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The genesis of phosphatic nodules in the Toyoma Formation, Northeastern Japan

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Various phosphatic and carbonate rocks occur in the Permian Toyoma Formation. Morikiyo & Matsunaga [1] reported the isotopic ratios of calcite, pyrite and apatite in these rocks and concluded that the rocks were formed by sulfate reduction. However, the δ^{34} S of pyrite was reported for the total rock; therefore, this reasoning requires reconsideration. The δ^{34} S of individual pyrite grains in the phosphatic rocks was re-analyzed using high-resolution secondary ion mass spectrometry. Intergranular differences in the δ^{34} S of rock samples are as large as 38‰. δ^{34} S is correlated to the morphology of pyrite; framboidal pyrite has low δ^{34} S, whereas euhedral pyrite has high δ^{34} S. It is thus possible to determine the stage of diagenesis in which each rock type was formed by considering the total rock pyrite δ ³⁴S and the modal proportion of the pyrite types. The phosphatic nodules were formed at the earliest stage in all of the rocks in the Formation. The source of phosphorus is inferred to be fish scales and bones, according to the Y, La, and Ce contents. Therefore, points to be clarified are: 1) the reason for dissolution of biogenic phosphate debris, and 2) the cause of apatite precipitation after the dissolution of the debris. At the time of deposition, the occurrence of trace fossils shows that surface sediments were not anoxic, rather that sediments inside were anoxic. Thus, the upward diffusion of H₂S from the underlying sediment resulted in surface sediment containing sulfur-oxidizing bacteria, which grew and produced sulfuric acid and caused the dissolution of phosphate debris. As burial of the sediment progressed, the former surface layer became anoxic, initiating sulfate reduction, which led to the precipitation of apatite.

[1] Morikiyo & Matsunaga (2001) Abstr. 11th Goldschmidt Conf. #3066.