

2) agrees within mutual error limits with the Sm-Nd age (Fig. 3). However, the Sm-Nd data show minor disturbances. Fig. 3 shows data obtained both in 2001 [3] and 2010 [7]. The Sm-Nd age obtained from these data is 4516 ± 37 Ma (2σ) when data for handpicked tridymite and two leachates are omitted from the isochron (open symbols). A bulk sample (WR) and pyroxene separate analysed in 2001 plot below data for corresponding samples analysed in 2010. An isochron for only 2010 data gives a slightly older calculated age of 4532 ± 57 Ma. Initial $\epsilon_{\text{Nd, HEDR}} = 0.4 \pm 0.4$ as calculated for either data set, and agrees within error limits with values of this parameter measured in the JSC lab for HED meteorites.

Rb-Sr Age: Fig. 4 shows Rb-Sr data obtained for Asuka 881394 in the 2001 and 2010 investigations. These data also show a probable isotopic disturbance. In this case, the most pronounced disturbance is for “Px1” a pyroxene-enriched sample with a density $> 3.32 \text{ g/cm}^3$ obtained as a heavy liquid separate. This sample and its leach residue align with handpicked tridymite from the 2001 investigation along an apparent 3.9 ± 0.1 Ga isochron. Because the measured $^{87}\text{Sr}/^{86}\text{Sr}$ ratio for the residue Px1(r) is in agreement with that measured in the first investigation for hand-picked Px, there is a strong suggestion that at ~ 3.9 Ga ago Rb was “sweated out” of pyroxene grain surfaces and migrated into late-stage tridymite, itself likely to have formed during granoblastic metamorphism of the rock. Plausibly this effect was minimal for the larger and purer handpicked pyroxene grains. Heating experiments for a eucrite showed that initial melting occurred along grain boundaries [12].

Discussion: The isotopic record in A881394 must be interpreted in the context of its granoblastic texture. The good concordance of the Al-Mg and Mn-Cr formation intervals suggests that the granoblastic texture formed before those isotopic systems closed. Later disturbance of the Sm-Nd and Rb-Sr systems appears

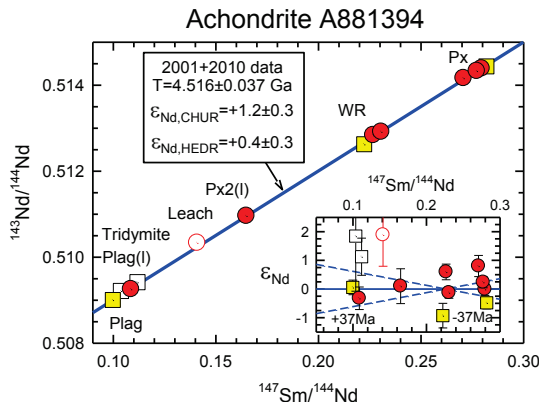


Figure 3. Sm-Nd data for Asuka 881394. Age calculated assuming $\lambda(^{147}\text{Sm}) = 0.00654 \text{ Ga}^{-1}$ [11].

to be limited to minor phases including tridymite and possibly Ca phosphate and ilmenite as well [12]. The Nd- and Sr-isotopic systems appear to have remained closed in the major mineral plagioclase. Thus, initial $^{87}\text{Sr}/^{86}\text{Sr} = 0.698989 \pm 14$ is indistinguishable from that for eucrites or angrites, whereas initial ϵ_{Nd} agrees with values measured for HED meteorites. These observations are consistent with the primordial Pb isotopic composition (CDT) in plagioclase reported by [6]. The observations suggest that the (most radiogenic+CDT) regression [6] yielding an age of 4566.05 ± 0.45 Ma should be preferred to one for which modern terrestrial Pb is included in the regression. However, the K-Ar system suggests that plagioclase was not closed to Ar-migration, and *in situ* SIMS analyses showed variations in $\delta^{26}\text{Mg}$ that were not well correlated with Al/Mg ratios in plagioclase [5]. Thus, caution in using Asuka 881394 data to evaluate the initial homogeneity of the distribution of short-lived nuclides in the early solar nebula is indicated.

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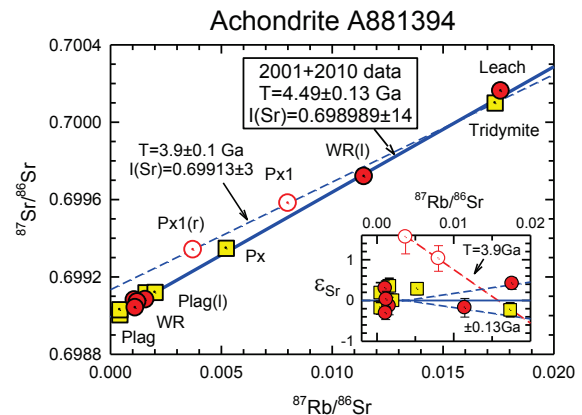


Figure 4. Rb-Sr data for Asuka 881394. Age calculated assuming $\lambda(^{87}\text{Rb}) = 0.01402 \text{ Ga}^{-1}$ [11].