

1 2.1: *In-situ* Ocean Observations: A brief history, present status and future directions

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24 Abstract

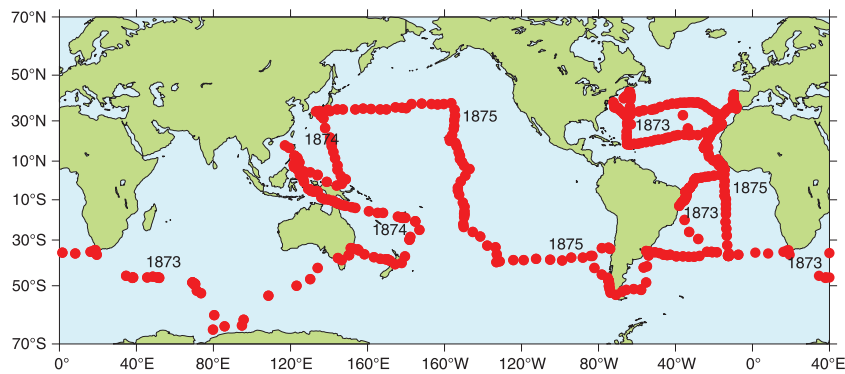
25 Observations at and below the surface of the oceans are essential for understanding the ocean system  
26 and the role played by the ocean in earth's climate, for documenting changes and for initialising,  
27 validating and improving ocean models. It is only since the late 20<sup>th</sup> century that, thanks to advances in  
28 microelectronics, battery technology and satellite communication *in-situ* observations, (together with  
29 satellite observations), have reached a volume and spatial distribution that allows us to track a wide  
30 range of global and regional phenomena. This review traces the development of *in-situ* ocean  
31 observations primarily from a physical standpoint and describes the internationally co-ordinated  
32 observing networks that now supply these observations. It considers the enormous changes that have  
33 occurred in the volume and distribution of these observations and the implication of these changes for  
34 defining the evolving state of the global ocean. Finally there is discussion of the prospects for further  
35 improving sustained ocean observations and for the delivery of integrated information from interrelated  
36 observing networks.  
37

38 1. Introduction

39 Observations of the interior of the ocean are fundamental to understanding ocean dynamics and  
40 properties, to monitoring changes in the oceans' state, (whether caused by natural or human  
41 influences), to quantifying the forcing at the atmosphere-ocean (in some areas, atmosphere-ice-ocean)  
42 boundary and for determining the role and importance of the ocean in the climate system. *In-situ* ocean  
43 observations also complement and provide ground truth for remotely-sensed observations of the ocean  
44 from earth-observing satellites (Chapter 2.2). Both satellite and *in-situ* observations are vital for ocean  
45 forecasting, ocean reanalysis and for assessing the fidelity of ocean and earth-system models and  
46 underpinning their future improvement (Chapters 5.2 and 5.3).

47

48 The technical and logistical challenges of making *in situ* ocean observations are legion; measurements  
49 often have to be made in areas far removed from land, in a corrosive liquid, at great pressure and in a  
50 fluid that is effectively opaque to electromagnetic radiation. Capturing the oceans' variability requires  
51 repeated measurements over wide areas and yet with small spatial resolution. Detecting change  
52 demands measurements of high precision and stability over decadal and longer time scales. For these  
53 reasons the history of scientifically-focussed, open ocean observations is relatively short: it may be said  
54 to have started with the voyage of *HMS Challenger* in the 1870s. (Wyville Thompson and Murray,  
55 1885)



56

57

58 **Figure 1.** Track of *HMS Challenger*. This was the first major scientific exploration of the global ocean. The  
59 voyage lasted 4 years and covered more than 68,000 nautical miles.

60

61 Through the 20<sup>th</sup> century, measurements became more accurate but remained relatively sparse and  
62 regionally focussed until the 1990s. During that century there were a number of initiatives and  
63 technical advances that, with hindsight, can be regarded as having been crucial steps in improving our  
64 ability to make systematic measurements within the global ocean. Often the driver for progress was the  
65 sequence of “*New observations lead to new understanding – new understanding points out the  
66 inadequacy of earlier observations – this understanding stimulates new technical development*”. The  
67 other major driver for progress was the application of advanced technologies to the oceans – solid state  
68 electronics in the 1960s and 70s, miniaturised computing power, and satellite communication and  
69 navigation from the 1990s to the present day. Innovative exploitation of these advances has led to  
70 major advances in our observational capability. In the following we briefly review the development of  
71 key observing technologies (section 2) and their impact on the number and distribution of ocean  
72 observations (section 3). Apart from the ever-present uncertainties of funding, the future for sustained  
73 ocean observations through the exploitation of emerging technologies (Section 3) and within the new  
74 Framework for Ocean Observations (section 5) is bright.

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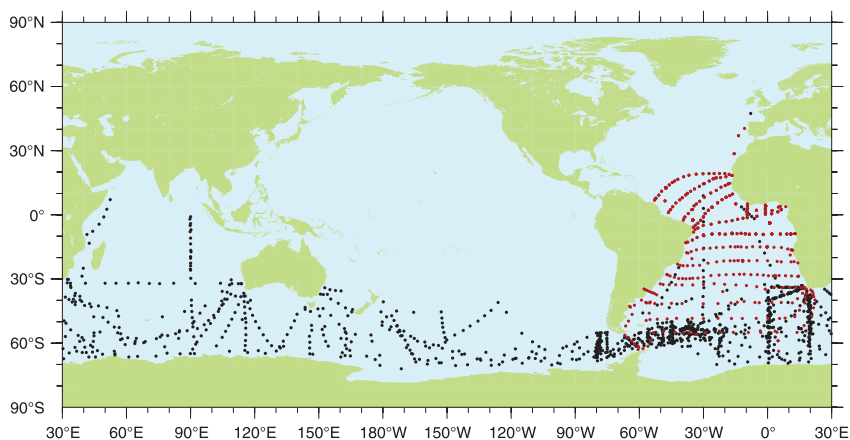
77 2. Development of present observational capability

78 Rather than considering the advances in observational capability on a parameter-by-parameter basis we  
79 will discuss the topic through a brief chronology of some of the most important technical developments  
80 that have enabled the establishment of the present multi-parameter ocean observing systems. In recent  
81 years progress has often been made in a number of key ocean parameters simultaneously through the  
82 mounting of major internationally coordinated observational programmes and/or the adoption of new  
83 generic observing platforms or technologies.

84  
85 2.1 Late 19<sup>th</sup> to mid 20<sup>th</sup> centuries

86 Improvements in navigation were the first drivers of systematic ocean observations. In the 19<sup>th</sup> century,  
87 following the introduction of Harrison's chronometer that enabled longitude to be determined, it  
88 became possible to estimate surface currents from a vessel's navigation and these were recorded in  
89 navigational logs. Such observations led to the compilation of surface currents in the Atlantic by James  
90 Rennell and published (1832) posthumously by his daughter and more widely by Mathew Fontaine  
91 Maury (1855). The major motivation for Maury's work was not scientific but was to use knowledge of  
92 surface currents to shorten sea voyages and hence gain commercial advantage. This was also a driver  
93 for the measurement of ocean temperatures since it was recognised by Benjamin Franklin that  
94 temperatures changed across the Gulf Stream and that by navigating into water of the correct  
95 temperature ships could speed their voyages between America and Europe. Early ocean surface  
96 temperature measurements were made by dipping a simple mercury-in-glass thermometer into, first,  
97 wooden and later canvas buckets of water collected from the sea surface: a technique that remained in  
98 common use until the mid-late 20<sup>th</sup> century. While wooden buckets were well insulated, canvas ones  
99 were less so and thus these temperature measurements are now known to be biased low due to  
100 evaporative cooling.

101  
102 The ability to make subsurface observations of temperature and salinity developed substantially  
103 between the pioneering voyage of *HMS Challenger* (1872-6), the *Meteor* expedition to the Atlantic  
104 (1925-27) (Wüst, 1935), and *Discovery* investigations in the Southern Ocean (starting in 1925)  
105 (Herdman, 1948). Most temperature measurements on *Challenger* were made with Six's  
106 maximum/minimum thermometers under the erroneous assumption of a monotonic decrease of  
107 temperature with depth. A small number of Negretti and Zambra reversing thermometers were also  
108 used on the Expedition (see Rice 2001) and these subsequently became the standard method for  
109 determining subsurface temperatures until the 1970s. Carefully calibrated reversing thermometers  
110 could determine temperature to at best about 5 millidegrees. The difference between paired  
111 thermometers (one protected against pressure effects and the other unprotected) allowed the depth of  
112 the measurement to be determined to within 10m.



113  
114 **Figure 2.** Stations worked during the Meteor Atlantic Expedition (red) and by the Discovery Investigations  
115 (black) (From NOAA WOD)

116 On *HMS Challenger*, salinity was determined by measuring density using a hydrometer and converting  
117 this value to a quantity of total dissolved solids. In 1902 the International Council for the Exploration  
118 of the Sea (ICES) was established to investigate the relationship between physical ocean properties and  
119 fisheries in the Northern Atlantic. One of ICES' earliest and most significant contributions to physical  
120 oceanography was the work of Martin Knudsen in standardising the determination of salinity by  
121 titration against silver nitrate solution and building on the investigation of seawater chemistry by  
122 Dittmar (1884) and Forchhammer (1856). A Standard Seawater Service was established by ICES, the  
123 successor organisations of which continue to provide the internationally accepted standard to the  
124 present. (<http://www.sea-technology.com/features/2011/0611/salinity.php>)  
125

126 Subsurface water samples were collected and thermometers deployed using strings of typically up to 12  
127 water sampling bottles clamped to a wire and sequentially triggered to close the bottles and reverse the  
128 thermometers by weights (messengers) sliding down the wire. In deep water multiple casts could be  
129 deployed at each station to sample the full ocean depth but even so the deepest measurements were  
130 commonly separated by several hundred metres. Many designs of sampling bottle were used but by the  
131 1920s, the Nansen bottle had become the generally-used standard. Water samples were drawn from  
132 each bottle (occasionally with duplicates) and stored for analysis by titration (for salinity) either at sea  
133 or the end of the voyage. The Nansen bottle/reversing thermometer combination was used virtually  
134 unchanged until the 1960s/70s when continuous profiling was introduced together with multisampler  
135 cassettes and plastic sample collection bottles. (See Section 2.2)  
136

137 Wüst (1935), using data from the Meteor Expedition that included dissolved oxygen (Winkler, 1888),  
138 developed the "core method" by which the spreading of subsurface water masses was inferred by  
139 tracing ocean property distributions. Velocity could not be determined in absolute terms below the  
140 ocean surface but vertical shear could be estimated by the dynamical method – geostrophy (Sandström  
141 and Helland-Hansen, 1903).  
142

143 The next technological development that added substantially to the inventory of ocean observations  
144 was the mechanical bathythermograph invented by Athelstan Spilhaus in 1937 (Spilhaus, 1938) and  
145 developed further by Al Vine of the Woods Hole Oceanographic Institution during the early 1940s.  
146 This instrument enabled the thermal stratification of the upper 150 m of the ocean to be determined,  
147 firstly from submarines and later from underway surface ships, and made possible predictions of  
148 underwater sound propagation. The technology was simple, a bourdon tube for pressure, a bi-metallic  
149 strip for temperature and the results scribed on a smoked glass slide. This instrument was eventually  
150 used systematically by the scientific community to make a major contribution to understanding the  
151 spatial variability, for example, of the Gulf Stream (Fuglister, 1963).  
152

## 153 *2.2 Second half of 20<sup>th</sup> century*

154 The advent of solid state electronics and the use of "O"- rings to seal pressure cases had a profound  
155 effect on oceanographic instrumentation. It enabled small, battery-powered electronics to be fitted into  
156 modestly sized instruments. One of the first applications in the field of ocean physics was the  
157 development of Swallow's neutrally buoyant floats with which he made the first, absolute  
158 measurements of deep currents.<sup>1</sup> (Swallow 1955) and then to discover (Crease 1962) the first evidence  
159 of the existence of an energetic ocean dominated by mesoscale variability.  
160

161 Technical development accelerated in the 1960s and early 1970s, leading to observational methods that  
162 we now regard as routine and greatly enhanced by improvements in navigation: LORAN and Decca  
163 Navigator in the 1960s followed by Transit satellite navigation in the 1970s. This meant for the first

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<sup>1</sup> Prior to 1955 measurements had been made by primitive recording current meters lowered from anchored ships. Such records were short and contaminated by navigational uncertainties. They could be summarized in a single one-page table (Bowden, 1954).

164 time that we knew (to an accuracy better than 1km) where observations had been made and thus it was  
165 possible to make greatly-improved estimates of surface currents. Chemical titrations for the  
166 determination of salinity were replaced by salinometers measuring electrical conductivity (Park and  
167 Burt, 1965) and the traditional water bottle/reversing thermometer combination started to be replaced  
168 by continuous temperature and salinity profilers - the Conductivity (Salinity), Temperature Depth  
169 (C(S)TD) instrument. The first were inductive STDs, such as that marketed commercially by Bissett  
170 Berman, and these were followed in the early 1970s by the Neil Brown CTD that used a conductivity  
171 sensor with greater stability. Calibrations of measurements from these profilers still depended on  
172 mercury-in-glass thermometers and, in the absence of multiple samples in the vertical, were often  
173 limited to only a deep and shallow calibration point in each profile.

174  
175 The limited measurements of deep ocean currents by Swallow's ship-tracked floats (lasting a few days  
176 and covering distances measured in tens of kilometres) were later enhanced by floats tracked over  
177 many months from fixed listening stations using low frequency sound propagation through the SOFAR  
178 channel. The CTD and SOFAR floats were developed specifically for use during the Mid-Ocean  
179 Dynamics Experiment conducted near Bermuda and allowed the first mapping of the ocean mesoscale  
180 velocity field (The MODE Group, 1978). Later the SOFAR float measurements were simplified by  
181 using fixed sound sources and receivers on drifting RAFOS (SOFAR spelled backwards) floats,  
182 (Rossby, Dorson and Fontaine, 1986).

183  
184 By the late 1960s the mechanical bathythermograph started to be replaced by the eXpendable  
185 bathythermograph (XBT). The XBT measured temperature with a thermistor and estimated depth as a  
186 function of time using a fall-rate algorithm. The probe relayed its data to the deploying ship through a  
187 thin 2-conductor copper wire spooled from reels on both the probe and the ship. The initial XBTs  
188 (model T4) reached approximately 450 m while later models reached greater depths. The majority of  
189 such probes were used by navies in anti-submarine operations but large-scale civilian use rapidly  
190 developed in experiments such as TRANSPAC. Such measurements greatly increased our knowledge  
191 of upper ocean variability (e.g. Koblinsky et al., 1984, Talley and White, 1987). However the reliance  
192 on fall rate algorithms was later revealed to be problematic when attempts were made to merge these  
193 data with other sources in which pressure (depth) was measured directly (Wijffels et al., 2008).

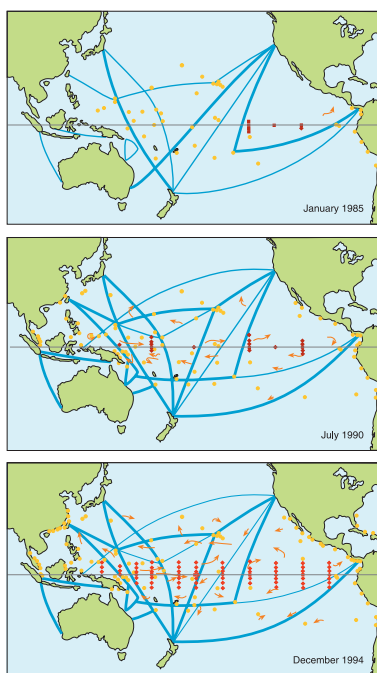
194  
195 Surface temperature measurements by ships were, by this time, being made using sensors inserted into  
196 the engine cooling water intake and recorded automatically. While generating more data, these  
197 measurements also introduced uncertainty due to the widely differing depths of these intakes on large  
198 and small ships together with thermal contamination from the ships' machinery. This latter factor was  
199 later reduced by the use of hull contact sensors (Kent et al., 1991). The process of defining what is  
200 meant by "sea surface temperature" is complex and was brought into sharp focus by the challenges of  
201 providing ground truth for satellite observations in which the measurement is of the skin temperature  
202 rather than a bulk interior value (Donlon et al., 2002). Following the advent of satellite SST  
203 measurements these and *in-situ* data were combined to produce global climatologies of which that by  
204 Reynolds (1988) was an early example. Today the GODAE High Resolution Sea Surface Temperature  
205 consortium (GHRSSST - [www.ghrsst.org](http://www.ghrsst.org)) seeks to produce and improve such climatologies.

206  
207 The period of the 1960s and 1970s also saw significant advances in the measurement of subsurface  
208 ocean currents using moored instruments. Micro-electronics allowed the development of internally  
209 recording (on magnetic tape) current meters. Two current meter designs dominated the field in the  
210 west; the Aanderaa RCM4 and the Geodyne, while in the Soviet Union, the mechanical Alekseev  
211 instrument was used extensively. However, up until the 1980s individual records, rarely exceeded  
212 30 days and significant differences in instrument response dependent on mooring type – with surface or  
213 subsurface buoyancy – were found. This led to the development of vector-averaging current meters that  
214 significantly reduced the wave-induced contamination of records from surface moorings. Together  
215 these technologies allowed major experiments (Polygon (1970), Mid-Ocean Dynamics Experiment

216 (MODE) and POLYMODE (1973-78)) to explore and map the oceans' mesoscale variability (The  
217 MODE Group, 1978), Kamenkovich, (1986), Freeland and Gould, (1976)).

218  
219 During the 1970s first the NIMBUS and later the more accurate TIROS series of satellites started to  
220 provide global instrument tracking by measuring the Doppler shift of radio signals. The method was  
221 used to reveal the paths of surface drifters and ultimately developed into the Argos tracking system.  
222 Regional experiments used this technology: NORPAX in the North Pacific starting in 1975 (McNally  
223 et al., 1983), followed by a Gulf Stream Experiment in 1978 (Richardson, 1983) and culminated in the  
224 internationally co-ordinated deployment of 300 drifters in the Southern Ocean in 1978-9 as a  
225 contribution to the First GARP (Global Atmospheric Research Project) Global Experiment (FGGE),  
226 (Garrett, 1980). While the FGGE buoys collected surface temperature and atmospheric pressure data,  
227 the quality of the near-surface velocity data suffered from the large size of the float bodies (windage),  
228 the poor performance of the "window-shade" drogues and the inability to detect without ambiguity if  
229 the drogue was still attached. (The technological developments during this era are described by Baker,  
230 1981).

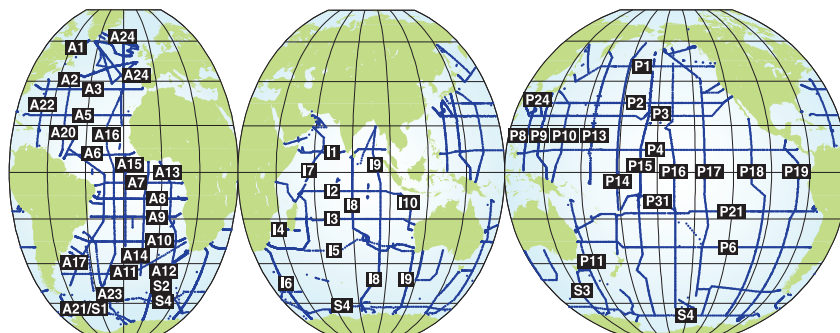
231  
232 Throughout the 1980s various combinations of moored current meters and neutrally buoyant floats,  
233 CTD profilers, expendable probes and surface drifters were used in a wide range of regional  
234 experiments. For example the Tropical Ocean Global Atmosphere (TOGA) project was a concerted  
235 effort to collect data from the equatorial Pacific combined with numerical modelling, aimed at  
236 understanding and predicting the evolution of the El Niño - Southern Ocean (ENSO) phenomenon.  
237 TOGA significantly enhanced the collection and distribution of *in-situ* sea level data, temperature  
238 profile data from XBT probes but most importantly led to the deployment of the Tropical Ocean  
239 Atmosphere (TAO, later TAO-TRITON) array of moorings measuring and reporting real-time upper  
240 ocean and atmospheric data (McPhaden et al 1998). The array has now expanded to cover the Atlantic  
241 and Indian Oceans (<http://www.pmel.noaa.gov/tao/global/global.html>). See section 2.3



242  
243 **Figure 3.** The evolution of the TAO-TRITON tropical Pacific Ocean observing system. (Top) The observing  
244 system at the beginning of the TOGA program in 1985, (middle) the evolving observing system in 1990 and  
245 (bottom) the sustained observing system that now comprises the TAO-TRITON observing system.  
246 Key. XBTs from volunteer observing ships; (blue lines), Coastal tide gauges (yellow dots); Drifting buoy  
247 (curved arrows); Current meter, temperature and salinity moorings and surface flux stations (red diamonds).  
248 (From McPhaden et al., 1998).

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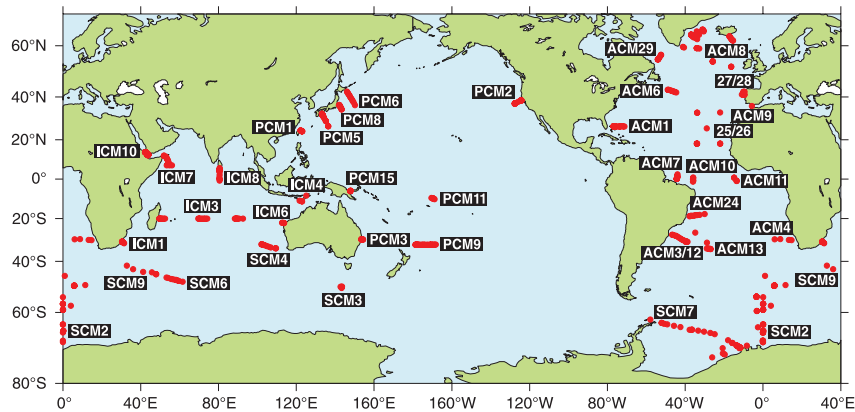
250 The insights into ocean variability gained during the 1960s, 1970s and 1980s highlighted the problems  
251 of interpreting sparsely sampled *in-situ* data and of detecting long-term change in the oceans (Wunsch,  
252 2001). It was the prospect of satellites carrying radar altimeters (as had been heralded by the brief 1978  
253 SeaSat mission (See Chapter 2.2)) that led for the first time to the planning of an almost global scale  
254 programme - the World Ocean Circulation Experiment (WOCE) – that aimed to improve models of the  
255 oceans’ role in climate by collecting comprehensive remote sensed and *in-situ* data (Thompson et al.,  
256 2001).  
257



258  
259  
260 **Figure 4.** The WOCE One-Time hydrographic sections. Each line consisted of full-depth stations at typically  
261 100km spacing measuring temperature, salinity, nutrients and a suite of chemical tracers.  
262

263 Improved understanding of the oceans’ role in climate-related variability from WOCE-derived data has  
264 been well-documented in the various chapters of Siedler, Church and Gould (2001) but it is worth  
265 noting that the strength of WOCE, in terms of producing a global scale data set, is that the resulting  
266 data serves as a baseline against which past and future change may be assessed. Given this,  
267 considerable effort was made during WOCE to ensure high quality and internal consistency of the  
268 project’s data sets. WOCE incorporated a number of sub-programs including ship based hydrographic  
269 sections, western boundary current mooring arrays and XBT and surface drifter programs (the latter  
270 jointly with TOGA). In particular the quality and comprehensiveness of temperature, salinity and ocean  
271 chemistry data in the WOCE Hydrographic Programme (WHP) (Figure 4) in conjunction with the Joint  
272 Global Ocean Flux Study (JGOFS) (King, Firing and Joyce, 2001) was unprecedented. These data  
273 were subsequently described in a series of four atlases (Sparrow, Chapman and Gould, 2012).  
274

275 The most innovative and significant technical advance during WOCE was the development of a  
276 neutrally buoyant float that did not depend on acoustic tracking and hence could be deployed on a  
277 global scale. The Autonomous, LAgrangian Circulation Explorer (ALACE) floats, (Davis, Webb,  
278 Regier and Dufour, 1992), were ballasted to drift at depths around 1000 m and programmed to surface  
279 at regular intervals by changing their buoyancy. Once at the surface, their positions were fixed by  
280 satellite (Argos tracking). Successive surfacing positions gave a measure of the time-averaged ocean  
281 currents for each drift segment. These floats allowed for the first time the collection of subsurface  
282 velocity data across entire ocean basins (Davis, 1998). Later in WOCE, the floats started the collection  
283 of, first, temperature and later temperature and salinity profiles acquired as the floats rose to the  
284 surface. In 1998 the Profiling ALACE (PALACE) float was envisaged as the means of building a  
285 global array that would provide unprecedented observations of the upper 2000m of the open oceans  
286 (Argo Science Team, 1998).



**Figure 5.** WOCE moored current meter arrays. Most mooring arrays were deployed for approximately two years. In many locations these arrays provided the first timeseries of ocean velocity, temperature and salinity.

In the 1990s the WOCE and TOGA programmes also provided an impetus for the systematic global-scale use of moorings. Since the 1980s technology had matured to allow for deployments of full-ocean-depth moorings carrying a range of self-recording instruments to observe variability of currents, temperature, salinity and pressure. In WOCE, arrays of moorings across several "choke points" (geographic constrictions of important ocean flows) documented the transports of deep and shallow boundary currents and flow through passages and gaps as well as contributing to trans-basin transport estimates. Retired ocean telephone cables complemented the choke point observations so as to observe changes in conductivity-weighted ocean transports. For example, more than 30 years of Florida Strait transport using this technique have made a significant contribution to documenting parts of the upper limb of the Atlantic subtropical gyre transport (Meinen et al., 2010). In the tropical Pacific, the TOGA moored array reached full implementation to record the state of the equatorial thermocline and provide real time information of winds to support seasonal El Niño forecasting.

The late 1980s and early 1990s also saw the development of moored and ship-mounted Acoustic Doppler Current Profilers (ADCP) for the observation of ocean velocity profiles. The ADCP measures ocean currents using sound waves to detect the Doppler effect from small scattering particles in the ocean; as particles move toward or away from the sound source, the frequency of the return signal is either higher or lower. Assuming that the particles are advected by the ocean currents, the frequency shift is proportional to the speed of the current along the axis of the acoustic beam. Combining information from 3 or more beams allows derivation of the ocean velocity in all three coordinates. As the emitted sound travels through the water column, the ADCP measures the current at many different depths simultaneously.

The full exploitation of ship-mounted ADCPs was dependent on the arrival of another enabling technology; the Global Positioning System (GPS). This became increasingly available to civilian users during the 1990s and as well as enabling absolute positions to be determined instantaneously to metre accuracy, ship's heading could also be determined with an accuracy far better than from previously-used gyro-compasses. This provided the information needed to accurately determine the speed and direction of the ship over the ground and enabled ADCPs to be mounted on research vessels to provide underway upper-ocean velocity (to 800 m) observations. From the 1990s onwards ADCPs were also incorporated in CTD/multisampler packages allowing velocity to be determined throughout the water column. (Fischer and Visbeck, (1993), King, Firing and Joyce, (2001)).

During the 1990s collaboration between WOCE and TOGA also led to an expansion of the collection of XBT data and significant improvements in the quality and quantity of data collected by surface drifters (See Chapter 4.2). As mentioned, the FGGE drifter extensively used in the 1970s did not



328 provide high quality surface ocean velocity data. A standardised WOCE/TOGA drifting buoy to suit  
329 observational requirements for meteorological and oceanographic applications was designed and  
330 deployed and is largely responsible for the improved data quality of Lagrangian surface measurements  
331 (Sybrandy and Niiler, 1991 and Chapter 4.2).

332  
333 Many of the physical observations established by WOCE and TOGA were eventually subsumed into  
334 the framework of the much broader CLIVAR (Climate Variability and Predictability) project,  
335 established in 1995 as a component of the World Climate Research Program (WCRP). CLIVAR's  
336 focus on the coupled ocean-atmosphere system and its interest in a broad range of time scales (from  
337 seasonal to centennial) inevitably diluted the momentum in full ocean depth observations gained during  
338 WOCE. However, a much broader agenda of sustained ocean observations was developed under the  
339 auspices of GOOS (Global Ocean Observing System) responding to the need to understand the ocean's  
340 role in climate as identified by WCRP, the Global Climate Observing System (GCOS) established in  
341 1992 and by the United Nations Framework Convention on Climate Change, UNFCCC. Observations  
342 were also increasingly required to be delivered in near-real-time for use in operational ocean  
343 information products to guide and safeguard marine operations and deliver short term ocean and  
344 weather forecasts. As the 20<sup>th</sup> century closed, there was a growing recognition of the potential to build  
345 on the observational capabilities that had been established for limited-lifetime scientific experiments so  
346 as to provide a framework for an emerging sustained ocean observing system.

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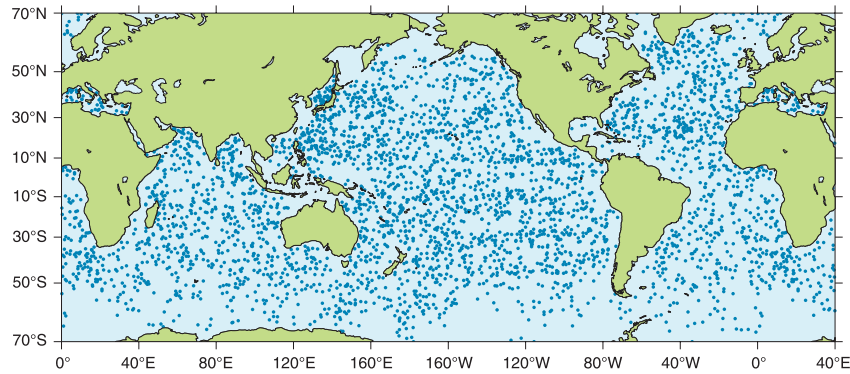
### 348 *2.3 21st century: consolidation of capabilities and growth of sustained observations*

349 The ocean observations collected during WOCE and TOGA were used to document the importance of  
350 the ocean in regional and global climate on short (days) to longer (decades and centuries) time scales  
351 (e.g. Siedler, Church and Gould, 2001). Moreover, WOCE established the strong international  
352 collaborations amongst national research institutes and funding agencies that would be required if a  
353 coordinated ocean observing programme were to be established together with the systems needed to  
354 collate, quality control and distribute data. Building on the success of WOCE, the aspirations of  
355 CLIVAR, and the requirements from GCOS and UNFCCC, and the continuing technological  
356 developments, in 1999 the ocean community held the first international conference solely focussed on  
357 sustained ocean observations, the OceanObs'99 conference (Koblinsky and Smith, 2001). The  
358 Conference's goal was to provide the framework and to set feasible objectives for the establishment of  
359 the first decade of a sustained ocean observing system. The network of sustained ocean observations  
360 that emerged from this conference was organised primarily around observing platforms that had been  
361 developed during WOCE, TOGA and CLIVAR.

362

363 The major observational programs that were delineated at OceanObs'99 were:

- 364 - a program called Argo that sought to establish and maintain a global-scale array of floats similar  
365 to the profiling ALACE float developed in WOCE. The array of 3000 instruments (roughly one  
366 every 300x300 km in the ice-free oceans deeper than 2000 m) would collect profile data  
367 (temperature and salinity) to 2000 m at nominal 10 day intervals. The data would be freely  
368 available in real-time and in a climate-quality-controlled data set with a 6 month lag. Argo  
369 would also produce subsurface velocity estimates (The Argo Science Team, 2001).

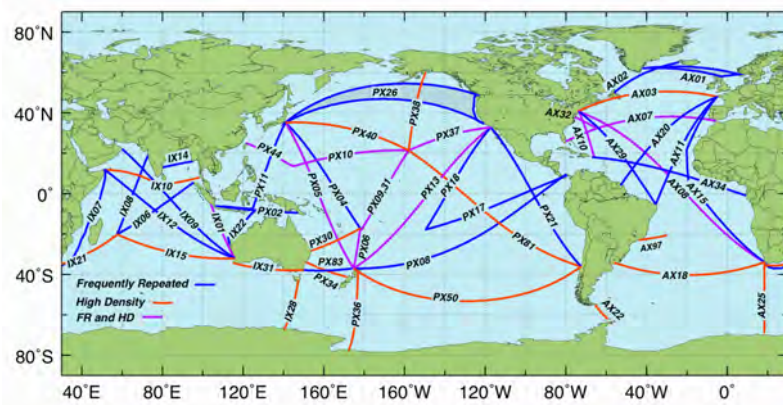


370

371 **Figure 6.** The array of Argo profiling floats that measure the temperature and salinity of the upper 2000 m of the  
 372 ocean. Positions of floats that had delivered data within the last 30 days of March 2013. (Data from [www-](http://www-argo.ucsd.edu)  
 373 [argo.ucsd.edu](http://www-argo.ucsd.edu)).

374 Argo reached its goal of 3000 operating floats by November 2007 and has since remained above  
 375 that level (due in large part to the steady improvement in float operational lifetimes), thus  
 376 providing continuous monitoring of the temperature, salinity, and velocity of the upper  
 377 temperate and tropical oceans. The initial Argo design criterion of an array density of one active  
 378 float per 300x300 km grid excluding regional seas and sea-ice zones has not yet been fully  
 379 achieved since some areas remain over populated while other regions are under sampled. The  
 380 success of Argo lies, not just in almost reaching design specification but in its data flow and  
 381 quality control systems. Argo has successfully established a real time data stream regardless of  
 382 the national provider of the particular floats and a delayed-mode stream of climate-quality  
 383 calibrated data.

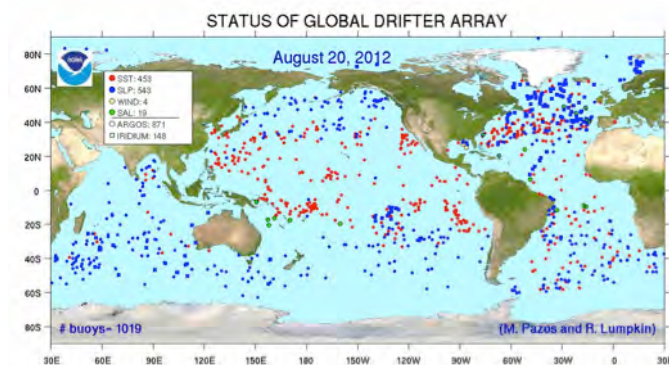
- 384 - A global XBT network that would focus on high resolution transects collecting temperature  
 385 profiles in the upper 800 to 1000 m across ocean basins mostly using commercial vessels  
 386 engaged in the Ship of Opportunity Program using semi-automatic XBT launchers. Although  
 387 some XBT transects have been maintained for 30 years, the program was redesigned after the  
 388 advent of Argo from a broad-scale sampling to a network with increased spatial and temporal  
 389 resolution focusing on boundary and choke points currents or regions of high seasonal  
 390 variability (tropical oceans) that would be complementary to the Argo global broad-scale array  
 391 (Smith et al, 2001). Currently two modes of XBT transects are in operation: Frequently  
 392 Repeated (RF), transects that are occupied 12-18 times per year with XBT deployment every  
 393 100-150 km, and High Density (HD) transects occupied 4 times per year with XBT deployment  
 394 every 25 km.



395

396 **Figure 7.** The present-day XBT network of transects across ocean basins. The XBTs are deployed from research  
 397 vessels and ships of the Ship of Opportunity Program. The transects are sampled in two modes: High Density (HD)  
 398 and Frequently repeated (FR). Some transects include time series with more than 30 years of data.  
 399 ([http://www.aoml.noaa.gov/phod/goos/xbt\\_network/index.php](http://www.aoml.noaa.gov/phod/goos/xbt_network/index.php)).

400 - Following on from WOCE and TOGA, a Global Surface Drifter Program was established that in  
 401 2005, reached the design density provided by 1250 drifters. The drifters are needed to anchor  
 402 satellite-based measurements of sea surface temperature as a critical component of the GODAE  
 403 High Resolution Sea Surface Temperature (GHRSSST - [www.ghrsst.org](http://www.ghrsst.org)). In addition they are  
 404 able to measure surface velocity. A subset of the drifters additionally observe atmospheric  
 405 pressure and surface wind speed and direction. A dedicated data centre assembles and provides  
 406 uniform quality controlled SST and surface velocity measurements. (This program and its  
 407 contribution to surface current measurement are described in Chapter 4.2.)



408  
 409 **Figure 8.** The global drifter array, as of August 20, 2012, that provides information of sea surface velocity,  
 410 temperature and/or pressure. (Image from <http://www.aoml.noaa.gov/phod/dac/index.php>)

411 - There are a small number of locations where regularly repeated observations over long time  
 412 periods have given insights into physical and biogeochemical ocean variability and also into the  
 413 processes that connect the upper ocean to its deep interior. Notable among these are  
 414 Hydrostation S (since 1954) and the Bermuda Atlantic Timeseries (BATS) (since 1988) near  
 415 Bermuda and the Hawaii Ocean Timeseries (HOT) (since 1988). More recent addition include  
 416 the Cape Verde Ocean Observatory (CVOO) and the European Station for Timeseries in the  
 417 Ocean (ESTOC) since 1994 near the Canary Islands. In addition some long-term observations  
 418 were conducted from ocean weather ships that were instrumental to support early  
 419 intercontinental air travel. The last weather ship, “Mike” in the central Norwegian Sea, was  
 420 decommissioned in January 2010. Several of those sites have been continued with moored  
 421 observatories, most notably ocean weather station “Bravo” in the central Labrador Sea and  
 422 “Papa” in the Northeast Pacific

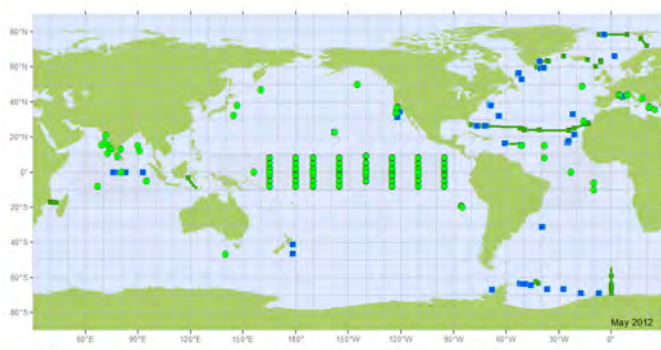
423 Building on these stations, and on the tropical surface mooring technology established during  
 424 TOGA and the WOCE boundary current mooring arrays, OceanSITES ([www.oceansites.org](http://www.oceansites.org))  
 425 was initiated at OceanObs’99 as a network of full-depth and surface time series at key climate-  
 426 relevant locations (Send, et al, 2001). It has since developed into a global network of moorings  
 427 at strategic locations in the ocean that measure a diverse range of ocean variables. This program  
 428 now incorporates the

- 429 • Tropical moored arrays (Pacific - TAO/Triton, Atlantic - PIRATA, Indian -RAMA)

430 • Arrays monitoring the North Atlantic overturning (MOVE, RAPID-WATCH, 16°N,  
431 53°N, Denmark Strait, Faroe-Shetland Channel and Fram Strait)

432 • Sites documenting water mass property changes such as those mentioned in the  
433 previous paragraph and similar long-term mooring/timeseries sites in the South  
434 Atlantic, Pacific, Indian, Arctic and Southern Oceans.

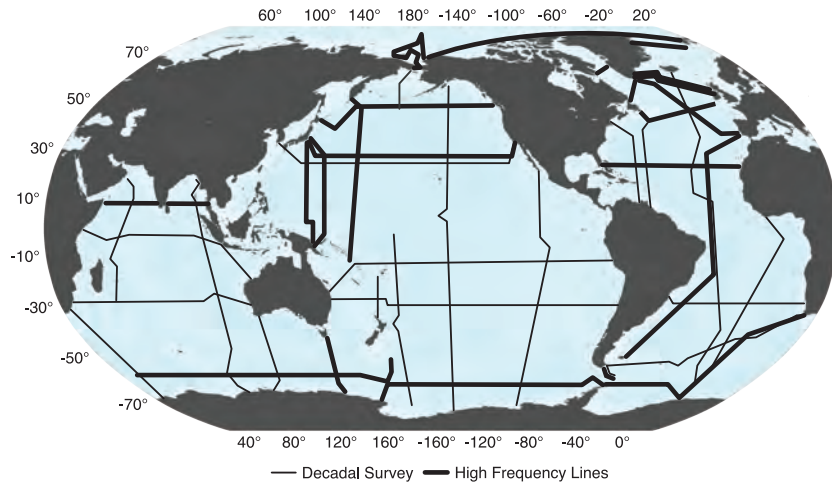
435 The time series from the observatories provide the means to develop accurate fields of  
436 watermass formation and transformation, as well as estimation of air-sea fluxes, and to allow  
437 quantification of the transports of major ocean current systems, and assessments of the  
438 variability of the vertical structure of the ocean and the role of eddy processes in the transport of  
439 heat and other properties.



440

441 **Figure 9.** Map of operating OceanSITES time-series stations as of May 2012. OceanSITES is a worldwide system  
442 of long-term, deepwater reference stations measuring many variables and monitoring throughout the water column  
443 from air-sea interactions down to 5000 meters. (Image from <http://www.oceansites.org/>). (To be improved with  
444 high resolution version)

445 - At OceanObs'99, Gould, Toole et al. (2001) articulated the need for continued systematic ship-  
446 based survey of the global ocean. This plan was supported by CLIVAR, GOOS and the  
447 International Ocean Carbon Coordination Project (IOCCP) and has since resulted in a program  
448 of hydrographic sections based on the WOCE lines and re-occupied at 5-10 year intervals. The  
449 Global Ocean Ship-based Hydrographic Investigations Program (GO-SHIP) (Hood et al., 2010)  
450 was formally established in 2007 to oversee the continuation of sustained ship-based  
451 hydrography and particularly addressed the needs of the community of scientists concerned with  
452 the ocean carbon and related chemical measurements. A global survey was completed in 2013  
453 and reoccupation of sections are continuing. GO-SHIP ([www.go-ship.org](http://www.go-ship.org)) is developing formal  
454 international agreements for a sustained international repeat ship-based hydrography program,  
455 including an internationally-agreed strategy and implementation plan; advocacy for national  
456 contributions to this strategy and participation in the global program and; providing a central  
457 forum for communication and coordination. GO-SHIP has brought up to date the manual of  
458 best practice for ship based hydrography first produced by WOCE. The establishment of GO-  
459 SHIP recognises that, despite numerous technological advances over the last several decades,  
460 ship-based hydrography remains the only method for obtaining the highest-quality, high spatial  
461 and vertical resolution measurements of a suite of physical, chemical, and biological parameters  
462 over the full water column. Ship-based hydrography (and the sparse array of deep moorings)  
463 provide essential contributions towards documenting ocean changes throughout the water  
464 column. This is particularly important for the deep ocean below 2 km (52% of global ocean  
465 volume) that cannot yet be sampled by profiling floats although such floats are under  
466 development.  
467

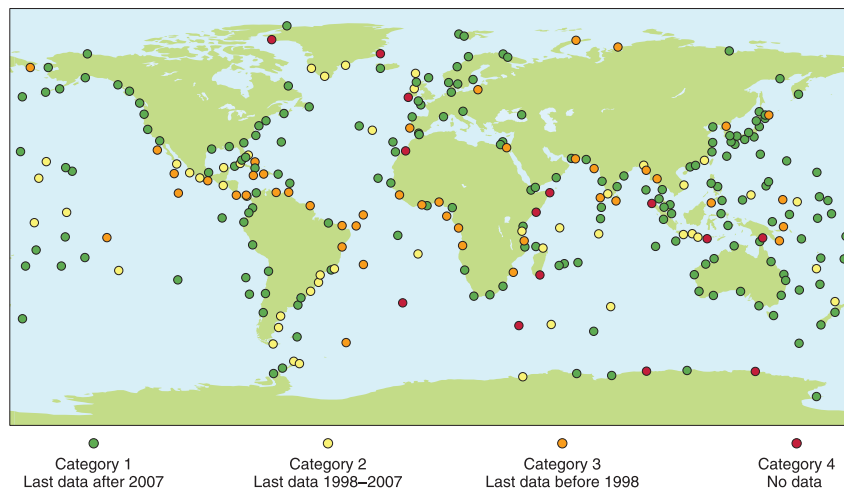


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**Figure 10.** Repeat hydrographic sections of the GO-SHIP program. These sections are maintained by the cooperation of a number of countries with the objective of completing a global survey every 10 years. (Image from [www.go-ship.org](http://www.go-ship.org)).

While these observing programs build on advances during the TOGA/WOCE era, two others have older roots.

- The systematic measurement of sea level extends back to the early 19th century and though such measurements were originally made in support of safe navigation, they now provide a valuable monitor of the consequences of changing ocean heat storage and an essential element of the ENSO monitoring and forecasting system. These measurements have, since the early 1990s been supplanted by systematic satellite altimetry but they remain an essential independent benchmark. (These topics are discussed in detail elsewhere in this book (Chapters 2.2, 4.5 and 6.1).)



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**Figure 11.** The present status of the GLOSS network of sea level stations showing the status of data delivery from the sea level sites that comprise the Global Core Network. (see <http://www.gloss-sealevel.org/>)

- All these observing systems are focussed on physical observations but ocean biogeochemistry also plays a key role in the climate system (Chapters 1.1, 5.7, 6.4, 6.5). Though biological and chemical ocean observations have a long history, it is only in the past three decades through programmes such as the JGOFS, IMBER and SOLAS components of the International Geosphere-Biosphere Program (IGBP), that these have been co-ordinated on a global scale and none yet match the density and global scale of physical measurements. The longest

494 timeseries of observations are those made since the 1930s by the continuous plankton  
495 recorders towed by research and commercial vessels (<http://www.globalcpr.org/>) These data  
496 have documented the changing patterns of zooplankton distribution many of which can be  
497 clearly linked to changes in ocean physics (Reid and Beaugrand, 2012).

498 In summary, progress over the first decade of the 21<sup>st</sup> century the *in-situ* observing system (detailed in  
499 plenary and community white papers published from OceanObs'09) includes:

- 500 • Complete implementation of the core Argo mission with more than 3500 (compared to a  
501 target of 3000) floats currently delivering high-quality data from the open and ice free oceans.  
502 (It should be noted that there are a number of developments being made that enable profiling  
503 floats to collect deeper data and data from ice-covered regions – See Section 3)
- 504 • Improvement in global distribution and number of surface drifters for sea surface temperature  
505 ground-truthing. (1100 drifters, 470 measuring atmospheric pressure and 700 measuring SST  
506 (March 2013))
- 507 • Maintenance of Pacific Tropical moored array and extended coverage of the moored array into  
508 the Atlantic and Indian Oceans.
- 509 • Established a growing network of multidisciplinary moored ocean observatories  
510 (OceanSITES).
- 511 • The reinvigoration of the science of sea surface temperature (SST) estimation based on the  
512 synthesis of the multiple satellite platform data streams and situ data; producing new and  
513 better SST products (with errors) via the Global High Resolution SST project (GHRSSST).
- 514 • Stabilisation of high quality sea level observations under the Global Sea Level Observing  
515 System (GLOSS) project.
- 516 • The transition of the global XBT network from broad-scale monitoring (taken over by Argo)  
517 to circulation monitoring via frequently repeated (FR) and high-density (HD) lines with a  
518 global design.
- 519 • The success in internationally coordinated efforts to reoccupy a subset of hydrographic and  
520 tracer transects (GO-SHIP).

### 521 3. Emerging and specialized ocean observing technologies

522 In the previous section there was a focus on ocean observations within globally co-ordinated observing  
523 networks that can be sustained for many years. However, during WOCE, TOGA and CLIVAR  
524 oceanographers have also developed and deployed other ocean observing systems, platforms and  
525 sensors, some of which may become part of sustained observing systems.

526  
527 CLIVAR in particular identified three type of observing activities.

- 528 • It encouraged and contributed to the establishment of the aforementioned global observing  
529 networks in collaboration with GOOS, with the expectation that the data they provided would  
530 support ocean analyses and the assessment of ocean variability and change from seasons to  
531 decades. Moreover, these data would provide the basis for ocean and climate prediction.
- 532 • Secondly CLIVAR sponsored several regional oceanographic experiments designed to test new  
533 observing capabilities in pilot mode. Examples are RAPID (monitoring the North Atlantic  
534 Meridional Overturning Circulation (MOC)), TACE (in the tropical Atlantic), VOCALS  
535 (around South America), SPICE (Southeast Pacific) and KESS (Kuroshio Current extension).  
536 Each of these sites, operated for a period of three to five years in regions of particular interest to  
537 CLIVAR's research objectives and were instrumented with a mix of traditional and new  
538 technologies. The intention was to improve the understanding and representation of a key ocean  
539 process in models but also to develop and field-test innovative ocean observing systems that  
540 could eventually become part of the sustained ocean observing system. (RAPID is now in its 9<sup>th</sup>  
541 year of observations, e.g. Johns et.al. 2011, and [www.noc.soton.ac.uk/rapidmoc/](http://www.noc.soton.ac.uk/rapidmoc/)).

542 • In addition there have been short-term ocean process experiments with ocean observations  
543 lasting from a few weeks to two years. These often employed observing technologies that were  
544 very likely to be too labor intensive or costly to become part of the sustained system.

545 Within these last two categories there have been deployments of new, innovative observing systems  
546 and sensors that we will now summarize. This is a rapidly developing field and the most appropriate  
547 and comprehensive references can be found in the White Papers submitted to the OceanObs'09  
548 conference (Hall, Harrison and Stammer, 2010).

### 549 *3.1 Advanced observing platforms*

550 An increasing demand for ocean observations, (in particular in remote areas, in winter and under  
551 extreme weather conditions), coupled with the high cost of research vessel operations have together  
552 stimulated an explosive growth of autonomous ocean observing platforms. Key climate processes occur  
553 in the Arctic and Southern Oceans, however sustained ocean observations in multi-year sea-ice, the  
554 marginal sea-ice region and under floating ice-sheets have been both a logistical and technological  
555 challenge. Modifications to and novel use of existing observations platforms are beginning to fill this  
556 data gap. Recent developments to Argo floats, both in software (ice detections algorithm and storage of  
557 multiple profiles) and hardware (rugged float bodies and antennae) are now extending the Argo  
558 program to the high latitude oceans (e.g. Wong and Riser, 2011). Moored profiling systems deployed  
559 on fast- and multi-year ice have also been developed to provide profiles of ocean properties under ice  
560 shelves and within the drifting ice pack. (e.g.. Timmerman et al., 2010). Novel use of miniature CTD  
561 systems deployed on animals, mainly seals, are now provide ocean observations in the high latitude  
562 oceans. (Boehme et al 2008)

563  
564 Self-propelled ocean gliders, using the same buoyancy engines as Argo floats and, with their short  
565 wings and satellite navigation, capable of underwater navigation are now used increasingly to routinely  
566 observe ocean property changes (e.g. <http://www.ego-network.org>). They can navigate the upper ocean  
567 (typically upper 1000m) with an effective speed of less than  $0.3 \text{ ms}^{-1}$ . Most gliders have battery  
568 endurance from several weeks up to almost a year depending on profiling depth and sensor suites.  
569 Gliders are expected soon to become part of the sustained ocean observing strategy. Other autonomous  
570 underwater vehicles (AUVs) include shallow and deep diving systems with propellers. A range of  
571 small vehicles with several hour endurance, well suited for coastal and near shore/ship applications, are  
572 available. Other systems can sample to 6000 m and have been used to make measurements of bottom  
573 boundary layer mixing processes.

574 New near-surface platforms have been developed. These include large moored platforms with fast  
575 satellite communication and significant onboard power capable of supporting new multidisciplinary  
576 ocean observatories and complementing regional sea floor cabled installations described below. Other  
577 systems use surface wave energy to propel the vehicle (Wave Glider) and unmanned sailing vessels  
578 may be expected to become available in the near future.

579 Consortia of commercial shipping companies and science organizations have come together to arrange  
580 for a wide range of instrumentation to be installed on commercial vessels (ferries, cruise ships and  
581 container vessels). This expands the capabilities of the Voluntary Observing Ship fleets from previous  
582 observations of surface temperature and salinity, meteorological parameters and the deployment of  
583 expendable (XBT) probes to include surface layer chlorophyll (fluorescence),  $\text{pCO}_2$  and upper ocean  
584 velocities. A summary of issues relating to the use of commercial vessels can be found in the final  
585 report of OCEANSCOPE project ([http://www.scor-int.org/Working\\_Groups/wg133.htm](http://www.scor-int.org/Working_Groups/wg133.htm)) co-sponsored  
586 by SCOR and IAPSO.

587 New platforms are under development for use in the near surface environment: Several designs for  
588 moored battery-powered winch systems are becoming available that allow sensitive sensor packages to

589 be kept for most of the time below the euphotic and surface layers to reduce bio-fouling and  
590 mechanical stress and to profile to the surface when required.

591 A growing number of cabled sea-floor observatories will enable the deployment of multi-disciplinary  
592 oceanographic sensor systems. The relatively high investment cost of sea-cables will limit the number  
593 of installations in the foreseeable future, but in particular new sensors and the observation of changes in  
594 real time near the sea floor will provide future opportunities for ocean observations. (See  
595 <http://www.oceanobservatories.org/>)

### 596 3.2 *Specialized observing systems and technologies*

597 Some key ocean parameters (ocean turbulence (mixing) and biogeochemical processes) that are crucial  
598 for improved understanding of fundamental ocean dynamics and development of sub-grid scale ocean  
599 parameterization in climate models cannot be adequately measured by techniques and observing  
600 platforms that are presently commonplace in the sustained observing system.

601  
602 Ocean turbulence (mixing) occurs sporadically and on short time scales and small spatial scales and so  
603 require highly specialized observing platforms. Observations of turbulence near the ocean surface are  
604 difficult due to the influence of wave motions, while turbulent observations in the ocean interior are  
605 sporadic and weak (Thorpe, 2004). Nevertheless, ocean turbulence is an important parameter in both  
606 the surface and deep ocean for dispersal of phytoplankton and stratification, respectively. Direct  
607 observations of the turbulence are obtained using fast thermistors and shear probes. These instruments  
608 can be fitted to a number of different platforms; free falling (or loosely tethered) profilers; gliders;  
609 mooring; submarines and AUV (Thorpe, 2004). Together with complementary observations of the  
610 ocean stratification rates of diapycnal mixing can be estimated. The highly episodic nature of  
611 significant mixing events put extra demands on the sampling and typically a large number of profiles  
612 need to be taken in a region to obtain reliable and representative estimates.

613  
614 An alternative approach has been to analyze the vertical dispersion of trace elements in the ocean. A  
615 particular successful method has been to inject an artificial inert tracer ( $\text{SF}_6$  or  $\text{SF}_5\text{CF}_3$ ) and follow its  
616 vertical and horizontal dispersion over many months to several years and large spatial scale (Ledwell et  
617 al., 1993). Several large scale tracer release experiments have been conducted in the Subtropical North  
618 Atlantic (NATRE, SaltFinger, GUTRE), Greenland Sea, Deep Brazil Basin (BBTRE) and in the  
619 Southern Ocean (DIMES) and have provided estimates of the time and space integrated ocean mixing  
620 rates.

621  
622 Because low frequency sound in the ocean can be detected over great distances and its speed is a  
623 function of ocean temperature and pressure, measuring acoustic travel times can be used to reveal the  
624 thermal structure of the ocean and its changes over time via a process known as acoustic tomography.  
625 The method uses cabled or moored arrays of sound sources and receivers to obtain tomographic images  
626 of ocean temperature distributions and their changes. High precision navigation of the moored source  
627 and receiver arrays, significant power consumption and potential impacts on marine mammals have so  
628 far limited the use of acoustic tomography in a sustained manner (Dushaw et al., 2010).

### 629 630 3.3 *New Sensors*

631 In general ocean sensors are required to have low power consumption, to be compact and to have low  
632 calibration drift. These requirements assume particular importance when deployment is made on  
633 autonomous or expendable platforms and much progress has been made in the past 20 years in  
634 improving these properties for temperature and salinity sensors. A diverse range of optical sensors are  
635 now available to measure ambient light, reflected light and fluorescence and are also used to measure  
636 dissolved oxygen, particulate carbon and chromophoric dissolved organic matter. A growing number of  
637 miniature wet chemistry analysis systems and even miniature mass spectrometers are becoming  
638 available.



639

640 Automatic image analysis and pattern detection will soon allow the routine identification of several  
641 types of plankton. First attempts to deploy gene-chips and highly specialized sensors will increase the  
642 range of multidisciplinary ocean observations in the coming decade. While the development of such  
643 novel biological and chemical sensors, now appears unrelated to the mainstream of climate science,  
644 measurements by such sensors may eventually lead to a better understanding of the impacts of climate  
645 change on the health and productivity of the oceans.

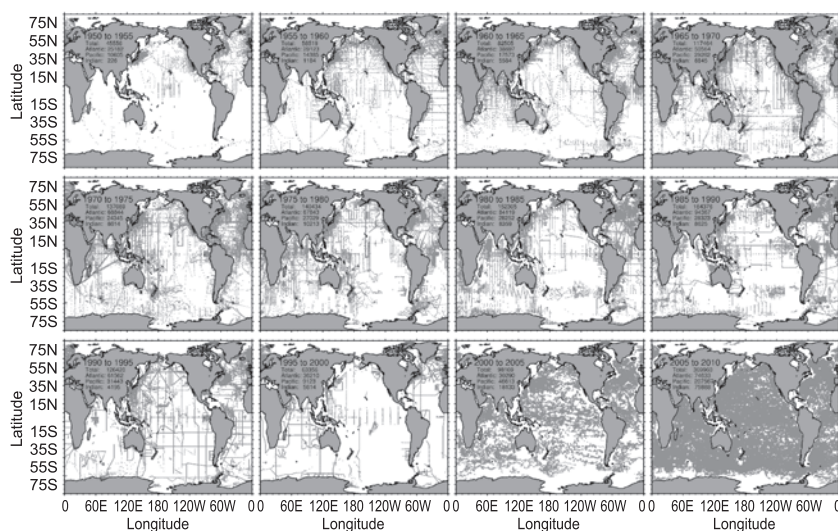
646

647 Robotic platforms, low power electronics and the possibility of small and capable sensors measuring a  
648 wide range of physical, chemical and biological parameters will allow for significant enhancements of  
649 the present-day sustained ocean observations. At the same time the improvement of ocean modelling  
650 capabilities and associated data assimilation techniques described in other chapters of this book will  
651 allow use of the full range of complementary *in-situ* and remotely sensed data to provide interpolated  
652 ocean information to an increasing range of marine applications. This wider view of ocean observations  
653 was a major driver for the OceanObs'09 Conference.

654

#### 655 4. Changes in data volume and coverage and implication for synthesis products

656 The diversity of in-situ observations and the enormous changes over time in their quantity, spatial  
657 distribution and precision create difficulties in producing the types of ocean syntheses that are required  
658 to assess global scale change in the oceans and to examine the significance of such changes for earth's  
659 climate.



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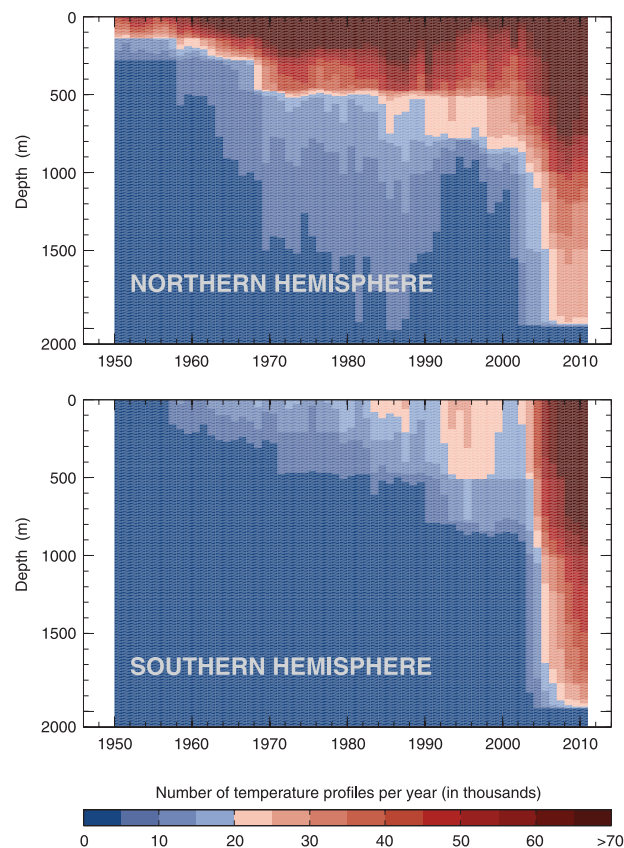
661 **Figure 12.** Number of ocean profiles from which data are available in 5 year temporal bins from 1950-  
662 1955 to 2005-2010. The increase in profiles in the last 10 years results from Argo..(From Durack,  
663 2011).

664 Figure 12 from Durack (2011) shows the number and spatial distribution of ocean profile data  
665 (measuring salinity) in 5 year time windows from 1950 to 2010. The North Atlantic is the only ocean  
666 with reasonable data coverage throughout the entire period – a result of the proximity of many  
667 laboratories in Europe and N. America with observational capability and in no small measure due to the  
668 long existence of the International Council for the Exploration of the Sea (ICES) (Went, 1972). Until  
669 the 1990s the coverage is due in large part to the accumulation of observations made in short-term  
670 regional experiments. The WOCE survey of the ocean is clearly visible as long coast-to-coast  
671 hydrographic sections between 1990 and 2000 (the WOCE data were collected between 1990 and  
672 1998). For many parts of the central South Pacific, Indian and Southern Oceans WOCE provided the

673 first ever observations of the ocean interior. From 2000 until present, we see the significant impact of  
674 Argo, in the first four years during its ramp-up phase that reached 1500 floats in 2004 and its design  
675 target of 3000 floats in 2007. Since 2007 no ice-free areas of the deep ocean have been devoid of  
676 measurements.

677 What the figure does not show is that there is also a distinct lack of wintertime observations  
678 particularly at high latitude. This results from ship-based operations having been preferentially  
679 scheduled in the summer season to avoid weather-related disruptions. This is true even for the  
680 relatively well-sampled north Atlantic where in an area bounded by 50-55°N and 35-45°W there are 4  
681 times as many salinity profiles in NOAA's World Ocean Data base in April through September than in  
682 the winter months.

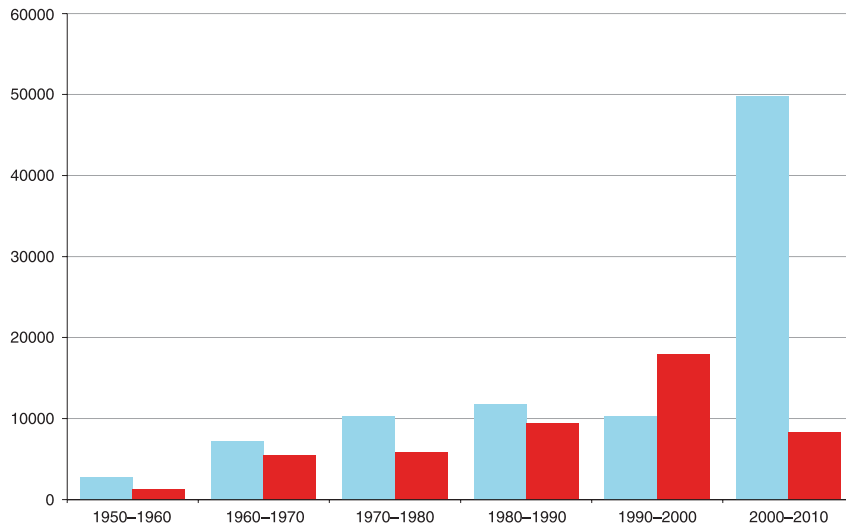
683 The evolution of the observing system has also resulted in significant changes in the depth distribution  
684 of observations (Figure 13, based on temperature profile measurements). The dominant data source in  
685 the early 1980s was the T4 (450 m) XBTs, these were then supplemented by 750 m (T7) probes and  
686 finally, post-2000, by Argo profiles.



687

688 **Figure 13.** The global number of temperature observations per month as a function of depth based on  
689 data from the UK Met Office's (EN3\_v2a) data set. (Ingleby and Huddleston, 2007).

690 While the full operation of the Argo program has led to an enormous increase in ocean observations  
691 from the sea-surface to 2000 m, the ocean below 2000 m is still poorly observed (Figure 14).



692

693 **Figure 14.** Number of salinity profiles per annum to 1000 m (blue) and to 4000 m per decade (red) in  
 694 NOAA World Ocean Database.

695 Here we see that the number of T/S profiles of the upper ocean (here defined as profiles to 1000m) has  
 696 increased from typically 10,000 per decade to 5 times that number (and with better temporal and spatial  
 697 distribution) thanks to Argo while the situation for the deep ocean remains poor. With the exception of  
 698 the decade of the 1990s where WOCE made a deliberate effort to sample the full water column, the  
 699 number of TS profiles to 4000m remains below 10,000 per decade (~3 per day!).

700 The fact that the data set of ocean observations is so heterogeneous, (changing since the 1950s from  
 701 discrete sampling levels, through continuous ship-based profiles to profiles from autonomous  
 702 instruments and with large areas unsampled for long periods) and with an ever-changing mix of data  
 703 sources presents enormous challenges to anyone trying to reconstruct the past state of the ocean and to  
 704 unambiguously determine how that state has changed. For that reason much of what is known about  
 705 ocean change has come from comparisons of re-occupations of trans-ocean sections (see for example  
 706 Wong, Bindoff and Church, 1999, Bryden, Longworth and Cunningham, 2005) or from the analysis of  
 707 ocean time series stations (e.g Joyce and Robbins, 1996). Here it is worth a note of caution. Each  
 708 observing system element applies calibration and quality assurance before archival of the data. The  
 709 critical analysis and interpretation of these quality assurance schemes is essential if analysis and  
 710 interpretation are to be rigorous. The individual observing systems each strive to maximize data  
 711 quantity and quality while delivering datasets as quickly and efficiently as is practical for each data  
 712 type. System interoperability in data formats, metadata protocols, and modes of data delivery is clearly  
 713 an enormous advantage but is not always achieved. There is also a need for data products, for example  
 714 gridded datasets with uncertainty estimates, in addition to the observational datasets. The  
 715 documentation and characterization of products and datasets is essential along with guidance on  
 716 suitability of datasets for a range of applications.

717 The synthesis and delivery of high-quality data and products for climate applications are major  
 718 undertakings that have historically been under-resourced. However, the production of these climate-  
 719 quality observational products is vital for assessing global ocean change and variability, data  
 720 assimilation model, initialization of climate and ocean only model and for the assessment of these  
 721 models (See Chapters 5.2, 5.3). As an example, calculations of the global ocean heat content and  
 722 freshwater (salinity) change over the observational period (Chapters 6.1 and 6.2) are derived from the  
 723 ocean temperature and salinity data. As we have noted, the temperature record is a synthesis of a  
 724 number of observation platforms – ship bucket, ship-based CTD, XBT and more recently Argo.

725 Between 1967 and 2001 the ocean temperature profile data was mainly comprised of XBT data (56%),  
726 but since the advent in 2001 of Argo these profiles have increasingly dominated the data record. The  
727 comparison of overlapping data sources allows the correction of instrumental shortcomings to be  
728 corrected as for example with the fall rate of XBT probes (Cowley et al. 2012).

729 Aside from the change in observations platform, there has been an evolution of the temperature scale  
730 definitions dating back to the early 1900s (Preston-Thomas, 1990) requiring care when considering the  
731 smallest temperature changes. Similarly, the methods for the determination of salinity and the  
732 definition of salinity have evolved over the observational period (Chapter 3.2).

733 Ocean data products commonly use the World Ocean Database 09 (WOD09) (Boyer et al., 2009,  
734 [http://www.nodc.noaa.gov/OC5/WOD09/pr\\_wod09.html](http://www.nodc.noaa.gov/OC5/WOD09/pr_wod09.html)) as their source of ocean observations. The  
735 WOD is an initiative of the National Oceanographic Data Center (NODC) and World Data Center for  
736 Oceanography (WDC) that collates ocean observations submitted to NODC/WDC by individual  
737 scientists and observation programs (i.e. Argo, global XBT, drifter and GO-SHIP), and national and  
738 regional data centers. WOD09 provides a centralized quality-controlled database of all ocean  
739 observations and metadata by observational platform. This database is used in the production of  
740 observational-based, objectively mapped ocean climatologies (e.g. World Ocean Atlas, 2009  
741 ([http://www.nodc.noaa.gov/OC5/WOA09/pr\\_woa09.html](http://www.nodc.noaa.gov/OC5/WOA09/pr_woa09.html)) and CSIRO Atlas for Region Seas  
742 (<http://www.marine.csiro.au/~dunn/cars2009/>)). Ocean reanalysis models (e.g. SODA, ECCO-  
743 GECCO, and others) combine ocean observational data and an ocean general circulation model driven  
744 by known forcing (wind, surface forcing) to reconstruct an ocean climatology that is consistent with the  
745 observational record and dynamically balanced. Data assimilation methods recognize that the  
746 historical data is sparse and apply a general circulation model to reconstruct the time evolving ocean  
747 properties and circulation. Ocean reanalyses are used for short-term (seasonal) to decadal ocean  
748 forecasting efforts. Many of these products employ additional data quality control steps further  
749 improving the reliability of the database (<http://www.icdc.zmaw.de/wohp.html>).

750 Such ocean climatologies are a great help in the documentation of temporal change since they provide  
751 baselines against which even sparse data sets may be compared. A recent example of such a  
752 comparison is that between the *Challenger* data from the 1870s and the recent Argo climatology  
753 (Roemmich, Gould and Gilson, 2012).

## 754 5. The Future –outstanding issues and a new framework for global ocean observing

755

756 We have explored the evolution of *in-situ* ocean observations and have noted that the rapid advances of  
757 the past two decades have brought us to the point where a sustained observing system can be envisaged  
758 that largely meets the needs of a wide range of users. Earth's climate has entered a phase in which  
759 impacts are increasingly attributable to human activities (sometimes referred to as the anthropocene)  
760 and this has greatly increased the socio-economic as well as scientific justification for ocean  
761 measurements. Ocean observations have shown that more than 90% of the extra heat energy stored by  
762 the Earth in the last 50 years is found in the oceans and that ocean salinity is a direct monitor of the  
763 global hydrological cycle (Solomon et al 2007, Durack, 2011). It is, therefore critically important that  
764 the observing system is capable of detecting and documenting global climate change so that policy  
765 makers (and the general public) can have access to climate observations and products to assess the  
766 present state of the ocean, cryosphere, atmosphere, and land and place them in context with the past. To  
767 be of both societal and scientific value, these observations need to be sustained over many decades and  
768 be of a quality that is adequate to address present-day concerns and those that may be of concern in the  
769 future.

770 Ocean observations are also needed to initialize and evaluate climate models and to improve  
771 predictions of climate change. Such assessments are essential for guiding national and international

772 policies that relate to resources (such as fisheries, agriculture and water supply) that are impacted by  
773 climate variability and change and efforts aimed at mitigating long-term climate change. The quality of  
774 the climate services provided to the public and policy makers is founded on advances in fundamental  
775 research for which comprehensive observations are also needed and these go beyond physics to include  
776 the mechanisms involved in ocean acidification, ocean deoxygenation and the loss of marine  
777 biodiversity. All require globally consistent, regionally coordinated and systematic ocean observations  
778 of climate-relevant properties.

779

780 In order to guide the development and implementation of the ocean observing system, a framework is  
781 required that integrates the present rather independent networks described earlier – Argo, OceanSites,  
782 GO-SHIP, XBT, Global Drifter Program, Sea Level (GLOSS). Currently these elements fall under the  
783 auspices of several programs each with their own priorities and agendas (Climate Variability and  
784 Predictability Experiment (WCRP/CLIVAR), Global Observing System (GOOS) as part of the Global  
785 Climate Observing System (GCOS) which is in turn an element of the Global Earth Observing System  
786 of Systems (GEOSS)). While all are fundamental elements of the Ocean Observing System (see figure  
787 15 for an artist’s view).

788

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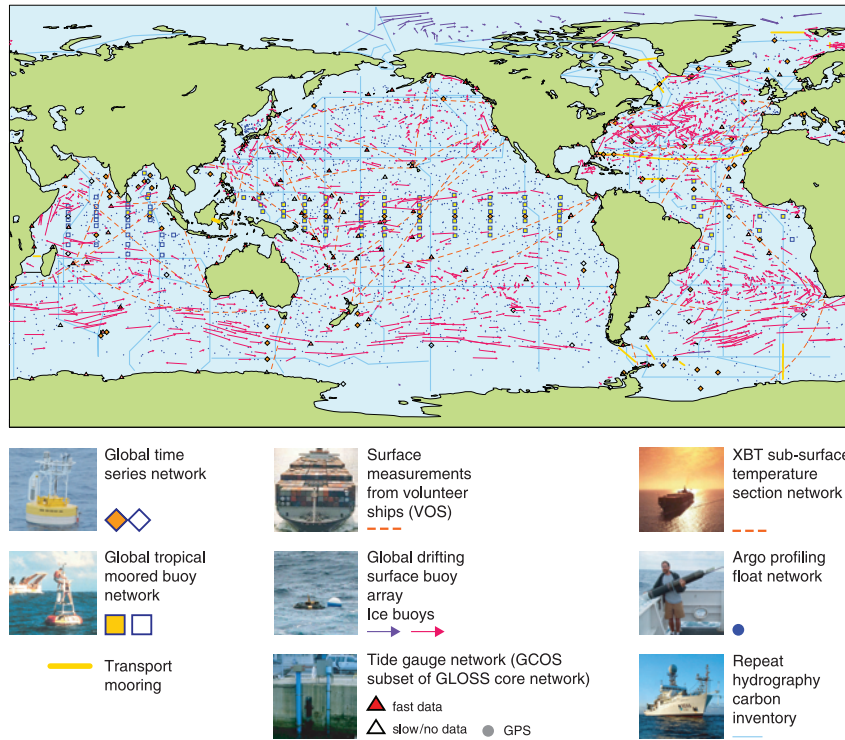
790

791 **Figure 15.** A vision the integrated Ocean Observing system incorporating many different observation  
792 platforms to provide the sustained high-quality routine observations of the ocean. (Image from GOOS)

793

#### 794 5.1 Building on OceanObs’09

795 Since the OceanObs’99 conference tremendous progress has been made towards deploying a truly  
796 global ocean observing system. Some of the elements of the *in situ* observing system have managed to  
797 reach their design specification (Argo and the Global Drifter Program), others still have room for  
798 improvement. The status of the present ocean observing system (Figure 16) and community  
799 recommendations for its enhancement were reviewed by the OceanObs’09 Conference and related  
800 activities were set in place. (Hall, Harrison and Stammer, 2010,  
801 <http://www.oceanobs09.net/proceedings>).



803

804 **Figure 16.** The global ocean observing system is comprised of numerous observing platforms that  
 805 provide sustained ocean data from the sea surface to the abyssal ocean at varying temporal and spatial  
 806 resolution. This figure shows the observing system status on May 2013.

807 Despite some significant inadequacies the present system has allowed us to produce estimates of  
 808 quantities that are essential for monitoring the oceans’ role in climate (heat, freshwater and carbon  
 809 storage, upper ocean circulation and, with less confidence, changes in the ecological cycle). From  
 810 these estimates and from our knowledge of the climate system it is now possible to identify a set of  
 811 Essential Ocean Variables (EOVs). (Table 1) Some of them are already included in the list of the  
 812 Essential Climate Variables (ECVs) defined by the Global Climate Observing System (World  
 813 Meteorological Organisation, 2010), while others have to be confirmed and specified in detail. Once  
 814 the community has agreed to a list, the intention will be for EOVs to be monitored in a sustained and  
 815 consistent manner and with adequate precision for climate applications. The most cost effective  
 816 network arrangements have to be determined and financial and governance arrangements have to be  
 817 developed.  
 818

|   |  |
|---|--|
| <p><b>Near surface atmospheric variables</b></p> <ul style="list-style-type: none"> <li>Air temperature</li> <li>Precipitation</li> <li>Atmospheric pressure</li> <li>Surface radiation budget</li> <li>Wind speed/direction</li> <li>Water vapour</li> </ul> | <p><b>Ocean surface variables</b></p> <ul style="list-style-type: none"> <li>Sea surface temperature.</li> <li>Sea surface salinity</li> <li>Sea level</li> <li>Sea state</li> <li>Sea ice coverage</li> <li>Current speed/direction</li> <li>Ocean colour (biological activity)</li> <li>Carbon dioxide (<math>pCO_2</math>)</li> </ul> |
|   | <p><b>Ocean subsurface variables</b></p> <ul style="list-style-type: none"> <li>Temperature</li> <li>Salinity</li> <li>Current speed/direction</li> <li>Nutrients</li> <li>Carbon</li> </ul>   |

819

820 **Table 1.** The subset of the Essential Climate Variables that may comprise the Essential Ocean  
 821 Variables. Sustained observations of these variables are required for the generation of global climate  
 822 products.

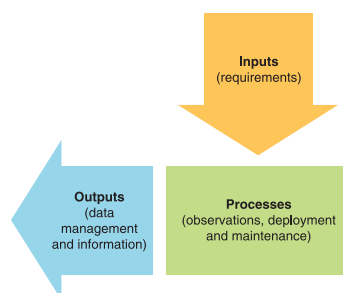
823 A key development between the two OceanObs conferences was an emphasis in 2009 on  
 824 interdisciplinary ocean observations and a call for more effective international coordination and tighter  
 825 integration between observing activities and the data and information product delivery. In response to  
 826 the new challenges, a systems approach to ocean observations (Figure 17) was developed and  
 827 encapsulated in the Framework for Ocean Observing (FOO, <http://www.oceanobs09.net/foo/>). The  
 828 FOO concept starts with societal drivers and the demands these generate for ocean observations and  
 829 include:

- 830 • The need to document ocean change (measuring the responses to climate change, overfishing  
 831 and pollution)
- 832 • Initializing ocean models for climate predictions (e.g. El Niño, Tropical Atlantic Variability,  
 833 Indian Ocean Dipole and their respective impacts on monsoon systems and decadal  
 834 predictability)
- 835 • Initializing short-term ocean forecasts for marine operations (e.g. oil spill and pollution  
 836 tracking, search-and-rescue)
- 837 • Regulatory matters of coastal states (e.g. Climate Change Convention, Convention of  
 838 Biodiversity, Marine Spatial Planning and associated demands).

839 The Framework would guide the whole ocean observing community and be organized around a set of  
 840 “Essential Ocean Variables (EOVs); an approach shown by GCOS to break down barriers to  
 841 cooperation amongst funding agencies and observing networks.

842 Implementation would be guided by the level of “readiness” with immediate implementation of  
 843 components that have already reached maturity while encouraging innovation and capacity building for  
 844 less mature observation streams and methods.

845 By taking a systems engineering approach, the FOO input requirements (observations) will be  
 846 identified as the information needed to address a specific scientific problem or societal issue. The  
 847 societal issues span from short-timescale needs such as hazard warning to such long-timescale needs as  
 848 knowledge of ecosystem limits appropriate to the sustainable exploitation of ocean resources. The  
 849 mechanisms to deliver these observation elements will then be identified in terms of technologies and  
 850 observing networks. The outputs (data and information products) will consist of the most appropriate  
 851 syntheses of ocean observation streams to provide services, address scientific problems or permit  
 852 informed decisions on societal issues.



853

854 **Figure 17.** A systems approach to ocean observations, the guiding principle of the Framework for  
855 Ocean Observations.

856 The vastness, remoteness, and harshness of the oceans means that collecting any *in situ* observations is  
857 expensive. As a consequence, observing systems have been and will continue to be designed to  
858 measure as many variables as possible so as to take full advantage of the limited number of observing  
859 platforms. These multiple sensors place demands on energy and thus a focus for FOO will be the  
860 avoidance of duplication between observing platforms and networks. However, the complementarity  
861 of observing networks (for instance between Argo and ship-based CTD observations) has enormous  
862 benefits in allowing intercalibration and eliminating systematic bias. Common standards for data  
863 collection and dissemination of EOVS data will be adopted so as to maximize the utility of data.

864 The Framework approach will be used to encourage partnerships between the research and operational  
865 communities so as to assess and improve the readiness levels of observation elements and data systems  
866 appropriate for each EOVS. Similar partnerships will refine requirements. The Framework should also  
867 enhance collaboration between developed and developing regions and promote the use of common  
868 standards and best practices.

869 In summary the Framework will promote a more consistent and integrated approach to the assessment  
870 of readiness, implementation and setting standards for information sharing among the varied and  
871 largely autonomous observing elements. It should also lead to a well-defined set of requirements and  
872 goals, facilitate coordination between observing system elements, streamline implementation of  
873 sustained global-scale observations by applying a systems engineering approach and identifying best  
874 practices.

## 875 6. Conclusions

876  
877 Since the early 1980s and 90s enormous progress has been made in gathering *in situ* ocean  
878 observations at global scales that complements information from ocean-focussed earth observing  
879 satellites. Building on global-scale research programs and stimulated by the impetus of the 1999  
880 OceanObs Conference, the international community has implemented the first elements of what may  
881 become a sustained *in situ* global ocean observing system. Many ocean and climate analysis and  
882 forecasting centres are now dependent on measurements and derived products provided by the present  
883 observing system. This period of development has coincided with a growing awareness, not only by  
884 scientists but by some politicians and a large part of the general public of the need to monitor the  
885 oceans in order to quantify the impacts and progress of climate change and of the impacts of pollution  
886 and exploitation of the oceans' living resources.

887  
888 The second OceanObs conference in 2009 built on this progress but also expanded the horizons of  
889 systematic ocean observing from its earlier largely physical and science-based perspective to one that  
890 recognised the lack of emphasis on systematic biological and biogeochemical observations and that  
891 recognised the need to underpin the case for sustained observations by reference to societal drivers.

892  
893 The coming decade will see a more systematic approach to the sustained observing of the ocean than  
894 has been common in the past decades through the developing Framework for Ocean Observations.  
895 This Framework has the potential to optimise the assembly of data from multiple and diverse platforms  
896 and eventually to optimise observing system design. If we continue to follow the example of the past  
897 two decades, innovative observation techniques made possible by experimental technologies will be  
898 incorporated into the mainstream of ocean observations.

899  
900 Of course, any future development, expansion and global-scale implementation of a sustained  
901 observing system will require commitment of funds from national and multi-national sources and there



902 will be competition for those funds within the ocean observing community and between the physical,  
903 biological and solid-earth science communities. Success in competing for these fund will require a  
904 clear demonstration of the benefits, not just to the work of the ocean and climate science communities  
905 but also to societal issues such as better predicting the likely progress of sea-level rise and the oceans'  
906 role in predictions of seasonal climate.

907  
908 While our ability to observe the oceans and to understand their role in the earth's climate system have  
909 advanced dramatically, national and international oversight structures and funding streams have  
910 changed much less – only a slowly growing number of governments or agencies are willing to  
911 acknowledge the need for a long-term (~decades) commitment of funds to observing programmes and  
912 networks.

913  
914 While these constraints may hinder progress, the dedication, persistence and innovative nature of the  
915 observational marine science community that has already made such remarkable progress might be  
916 expected to overcome such obstacles.

917  
918 The highest priority for the coming decade must be to sustain the present ocean observing system,  
919 while improving its coverage and data quality. The system can also be significantly enhanced by the  
920 following extensions of existing elements and by integration across elements:

- 921
- 922 • The sampling domain of autonomous platforms can become truly global through extensions to  
923 higher latitude, into marginal seas and the deep ocean, and through higher resolution  
924 observations in boundary current regions. Incremental technology developments and  
925 definition of new sampling requirements are needed for these extensions.
  - 926 • Multi-decadal ocean warming and ocean acidification have impacts on marine ecosystems  
927 with severe socio-economic consequences. Given the value of ocean ecosystems to human  
928 health and welfare, it is important to understand the links between ocean and climate  
929 variability, marine chemical process and their impact on marine ecosystems. Thus, there is an  
930 urgent need to *fully integrate biogeochemical and biological observations* into the ocean  
931 observing system.
  - 932 • The global network measuring the physical state of the oceans provides a platform for multi-  
933 disciplinary observations of biogeochemical and ecosystem impacts of climate change. Key  
934 requirements are further developments in low-power sensor accuracy and stability, and  
935 effective integration between autonomous and shipboard observational networks (e.g.  
936 definition of core variables; ensuring a sufficient quantity of reference-quality data for quality  
937 assurance of autonomous sensors).
  - 938 • Improvements in the observation of the ocean surface layer and of air-sea exchanges require  
939 better utilization of research vessels and commercial shipping, improvements to automated  
940 measurement systems, better coordination across networks, and a review of sampling  
941 requirements for marine meteorology and ocean surface velocity.
  - 942 • Strong commitment to preserve continuity, or in some cases like satellite measurements of the  
943 air-sea momentum flux from scatterometers and variations in the ocean mass field from  
944 gravity satellites, reinstate measurement missions. The principal challenge remains to  
945 advocate, plan and finance and press for executing the transition of the critical satellite sensors  
946 to sustained status.

947  
948 A major effort is needed to ensure that data quality is maximized, that data access is simplified  
949 (including for data types extending across multiple observational networks), and that data products are  
950 useful and available. All measurements need to be documented, calibrated and stored in internationally  
951 accessible data base systems. Most of the raw data from different observing networks will be combined  
952 using a range of techniques from simple statistical tools to full 4DVAR ocean data assimilation  
953 (Chapters 5.2 and 5.3) to produce ocean information products that are capable of addressing the user

954 requirements outlined above. This information flow should be implemented by an integrated network  
955 of data information centers, data assimilating systems and a number of routine and near real time  
956 assessments of key ocean quantities with adequate time and space resolution.  
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