(j) = c(j-1)/bet
  bet = b(j)-a(j)*gam(j)
  if (abs(bet).lt.1.0E-06) pause ', tridag failed'
  if this happens then you should rewrite
  your equations as a set of order n-1, with
  u2 trivially eliminated.
  u(j) = (r(j)-a(j)*u(j-1))/bet
10 continue
do 20 j = n-1, 1, -1
  backsubstitution
  u(j) = u(j)-gam(j+1)*u(j+1)
20 continue
return
end
subroutine wbmb
** SUBROUTINE:
** wbmb.for - the mass balance is checked for the sediment layer of
** the TOXSWA program
** author - P.I. Adriaanse
** DESCRIPTION:
** All the incoming and outgoing mass terms are calculated for the
** sediment layer and it is checked that no mass is lost due to
** errors in the calculation process. The user receives an error
** message when more than 0.1% of the originally applied mass plus
** the mass incoming from the water layer is lost. Calculations are
** done per sediment subsystem first and afterwards combined to form
** the entire sediment layer.

```

```

** HISTORY:
** 10 April 1996 - P.I. Adriano
** version 1.0
**
** COPYRIGHT:
** DLO Winand Staring Centre for Integrated Land, Soil and Water
** Research (SC-DLO), 1996.
**
C include 'common.for'      common variables
C                          local variables
integer k, kk, itlto
real cwbj(mxznxnotot)
real rsnwl, rsum, rsnoutw, rseper, totrstf,
& totmwb, totmwb, mwbtop, delstot,
& cuinwl(mxknnotot), cus(mxknnotot), cuoutwl(mxknnotot),
& cuspr(mxknnotot), cutf(mxknnotot), qumwb, qumwbr,
& allcumbst, allcumb, allmwbtop,
& allcumb, allcumbpr,
& allcumb, allcumb, allcumb, allcumb, allcumb, allcumb,
& character finame*12

C----- Calculation of incoming substance from water layer and of
C substance lost to water layer
C (this is P*Ulb at time j and k=1/2)
if (qseifj.ge.zero) then
  rsnwl = leptot*qseifj*cowlj(ixnotot) - pemh(1)*pormh(1)*
& kdfwmbh(1)*((cwbj(1)-cowlj(ixnotot))/(.5*delz(1)))
& else
  rsnwl = leptot*qseifj*cowlj(ixnotot) - pemh(1)*pormh(1)*
& kdfwmbh(1)*((cwbj(1)-cowlj(ixnotot))/(.5*delz(1)))
& end if
rsoutw = amin(0.0, rsnwl)
rsinwl = amax(0.0, rsnwl)

C----- Calculation of incoming substance by upward seepage
C (this is P*Ulb at time j and k=znnotot+1/2)
rsum = leptot * amin(0.0, qseifj) * colotj

C----- Calculation of substance percolated out of wb-subsystem
C (this is P*Ulb at time j and location k=znnotot)
rseper = amax(1,0.0)*leptot*cwbj(znnotot)

C----- Calculation of substance lost to water layer
C (this is P*Ulb at time j and k=1/2)
if (qseifj.ge.zero) then
  rsnoutw = leptot*qseifj*cowlj(ixnotot) - pemh(1)*pormh(1)*
& kdfwmbh(1)*((cwbj(1)-cowlj(ixnotot))/(.5*delz(1)))
& else
  rsnoutw = leptot*qseifj*cowlj(1) - pemh(1)*pormh(1)*
& kdfwmbh(1)*((cwbj(1)-cowlj(1))/(.5*delz(1)))
& end if
rsoutw = amin(0.0, rsnoutw)

C----- Calculation of transformation of substance
do 10 k = 1, nznoub
  cwbj(k) = por(k)*cwbj(k) + bsdwb(k)*kdfwrb(k)*coobkombw*
& totrstf + totrstf - krtfwb*cwbj(k)*pe(k)*delz(k)
10 continue

C----- Initial conditions for mass balance wb-system
totmwbst = 0.0
do 20 k = 1, nznnotot
  totmwbst = totmwbst + castwb(k)*pe(k)*delz(k)
20 continue

if (its.eq.1) then
  cuinwl(ixnotot) = 0.0
  cus(ixnotot) = 0.0
  cuspr(ixnotot) = 0.0
  cuoutwl(ixnotot) = 0.0
  cutf(ixnotot) = 0.0
end if

C----- Actual lineic mass of substance in selected top layer
mwbtop = 0.0
do 30 k = 1, ktop
  mwbtop = mwbtop + cawbjpl(k)*pe(k)*delz(k)
30 continue

C----- Actual total lineic substance mass in wb-system
at end of timestep j
totmwb = 0.0
do 40 k = 1, nznnotot
  totmwb = totmwb + cawbjpl(k)*pe(k)*delz(k)
40 continue

C----- Integration with respect to time
cuinwl(ixnotot) = cuinwl(ixnotot) + deltwb*rsinwl
cus(ixnotot) = cus(ixnotot) + deltwb*rsum
cuspr(ixnotot) = cuspr(ixnotot) + deltwb*rseper
cuoutwl(ixnotot) = cuoutwl(ixnotot) + deltwb*rsoutw
cutf(ixnotot) = cutf(ixnotot) + deltwb*totrstf

C----- Mass balance checked
C (lineic quantity (g/m') to check mass balance wb)
qumwb = totmwb + cuinwl(ixnotot) - cus(ixnotot) -
& cuspr(ixnotot) + cuoutwl(ixnotot) + cutf(ixnotot) - totmwb

C----- Mass balance for all wb-subsystems totalized of ditch
C-----
C----- Initial conditions
if (ixnotot.eq.nozdit) then
  alltotmwbst = 0.0
  allmwbtop = 0.0
  allcumbst = 0.0
  allcumb = 0.0
  allcumbpr = 0.0
  allcumb, allcumb, allcumb = 0.0
end if

C
C cu..terms accumulated over time so
allcu..terms = 0.0
if (ixnotot.eq.nozdit) then
  allcuinwl = 0.0
  allcus = 0.0
  allcuspr = 0.0
  allcuoutw = 0.0
  allcutf = 0.0
end if

C----- Integration in direction of flow in ditch
alltotmwbst = alltotmwbst + totmwbst*delx(ixnotot)
allmwbtop = allmwbtop + mwbtop*delx(ixnotot)
alltotmwb = alltotmwb + totmwb*delx(ixnotot)
allcuinwl = allcuinwl + cuinwl(ixnotot)*delx(ixnotot)
allcus = allcus + cus(ixnotot)*delx(ixnotot)
allcuspr = allcuspr + cuspr(ixnotot)*delx(ixnotot)
allcuoutw = allcuoutw + cuoutw(ixnotot)*delx(ixnotot)
allcutf = allcutf + cutf(ixnotot)*delx(ixnotot)

C----- Check of mass balance for all wb-subsystems totalized of ditch
if (ixnotot.eq.nznofb-xnnotot) then
  allqumwb = alltotmwbst + allcuinwl - allcus - allcuspr +
& allcuoutw + allcutf - alltotmwb
end if

C----- Output
C----- Output for selected wb-subsystems
C----- Output

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