

# 1 A preliminary assessment of climate change impacts on sugarcane in Swaziland

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## 8 **Abstract**

9 The spatial and temporal impacts of climate change on irrigation water requirements and yield  
10 for sugarcane grown in Swaziland have been assessed, by combining the outputs from a  
11 general circulation model (HadCM3), a sugarcane crop growth model and a GIS. The  
12 CANEGRO model (embedded with the DSSAT program) was used to simulate the baseline  
13 and future cane net annual irrigation water requirements ( $IR_{net}$ ) and yield ( $t\ ha^{-1}$ ) using a  
14 reference site and selected emissions scenario (SRES A2 and B2) for the 2050s (including  
15  $CO_2$ -fertilisation effects). The simulated baseline yields were validated against field data from  
16 1980-1997. An aridity index was defined and used to correlate agroclimate variability against  
17 irrigation need to estimate the baseline and future irrigation water demand (volumetric). To  
18 produce a unit weight of sucrose equivalent to current optimum levels of production, future  
19 irrigation needs were predicted to increase by 20-22%. With  $CO_2$ -fertilisation, the impacts of  
20 climate change are offset by higher crop yields, such that  $IR_{net}$  is predicted to increase by 9%.  
21 The study showed that with climate change, the current peak capacity of existing irrigation  
22 schemes could fail to meet the predicted increases in irrigation demand in nearly 50% of years  
23 assuming unconstrained water availability.

24 **Keywords:** *CANEGRO; GIS; irrigation; sugarcane; water; yield.*

## 25 **1. Introduction**

26 Many studies in the research literature describe how agricultural production in Africa will be  
27 one of the sectors most vulnerable to climate change and variability (Challinor *et al.*, 2005).

28 This is because a significant proportion of the African economy is dependent on agriculture,  
29 most of Africa's water (85%) is used for agriculture (Downing *et al.*, 1997), farming  
30 techniques are relatively primitive and the majority of the continent is already hot and dry  
31 (Kurukulasuriya and Mendelsohn, 2008). Spatial and temporal changes in precipitation and  
32 temperature patterns will thus have major impacts on the viability of both dryland and  
33 irrigated farming (Benhin, 2008). For an important commodity crop such as sugarcane where  
34 water is a limiting factor in production, the priorities are to assess the impacts of climate  
35 change on both resource availability (for irrigation abstraction) and water demand (for crop  
36 production). However, most studies to date have focussed on agriculture and rural livelihoods,  
37 with limited attention to impacts on sugarcane in southern Africa (Deressa *et al.*, 2005).

38 In many African countries, including Nigeria (Binbol *et al.*, 2006), South Africa (Hassan and  
39 Olbrich, 2000), Zambia and Zimbabwe, sugarcane forms the mainstay of the economy. In  
40 Swaziland, production dates back to the mid-1950s, with the establishment of mills at Big  
41 Bend, Mhlume and Simunye. Sugarcane production has grown steadily, and in 2007  
42 accounted for 59% of Swaziland's agricultural output and 24% of gross domestic product  
43 (GDP). In 2008 production was reported to be 5,100,456 tonnes with an average annual yield  
44 of 102 t ha<sup>-1</sup> (SSA, 2009). Most cultivation is concentrated on large plantations in the  
45 Lowveld and Lower Middleveld regions, where fertile soils and high temperatures provide  
46 ideal conditions for production, although all are dependant on irrigation to supplement low  
47 rainfall during the growing season. In this context, Swaziland is unique, as sugarcane cannot  
48 be grown without irrigation, in contrast to neighbouring countries such as South Africa where  
49 40% of the total cropped area is irrigated (Inman-Bamber and Smith, 2005). As a  
50 consequence, the majority of water abstracted for agriculture (96%) in Swaziland is used for  
51 sugarcane production (Matondo *et al.*, 2005).

52 The total cane cropped area is currently 52,071 ha (SSA, 2009) having increased from 14,500  
53 ha in the late 1960s (Murdoch, 1968); further irrigation developments are underway which

54 will result in an additional 19,000 ha being cultivated. This will add pressure on already  
55 strained water resources, and is likely to lead to increased tensions with neighbouring riparian  
56 states regarding water allocations for agriculture (Nkomo and van der Zaag, 2004). At present,  
57 sugarcane irrigation needs vary between 10,000 to 14,000 m<sup>3</sup> ha<sup>-1</sup> depending on variety, soil  
58 and agroclimate conditions. Although traditional methods such as furrow are still popular (39%  
59 of the total area), sprinklers (54%), centre pivots (3%) and drip (3%) are gaining favour  
60 (Nkomo and van der Zaag, 2004) as estates switch technology to improve water efficiency  
61 (more 'crop per drop'), coupled with concerns regarding labour availability (Merry, 2003).  
62 Deressa *et al* (2005) assessed the economic impacts of climate change on sugarcane in South  
63 Africa using a Ricadian approach. By combining critical damage point analyses with  
64 information on agroclimate variability their analyses showed that sugarcane revenue is more  
65 sensitive to predicted increases in temperature, rather than rainfall. Their analysis excluded  
66 the impacts of CO<sub>2</sub> fertilisation on productivity. Previous studies have investigated the  
67 impacts of climate change on water resources in Swaziland but have not considered sugarcane  
68 production (Matondo *et al.*, 2004). Other studies have assessed agronomic impacts and the  
69 potential for using spatial (GIS) modelling for yield prediction (Kiker, 2000). The objective of  
70 this study was to conduct a preliminary assessment of climate change impacts on sugarcane  
71 production in Swaziland.

## 72 **2. Methodology**

73 In summary, the outputs from a general circulation model (GCM), a sugarcane crop growth  
74 model and a geographical information system (GIS) have been combined to assess the spatial  
75 and temporal impacts of climate change on cane yield and irrigation needs. Using selected  
76 IPCC SRES scenarios for the 2050s (Nakicenovic *et al.*, 2000), future climate datasets were  
77 derived for a reference site using outputs from the HadCM3 model. The net annual irrigation  
78 water requirements (IR<sub>net</sub>) and crop productivity (t ha<sup>-1</sup>) for the baseline and selected IPCC

79 scenario were then simulated using the CANEGRO model embedded within the DSSAT  
80 (Decision Support System for Agrotechnology Transfer) program (Jones *et al.*, 2003). The  
81 crop simulations considered future emissions scenarios both with and without CO<sub>2</sub>  
82 fertilisation effects. Using potential soil moisture deficit (PSMD) as an aridity index, maps  
83 showing future changes in agroclimate were produced. Finally, a linear regression analysis  
84 between agroclimate variability and irrigation need was used to estimate current and future  
85 volumetric water demand for sugarcane. A brief description of the study site, the climate  
86 change scenarios and datasets, crop modelling and GIS mapping, is provided below.

### 87 *2.1 Study site*

88 The study site was Mhlume (Lon: 26:03:02S; Lat: 31:50:05E), in the eastern Lowveld, an area  
89 in which nearly half the total area of irrigated sugarcane in Swaziland is located. The  
90 Commonwealth Development Corporation (CDC) established a sugar mill at Mhlume in the  
91 1950s, now owned by the Royal Swaziland Sugar Corporation (RSSC) who manage  
92 approximately 20,000 ha of sugarcane which is milled at Mhlume and Simunye factories.  
93 RSSC is one of the largest companies in Swaziland, producing two-thirds of the country's  
94 sugar. Mhlume has a sub-tropical steppe climate and compared to other parts of the country,  
95 the Lowveld region is characterised by low rainfall and high temperatures. For this study,  
96 daily weather records for 1969-1996 for the site were available. In January, the mean monthly  
97 temperature is 31°C, but with daily maximum temperatures as high as 39°C. The minimum  
98 monthly mean temperature (9°C) occurs in winter (June to July), but on some days can be as  
99 low as 3°C. Nearly 80% of annual rainfall occurs between October and March. Reference  
100 evapotranspiration (ET<sub>o</sub>) (Allen *et al.*, 1998) exceeds rainfall in all months, with the greatest  
101 moisture deficits occurring between May and September (Figure 1). A cropping database for  
102 RSSC provided detailed field records on planting and harvest dates, varieties grown, soil  
103 types, ratoon periods, irrigation methods, and yields (harvested cane and sucrose) from 1980  
104 to 2007. This database was used for validating the CANEGRO simulation outputs.

## 105 2.2 Climate change scenarios and datasets

106 Climate projections were based on the HadCM3, a third generation coupled atmosphere-ocean  
107 general circulation model developed at the Hadley Centre for Climate Prediction and  
108 Research (Johns *et al.*, 1997). It was developed from the earlier HadCM2 model, used to  
109 generate predictions of climate change for the IPCC 3rd and 4th Assessment Reports, and has  
110 been widely used in Africa for impact assessments. For example, Tanser *et al* (2003) studied  
111 the effects of climate change on malaria transmission in Africa using HadCM3 and three  
112 climate scenarios. Thomas *et al* (2005) studied the mobilization of southern African desert  
113 dune systems using outputs from three GCMs (HadCM3, HadCM2 and CGCM1). They  
114 showed that for the HadCM3 model, index values for dune activity bore a very close  
115 relationship to those derived from observed data for the 1961–90 period.

116 The HadCM3 has a higher spatial resolution than previous versions (2.5° x 3.75°, latitude by  
117 longitude) and allows the radiative effects of CO<sub>2</sub> and other minor greenhouse gases,  
118 including water vapour and ozone to be represented. In order to provide information on  
119 possible changes in global climate, the model is forced to consider future scenarios where  
120 changes in atmospheric CO<sub>2</sub> concentration are assumed depending on anthropogenic activity  
121 for three 30-year mean periods (2020s, 2050s, and 2080s). The scenarios reflect different  
122 ‘storylines’ based on differing rates of demographic change, industrial activity, dependence  
123 on fossil fuels, and other socio-economic indicators. These represent mutually consistent  
124 characterisations of future states of the world during the 21st century, and are neither  
125 predictions nor forecasts of future conditions. Rather, they describe alternative plausible  
126 futures that conform to sets of circumstances or constraints within which they arise. The true  
127 purpose of scenarios is thus to determine the possible ramifications of climate change along  
128 one or more plausible (but indeterminate) paths.

129 The emissions are based on those developed by the IPCC (Nakicenovic *et al.*, 2000) and  
130 known as SRES (Special Report on Emission Scenarios). In simple terms, there are four  
131 ‘marker scenarios’ that combine two sets of divergent tendencies. One set varying between  
132 strong economic values and strong environmental values, the other set varying between  
133 increasing globalisation and increasing regionalisation (IPCC-TGCI, 1999). The scenarios  
134 are commonly referred to as A1 (economic-global), B1 (environmental-global), A2  
135 (economic-regional), and B2 (environmental –regional). For this research, the A2 and B2  
136 scenarios for the 2050s were chosen. The A2 scenario has the higher atmospheric CO<sub>2</sub>  
137 concentration and temperature increase with the highest population increase, the B2 is less  
138 extreme, assuming greater efforts to control global CO<sub>2</sub> emissions (de Silva *et al.*, 2007)  
139 (Table 1). Strzepek and McCluskey (2006) assessed the impacts of climate change on regional  
140 water resources and agriculture in Africa using five different models (CSIRO2, HadCM3,  
141 CGCM2, ECHAM and PCM) using the same emission scenarios. An approach involving  
142 downscaling the HadCM3 outputs for each scenario was chosen in preference to using a  
143 regional climate model (RCM) as previously used in South Africa (Hudson and Jones, 2002;  
144 Tadross *et al.*, 2005) since these studies considered only one socioeconomic scenario (SRES  
145 A2) for 2100. The challenges of choosing an appropriate GCM, a representative number of  
146 emissions scenarios and time slices and the downscaling approach in order to capture an  
147 appropriate degree of uncertainty in the modelling are considered under the methodological  
148 limitations section.

149 When downscaling, changes in climate need to be considered relative to a ‘baseline’. In this  
150 study, a baseline climatology developed by the International Water Management Institute  
151 (IWMI) was used (New *et al.*, 2002). This 10’ resolution dataset includes gridded mean  
152 monthly surface climate data, derived from observed data for 1961 to 1990, to match the  
153 World Meteorological Organisation (WMO) standard. However, it is important to check that  
154 the baseline (historical) climate for a study site is consistent with the equivalent gridded

155 baseline climatology data. For Mhlume, the historical baseline referred to 1969 to 1996. A  
156 comparison between Mhlume and the equivalent IWMI grid pixel (1961-90) using mean  
157 monthly data for rainfall and reference evapotranspiration (ET<sub>o</sub>) is shown in Figure 2.  
158 Although the time series are different, linear regression analyses showed a very high  
159 correlation between the two datasets (Rainfall  $R^2 = 0.96$  and ET<sub>o</sub>  $R^2 = 0.98$ ) confirming that  
160 the IWMI baseline climatology was appropriate for the downscaling process.

161 It is acknowledged that GCMs do not simulate the present climate perfectly, and that model  
162 changes predicted from the present to the future are generally more reliable than the present or  
163 the future climate predicted alone (Carter *et al.*, 1994). Downscaling GCM outputs for the  
164 study site was undertaken using a well established procedure using ‘change factors’ (Diaz-  
165 Nieto and Wilby, 2005). A baseline climatology for the site was first established. Changes in  
166 the equivalent climate variables for the GCM grid box closest to the target site (Mhlume)  
167 were calculated by taking the difference between the transient HadCM3 GCM runs with the  
168 IWMI observed climate data from the 30 year baseline period (Table 2). Finally, these  
169 ‘change factors’ (CF) were applied to the historical baseline – adding the changes in  
170 temperature to the observed temperature, and multiplying ratio changes for precipitation and  
171 other variables (e.g. solar radiation, wind, ET<sub>o</sub>) by their observed daily values during the  
172 period 1961-90 (Alexandrov and Hoogenboom, 2000). Two new datasets were generated, to  
173 represent the future climate at Mhlume under each SRES scenario (2050\_A2 and 2050\_B2).  
174 Using this approach, all the daily climate values in each month are altered by the same  
175 percentage, each day and in each year of record. This approach has the virtue of simplicity  
176 and maintains a realistic temporal structure of climate data. It also assumes that the relative  
177 variability in climate from day to day and year to year (the shape of the frequency distribution)  
178 remains constant. Whilst it is recognised that this is not necessarily true of future climate, it  
179 avoids introducing additional uncertainty into the analysis. Similar CF approaches for  
180 downscaling have been applied in the UK (Pilling and Jones, 1999), Bulgaria (Alexandrov

181 and Hoogenboom, 2000), Spain (Rodriguez Diaz *et al.*, 2007) and Sri Lanka (de Silva *et al.*,  
182 2007). The historical baseline and perturbed future climate datasets for Mhlume were used as  
183 inputs for the sugarcane crop modelling.

### 184 *2.3 Modelling sugarcane yield and water use*

185 For simulating baseline and future sugarcane yield and irrigation needs, the CANEGRO  
186 model was used; this is one of 16 crop models embedded within the DSSAT (v4.0) program  
187 (Jones *et al.*, 2003). A brief description of the CANEGRO model is given below, but readers  
188 interested in a detailed description are referred to Inman-Bamber (1991; 1995) and O’Leary  
189 (2000). The CANEGRO model was originally developed by the South African Sugar  
190 Association Experiment Station (SASEX) to determine optimal harvest age because of risks  
191 from the stalk borer *Eldana sacchararina* (Inman-Bamber, 1995). It has since been embedded  
192 into DSSAT and used in Africa (Inman-Bamber and Kiker, 1997), Asia (Jintrawet and  
193 Prammanee, 2005) and America. The model contains carbon simulation, crop development,  
194 energy and water simulation components. Although it was coupled to a soil and plant nitrogen  
195 model from the CERES-Maize model (Jones and Kiniry, 1986) in the DSSAT program this  
196 has not yet been validated, and hence CANEGRO remains a radiation-water-temperature  
197 limited model that takes no account of nutrient status (O’Leary, 2000). The model has,  
198 however, been extensively tested for simulating above ground biomass and water status for  
199 NCo376, a popular cultivar grown in Swaziland. Keating *et al* (1995) has shown CANEGRO  
200 to be robust for simulating biomass with water or nitrogen stress and Inman-Bamber (1994,  
201 1995) reported on its application at two different locations. The model requires input data  
202 relating to the local weather, crop and soil characteristics, and management practices  
203 (fertilizer and irrigation regimes) and runs on a daily time-step to calculate crop phasic and  
204 morphological development using temperature, day length and genetic characteristics. The  
205 weather, crop and soil datasets and assumptions used for parameterising CANEGRO are  
206 outlined below.



207 Three weather datasets were used. A historical baseline dataset containing daily maximum  
208 and minimum temperature, wind speed, solar radiation, rainfall, and relative humidity for  
209 Mhlume for 1969-1996, and two equivalent perturbed datasets for the SRES 2050\_A2 and  
210 2050\_B2 scenario, respectively, as described previously. Crop modelling was based on  
211 NCo376, a cultivar which is grown extensively in Swaziland. An analysis of RSSC field data  
212 for 1980-2007 showed that on average this variety accounts for 66% of the total cropped area.  
213 The study assumed a plant cane crop; however, in reality, sugarcane is ratooned and only a  
214 small proportion (typically 10%) is plant cane. At Mhlume, over three-quarters (77%) of the  
215 annual cropped area is ratooned cane aged 1 and 6 years (Figure 3). This was acknowledged  
216 to be a methodological limitation as plant cane yields are higher than ratooned cane. However,  
217 analysis of RSSC field data actually showed that the average yield for plant cane was not  
218 significantly different from ratooned cane aged 1-6 years (Figure 3). It was therefore assumed  
219 that simulating plant cane yield would provide a reasonable indication of 'typical' yield for  
220 cane under both current and future climates.

221 Planting and emergence dates were assumed to be identical. This is because in ratoon cane the  
222 stems are cut to ground level and the stumps appear above ground, as in emergence. Normal  
223 practice is to stagger planting in order to optimise cane supplies to the factory. For the  
224 modelling exercise, November planting was chosen as this coincides with higher temperatures  
225 and rainfall (Figure 1) which is the ideal condition for germination and filleting (Doorenbos  
226 and Kassam, 1979). The assumed irrigation method was furrow as this represented 52% of the  
227 irrigated area in the region. An automatic irrigation schedule (defining the timing and amount  
228 of irrigation) was chosen, with irrigation scheduled to return the soil back to field capacity  
229 when the profile soil water content dropped below 65% of total available water. This is  
230 assumed typical of current irrigation management practices in the region. Irrigation efficiency  
231 was assumed to be 100%, as net irrigation water requirements were being modelled, although  
232 in practice surface irrigation efficiencies are considerably lower.

233 At RSSC, the soils are grouped into three classes ranging from good (1) to poor (3) in terms  
234 of sugarcane suitability. For modelling, the fields were assumed to have 'R-set' soils. These  
235 are Class 1 soils, equivalent to heavy textured Shortlands and Hutton Forms in the South  
236 African Binomial Soil Classification, with an effective rooting depth of 1 m and an available  
237 water capacity (AWC) of 140-180 mm m<sup>-1</sup> (SASEX, 1999). They are defined as moderate to  
238 well structured red or reddish brown clay loam to clay soils with moderate organic matter  
239 content, and usually occur in mid-slope positions on well draining gentle slopes (Nixon,  
240 2006). They are one of the best soils, giving higher cane yields than other local soils  
241 (Murdoch, 1968). An analysis of RSSC field data showed that 65% of the cropped area were  
242 on Class 1 soils, and 76% of all fields contained R-set soils.

243 The CANEGRO model was parameterised and used to simulate annual sugarcane yield and  
244 irrigation needs for a baseline 'scenario' using data from 1980-96. The model was then re-run  
245 for each SRES scenario (with and without CO<sub>2</sub> fertilisation effects) using the same crop and  
246 soil files, but with the future climate datasets. For each year of simulation, model outputs  
247 included biomass yield (t ha<sup>-1</sup>), sucrose yield (t ha<sup>-1</sup>), irrigation needs (mm), and water use  
248 efficiency (WUE) defined as kilograms of sucrose production per cubic metre of irrigation  
249 water usefully applied (kg<sup>-1</sup> m<sup>-3</sup>).

#### 250 *2.4 Model validation*

251 It is important to have confidence that a crop model can predict with reasonable accuracy  
252 historical variations in yield, before imposing further uncertainty through climate change. The  
253 CANEGRO model was used to simulate yields for 1980-96. For validation purposes, RSSC  
254 field data for the same period were obtained. These contained information on cane yield,  
255 including variety, ratoon year, planting and harvest dates, and soil type (18,000 records in  
256 total) on a field by field basis. From this, a validation dataset was produced (based on 1549  
257 fields) containing yields for all fields growing plant cane (variety NCo376) on R-set soils. A

258 comparison between the CANEGRO modelled and RSSC observed cane yields was  
259 completed (Figure 4). Visually, for most years, the modelled yield compared well to the  
260 average observed yield and within  $\pm 1$  SD (as shown by the error bars). In some years, the  
261 modelled and observed average yields were very similar. To assess whether bias of modelled  
262 yields versus observed yields were statistically significant, the model outputs were analysed  
263 for lack of fit (LOFIT) with the observed data using a method described by Whitmore (1991).  
264 This test was chosen in preference to more widely used goodness-of-fit statistics such as the  
265 correlation coefficient ( $r$ ) and root mean squared error (RMSE) because rather than  
266 comparing a single modelled value against a single observed value, it considers multiple  
267 observed values and differing numbers of observed values in a temporal series. The calculated  
268 F value (1.50) was not significant, confirming there was no evidence to suggest that the  
269 modelled and observed data were statistically different.

### 270 *2.5 Modelling agroclimate and irrigation demand*

271 The variables that directly influence soil moisture and hence irrigation are rainfall and  
272 reference evapotranspiration (ET<sub>o</sub>). To assess the spatial impacts of climate variability on  
273 sugarcane irrigation needs, an approach was needed to extrapolate the CANEGRO modelled  
274 outputs for a single site (Mhlume) across Swaziland. Previous research has shown that a  
275 strong relationship exists between irrigation need and potential soil moisture deficit (PSMD)  
276 for a range of crops and climates, including rice in Sri Lanka (de Silva *et al.*, 2007) and  
277 horticulture in Spain (Rodriguez Diaz *et al.*, 2007). The advantage of this index over others  
278 such as the Wetness Index (ratio of total annual rainfall and total annual evapotranspiration) is  
279 that the distribution of rainfall and ET throughout the year is taken into account. Furthermore,  
280 in many African countries where spatial information is sparse or non-existent, the PSMD  
281 agroclimate index is more appropriate than the Palmer Drought Severity Index (PDSI) which  
282 requires detailed spatial soils information (Narasimhan and Srinivasan, 2005). To assess  
283 agroclimate (PSMD) variability across Swaziland a water balance model was used, working

284 from mean monthly rainfall and ETo gridded data from the IWMI baseline climatology. The  
285 PSMD for each grid pixel at the end of each month is calculated from:

$$286 \quad PSMD_i = PSMD_{i-1} + ET_i - P_i \quad [1]$$

287 Where

288  $PSMD_i$  = potential soil moisture deficit in month i, mm

289  $ET_i$  = reference evapotranspiration in month i, mm

290  $P_i$  = rainfall in month i, mm

291 At the start of the sugarcane irrigation season the PSMD is assumed to be zero. In months  
292 where  $P_i > (PSMD_{i-1} + ET_i)$ , no soil moisture deficit is assumed to occur and  $PSMD_i = 0$ . In  
293 Swaziland, soil moisture deficits start to build up each month as  $ET > P$ , peak in late summer  
294 (August) and then continue through the autumn and winter. Therefore in Swaziland, the  
295 estimation of PSMD starts with January as month  $i = 1$ . The maximum PSMD of the 12  
296 months of the year is the  $PSMD_{max}$  for that grid pixel. A gridded dataset containing the  
297  $PSMD_{max}$  for each grid pixel at a resolution of 10 min latitude/longitude (16 km x 16 km) for  
298 Swaziland was produced.

299 A modified approach was required to generate an equivalent  $PSMD_{max}$  dataset for each SRES  
300 scenario. This involved using a GIS to first downscale the HadCM3 GCM data from a grid  
301 mesh of  $2.5^\circ \times 3.75^\circ$  (latitude by longitude) down to a 10 minute grid to match the IWMI  
302 baseline climatology (New *et al.*, 2002) using a krigging interpolation technique. Tanser *et al*  
303 (2003) used a similar approach to interpolate the future climate scenario surfaces to the  
304 resolution of their long-term mean data using bilinear interpolation. The relative change  
305 between the baseline and future for each scenario, grid pixel and climate variable was then  
306 calculated:

$$307 \quad CF_{v-m,j} = \frac{V_{HadCM3\_fut\_m,j}}{V_{HadCM3\_bas\_m,j}} \quad [2]$$

308 Where:

309  $CF_{v,m,j}$  is the change factor for variable  $v$  in month  $m$  for pixel  $j$  from the HadCM3 model;

310  $V_{HadCM3\_fut\_m,j}$  is the predicted value for a climate variable from the HadCM3 model, and;

311  $V_{HadCM3\_bas\_m,j}$  is the baseline value for a climate variable in the HadCM3 model.

312 Using krigging interpolation techniques, the change factors calculated in Equation 2, were

313 then interpolated to the grid pixels in the IWMI baseline climatology. The relative change

314 between the IWMI baseline climatology and the HadCM3 future scenario for each climate

315 variable (temperature, precipitation, solar radiation, wind speed and relative humidity) for

316 each month and grid pixel was then calculated. These ‘change factors’ were then applied to

317 the IWMI baseline climatology to derive two future climate datasets at 10’ resolution:

$$318 \quad V_{10'\_fut\_m,i} = CF_{v\_m,i} \cdot V_{10'\_bas\_m,i} \quad [3]$$

319 Where:

320  $CF_{v,m,i}$  is the interpolated change factor for variable  $v$  in month  $m$  and pixel  $i$ ;

321  $V_{10'\_bas}$  is the pixel value for a climate variable in the IWMI baseline climatology, and;

322  $V_{10'\_fut}$  is the predicted value for a climate variable in the IWMI baseline climatology.

323 Two datasets containing gridded  $PSMD_{max}$  values for each SRES scenario at 10’ resolution

324 were produced. Using a GIS, the  $PSMD_{max}$  data were classified and mapped to show the

325 spatial variability in agroclimate across Swaziland for the baseline and each SRES scenario.

326 To assess the impacts of climate change on volumetric irrigation demand, a correlation

327 between irrigation need and agroclimate is necessary. One of the outputs from the

328 CANEGRO model is annual irrigation need (mm). Using Equation 1 and the climate data for

329 Mhlume, the  $PSMD_{max}$  in each simulated year (1980-96) was calculated. A correlation

330 between annual  $PSMD_{max}$  and annual irrigation need was derived by linear regression analysis

331 (Figure 5). Using a GIS, the  $PSMD_{max}$  dataset for Swaziland was combined with the  
332 regression equation (Figure 5) to estimate the irrigation need (mm) in each grid pixel. Data on  
333 the location and cropped area (ha) of sugarcane in Swaziland was obtained and imported into  
334 the GIS. The volumetric irrigation demand ( $m^3$ ) was then calculated by multiplying the  
335 reported sugarcane cropped area (ha) with the estimated irrigation need (mm) in each grid  
336 pixel. The total volumetric irrigation demand for sugarcane grown in Swaziland taking into  
337 account the agroclimate variability across the country, for the baseline and each SRES  
338 scenario, was estimated.

### 339 **3. Results and Discussion**

#### 340 *3.1 Impacts on sugarcane yield and water use efficiency*

341 The estimated changes in sucrose, biomass yield and WUE from the baseline for each SRES  
342 scenario (with  $CO_2$  fertilisation for the 2050\_A2 SRES scenario) are summarised in Table 3.  
343 With climate change, relatively minor increases in productivity are estimated, principally due  
344 to increased radiation levels and higher temperatures (1-6% and 10-29% above the baseline,  
345 respectively). This is consistent with Batchelor (1992) who observed trends of increasing  
346 growth with increasing temperature. Whilst predicted increases in sucrose yield are small (2-  
347 3%), ET was estimated to increase by between 11-14%. This results in a reduction in WUE by  
348 10% for both SRES scenarios. However, when the  $CO_2$  concentration for the baseline (330  
349 ppmv) was increased (600 ppmv) for the 2050s, there was a noticeable increase in biomass  
350 and sucrose yield. This is consistent with Watson *et al.* (1996) who reported that a doubling of  
351  $CO_2$  concentration from present levels would increase biomass by 10-30%.  $CO_2$  enrichment  
352 of the atmosphere increases the rate of photosynthesis, and thus yields, and is expected to  
353 reduce water use. In this study, the crop modelling suggests that sucrose yield under the SRES  
354 2050\_A2 scenario, with  $CO_2$  fertilisation would be 15% higher than the baseline yield. There  
355 seems to be only a minor effect of  $CO_2$  fertilisation on WUE. According to Downing *et al.*

356 (1997) a doubling of CO<sub>2</sub> concentration may increase WUE by up to 50%, with stronger  
357 effects for plants with C<sub>3</sub> pathways. The 5% WUE increase in this study is low, and possibly  
358 due to sugarcane having a C<sub>4</sub> pathway, which is less water-efficient.

359 The beneficial effects of climate change on yield due to increased CO<sub>2</sub> concentration might  
360 offset the potentially negative impacts of increased irrigation need, particularly in countries  
361 where water resources are scarce. Defining any increase in irrigation need is thus important,  
362 because if the increase in irrigation need is accompanied by an increase in yield, then in  
363 producing a unit weight of economic yield, the same amount of water may still be used, or  
364 even less. Therefore a net increase in irrigation will be when more irrigation water is required  
365 for the same unit weight of yield (defined as a standard yield). Figure 6 shows the ranked  
366 annual irrigation needs for the baseline and each future scenario for a 'standard' yield. The  
367 'standard' yield is defined as one obtained when the crop has no limitations of water. The  
368 results show that for all scenarios, there is an average increase in irrigation need from the  
369 baseline of between 19-21%. However, with CO<sub>2</sub> fertilisation, the increase in irrigation need  
370 is nearly halved (9%).

### 371 *3.2 Impacts on sugarcane irrigation water requirements*

372 The predicted changes in seasonal irrigation need (depths applied, mm) from the baseline for  
373 each SRES scenario are summarised in Figure 6. The crop modelling suggested an increase in  
374 crop water requirement (ET<sub>crop</sub>) of between 11-14% (Table 3). With climate change, the  
375 combined effect of reduced summer rainfall and increased evapotranspiration rates, results in  
376 an increase in average irrigation need of 22-26%, depending on scenario. This could have  
377 major implications for both existing sugarcane plantations and new developments because  
378 irrigation schemes (pipe distribution and canal networks, and application equipment) are  
379 designed to meet a certain 'peak' daily and seasonal need. Designing for an 'average' year  
380 would result in under-capacity in dry years when returns from irrigation are highest. Similarly,

381 designing for the driest year would lead to unnecessarily large pipes and canals and would be  
382 uneconomical. Hence most irrigation schemes are designed to meet peak need for a ‘design’  
383 dry usually defined as a return period equivalent to an 80% probability of non-exceedance.  
384 However, with increasing reliance on irrigation to attain high quality production (rather than  
385 just yield increment), combined with concerns regarding the increased likelihood of future dry  
386 years, many new irrigation schemes are now being designed to cope with more extreme  
387 events (greater than the 80% probability of non-exceedance). Figure 6 shows the potential  
388 increase in ‘design’ dry year need from the baseline for each scenario. The important point is  
389 that a future ‘average’ year in irrigation terms could well be more akin to a current ‘design’  
390 dry year, meaning that with climate change future peak irrigation needs could well exceed  
391 current design criteria for existing irrigation schemes, and in approximately 50% of years.

### 392 *3.3 Impacts on agroclimate and irrigation demand*

393 The spatial variability in agroclimate for the baseline and each SRES scenario are shown in  
394 Figure 7. For the baseline, the agroclimate zones show a strong north-south delineation, with  
395 the highest aridity values observed in the west around Big Bend (700-800 mm) then declining  
396 westwards towards Malkerns (300-400 mm). The highest aridity values correspond to where  
397 irrigation needs are highest. With climate change, the zones of highest aridity are predicted to  
398 increase in area and magnitude, moving further north towards the sugarcane growing areas of  
399 Mhlume and Simunye. Under both SRES scenarios, major changes in the spatial variability in  
400 agroclimate are predicted, with large regions of the country predicted to experience conditions  
401 more arid than those currently experienced anywhere in the country.

402 For the baseline, the total theoretical volumetric irrigation water demand is estimated to be  
403  $24138 \times 10^6 \text{ m}^3 \text{ year}^{-1}$ , with nearly three quarters (72%) concentrated within three production  
404 areas of Mhlume, Simunye, and Big Bend. With climate change, the volumetric irrigation  
405 demand in these areas is projected to increase by 18-21%. This could have major



406 repercussions on other abstractors, particularly in water scarce and trans-boundary catchments.  
407 For example, Nkomo and van der Berg (2004) investigated water availability and abstraction  
408 in the Komati river basin, which is shared by Swaziland, South Africa and Mozambique, and  
409 where irrigated sugarcane constitutes the dominant land use. They investigated the impacts of  
410 two new dams on water reliability, and found that improved water supply for irrigation in  
411 Swaziland and South Africa had been achieved. However, they reported that future water  
412 demands in 2015 even without climate change would result in appreciable shortages for  
413 irrigation. The preferred adaptation options included increasing irrigation efficiency (from  
414 surface to micro-irrigation) and reducing the sugarcane cropped area in favour of other less  
415 water demanding higher value crops, including horticulture and flowers.

#### 416 **5. Methodological limitations**

417 Inevitably, the approaches developed in this study which have linked climate, crop and GIS  
418 modelling have numerous limitations. The crop and agroclimate modelling were based on one  
419 GCM, two scenarios and one time-slice. Although the HadCM3 and SRES scenarios (A2, B2)  
420 have previously been used in various African studies (Hulme *et al.*, 2001) a more detailed  
421 assessment would need to consider a range of GCM outputs (to account for individual model  
422 error), additional time slices (2030s, 2080s) and the full ensemble of SRES scenarios (to  
423 consider alternative demographic, socio-economic and technological changes). By  
424 considering only one GCM the level of uncertainty in the model outputs cannot be easily  
425 quantified. For example, the ECHAM4 GCM has been shown to significantly increase  
426 predicted changes in irrigation demand for some regions compared to the HadCM3 GCM  
427 (Doll, 2002).

428 Arnell *et al* (2003) analysed different ways of constructing climate change scenarios using  
429 output from three climate models (HadRM3H, HadCM3, HadAM3H). Sixteen scenarios were  
430 constructed, representing different combinations of model scale (GCM, RCM), whether the

431 simulations were used directly or changes were applied to an observed baseline, and whether  
432 observed or simulated variations from year-to-year were used. The different ways of deriving  
433 climate scenarios resulted in a range in change in average annual runoff of between 10-20%  
434 by 2071-2100, depending on the model and approach used. Using a regional climate model  
435 over-estimated rainfall across much of southern Africa and resulted in excessive simulated  
436 runoff. This led to smaller estimates of change in future runoff than when changes in climate  
437 were applied to an observed climate baseline. Arnell *et al* (2003) concluded that it was  
438 preferable to apply modelled changes in climate to observed data to construct climate  
439 scenarios (as used in this study) rather than derive these directly from the regional climate  
440 model (RCM) simulations.

441 Although there has been a marked increase in the number of RCM simulations, very few  
442 studies have been conducted over southern Africa as most research institutions in this region  
443 lack access to the necessary technology. Regional models, such as HadRM3 are also able to  
444 resolve tropical cyclones, which affect eastern tropical regions of southern Africa in summer.  
445 The hydrological cycle is stronger in the RCM, with consequent increases in the intensity of  
446 rainfall, in the magnitude of the moisture fluxes and in soil moisture compared to the driving  
447 GCM (Hudson and Jones, 2002). Further studies should therefore investigate the differences  
448 in climate change signal derived from using a suitable RCM compared against using  
449 established GCM outputs to provide a better assessment of the uncertainty associated with the  
450 climate change modelling aspects of this work. Linked to this, is the method of downscaling.  
451 In this study, a popular approach using change factors (CF) was used, but this has limitations  
452 compared to statistical downscaling (SD) using transfer functions and stochastic weather  
453 generators (Diaz-Nieto and Wilby, 2005). The problem is that the future temporal pattern of  
454 wet and dry days remains unchanged, and so changes in the intensity and frequency of rainfall  
455 events can not be investigated. Further studies should consider using an alternative SD  
456 approach which would allow more detailed analysis of climate change uncertainty and

457 exploration of temporal sequencing of meteorological events (e.g. droughts, rainfall). The  
458 effect of different resolution between the HadCM3 model (2.5 x 3.75 degrees) and the IWMI  
459 baseline climatology (10' latitude/longitude) datasets, and choice of interpolation may also  
460 have introduced some distortion. Finally, the GCM outputs were used to generate future  
461 datasets based on predicted 'average' changes in climate. However, in agricultural irrigation,  
462 a statistically defined 'design' dry year with a defined probability of non-exceedance is used,  
463 rather than an 'average' year. The predicted future 'average' irrigation needs presented in this  
464 study are thus likely to significantly under-estimate future 'dry' year irrigation demand.

465 The crop model outputs are of course sensitive to model parameterisation. Further modelling  
466 would benefit from a sensitivity analysis of the key variables known to influence water use  
467 and cane yield, including modifying crop characteristics to capture the effects of varying  
468 planting dates for different ratooned cane, simulating different soil types (textures and depths),  
469 assessing the proportion of effective rainfall, and assessing the impacts of different irrigation  
470 scheduling strategies to reflect either traditional (furrow) or more efficient (micro) application  
471 methods. Modelling could also investigate the impacts of future changes in reliability of water  
472 supply; this study assumed unconstrained demand, but reducing the availability of water for  
473 irrigation at differing times during the season (for example, due to low flows or seasonal  
474 droughts) would impact on cane development and yield.

## 475 **6. Conclusions**

476 To produce a unit weight of sucrose equivalent to current optimum levels of production,  
477 future irrigation needs were predicted to increase by 20-22%. With CO<sub>2</sub>-fertilisation, the  
478 impacts of climate change are offset by higher crop yields, such that IR<sub>net</sub> is predicted to  
479 increase by 9%. The study showed that with climate change, the current peak capacity of  
480 existing irrigation schemes could fail to meet the predicted increases in irrigation demand in  
481 nearly 50% of years assuming unconstrained water availability.

482 GIS modelling confirmed that climate change will impact strongly on the spatial variability in  
483 agroclimate and hence demand for irrigation. Although the study was based on only one  
484 GCM, and considered a limited number of scenarios, these preliminary findings do highlight  
485 some of the potential risks that climate change could impose on sugarcane production in  
486 Southern Africa. The approaches developed in this paper and results serve to provide a useful  
487 baseline from which more detailed investigations should be undertaken, from which more  
488 strategic interventions, including adaptations could then be planned.

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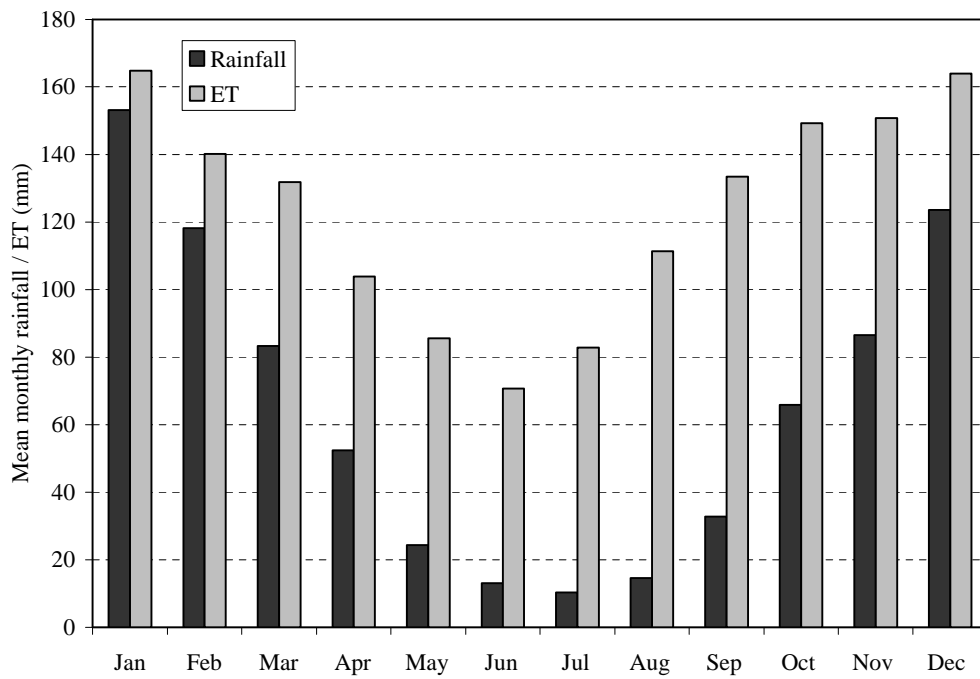
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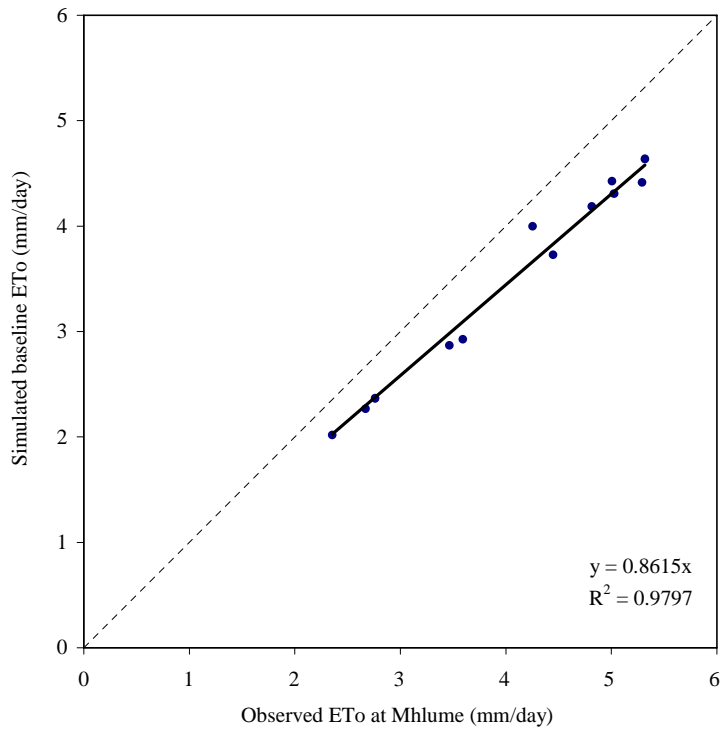
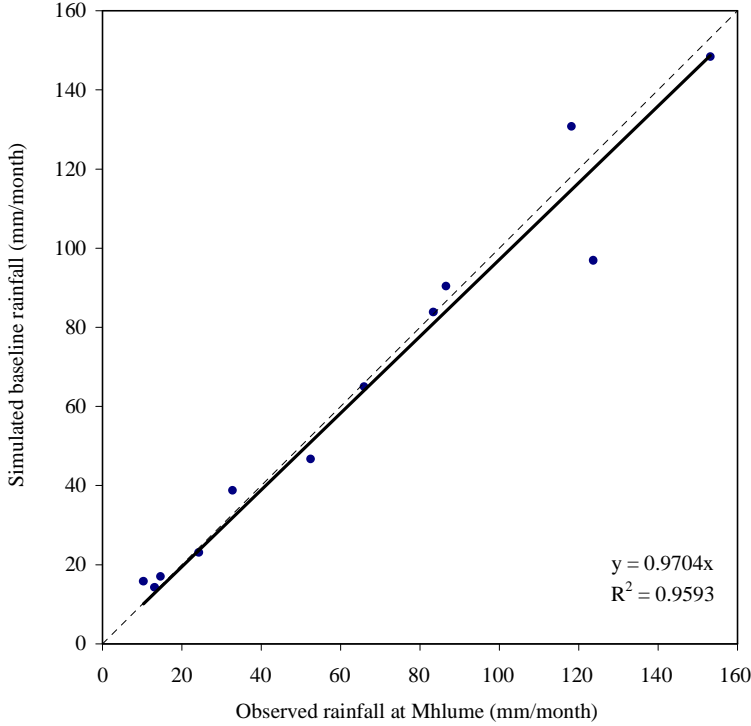
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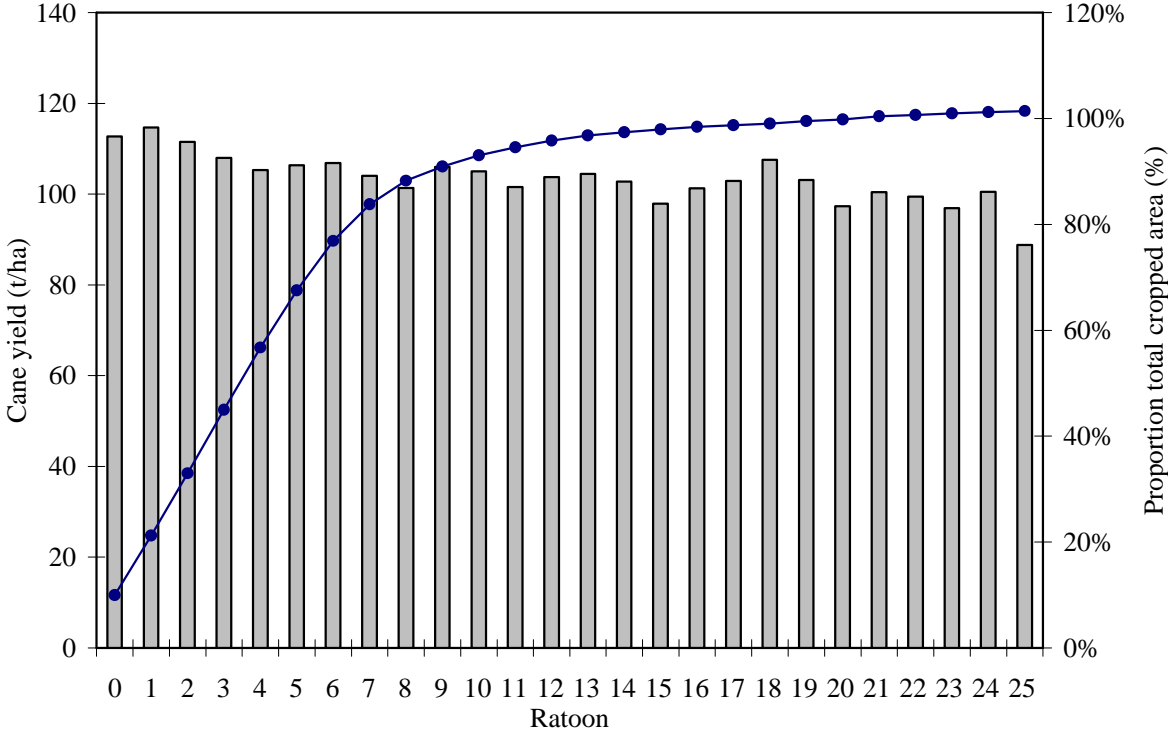
**Figure 1** Mean monthly rainfall and reference evapotranspiration (ET<sub>o</sub>) at Mhlume, Swaziland, based on daily historical data from 1969 to 1996.



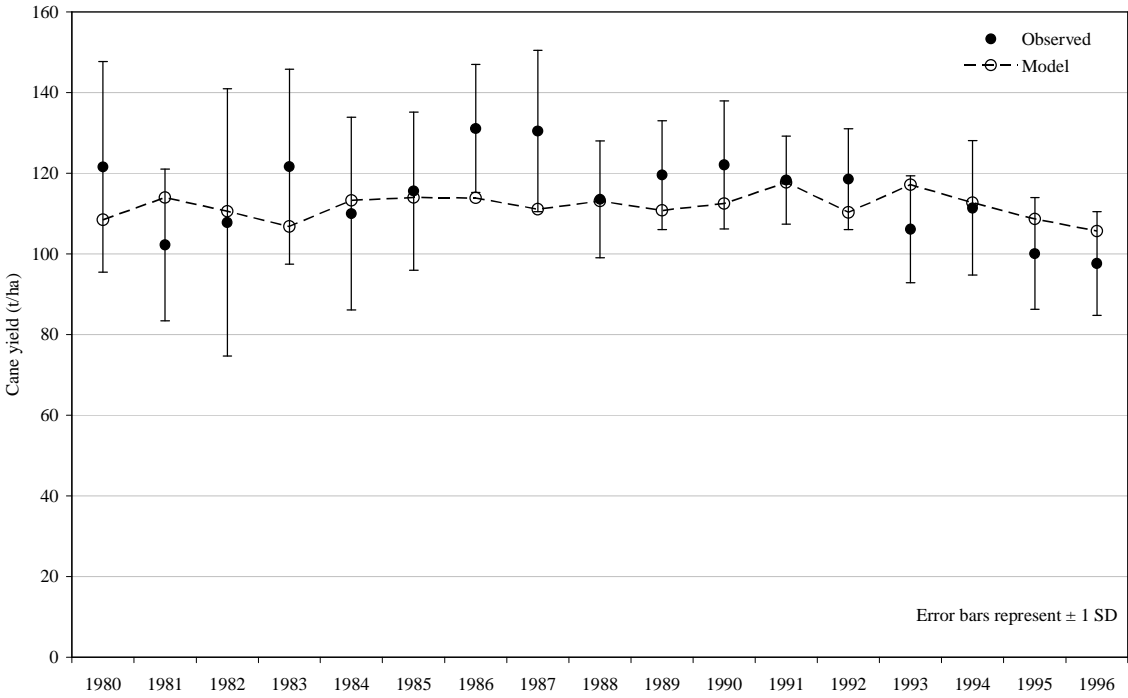
**Figure 2** Comparison of observed mean monthly rainfall (mm/month) and mean daily (mm/day) reference evapotranspiration (ETo) for Mhlume against simulated grid pixel data from the IWMI baseline climatology.



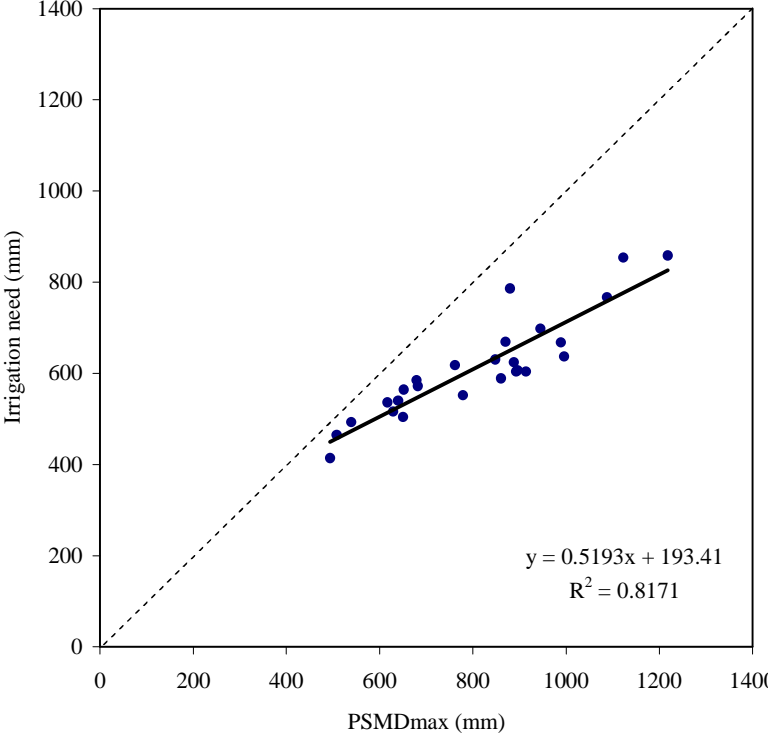
**Figure 3** Reported average cane yield (t/ha) and cumulative proportion (%) of total cropped area, by ratoon year, at Mhlume, based on data for 1980-2007.



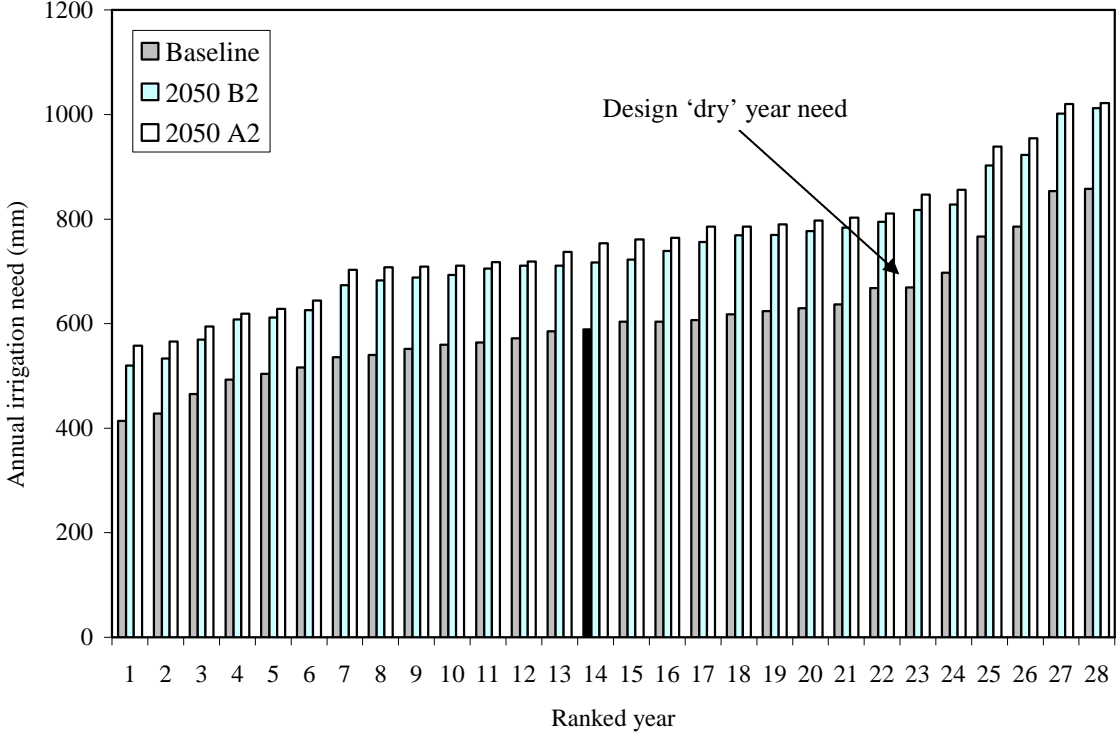
**Figure 4** Comparison between CANEGRO simulated average annual yield (t/ha) and RSSC average annual field yield (observed) between 1980-1996.



**Figure 5** Relationship between annual maximum potential soil moisture deficit (calculated using a monthly water balance Eq. 1) and annual irrigation need (calculated using CANEGRO) for Mhlume for the baseline. A linear regression was fitted to the data points.

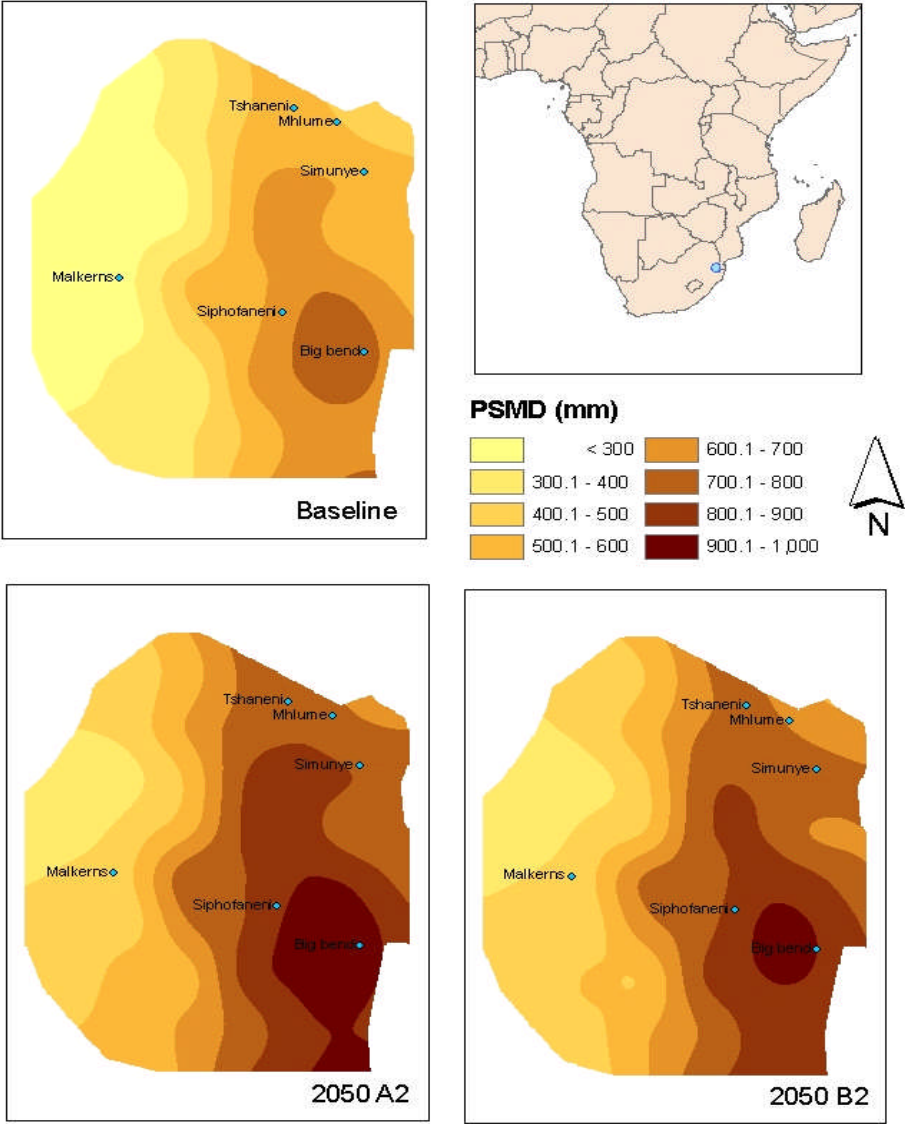


**Figure 6** CANEGRO simulated annual irrigation needs (ranked) for sugarcane at Mhlume, for the baseline and each SRES scenario (2050\_A2 and 2050\_B2). The average irrigation need for the baseline is shown in black.





**Figure 7** Spatial variability in agroclimate (PSMD) for Swaziland for the baseline (1961-90) and IPCC SRES 2050\_A2 and 2050\_B2 scenarios.



**Table 1** IPCC defined climate change scenarios (A2 and B2) and their characteristics for the 2050s (Source IPCC, 1999).

Characteristic	IPCC scenario	
	A2	B2
Population growth	High	Medium
GDP growth	Medium	Medium
Energy use	High	Medium
Global CO <sub>2</sub> emissions (GtC/yr)	17.43	11.01
Atmospheric CO <sub>2</sub> concentration (ppmv)*	547	601
Land-use changes	Medium /high	Medium
Resource availability	Low	Medium
Technological change	Slow	Medium
Change favouring	Regional	Dynamics as usual

\* Bern-CC model predictions

**Table 2.** Derived changes in mean monthly climate, between the baseline and each SRES emissions scenario, by variable and month, for Mhlume.

<b>Scenario</b>	<b>Variable</b>	<b>Jan</b>	<b>Feb</b>	<b>Mar</b>	<b>Apr</b>	<b>May</b>	<b>Jun</b>	<b>Jul</b>	<b>Aug</b>	<b>Sep</b>	<b>Oct</b>	<b>Nov</b>	<b>Dec</b>
2050 A2	Temp (° C)	2.70	2.44	2.33	2.91	3.89	4.80	4.96	3.77	3.33	3.87	2.30	2.68
	Rainfall (%)	-5.80	18.60	-23.01	-7.35	22.62	5.07	-32.71	-26.21	-21.91	-28.82	-2.73	-2.73
	Solar radiation (%)	5.06	-1.07	4.31	1.56	-3.86	-4.35	3.89	1.57	1.99	6.30	-4.69	1.67
	Wind (%)	3.03	-1.18	0.79	0.80	4.03	-7.08	1.73	5.84	6.78	9.13	0.51	0.51
	RH (%)	-3.31	-4.63	-2.61	-2.70	3.93	6.49	1.98	-8.35	-10.05	-12.20	-17.62	-2.48
	ETo (%)	12.47	7.56	10.65	11.29	9.39	8.42	18.22	20.23	19.07	24.06	12.72	9.34
2050 B2	Temp (° C)	1.87	2.19	2.21	1.98	3.30	4.33	4.81	3.36	2.97	2.54	1.45	1.92
	Rainfall (%)	-5.80	18.60	-23.01	-7.35	22.62	5.07	-32.71	-26.21	-21.91	-28.82	-2.73	-1.03
	Solar radiation (%)	2.68	2.37	4.37	3.17	-3.01	-2.83	-1.14	1.04	2.34	2.89	-3.23	4.56
	Wind (%)	-2.93	-0.28	1.43	0.09	-3.50	-3.99	7.00	3.80	6.41	7.60	1.09	-0.99
	RH (%)	-1.90	-3.25	-2.61	-2.58	4.66	7.04	0.44	-4.78	-9.24	-11.62	-12.18	-1.51
	ETo (%)	7.15	8.86	10.45	9.19	5.49	8.08	18.76	15.58	17.46	17.09	8.51	8.60

**Table 3** Modelled cane yield (t/ha), ‘design’ dry year irrigation need (mm/year) and water use efficiency ( $\text{kg}^{-1} \text{m}^3$ ) for the baseline (BL) and each climate change scenario.

<b>Output</b>	<b>BL</b>	<b>2050_A2</b>		<b>2050_B2</b>		<b>2050_A2 fert</b>	
	mm	mm	%	mm	%	mm	%
Average annual rainfall (mm)	778	738	-5	737	-5	738	-5
Average annual ETo (mm)	1161	1320	14	1292	11	1320	14
ET <sub>crop</sub> (mm)	1162	1320	14	1292	11	1320	14
IR <sub>net</sub> (mm)	605	761	26	738	22	761	26
Design irrig. need (mm)	668	811	21	795	19	811	21
Sucrose yield ( $\text{kg ha}^{-1}$ )	24747	25466	3	25168	2	28429	15
Biomass yield ( $\text{kg ha}^{-1}$ )	65835	69056	5	68202	4	75520	15
Stalk yield ( $\text{kg ha}^{-1}$ )	45457	47911	5	47262	4	52790	16
WUE ( $\text{kg/m}^3$ )	2.1	1.9	-10	1.9	-10	2.2	5
IR <sub>net</sub> ‘standard’ yield (mm)	605	739	22	726	20	662	9