Assessing the Liquefaction Hazard in the Groningen Region of the Netherlands due to Induced Seismicity: Limitations of Existing Procedures and Development of a Groningen-Specific Framework

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Abstract The Groningen gas field is one of the largest in the world and has produced over 2000 billion m^3 of natural gas since the start of production in 1963. The first earthquakes linked to gas production in the Groningen field occurred in 1991, with the largest event to date being **M** 3.6. As a result, the field operator is leading an effort to quantify the seismic hazard and risk resulting from the gas production operations, including the assessment of liquefaction hazard. However, due to the unique characteristics of both the seismic hazard and the geological subsurface, particularly the unconsolidated sediments, direct application of existing liquefaction evaluation procedures is deemed inappropriate in Groningen. Specifically, the depth-stress reduction factor (r_d) and the Magnitude Scaling Factor (MSF) relationships inherent to existing variants of the simplified liquefaction evaluation procedure are considered unsuitable for use. Accordingly, efforts have first focused on developing a framework for evaluating the liquefaction potential of the region for magnitudes ranging from **M** 3.5 to 7.0. The limitations of existing liquefaction procedures for use in Groningen and the path being followed to overcome these shortcomings are presented in detail herein.

Keywords Liquefaction, liquefaction hazard, induced seismicity, Groningen gas field

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1 Introduction

The Groningen gas field is located in the northeastern region of the Netherlands and is one of the largest in the world. It has produced over 2000 billon m^3 of natural gas since the start of production in 1963. The first earthquakes linked to gas production in the Groningen field occurred in 1991, although earthquakes were linked to production at other gas fields in the region since 1986. To date the largest induced earthquake due to production at the Groningen field is the 2012 moment magnitude (**M**) 3.6 Huizinge event, and the largest recorded peak ground acceleration (PGA) is 0.11 g which was recorded during a more recent, smaller (local magnitude, M_L, 3.4) event. In response to concerns about the induced earthquakes, the field operator Nederlandse Aardolie Maatschappij (NAM) is leading an effort to quantify the seismic hazard and risk resulting from the gas production operations (Bourne et al. 2015, van Elk et al. 2017). In view of the widespread deposits of saturated sands in the region, the risk due to earthquake-induced liquefaction is being evaluated as part of this effort. Although an almost negligible contributor to earthquake fatalities, liquefaction triggering is an important threat to the built environment and in particular to infrastructure and lifelines (e.g., Bird and Bommer 2004).

Central to the liquefaction hazard/risk assessment of the Groningen field is the stress-based “simplified” liquefaction evaluation procedure, which is the most widely used approach to evaluate liquefaction potential worldwide. While most of the recently proposed variants of this procedure yield similar results for scenarios that are well represented in the liquefaction case history databases (e.g., Green et al. 2014), their predictions deviate, sometimes significantly, for other scenarios (e.g., low magnitude events; very shallow and very deep liquefiable layers; high fines content soils; medium dense to dense soils). These deviations result partly because existing variants of the simplified procedure are semi-empirical, hence they are apt for replicating existing data but lack proper extrapolation power. The empirical elements of existing procedures are derived from data from tectonic earthquakes in active shallow-crustal tectonic regimes such as California, Japan, and New Zealand. These conditions are different from those that of the Groningen field. Moreover, the geologic profiles/soil deposits in Groningen differ significantly from those used to develop the empirical aspects of the simplified procedure. As a result, the suitability of existing variants of the simplified procedure for direct use to evaluate liquefaction in

Groningen is questionable. Accordingly, prior to assessing the liquefaction hazard in Groningen, efforts have first focused on developing a framework for performing the assessment. This actually required a step backwards to develop an "unbiased" liquefaction triggering procedure for tectonic earthquakes, due to biases in relationships inherent to existing variants of the simplified procedure (e.g., Boulanger and Idriss 2014).

In the following sections, the shortcomings in current variants of the simplified procedures for use in Groningen are detailed. Then, the efforts to develop a new "unbiased" variant of the simplified liquefaction evaluation procedure are presented. An outline of how this procedure is being modified for use in Groningen is presented next, followed by a brief overview of how the liquefaction hazard of Groningen will be assessed.

2 Shortcoming in existing variants of the simplified liquefaction evaluation procedure for use in Groningen

2.1 Overview of the simplified procedure

As mentioned in the Introduction, the stress-based simplified liquefaction evaluation procedure is central to the approach adopted to assess the liquefaction hazard in the Groningen region. The word "simplified" in the procedure's title originated from the proposed use of a form of Newton's Second Law to compute cyclic shear stress (τ_c) imposed at a given depth in the soil profile, in lieu of performing numerical site response analyses (Whitman 1971; Seed and Idriss 1971). Inherent to this approach to computing the seismic demand is a depth-stress reduction factor (r_d) that accounts for the non-rigid response of the soil profile and a Magnitude Scaling Factor (MSF) that accounts for the effects of the shaking duration on liquefaction triggering. For historical reasons the duration of an **M** 7.5 earthquake is used as the reference for MSF.

Case histories compiled from post-earthquake investigations were categorized as either "liquefaction" or "no liquefaction" based on whether evidence of liquefaction was or was not observed. The seismic demand (or normalized Cyclic Stress Ratio: CSR*) for each of the case histories is plotted as a function of the corresponding normalized *in situ* test metric, e.g., Standard

Penetration Test (SPT): $N_{1,60cs}$; Cone Penetration Test (CPT): q_{c1Ncs}; or small strain shear-wave velocity (V_S): V_{S1}. In this plot, the "liquefaction" and "no liquefaction" cases tend to lie in two different regions of the graph. The "boundary" separating these two sets of case histories is referred to as the Cyclic Resistance Ratio ($CRR_{M7.5}$) and represents the capacity of the soil to resist liquefaction during an **M** 7.5 event. This boundary can be expressed as a function of the normalized *in situ* test metrics.

Consistent with the conventional definition for factor of safety (FS), the FS against liquefaction (FS_{liq}) is defined as the capacity of the soil to resist liquefaction divided by the seismic demand:

$$FS_{liq} = \frac{CRR_{M7.5}}{CSR^*} \quad (1)$$

The Dutch National Annex to the Eurocode for the seismic actions (i.e., NPR 9998 2017), recommends the use of the Idriss and Boulanger (2008) variant of the simplified liquefaction evaluation procedure, but allows other variants to be used if they are in line with the safety philosophy of the NPR 9998-2017. As a result, the Idriss and Boulanger (2008) variant and the updated variant (Boulanger and Idriss 2014) have been used in several liquefaction studies in Gronginen, resulting in predictions of potentially catastrophic liquefaction effects that have severe implications for buildings and for infrastructure such as dikes.

2.2 Depth-stress reduction factor: r_d

As stated above, r_d is an empirical factor that accounts for the non-rigid response of the soil profile. Both the Idriss and Boulanger (2008) and Boulanger and Idriss (2014) variants of the simplified liquefaction evaluation procedure use an r_d relationship that was developed by Idriss (1999). As shown in Figure 1, the Idriss (1999) r_d relationship is a function of earthquake magnitude and depth, with r_d being closer to one for larger magnitude events (note that r_d = 1 for all depths corresponds to the rigid response of the profile). This is because larger magnitude events have longer characteristic periods and, hence, ground motions with longer wave lengths. As a result, even a soft profile will tend to respond as a rigid body if the characteristic wave length of the ground motions is significantly longer than the overall thickness of the profile. Accordingly, the correlation between earthquake magnitude and the frequency content of the earthquake motions significantly influences the r_d relationship. This raises questions regarding the appropriateness of the Idriss (1999) relationship, which was developed using motions recorded during tectonic events, for evaluating liquefaction potential in Groningen where the seismic hazard is dominated by induced earthquakes having magnitudes less than **M** 5.

Another issue with the Idriss (1999) r_d relationship is that it tends to predict overly high CSR* values at depth in a soil profile for tectonic events. This bias is illustrated in Figure 1 and is pronounced for depths between ~3 to 20 m below the ground surface. As a result, when used to evaluate case histories to develop the $CRR_{M7.5}$ curves that are central to the procedure, the biased r_d relationship results in a biased positioning of the $CRR_{M7.5}$ curve. The significance of this issue is mitigated to some extent when the same r_d relationship used to develop the $CRR_{M7.5}$ curve is also used in forward analyses (i.e., the bias cancels out). However, this will not be the case if site/region-specific r_d relationships are developed and used in conjunction with a $CRR_{M7.5}$ curve that was developed using a “biased” r_d relationship.

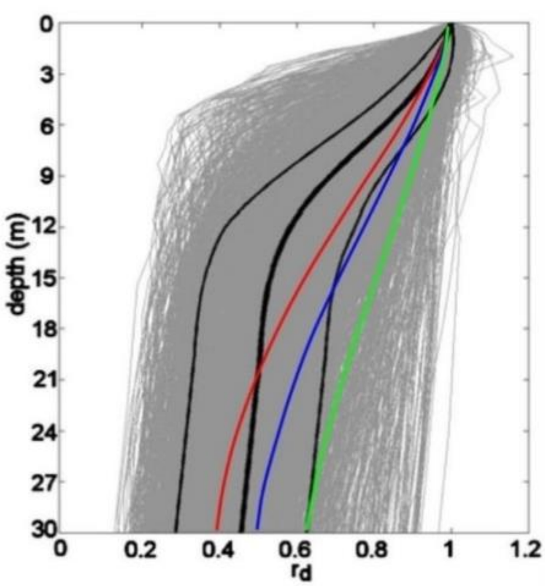

Fig. 1 The red, blue, and green lines were computed using the Idriss (1999) r_d relationship for **M** 5.5, **M** 6.5, and **M** 7.5 events, respectively. The grey lines were computed by Cetin (2000) from equivalent linear site response analyses performed using a matrix of 50 soil profiles and 40 motions. The black lines are the median (thick line) and median plus/minus one standard deviation

(thinner lines) for the Cetin (2000) analyses.

2.3 Magnitude Scaling Factor: MSF

As stated above, MSFs account for the influence of the strong motion duration on liquefaction triggering. MSFs have traditionally been computed as the ratio of the number of equivalent cycles for an **M** 7.5 event to that of a magnitude **M** event, raised to the power *b* [i.e., MSF=$(n_{eqM7.5}/n_{eqM})^{b}$]. Both the Idriss and Boulanger (2008) and Boulanger and Idriss (2014) procedures used the Seed et al. (1975) variant of the Palmgren-Miner (P-M) fatigue theory to compute $n_{eq\,M7.5}$ and n_{eqM} from earthquake motions recorded at the surface of soil profiles. Furthermore, they obtained the value of *b* from laboratory test data. The parameter *b* is the negative of the slope of a plot of log(CSR) versus log(N_{liq}), as shown in Figure 2; N_{liq} is the number of cycles required to trigger liquefaction in a soil specimen subjected to sinusoidal loading having an amplitude of CSR, typically determined using cyclic triaxial or cyclic simple shear tests.

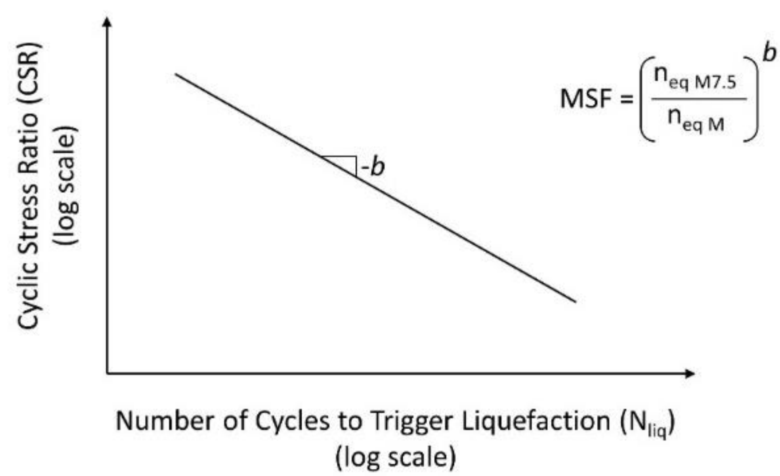

Fig. 2 Relationship between laboratory CSR vs. N_{liq} and MSF.

There are several shortcomings inherent to the approach used by Idriss and Boulanger (2008) and Boulanger and Idriss (2014) to compute the number of equivalent cycles and MSF. These include:

- Both the magnitude and uncertainty of n_{eq}, and hence MSF, are assumed to be constant with depth. However, Green and Terri (2005) have shown that n_{eq} can vary with depth in a given profile and Lasley et al. (2017) showed that while the median value for n_{eq} computed for a large number of soil profiles and ground motions is relatively constant with depth, the uncertainty in n_{eq} varies with depth.

- Pulses in the acceleration time history having an amplitude less than $0.3 \cdot a_{max}$ are assumed not to contribute to the triggering of liquefaction, and thus are not considered in the computation of n_{eq}. Using a relative amplitude criterion to exclude pulses is contrary to the known nonlinear response of soil which is governed by the absolute amplitude of the imposed load, among other factors. The use of a relative amplitude exclusion criterion with tectonic earthquake motions may inherently bias the resulting MSF.
- Each of the two horizontal components of ground motion is treated separately, inherently assuming that both components have similar characteristics. However, analysis of recorded motions has shown this is not always the case, particularly in the near fault region (e.g., Green et al. 2008; Carter et al. 2016). Groningen ground-motions recorded at short source-to-site distances often display pronounced polarization (Stafford et al. 2018).
- The *b* values used by Boulanger and Idriss (2014) were derived from several laboratory studies performed on various soils and it is uncertain whether all these studies used a consistent definition of liquefaction in interpreting the test data. As a result, the *b* values proposed by Boulanger and Idriss (2014) entail considerable uncertainty (Ulmer et al. 2018), with the proposed values not being in accord with those inherent to the shear modulus and damping degradation curves used in the equivalent linear site response analyses to develop the r_d correlations (a point elaborated upon subsequently).
- Recent studies have shown that the residuals of the amplitude and duration of earthquake ground motions are negatively correlated (e.g., Bradley 2011) and this feature is clearly observed in the Groningen data (Bommer et al. 2016). None of the MSF correlations developed to date, to include the one proposed by Boulanger and Idriss (2014), have considered this.

Some of the shortcomings listed above will be more significant to the Groningen liquefaction hazard assessment than others, but it is difficult to state *a priori* which ones these are. Furthermore, even for tectonic earthquakes the validation of MSF relationships is hindered by the limited magnitude range of case histories in the field liquefaction databases, with the majority of the cases being for events having magnitudes ranging from **M** 6.25 to **M** 7.75 (NRC 2016). Specific to the Groningen liquefaction hazard assessment, MSFs for small magnitude events are very important, particularly given that published MSF relationships vary by a factor of 3 for **M** 5.5 (Youd et al. 2001), with this factor increasing if the proposed MSF relations are extrapolated to lower magnitudes.

3 Removing bias from the simplified liquefaction evaluation procedure for tectonic earthquakes

3.1 Depth-stress reduction factor: r_d

A new relationship for r_d was developed by Lasley et al. (2016) using an approach similar to that used by Cetin (2000). Equivalent linear site response analyses were performed on 50 soil profiles compiled by Cetin (2000) that are representative of those in the liquefaction case history databases. However, Lasley et al. (2016) used a larger set of recorded input motions in their analyses than were available at the time of the Cetin (2000) study. Several functional forms for r_d were examined by Lasley et al. (2016) in regressing the results from the site response analyses, with the following form selected because of its simplicity and fit of the data (i.e., relatively low standard deviation of the regressed data):

$$r_d = (1-\alpha)exp\left(\frac{-z}{\beta}\right) + \alpha + \varepsilon_{r_d} \quad \text{(2a)}$$

where z is depth in meters, α is the limiting value of r_d at large depths and can range from 0 to 1, the variable β controls the curvature of the function at shallow depths, and ε_{r_d} is a zero-mean random variable with standard deviation σ_{r_d}. Expressions for α and β are:

$$\alpha = exp(-4.373 + 0.4491 \cdot \boldsymbol{M}) \quad \text{(2b)}$$

$$\beta = -20.11 + 6.247 \cdot \boldsymbol{M} \quad \text{(2c)}$$

and σ_{r_d} is defined as:

$$\sigma_{r_d} = \frac{0.1506}{[1+exp(-0.4975 \cdot z)]} \quad \text{(2d)}$$

Relative to the other r_d relationships inherent to commonly used variants of the simplified procedure, the Lasley et al. (2016) model was developed using more site response data and more rigorous regression analyses. So while all relationships inherently have some bias, a strong argument can be made that Lasley et al. (2016) has the least bias of commonly used relationships and was therefore adopted for use herein.

3.2 Magnitude Scaling Factor: MSF

Development of a MSF relationship that overcomes all the shortcomings listed above for the Idriss and Boulanger (2008) and Boulanger and Idriss (2014) relationships is not as straightforward as developing the new r_d relationships. The reason for this is that there are many more issues with existing MSFs than there are with the r_d relationships. As a result, a new approach needed to be used to compute MSFs, as opposed to implementing an existing approach using a more comprehensive dataset and a more rigorous regression analysis.

As mentioned previously and shown in Figure 2, MSFs are computed from equivalent number of cycles, n_{eq}. Well-established fatigue theories have been proposed for computing n_{eq} for materials having varying phenomenological behaviour; reviews of different approaches for computing n_{eq} are provided in Green and Terri (2005) and Hancock and Bommer (2005), among others. Developed specifically for use in evaluating liquefaction potential, the approach proposed by Green and Terri (2005) was selected for developing an n_{eq} relationship for the Groningen project. This approach is an alternative implementation of the P-M fatigue theory that better accounts for the nonlinear behaviour of the soil than the Seed et al. (1975) variant. In this approach, dissipated energy is explicitly used as the damage metric. n_{eq} is determined by equating the energy dissipated in a soil element subjected to an earthquake motion to the energy dissipated in the same soil element subjected to a sinusoidal motion of a given amplitude and a "duration" of n_{eq}. Dissipated energy was selected as the damage metric because it has been shown to correlate with excess pore pressure generation in saturated cohessionless soil samples subjected to undrained cyclic loading (e.g., Green et al. 2000; Polito et al. 2008). Furthermore, from a microscopic perspective, the energy is thought to be predominantly dissipated by the friction between sand grains as they move relative to each other as the soil skeleton breaks down, which is requisite for liquefaction triggering.

Conceptually, the Green and Terri (2005) approach for computing n_{eq} is shown in Figure 3. Stress and strain time-histories at various depths in the soil profile are obtained from a site response analysis. By integrating the variation of shear stress over shear strain, the cumulative dissipated energy per unit volume of soil can be computed (i.e., the cumulative area bounded by the shear stress-shear strain hysteresis loops). n_{eq} is then determined by dividing the cumulative dissipated energy for the entire earthquake motion by the energy dissipated in one equivalent cycle. For historical reasons, the shear stress amplitude of the equivalent cycle (τ_{avg}) is taken as $0.65 \cdot \tau_{max}$ (where τ_{max} is the maximum induced cyclic shear stress, τ_c, at a given depth), and the dissipated energy associated with the equivalent cycle is determined from the constitutive model used in the site response analysis.

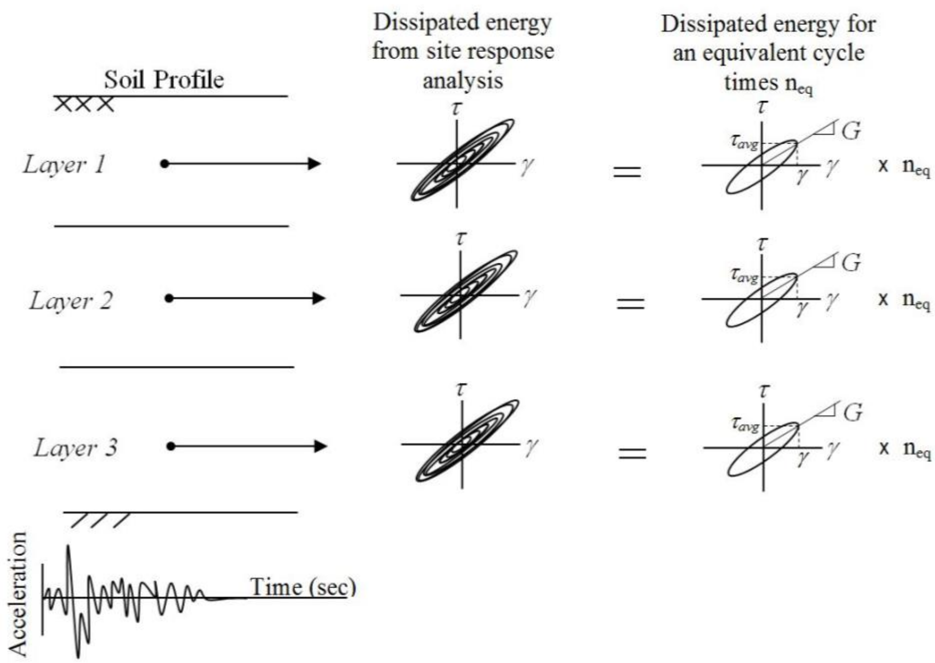

Fig. 3 Illustration of the proposed procedure to compute n_{eq}. In this procedure, the energy dissipated in a layer of soil, as computed from a site response analysis, is equated to the energy dissipated in an equivalent cycle of loading multiplied by n_{eq}.

As noted above, one of the shortcomings of the Seed et al. (1975) variant of the P-M fatigue theory is the way in which multi-directional shaking is taken into account. Specifically, each of the two horizontal components of ground motion is treated separately, inherently assuming that both components have similar characteristics. However, analysis of recorded motions has shown this is not always the case, particularly in the near fault region (e.g., Green et al. 2008; Carter et al. 2016). In contrast, Green and Terri (2005) accounted for multi-directional shaking by performing separate site response analyses for each horizontal component in a pair of motions, adding the energy dissipated at the respective depths for each component of motion, and setting the amplitude of the equivalent cycle as 0.65 times the geometric mean of the maximum shear stresses experienced at a given depth. This approach is referred to as "Approach 2" in Lasely et al. (2017) and is used herein because it better accounts for differences in the characteristics in the two horizontal components of motion.

Lasley et al. (2017) implemented the Green and Terri (2005) approach for computing n_{eq} using the same motions and profiles used by Lasley et al. (2016) to develop their r_d relationship. Their proposed n_{eq} relationship is:

$$\ln(n_{eq}) = 0.4605 - 0.4082 \cdot \ln\left(\frac{a_{max}}{g}\right) + 0.2332 \cdot \boldsymbol{M} + \varepsilon_{Total} \qquad (3a)$$

where a_{max} is in units of g and ε_{Total} is a zero-mean random variable with standard deviation σ_{Total} given by:

$$\sigma_{Total}(z) = \max\left[0.5399 - \frac{z}{26.4}(0.5399 - 0.4626), 0.4626\right] \qquad (3b)$$

where z is depth in meters. The dependency of n_{eq} on a_{max} in Eq. 3 was chosen because of the observed negative correlation of strong ground-motion duration with a_{max} (e.g., Bradley 2011). Also, the functional form of this correlation is not an impediment to implementation because the simplified liquefaction evaluation procedures require both the magnitude (for MSFs and r_d) and a_{max} as input variables.

The *b* value that is needed to relate n_{eq} to MSFs (e.g., Figure 2) can also be determined from the constitutive model used in the site response analysis, by assuming that the CSR vs. N_{liq} curve shown in Figure 2 is a contour of constant dissipated energy (Figure 4). In Figure 4, the dissipated energy for a **M** 7.5 earthquake, $\Delta W_{M7.5}$, is computed using:

$$\Delta W_{M7.5} = \frac{2\pi \cdot D_\gamma \cdot \tau_c^2}{G_{max} \cdot \left(\frac{G}{G_{max}}\right)_\gamma} \cdot n_{eq\ M7.5} \tag{4}$$

where D_γ is the damping ratio for the induced shear strain γ, τ_c is the cyclic shear stress and *G* is the secant shear modulus. This equation is based on the assumption that the soil can be modelled as a visco-elastic material, consistent with the assumption inherent to the equivalent linear site response algorithm. For liquefaction evaluations, τ_c used to compute $\Delta W_{M7.5}$ can be determined from the $CRR_{M7.5}$ curve from the simplified liquefaction evaluation procedure (e.g., Boulanger and Idriss 2014). Accordingly, the computed CSR vs. N_{liq} curve corresponds to a soil having a given q_{c1Ncs} and confined at an initial effective overburden stress (σ'_{vo}) (i.e., $\tau_c = CRR_{M7.5} \times \sigma'_{vo}$); the small strain shear modulus (G_{max}) for the soil should be consistent with the penetration resistance used to determine $CRR_{M7.5}$. The damping (D_γ) and the degraded secant shear modulus, $G_{max} \cdot (G/G_{max})_\gamma$, values in Eq. (4) are commensurate with the induced shear strain (γ) in the soil and can be determined iteratively from the shear modulus and damping degradation curves used to model the soil response (e.g., Darendeli and Stokoe 2001). Once the value of $\Delta W_{M7.5}$ is determined, a contour of constant dissipated energy can be computed for different amplitudes of loading by simply computing the number of cycles for the assumed loading amplitude required for the dissipated energy to equal $\Delta W_{M7.5}$. The parameter *b* is assumed equal to the negative of the slope of the contour of constant dissipated energy. The assumption that the CSR vs. N_{liq} curve is a contour of constant dissipated energy inherently implies that the energy dissipated in a given element of soil at the point of liquefaction triggering is unique and independent of the imposed loading characteristics. Several studies have shown that this is a reasonable assumption (e.g., Kokusho and Kaneko 2014; Polito et al. 2013).

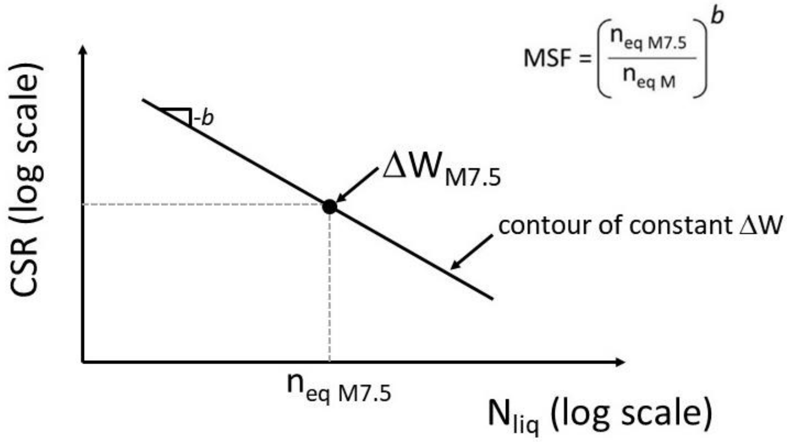

Fig. 4 A CSR vs. N_{liq} curve can be computed from shear modulus and damping degradation curves assuming the curve is a contour of constant dissipated energy. $\Delta W_{M7.5}$ can be computed using Eq. (4) and the remaining portions of the curve can be computed for different amplitudes of loading by simply computing the number of cycles for the assumed loading amplitude required for the dissipated energy to equal $\Delta W_{M7.5}$.

The degradation curves proposed Darendeli and Stokoe (2001) were used herein to determine the *b* values following the procedure illustrated in Figure 4 for a range of effective confining stresses and soil densities, with the resulting values ranging from 0.33 to 0.35. However, b = 0.34 for the vast majority of the confining stress-density combinations considered and was thus used herein to compute MSFs from n_{eq}. Additionally, b = 0.34 is consistent with laboratory curves developed from high-quality undisturbed samples obtained by freezing (Yoshimi et al. 1984). Accordingly, MSFs herein are computed as:

$$MSF = \left(\frac{n_{eq\,M7.5}}{n_{eq\,M}}\right)^b = \left(\frac{14}{n_{eq\,M}}\right)^{0.34} \leq 2.02 \quad (5a)$$

$$\sigma_{\ln(MSF)} = b \cdot \sigma_{\ln(\mathrm{n_{eq\,M}})} = 0.34 \cdot \sigma_{\ln(\mathrm{n_{eq\,M}})} \quad (5b)$$

where $\sigma_{\ln(MSF)}$ is a first order approximation for the standard deviation of the natural log of the MSF, and $n_{eq\,M}$ and $n_{eq\,M7.5}$ are computed using Eq. (3).

To compute $n_{eq\,M7.5}$ using Eq. (3), **M** is set to 7.5 and a corresponding value for a_{max} needs to be assumed (i.e., $a_{max7.5}$). The value of $a_{max7.5}$ was determined by computing the average a_{max} for the case histories in the Boulanger and Idriss (2014) SPT and CPT liquefaction case history databases ranging in magnitude from **M** 7.4 to 7.6. The average a_{max} for the 116 case histories that fell within this magnitude range was ~0.35 g. Using this value for $a_{max7.5}$, $n_{eq\ M7.5}$ was computed to be ~14. This value is similar to that determined by Seed et al. (1975), i.e., $n_{eq\ M7.5}$ = 15. However, the value reported by Seed et al. (1975) represents the average for two horizontal components of motion, while the value computed herein represents the combined influence of both components of motion (Approach 2, Lasley et al. 2017). As a result, the value computed herein is approximately half of that computed by Seed et al. (1975). This difference is due both to the significantly larger ground motion database used by Lasley et al. (2017) to develop Eq. (3), where the motions used by Lasley et al. (2017) represented a broader range of magnitudes and site-to-source distances compared to those used by Seed et al. (1975), and to the differences in the approaches used to compute n_{eq}. However, both of these differences also influence the denominator in Eq. (5a), which minimizes their influence on the resulting MSF. The upper limit on the MSF (i.e., 2.02) corresponds to a scenario where the earthquake motions consist of a single shear stress pulse in one of the horizontal components of motion. A plot of Eq. (5a) is shown in Figure 5 for magnitudes ranging from **M** 5.0 to 8.5 and a_{max} ranging from 0.1 to 1.0 g.

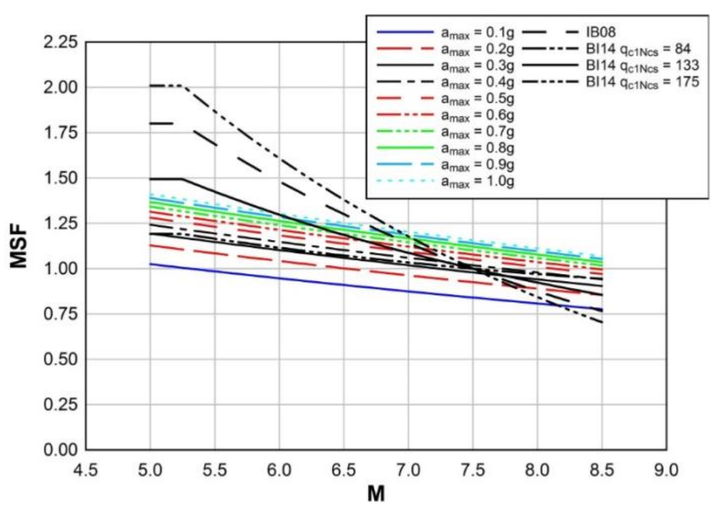

Fig. 5 For a given magnitude earthquake, MSF developed herein increases as a_{max} increases. Also, for comparison, the MSFs proposed by Idriss and Boulanger (2008) (IB08) and Boulanger and Idriss (2014) (BI14) are also shown.

Figure 5 also shows a comparison of the MSF developed herein with those proposed by Idriss and Boulanger (2008) and Boulanger and Idriss (2014), where the latter is shown for q_{c1Ncs} = 84, 133, and 175 atm. As may be observed from this figure, for a given value of a_{max} the MSF developed herein has about the same dependency on magnitude as the MSF proposed by Boulanger and Idriss (2014) for q_{c1Ncs} = 84 atm (i.e., medium dense sand). However, the difference between the two is that the former is a function of a_{max}, with MSF for a given magnitude increasing as a_{max} increases.

3.3 "Unbiased" $CRR_{M7.5}$ curve

The Lasley et al. (2016) r_d relationship and the MSF relationship developed herein were used to reanalyse the CPT liquefaction case history database compiled by Boulanger and Idriss (2014); all other parameters/relationships used to analyse the case history data were the same as those used by Boulanger and Idriss (2014). These case histories were then used to regress a new "unbiased" deterministic liquefaction triggering curve (i.e., $CRR_{M7.5}$ curve), which is shown in Figure 6. This curve approximately corresponds to a probability of liquefaction [P(liq)] of 35% (total uncertainty) and is given by:

$$CRR_{M7.5} = exp\left\{\left(\frac{q_{c1Ncs}}{113}\right) + \left(\frac{q_{c1Ncs}}{1000}\right)^2 - \left(\frac{q_{c1Ncs}}{140}\right)^3 + \left(\frac{q_{c1Ncs}}{137}\right)^4 - 2.8118706\right\} \leq 0.6 \qquad (6)$$

where q_{c1Ncs} is computed using the procedure outlined in Boulanger and Idriss (2014).

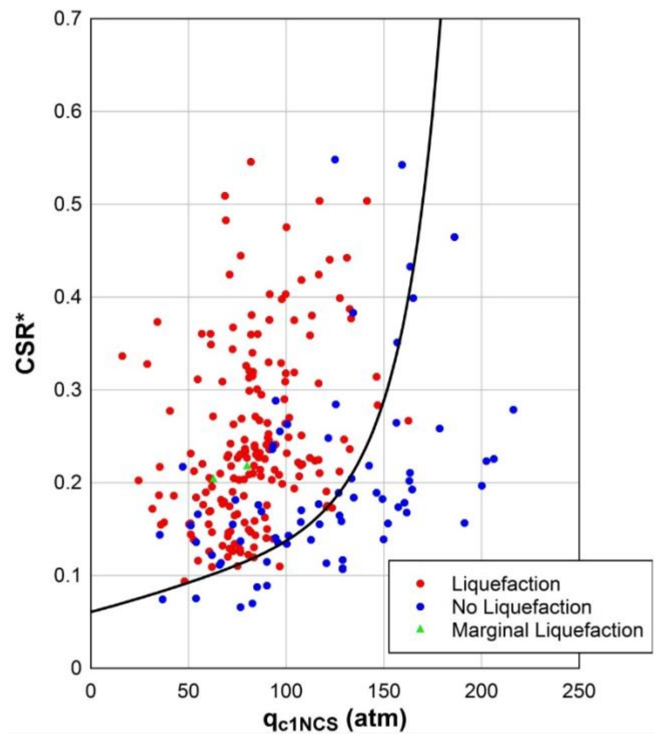

Fig. 6 "Unbiased" deterministic $CRR_{M7.5}$ curve regressed from liquefaction case history data from Boulanger and Idriss (2014) that were reanalysed using Lasley et al. (2016) r_d relationship and MSF developed herein.

4 Assessment of liquefaction hazard in Groningen

To determine whether a Groningen-wide liquefaction hazard assessment is warranted, a liquefaction hazard pilot study is being performed first, wherein the study area was selected to simultaneously satisfy three criteria: (a) proximity to the region of highest shaking hazard; (b) sampling of areas with sand deposits that are thick, shallow, young, and loose; and (c) sampling of multiple site-response zones used in developing the Groningen-specific ground-motion model (Rodriguez-Marek et al. 2017). The location of the pilot study area is shown in Figure 7, along with the cumulative thicknesses of the Holocene sand deposits that comprise the Naaldwijk formation which is considered to have the highest liquefaction potential in the region (Korff et al. 2017). However, before the liquefaction pilot study can be performed, Groningen-specific r_d and MSF relationships must be developed following the approaches used by Lasley et al. (2016, 2017) and presented above. The soil/geologic profiles and ground motions used to develop the Groningen-specific relationships are detailed below.

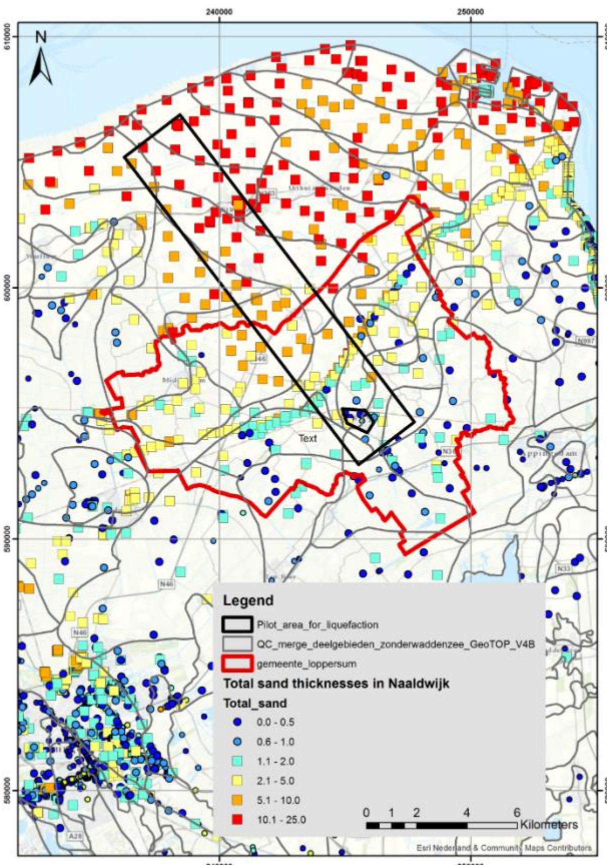

Fig. 7 Location of the liquefaction pilot study area across the Groningen gas field. Also shown are the cumulative thicknesses of the Holocene sand deposits that comprise the Naaldwijk formation.

4.1 Groningen-specific r_d and MSF relationships

The geological setting of Groningen, including detailed cross sections, is described in Kruiver et al. (2017a), and the velocity model from the selected reference rock horizon (at ~ 800 m depth) to the ground surface is described in detail by Kruiver et al. (2017b). An example of the resulting V_s profiles is shown in Figure 8. The unit weights of the strata in the profiles are also needed for the site response analyses. Towards this end, the assignment of unit weight is based on representative values for stratigraphic lithological units derived from CPTs using Lunne et al. (1997). For some of the deeper formations, the density is assumed to be constant, consistent with the borehole logs from two deep boreholes (Kruiver et al., 2017a, b).

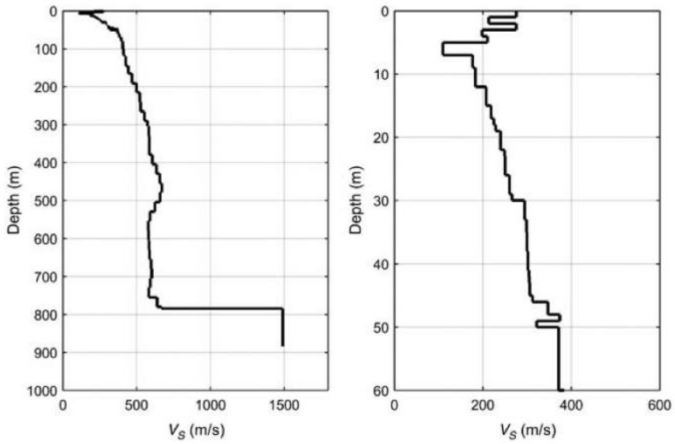

Fig. 8 Sample V_S profile at the location of one of the many ground-motion recording stations in the field. The plot on the left is the full profile down to reference rock horizon (depth of ~800 m), and the plot on the right is an enlarged view of the upper 60 m of the profile. (Rodriguez-Marek et al. 2017)

The software SMSIM (Boore 2005, version 16/12/2009) was used in conjunction with the Groningen-specific model parameters to generate motions at the reference horizon (Bommer et al. 2017) for magnitudes ranging from **M** 3.5 to 7.0 and epicentral distances ranging from 0.1 to 60 km. The lower bound was chosen on the basis of no liquefaction having been observed in the field to date and to explore the full range of potential triggering events, despite the fact that globally there is no reliable evidence of liquefaction triggering by earthquakes smaller than **M** 4.5 (Green and Bommer 2018). The upper value in the maximum magnitude distribution is **M** 7.25 as determined by an expert panel (Bommer and van Elk 2017).

Once developed, the Groningen-specific r_d and MSF relationships can be used in conjunction with the $CRR_{M7.5}$ curve shown in Figure 6 to compute the FS_{liq} at depth in profiles in Groningen subjected to induced earthquake motions. The computation of liquefaction hazard curves that will be used to determine whether the hazard due to liquefaction is significant enough to require the consequences from liquefaction to be assessed is discussed next.

4.2 Planned output from the liquefaction hazard study

The liquefaction hazard will be calculated using a Monte Carlo method (Bourne et al. 2015) wherein probability distributions for activity rates (Bourne and Oates 2017), event locations and magnitudes, and resulting ground motions will be sampled such that the simulated future seismic hazard is consistent with historical seismic and reservoir compaction datasets. For each event scenario, the developed Groningen-specific relationships will be used to compute the FS_{liq} as a function of depth for ~100 profiles across the pilot study area.

The "Ishihara inspired LPI" (LPI_{ish}) framework will be used to relate computed FS_{liq} to the predicted the severity of surficial liquefaction manifestation, which has been shown to correlate to liquefaction damage potential for level ground sites. The LPI_{ish} framework was proposed by (Maurer et al. 2015a) and is a conceptual and mathematical merger of the Ishihara (1985) H_1-H_2 chart and Liquefaction Potential Index (LPI) framework (Iwasaki et al. 1978). The most notable differences between the original LPI and LPI_{ish} frameworks are that the latter better accounts for the influence of the non-liquefiable crust on the severity of surficial liquefaction manifestations (Green et al. 2018) and more appropriately weights the contribution of shallower liquefied layers to surficial manifestations (van Ballegooy et al. 2014). The LPI_{ish} framework was chosen for this study because it has been shown to yield more accurate predictions of the severity of surficial liquefaction manifestations than competing indices (Maurer et al. 2015a, b): LPI (Iwasaki et al. 1978) and LSN (van Ballegooy et al. 2014).

The output from the liquefaction pilot study will be liquefaction hazard curves for the ~100 sites in the study area, where the hazard curves show the annual frequency of exceedance (AFE) of varying LPI_{ish} values for a site. Consistent with the requirements of NPR 9998-2017 (NPR 9998 2017), which was specifically for the Groningen field, LPI_{ish} values corresponding to an AFE of $\sim 4\times10^{-4}$ (or a 2475-year return period) will be of interest. The results from this pilot study will differ from previous liquefaction studies performed for Groningen, where liquefaction was evaluated in previous studies for earthquake scenarios (i.e., ground motions and magnitudes) corresponding to a given return period (i.e., a "pseudo-probabilistic" approach).

The optimal LPI_{ish} thresholds corresponding to different severities of surficial liquefaction manifestations are dependent on the liquefaction triggering procedure used to compute FS_{liq} and the characteristics of the profile. However, without liquefaction case history data to develop Groningen-specific thresholds, the thresholds proposed by Iwasaki et al. (1978) will be conservatively (Maurer et al. 2015c) used in the pilot study with the LPI_{ish} framework (i.e., LPI_{ish} < 5: none to minor surficial liquefaction manifestations are predicted; LPI_{ish} > 15: severe surficial liquefaction manifestations are predicted).

5 Discussion and conclusions

The presence of saturated loose deposits of young sands in the Groningen field region creates the necessity to assess the potential for liquefaction triggering by the earthquakes being induced by the gas production as an integral component of the seismic risk analysis. The application of liquefaction hazard assessment procedures calibrated for larger-magnitude tectonic earthquakes in other regions has resulted in predictions of potentially catastrophic liquefaction effects, with severe implications for buildings and for infrastructure such as dikes. Despite the fact these estimates, sometimes associated with earthquake scenarios only fractionally greater than the lower bound for events that have been observed globally to trigger liquefaction that poses a threat to the built environment (Green and Bommer 2018), the dissemination of such results has raised great concern regarding liquefaction hazard in Groningen.

Due to the unique characteristics of both the seismic hazard and the geologic profiles/soil deposits in Groningen, direct application of existing variants of the simplified liquefaction evaluation procedure is deemed inappropriate for assessing the liquefaction hazard of the region, including the Idriss and Boulanger (2008) procedure recommended in the NPR 9998-2017 and the updated variant, Boulanger and Idriss (2014). Accordingly, efforts were first focused on re-analyzing the liquefaction case histories that were compiled for natural earthquakes to remove bias in their interpretation. Towards this end, new a depth-stress reduction factor (r_d) and number of equivalent cycles (n_{eq})/magnitude scaling factor (MSF) relationships for shallow crustal active tectonic regimes were developed and used in the reanalysis of the cone penetration test (CPT) "liquefaction" and "no liquefaction" case histories compiled by Boulanger and Idriss (2014). These case histories were then used to regress a new "unbiased" deterministic liquefaction triggering curve (or cyclic resistance ratio curve: $CRR_{M7.5}$). The "unbiased" procedure can be readily adapted to evaluate liquefaction potential in regions with unique profiles and/or ground motions, such as Groningen. This is being achieved by using similar approaches to those employed to develop the new r_d and MSF relationships for tectonic earthquakes (Lasley et al. 2016, 2017) to develop Groningen-specific relationships using motions and soil profiles characteristic to Groningen.

The liquefaction hazard will be calculated using a Monte Carlo method wherein probability distributions for activity rates, event locations and magnitudes, and resulting ground motions are sampled such that the simulated future seismic hazard is consistent with historical seismic and reservoir compaction datasets for events having magnitudes ranging from **M** 3.5 to 7.0. For each event scenario, the Groningen-specific relationships will be used to compute the factor of safety (FS_{liq}) against liquefaction as a function of depth for ~100 profiles across the liquefaction pilot study area and corresponding Ishihara inspired Liquefaction Potential Index (LPI_{ish}) (Maurer et al. 2015a) hazard curves are being computed for each profile. The hazard curves specify the return periods of different severities of surficial liquefaction manifestations, with the severities corresponding to a return period of 2475 years being of interest per the NPR 9998-2017. This is in marked contrast to previous liquefaction hazard studies performed for Groningen that used a pseudo-probabilistic approach, where the FS_{liq} or LPI is computed for an earthquake scenario (i.e., ground motions and magnitude) corresponding to a given return period.

The framework of the liquefaction hazard pilot study is in complete accord with the safety philosophy of the NPR 9998-2017 and is particularly well suited to the specific nature of the time-dependent induced seismicity being considered. The results of the study will form the basis on which decisions will be made regarding the need for implementing mitigation measures. The liquefaction hazard study is benefiting significantly from the broader efforts to assess the regional seismic hazard in Groningen, to include the development of a regional velocity model (Kruiver et al. 2017a, b), site response model (Rodriguez-Marek et al. 2017), and ground-motion prediction model (Bommer et al. 2017).

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