

Experimental Evaluation of Flux Footprint Models

K. Heidbach, M. Mauder, H.P. Schmid

Karlsruhe Institute of Technology, Institute of Meteorology and Climate Research (IMK-IFU), Germany

Contact: katja.heidbach@kit.edu

1. Motivation

In most footprint studies computationally inexpensive models are applied

BUT:

Assumption of horizontally homogeneous turbulence can usually not be fulfilled in reality

Increased uncertainties

Evaluation at real-world flux sites

Up to now, there are no experiments that evaluate the 2D flux footprint directly

2. TERENO-Research Site "Graswang", Bavaria

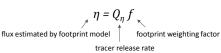


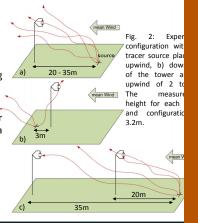
Fig. 1: a) Grassland site in Graswang, southern Germany (47.57° N 11.03° E: 870 m a.s.l.) with b) the CH₄ flux measurement system and c) the tracer gas diffuser of ~1 m2 size, c) frequency distribution of wind direction, July-Oct 2013.

- Located on a flat valley bottom (~1 km wide). flanked by steep sides
- Distinct mountain-valley breeze
- Natural flux of methane almost zero

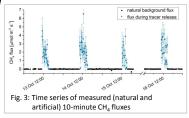
3. Methods

- Eddy covariance measurements (CSAT3, LI7700, LI7500)
- Surface source of ~1 m² size (Fig. 1c)
- Tracer gas: CH₄
- Release rate: 6-8 | min⁻¹ continuously over one averaging period (10 minutes)
- 3 different experiment configurations (Fig. 2)
- Evaluation of 3 footprint models: Kormann and Meixner (2001), Hsieh et al. (2000) and a parameterization of a backward Lagrangian footprint model (Kljun et al., 2004)
- Flux estimated by the model is determined with





4. Results – Experiments with upwind source



- Artificial flux in most cases ~100 times larger than the natural flux (Fig. 3)
- → Surface source of just 1 m² is a good possibility to precisely validate the 2D footprint
- Daytime-experiments
- → Mostly unstable conditions
- Hsieh et al. (2000) matches observations best (mode of frequency distribution closest to 1, Fig. 4 left)
- underestimate the maximum of the footprint (Fig. 4

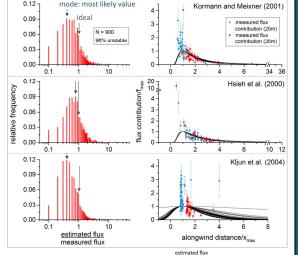
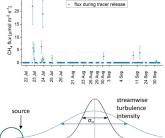


Fig. 4: Left: Frequency distribution of the ratio $\frac{\text{estimated nux}}{\text{measured flux}}$, data of experiment configurations a+c are included. Right: Measured and estimated flux contributions standardized with footprint maximum as a function of along-wind distance standardized with distance of footprint maximum for 3 different footprint models; only data of experiment configuration c are shown. Vertical bars denote the turbulence sampling error estimated following Finkelstein and Sims (2001)

5. Results – Experiments with downwind source

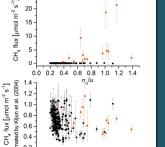
- Flux contribution from downwind source is measurable only occasionally (Fig. 5).
- Downwind contribution depends on streamwise turbulence intensity σ_{\parallel}/u (Fig. 6+7a).
- Kormann and Meixner (2001) and Hsieh et al. (2000) do not consider downwind contribution.
- Kljun et al. (2004) estimates a downwind contribution for any time period, even when along-wind turbulence intensity is low (Fig. 7b+c).



advection

Fig 5: Discontinuous series of measured 10-minute CH₄ fluxes during periods of tracer release (downwind source).

Fig. 6: The effect of streamwise turbulence intensity on downwind contribution with low (black) and high (blue)



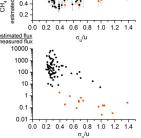


Fig. 7: a) measured 10-minute CH₄ fluxes, b) CH₄ fluxes estimated by Kljun et al. (2004) and c) model performance of Kljun et al (2004) as a function of σ₁₁/u.

6. Summary

- Tracer experiments aimed assessing the applicability utility of commonly u footprint models at observation conditions.
- Overall, the three evalua models match observati roughly, but all und estimate the flux.
- found a measura contribution to the flux fi downwind sou depending on streamy turbulence intensity.
- Downwind contribution occ only intermittently and continuously.
- The downwind estimate of the Kljun et (2004) model needs to optimized.







Finkelstein, P. L., and Sims, P. F., 2001, Sampling error in eddy correlation flux measurements, Journal of Geophysical Research-Atmospheres, 106, p. 3503-3509 Hsieh, C. I., Katul, G., and Chi, T., 2000. An approximate analytical model for footprint estimation of scaler fluxes in thermally stratified atmospheric flows. Advances in Water Resources, 23, p.765-772. Kljun, N., P. Calanca, M. W. Rotach and H. P. Schmid, 2004. A simple parameterisation for flux footprint predictions, Boundary-Layer Meteorology, 112(3), 503-523. formann, R. and F. X. Meixner, 2001, An analytical footprint model for non-neutral stratification, Boundary-Layer Meteorology, 99(2), 207-224

Funding for TERENO was provided by the Federal M Education and Research (BMBF). The support by the Ba Staatsforsten and our co-worker Elisabeth Eckart is app