

Analysis of the ULF electromagnetic emission related to seismic activity, Teoloyucan geomagnetic station, 1998–2001

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Abstract. Results of ULF geomagnetic measurements at station Teoloyucan (Central Mexico, 99.11°35.735' W, 19.44°45.100' N, 2280 m height) in relation to seismic activity in the period 1998–2001 and their analysis are presented. Variations of spectral densities for horizontal and vertical components, polarization densities and spectrograms of magnetic field, their derivatives are analyzed as a part of traditional analysis in this study. Values of spectral density were calculated for 6 fixed frequencies $f=1, 3, 10, 30, 100$ and 300 mHz. Fractal characteristics of spectra were analyzed in the conception of SOC (Self-Organized Criticality). 2 nighttime intervals, 0–3 and 3–6 h by local time have been used to decrease the noise interference in raw data. In order to exclude the intervals with a high geomagnetic activity from analysis we referred to A_p indices, calculated for corresponding time intervals. The contribution of seismic events to geomagnetic emission was estimated by seismic index $k_s=10^{0.75M_s}/10D$, where M_s is the amplitude of the earthquake and D is the distance from its epicenter to the station.

1 Introduction

Different long-term observatories in Japan (Hayakawa, 1999), Russia (Molchanov, 2001, 2003), USA (Fraser-Smith, 1990, 1993), prove that the ULF EM band is very promising for monitoring of EM emission generated in the period preceding strong earthquakes and volcano eruptions. The ULF part of electromagnetic emission, generated by the earthquake, can be recorded at the Earth's surface by magnetometers without significant attenuation, if they are generated at typical earthquake nucleation depths (~ 10 km). Higher frequencies would have smaller skin depths and therefore, greater attenuation before reaching the Earth's surface. In addition to that, the ULF geomagnetic equipment does not

require complicated maintenance. Finally, the existing powerful methods of the data processing, such as spectral and polarization analysis (Hayakawa, et al., 2002), gradient metering and triangulation analysis (Ismaguilov et al., 2001, 2003), blind de-convolution technique (PCA analysis, Gotoh, Hayakawa et al., 2001, 2003; Kotsarenko, 2001), multifractal analysis (Smirnova et al., 2001) give a variety of application of mentioned methods of data processing and analysis of the results. The main aim of our study was to analyze the geomagnetic data observed at the station Teoloyucan during the period 1998–2001 and to find a statistical correlation between observed ULF EM emission and strong earthquakes occurred during that period of time.

2 Experiment and methodology

The geomagnetic data were collected at the station Teoloyucan (99.11°35.735' W, 19.44°45.100' N, 2280 m height) by 3-coordinate fluxgate magnetometer designed at UCLA, operated at 1 Hz sampling rate frequency, equipped by a GPS system for data synchronization (Fig. 1).

Seismic data have been classified by 3 criteria: magnitude of earthquake M_s , distance from epicenter of EQ D , and the seismic index (Molchanov et al., 2003) $k_s=10^{0.75M_s}/10D$, which was the main one (Fig. 2).

As well we referred to A_p indices, calculated for corresponding time intervals of seismic events in order to exclude from our analysis the events occurred during high geomagnetic activity. Finally, there were only 5 earthquakes with $k_s > 10$ occurring under quiet geomagnetic conditions. Their characteristics, the map of the epicenters and the station location are shown at the Table 1 and in a Fig. 3.

The obtained results are distributed in three main blocks according to their applied technique.



Fig. 1. Teoloyucan magnetic station and 3-component ULF magnetometer.

2.1 Spectral and polarization analysis

We have chosen 2 nighttime intervals, 0–3 and 3–6 h by local time to decrease the noise interference in raw data. Preliminary data processing includes filtering by running averaging with 2σ screening and de-spike procedure in order to exclude spike-like noise (instrumental errors and human activity) and de-trended procedure to relieve local disturbances from the slow diurnal variation.

We calculated spectral values S_Z and S_G (where $G^2 = H^2 + D^2$, G – total horizontal component), by PSD algorithm in 6 frequency sub-bands: $f=1, 3, 10, 30, 100$ and 300 mHz, with NFFT=10 800 (3 h); constructed polarization ratios $R = S_Z/S_G$, and their derivatives $R(\Delta f_1)/R(\Delta f_2)$; calculated 5400 daily values for S_Z and S_G and R (with the same NFFT=10 800) for a construction of the annual spectrograms; and combined all the mentioned results with corresponding variation of A_p with a purpose to isolate EM emissions generated by seismic sources from global geomagnetic disturbances such as magnetic storms.

2.2 The fractal analysis

This procedure requires the graph of daily PSD $S_{H,G} = S_{H,G}(f)$ to be presented with a logarithmic scale for both x- and y-axes. Mentioning the frequency dependence as $S \sim f^{-\beta}$, we calculated the exponent β by linear fit of the plotted dependence using least-squares. The shape of the PSD graphs was a quite different for the different 3-time intervals of data.

Some fractures and bends at definite frequencies appeared at most of noon and evening spectra, so the same 0–3 and 3–6 local time intervals have been used for this analysis.

2.3 The retrospective analysis of geomagnetic micropulsations in relation to seismic activity

We analyzed 5 seismic events, mentioned before, for a possible correlation with geomagnetic micropulsations, both continued and irregular ones, which occurred at the day of the earthquake, days before and after earthquake. Exception was the earthquake occurred on 30 September 1999, apparently none of the observed micropulsation has relation with geomagnetic activity and most probably they have been generated locally.

Table 1. The 5 strong and moderate earthquakes chosen for analysis.

Date, GMT time	Magnitude, M_s	Depth, km
15 June 1999, 20:42	7.0	69
30 September 1999, 16:31	7.5	16
9 August 2000, 11:41	6.5	16
19 May 2001, 23:21	6.5	20
7 October 2001, 21:39	6.1	10

We have used Matlab-based software (© Hannes Schwarzl, UCLA University) to filtrate data in corresponding frequency sub-band, related to proper geomagnetic pulsations, process and analyze them using powerful a interface.

As well we referred to the data of other geomagnetic stations in order to verify the character of the observed effects, global or local ones.

3 Description analysis of the obtained results

The analysis of the temporal dynamics of spectral densities $S_{Z,G}$, polarization ratios $R = S_Z/S_G$, their derivatives and annual spectrograms reveals following features: noticeable increase for the values of most of the mentioned characteristics in periods about 14–30 days before 3 of 5 earthquakes: moderate increase expressed at lower frequencies for the earthquake occurred on 15 June 1999, and high, multispike-like increase at all frequency sub-bands for earthquakes occurred on 30 September 1999 and 19 May 2001 (Fig. 4), especially clearly seen in the polarization ratio variations. Almost all the mentioned changes in spectra and polarization took place during quiet local geomagnetic activity (A_p low). These events have duration from 1 to 2 weeks, corresponding change of values are: up to tens nano Tesla in absolute values, up to tens of nT^2/Hz for spectral values and up to 10 time in polarization ratio. Such a feature can be interpreted as a possible ULF geomagnetic emission generated in the earthquake preparation site. There was no noticeable event observed for the other 2 earthquakes. All other changes in spectral variation can be explained by moderate and high magnetospheric activity.

Variations of fractal indices for the vertical (Z) and total horizontal (G) components and their ratio (Z/G), calculated for 3 years, reveal only 1 significant change in their dynamics, related to earthquake occurred on 19 May 2001. Noticeable noise-like behavior with slight increase in Z-component and some drops in G-component occurred in about 1 month before that earthquake with approximate duration about 20 days (Fig. 5).

The example of observed micropulsations related to the earthquake occurred on 15 June 1999 at 20:42 GMT is presented at Fig. 6. The complete description of the observed events one can see in the Table 2.

Table 2. The geomagnetic activity (Pc1-5, Pi1,2) during 5 strong and moderate earthquakes chosen for analysis.

DATE	15 June 1999	30 Sept. 1999	9 Aug. 2000	19 May 2001	7 Oct. 2001
DST	Magnetic silence during all the June	Magnetic storm 7 days before the EQ, low activity at the day before the EQ	Low activity the day before EQ and 4 h before the EQ	Low activity during the day of the EQ and about 2 days before the EQ	Low activity during the October
AE	Quiet magnetically during the EQ, hours before and after the EQ	Low activity during the day of the EQ at the auroral zone	Quiet magnetically during the EQ, several hours before and hours after the EQ	Low activity during the day of EQ and some hours after the EQ	Quiet magnetically during the EQ, hours before and after the EQ
K _p	Quiet period 5 days before and 10 days after EQ	Moderate magnetic activity at the day of the EQ and preceding days	Low activity at the day of the EQ and 3 days before	Low activity during the day of EQ and some days after the EQ	Low activity at the day of the EQ, 3 days before and 4 days after
Pc1-2	Activity of B_z the day before EQ, 4 h before, and 2 h after the EQ	No activity	Low activity of B_z 7 h. before the EQ and the day before at 09:00	No activity before the EQ and low activity 1 h after the EQ in component B_z	Activity 2 h before the EQ in B_x and B_z , and 4 h after
Pc3	Activity of B_z the day preceding EQ 4 h and 2 h after the EQ.	No activity	Low activity of B_z 7 h before the EQ and the day before at 09:00	No activity before the EQ and low activity 1 h after the EQ in component B_z	Activity 2 h. before the EQ in B_x y B_z
Pc4	Activity of B_z at the previous day at 4 h, the day of the EQ at 0 h and at 12 h, 2 h after the EQ	Activity of 3 components the day before at 9 h and 6 h before the EQ	Activity of B_z the day before at 9 h and 2 h before the EQ	No activity at days of EQ neither before and after EQ	Activity the day before the EQ in 3 components, several minutes after the EQ in B_z , and 6 h after the EQ
Pc5	Activity of 3 components at the previous day at 4 h and the day of the EQ at 12 h. 3 components and 2 h after activity of B_z	Activity of 3 components the day before at 9 h, high activity during the day of the EQ, activity of B_z the day after EQ	Activity of B_z the day before at 9 h, 7 h before the EQ and 1h after	Activity 14 h. Before the EQ - 3 components, 8 h. before - in B_z , and 1 day after - in B_z	Activity the day before the EQ at 4 h and 12 h in 3 components, 11 h after the EQ in B_z , 6 h after the EQ in B_z
Pi1	Activity of B_z at the previous day at 4 h and 2 h after the EQ	No activity	Activity of B_z the day before at 9 h, 7 h before the EQ	No activity at days of EQ neither before and after EQ	Activity at 2 h before the EQ and 4 h after in B_x and B_z
Pi2	Activity of B_z at the previous day at 4 h, the day of the EQ at 0 h and 12 h, 2 h after the EQ	Low activity of 3 components day before and 6 h before the EQ and the next day after the EQ	Activity of B_z the day before at 9 h, 7 h before the EQ	No clear activity	Activity during the EQ and several minutes after
Correlation with EQ	Positive	Low Positive	Positive	Slightly Positive	Positive

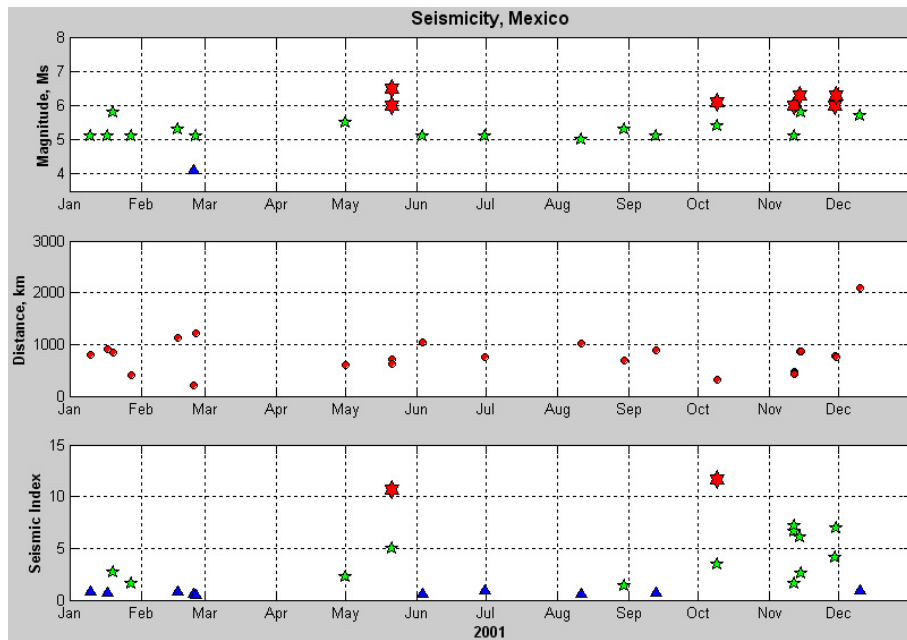


Fig. 2. Seismicity in Mexico in 2001. M_s are magnitudes of the EQs; distances from the epicenter D and seismic index k_s are calculated for Teoloyucan geomagnetic station.



Fig. 3. Map of epicenters (stars) of earthquakes discussed and Teoloyucan magnetic station (triangle).

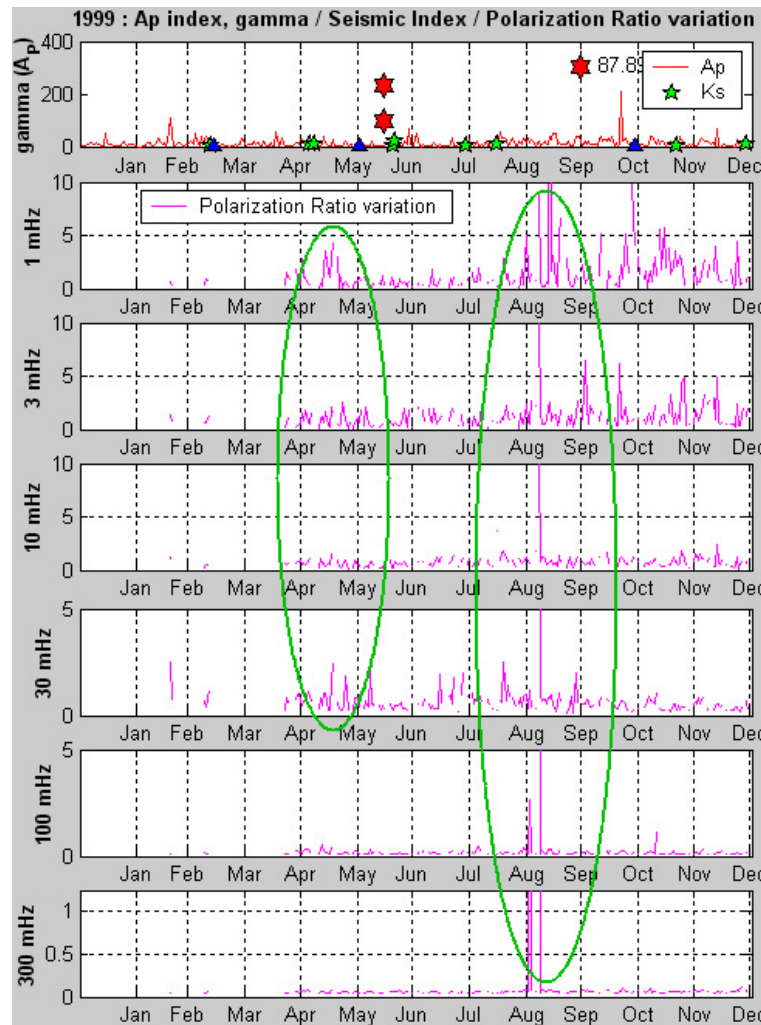


Fig. 4. Temporal variation of polarization ratio calculated for 6 frequency sub-bands, upper panel – seismic activity (k_S index) and geomagnetic activity (A_P index) dynamics, 1999.

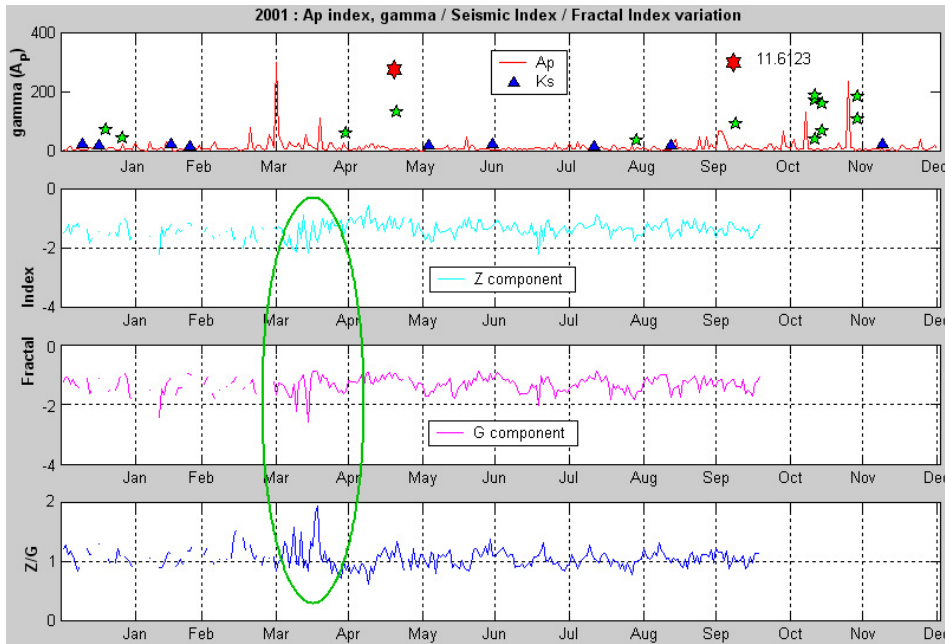


Fig. 5. Temporal variation of fractal index, upper panel – seismic activity (k_S index) and geomagnetic activity (A_P index) dynamics, 2001.

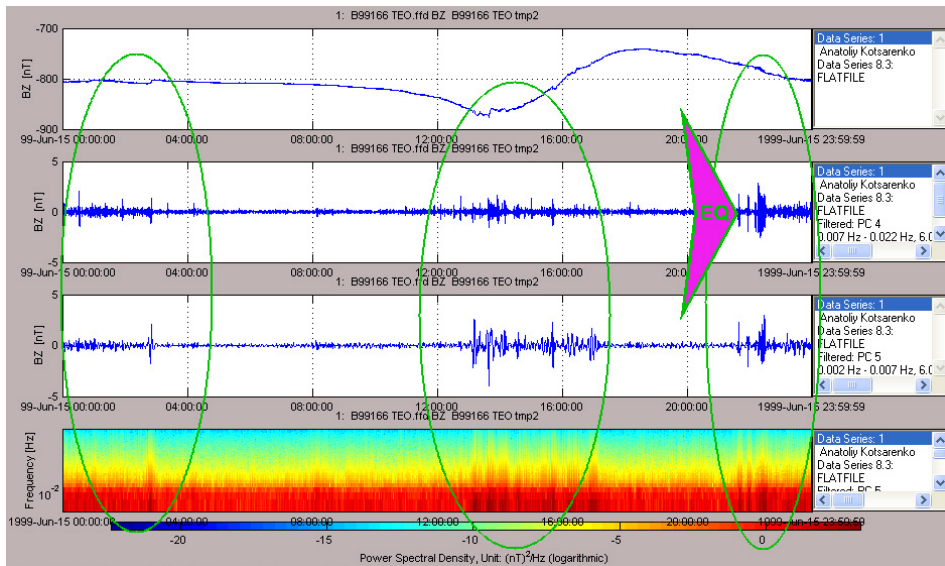


Fig. 6. Micropulsations, observed on 15 June 1999 in component Z. Upper panel – B_Z time series, second panel – PC4 sub-band, third panel – PC5 sub-band, lower panel – daily spectrogram for PC5. The earthquake time is marked with a violet mark.

Generally all the observed events can be classified in 3 blocks due to their occurrence time: pre-seismic (before the earthquake), co-seismic (at the exact time of the earthquake and during it) and post-seismic events observed at different sub-bands. Most of events, classified in the table, have no relation neither with global nor with local geomagnetic activity, which was proven by checking with DST , K_P and AE geomagnetic indices and by comparing obtained results with the results based on the data of another geomagnetic stations.

Moreover, observed micropulsations normally do not occur at all during the days with absent or low geomagnetic

activity. For these reasons we can testify the probable relation of the observed events with the effects of the preparation of the earthquakes.

4 Conclusions

In this paper we have studied geomagnetic data performing analysis in 3 different conceptions. The traditional one, such as construction of spectral long time series (S_Z , S_G , $R=S_Z/S_G$, their derivatives and spectrograms), calculated

for different ULF sub-bands; Fractal analysis and multi-analysis of geomagnetic micropulsation data corresponding to 5 strong and moderate earthquakes with magnitudes greater than 6.0, that occurred in Mexico during 1999–2001.

We found the following:

1. In the case of 3 of the mentioned events studied within the concept of traditional analysis, most characteristics show noticeable increase, taking place from 14 to 30 days in almost all frequencies, mainly at the lower ones.
2. For the fractal analysis, there was only one significant change with 20 days duration, that occurred in about one month prior to the corresponding earthquake.
3. Some micropulsations not associated to magnetospheric activity were present before, during and after some major earthquakes, some of them shown in all three components of the magnetic field.

We are aware that these results are still with the possibility of improvement with more extended network observations and therefore, with the possibility of applying other advanced techniques.

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