

Using Bayesian Model Averaging (BMA) to calibrate probabilistic surface temperature forecasts over Iran

I. Soltanzadeh¹, M. Azadi², and G. A. Vakili²

¹Institute of Geophysics, University of Tehran, Tehran, Iran

²Atmospheric Science & Meteorological Research Center (ASMERC), Tehran, Iran

Received: 25 October 2010 – Revised: 8 July 2011 – Accepted: 11 July 2011 – Published: 21 July 2011

Abstract. Using Bayesian Model Averaging (BMA), an attempt was made to obtain calibrated probabilistic numerical forecasts of 2-m temperature over Iran. The ensemble employs three limited area models (WRF, MM5 and HRM), with WRF used with five different configurations. Initial and boundary conditions for MM5 and WRF are obtained from the National Centers for Environmental Prediction (NCEP) Global Forecast System (GFS) and for HRM the initial and boundary conditions come from analysis of Global Model Europe (GME) of the German Weather Service. The resulting ensemble of seven members was run for a period of 6 months (from December 2008 to May 2009) over Iran. The 48-h raw ensemble outputs were calibrated using BMA technique for 120 days using a 40 days training sample of forecasts and relative verification data.

The calibrated probabilistic forecasts were assessed using rank histogram and attribute diagrams. Results showed that application of BMA improved the reliability of the raw ensemble. Using the weighted ensemble mean forecast as a deterministic forecast it was found that the deterministic-style BMA forecasts performed usually better than the best member's deterministic forecast.

Keywords. Meteorology and atmospheric dynamics (Mesoscale meteorology)

1 Introduction

Ensemble forecasting is a numerical prediction method that samples the uncertainties in initial conditions and model formulation. Thus, rather than producing a single deterministic forecast, multiple forecasts are produced by making small alterations to either the initial conditions or the forecast model,

or both. Ensemble forecasts have been operationally implemented on the synoptic scale (Toth and Kalnay, 1993; Houtekamer and Derome, 1996; Molteni et al., 1996) and on the mesoscale (Stensrud et al., 1999; Wandishin et al., 2001; Gritmit and Mass, 2002; Eckel and Mass, 2005). Despite their relatively high skill, they tend to be under-dispersive and thus uncalibrated, especially for weather quantities at the surface.

In the last couple of years various statistical methods such as logistic regression (Wilks, 2006), Bayesian Model Averaging (Raftery et al., 2005), non-homogeneous Gaussian regression (Gneiting et al., 2005) and Gaussian ensemble dressing (Roulston and Smith, 2003; Wang and Bishop, 2005), among others, have been developed for calibrating the raw ensemble forecasts.

In this study the Bayesian Model Averaging (BMA) technique, proposed by Raftery et al. (2005), has been used to calibrate the raw outputs of a multi-model multi-analysis ensemble for 2-m temperature at 299 meteorological stations over Iran. In BMA, parameters (weights and variances) for a mixture of distributions (e.g. Gaussians) are estimated over a sliding-window training period of forecasts and observational data. Parameter estimation is accomplished by maximizing the log-likelihood or minimizing the Continuous Ranked Probability Score (CRPS) (Hersbach, 2000). A BMA weight is determined for each individual-member PDF allowing for unequal weighting of the component forecasts, if appropriate. One advantage of BMA technique is that it gives a full probability distribution as forecast. The predictive probability density function (PDF) of BMA is a weighted average of distributions centered on the ensemble member's forecasts after bias correction. The weights are posterior probabilities of the component models constructing the ensemble and reflect the forecasts' relative contribution to overall predictive PDF skill over a training period. The BMA method is robust to exchangeability assumptions and the BMA post-processed combined ensemble shows better verification results than raw ensemble systems. These results



Correspondence to: I. Soltanzadeh
(soltanz@ut.ac.ir)

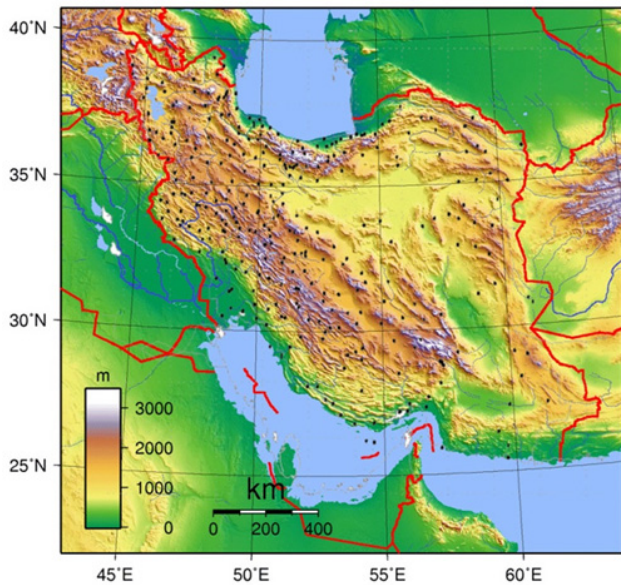


Fig. 1. Topography and Synoptic stations distribution over Iran.

suggest that statistically post-processed multi-model ensembles can outperform individual ensemble systems, even in cases in which one of the constituent systems is superior to the others. Beside the predictive PDF one can get a point or deterministic-style forecast by using just a weighted average of the individual forecasts in the ensemble.

Raftery et al. (2005) proposed and applied the BMA to surface temperature and mean sea level pressure forecasts of the University of Washington short-range ensemble with five members. They got calibrated and sharp predictive PDFs. This method has been used for post-processing the ensemble outputs for temperature, wind direction and precipitation in some other parts of the world with success. For example, Wilson et al. (2007) applied the BMA to calibrate surface temperature forecasts from the 16-member Canadian ensemble system. Assessment of the post-processed ensemble outputs with around 40 days training period showed a good calibration of the ensemble dispersion.

BMA was originally developed for weather quantities in which PDFs could be approximated by normal distributions, such as temperature and sea level pressure. Bao et al. (2010) extended and applied the BMA to 48-h forecasts of wind direction using von Mises densities as the component distributions centered at the individually bias-corrected ensemble surface wind direction and could get consistent improvements in forecasts. Sloughter et al. (2007) used a mixture of a discrete component at zero and a gamma distribution as predictive PDF for individual ensemble members and applied the BMA to daily 48-h forecasts of 24-h accumulated precipitation in the North American Pacific Northwest in 2003–2004 using the University of Washington mesoscale ensemble. They could get PDFs corresponding to probability of

precipitation forecasts that were much better calibrated compared to consensus voting of the ensemble members. The results of BMA were also better estimations of the probability of high-precipitation events than logistic regression on the cube root of the ensemble mean.

The present study aims at producing calibrated surface temperature forecasts at 299 meteorological stations scattered across Iran (Fig. 1) using a multi-model multi-analysis ensemble for the period from 15 December 2008 to the 11 June 2009. Predictive forecast PDFs' performances are evaluated using reliability and ROC diagrams. Point or deterministic-style BMA forecasts are compared with deterministic forecast of individual members using standard scores.

The paper is organized as follows: the BMA procedure is described briefly in Sect. 2, while the implementation details are presented in Sect. 3. Verification results are discussed in Sect. 4 and finally, conclusions and proposal for further works are drawn in Sect. 5.

2 Calibration method

Bayesian Model Averaging (BMA) was proposed by Raftery et al. (2005) as a statistical post-processing approach for combining different model forecasts and producing full predictive PDFs from ensembles, subject to calibration and sharpness. In the following a brief description of the method is presented, but the reader is referred to Raftery et al. (2005) for full details. The BMA predictive PDF of the weather quantity to be forecast is a weighted average of PDFs defined around each individual bias-corrected ensemble member. The individual PDFs for the ensemble members need not to be Gaussian or even the same. In this study, as in Raftery et al. (2005), the weather quantity to be forecast, y , is temperature whose behavior can be estimated by a normal distribution. Hence a Gaussian distribution, $h_k(y|f_k)$, is defined around each individual forecast, f_k , conditional on f_k being the best forecast in the ensemble. The BMA predictive PDF is then a conditional probability for a forecast quantity y given K model forecasts f_1, \dots, f_k , and is given by:

$$p(y|f_1, \dots, f_k) = \sum_{k=1}^K w_k g_k(y|f_k) \quad (1)$$

where w_k is the posterior probability of forecast k being the best one, and is based on forecast k 's relative performance over a training period. The w_k 's are nonnegative probabilities and their sum is equal to 1, that is $\sum_{k=1}^K w_k = 1$. Here K , the number of ensemble members, is equal to 7. $g_k(y|k)$ is a univariate normal PDF with mean $f_k = a_k + b_k f_k$, (bias-corrected forecast) that is a linear function of forecast f_k , and standard deviation σ^2 assumed to be constant across ensemble members. This situation is denoted by:

$$y|f_k \sim N(a_k + b_k f_k, \sigma^2) \quad (2)$$

Table 1. Ensemble members configuration.

| Member name | Cumulus | Planetary boundary layer | Microphysics | Long wave radiation | Short wave radiation | Surface layer | Land surface | Initial conditions |
|-------------|-------------------------------|--------------------------|--------------------|-----------------------------|----------------------|---------------------------------------|---------------------------|--------------------|
| WRF1 | Kain-Fritsch (new Eta) scheme | MRF scheme | Kessler scheme | RRTM scheme | Dudhia scheme | Monin-Obukhov scheme | Noah land-surface Model | GFS |
| WRF2 | Grell-Devenyi ensemble scheme | MRF scheme | Kessler scheme | RRTM scheme | Dudhia scheme | Monin-Obukhov scheme | Noah land-surface Model | GFS |
| WRF3 | Betts-Miller-Janjic scheme | MRF scheme | Kessler scheme | RRTM scheme | Dudhia scheme | Monin-Obukhov scheme | Noah land-surface Model | GFS |
| WRF4 | Kain-Fritsch | Mellor-Yamada-Janjic | Lin | RRTM scheme | Goddard | Monin-Obukhov–Janjic scheme | Noah | GFS |
| WRF5 | Grell-Devenyi ensemble | Mellor-Yamada-Janjic | WSM3 | RRTM scheme | Dudhia scheme | Eta similarity | 5-layer thermal diffusion | GFS |
| MM5 | Betts-Miller-Janjic scheme | Mellor-Yamada-Janjic | Mixed phase | CCM2 | CCM2 | Monin-Obukhov scheme | 5-layer soil model | GFS |
| HRM | Mass flux (Tiedtke) | Mellor-Yamada-Janjic | Doms and Schättler | δ-2 stream radiation scheme | | Level-2 scheme and 7-layer soil model | | GME |

Coefficients a_k and b_k in the mean of the individual PDFs vary with time and location and are estimated by a linear regression of observed temperature, y , on model k forecasts, f_k , in the training period, for each time and location separately. This regression can be considered as a preliminary debiasing of the deterministic forecasts in the ensemble. The K weights or posterior probabilities w_k and variance σ^2 are estimated using maximum likelihood (Fisher et al., 1922). For a fixed set of training data and underlying normal probability model, the method of maximum likelihood selects values of the model parameters that maximize the likelihood function, that is, the value of the parameter vector under which the observed data were most likely to have been observed. For both mathematical simplicity and numerical stability, usually the log-likelihood function is used for maximization rather the likelihood function itself. Usually, the maximum value of the log-likelihood function is evaluated using the expectation-maximization (EM) algorithm (Dempster et al., 1977; McLachlan and Krishnan, 1997). The BMA deterministic forecast also can be calculated by weighted averaging of the K deterministic forecasts using w_k as weights, that is:

$$\sum_{k=1}^K w_k (a_k f_k + b_k) \quad (3)$$

3 The ensemble system and data

Forecasts of the Weather Research and Forecasting (WRF: Skamarock et al., 2008) model with five different configurations, The fifth-generation Pennsylvania State University–National Center for Atmospheric Research Mesoscale Model (MM5: Dudhia, 1993; Grell et al., 1994) and High Resolution Model (HRM: Majewski, 1991, Majewski and Schrodin, 1994) of the Deutscher Wetter Dienst (DWD) both with one configuration for 2-m temperature, 48-h in advance are used in this study to build a seven-member ensemble. The model settings are presented in Table 1. As seen in the table the main differences between different model setups

pertain to convective and boundary layer parameterization schemes. WRF and MM5 are used with non-hydrostatic option whereas HRM is hydrostatic. The initial and boundary conditions come from the operational 12Z runs of the global forecasting system (GFS) of NCEP (National Center for Environmental Prediction) (Sela, 1980) for MM5 and WRF and of DWD’s global model (GME) for HRM models respectively. The integration period goes from 15 December 2008 to 11 June 2009.

The period of study includes spring and summer seasons. In the onset of spring, subtropical anticyclone migrates to northern latitudes and hence baroclinic mid-latitude systems are weakened over south and central part of Iran. During summer time the mid-latitude synoptic baroclinic systems enter Iran from west and north-west and influence the country only in the north-west. The cold air associated with these systems, causes rapid temperature changes in the northern part of Iran. As we enter from spring to summer, the subtropical anticyclone becomes dominant over most part of the country and many places experience their maximum temperature during summer. In some places the temperature exceeds 50 and reaches even 55 °C. In the southern part of the country, there is an interaction between Iranian heat low and the Indian monsoon low that causes the convective systems to become dominant and play an important role in the temperature changes over these regions.

MM5 and WRF were run with two nested domains, with the larger domain covering the south-west middle east from 10° to 51° north and from 20° to 80° east and the smaller domain covers Iran from 23° to 41° north and from 42° to 65° east. The spatial resolutions are 45 and 15-km for the coarser and finer domains, respectively. The inner domain in HRM considered here, covers from 25° to 40° north and from 43° to 63° east with spatial resolution of 14-km. Totally 1134 simulation have been performed and forecasts out to +48h ahead for the inner domains have been used to form the ensemble of forecasts.

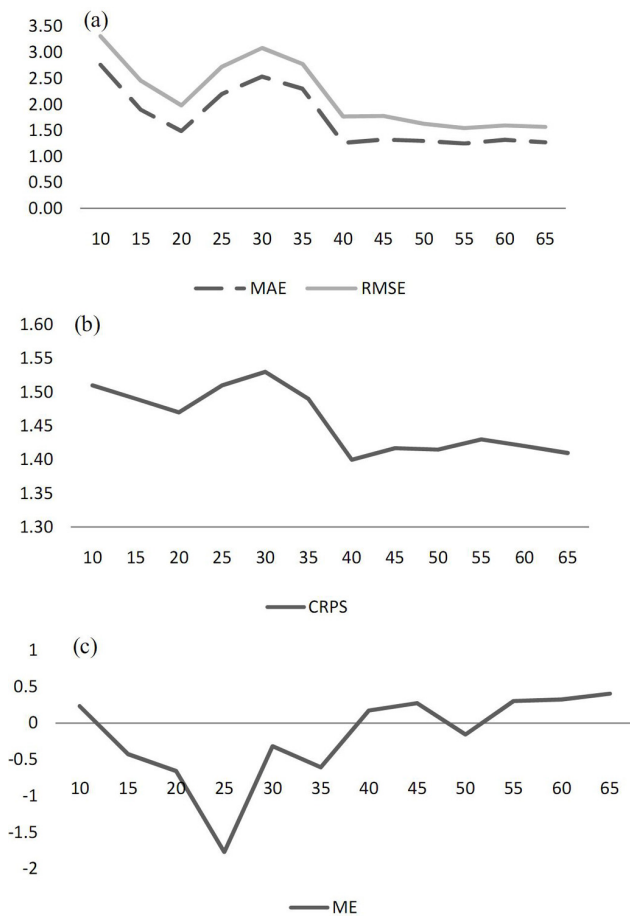


Fig. 2. Comparison of training period length for surface temperature: diagram of (a) MAE, RMSE, (b) CRPS, and (c) ME for from 10 to 65 days.

The data used in this study consist of 12Z observations of 2-m temperature at 299 irregularly spaced synoptic meteorological stations scattered all over the country from 15 December 2008 to the 11 June 2009 and corresponding 48-h forecasts from the above mentioned seven members of the ensemble bi-linearly interpolated to the observation sites. The geographical distribution of the synoptic stations is presented in Fig. 1. Using N days as training period, the BMA predictive PDF, Eq. (1), and BMA deterministic forecast, Eq. (3), for the 2-m temperature were evaluated for each station site and the remaining days.

4 Training period

The sample of past days, N , used as training period, in estimating the unknown parameters (a_k , b_k , w_k and σ) in Eq. (2) is a sliding training window, such that new coefficients are estimated for each day using the most recent N days as training period. In principle, length of the training period must be such that it does not lead to over-fitting. How-

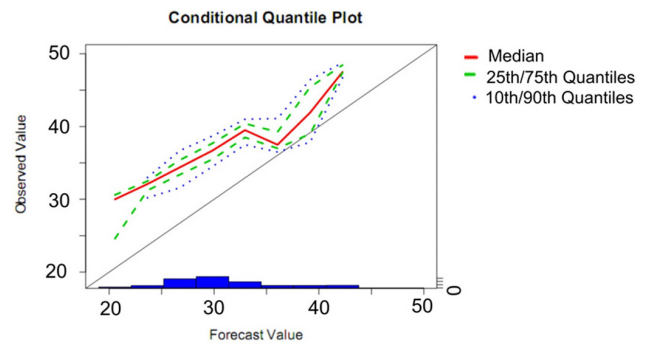


Fig. 3. Conditional Quantile Plot (CQP) diagram for MM5 2-m temperature forecasts.

ever, longer training periods increase statistical variability and shorter training periods make the forecast system to respond more quickly to changes in the model error patterns due to weather regime changes. It is clear that using short training periods have the advantage of data availability and ease of computations. A balance should be made in determining the length of the training period. In this study, as in Raftery et al. (2005), some experiments were conducted with training period changing from 10 to 65-day. Figure 2 shows the mean absolute error (MAE), root mean squared error (RMSE) and mean error (ME) of the BMA deterministic forecasts and the continuous rank probability score (CRPS) versus training days. Minimum errors can be seen on the figure around 20 and 40 days training period. But increasing the training days beyond 40 does not change the results significantly and the error remains almost constant or even increases. So, a 40-day window was selected as the training period.

5 Results and discussion

5.1 Raw ensemble

Figure 3 presents an example of the conditional quantile plot (CQP) (Wilks, 2006) for 48-h temperature forecast of one member (member 6: MM5) in the ensemble for the period mentioned in Sect. 2. It is clear that deterministic forecasts, corresponding to the 7th ensemble member show a cold bias and thus a debiasing of the forecasts is needed. The CQPs related to other ensemble members (not shown here) show similar results, i.e. a cold bias.

Rank histograms are very useful tools for evaluating an ensemble forecast system performance (e.g. Hamill and Colucci, 1997, 1998; Hou et al., 2001; Stensrud and Yusuf, 2003). A rank histogram is a histogram of the observation ranks when pooled in the sorted forecasts of the ensemble members. Hamill (2001) shows how to use rank histograms for evaluating ensemble forecasts appropriately. The rank histogram of the raw ensemble is presented in Fig. 4.

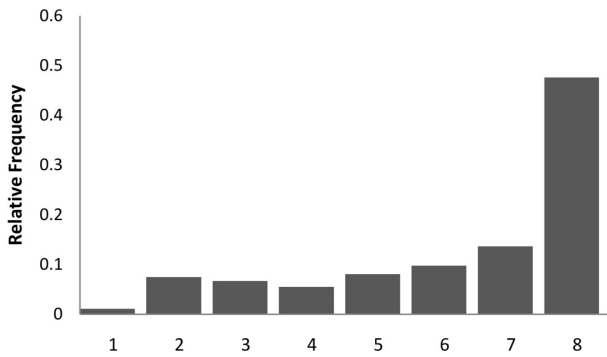


Fig. 4. Rank histogram of the raw ensemble for 48-h surface temperature forecasts.

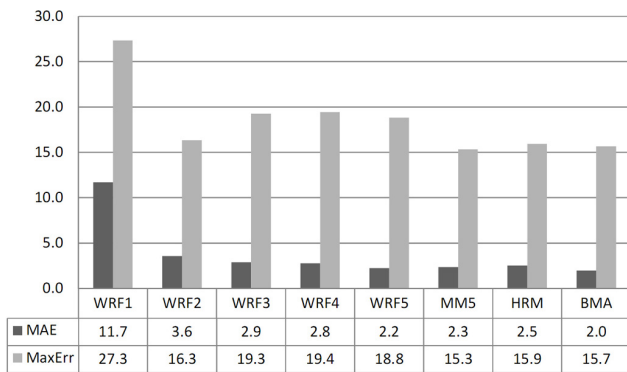


Fig. 5. Minimum Absolute Error (MAE) and maximum error for all ensemble members and calibrated BMA outputs.

Under-dispersion of the ensemble is reflected in the non-uniform shape of its rank histogram. As it is seen from the figure, the rank histogram for the raw ensemble is a sloped one, showing a consistent bias in the ensemble forecast. The ensemble members have under-forecast or cold bias, such that around 50 % of the times the observed temperature was greater than all the ensemble member values. This result is consistent with the results presented in Fig. 3. The under-dispersion for the raw ensemble has been reported in many other studies. The goal of post-processing is to correct for such known forecast errors, i.e. to construct a calibrated ensemble with statistical properties similar to the observations.

5.2 Deterministic BMA forecast

For comparing the deterministic BMA forecast (Eq. 3) with the deterministic forecast corresponding to the best member of the ensemble, the mean absolute error (MAE) and percentage of successful forecasts (forecasts with less than 2 °C difference from the verifying observation) are considered. In terms of MAE, Fig. 5 shows that MAE of the deterministic forecasts of the seven members of the ensemble system are between 2.2 to 6.7 °C, while that of the BMA deterministic forecast is lower and around 2 °C.

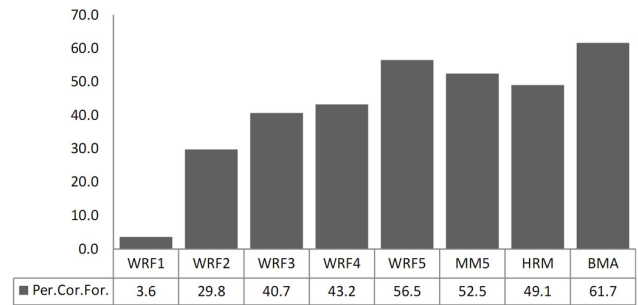


Fig. 6. Histogram of percentage of acceptable forecast of raw data and BMA outputs for ± 2 °C error.

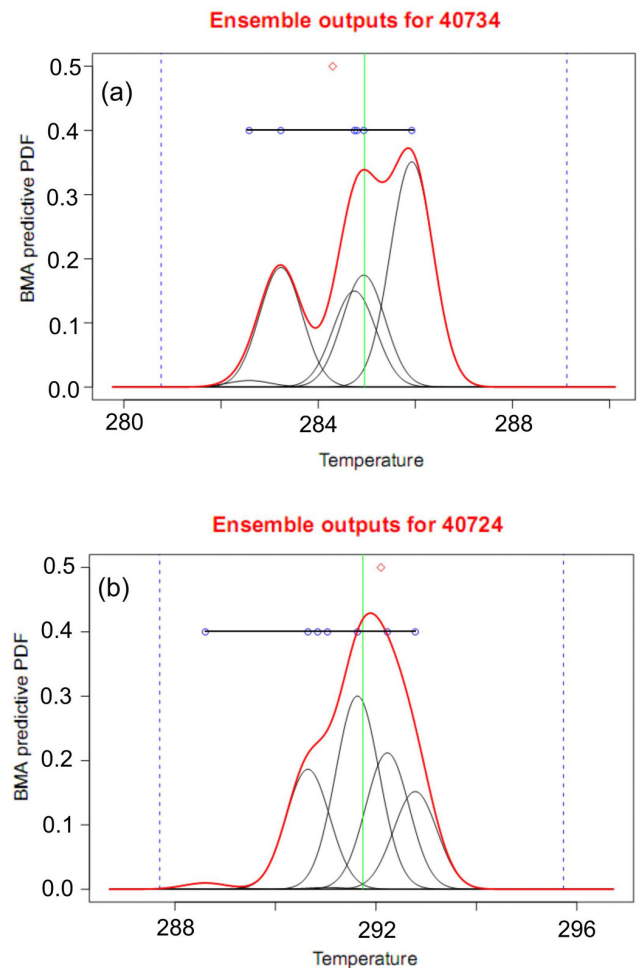


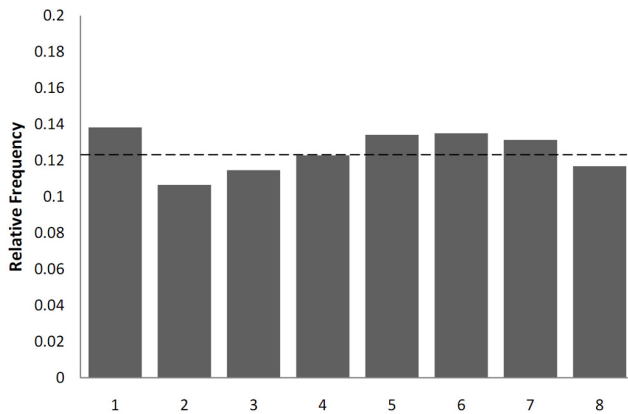
Fig. 7. BMA predictive PDF (thick red curve) and its seven components (thin black curves) for the 48-h surface temperature for two synoptic stations of (a) Piranshahr and (b) Noshahr in initialized at 12:00 UTC on first February and first May 2009, respectively. Also shown are the ensemble member forecasts and range (solid horizontal line and blue bullets), observation (red bullets), the BMA 90 % prediction interval (dotted lines) and the median of the PDF (solid green vertical line).

Table 2. Weights of two synoptic stations of Piranshahr and Noshahr in initialized at 12:00 UTC on 1 February and 1 May 2009, respectively.

| HRM | MM5 | WRF5 | WRF4 | WRF3 | WRF2 | WRF1 | Date | Station |
|-------|-------|-------|-------|-----------------------|-----------------------|----------------------|----------|------------|
| 0.348 | 0.175 | 0.246 | 0.216 | 3.02×10^{-3} | 7.1×10^{-6} | 0.0112 | 20090401 | Noshahr |
| 0.20 | 0.170 | 0.399 | 0.212 | 0.0115 | 2.98×10^{-6} | 6.8×10^{-3} | 20090501 | Piranshahr |

Table 3. Performance scores for eight categories ranging from “equal to 0 °C” out to “above 30 °C” with 5 °C interval.

| | Freezing point | 0 to 5 | 5 to 10 | 10 to 15 | 15 to 20 | 20 to 25 | 25 to 30 | above 30 |
|------------------------|----------------|--------|---------|----------|----------|----------|----------|----------|
| Brier Score (BS) | 0.017 | 0.052 | 0.102 | 0.123 | 0.102 | 0.069 | 0.035 | 0.013 |
| Brier score – baseline | 0.024 | 0.076 | 0.158 | 0.188 | 0.162 | 0.124 | 0.063 | 0.027 |
| Skill score | 0.305 | 0.324 | 0.359 | 0.345 | 0.374 | 0.447 | 0.443 | 0.543 |
| Reliability | 0.002 | 0.002 | 0.001 | 0.000 | 0.002 | 0.001 | 0.002 | 0.003 |
| Resolution | 0.010 | 0.027 | 0.057 | 0.065 | 0.062 | 0.057 | 0.030 | 0.017 |
| Uncertainty | 0.024 | 0.076 | 0.158 | 0.188 | 0.162 | 0.124 | 0.063 | 0.027 |

**Fig. 8.** Probability Integral Transform (PIT) histogram of calibrated BMA outputs.

Percentage of the successful forecasts for seven members of the ensemble, shown in Fig. 6, ranged from 3.6 % to 56.5 % for the first and fifth members, respectively. This score for the BMA deterministic forecast is close to 62 % which is again better than those of all ensemble members. It is thus seen that the BMA deterministic forecast outperforms all other seven deterministic forecasts corresponding to the ensemble members.

5.3 Calibrated ensemble

One main aim of the ensemble forecasting is to account for various uncertainties in the ensemble system for issuing probabilistic forecasts. Figure 7a and b shows two examples of the final predictive BMA for 48-h forecast of 2-m temperature valid at first February and first May 2009, along with

its seven normal PDF components issued for two synoptic stations of Piranshahr and Noshahr located in the west and north of the country. As is seen, the calibrated BMA PDF that is a weighted sum of its seven components, is a non-normal distribution. Table 2 shows the calculated weights given to each member of the ensemble. It is seen that for Noshahr, the weights in descending order are given to HRM, WRF5, WRF2 and MM5 members respectively, while for Piranshahr are given to WRF5, WRF4, HRM and MM5. As mentioned above, a higher weight given to a member means that member is more useful. But, as mentioned by Gneiting et al. (2005), low weight for a member does not mean necessarily a lower performance of that member. If there are colinearities between two (or more) ensemble members, then their informations are similar and one of them might be given low weight though this member alone might be skillful.

One important aim of applying the BMA technique is to obtain a well calibrated ensemble with reduced underdispersion. The fact that the predictive PDFs are calibrated is reflected in the uniformity of the post-processed ensemble rank histogram for 48-h forecasts, presented in Fig. 8. As the figure shows, the BMA has been very successful in calibrating the raw ensemble forecasts.

Performance scores such as Brier score and Skill score for eight temperature categories are presented in Table 3. It is apparent that, BMA provides reliable and skillful probabilistic forecasts for most quantiles for all 299 station locations over Iran. More detailed comparison of BMA calibrated forecasts with the raw ensemble can be obtained from the attribute diagram for probability forecast of particular quantiles. Attribute diagram shows how well the predicted probabilities of an event correspond to their observed frequencies. Figure 9a shows the attribute diagram for raw ensemble which has been derived from democratic voting method for 20 to

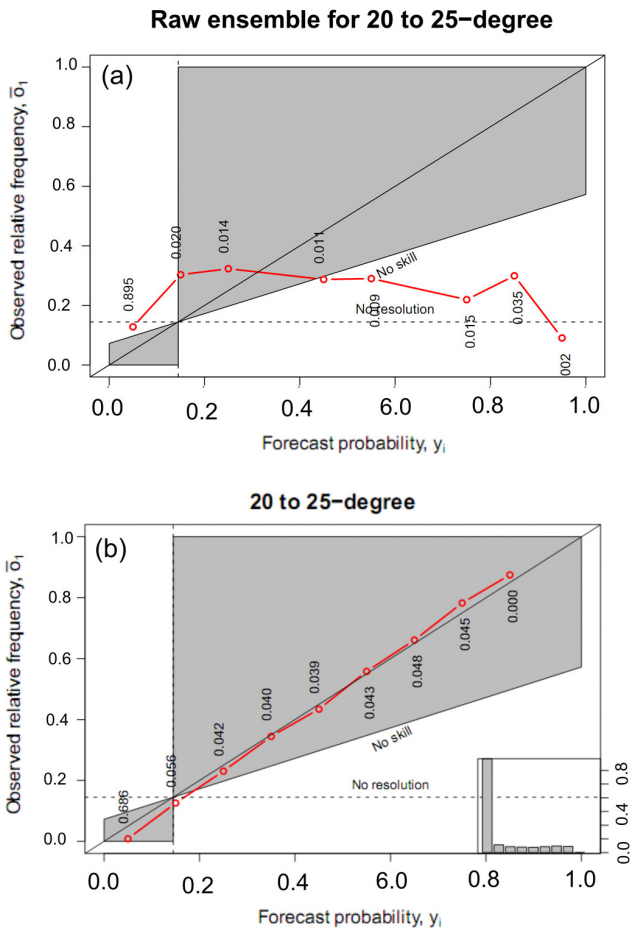


Fig. 9. Attribute diagrams for (a) raw ensemble and (b) BMA for “20 to 25 °C” interval. In this figure a horizontal “no resolution” line is drawn at the climatological frequency. A vertical line is also drawn for a forecasted probability at the observed frequency, a “perfect reliability line” is also drawn with a slope of 1.

25 °C interval which indicates for most of probabilities raw ensemble has no skill. Attribute diagrams of BMA weighted mean for the same interval in Fig. 9b shows significant improvement over the raw ensemble (Fig. 9a and b).

The ROC (Relative Operating Characteristic or Receiver Operating Characteristic) diagram is another discrimination-based graphical forecast verification display (Wilks, 2006). The ROC was first introduced into meteorology by Mason (1982), is a graphical plot of hits against false alarms, using a set of increasing probability thresholds. The area under the ROC curve is a measure of performance. An area equal to 1 implies perfect performance, while an area of 0.5 corresponds to random forecasts or to the climatological forecasts. Results from the calibrated ensemble for various threshold values, as those used for attribute diagrams, for 48-h temperature forecasts indicate that the ensemble exhibit good event discrimination (Fig. 10). It is seen that for warmer thresholds (greater than 20 °C) and freezing point, the ROC curve bows

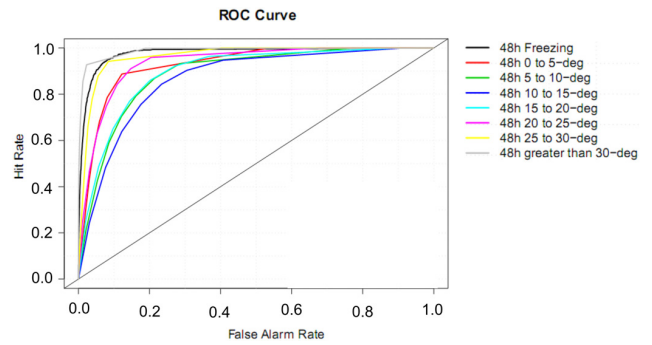


Fig. 10. Relative operating characteristic (ROC) diagrams for eight temperature categories ranging from “equal to 0 °C” out to “above 30 °C” with 5 °C interval.

the most beyond the diagonal line and thus the ensemble for these thresholds shows more discriminating than others.

Figure 11 shows BMA predictive PDF for two extreme events one with minimum and another with maximum 2-m temperature. Minimum surface temperature during the study period happened on 2 January 2009 at Jolfa located in the north of Iran, while Nik-Shahr in the south-east of the country experienced the maximum temperature record on 28 April 2009. In both cases, most of the ensemble members have not simulated the surface temperature; this is probably because mesoscale models have difficulty in correctly forecasting extreme temperatures. BMA gave the highest weight to the member that had better surface temperature prediction, WRF5 (Fig. 11a), while for other case, minimum temperature event, the best forecast of surface temperature was given by MM5 and WRF5, for this case, MM5 underestimated the temperature and was assigned the least weight while WRF5 obtained the highest weight.

6 Conclusion

This paper describes the results of 48-h probabilistic surface temperature forecasts over Iran for the period of 15 December 2008 to 11 June 2009 using Bayesian Model Averaging for calibration of the ensemble outputs. The ensemble system consists of the WRF model with five different configurations, MM5 and HRM both with one configuration. The initial and boundary conditions come from the operational 12Z runs of GFS for MM5 and WRF, and GME for HRM models respectively.

The probabilistic forecasts were accomplished for 299 synoptic station locations scattered across Iran. The experiment was set up in such a way that it could be run in real time operations; the BMA was trained on recent realizations of the forecast errors, and then applied to the subsequent forecasts in the 3 months test period. Based on the results of several experiments with different training sample sizes from 10 to 65-day, a 40-day window was selected as the training period.

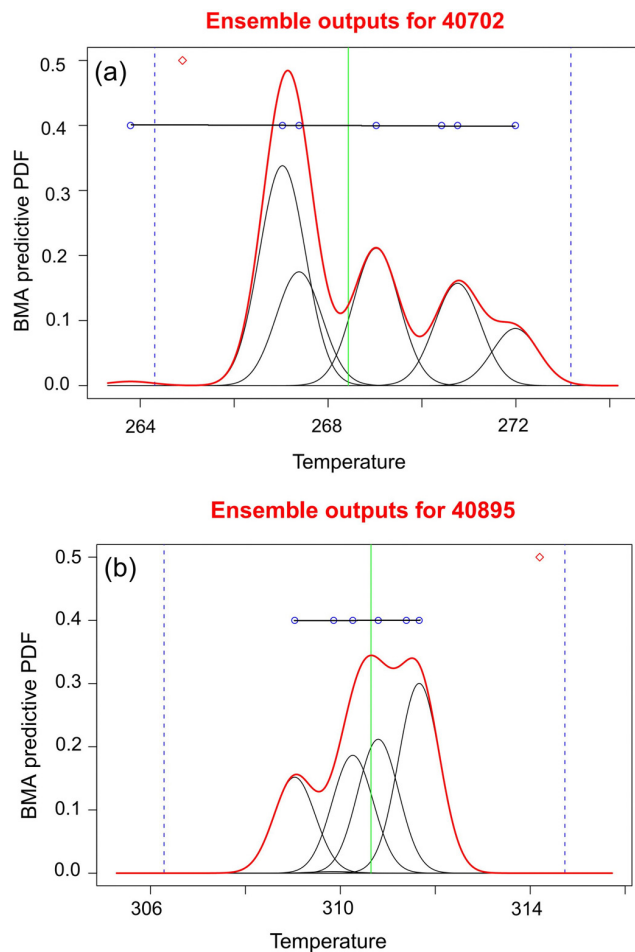


Fig. 11. Same as Fig. 7, but for (a) Jolfa on 2 January 2009 and (b) Nik-Shahr on 28 April 2009.

Overall results showed that the BMA technique is almost successful at removing most, but not all of the underdispersion exhibited by the raw ensemble and thus attaining higher reliability in the probabilistic forecasts that could be used in an operational framework. Using the weighted ensemble mean forecast as a deterministic forecast it was found that the deterministic-style BMA forecasts performed almost always better than the best member's deterministic forecast in the ensemble.

Acknowledgements. Financial support for this research is gratefully acknowledged from Khuzestan Water & Power Authority Co., Iran. The authors also appreciate the useful discussions and suggestions from D. Parhizkar from Atmospheric Science and Meteorological Research Center (ASMERC) and M. Khodadi from I. R. of Iran Meteorological Organization (IRIMO).

Topical Editor P. Drobinski thanks C. Santos and another anonymous referee for their help in evaluating this paper.

References

- Bao, L., Gneiting, T., Grimit, E. P., Guttorp, P., and Raftery, A. E.: Bias Correction and Bayesian Model Averaging for Ensemble Forecasts of Surface Wind Direction, *Mon. Weather Rev.*, 138, 1811–1821, 2010.
- Dempster, A., Laird, N., and Rubin, D.: Maximum likelihood from incomplete data via the EM algorithm, *J. R. Stat. Soc., Series B*, 39(1), 1–38, 1977.
- Dudhia, J.: A nonhydrostatic version of the Penn State/NCAR mesoscale model: Validation tests and the simulation of an Atlantic cyclone and cold front, *Mon. Weather Rev.*, 121, 1493–1513, 1993.
- Eckel, F. A. and Mass, C. F.: Aspects of effective mesoscale, short-range ensemble forecasting, *Wea. Forecasting*, 20, 328–350, 2005.
- Fisher, R. A., Thornton, H. G., and Mackenzie, W. A.: The accuracy of the plating method of estimating the density of bacterial populations, CP22 in Bennett 1971, vol. 1., *Annals of Applied Biology*, 9, 325–359, 1922.
- Gneiting, T., Raftery, A. E., Westveld, A. H., and Goldman, T.: Calibrated probabilistic forecasting using ensemble Model Output Statistics and minimum CRPS estimation, *Mon. Weather Rev.*, 133, 1098–1118, 2005.
- Grell, G. A., Dudhia, J., and Stauffer, D. R.: A description of the fifth-generation Penn State/NCAR mesoscale model (MM5), NCAR Tech. Note TN-398+STR, 122 pp., 1994.
- Grimit, E. and Mass, C.: Initial results of a mesoscale short-range ensemble forecasting system over the Pacific Northwest, *Weather Forecast.*, 17, 192–205, 2002.
- Hamill, T. M.: Interpretation of rank histograms for verifying ensemble forecasts, *Mon. Weather Rev.*, 129, 550–560, 2001.
- Hamill, T. M. and Colucci, S. J.: Verification of Eta-RSM short-range ensemble forecasts, *Mon. Weather Rev.*, 125, 1312–1327, 1997.
- Hamill, T. M. and Colucci, S. J.: Evaluation of Eta-RSM ensemble probabilistic precipitation forecasts, *Mon. Weather Rev.*, 126, 711–724, 1998.
- Hersbach, H.: Decomposition of the Continuous Ranked Probability Score for Ensemble Prediction Systems, *Weather Forecast.*, 15, 559–570, 2000.
- Hou, D., Kalnay, E., and Droegemeier, K. K.: Objective verification of the SAMEX'98 ensemble forecasts, *Mon. Weather Rev.*, 129, 73–91, 2001.
- Houtekamer, P. L. and Derome, J.: The RPN ensemble prediction system. Proceedings, ECMWF Seminar on Predictability. Vol. II. ECMWF, 121–146, available from ECMWF, Shinfield Park, Reading, Berkshire RG2 9AX, United Kingdom, 1996.
- Majewski, D.: The Europa-Model of the DWD, ECMWF Seminar on numerical methods in Atmospheric Science, 2, 147–191, 1991.
- Majewski, D. and Schrodin, R.: Short description of Europa-Model (EM) and Deutschland Model (DM) of the DWD, *Q. Bull.*, 1994.
- Mason, I. B.: A model for assessment of weather forecasts, *Austral. Met. Mag.*, 30, 291–303, 1982.
- McLachlan, G. and Krishnan, T.: The EM algorithm and extensions. Wiley series in probability and statistics, John Wiley & Sons, 1997.
- Molteni, F., Buizza, R., Palmer, T. N., and Petroliagis, T.: The ECMWF ensemble prediction system: Methodology and val-

- idation, *Q. J. Roy. Meteorol. Soc.*, 122, 73–119, 1996.
- Raftery, A. E., Gneiting, T., Balabdaoui, F., and Polakowski, M.: Using Bayesian model averaging to calibrate forecast ensembles, *Mon. Weather Rev.*, 133, 1155–1174, 2005.
- Roulston, M. S. and Smith, L. A.: Combining dynamical and statistical ensembles, *Tellus*, 55A, 16–30, 2003.
- Sela, J. G.: Spectral Modeling at the National Meteorological Center, *Mon. Weather Rev.*, 108, 1279–1292, 1980.
- Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D., Barker, D., Wang, W., and Powers, J. G.: A description of the Advanced Research WRF Version 3, NCAR Tech. Note NCAR/TN-475+STR, 2008.
- Sloughter, J. M., Raftery, A. E., Gneiting, T., and Fraley, C.: Probabilistic quantitative precipitation forecasting using Bayesian model averaging, *Mon. Weather Rev.*, 135, 3209–3220, 2007.
- Stensrud, D. J. and Yussouf, N.: Short-range ensemble predictions of 2-m temperature and dewpoint temperature over New England, *Mon. Weather Rev.*, 131, 2510–2524, 2003.
- Stensrud, D. J., Brooks, H. E., Du, J., Tracton, M. S., and Rogers, E.: Using ensembles for short-range forecasting, *Mon. Weather Rev.*, 127, 433–446, 1999.
- Toth, Z. and Kalnay, E.: Ensemble forecasting at NMC: The generation of perturbations, *B. Am. Meteorol. Soc.*, 74, 2317–2330, 1993.
- Wandishin, M. S., Mullen, S. L., Stensrud, D. J., and Brooks, H. E.: Evaluation of a short-range multimodel ensemble system, *Mon. Weather Rev.*, 129, 729–747, 2001.
- Wang, X. and Bishop, C. H.: Improvement of ensemble reliability with a new dressing kernel, *Q. J. Roy. Meteorol. Soc.*, 131, 965–986, 2005.
- Wilks, D. S.: *Statistical Methods in the Atmospheric Sciences*, 2nd Edition, Academic Press, 627 pp., 2006.
- Wilson, L. J., Beauregard, S., Raftery, A. E., and Verret, R.: Calibrated surface temperature forecasts from the Canadian ensemble prediction system using Bayesian model averaging, *Mon. Weather Rev.*, 135, 1364–1385, 2007.