

Increased atmospheric ammonia over the world's major agricultural areas detected from space

J. X. Warner¹, R. R. Dickerson¹, Z. Wei¹, L. L. Strow², Y. Wang³, and Q. Liang^{4,5}

¹Department of Atmospheric and Oceanic Science, University of Maryland College Park, College Park, MD 20742, U.S.A.

²Department of Physics and Joint Center for Environmental Technology, University of Maryland Baltimore County, Baltimore, MD 21250, U.S.A.

³Department of Earth and Atmospheric Sciences, University of Houston, Houston, TX 77204, U.S.A.

⁴NASA Goddard Space Flight Center, Atmospheric Chemistry and Dynamics, Greenbelt, MD, U.S.A.

⁵Universities Space Research Association, GESTAR, Columbia, MD, USA.

Corresponding author: Juying Warner (juying@atmos.umd.edu)

Key Points:

- First decade-long ammonia records (2002 – 2016) were retrieved from AIRS satellite daily measurements.
- Substantial increases in ammonia concentrations are observed over several of the world's major agricultural regions.
- Causes of ammonia increase include increased fertilizer use, increasing temperatures, and decreased loss to aerosols.

Abstract

This study provides evidence of substantial increases in atmospheric ammonia (NH_3) concentrations (14-year) over several of the world's major agricultural regions, using recently available retrievals from the Atmospheric Infrared Sounder (AIRS) aboard NASA's Aqua satellite. The main sources of atmospheric NH_3 are farming and animal husbandry involving reactive nitrogen ultimately derived from fertilizer use; rates of emission are also sensitive to climate change. Significant increasing trends are seen over the US ($2.61\% \text{ yr}^{-1}$), the European Union (EU) ($1.83\% \text{ yr}^{-1}$), and China ($2.27\% \text{ yr}^{-1}$). Over the EU, the trend results from decreased scavenging by acid aerosols. Over the US, the increase results from a combination of decreased chemical loss and increased soil temperatures. Over China, decreased chemical loss, increasing temperatures, and increased fertilizer use all play a role. Over South Asia, increased NH_3 emissions are masked by increased SO_2 and NO_x emissions, leading to increased aerosol loading and adverse health effects.

Index terms

Constituent sources and sinks; Troposphere: composition and chemistry; Pollution: urban and regional

Keywords

Ammonia trends; NH_3 ; ammonium aerosols; SO_2 ; NO_2 ; emission

53

54

55 1 Introduction

56 Atmospheric ammonia (NH_3) is an important component of the global nitrogen cycle
57 [Galloway *et al.*, 2002, 2008; Sutton *et al.*, 2007, 2008; Erisman *et al.*, 2008, 2013; Fowler *et al.*,
58 2013, 2015]. In the troposphere ammonia reacts rapidly with acids such as sulfuric (H_2SO_4),
59 nitric (HNO_3) to form fine particulate matter ($\text{PM}_{2.5}$) [Malm *et al.*, 2004]. These ammonium
60 (NH_4^+) containing aerosols affect Earth's radiative balance, both directly by scattering incoming
61 radiation [Adams *et al.*, 2001; Martin *et al.*, 2004; Henze *et al.*, 2012] and indirectly as cloud
62 condensation nuclei [Abbatt *et al.*, 2006]. $\text{PM}_{2.5}$ endangers public health by penetrating the
63 human respiratory systems, depositing in the lungs and alveolar regions [Pope *et al.*, 2002], and
64 causing premature mortality [Lelieveld *et al.*, 2015]. A precursor of these inorganic aerosols,
65 gaseous NH_3 is often the limiting species in their formation [Wang *et al.*, 2013; Lelieveld *et al.*,
66 2015]. Excess reactive nitrogen reduces biodiversity and causes harmful algal blooms and anoxic
67 conditions. Dry deposition of gaseous ammonia may have substantially greater adverse impacts
68 on ecosystem health than deposition of ammonium in aerosols or precipitation [Sheppard *et al.*,
69 2011]. In contrast, $\text{PM}_{2.5}$ has greater impact on human morbidity and mortality. In this article we
70 quantify recent (~14-year) increases in tropospheric ammonia and suggests likely causes for
71 these trends.

72 Major sources of atmospheric ammonia involve agricultural activities including animal
73 husbandry, especially concentrated animal feeding operation, and fertilizer use [Streets *et al.*,
74 2003; Huang *et al.*, 2012; Hauglustaine *et al.*, 2014; Riddick *et al.*, 2016]. Ammonium fertilizers
75 are essential in high-yield crop production, and contribute substantially atmospheric NH_3 .
76 Fertilizer usage in China (~31.2 TgN yr⁻¹ and ~+2.7% yr⁻¹) and India (~18.8 TgN yr⁻¹ and
77 ~+3.6% yr⁻¹) has increase several fold in the last two decades, from Earth Policy Institute
78 (http://www.earth-policy.org/data_highlights/2014/highlights43) and according to the World
79 Bank (<http://data.worldbank.org/indicator/AG.CON.FERT.ZS>). It is estimated that 50% of the
80 total NH_3 emission in 2000 in China came from fertilizer application and another 38% from
81 other agricultural sources [Streets *et al.*, 2003]. Ammonia emissions increase with increasing
82 nitrogen content and pH of soils and manure storage facilities, and increase exponentially with
83 temperature (emissions roughly double between 300 and 306 K), except below freezing when
84 emissions are near zero [Riddick *et al.*, 2016]. A minimum level of soil moisture is also required
85 for the microbial activities, such as urea hydrolysis, that generate NH_3 . Biomass burning, highly
86 episodic in nature, accounts for <10% of the global total, but can be a locally important source
87 [Dentener and Crutzen, 1994; Roelle and Aneja, 2002; Galloway *et al.*, 2004].

88 Major sinks of atmospheric ammonia involve dry deposition and wet removal by
89 precipitation, as well as conversion to particulate ammonium by reaction with acids. These acids
90 arise primarily from the oxidation of pollutants SO_2 and NO_x ($\text{NO} + \text{NO}_2$) generated in the
91 combustion of fossil fuels. Ammonium sulfate is generally removed by precipitation. Condensed
92 ammonium nitrate (NH_4NO_3) exists in equilibrium with NH_3 and gas-phase HNO_3 . Lower
93 temperatures favor the aerosol phase.

94 Measurements of ambient NH_3 are sparse, but satellites provide a means to monitor
95 atmospheric composition globally. Through recent improvements in retrieval algorithms, the
96 Atmospheric Infrared Sounder (AIRS) aboard NASA's Aqua satellite now provides daily global
97 measurement of atmospheric NH_3 . Warner *et al.* [2016] described global NH_3 concentrations
98 using the averaged 13-year satellite data record (2003-2015) from AIRS and provided a global
99 perspective on its emissions, distributions, and spatial variability. They also discussed the

100 retrieval algorithm, preliminary validation, and qualitative comparisons to measurements from
101 other sensors. In this study, we focus on the NH₃ temporal variability, or trends, from September
102 2002 to August 2016 and discuss possible mechanisms underlying these trends. These AIRS NH₃
103 retrievals have greater daily coverages and a longer record than those from the Tropospheric
104 Emission Spectrometer (TES) [Beer et al., 2008]; and are based on higher channel sensitivities,
105 due to the afternoon overpasses, than the Infrared Atmospheric Sounding Interferometer (IASI)
106 [Clarisse et al., 2009]. Van Damme et al. [2015] showed six-year time series of NH₃ total column
107 values over six regions of the world from IASI's early morning (9:30am local time overpass) and
108 evening (9:30pm local time overpass) measurements. Whereas they identified the relatively large
109 emission peaks in the time series as resulting from biomass burning events, their study did not
110 indicate clear increasing or decreasing trends. Schiferl et al. [2016] used a combination of
111 observations (including IASI) and a model to evaluate variability in NH₃ over the US and
112 concluded that variability in meteorology and reduced SO₂ and NO_x emissions drive NH₃
113 changes observed between 2008 and 2012.

114 In Section 2, we describe the methods and data used in the analyses, and in Section 3, we
115 present global ammonia trends. In Section 4, we focus on the ammonia trends in the primary
116 regions of interest and discuss the driving mechanisms.

117 **2 Methods and data**

118 **2.1 AIRS NH₃ VMRs**

119 Warner et al. [2016] discussed in detail the AIRS NH₃ retrieval method, quality
120 assurance, global NH₃ distributions, and preliminary validation. We applied additional thresholds
121 for the trend computations. We used only NH₃ data with Degree Of Freedom for Signal (DOFS)
122 greater than 0.1, in addition to other retrieval quality assurance flags (e.g., χ^2 , retrieval residual,
123 cloud-cleared-radiance quality flags, etc.). All retrieval results were screened by a minimum
124 thermal contrast determined by AIRS L2 products. If the lower layers of the atmosphere and
125 Earth's surface have similar temperatures, (low thermal contrast), they emit similar amounts of
126 thermal radiation [Deeter et al., 2007], and AIRS cannot quantify NH₃ in these layers. While
127 AIRS NH₃ products are outputted at multiple levels from 500 hPa to the surface, here we use
128 NH₃ VMRs at 918 hPa, where the peak sensitivity is, for this study.

129 For the seasonal cycles, we used a 7-day average for each region and applied an n-point
130 smoothing [Garcia, 2010]. We averaged the seasonal cycles into three periods: from 2002
131 through 2008, from 2009 through 2013, and from 2014 through 2016, with the mean and the 1- σ
132 standard deviations. Since we only included NH₃ concentrations from frequent sources with
133 elevated NH₃ VMRs, the values shown in all figures maybe higher than the average
134 concentrations in a region.

135 **2.2 Meteorological data sources**

136 As meteorological conditions influence the rate of ammonia emission and deposition, we
137 examine surface skin temperature and total precipitation anomalies using European Centre for
138 Medium-Range Weather Forecasts (ECMWF) era-interim reanalysis (EI) [Berrisford et al.,
139 2011; Dee et al., 2011]. The precipitation anomalies were computed using the 12-hour forecast
140 accumulated in each month for each ECMWF grid (0.75°x0.75°). Only grids containing AIRS
141 NH₃ retrievals were considered.

142 For surface skin temperatures anomalies and trends, we selected ECMWF EI daytime
143 data to match the NH₃ daytime product by only using the model outputs between 9am and 3pm
144 local times. Skin temperatures trends were based on daily means and only included the cases
145 with NH₃ retrievals. Additionally, to avoid partial year trends, we use only skin temperatures for
146 the period from March to August when it is sufficiently high to be most relevant to the Northern
147 Hemisphere NH₃ emissions. The linear fits of surface skin temperatures, however, are not
148 statistically significant as indicated by high p-values; they are discussed as references only. We
149 used the published ammonia temperature dependence [*Dentener and Crutzen, 1994; Galloway et*
150 *al., 2004; Riddick et al., 2016*], with the expression: $E_2 / E_1 = \{ \exp [-10380 (1/T_2 - 1/T_1)] - 1 \} * 100\%$,
151 where T_1 is assumed to be 300K, T_2 is T_1 plus the observed annual temperature increase,
152 E_1 and E_2 are NH₃ volume mixing ratios corresponding to T_1 and T_2 .

153 2.3 Thermal contrasts

154 Remote sensing measurement sensitivities depend on surface thermal contrasts of the
155 target areas, and in the case of NH₃, higher thermal contrast generally results in a higher
156 retrieved NH₃ concentrations. Surface thermal contrasts are defined as the differences between
157 surface skin temperatures and surface air temperatures, however, we approximate the ECMWF
158 2-meter air temperatures as surface air temperatures. While sufficient thermal contrast is needed
159 for good signal-to-noise ratio, the influence of thermal contrast on the retrieved NH₃
160 concentrations and variability needs to be addressed, especially in trend related studies. When we
161 examine retrievals using separate ranges of thermal contrasts in 2°K degree increments, we
162 found that, although the rates of increase/decrease in NH₃ are different in each thermal contrast
163 range especially during winter season, the tendencies and the magnitudes are similar.
164 Furthermore, the 14-year thermal contrast from ECMWF EI in the regions of our study shows
165 slight decreases (i.e., -0.016°K yr⁻¹, p=0.002, for the US Midwest, -0.021°K yr⁻¹, p=0.000, for the
166 EU, -0.011°K yr⁻¹, p=0.053, for China, and -0.069°K yr⁻¹, p=0.000, for South Asia). This
167 indicates that the increasing trends of the NH₃ concentrations are not the results of the thermal
168 contrast increasing, since the thermal contrast has decreased. We used thermal contrast daily
169 means in March to August for trend computations and only included regions where there are
170 NH₃ pixels.

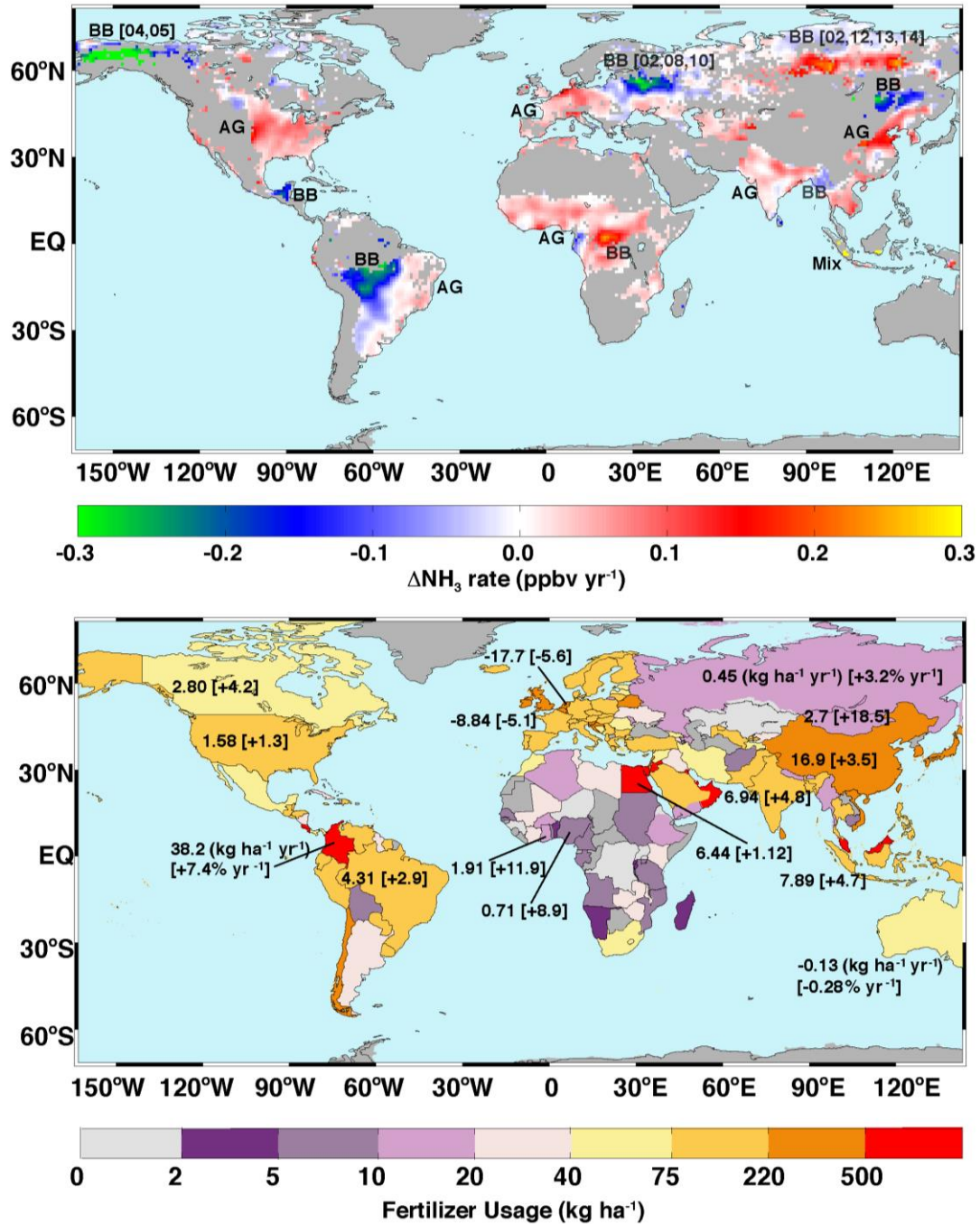
171 2.4 OMI SO₂ and NO₂

172 We examine ammonia trends jointly with SO₂ and NO₂ changes to determine the
173 scavenging by acid aerosols. We used OMI Level-3 SO₂ planetary boundary layer Volume
174 Column Density (VCD, in DU) spanning October 2005 through August 2016
175 (http://disc.sci.gsfc.nasa.gov/Aura/data-holdings/OMI/omso2e_v003.shtml). Strong volcanic
176 emissions are removed from the dataset by examining the daily region-wide 99.9-percentile of
177 SO₂ VCDs. If the percentile is found to exceed a threshold value (US 5 DU, Europe 8 DU, China
178 10 DU, India 8 DU), all data from that day were excluded [*Krotkov, et al., 2016*]. Note that there
179 are still a few large jumps of SO₂ values associated with volcanic eruptions (e.g., in 2006 and
180 2008), but volcanoes with a large spatial impact tend to send SO₂ into the upper troposphere
181 where there is little NH₃.

182 We used OMI cloud-screened tropospheric column NO₂ (molecules/cm²) datasets
183 similarly to SO₂. As for meteorological variables, SO₂ and NO₂ concentrations were averaged
184 only in the areas where NH₃ retrievals were used. For OMI SO₂, the trends are only significant at

185 95% confidence level over China. The OMI NO₂ trends are significant at 95% confidence level
 186 for the Western US and the EU. The SO₂ concentrations over the US are often below OMI
 187 detection limit, and therefore, not used.

188 **3 Observed global ammonia trends**



189 **Figure 1. Trend in AIRS NH₃.** Top: Temporal trends (+ 0.3 to - 0.3 ppbv yr⁻¹). The locations and the last two
 190 digits of the years of biomass burning events (in square brackets) are marked as BB. The main agricultural regions
 191 are marked as AG. Bottom: national averaged annual N fertilizer usage in 2002 – 2013 in kg ha⁻¹ and trends
 192 [percentage changes] in kg ha⁻¹ yr⁻¹ [% yr⁻¹]; see text for details.

193 Figure 1 top panel depicts a global map of the rate of change of NH₃ volume mixing ratio
 194 (VMR) in parts-per-billion by volume (ppbv) per year computed using linear regression of daily
 195 mean values in each 1°x1° latitude-longitude grid cell. We used daily mean values in each grid to
 196 obtain Fig. 1 top panel. We further smoothed the results using a 2-dimensional penalized least
 197 squares method allowing fast smoothing of data in one and higher dimensions by means of the
 198 discrete cosine transform [Garcia 2010]. In addition to the general quality assurance described in
 199 Methods and Data, we added constraints of only including grids with at least 10% of the pixels
 200 greater than 2 ppbv and records longer than 10 years. The years and locations of the biomass
 201 burning (BB) events are determined from Moderate Resolution Imaging Spectroradiometer
 202 (MODIS) fire products [Giglio et al., 2010] and shown on Fig. 1 top. The main agricultural (AG)
 203 regions are determined from Friedl et al. [2010]. The years of the regular fires are not shown.
 204 The fertilizer information (i.e., annual usage, trends between 2002 and 2013, and percentage
 205 trends) from the World Bank (<http://data.worldbank.org/indicator/AG.CON.FERT.ZS>) is plotted
 206 on Fig. 1 bottom.

207 AIRS reveals both increases and decreases over disparate parts of the globe. Biomass
 208 burning related decreases are seen over Alaska, the central district of Russia and Eastern Europe,
 209 Mongolia, Inner Mongolia and NE China, the Yucatan of Mexico, and the Amazon in western
 210 Brazil; increases are seen over Siberia and Indonesia. These large fire events are highly episodic,
 211 often driven by one or two outlier years (see Fig. 1 top), and are not statistically robust. We focus
 212 on regions where such events have minimal influence. Significant NH₃ increases are seen over
 213 the American Midwest and southern California (SoCA) (US), east central China, the European
 214 Union (EU) countries (e.g., The Netherlands, Germany, Denmark, and Po Valley, Italy), and
 215 parts of South Asia (i.e., Bangladesh, India, Pakistan, Cambodia, and Viet Nam), South America
 216 (Brazil, Colombia, Ecuador, and parts of Peru), and central Africa (Nigeria, Ghana, Sierra Leone,
 217 and Guinea), the Nile Delta of Egypt, and Fergana Valley, Uzbekistan. Regions with increasing
 218 trends are generally associated with anthropogenic emissions due to intense agricultural (e.g.,
 219 related to NH₃ emissions) and changing acid precursor (SO₂ and NO_x) emissions.

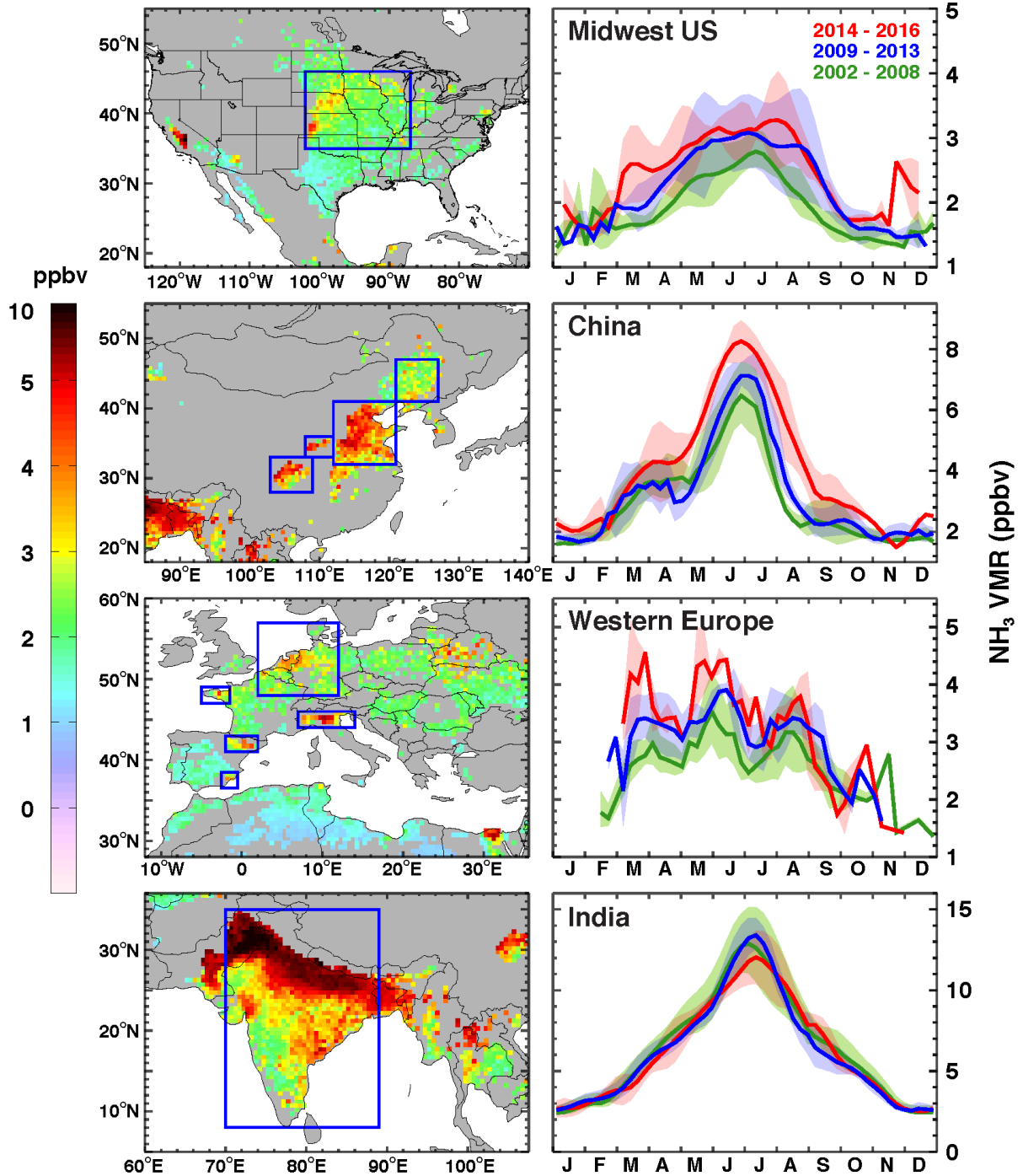
220 Parts of Brazil and parts of Africa have seen substantial increases in ammonia
 221 concentrations, but trends in the tropics are complicated by their proximity to major areas of
 222 biomass burning, a highly variable source of NH₃. Fertilizer use increase in 2002 – 2013, e.g.,
 223 4.3 kg ha⁻¹ yr⁻¹ (+2.9% yr⁻¹) over Brazil and 0.71 kg ha⁻¹ yr⁻¹ (+8.9% yr⁻¹) over Nigeria (see Fig.
 224 1 bottom), can largely explain the NH₃ increase over these regions.

225 **4 Regional ammonia trends and their driving mechanisms**

226 Figure 2 left panels show 14-year mean NH₃ VMRs for four regions with intense
 227 agricultural activities, the US, China, the EU, and South Asia, where the blue boxes in each
 228 panel outline the areas used in the trend computations. The underlying maps were selected using
 229 previously defined frequent occurrences of elevated NH₃ concentrations and good measurement
 230 sensitivities of AIRS [Warner et al., 2016]. Highest concentrations were observed over densely
 231 populated and heavily farmed South Asia followed by Northeast China where a high percentage
 232 of land is used for fertilized crops [Huang et al., 2012].

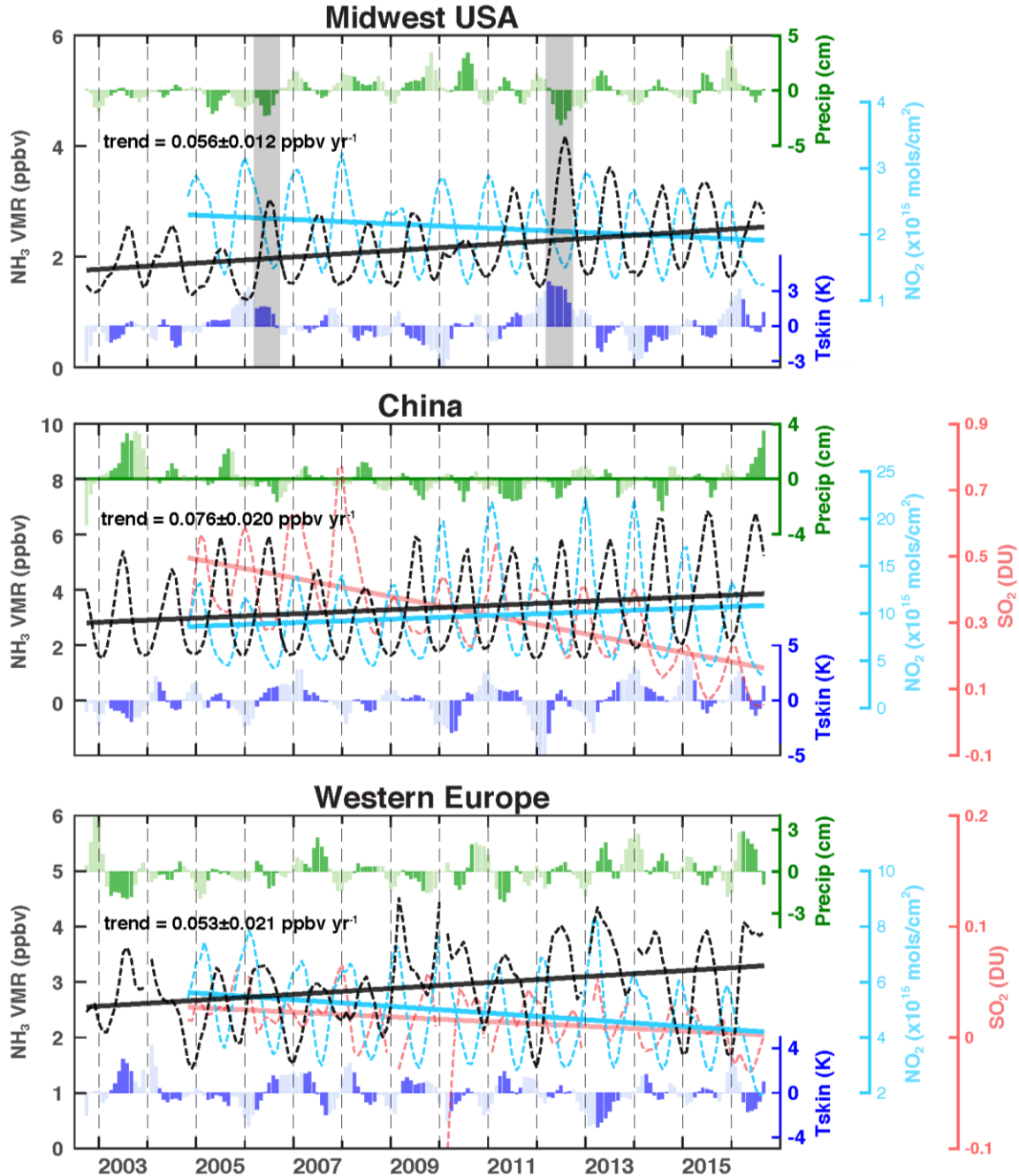
233 AIRS trend analysis (Fig. 3) shows that NH₃ concentrations have increased over the US
 234 by 0.056 ± 0.011 ppbv yr⁻¹ (~2.61% yr⁻¹), over China by 0.076 ± 0.020 ppbv yr⁻¹ (~2.27% yr⁻¹),
 235 and the EU by 0.053 ± 0.021 ppbv yr⁻¹ (~1.83% yr⁻¹); error bars represent $\pm 1 \sigma$ standard
 236 deviations and the percent increase is based on the mean concentration. The increasing trends in

237 Fig. 3 are significant at the 95% confidence level for NH_3 over the American Midwest, the EU,
 238 and China with p-values at 0.0003, 0.026, and 0.0028, respectively. South Asia shows only a



239
 240 **Figure 2. Left: Regions with intense agricultural activities.** The averaged 14-year NH_3 VMRs are shown for the
 241 American Midwest (top left), China (2nd left), EU (3rd left), and South Asia (bottom left). The blue boxes outline
 242 areas used in the trend studies. **Right: The seasonal variability in Midwest U.S. (top right), China (2nd right),**
 243 **EU (3rd right), and South Asia (bottom right).** The 7-day means of NH_3 VMRs at 918 hPa are averaged in 2002 –
 244 2008 (green color), 2009 – 2013 (blue color), and 2014 – 2016 (red color) temporal bands, where broad solid lines
 245 represent the averages; and shaded areas are within 1-sigma standard deviations.

246 slight increase (0.0098 ± 0.019 ppbv yr^{-1}) that is not statistically significant (p -value = 0.61). In
 247 Fig. 3, the monthly mean AIRS NH_3 VMRs at 918 hPa are plotted from September 2002 to
 248 August 2016 with proper quality assurance and screening by frequent occurrences. The trends,
 249 however, were linearly fitted using the yearly averaged values.



250 **Figure 3. The recent (~14 year) trends of AIRS NH_3 concentrations.** The NH_3 concentrations (i.e., VMRs in
 251 ppbv) at 918 hPa are shown in black dashed curves, with linear fits in solid lines, for Midwest US (top panel), China
 252 (middle panel), and EU (bottom panel) averaged in each region, respectively. The NH_3 increasing trends are
 253 correlated with OMI SO_2 decreasing trends (red color curve, shown only for China and EU) and NO_2 trends (cyan

254 color). Also shown are the surface skin temperature anomaly (K, in blue color bars) and total precipitation anomaly
255 (cm, in green color bars) from the ECMWF era-interim reanalysis.

256 Over the US, the Midwest shows significant increases in NH_3 , but fertilizer use has
257 grown only modestly ($\sim 1.3\% \text{ yr}^{-1}$) and total food consumption has remained constant within
258 observational uncertainty. Additionally, wet deposition of NH_4^+ does not show any discernable
259 trend [Lajatha and Jones, 2013].

260 What, then is the cause of the definitive growth in the concentration of NH_3 ? Increases
261 can be attributed to a larger fraction remaining in the gas phase due to decreased removal to the
262 condensed phase. For the period of 2003 to 2014 (latest year for which data are available) the
263 USEPA reports an average NO_2 emissions decreasing at an average of $5.3\% \text{ yr}^{-1}$ and SO_2
264 decreasing at $9.0\% \text{ yr}^{-1}$ (<https://www.epa.gov/air-emissions-inventories/air-pollutant-emissions-trends-data>). Measurements from the Interagency Monitoring of Protected Visual Environments
265 (IMPROVE) network indicate that US annual mean ambient SO_2 and SO_4^{2-} concentrations have
266 demonstrated consistent decreases [Hand *et al.*, 2012]. Satellite measurements from the Ozone
267 Monitoring Instrument (OMI) [Krotkov *et al.*, 2016] of tropospheric NO_2 indicated a steady
268 decrease ($-1.54\% \text{ yr}^{-1}$; $p\text{-value}=0.003$) from 2004 to 2016 (Fig. 3, top panel). Emissions of SO_2
269 and NO_x from the electric power sector in 2012 declined to their lowest level since the passage
270 of the Clean Air Act Amendments of 1990. The sum of emission level of SO_2 and NO_x in 2012 is
271 approximately half of those in 2005 when Clean Air Interstate Rule, a cap-and-trade program
272 intended to reduce SO_2 and NO_x beyond the levels defined by the acid rain program in the
273 eastern half of the US, was announced. The decline in emissions is due primarily to an increasing
274 number of [coal-fired units retrofitted](#) with scrubbers, to coal plants switching to lower sulfur coal
275 and low NO_x burners to limit NO_x emissions. Much of the increase in NH_3 over the US,
276 especially after 2011, is thus an unintended consequence of successful measures to control acid
277 deposition.
278

279 Year-to-year NH_3 variations over the US are also affected by meteorological conditions.
280 The highest NH_3 concentrations occurred in 2012 when the surface skin temperature anomaly
281 was up to 4 K in the spring and summer months and the total precipitation anomaly was the
282 lowest in the 14-year history (Fig. 3; negative 3 cm; shaded areas). Increased skin temperatures
283 facilitate higher NH_3 emission rates, and decreased precipitation reduces scavenging of NH_3 gas,
284 although a minimal level of soil moisture is necessary for NH_3 release. Hot, dry summers were
285 conducive to high NH_3 concentrations in 2012 and 2006. From September 2002 to August 2016,
286 the average surface skin temperatures in the US Midwest increased by an average of 0.056 K yr^{-1}
287 using March – August data. Using the published ammonia temperature dependence (see section
288 2.2), we can approximate the contribution from increasing temperatures during this period as
289 $0.65\% \text{ yr}^{-1}$ or approximately 25% of the total ($2.61\% \text{ yr}^{-1}$) increase observed. Note that these
290 temperature trends are highly uncertain considering the fitting is not significant ($p\text{-value} = 0.19$).
291 The effects of variability in climate, which strongly influence NH_3 emission, also influence
292 deposition of NH_3 and the partitioning between gas and aerosol phase [Fowler *et al.*, 2015].
293 Higher air temperatures will reduce the stability of ammonium nitrate aerosols, leading to higher
294 VMRs of NH_3 .

295 The seasonal cycle for the US (Fig. 2 top right panel) shows a broad peak from April to
296 September corresponding to growing season and warmest temperatures; most of the increase is
297 seen in these months. This figure also shows a recent broadening of the maximum, possibly
298 related to warmer temperatures in spring and fall.

299 Over the EU region, increased NH₃ concentrations appear to be due almost entirely to the
 300 decline of SO₂ (-0.0021 ± 0.0007 DU yr⁻¹ and -14.12% yr⁻¹) and NO₂ ($-0.12 \pm 0.0196 \times 10^{15}$
 301 molecules/cm² yr⁻¹ and -2.44% yr⁻¹), according to OMI observations (Fig. 3 bottom panel);
 302 aerosol loading over Europe has likewise declined between 1998 and 2010 [Hsu *et al.*, 2012]. A
 303 number of regulations on air quality protection existed in EU since 1980, yet a Directive of the
 304 European Parliament and of the Council on ambient air quality and cleaner air for Europe was
 305 established in 2008 as a basic legal instrument regulating air quality management [Kuklinska *et al.*, 2015]. This new directive obligates the EU members to implement plans to meet the
 306 permissible levels of certain substances. Possibly due to this action, as well as the economic
 307 crisis, the OMI SO₂ showed a sudden reduction and AIRS NH₃ showed a sudden increase in
 308 2009 and remained at nearly the same level. The relatively small NH₃ increase compared to other
 309 major agricultural regions may be partially due to decreases in nitrogenous fertilizer use, e.g., –
 310 5.2 kg ha⁻¹ yr⁻¹ (-0.3% yr⁻¹) from the World Bank 2016 data base and USDA, World Fertilizer
 311 Consumption Statistics, and International Fertilizer Industry Association, Paris, 2015
 312 (<http://www.fertilizer.org/Statistics>). Concentrated animal feeding operations are another major
 313 contributor to NH₃ emissions and have increased in the EU countries (e.g., Food And Agriculture
 314 Organization Of The United Nations, <http://faostat.fao.org>) and in the US
 315 (<http://www.factoryfarmmap.org/-animal:cattle;location:US;year:2012>). The surface skin
 316 temperature influence on the NH₃ trends in the EU regions was not studied, due to low
 317 temperature trends and lack of significance (e.g., -0.0125 K yr⁻¹ with p-value = 0.70). The
 318 seasonal cycle over the EU (Fig. 2 3rd right panel) is modest, reflecting the weak seasonality in
 319 temperature and precipitation.
 320

321 Over China, the increasing trend of NH₃ (0.076 ppbv yr⁻¹) appears to be related to
 322 decreased sulfur emissions, increased fertilizer use, and increasing local temperatures. The OMI
 323 SO₂ indicate an irregular but discernable decreasing trend (Fig. 3) while fertilizer application has
 324 increased at a rate of 3.5% yr⁻¹ over roughly the time period of our AIRS measurements,
 325 according to the World Bank (<http://data.worldbank.org/indicator/AG.CON.FERT.ZS>).
 326 Decreases in OMI SO₂ (-0.028 ± 0.0052 DU yr⁻¹, or -8.48% yr⁻¹) (Fig. 3 middle panel) generally
 327 track NH₃ increases. An exception to the anti-correlation between SO₂ and NH₃ is the reduction
 328 of both species in 2008 driven by aggressive pollution reduction measures associated with the
 329 Beijing Olympic Games [Wang *et al.*, 2009]. Chinese pollution control legislations are in 5-year
 330 increments, with the 11th five-year plan (FYP) in 2005 – 2010 aiming to reduce SO₂, while the
 331 12th between 2011 and 2015 aiming to reduce SO₂ and NO_x
 332 ([http://wenku.baidu.com/link?url=LdcQKxIkl-HYhK7uONVne4e5-
 333 ikl5Ukvg3iiMVMX37E4LLbIYYfR0s0kdRUbwxydVmZYUcVCFbKyytqxxJPG4kbMQqiUyV
 334 ahVdc95ZKTiG](http://wenku.baidu.com/link?url=LdcQKxIkl-HYhK7uONVne4e5-ikl5Ukvg3iiMVMX37E4LLbIYYfR0s0kdRUbwxydVmZYUcVCFbKyytqxxJPG4kbMQqiUyVahVdc95ZKTiG)). The significant SO₂ emission drop in Northeastern China in 2008, and later,
 335 are because the 11th FYP mandates the installation of emission control devices for power and
 336 steel plants. The NH₃ concentrations increased in 2009 and stayed consistently higher than
 337 before 2009 period (see Fig. 2 2nd right panel), correlating well with the SO₂ reduction. The OMI
 338 NO₂ shows increases over China through 2014, then a fast decreasing in 2015 and 2016 at the
 339 end of the 12th FYP. The NH₃ concentration has reached the highest values in China in 2014,
 340 2015, and 2016 in response to the lowest SO₂ and NO₂ values in our data records. Studies of the
 341 response of ammonia emissions to temperature have been conducted primarily in North America
 342 and Europe, making extrapolation to Asian soils more uncertain, but if we apply these factors to
 343 China, increases in surface skin temperature of 0.0969 K yr⁻¹ (p-value = 0.0057) in spring and
 344 summer explain an increase of approximately 1.12% yr⁻¹ NH₃.

345 The seasonal cycle over China (Fig. 2 2nd top right panel) shows a sharp peak in June and
346 July corresponding to the warmest temperatures and local precipitation maximum. The
347 monsoons generate a strong seasonal cycle in precipitation not seen over the US or EU. A
348 secondary maximum in spring corresponds to peak fertilizer application. The growing season has
349 broadened over China during the period of study, as was observed for the US.

350 South Asia (Figs. 2 & 4) shows the highest concentrations and strongest seasonal cycle,
351 but no significant trend over the past 14 years. Heavy fertilizer use and the highest reported
352 number of cattle of any country lead to strong emissions in the warmer months. South Asia has a
353 distinctive monsoon with ~80% of the precipitation falling in between June and September in
354 Delhi. While winters are warm relative to the US, EU, and China, lack of soil moisture inhibits
355 NH₃ production and release. Fertilizer use in South Asia has increased by 6.9 kg ha⁻¹ yr⁻¹ (+4.8%
356 yr⁻¹) from 2002 to 2013, according to the World Bank. The surface skin temperature influence on
357 the NH₃ trends in South Asia is not studied, due to missing data arising from the uncertainties in
358 summer monsoons, and the small NH₃ trends. Increased emissions are not reflected in the AIRS
359 observations because recent increases in SO₂, 3.25% yr⁻¹ (p-value = 0.047), and NO_x, 1.22% yr⁻¹
360 (p-value = 0.0002), from uncontrolled coal combustion and other sources have led to greater
361 conversion of gaseous NH₃ into particulate sulfates and nitrates. Monitoring with sun
362 photometers indicates a substantial increase in aerosol concentrations over recent years [*Hsu et*
363 *al.*, 2012].

364 **5 Conclusions**

365 The 14-year AIRS satellite record indicates substantial, statistically significant increases
366 in ammonia over several of the world's major agricultural regions, with deleterious effects on
367 vegetation and ecosystem health. Over the US, increases in NH₃ appear to be due to control of
368 SO₂ and NO_x (an unintended consequence of successful acid rain regulations), and to regionally
369 warming temperatures. Over the EU, NH₃ concentrations have increased despite reduced
370 fertilizer use, again due to improved control of sulfur and nitrogen oxide emissions. Over China,
371 a combination of expanded agricultural activities, nascent SO₂ control measures, and increasing
372 temperatures cause the observed increases in ammonia. Over South Asia, increased NH₃
373 emissions from growing fertilizer use are likely masked by simultaneous increases in SO₂ and
374 NO_x emissions, resulting in increased concentrations of fine aerosols with adverse health effects.

375 The observed trends deduced here can guide numerical simulation of tropospheric
376 ammonia and inform policy to mitigate disruption of biogeochemical nitrogen cycles and
377 improve air quality. Complete validation of this satellite ammonia product is needed using long-
378 term ground, as well as new airborne measurements as they become available. Ammonia trend
379 monitoring efforts will continue through the lifetime of AIRS sensor, and with current and future
380 operational sensors such as IASI and CrIS (Cross-track Infrared Sounder) preferably using
381 consistent algorithms.

382 **Acknowledgements**

383 This study was funded by NASA's The Science of Terra and Aqua program under grant numbers
384 NNX11AG39G and NNX12AJ05G. We wish to acknowledge the AIRS, OMI, GEOS-Chem,
385 and ECMWF science teams. RRD was a member of ACAST. MERRA data used in this
386 study/project have been provided by GMAO at NASA Goddard Space Flight Center through the

387 NASA GES DISC online archive. Computations were performed on the NASA Center for
388 Climate Simulation (NCCS) super computing system.

389 The observational data that support the findings of this study are available on the website
390 via the corresponding author (J.X.W.) (http://atmos.umd.edu/~juying/GRL_2017_AIRS_NH3).

391 References

- 392 Abbatt, J. P. D., S. Benz, D.J. Cziczo, Z. Kanji, U. Lohmann, O. Mohler (2006), Solid
393 Ammonium Sulphate Aerosols as Ice Nuclei: A Pathway for Cirrus Cloud Formation,
394 *Science*, 313, 1770 doi:10.1126/science1129726.
- 395 Adams, P. J., J. H. Seinfeld, D. Koch, L. Mickley, and D. Jacob (2001), General circulation
396 model assessment of direct radiative forcing by the sulfate–nitrate–ammonium–water
397 inorganic aerosol system, *J. Geophys. Res.-Atmos.*, 106, 1097–1111,
398 doi:10.1029/2000JD900512.
- 399 Beer, R., M. W. Shephard, S. S. Kulawik, S. A. Clough, A. Eldering, K. W. Bowman, S. P.
400 Sander, B. M. Fisher, V. H. Payne, M. Luo, G. B. Osterman, and J. R. Worden (2008), First
401 satellite observations of lower tropospheric ammonia and methanol, *Geophys. Res. Lett.*, 35,
402 L09801, doi:10.1029/2008GL033642.
- 403 Berrisford, P., D. Dee, P. Poli, R. Brugge, K. Fielding, M. Fuentes, P. Kallberg, S. Kobayashi, S.
404 Uppala, and A. Simmons (2011), The ERA-Interim Archive Version 2.0, *ERA Report Series*
405 *1, ECMWF*, Shinfield Park. Reading, UK 13177.
- 406 Clarisse, L., C. Clerbaux, F. Dentener, D. Hurtmans, and P.-F. Coheur (2009), Global ammonia
407 distribution derived from infrared satellite observations, *Nature Geosci.*, 2(7), 479–483,
408 doi:10.1038/ngeo551.
- 409 Dee, D. P., S. M. Uppala, A. J. Simmons, P. Berrisford, P. Poli, S. Kobayashi, U. Andreae, M. A.
410 Balmaseda, G. Balsamo, P. Bauer, P. Bechtold, A. C. M. Beljaars, L. van de Berg, J. Bidlot,
411 N. Bormann, C. Delsol, R. Dragani, M. Fuentes, A. J. Geer, L. Haimberger, S. B. Healy, H.
412 Hersbach, E. V. Hólm, L. Isaksen, P. Kållberg, M. Köhler, M. Matricardi, A. P. McNally, B.
413 M. Monge-Sanz, J.-J. Morcrette, B.-K. Park, C. Peubey, P. de Rosnay, C. Tavolato, J.-N.
414 Thépaut, F. Vitart (2011), The ERA-Interim reanalysis: Configuration and performance of
415 the data assimilation system. *Quart. J. R. Meteorol. Soc.*, 137, 553-597. DOI: 10.1002/qj.828.
- 416 Deeter, M. N., D. P. Edwards, J. C. Gille, and J. R. Drummond (2007), Sensitivity of MOPITT
417 observations to carbon monoxide in the lower troposphere, *J. Geophys. Res.*, 112, D24306,
418 doi:10.1029/2007JD008929.
- 419 Dentener, F. J., and P. J. Crutzen (1994), A three-dimensional model of the global ammonia
420 cycle, *J. Atmos. Chem.*, 19, 331-369.
- 421 Erisman, J. W., M. A. Sutton, J. Galloway, Z. Klimont, and W. Winiwarter (2008), How a
422 century of ammonia synthesis changed the world, *Nature Geosci.*, 1(10), 636–639,
423 doi:10.1038/ngeo325.
- 424 Erisman, J. W., J. N. Galloway, S. Seitzinger, A. Bleeker, N. B. Dise, R. Petrescu, A. M. Leach,
425 and W. de Vries (2013), Consequences of human modification of the global nitrogen cycle,
426 *Philos. T. R. Soc. B*, 368, 1621, doi:10.1098/rstb.2013.0116.

- 427 Fowler D, M. Coyle, U. Skiba, M. A. Sutton, J. N. Cape, S. Reis, L. J. Sheppard, A. Jenkins, B.
428 Grizzetti, J. N. Galloway, P. Vitousek, A. Leach, A. F. Bouwman, K. Butterbach-Bahl, F.
429 Dentener, D. Stevenson, M. Amann, and M. Voss (2013), The global nitrogen cycle in the
430 twenty-first century. *Phil Trans R Soc B* 368, 20130164, doi:10.1098/rstb.2013.0164
- 431 Fowler, D., C. E. Steadman, D. Stevenson, M. Coyle, R. M. Rees, U. M. Skiba, M. A. Sutton, J.
432 N. Cape, A. J. Dore, M. Vieno, D. Simpson, S. Zaehle, B. D. Stocker, M. Rinaldi, M. C.
433 Facchini, C. R. Flechard, E. Nemitz, M. Twigg, J. W. Erisman, K. Butterbach-Bahl, and J. N.
434 Galloway (2015), Effects of global change during the 21st century on the nitrogen cycle,
435 *Atmos. Chem. Phys.*, 15, 13849–13893, doi:10.5194/acp-15-13849-2015
- 436 Friedl, M. A., D. Sulla-Menashe, B. Tan, A. Schneider, N. Ramankutty, A. Sibley, and X. M.
437 Huang (2010), MODIS Collection 5 global land cover: Algorithm refinements and
438 characterization of new datasets, *Remote Sens. Environ.*, 114(1), 168–182,
439 doi:10.1016/j.rse.2009.08.016.
- 440 Galloway, J. N., and E. B. Cowling (2002), Reactive Nitrogen and The World: 200 Years of
441 Change, *Ambio.*, 31(2), 64–71, 5 doi:10.1579/0044-7447-31.2.64.
- 442 Galloway, J. N., F. J. Dentener, D. G. Capone, E. W. Boyer, R. W. Howarth, S. P. Seitzinger,
443 G. P. Asner, C. C. Cleveland, P. A. Green, E. A. Holland, D. M. Karl, A. F. Michaels,
444 J. H. Porter, A. R. Townsend, C. J. Vöosmarty (2004), Nitrogen cycles: past, present, and
445 future, *Biogeochemistry*, 70(2), 153-226.
- 446 Galloway, J. N., A. R. Townsend, J. W. Erisman, M. Bekunda, Z. Cai, J. R. Freney, L. A.
447 Martinelli, S. P. Seitzinger, and M. A. Sutton (2008), Transformation of the nitrogen cycle:
448 Recent trends, questions, and potential solutions, *Science*, 320, 889–892.
- 449 Garcia, D. (2010), Robust smoothing of gridded data in one and higher dimensions with
450 missing values. *Comput Stat Data Anal* 54, 1167–1178, doi:10.1016/j.csda.2009.09.020.
- 451 Giglio, L., J.T., Randerson, G.R., Van der Werf, P.S., Kasibhatla, G.J., Collatz, D.C., Morton,
452 R.S., DeFries, (2010). Assessing variability and long-term trends in burned area by merging
453 multiple satellite fire products. *Biogeosciences* 7, 1171e1186. [http://dx.doi.org/10.5194/bg-7-](http://dx.doi.org/10.5194/bg-7-1171-2010)
454 1171-2010.
- 455 Hand, J. L., B. A. Schichtel, W. C. Malm, and M. L. Pitchford (2012), Particulate sulfate ion
456 concentration and SO₂ emission trends in the United States from the early 1990s through
457 2010, *Atmos. Chem. Phys.*, 12, 10353-10365, doi:[10.5194/acp-12-10353-2012](https://doi.org/10.5194/acp-12-10353-2012).
- 458 Hauglustaine, D. A., Y. Balkanski, and M. Schulz (2014), A global model simulation of present
459 and future nitrate aerosols and their direct radiative forcing of climate, *Atmos. Chem. Phys.*,
460 14, 11031–11063, doi:10.5194/acp-14-11031-2014.
- 461 Henze, D. K., D. T. Shindell, F. Akhtar, R. J. D. Spurr, R. W. Pinder, D. Loughlin, M. Kopacz,
462 K. Sing, and C. Shim (2012), Spatially refined aerosol direct radiative forcing efficiencies,
463 *Environ. Sci. Technol.*, 46, 9511–9518, doi:10.1021/es301993s.
- 464 Hsu, N. C., R. Gautam, A. M. Sayer, C. Bettenhausen, C. Li, M. J. Jeong, S.-C. Tsay, and B. N.
465 Holben (2012), Global and regional trends of aerosol optical depth over land and ocean using
466 SeaWiFS measurements from 1997 to 2010, *Atmos. Chem. Phys.*, 12(17), 8037-8053.

- 467 Huang, X., Y. Song, M. Li, J. Li, Q. Huo, X. Cai, T. Zhu, M. Hu, and H. A. Zhang (2012), High-
468 resolution ammonia emission inventory in China, *Global Biogeochem. Cy.*, *26*, GB1030,
469 doi:10.1029/2011GB004161.
- 470 Krotkov, N. A., C. A. McLinden, C. Li, L. N. Lamsal, E. A. Celarier, S. V. Marchenko, W. H.
471 Swartz, E. J. Bucsela, J. Joiner, B. N. Duncan, K. F. Boersma, J. P. Veefkind, P. F. Levelt, V.
472 E. Fioletov, R. R. Dickerson, H. He, Z. Lu, and D. G. Streets (2016), Aura OMI observations
473 of regional SO₂ and NO₂ pollution changes from 2005 to 2015, *Atmos. Chem. Phys.*,
474 *16*, 4605-4629, doi:[10.5194/acp-16-4605-2016](https://doi.org/10.5194/acp-16-4605-2016).
- 475 Lajtha, K., and J. Jones (2013), Trends in cation, nitrogen, sulfate and hydrogen ion
476 concentrations in precipitation in the United States and Europe from 1978 to 2010: a new
477 look at an old problem, *Biogeochemistry*, *116*(1-3), 303-334.
- 478 Lelieveld, J., J. S. Evans, M. Fnais, D. Giannadaki, and A. Pozzer (2015), The contribution of
479 outdoor air pollution sources to premature mortality on a global scale. *Nature*, *525*.7569,
480 367-371.
- 481 Malm, W. C., B. A. Schichtel, M. L. Pitchford, L. L. Ashbaugh, and R. A. Eldred (2004), Spatial
482 and monthly trends in speciated fine particle concentration in the United States, *J. Geophys.*
483 *Res.* *109*, D03306, doi:10.1029/2003JD003739.
- 484 Martin, S. T., Hung, H.-M., Park, R. J., Jacob, D. J., Spurr, R. J. D., Chance, K. V., and Chin, M.
485 (2004), Effects of the physical state of tropospheric ammonium-sulfate-nitrate particles on
486 global aerosol direct radiative forcing, *Atmos. Chem. Phys.*, *4*, 183–214, doi:10.5194/acp-4-
487 183-2004.
- 488 Pope, C. A., Burnett, R. T., Thun, M. J., Calle, E. E., Krewski, D., Ito, K., and Thurston, G. D.
489 (2002), Lung cancer, cardiopulmonary, mortality, and long-term exposure to fine particulate
490 air pollution, *J. Am. Med. Assoc.*, *287*, 1132–1141.
- 491 Riddick, S., Ward, D., Hess, P., Mahowald, N., Massad, R., and Holland, E. (2016), Estimate of
492 changes in agricultural terrestrial nitrogen pathways and ammonia emissions from 1850 to
493 present in the Community Earth System Model, *Biogeosciences*, *13*(11), 3397-3426.
- 494 Roelle, P. A. and Aneja, V. P. (2002), Environmental Simulation Chambers: Application to
495 Atmospheric Chemical Processes, *Science*, *457 pp.* (Springer, 13 January 2006).
- 496 Schiferl, L. D., Heald, C. L., Van Damme, M., Clarisse, L., Clerbaux, C., Coheur, P.-F., Nowak,
497 J. B., Neuman, J. A., Herndon, S. C., Roscioli, J. R., and Eilerman, S. J. (2016): Interannual
498 variability of ammonia concentrations over the United States: sources and implications,
499 *Atmos. Chem. Phys.*, *16*, 12305-12328, doi:[10.5194/acp-16-12305-2016](https://doi.org/10.5194/acp-16-12305-2016).
- 500 Sheppard L. J., Leith, I. D., Mizunuma, T., Cape, J. N., Crossley, A., Leeson, S., Sutton, M. A.,
501 Dijk, N. V., and Fowler, D. (2011), Dry deposition of ammonia gas drives species change
502 faster than wet deposition of ammonium ions: evidence from a long-term field manipulation,
503 *Global Change Biology*, *17*, 3589–3607, doi:10.1111/j.1365-2486.2011.02478.x.
- 504 Streets, D. G., Bond, T. C., Carmichael, G. R., Fernandes, S. D., Fu, Q., He, D., Klimont, Z.,
505 Nelson, S. M., Tsai, N. Y., Wang, M. Q., Woo, J. H., and Yarber, K. F. (2003), An inventory
506 of gaseous and primary aerosol emissions in Asia in the year 2000, *J. Geophys. Res.-Atmos.*,
507 *108*(D21), 8809, doi:10.1029/2002JD003093.

- 508 Sutton, M. A., Nemitz, E., Erisman, J. W., Beier, C., Bahl, K. B., Cellier, P., de Vries, W.,
509 Cotrufo, F., Skiba, U., Di Marco, C., Jones, S., Laville, P., Soussana, J. F., Loubet, B.,
510 Twigg, M., Famulari, D., Whitehead, J., Gallagher, M. W., Neftel, A., Flechard, C. R.,
511 Herrmann, B., Calanca, P. L., Schjoerring, J. K., Daemmgen, U., Horvath, L., Tang, Y. S.,
512 Emmett, B. A., Tietema, A., Penuelas, J., Kesik, M., Brueggemann, N., Pilegaard, K.,
513 Vesala, T., Campbell, C. L., Olesen, J. E., Dragosits, U., Theobald, M. R., Levy, P., Mobbs,
514 D. C., Milne, R., Viovy, N., Vuichard, N., Smith, J. U., Smith, P., Bergamaschi, P., Fowler,
515 D., and Reis, S. (2007), Challenges in quantifying biosphere-atmosphere exchange of
516 nitrogen species, *Environ. Pollut.*, *150*, 125–139, doi:10.1016/j.envpol.2007.04.014.
- 517 Sutton, M., Erisman, J., Dentener, F., and Moller, D. (2008), Ammonia in the environment: From
518 ancient times to the present, *Environ. Pollut.*, *156*, 583–604,
519 doi:10.1016/j.envpol.2008.03.013.
- 520 Van Damme, M., Erisman, J. W., Clarisse, L., Dammers, E., Whitburn, S., Clerbaux, C.,
521 Dolman, A. J., and Coheur, P.-F. (2015), Worldwide spatiotemporal atmospheric ammonia
522 (NH₃) columns variability revealed by satellite, *Geophys. Res. Lett.*, *42*(20), 8660–8668,
523 doi:10.1002/2015GL065496.
- 524 Wang, Y., Hao, J., McElroy, M. B., Munger, J. W., Ma, H., Chen, D., & Nielsen, C. P. (2009),
525 Ozone air quality during the 2008 Beijing Olympics: effectiveness of emission restrictions,
526 *Atmos. Chem. Phys.*, *9*(14), 5237–5251.
- 527 Wang, Y., Zhang, Q. Q., He, K., Zhang, Q., and Chai, L. (2013), Sulfate-nitrate-ammonium
528 aerosols over China: response to 2000–2015 emission changes of sulfur dioxide, nitrogen
529 oxides, and ammonia, *Atmos. Chem. Phys.*, *13*, 2635–2652, doi:10.5194/acp-13-2635-2013.
- 530 Warner, J. X., Wei, Z., Strow, L. L., Dickerson, R. R., Nowak, R. (2016), Global Ammonia
531 Sources Seen by AIRS 13-years Measurements, *Atmos. Chem. Phys.* *16*, 5467-
532 5479, doi:10.5194/acp-16-5467-2016.