

# Study of Sea level changes with Leckie method of the Abderaz Formation (Middle Turonian-Early Campanian) at type section based on foraminifera

Meysam Shafiee Ardestani<sup>1\*</sup>, Abbas Abbasi<sup>1</sup>

Department of Geology, Faculty of Science, Ferdowsi University of Mashhad, Mashhad, Iran

\* Corresponding author: m\_shafiee@khayam.ut.ac.ir

## Abstract

In this research the Abderaz Formation at its type section with an age of Turonian-early Campanian and a thickness of 300 m containing light grey shale and marls was investigated. Statistical Studise of the morphotype groups of planktonic foraminifera shows that the majority of them are of shallow water forms (SWF) and deep water forms (DWF) and planktonic to benthic ratio is high indicating specific condition of oligotrophy and sedimentation in a relatively deeper marine condition

**Keywords:** Planktonic foraminifera, Morphotype, Abderaz Formation, Sea level changes

## Introduction

The Study on the morphotypes and planktonic to benthic ratio was the major aim of the research. This study was intended to explore the marine sedimentation of Abderaz Formation in (outer neritic-upper bathyal) restrict. Then 84 SEM images have been obtained and demonstrated in frame of 2 plate.

## Material and Method

The section studied is located about 1 km to the Muzduran, north eastern Mashhad (a city of Iran), Kopet Dagh basin. At this locality (E: 60, °33', 00'', N: 36°, 10', 40'')(Fig1). Type section of Abderaz Formation has 300m thickness. At the typical gap such as all regions under the surface sub-contact of Abderaz Formation are un-correlated with Aitamir Formation. But its upper layer with Abtalkh Formation is in continuous correlation. The upper layer has elected as chalk limestone upper border. A total of 130 samples were collected from the section, but Only 102 samples were included in study, 7 samples due to the existence of salvation effects and 21 samples was obtained from reworking damages that were excluded from the study. Which were soaked in water with diluted hydrogen peroxide, washed through 63µm, 150µm and 250µm sieves, and dried until clean foraminiferal residues were recovered. About 200-300 individuals were picked up for each sample in two size fractions (63-150µm and >150µm) and mounted on dark cardboard slides for identification. These two size fractions were analyzed in order to obtain statistically significant representatives of the small and large groups. Species identifications are based on (Caron, 1985, Robaszynski and Caron, 1983-1984, 1995 Loeblich and Tappan, 1950, 1988 and Tappan 1940, 1943).

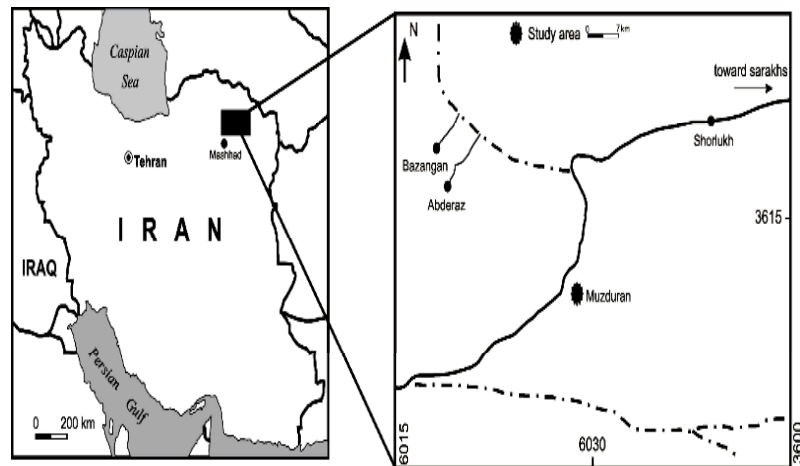


Fig1. The geographical map and the ways to the region of the study.

## Discussion

Groups of planktonic morphotypes are distinguished by depth of living (Leckie 1987) (Fig2). Those are consisting of:

1- **Epicontinental Sea Forms= ESF:** *Heterohelix* and *Guembelitra* are related to faunas of shallow epicontinental sea or the border sea (Leckie 1987, PremoliSilva & Sliter 1999, Keller, 2002).

2- **Shallow Water Forms= SWF**

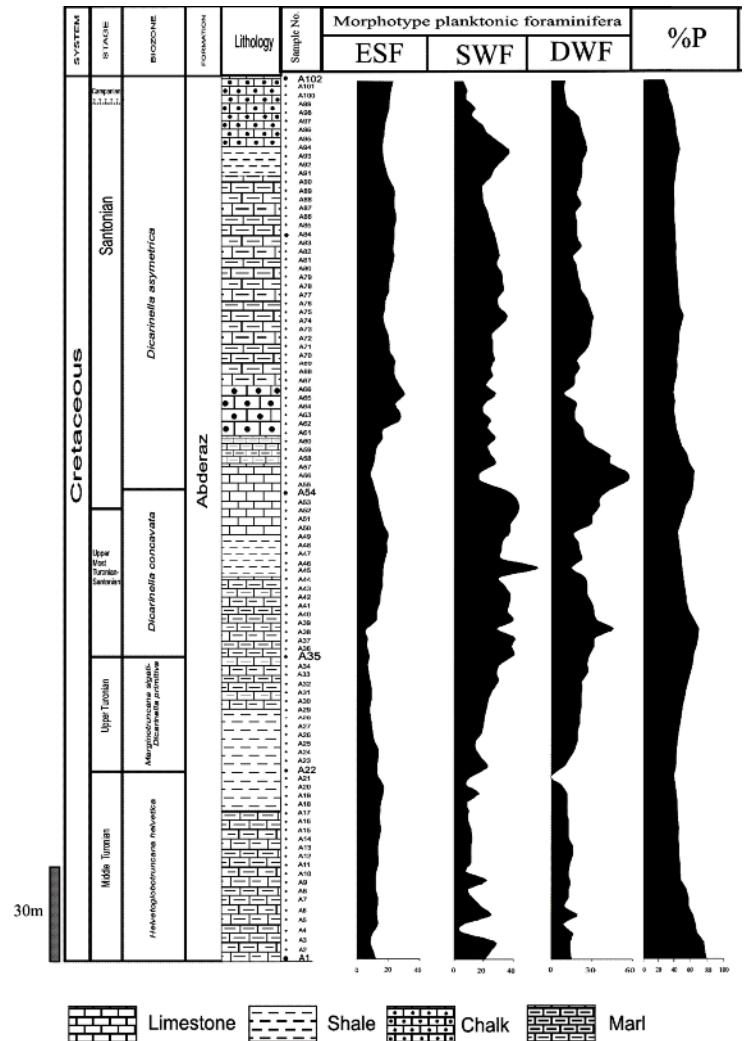
*Hedbergella delrioensis* and *Whiteinella* are related to this fauna (Leckie 1987, PremoliSilva & Sliter 1999).

3- **Deep Water Forms= DWF**

These faunas were counted like keeled shapes

there were 300 samples in the size of 120 mesh completely by chance, from every samples were counted that the result of this count shows at the first of successions and the time middle Turonian **ESF** was conquering and the amount of the **DWF** and **%P** was less in the area that this paragraph. in the late Turonian the group of **DWF** was increasing in the area that it indicated the proportional increasing of depth in the area and by this time portici structure has been larger and in umbilical structure is born in this unicellular, and in Coniacian time has decreased the amount of **DWF** in the area again and the members of **ESF** increased with the less **%P** in the area again and during Coniacian to Santonian the members of **DWF** with **%P** increase in the area for another time and in Santonian time, sea water shows a vacillation mood in the above-mentioned section . Also the planktonic to benthic ratio which explains that at the deposits 400 meter at total part of in this Formation. This study was intended to explore the marine sedimentation of **Abderaz** Formation in (outer neritic-upper bathyal) restrict and the provided curves from morphotypes changes are in full agreement with the curves of the sea level changes and planktonic foraminifer to epifauna benthic ratio(Fig2).

Fig2-Comparison of planktonic morphotype curves with **%P** )**ESF** = Epicontinental Sea Forms, **SWF** = Shallow Water Forms, **DWF** = Deep Water Forms, **%P**=percentage of planktonic foraminifera)



## Result

Groups of planktonic morphotypes are distinguished by depth of living that is consisting of:

- 1- Shallow area faunas= **ESF** = Epicontinental Sea Forms
- 2- Middle water faunas = **SWF** = Shallow Water Forms
- 3- Deep water faunas (lower than 100) = **DWF** = Deep Water Forms

In the time of middle Turonian simultaneous with subtraction of the percent of **DWF** that indicates the dwindling of proportional in mentioned section. **%P** increases but in the late Turonian that the percent of **DWF** increases that it would indicated the proportional of depth increasing in area and the structured shapes in vicinity has increased and the structured shaped (tegilla) recently has born and in Coniacian time the **DWF** diminished again and **%P** increase and in Coniacian -Santonian boundary by increasing the shapes of **DWF** and **%P** became the most in this time. That this affair it is because of the advent of *Globotrancana* and increasing the

number of them in Santonian time but in the late Santonian and the early Campanian by diminishing the percent of **DWF** and increasing **ESF**, the lip shapes became more in area. This study was intended to explore the marine sedimentation of Abderaz Formation in (outer neritic-upper bathyal) restrict and the provided curves from morphotypes changes are in full agreement with the curves of the sea level changes and planktonic foraminifer to epifauna benthic ratio.

### Acknowledgments

This study was supported by Ferdowsi University of Sciences Research Council, National Museum of Mashhad University, Mashhad, Iran and Razi Institute of Metalogy, Division of SEM, Karaj, Iran.

### Reference:

Caron, M. (1985). Cretaceous planktic foraminifera. In: Bolli, H.M., Saunders, J.B., Perch Nielsen, K. (Eds), *Plankton stratigraphy*. Cambridge University Press. pp 17-86.

Caron, M., Homewood, P. 1983 Evolution of early planktic foraminifers. *Mar. Micropaleontology*. 7, 435-462. -Marcel & A. de Vernal, Elsevier, 843 pp.

Keller, G., Adatte, T., Stinnesbeck, W., Luciani, V., Karoui, N., Zaghbib-Turki, D., 2002 Tertiary mass extinction in planktic foraminifera. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* 178, 257-298.

Leckie, R.M., 1987. Paleocology of mid-Cretaceous planktonic foraminifera: A comparison of open ocean & Epicontinental Sea assemblages. *Micropaleontology*. 33, 164-176

Loeblich, A. R., JR., & Tappan, Helen 1950. Foraminifera from the type Kiowa Shale,

Lower Cretaceous of Kansas. *Kansas, Univ., Pal. Contr.*, no. 6 (Protozoa art. 3), pp. 1-15, pls. 1-2

Loeblich, A. R. Jr & Tappan, E. 1988. *Foraminiferal genera & their classification*, 970 pp. (Van Nostrand Reinhold Company, New York).

Premoli Silva, I., Sliter, W.V., 1999. Cretaceous paleoceanography: evidence from planktonic foraminiferal evolution. *Geology. Soc Am. Spec. Pap.*, vol. 332, pp. 301-328.

Robaszynski, F., Caron, M., 1979. Atlas de foraminifères planctoniques du Crétacé moyen (Mer Boreale et Tethys), *première partie. Cahiers de Micropaleontologie* 1 (1-185 pp.).

Robaszynski, F., & Caron, M., 1995, Foraminifères planctoniques du crétacé: *Bulletine Société Géologique de France*, t.166, p.681-698

Tappan, H., 1940, Foraminifera from Thengrayson Formation of northern Texas. *Journal of Paleontology*, v.17, p.93-126

Tappan, H., 1943, Foraminifera from the duck Creek Formation of Oklahoma & Texas. *Journal of Paleontology*, v.17, p.93-126

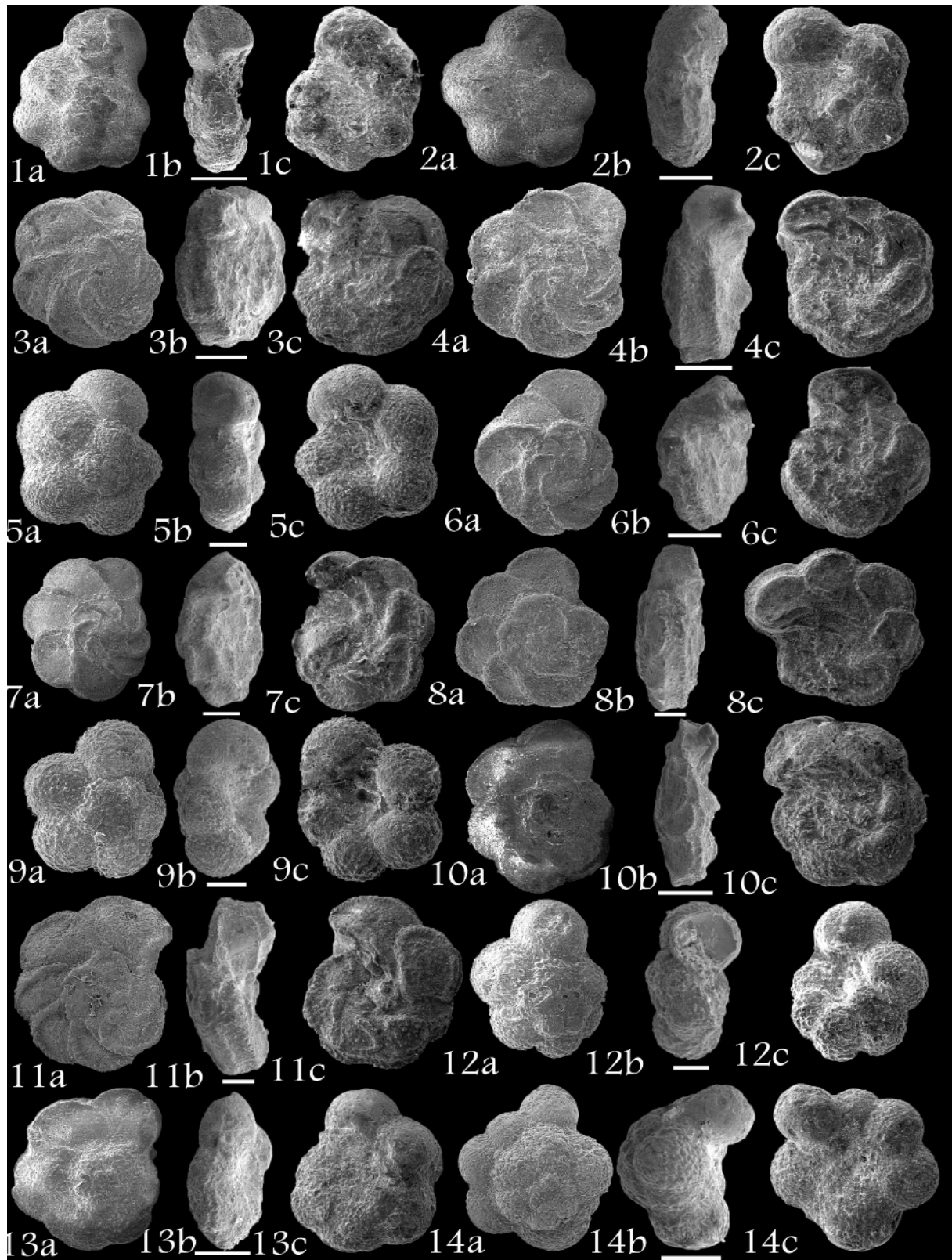
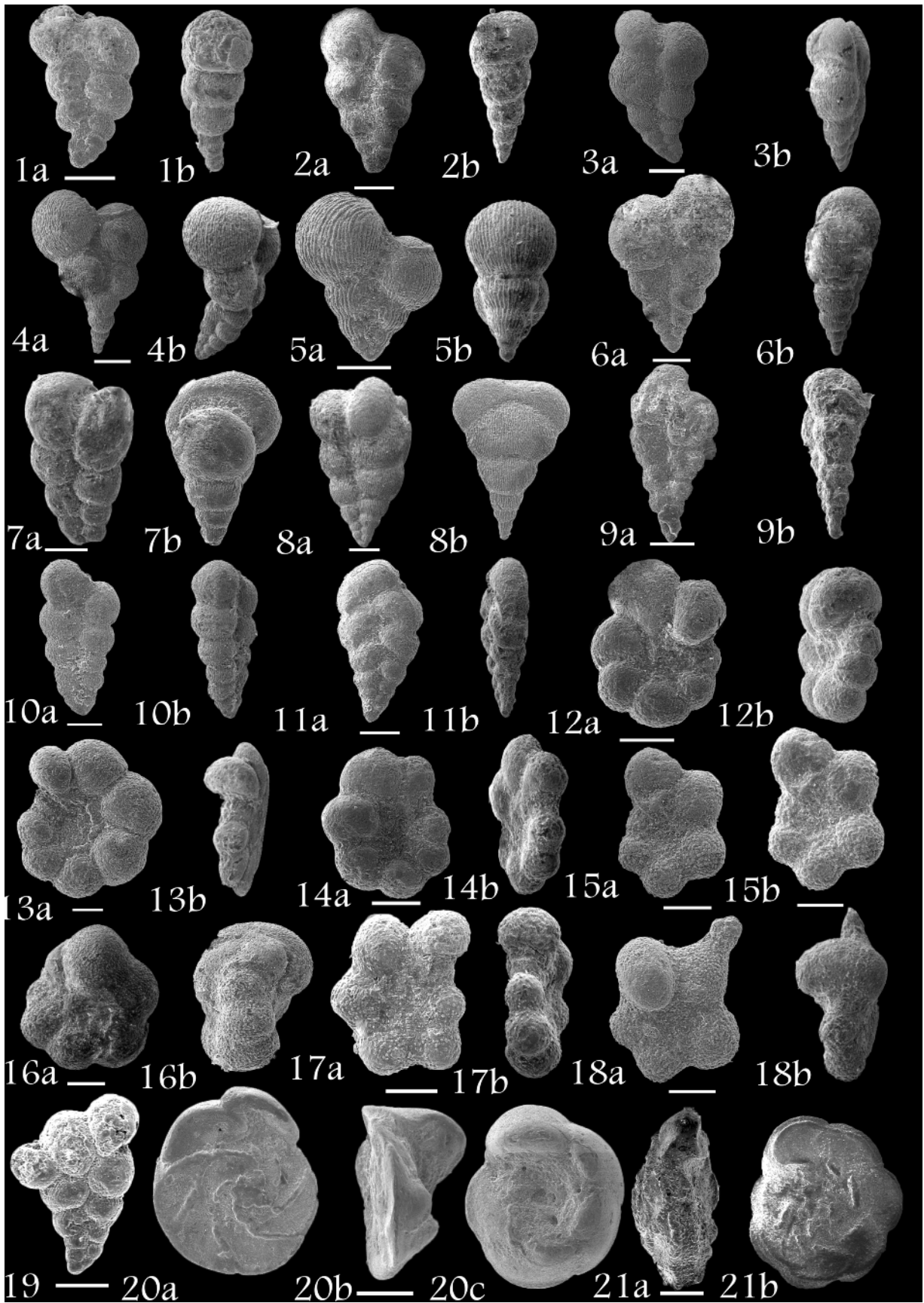


Plate 1: 1-*Whiteinella brittonensis*, Sample 17. 2-*Whiteinella archaeocretacea*, Sample 17. 3, 4-*Marginotruncana pseudolinneiana*, Samples 29, 42. 5-*Whiteinella archaeocretacea*, Sample 25. 6- *Marginotruncana coronata*, Sample 20. 7- *Marginotruncana pseudolinneiana*, Sample 31. 8- *Marginotruncana coronata*, Sample 40. 9- *Whiteinella aprica*, Sample 1. 10, 11- *Marginotruncana coronata*, Sample 47, 48. 12-*Whiteinella brittonensis*, Sample 12. 13-*Dicarinella cf hagni*, Sample 16. 14-*Whiteinella paradubia*, Sample 20. Scale bar represents 200µm except for Samples 8, 9, 11, 12 which represents 100µm.



**Plate2:** 1-6-*Heterohelix globulosa*, Samples 4, 15, 22, 25, 76, 95, 7, 8-*Pseudotextularia nutalli*, Samples 55, 96, 9-*Heterohelix moremani*, Sample 15, 10-*Heterohelix papula*, Sample 14, 11-*Laeviheterohelix pulchra*, Sample 54, 12-*Globigerinelloides prairiehillensis*, Sample 55, 13-*Globigerinelloides* sp, Sample 22, 14-*Globigerinelloides* sp, Sample 50, 15-*Globigerinelloides* sp, Sample 56, 16-*Globigerinelloides prairiehillensis*, Sample 56, 17-*Globigerinelloides* sp, Sample 56, 18-*Globigerinelloides* sp, Sample 50, 19- *Ventilabrella* cf *austiniana*, Sample 40, 20-*Globotruncanita elevata*, Sample 99, 21-*Globotruncana* sp, Sample 85. Scale bar represents 100µm except for Samples 20, 21 which represents 200µm and Sample 13 which represents 50µm.