POPULAR SUMMARY 1

Satellite observation highlights of the 2010 Russian Wildfires

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From mid-June 2010 through August western Russia experienced an unprecedented heat wave characterized by prolonged high temperatures (~ 40degC) and drought conditions. The heat wave set-up the ideal meteorological conditions for an unprecedented wildfire event. Plumes of thick smoke and burning pollution products were reported over highly populated regions including the capitol city of Moscow. The negative human and economic impacts were severe and extensively covered by the local and international media.

Our study took advantage of the large complement of NASA's Earth Observing System (EOS) sensors to track and quantify the source of the thick smoke and wildfire byproducts, such as carbon monoxide (CO), which settled over Moscow and nearby cities. A typical tracer of carbonaceous (or smoke) aerosols produced from wildfires smoke aerosols is the Aerosol Index (AI) that is measured by the Ozone Monitoring Instrument (OMI) on-board the Aura satellite. Over Moscow, OMI measured unprecedented levels of smoke aerosols between the end of July and mid-August that were an order of magnitude higher than previous summers going back to the earliest record in 2005. Likewise, CO, measured by the Atmospheric Infrared Sounder (AIRS) on-board Aqua, showed exceptionally high levels over Moscow. Previous summers going back to 2003 typically have an average CO concentration of $20x10^{17}$ molecules/cm². To put this wildfire event into perspective, the magnitude of the CO we observed over Moscow was equivalent to the 2006 El Nino wildfire event over Indonesia where some of the most intense wildfires have been documented.

Using the MODIS fire count data on-board the Aqua and Terra satellites, we observed numerous wildfires throughout western Russia and Eastern Europe that raged for almost three weeks between the end of July and mid-August. During this time period air-parcel back-trajectories initiated from Moscow traced the origin of the enhanced smoke pollution from wildfires raging in the southeast. The MODIS Fire Radiative Power (FRP) product measured very intense of the fires clustered in that same region and AIRS CO was also historically high.

1 Satellite observation highlights of the 2010 Russian Wildfires 2 Jacquelyn C. Witte¹, Anne R. Douglass², Bryan N. Duncan², Arlindo M. da Silva², and 3 Omar Torres² 4 5 6 ¹Science Systems and Applications Inc. Lanham, MD, 20703, USA ²NASA Goddard Space Flight Center, Code 613.3, Greenbelt, MD 20771, USA 7 8 9 Abstract 10 From late-July through mid-August 2010, wildfires raged in western Russia. The 11 resulting thick smoke and biomass burning products were transported over the highly 12 13 populated Moscow city and surrounding regions, seriously impairing visibility and affecting human health. We demonstrate the uniqueness of the 2010 Russian wildfires by 14 using satellite observations from NASA's Earth Observing System (EOS) platforms. 15 16 Over Moscow and the region of major fire activity to the southeast, we calculate unprecedented increases in the MODIS fire count record of 178 %, an order of magnitude 17 18 increase in the MODIS fire radiative power (308%) and OMI absorbing aerosols (255%), 19 and a 58% increase in AIRS total carbon monoxide (CO). The exceptionally high levels of CO are shown to be of comparable strength to the 2006 El Niño wildfires over 20 Indonesia. Both events record CO values exceeding 30×10^{17} molec-cm². 21 22 23 1. Introduction 24

Forest fires are both a source and sink of carbon, releasing carbon dioxide (CO₂). 25 26 and CO to the atmosphere while burning, and removing of CO₂ during post-fire re-27 growth, thus playing an important role in the global carbon cycle [Olson et al., 1983; 28 Crutzen and Andreae, 1990; Kasischke et al., 2005]. Russia includes approximately 30% 29 of the world's total forested area, and forest fires are common [Alimov et al., 1989]. 30 Despite improvements in spatio-temporal coverage of fire events due to satellite 31 monitoring, the behavior of forest ecosystems under fire conditions remains uncertain 32 [Mottram et al., 2005]. Forecasting the influence of forest fires on regional and global 33 scales remains a challenge.

34 The 2010 Russian wildfires was an unprecedented forest fire event that spread 35 dangerously towards populated regions, significantly impacting human health and 36 livelihood. Media coverage was extensive and socio-economic statistics and impacts are 37 readily available in the on-line literature. A blocking high-pressure system over western 38 Russia and parts of Eastern Europe resulted in anomalously high temperatures and dry 39 conditions in July and August 2010 [Lau and Kim, 2010]. Our study will show that from 40 late-July through mid-August the circulation produced by the blocking event transported 41 heavy wildfire smoke and burning byproducts, such as CO, over Moscow and surrounding regions. Consequently, the city experienced impaired visibility and 42 43 unhealthy levels of smoke and smog, compounded by local pollution sources. We use 44 observations of fire activity, smoke, and CO from several sensors on NASA's EOS 45 platforms including Aura, Aqua, and Terra to show that the 2010 Russian wildfires are 46 unique in the observing records of these sensors.

The next section describes the observations used in this study, followed by the analyses in section 3. Included is an overview of the regional meteorological conditions from daily radiosondes at various locations in the western Russia, including Moscow. We also compare the 2010 Russian wildfires with the 2006 El Niño wildfires in Indonesia.

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52 2. Satellite Data

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2.1 Active Fire Counts and Fire Radiative Power

56 Active fire count data are taken from the Moderate Resolution Imaging 57 Spectroradiometer (MODIS) instruments on NASA's Terra and Aqua satellites that were 58 launched in December 1999 and May 2002, respectively. The Level 2 active fires 59 products MOD14 (Terra) and MYD14 (Aqua) have a pixel size of 1 km² at nadir 60 covering an area of 2340×2030 km in the across- and along-track directions, respectively. 61 The fire detection algorithm is described in *Giglio et al.*, [2003] and has been shown to 62 provide useful information about the spatial and temporal dynamics of fire activity 63 [Giglio et al, 2006 and references therein]. The fire detection strategy is based on 64 absolute detection of a fire (when the fire strength is sufficient to detect), and on 65 detection relative to its background (to account for variability of the surface temperature 66 and reflection by sunlight). The Fire Radiative Power (FRP, in Megawatts) measures the 67 radiant heat output of the detected fires. Kaufman et al. [1996] developed an empirical 68 non-linear relationship between the MODIS mid-infrared channel brightness 69 temperatures at an active fire pixel, and the fire FRP over all wavelengths.

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71 2.2 UV Aerosol Index

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73 The Dutch-Finnish OMI instrument is a nadir-viewing moderate resolution 74 UV/Vis spectrometer on NASA's Aura satellite, launched on July 2004 into a sun-75 synchronous orbit with an equator crossing-time of 13:38 in the ascending node. OMI has a full cross-track swath of 2600 km, containing 60 pixels ranging from 15×30 km² at 76 77 nadir to 42×162 km² at the edge of the swath. The OMAERUV aerosol algorithm uses 78 the radiances measured at 354 and 388 nm to retrieve the UV Aerosol Index (UVAI). 79 *Torres et al.* [2007] and references therein describe the algorithm that was originally 80 developed for TOMS (Total Ozone Mapping Spectrometer). All UVAI data have been 81 filtered for the row anomalies that have affected the Level 2 data since 2007. Detailed 82 information on the OMI row anomaly can be found at

83 http://www.knmi.nl/omi/research/product/rowanomaly-background.php.

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85 2.3 Total Column CO

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87 Aqua's Atmospheric Infrared Sounder (AIRS) is a cross-track scanning grating 88 spectrometer that provides total column CO (CO_{TC}) data with a nadir 45 km field of view 89 across a 1650 km swath [*Aumann et al.*, 2003]. CO_{TC} has an estimated uncertainty for an 90 individual measurement of 7–8% with standard deviations between ±2 and ±6% 91 [*Yurganov et al.*, 2002]. AIRS Science Team Version 5 Level 2 daytime swaths are used

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94 3 Analyses

9596 3.1 Unique Fire Event

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Figure 1a shows the location of all the active fires detected by MODIS Aqua and Terra for August 2010. We observe that FRP values greater than 500 Mwatts (yellow circles) are clustered southeast of Moscow (black star). This is a region of mixed and deciduous forest with a portion consisting of peatland (USSR State Forestry Committee, 1990). We focus on the southeast domain (cyan box, referred to as SE) covering 51-57°N and 37-50°E and tally the fire counts and FRP within that domain. Results are plotted in Figures 1a and 1b for June through August since 2003. We observe the following:

a) The fires are triggered earlier in 2010 than any previous year. On July 25th 2010,
the fires ramp-up and sustain very high levels of activity and intensity through midAugust. By August 14th the fires begin to wane, while prior years show the fire products
ramping up at this time and peaking later in the month.

b) Compared to prior years, the fires from late-July to mid-August 2010 are the
most numerous and intense (two exceptions in FRP in 2008). Table 1 summarizes fire
counts and FRP between July 25th and August 31st to include the 2010 fire event and
seasonal fires that prior years show occurring throughout August. The statistics reveal
exceptionally high values in 2010; FRP is an order of magnitude larger than previous
years and the fire counts are almost doubled. The percentage increases in 2010 relative to
the 2003-2009 mean are exceptionally high: 178% for fire count, and 309% for FNR.

116 At present, satellite measurements of fire activity are the best estimates of fire 117 detection and strength [Mottram et al., 2005; Rov et al., 2008], however, it is important to 118 keep the limitations of this data set in mind. In the vicinity of heavy clouds and very large 119 fires the MODIS FRP may be less intense or not detected, resulting in a systematic low-120 bias in the measurements [Giglio et al., 2006]. Ground fires, such as peat fires in our SE 121 domain, generally do not produce sufficient heat to be detected by MODIS [Rov et al., 2008]. Only subsets of fires are captured due to the relatively large viewing geometry, i.e. 122 pixel sizes ranging from 1 km² at nadir to 4-5 km² at edge of the swath. Thus, although 123 124 MODIS captures record fires over western Russia, the actual fire detected and intensity 125 (in FRP) may be much higher.

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127 Table 1. Combined MODIS Aqua and Terra fire counts and FRP in the SE domain [51-

128 $57^{\circ}N$, $37-50^{\circ}E$] between July 25^{th} and August 31^{st} per year.

Year	Fire Count	FRP [$\times 10^4$ Mwatt]
2010	26,876	166.568
2003-2009	9683	40.729
2009	6,784	33.270
2008	18,004	94.180
2007	11,206	49.842
2006	8,582	32.320
2005	12,873	43.224
2004	8,973	28.839
2003	1.358	3.428

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130 3.2 Anomalous Surface Temperatures and Relative Humidity

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Anomalous Surface Temperatures and Relative Humberry

132 Lau and Kim [2010] provides a thorough analysis of the synoptic weather patterns over western Russia that set-up the ideal conditions for the wildfires to thrive, spread and 133 134 intensify for a prolonged period of time. Trajectory results from the NOAA HYSPLIT 135 trajectory model [Draxler and Rolph, 2010] reveal the circulation pattern produced by this blocking event. Daily clusters of backward and forward trajectories initiated during 136 137 the Aqua and Terra overpass time's, for the first week in August (during the height of the 138 fires activity: see Fig. 1) from the Moscow city center $(37.6^{\circ}N, 55.7^{\circ}E)$ at levels ranging 139 from 0.5km up to 5km show a general clockwise motion indicative of a high-pressure 140 system (not shown). Forward trajectories from clusters of fires SE of the city show transport of air toward Moscow. Surface temperate (T_{sfc}) and relative humidity (RH_{sfc}) 141 anomalies from 12Z daily radiosonde measurements are plotted in Figure 2 at Moscow 142 143 (red) and nearby stations (locations in Fig. 2c). Data were taken from the NOAA/Earth 144 System Research Laboratory archive (http://www.esrl.noaa.gov/raobs/) going back to 1994. The daily anomalies are calculated by subtracting the 1994-2009 T_{sfc} and RH_{sfc} 145 mean from their respective 2010 values. Focusing on the summertime period accentuates 146 147 the anomalously high (positive) T_{sfc} and low (negative) RH_{sfc}, relative to 2010, associated 148 with the blocking high-pressure system. From mid-June to mid-August the range of 149 maximum T_{sfc} and minimum RH_{sfc} at these sites is 35-41°C and 9-25%, respectively. 150 These anomalous meteorological parameters are coincident with the maximum time period of the fires, observed in Figure 1. Wind directions from the surface up to 700 hPa 151 152 from late-July to early-August 2010 are predominantly from the SE quadrant (Fig. 2d, 153 grey shaded). This further substantiates our claim that the smoke plumes reported over 154 Moscow during the height of the fire activity (Fig. 1) originated from wildfires largely 155 within the SE domain.

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157 3.3 Exceptional Smoke and CO_{TC}

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5.5 Exceptional Shloke and CO_{1C}

159 The OMI UVAI is a useful parameter for tracking absorbing aerosols (i.e. smoke) 160 even in the presence of clouds [deGraaf et al., 2005] and a few recent studies have used 161 UVAI observations to link smoke plumes to biomass burning regions [Fromm et al., 162 2005; Torres et al., 2007; Christopher et al., 2008]. The UVAI 3-day running mean over the Moscow domain $(1^{\circ} \times 1^{\circ})$ area average around the city center) is plotted in Figure 3a 163 164 measuring exceptionally high smoke (>> 1) in early August 2010 (red). Values greater than 1 rarely occur in previous years. Coincident with the start of the fire activity (Fig. 1), 165 UVAI builds from July 25th, returning to mean values after mid-August. Between August 166 167 5^{th} and 10^{th} UVAI values > 2 and large1-sigma standard deviations > ±0.4 are observed, 168 not seen in previous years. We do not show the UVAI within our SE domain because of 169 significant under-sampling due to the row anomalies which, since 2009, affects almost 170 half the OMI cross-track positions. The sparse data that are available qualitatively 171 support the presence of elevated levels of UVAI in the SE domain.

172 The AIRS CO_{TC} over the Moscow and SE domains is plotted in Figure 3b. Again 173 we see unprecedented levels of CO_{TC} peaking on August 7th of 37.1±5.1×10¹⁷ molec-cm² 174 in the Moscow domain (red) and 39.1±4.4×10¹⁷ molec cm² in the SE domain (purple). 175 We also observe elevated levels at the end of July 2006 (crosses) coincident with the 176 UVAI in Fig. 3a indicating another fire event, although short-lived compared to what we 177 observe in 2010. Interestingly, as with the UVAI, the 1-sigma standard deviations are also large in both domains. The highest estimate occurs on August 1^{st} at $\pm 7.4 \times 10^{17}$ 178 molec-cm² and $\pm 5.4 \times 10^{17}$ molec-cm² over the Moscow and SE domains, respectively. 179 180 CO_{TC} values over both domains are comparable in magnitude to that over Sumatra and Borneo, Indonesia during the 2006 El Niño wildfires (Fig. 3b, grey dotted and dashed 181 lines, respectively). Exceptionally high values exceeding 30×10^{17} molec-cm² are 182 observed during both events. 183

Table 2 highlights the record high levels of CO_{TC} and UVAI in 2010 (bold) 184 relative to prior years during the August 1-18 peak period. CO_{TC} over Moscow and SE 185 domains increases 53% and 58%, respectively, in 2010 relative to the 2003-2009 mean. 186 187 UVAI over the Moscow domain increases an order of magnitude ($\sim 255\%$), relative to the 188 2005-2009 mean. Values of CO_{TC} in 2010 over both domains, including their 1-sigma 189 standard deviations are similar to what we calculate over Sumatra and Borneo. During their peak period between October 10 and November 11, 2006 we estimate Sumatra 190 CO_{TC} to be 34.7±3.9 ×10¹⁷ molec-cm² and 34.9±5.3 ×10¹⁷ molec-cm² over Borneo. 191

192 There is a dip in the CO_{TC} measurements over the Moscow domain on August 193 11th and 12th, followed by a second peak in mid-August (13th-18th). The UVAI also shows 194 a slight secondary peak (Fig. 3a). Trajectory analyses on the 11th and 12th show winds 195 from the SE domain transporting smoke eastward, away from Moscow, while the city 196 receives air primarily from the south and southwest where fires continue to erupt and 197 transport smoke (Fig. 1). However, CO_{TC} remains much higher relative to previous years.

After August 18th CO_{TC} and UVAI return to values typical of previous years. 198 199 Noteworthy is the absence of elevated CO_{TC} in the SE domain in August from 2003 to 200 2009 concurrent with elevated active fires (Fig. 1). This may be due to the type of 201 vegetation being burned in this region and/or that CO_{TC} is largely confined in the 202 boundary layer where AIRS retrievals are less sensitive [Yurganov et al., 2007]. Beside 203 wildfires, peat fires may be contributing to the exceptionally high CO_{TC} in 2010. A 204 significant portion of peat in Russia (60 Mt yr -1) is used as fuel [Kolchugina and Vinson, 205 1993]. In particular, peat fires are known to smolder for prolonged periods of time and emit large reservoirs of carbon, in the form for CO [Immirzi et al., 1992; Kasischke et al., 206 207 2005]. The degree of involvement of peat fires during this event and the altitude of the 208 fire plumes prior to 2010 requires further study.

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210	Table 2. August 1-18 mean per year of CO_{TC} [×10 ¹⁷ molec-cm ²] and UVAI [unitless] for
211	the Moscow domain and SE domain (CO_{TC} only).

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SE Domain		Moscow Domain	
	[51-57°N and 37-50°E]	$[1^{\circ} \times 1^{\circ}$ mean around the city center]	
Year	$CO_{TC} \pm 1 - \sigma$	$CO_{TC} \pm 1 - \sigma$	UVAI
2010	32.43 ± 5.05	29.47 ± 2.62	1.49 ± 0.58
2003-2009	20.48 ± 1.05	19.24 ± 0.58	$2005-2009: 0.42 \pm 0.19$
2009	18.78 ± 0.98	17.34 ± 0.50	0.42 ± 0.13
2008	19.69 ± 1.04	17.91 ± 0.38	0.44 ± 0.23
2007	20.34 ± 1.01	18.69 ± 0.57	0.38 ± 0.18
2006	20.99 ± 1.19	19.62 ± 0.53	0.40 ± 0.23
2005	20.15 ± 0.88	18.64 ± 0.68	0.44 ± 0.20

2004	21.30 ± 1.19	20.58 ± 0.76	N/A
2003	22.09 ± 1.05	21.90 ± 0.65	N/A

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213 4. Summary

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215 EOS satellite multi-sensor data has enhanced our capability of tracking major 216 atmospheric events, such as the 2010 Russian wildfires. Radiosondes stationed in western 217 Russia measure anomalously high T_{sfc} (> 35°C) and low RH_{sfc} (< 25%) during the 218 summer months. From late-July to mid-August 2010, record fires south and east of 219 Moscow were observed by MODIS. The percentage increases in fire counts and FNR, 220 relative to the 2003-2009 mean, are 178% and 309%, respectively. OMI UVAI over 221 Moscow is an order of magnitude higher than previous years (255% increase). Likewise, 222 AIRS CO_{TC} is 53% and 58% higher over the Moscow and SE domains, respectively. Exceptionally high CO_{TC} during the peak period of the Russian wildfires are comparable 223 224 to what is observed during the 2006 El Niño wildfires over Sumatra and Borneo. Both events showed values exceeding 30×10^{17} molec-cm². After mid-August, MODIS fire 225 226 activity drops well below what is typically seen in previous years, while UVAI and CO_{TC} 227 return to mean values. 228 229 Acknowledgement: This work is supported by NASA's Atmospheric Chemistry, 230 Modeling and Analysis, and Applied Sciences Air Quality Programs. 231 232 References 233 Alimov, Y.P., I.V. Golovikhin, L.B. Zdanevich, and I.V. Yunov (1989). Dynamics 234 235 of forests under forest management organization regarding the main forest forming 236 species in 1966-1988, U.S.S.R State Forestry Committee, pp. 159, Moscow. 237 238 Aumann H. H. et al. (2003), AIRS/AMSU/HSB on the Aqua Mission: Design, 239 Science Objectives, Data Products, and Processing Systems, IEEE Trans. Geo. Rem. 240 Sens., 41, 253-264. 241 242 Christopher, S. A., P. Gupta, J. Haywood, and G. Greed (2008), Aerosol optical 243 thicknesses over North Africa: 1. Development of a product for model validation using 244 Ozone Monitoring Instrument, Multiangle Imaging Spectroradiometer, and Aerosol 245 Robotic Network, J. Geophys. Res., 113, D00C04, doi:10.1029/2007JD009446. 246 247 Crutzen, P. J. and M. O. Andreae (1990), Biomass Burning in the Tropics: Impact on 248 Atmospheric Chemistry and Biogeochemical Cycles, Science, 250, 1669–1678. 249 250 deGraaf, M., P. Stammes, O. Torres, and R. B. A. Koelemeijer (2005), Absorbing 251 Aerosol Index: Sensitivity analysis, application to GOME and comparison with TOMS, J. 252 Geophys. Res., 110, D01201, doi:10.1029/2004JD005178. 253 254 Draxler, R.R. and Rolph, G.D. (2010), HYSPLIT (HYbrid Single-Particle Lagrangian 255 Integrated Trajectory) Model access via NOAA ARL READY Website

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317 Figure 1. Combined Aqua and Terra fire counts and FNR over the SE domain are plotted

in (a) and (b), respectively. Map inset in (a) of Aqua and Terra fire counts (red circles).

319 Yellow circles indicate FNR > 500 Mwatts. Star marks the location of Moscow. The cyan

box defines the SE domain covering $51-57^{\circ}N$ and $37-50^{\circ}E$.



Figure 2. (a) T_{sfc} and (b) RH_{sfc} anomalies. Radiosonde locations, including symbol legend for (a) and (b) are mapped in (c). Wind directions in July and August 2010 are shown in (d) at the surface (red), 850hPa (blue), and 700hPa (green). Grey shading highlights the southeast quadrant.





- 328 Figure 3. (a) UVAI plotted per year over the Moscow domain between June and August.
- 329 2010 is highlighted in red. CO_{TC} in (b) plotted similar to (a), also including the SE
- domain (purple). The 1- σ standard deviations are plotted for 2010 only in grey shading.
- 331 Sumatra [96-107°E, 7°S-5°N] (grey dotted) and Borneo [109-119°E, 5°S-5°N] (grey
- dashed) are overlaid in (b) for September November 2006.

