

Seasonal to hour variation scales in abundance and production of total and particle-attached bacteria in the open NW Mediterranean Sea (0–1000 m)

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Abstract. We present the vertical and temporal dynamics of total vs. particle-attached bacterial abundance and activity over a 5 week period under summer to autumn transition in NW Mediterranean Sea. At a weekly time scale, total bacterial biomass and production in the euphotic layers was significantly correlated with phytoplanktonic biomass. At an hourly time scale, total bacterial biomass responded very rapidly to chlorophyll *a* fluctuations, suggesting a tight coupling between phytoplankton and bacteria for resource partitioning during the summer-autumn transition. In contrast, no influence of diel changes on bacterial parameters was detected. Episodic events such as coastal water intrusions had a significant positive effect on total bacterial abundance and production, whereas we could not detect any influence of short wind events whatever the magnitude. Finally, we show that particle-attached bacteria can represent a large proportion (up to 49%) of the total bacterial activity in the euphotic layer but display rapid and sporadic changes at hourly time scales. In the mesopelagic layers, bacterial abundance and production linearly decreased with depth, except some production peaks at 400–750 m. This study underlines the value of large datasets covering different temporal scales to clarify the biogeochemical role of bacteria in the cycling of organic matter in open seawater.

1 Introduction

At concentrations ranging between 10^4 – 10^6 cells ml⁻¹, marine bacteria represent the most abundant and biogeochemically important organisms in the oceans. One-half of oceanic primary production on average is channelled via bacteria into the microbial loop (Azam et al. 1983; Cole et al. 1988). In oligotrophic systems, heterotrophic bacterial biomass is generally equal to that of the phytoplanktonic or often greater (Robarts et al., 1996; Socal et al., 1999). Microbes consume an estimated 75% of the sinking particulate organic carbon flux in the upper 500 m (Karl et al., 1988; Cho and Azam, 1990). Beyond bulk abundance and activity estimates, the qualitative attributes and activities of the bacterioplankton are crucial to ecosystem function, and an especially key role of the microbial loop is as a major pathway of carbon transfer in marine systems (Azam et al., 1983). The fate of dissolved organic carbon (DOC) pool is mainly determined by the activity of heterotrophic bacteria, which act as a link or sink of DOC for higher trophic levels. The fraction of primary production used by bacteria is highly variable over various time and space scales (Ducklow et al., 1993). Recently, a network of “Microbial Observatories” has been developed in different habitats and across environmental gradients to explore the magnitude and the variation scales of the fluxes driven by microorganisms (www.nsf.gov/bio/pubs/awards/mo03.htm). However, marine microbial observatories generally rely on coastal areas. Indeed, numerous studies have shown that bacterial abundance and activity vary at a seasonal scale



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in a large number of coastal waters (Shiah and Ducklow, 1994; Pinhassi and Hagström, 2000; Gerds et al., 2004; Pearce et al., 2007; Sapp et al., 2007), but considerably less information is available in open Ocean (Ducklow and Carlson, 1992). Substrate availability and, to a lesser extent, temperature have been to be the main driving forces of bacterial abundance and production at a seasonal scale (Wikner and Hagström, 1999; Lemée et al., 2002). Short time scale variations of bacterial abundance and production have been explored in more detail in both coastal and open oceans. Examples of significant diel variations of bacterioplankton activity in response to varying photosynthetic rates have been reported from various marine environments (Turley and Lochte, 1986; Herndl and Malacic, 1987; Gasol et al., 1998; Kuipers et al., 2000; Ghiglione et al., 2007). While these studies indicate that phyto- and bacterio-plankton activities can change on the scale of hours, cell abundance and biomass are often more constant, suggesting an efficient recycling mechanism of carbon and nutrients within the microbial food web during such diel cycles. To our knowledge, the effect of episodic events in open oceans, such as salinity changes due to coastal water intrusion or wind events has been poorly investigated. These interactive and successive rapid processes frequently occur during the seasonal transition period. For example, if a storm is sufficiently strong, it may induce vertical mixing by increasing the depth of the surface mixed layer and upwelling NO_3^- from the deep reservoir. The nitrate so entrained could stimulate photosynthesis then secondary production in the euphotic zone (Platt et al., 1992; Wu et al., 2007).

In this study, we present the vertical and temporal dynamics of heterotrophic bacteria under summer-autumn transition at a site located near the DYFAMED station (NW Mediterranean Sea). This site was chosen because (i) it is far enough away from the Ligurian Current to be sufficiently protected from lateral transport, thereby permitting a 1D study and (ii) it is very close to the JGOFS time-series station DYFAMED, which means that long time series data set of biological, biogeochemical and physical parameters are available. The seasonal variations of the biogeochemical production regimes have been well studied in this area: deep convection occurs during winter leading to a spring bloom; oligotrophic conditions prevail during summer while perturbations in the meteorological forcing generate a secondary bloom in autumn (Marty, 2002). The vertical and temporal dynamics of heterotrophic bacteria (Van Wambeke et al., 2001; Lemée et al., 2002; Tanaka and Rassoulzadegan, 2004; Misic and Fabiano, 2006) and the contribution of particle-attached bacteria (Turley and Stutt, 2000; Harris et al., 2001) have already been reported in the NW Mediterranean Sea. In a previous study, we have shown that the contribution of attached bacteria to total bacterial activity can reach up to 83% under mesotrophic condition, then reinforcing the biogeochemical role of this fraction in the cycling of particulate organic carbon in the NW Mediterranean Sea (Ghiglione et al., 2007). However,

the temporal scales at which bacterial abundance and production vary through the water column are still unknown. A companion paper (Ghiglione et al., 2008) explored the community composition of the bacterioplankton along the depth gradient with molecular biology tools but do not investigate the effect of episodic events because of methodological considerations. In this study, we provide a large set of data on the vertical (0–1000 m) and temporal (from hour to week scales, during 5 weeks) dynamics of total vs. particle-attached bacterial abundance and activity under summer-autumn transition at a site very close to the JGOFS-DYFAMED station in NW Mediterranean Sea. The seasonal pattern was investigated by comparison with results previously obtained at the station DYFAMED (Ghiglione et al., 2007).

2 Materials and methods

2.1 Study site and samplings

Sampling was carried out during the cruise DYNAPROC-2 (DYNamics of the rapid PROCesses in the water column) conducted from 14th September (julian day, JD 258) to 18th October 2004 (JD 292) at an offshore station (28 miles offshore, $43^\circ 25' 4'' \text{N}$, $8^\circ 00' 5'' \text{E}$) located near the permanent DYFAMED station ($43^\circ 25' 2'' \text{N}$, $07^\circ 51' 8'' \text{E}$) in the NW Mediterranean Sea on board RV “Thalassa”. Samples were collected with a rosette system equipped with twelve 24-l Niskin bottles and a conductivity-temperature-fluorescence-depth profiler (CTD – *SeaBird SBE 911 plus*). For measurements at week-time scales, samples were taken at least every day at noontime at 6 depths (5, 20, 40, 60, 80, 150 m) from JD 261 to JD 273 (17th to 29th September) and from JD 277 to JD 289 (3th to 15th October). For studies at hourly time-scales, samples were taken every 6 h at the same depths from JD 268 to JD 273 (24th to 29th September) and from JD 284 to JD 289 (10th to 15th October). In addition, surface to 1000 m depth profiles (5, 20, 40, 60, 80, 150, 200, 400, 500, 750, 1000 m) were performed 12 times at day and night during the cruise (18th, 19th and 26th September=JD 262–263 and JD 270 and 5th, 6th and 12th October=JD 279–280 and JD 286). Low salinity water masses (LSW) percentage in the water column was depicted by the *P* index, based on average salinity *S* in the 40–70 m water layer, according to the following formula: $P \text{ index} = (S_{\text{max}} - S(\text{sta})) / (S_{\text{max}} - S_{\text{min}})$; *S*(sta) is the average salinity in the 40–70 m layer, *S*_{min} is that of the water taken as reference for desalinated water (coastal waters in our case), *S*_{max} is that of the salinity waters without anomaly (see <http://www.obs-vlfr.fr/proof/vt/op/ec/peche/pec.htm> – data DYNAPROC log & basic files – for more detailed explanation). *In situ* fluorescence was converted to chlorophyll-*a* (Chla) using a regression between *in situ* fluorescence and measurements of the water column Chla concentrations (mg Chla m^{-3}) from selected

depths ($\text{Chl}a=2.0740 \times \text{Fluo}$ ($R^2=+0.970$, $p<0.01$, $n=453$) from JD 261–273 and $\text{Chl}a=1.7807 \times \text{Fluo}$ ($R^2=+0.960$, $p<0.01$, $n=466$) from JD 277–289 (Andersen et al., this issue¹). $\text{Chl}a$ concentrations were converted into carbon equivalents by applying the average conversion factors 32.5 for the upper mixed layer, 18.5 for the upper part of deep chlorophyll maximum (DCM) and 12.7 below (Van Wambeke et al., 2002).

2.2 Bacterial abundance and biomass

1.5 ml sub-samples for total bacteria counts were put in a cryo-vial, fixed for at least 20 min at room temperature with 1% (v/v) glutaraldehyde (final concentration), frozen in liquid nitrogen and later stored at -80°C freezer for later analysis. Bacteria cells were enumerated using SYBR Green staining and flow cytometry according to the method described by Marie et al. (1997). Briefly, glutaraldehyde-fixed sub-samples were thawed and incubated with SYBR Green (Molecular Probes, Eugene, OR, USA) at a final concentration of 10^{-4} (v/v) for 15 min at room temperature in the dark. Analyses were performed with a FAC Sort flow cytometer (Becton Dickinson, San Jose, CA, USA) equipped with an air-cooled argon laser (488 nm, 15 mW). Samples were run at low speed (approx. $20 \mu\text{l min}^{-1}$) and data were acquired in log mode until around 10 000 events had been recorded. We added $10 \mu\text{l}$ per sample of a 10^6 ml^{-1} solution of yellow-green fluorescent micro spheres ($0.95 \mu\text{m}$ diameter beads – Polysciences Inc., Warrington, Pa) as an internal standard.

Total bacteria were detected by their signature in a cytogram of side scatter (SSC) versus green fluorescence (FL1). In the same cytograms, two main bacterioplankton groups can be discriminated: HNA bacteria with high nucleic acid content (high FL1 value) and LNA bacteria with low nucleic content (low FL1 value). The SSC/FL1 cytogram for each sample was individually inspected, and HNA and LNA regions manually moved to conform to the appropriate areas of the bacterial dot-plot. Cytometric noise was discarded both by setting a threshold on FL1 and by manually separating noise from cells in the FL1 versus FL3 plot (Gasol et al., 1999). Data acquisition was performed with the Cell Quest system (Becton Dickinson) and data analysis with a custom-designed software CYTOWIN version 4.3 (Vaulot, 1989). HNA bacteria abundance was expressed as the percentage of total bacteria population.

Bacteria attached to particles were enumerated in the $>0.8 \mu\text{m}$ size fraction. Immediately after collection, 5 ml samples were filtered onto $0.8 \mu\text{m}$ pore-size polycarbonate membranes (Isopore Filters, Millipore). Then, filters were placed in 5 ml 2% glutaraldehyde and kept at -80°C . Thawed filters were sonicated back in the laboratory ($306 \mu\text{m}$ amplitude, 50% duty cycle, 2 min, cooling in

water bath) with a Sonifier 250 (Branson Ultrasonics Corp., Danbury, CO, USA) to disperse bacteria from the particles (Velji and Albright, 1993). Bacteria contained in the $>0.8 \mu\text{m}$ size fractions (attached bacteria) were enumerated using SYBR Green staining and flow cytometry as described above. Attached bacteria abundance was expressed as the percentage of the total bacteria population. Four deep profiles (11 levels between 0–1000 m) were studied using both epifluorescence microscopic method (Hobbie et al., 1977) and flow cytometry. Comparison of microscopic observations before and after sonication showed that our sonication procedure on $0.8 \mu\text{m}$ filters do not damage the cells and resulted in $>95\%$ recovery of particle-attached bacterial cells. In addition, 44 samples analyzed by both methods showed no significant differences (mean= 5.9 ± 1.8 and $6.8 \pm 1.3 \times 10^5 \text{ cells ml}^{-1}$ by epifluorescence and flow cytometry, respectively) and were closely correlated ($R=+0.68$, $p<0.01$, $n=44$) that confirmed the validity of using sonication before flow cytometric enumerations as shown previously (Riemann and Winding, 2001; Worm et al., 2001).

Heterotrophic bacterial biomass (BB) was calculated by using a carbon content per cell of $15 \text{ fg C cell}^{-1}$ generally used in oligotrophic ecosystems (Fukuda et al., 1998; Caron et al., 1999), and conversion factors taking into account photoacclimation were used to evaluate the autotrophic biomass (Van Wambeke et al., 2002).

2.3 Bacterial production

Bacterial heterotrophic production (BP) was estimated from ^3H -leucine (Amersham, 161 Ci mmol^{-1}) incorporation rates into bacterial proteins as described by Kirchman et al. (1985). The centrifugation method of Smith and Azam (1992) was used routinely to estimate total bacterial production (TBP) in all profiles (0–150 and 0–1000 m) whereas the classical filtration method (Kirchman, 1993) was used to estimate bacterial production of particle-attached bacteria (ABP). For TBP, 1.5 ml samples were dispersed into four 2 ml screw cap micro-centrifuge tubes. A mixture of ^3H -leucine and non-radioactive leucine was added to final concentrations of 4 and 16 nm l^{-1} , respectively. These concentrations were experimentally determined to be a saturation concentration in all depths. One of the replicates had already received trichloroacetic acid (TCA) to a 5% final concentration and acted as the dead control. Samples were incubated in the dark at the in situ temperature for 2 to 8 h, depending on sampling depth. Previous experiments had shown leucine incorporation to be linear over these time periods (data not shown). Incubations were stopped by the addition of 50% TCA (5% final concentration). The samples were centrifuged at 16.000 g for 10 min. The supernatant was discarded and 1.5 ml of 5% TCA was added. The samples were shaken vigorously using a vortex mixer and centrifuged again. The supernatant was discarded and 1.5 ml of scintillation cocktail (NBS 104, Amersham) was added.

¹Andersen, V., Goutx, M., Prieur, L., and Dolan, J.: Short-scale temporal variability in the Mediterranean, submitted in BGS, 5, 2008.

Table 1. Seasonal evolution of total bacteria biomass (TBB), total bacterial production (TBP) and total specific activity (TSA) in the upper layer (0–150 m) at the Dyfamed site. The data (mean \pm SD) are integrated over the upper 0–150 m depth. N=number of samples analysed. March and June 2003 data originated from previous cruises at the station DYFAMED (see Ghiglione et al., 2007).

	TBB (mg C m ⁻²)	TBP (mg C m ⁻² h ⁻¹)	TSA (fg C cell ⁻¹ day ⁻¹)
March 2003 (N=24)	2757.6 \pm 81.6	52.8 \pm 16.8	0.29 \pm 0.08
June 2003 (N=24)	472.8 \pm 211.2	20.4 \pm 3.6	0.75 \pm 0.46
Sep–Oct 2004 (N=324)	720.0 \pm 111.6	28.8 \pm 6.0	0.53 \pm 0.09

In the filtration method, three replicates (10 to 30 ml according to sampling depth) and one control pre-killed with formalin (2% final concentration), received 20 nm l⁻¹ leucine mix (2 nm l⁻¹ ³H-leucine and 18 nm l⁻¹ cold leucine). Samples were incubated as for the centrifugation method and the incubations were stopped using formalin (2% final concentration). The samples were filtered through 0.2 μ m cellulose ester filters (Millipore, type GS) and were extracted with 5% TCA for 10 min followed by three 3 ml rinses with 5% TCA. Filters were then placed in scintillation vials and dissolved with 0.5 ml ethyl acetate prior to the addition of 5 ml of NBCS 104 Amersham scintillation cocktail. Because TCA passed through 0.8 μ m pore-size filters during the protein extraction, a direct estimation of Bacterial Production by attached bacteria (ABP) was not possible. Consequently, two samples sets were incubated for each measurement: a first set was treated as above for estimation of total BP and an additional set was filtered through 0.8 μ m after incubation and stop with formalin. The proteins extracted from the <0.8 μ m fraction permitted to estimate the free-living bacterial production. The production in <0.8 μ m fraction was always smaller than the total production. The ABP was calculated as the difference between TBP and BP by free-living bacteria.

Radioactivity of all micro-centrifuge tubes and 0.2 μ m filters were counted on a tri-CARB 1500 Packard liquid scintillation counter. Quenching was corrected by internal standard and control counts were subtracted. The mean coefficient of variation of the triplicate measurements was 8,7 and 12,5% for centrifugation and filtration method, respectively. Centrifugation and filtration methods were compared on the basis of 84 samples and we did not find any significant difference between the two methods (mean TBP=22.52 \pm 5.50 from 1.5 ml samples and 20.01 \pm 4.85 ng C l⁻¹ h⁻¹ from 10–30 ml samples with centrifugation and filtration methods, respectively). Rates of bacterial production (BP) were calculated from leucine incorporation rates using a conversion factor of 1.5 kg C mol⁻¹ leucine (Kirchman, 1993) and were expressed as ng C l⁻¹ h⁻¹. The per-cell specific activity was calculated from the BP/BA ratio and expressed as fg C cell⁻¹ day⁻¹.

2.4 Statistical analyses

Statistical analyses were performed with STATISTICA software. With the exception of the percentages of HNA and attached bacteria, all variables were log-transformed in order to attain normality and homogeneity of variables. The relationships between variables were explored by use of Pearson's correlation coefficient. A long-term trend was determined by linear regression of parameters with time and a F-test on slopes was performed. Student's *t*-test was used to assess significant differences between paired means. The whole integrated values were calculated according to the classical trapezoidal method. All data were reported as means \pm SD.

3 Results

3.1 Environmental conditions

The main hydro-biological characteristics of the studied site during the DYNAPROC-2 cruise (14th September to 18th October 2004 – JD 258 to JD 292) are presented by Andersen et al. (this issue¹). Briefly, this seasonal transition period was marked by a strong water column stratification partially disrupted at the end of the cruise and low nutrients stocks. The apparent stability of the hydro-biological structure of the water column prevailing during the five week sampling period was disturbed by various episodic meteorological events. The most outstanding events were the intrusion of low salinity water masses (LSW) (<38.2‰) occurring below the thermocline (between 40 and 80 m) during 10 days from JD 265 (21th September) and for 5 days from JD 283 (9th October). The increase of LSW percentage in the 20–70 m water column is depicted by the index of low salinity water (see <http://www.obs-vlfr.fr/proof/vt/op/ec/peche/pec.htm>). Three strong wind events (speed >20 nds) took place for 12, 24 and 60 h on JD 269 (25th September), JD 284 (10th October) and JD 286 (12th October), respectively. The two last wind events induced a strong decrease of air temperature, a beginning of de-stratification and the mixed layer deepened (Andersen et al., this issue¹). At the beginning of the cruise, two deep-chlorophyll maxima (DCM, 50–60 m and 90 m depth) were detected, resulting in a phytoplankton

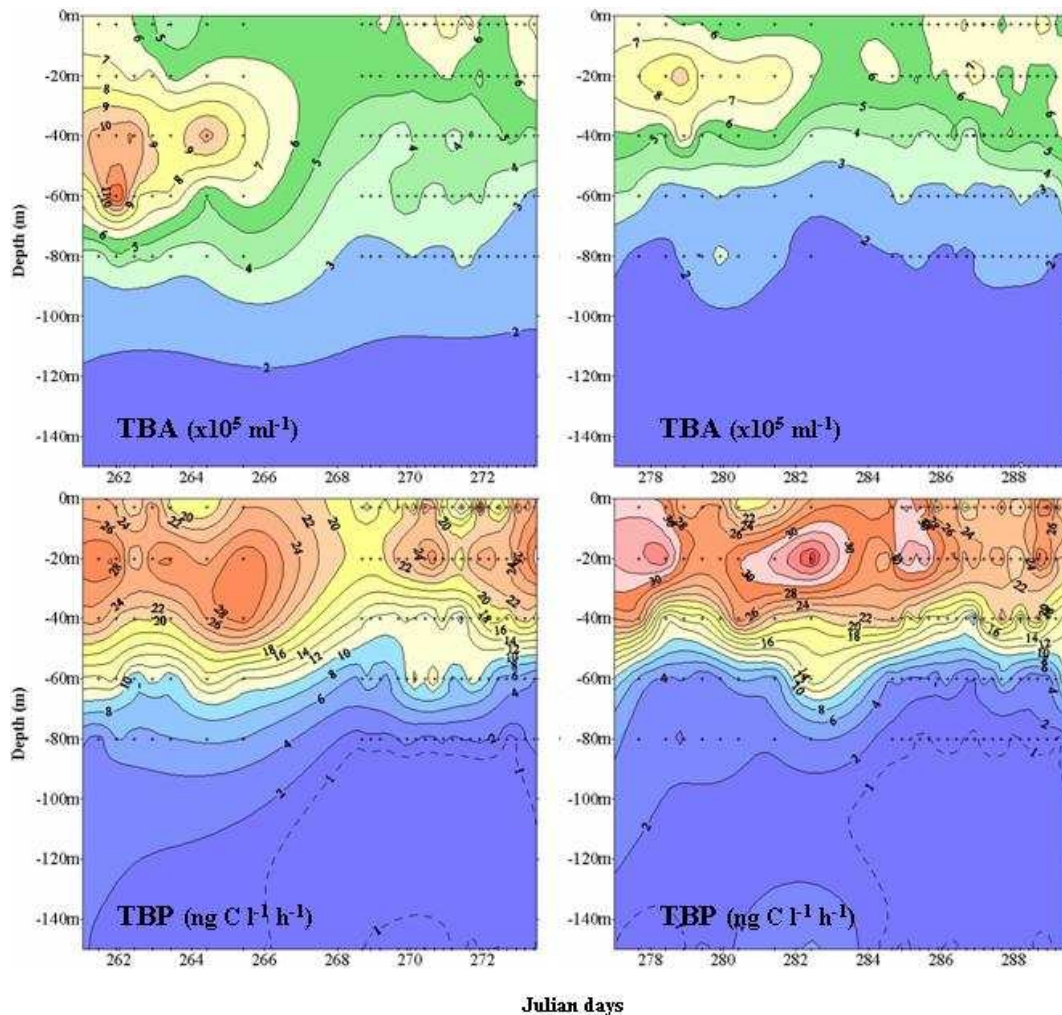


Fig. 1. Five week variations of total bacterial abundance (TBA) and total bacterial production (TBP) in the upper zone (0–150 m depth). Crosses indicate sampling times and depths.

biomass exceptionally high for the time period (Chl_a concentration of 35–40 mg m⁻²). After JD 263 (19th September), only one DCM was observed at 40–50 m depth with Chl_a concentration of 20–25 mg m⁻² (Marty et al., 2008)

3.2 Temporal variation scales of bacterial abundance and production within 0–150 m depths

3.2.1 Seasonal values of total bacterial abundance and production

The results we obtained during the September–October 2004 sampling period were compared to those obtained for two previous cruises conducted at the same site in March and June 2003 (Ghiglione et al., 2007) (Table 1). In spring condition (March 2003), total bacterial biomass (TBB) was 5.9 and 3.8 time more important, total bacterial production (TBP) was 2.6 and 1.8 time more important and total spe-

cific activity (TSA) was 2.6 and 1.8 time less important than in summer (June 2003) and summer-autumn (September–October 2004) conditions, respectively. Then, September–October 2004 bacterial parameter values were much closer to the oligotrophic summer 2003 than the mesotrophic spring 2003 conditions.

3.2.2 Weekly variation of total bacterial abundance and production

Total bacterial abundance (TBA) generally increased from surface to the DCM and decreased below with maximal values observed at the surface waters at the end of the two studied periods (JD 269–273 and JD 286–289) (Fig. 1). TBA ranged from a minimum of 4.6×10^5 cells ml⁻¹ in subsurface to a maximum of 14.8×10^5 cells ml⁻¹ at the DCM and decreasing to a minimum of 2.1×10^5 cells ml⁻¹ at 150 m (mean = $6.1 \pm 2.4 \times 10^5$ cells ml⁻¹, n=324). We

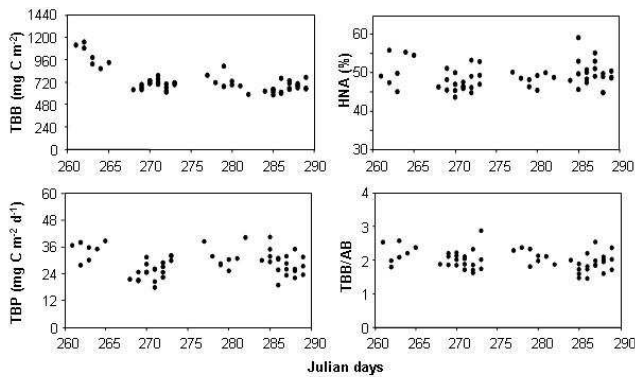


Fig. 2. Evolution of bacterial parameters integrated over the upper 0–150 m. TBB=total bacterial biomass; TBP=total bacterial production; HNA=relative abundance of cells with high nucleic acid content to the total bacterial abundance; TBB/AB=TBB to autotrophic biomass ratio.

found a good correlation between TBA and *Chla* integrated concentration in the first 150 m depth ($R=+0.44$, $p<0.05$, $n=54$). Total bacterial production (TBP) generally increased from surface to 20 m and decreased below, ranging from a minimum of $10.6 \text{ ng C l}^{-1} \text{ h}^{-1}$ in the sub-surface to a maximum of $38.6 \text{ ng C l}^{-1} \text{ h}^{-1}$ at 20 m and decreasing to a minimum of $0.3 \text{ ng C l}^{-1} \text{ h}^{-1}$ at 150 m depth (mean= $15.01 \pm 10.16 \text{ ng C l}^{-1} \text{ h}^{-1}$, $n=324$) (Fig. 1). TBP variability with depth and time was higher compared to TBA (CV=68 and 40% respectively, $n=324$) and these two parameters were strongly linked during the entire sampling period ($R=+0.59$, $p<0.01$, $n=324$).

Integrated total bacterial biomass (TBB) and production (TBP), as well as the relative contribution of cells with high nucleic acid content (HNA expressed in % of TBA) and total bacterial biomass to autotrophic biomass ratio (TBB/AB) were calculated on for the 0–150 m upper layer (0–150 m) (Fig. 2). Integrated TBB values showed a slight decrease from 1150.8 to 864.0 mg C m^{-2} from JD 261 to JD 265 (17th to 21th September) and remained almost stable during the rest of the sampling period (mean= $686.0 \pm 60.5 \text{ mg C m}^{-2}$, $n=47$). Integrated TBP values ranged from 17.9 to 40.4 $\text{mg C m}^{-2} \text{ day}^{-1}$ (mean= $28.8 \pm 5.5 \text{ mg C m}^{-2} \text{ day}^{-1}$, $n=54$). The relative contribution of HNA bacteria to TBA was rather stable during the 5 weeks sampling period (from 44 to 59%, mean= $49 \pm 3\%$, $n=54$) as well as the ratio TBB/AB (from 1.45 to 2.87, mean= 2.00 ± 0.29 , $n=54$). These four parameters (TBB, TBP, HNA and TBB/AB) fitted to a linear regression with time showed no long-term trend during the one-month sampling period ($p>0.05$, variance analysis F-test).

3.2.3 Diel variation scale of total bacterial abundance and production

In order to evaluate the influence of diel and episodic meteorological events (strong wind or intrusion of low salinity water masses), short time changes of the TBB and TBP were investigated. Samples were taken in the 0–150 m upper layer by sampling every 6 h during 120 h in two periods (JD 268–273 and JD 284–289=24th to 29th September and 10th to 15th October) and at 6 depths levels from surface to 150 m depth. The P Index (proxy of salinity anomalies) integrated between 20 and 70 m depth indicated that the first coastal water intrusion that appeared from JD 264 to JD 273 (Sept 20th to 29th) was more intense than the second (from JD 282 to JD 286=8th to 12th October) (Fig. 3). The temperature of the surface water (5 m) remained relatively stable during the first study period (JD 268–273) except for the first days and decreased with time during the second period (JD 284–289) in accordance to the strong wind period occurring in the middle of this sampling period.

During the two periods, 0–150 m integrated parameters (TBB_{0–150m}, TBP_{0–150m}, *Chla*_{0–150m} and HNA_{0–150m}%) showed strong variations on short-time scales (Fig. 3). No clear diel periodicity of these parameters could be detected. Interestingly, TBB_{0–150m} variations during the two 120 h periods presented the same pattern that *Chla*_{0–150m} variations ($R=+0.48$, $p<0.05$, $n=19$ for the first period and $R=+0.39$, $p<0.05$, $n=20$ for the second period) (Fig. 3). In addition, TBP_{0–150m} variations were significantly correlated with the relative abundance of the high nucleic acid content bacteria (HNA%) ($R=+0.71$, $p<0.01$, $n=19$ and $R=+0.47$, $p<0.05$, $n=20$, for the first (JD 268–273) and second (JD 284–289) study period, respectively) and these relationships were better than TBP/TBB ($R=+0.36$, $p<0.05$, $n=19$ and $R=+0.23$, $p<0.05$, $n=20$ for the two periods respectively). Finally, significant negative correlations were found between the P index (proxy of salinity anomalies) and both 0–150 m depth integrated TBA and TBP. These relationships were stronger during the first water intrusion from JD 265 to JD 273 ($R=-0.85$, $p<0.01$ and $R=-0.47$, $p<0.05$, $n=19$ for TBA and TBP, respectively) compared to the second from JD 283 to JD 286 ($R=-0.36$, $p<0.05$, $n=20$ only for TBA) when the low salinity water mass was lower, as shown by the index P values (Fig. 3).

3.3 Temporal variation scales of total vs. particle-attached bacterial abundance and production

3.3.1 Weekly to diel variation scales of total vs. particle-attached bacterial abundance and production in the 0–1000 m water column

Twelve 0–1000 m depth profiles were carried out every day and night by sampling at 11 depths to follow weekly to diel changes in the vertical distribution of the total and

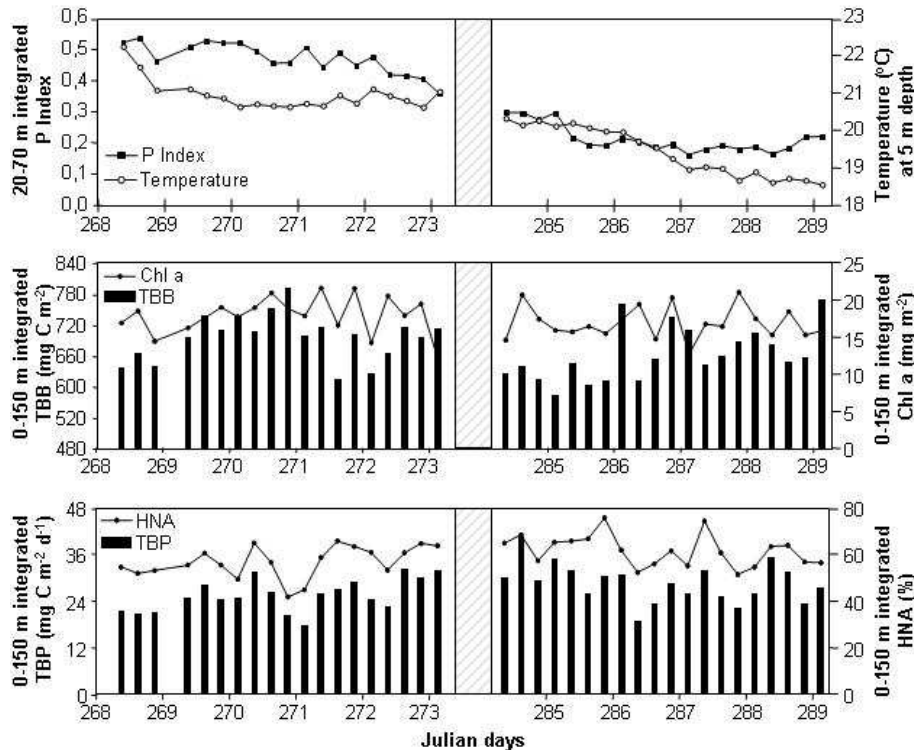


Fig. 3. Short time scale variability ($dt=6$ h) of 20–70 m depth integrated P index (a proxy of salinity anomalies), 5 m depth temperature, and 0–150 m depth integrated chlorophyll a , total bacterial biomass (TBB), relative abundance of cells with high nucleic acid content to the total bacterial abundance (HNA) and total bacterial production (TBP) within 2 periods (JD 268–273 and JD 284–289).

particle-attached bacterial abundance and production in the water column (Figs. 4 and 5). TBA exhibited a pronounced vertical gradient with maxima at 40 m depth at the beginning of the sampling period moving up to 5 m at the end of the sampling period with concentration values of 14.4 to 7.4×10^5 cells ml^{-1} , respectively (Fig. 4). TBA decreased drastically in the mesopelagic layers (150–1000 m) with the lowest values at the 1000 m depth, ranging from 0.5 to 0.7×10^5 cells ml^{-1} . No significant difference was observed between noontime and midnight. The abundance of particle-attached bacteria followed a similar pattern with a strong vertical gradient (Fig. 5). The main contribution of the attached bacteria to the TBA was observed in the upper 150 m layer with maximal values ranging from 14 to 25% of the TBA at the DCM, whereas their contribution was very low in the mesopelagic layers ($<4\%$ of the TBA). The day and night samples were compared for total and attached bacterial abundance (each depth for the twelve 0–1000 m profiles) and Student's t -test showed no significant difference between noontime and midnight. TBP exhibited also a pronounced vertical gradient with maximum values established at 20 m during the whole cruise, ranging from 20.3 and 30.9 $\text{ngC l}^{-1} \text{h}^{-1}$, and minima at 1000 m depth, ranging from 0.03 to 0.08 $\text{ngC l}^{-1} \text{h}^{-1}$ (Fig. 4). One exception was found on JD 279, when a relatively high TBP (8.10 $\text{ngC l}^{-1} \text{h}^{-1}$)

was observed at 500 m during the night. Other deep production peaks occurred in the 400–500 m layer during the second part of the sampling period but they do not appear on the graphs because of the low x-scale. The contribution of the particle-attached fraction on the TBP varied also drastically with depth and showed a similar vertical gradient as TBP (Fig. 5). The maximal contribution of the attached fraction reached values up to 49% of TBP, and generally situated between 20–40 m (just above DCM layer). In the mesopelagic waters, this contribution was generally negligible ($<2\%$), with some exceptions found between 400 and 750 m depths (from 8 to 15% of TBP). As observed for the abundance parameter, the day and night samples were compared for total and attached bacterial production and Student's t -test also showed no significant changes between day and night.

3.3.2 Hourly variation of total vs. particle-attached bacterial abundance and production at the subsurface

The influence of episodic meteorological events (strong wind or intrusion of low salinity water) on total vs. particle-attached bacterial abundance, production and specific activity was investigated by sampling every 6 h during 120 h in two periods (JD 268–273 and JD 284–289) (Fig. 6). For all of these parameters, both total vs. attached bacteria

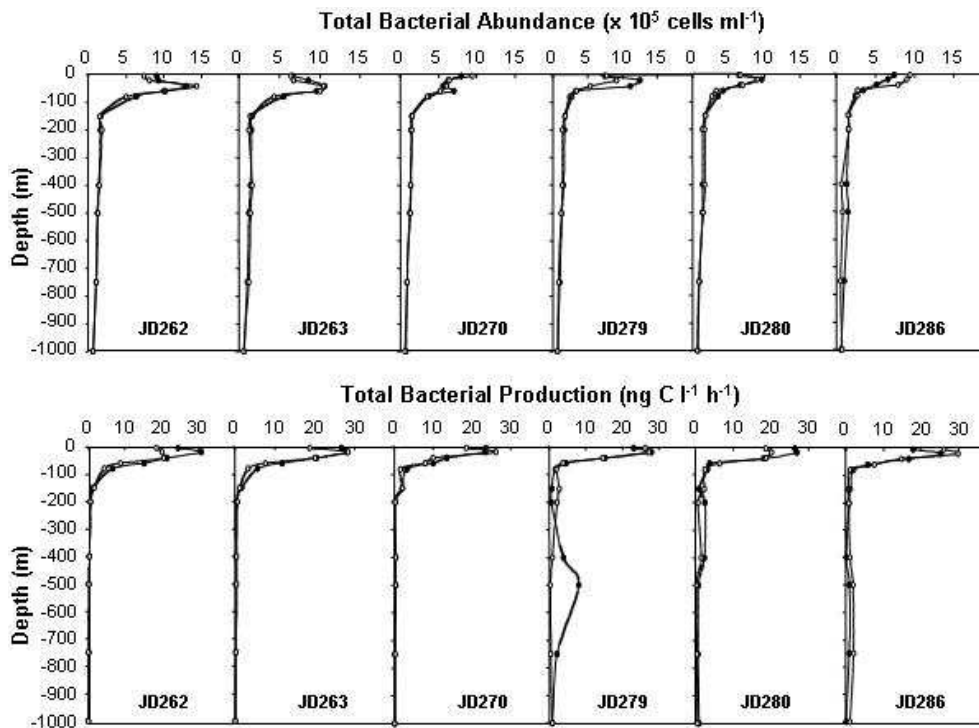


Fig. 4. Vertical changes in the total bacteria abundance and total bacterial production in the 0–1000 m water column during night (●) and day (○). Agreement with Julian days is: 18/09=262; 19/09=263; 26/09=270; 04/10=279; 05/10=280; 12/10=286.

exhibited sporadic temporal variations with neither diel pattern nor any clear relation with salinity or temperature events (Fig. 6). The contribution of the attached fraction represented from 7 to 17% of the TBA (mean=11±3%) in the first period (JD 268–273), while the dynamic of the attached fraction varied inversely to TBA ($R=-0.47$, $p<0.05$, $n=19$). In the second period (JD 284–289), the contribution of the attached fraction exhibited the same pattern than TBA and varied from 11 to 21% (mean=15.8±2.8%) of TBA. The contribution of the attached fraction to the TBP fluctuated from 25 to 55% (mean=34.7±9.8%) and from 18 to 63% (mean=37.7±13.2%) during the first and second periods, respectively (Fig. 6).

The ratio between bacterial production and bacterial abundance, namely specific activities (SA) was generally more important for the attached fraction compared to the total fraction. SA averaged from 1.82 ± 0.93 and 2.03 ± 1.01 fg C cell⁻¹ day⁻¹ for the attached fraction, while total SA averaged from 0.57 ± 0.15 and 0.79 ± 0.17 fg C cell⁻¹ day⁻¹ in the first and second period, respectively. As well as for bacterial production, total and particle-attached bacterial SA exhibited similar pattern with time, with higher BP and SA during the second period.

4 Discussion

4.1 Seasonal to hourly variation in bacterial abundance and production in the 0–150 m depth layer

One of the main contributions of our study is provision of a complete time series data set on heterotrophic bacterial abundance and production at seasonal to hourly scales at a central point in the Ligurian Sea close to the French JGOFS-DYFAMED station. To the best of our knowledge this is the first time that such a complex dataset is analyzed to attempt to unravel the temporal scales of variations in bacterioplankton abundance and activity in the NW Mediterranean Sea.

To evaluate the dynamics of total heterotrophic bacterial abundance and activity at a seasonal scale, data from this study (summer-autumn transition) were compared with previous data from two cruises conducted at the same site under spring bloom and summer stratification conditions (Ghiglione et al., 2007) (Table 1). As expected, total bacterial biomass and production values obtained in this study present slightly higher values than under summer oligotrophic condition but drastically lower values than under spring bloom conditions. This observation is in good agreement with the summer to autumn transition period characterizing the sampling period of the present study and with the strong seasonal variability of hydrological structures prevailing in this area (Andersen and Prieur, 2000; Vidussi et al., 2000).

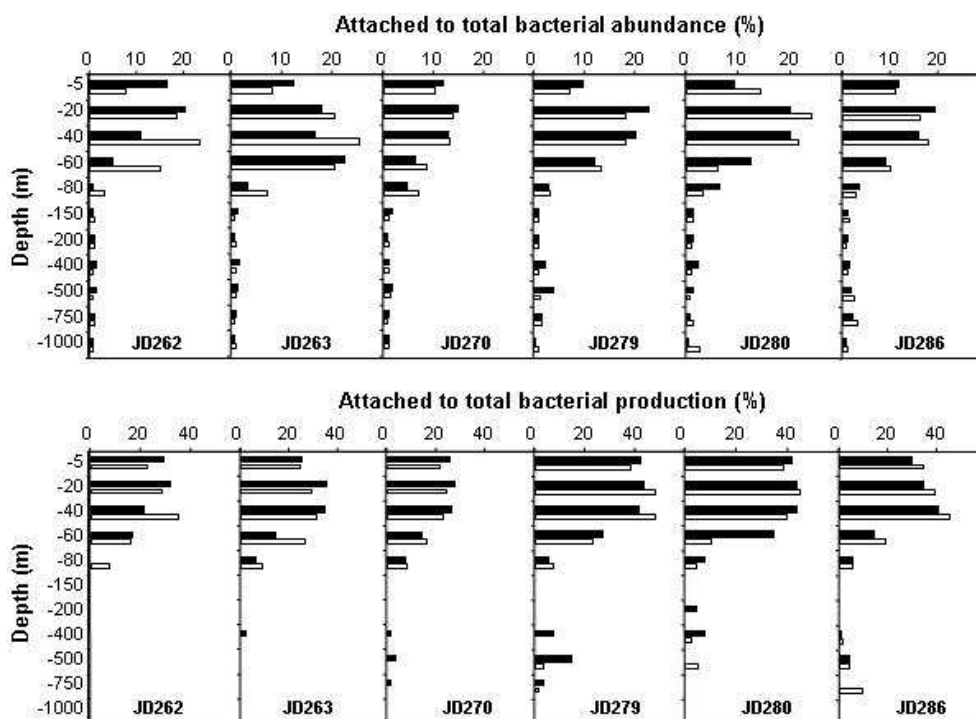


Fig. 5. Vertical changes of the attached to total bacterial abundance ratio in the 0–1000 m water column during night (■) and day (□). Agreement with Julian days is: 18/09=JD 262; 19/09=JD 263; 26/09=JD 270; 04/10=JD 279; 05/10=JD 280; 12/10=JD 286.

During the 5 week sampling period, total heterotrophic bacterial abundance and activity showed large fluctuation in the 0–150 m upper layer, and especially in relation to the deep chlorophyll maxima depths (Fig. 1). For example at 60 m depth, TBA varied from 14.8 to 2.5×10^5 cells ml^{-1} in relation to drastic changes in the Chla corresponding values found at the same depth (from 0.10 to 0.67 mg Chla m^{-2}). Previous investigations have shown that when bacterial biomass is limited solely by nutrients of phytoplankton origin, a strong correlation is observed between bacterial and phytoplankton biomass (Cole et al., 1988) and between bacterial abundance and production (Billen et al., 1990). Integrated 0–150 m depth values of bacterial biomass to autotrophic biomass ratios remained relatively constant and always >1 (Fig. 2). Such dominance of bacterial biomass relative to autotrophic biomass agrees with other results for the open Ocean (Ducklow and Carlson, 1992). These authors suggested that such situation occurs when phytoplankton biomass is low ($\text{Chla} < 1$ mg m^{-3}) and when bacteria are sustained by phytodetritus or by-products from grazing. For the entire studied period, Chla concentration was low (from 0.16 to 0.40 mg m^{-3} in the upper 150 m) and we found a significant correlation between 0–150 m integrated bacterial biomass and Chla concentration ($R=+0.44$, $p<0.05$, $n=54$) but this relation was more significant in the 40–80 m layer ($R=+0.67$, $p<0.01$, $n=54$) compared to 0–20 m layer ($R=+0.13$, $p<0.05$, $n=54$). In a same way, bacterial abun-

dance and bacterial production measured into the Chla maximum layer (40–80 m) were also strongly linked during the whole sampling period ($R=+0.65$, $p<0.01$, $n=162$), whereas no significant relation was observed in the surface mixed layer (0–20 m). Linear regression analysis calculated in the 40–80 m layer between integrated BB and Chla concentration ($Y=0.24X+5.25$, $R^2=0.65$, $n=54$) suggest that bacterioplankton living in the Chla maximum layer was regulated by C-compounds derived from phytoplankton activity. In addition, the slope <1 indicated that bacterial responses to resource availability tend to be attenuated by predation or viral pressure (Ducklow and Carlson, 1992; Dufour and Torretton 1996; Weinbauer et al., 2003). Nanoflagellates and ciliates populations have been previously observed in our study site and predation over bacterioplankton considered (Tanaka and Rassoulzadegan, 2002). Thus, in our conditions both bottom-up and top-down processes were controlling bacterial populations inhabiting the 0–150 m upper layer but, further grazing experiments are needed to estimate the grazing rates and the relative importance of bottom-up and top-down control in such conditions.

Investigations at hour time scales allowed us to enlighten the influence of diel and episodic events (coastal water intrusions and wind events) on total bacterial abundance and production (Fig. 3). By sampling every 6 h during 5 days within two periods at 9 days intervals, we did not found any diel periodicity on the 0–150 m integrated total bacterial abundance

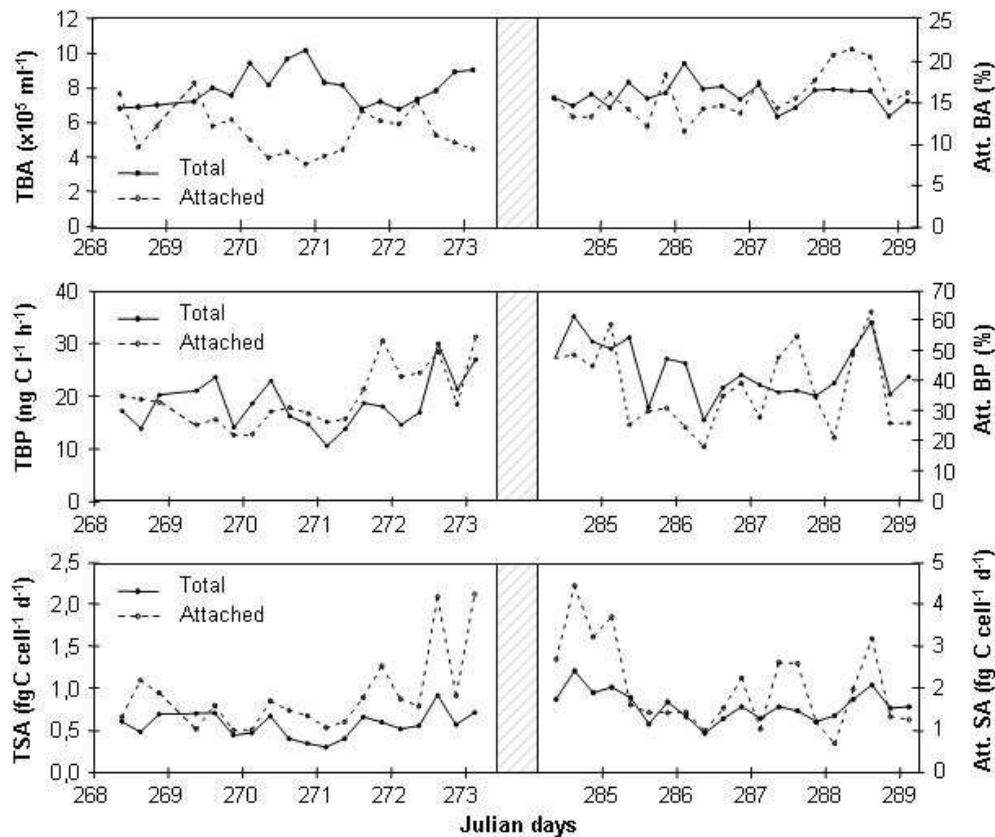


Fig. 6. Total vs. attached bacterial abundance, production and specific activity at 5 m depth. TBA=total bacterial abundance, TBP=total bacterial production, TSA=total specific activity, Att. BA=relative attached bacterial abundance on the TBA, Att. BP=relative attached bacterial production on the TBP, Att. SA=relative attached bacterial specific activity on the TSA.

and production. We did not expect such result since diel changes of bacterial production were observed in several open Ocean, including the NW Mediterranean Sea (Gasol et al., 1998; Kuipers et al., 2000; Winter et al., 2004). These results together with hypothesis mentioned above suggest that bacterial abundance and production was not directly linked to the diel changes in dissolved organic matter produced by phytoplankton (bottom-up control), thus reinforcing the role of top-down control by predation or viral lysis in our sampling conditions. At the DYFAMED station, diel changes in bacterial production was also observed during spring phytoplanktonic bloom (Ghiglione et al. 2007), but not during spring to summer transition (Van Wambeke et al., 2001). Such discrepancy can result from seasonal factors or meteorological conditions (rain, cloudy days, wind) that could superimpose their effects on diel variability and consequently, probably prevent us from obtaining reproducible diel patterns (Coffin et al., 1993). This is probably the case in our study, as shown by the influence of episodic events on the bacterial compartment during summer to autumn transition conditions. First, desalted water intrusions occurred twice during our sampling period as described by the P index, a proxy of salinity anomalies (Andersen et al., this issue¹). We found

a significant negative relationship between the 0–150 m integrated total bacterial abundance and production and the P index, especially during the first compared to the second intrusion ($R=-0.85$, $p<0.01$, $n=19$ compared to $R=-0.36$, $p<0.05$, $n=20$ for TBA, and $R=-0.47$, $p<0.05$, $n=19$ compared to $R=-0.18$, $p<0.05$, $n=20$ for TBP) (Fig. 3). These results suggest a proportional impact of salinity anomalies on total bacterial abundance and activity together with the magnitude of the low salinity water masses intrusion. The effect of wind events on total bacterial abundance and production was not detectable in our conditions. We did not find any significant correlation between total bacterial abundance and production neither following low magnitude (<10 nds) or stronger (>20 nds) wind events. During the second period of our hour scale sampling (JD 284–289), the strong wind created an important decrease of temperature and the upper-mixed layer moved down from 20 to 40 m depth (see Andersen et al., this issue¹). Several authors suggested that such wind events could result in changes in total bacterial abundance and production, as a result of an increase of NO_3^- availability, biomasses and particle fluxes in the upper layer (Haury et al., 1990; Marra et al., 1990; Kiorboe, 1993). Mediterranean surface waters being limited by P and

N during the stratified period, an increased NO_3^- availability could result in changes of bacterial abundance and production. Because of cruise logistics, the hour scale sampling was stopped three days after the stronger wind event, when the euphotic layer de-stratification begins. Thus, even if we did not find any influence of low magnitude wind events (<10 nds) on total bacterial abundance and production, our results suggest also that the response of the bacterial compartment may take more than three days to react to stronger wind events (>20 nds).

4.2 Temporal evolution of high nucleic acid content cells in the euphotic layer

In aquatic systems, it is possible to distinguish populations with high (HNA cells) and low nucleic acid content (LNA cells) by flow cytometry. In our study, the percentage of HNA cells to the total bacterial count (% HNA) varied from 44% to 59% (average= $49\pm 3\%$, $n=324$) and remained stable within the whole 0–150 m upper layer (Fig. 3). These values are in agreement with previous results found in various oceans (Sherr et al., 2006; Zubkov et al., 2006) and in the NW Mediterranean (Scharek and Latasa, 2007). The relative contribution of HNA bacteria to total abundance slightly fluctuated during our sampling period (CV=14%) but neither temporal trend nor depth significant pattern was observed. In our large set of samples ($n=324$), we found a significant and positive correlation between the percentage of HNA in the total bacterial population and the total bacterial production ($R=+0.75$, $p<0.01$) while a weaker relationship was observed between total bacterial abundance and production ($R=+0.59$, $p<0.01$). HNA cells are generally considered to represent active members of the bacterial community, whereas LNA may be dead or dying cells (Gasol et al., 1999; Lebaron et al., 2002; Servais et al., 2003). However, more recent studies question the use of HNA cell abundance as a proxy for activity in natural systems (Longnecker et al., 2005 and 2006; Sherr et al., 2006; Bouvier et al., 2007; Moran et al., 2007). Such discrepancies could be explained by the variability in the cytometric characteristics of HNA/LNA linked to different ecosystems (Bouvier et al., 2007). In our study the dichotomous view of HNA and LNA appears reasonable as all observations were made at the same location. As the HNA cells, were strongly correlated with total bacterial production, they can be considered as the active members of the bacterial community in our study site.

4.3 Dynamics of total bacterial abundance and production in the mesopelagic layer (150–1000)

As already reported, we found that total bacterial abundance and production linearly decreased with depth in the mesopelagic layer (Turley and Stutt, 2000; Tanaka and Rassoulzadegan, 2004). In parallel, clear changes in bacterial diversity by vertical zonation in three layers, above, in or

just below the chlorophyll maximum and deeper, that remained stable during the entire sampling period was also observed by Ghiglione et al. (2008). Compared to previous studies at the same JGOFS-DYFAMED station, total bacterial abundance and production values measured during the sampling period (0.53 to 1.83×10^5 cells ml^{-1} and 0.07 to 8.10 ng $\text{C l}^{-1} \text{h}^{-1}$ for total bacterial abundance and production, respectively; Fig. 4) are comparable to the previous studies driven in the mesopelagic layers of the NW Mediterranean Sea (Turley and Stutt, 2000; Harris et al., 2001; Tamburini et al., 2002; Tanaka and Rassoulzadegan, 2004). The magnitude of depth-dependent decrease of total bacterial abundance was relatively constant for our study period. The calculated slopes of log-log linear regression varied from -0.47 to -0.62 and the mean slope ($=-0.50\pm 0.07$, $n=12$) was slightly smaller compared to those previously described in the same zone ($=-0.66\pm 0.13$, $n=10$ in Tanaka and Rassoulzadegan, 2004). In contrast, the slopes relative to total bacterial production fluctuated for the one-month study period from -0.70 to -1.38 suggesting short-term variations in depth-dependent decrease of total bacterial production during the summer-autumn transient period. Moreover, compared to slopes previously described by Tanaka and Rassoulzadegan (2004) in their study based on seasonal sampling (mean slope= -1.15 ± 0.30 , $n=5$), our mean slope for TBP (mean slope= -0.73 ± 0.49 , $n=12$) shows no significant difference suggesting a low seasonal variation in the vertical distribution of bacterioplankton. In addition, the high variability of our TBP slopes can be due to episodic deep peaks of production observed between 400 to 750 m depths. Interestingly, such episodic deep peaks of bacterial production were already reported in the NW Mediterranean (Mistic and Fabiano, 2006). Given the mean CV of triplicate BP measurements (12.5%) these deep peaks show a real bacterial production. They suggest a rapid adaptation of bacteria to the episodic occurrence of organic matter probably originated from zooplankton metabolism (egestion of fecal material).

4.4 Temporal trends in total vs. particle-attached bacterial abundance and production in the 0–1000 m water column

We defined the terms “free-living” and “particle-attached” bacteria on the basis of a $0.8\text{-}\mu\text{m}$ -pore-size filter fractionation. A similar cut-off was used by several authors (Hollibaugh et al., 2000; Ghiglione et al., 2007) used also GF/C filters with a nominal pore size of $1.2\text{ }\mu\text{m}$ (Gasol and Morán, 1999). Microscopic observations showed that free-living bacteria retained on the $0.8\text{ }\mu\text{m}$ filters represented less of 5% of $>0.8\text{ }\mu\text{m}$ fraction. So, the over-estimation of attached bacteria was very low in the upper layer ($<1\%$ and 2% for BA and BP, respectively) and negligible in mesopelagic layers. In addition, to minimize the bacteria-particles dislodging and consequently the under-estimation of attached bacteria,

all filtrations were performed at a very low vacuum pressure. So, the size fractionation process used in our study can be considered as a clear-cut distinction between free-living and attached bacteria.

The contribution of attached bacteria to total bacterial abundance varied from 5 to 25% (average = $15 \pm 5\%$, $n=48$) in the first 60 m depth and remained in the same range during the 5 weeks studied period (Fig. 5). These values are consistent with previous reports on various pelagic environments (see review by Simon et al., 2002). We found a significant correlation between attached to total bacterial abundance ratio and Chl *a* concentration in the first 150 m depth ($R=+0.54$, $p<0.05$, $n=72$), suggesting a tight coupling between the vertical distribution of the attached bacteria and phytoplankton biomass. However, no significant relation was found between the contribution of attached bacterial abundance and diel or episodic events (Fig. 6). In the water column, the contribution of the attached fraction to total bacterial abundance decreased rapidly in the mesopelagic layer (<4%), as already described in other studies (see review by Simon et al., 2002).

The contribution of the attached fraction to the total bacterial production varied from 11 to 49% of the total bacterial production in the first 60 m depth with higher values at the end of the sampling period (Fig. 5). These results are in between the highest and lowest values reported in spring and summer 2003 (80% and 5%, respectively) at the same station (Ghiglione et al. 2007), but lower values (<30%) are usually reported in marine systems (see review by Simon et al., 2002). Such discrepancy could be due to differences in the studied systems (mainly in the size, nature and concentration of particles) or in protocols used to measure total bacterial production and to separate attached bacteria from free-living cells. In mesopelagic waters, the contribution of the attached fraction to the total bacterial production was generally low (less than 2%), except in some occasion (maximum of 18% at 500 m depth), but without any relation to exceptional increase of total bacterial production described above. At a shorter time scale, the contribution of attached fraction to the total bacterial production varied rapidly and significantly (from 18 to 63% within a few days) when samples were taken every 6 h. We did not find any significant relation between the contribution of attached fraction to the total bacterial production or the attached bacterial specific activity and diel or episodic events (Fig. 6). These results suggest that such indirect parameters were not sufficient to explain the rapid and chaotic variation of the total vs. attached bacterial production at hour time scale. These results reinforce the biogeochemical role of attached bacteria in the cycling of organic matter and reveal rapid and sporadic changes in their activity at hour time scale that influence drastically the total bacterial production in the NW Mediterranean Sea.

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