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Clyde superficial deposits and bedrock models released to the ASK Network 2013 : a guide for users

Geology and Landscape Programme

Open Report OR/13/002

BRITISH GEOLOGICAL SURVEY

GEOLOGY AND LANDSCAPE PROGRAMME

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S Arkley, K Whitbread, A Monaghan

Contributor D Lawrence

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British Geological Survey offices

BGS Central Enquiries Desk

Tel 0115 936 3143 Fax 0115 936 3276
email enquiries@bgs.ac.uk

Environmental Science Centre, Keyworth, NottinghamNG12 5GG

Tel 0115 936 3241 Fax 0115 936 3488
email sales@bgs.ac.uk

Murchison House, West Mains Road, EdinburghEH9 3LA

Tel 0131 667 1000 Fax 0131 668 2683
email scotsales@bgs.ac.uk

Natural History Museum, Cromwell Road, LondonSW7 5BD

Tel 020 7589 4090 Fax 020 7584 8270
Tel 020 7942 5344/45 email bgslondon@bgs.ac.uk

Columbus House, Greenmeadow Springs, Tongwynlais, CardiffCF15 7NE

Tel 029 2052 1962 Fax 029 2052 1963

Maclean Building, Crowmarsh Gifford, Wallingford OX10 8BB

Tel 01491 838800 Fax 01491 692345

Geological Survey of Northern Ireland, Colby House, Stranmillis Court, BelfastBT9 5BF

Tel 028 9038 8462 Fax 028 9038 8461

www.bgs.ac.uk/gsni/

Parent Body

Natural Environment Research Council, Polaris House, North Star Avenue, SwindonSN2 1EU

Tel 01793 411500 Fax 01793 411501
www.nerc.ac.uk

Website www.bgs.ac.uk

Shop online at www.geologyshop.com

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Summary

This report draft provides an overview of the Clyde superficial deposits models to be released in 2013 and detail on the Central Glasgow Superficial Deposits Model currently released to the ASK network. It also provides information on the Central Glasgow Bedrock Model currently

released. The geological models are an interpretation of digital datasets held by the British Geological Survey.

A summary of the construction and limitations of the models and a brief description of the modelled units is given. The report will be updated and revised as more models become available for release to the ASK network.

More details on the models can be found in the previous reports Merritt et al. (2009), Monaghan (2012a) and Monaghan et al. (2012).

1 Modelled volume, purpose and scale

1.1 SUPERFICIAL DEPOSITS

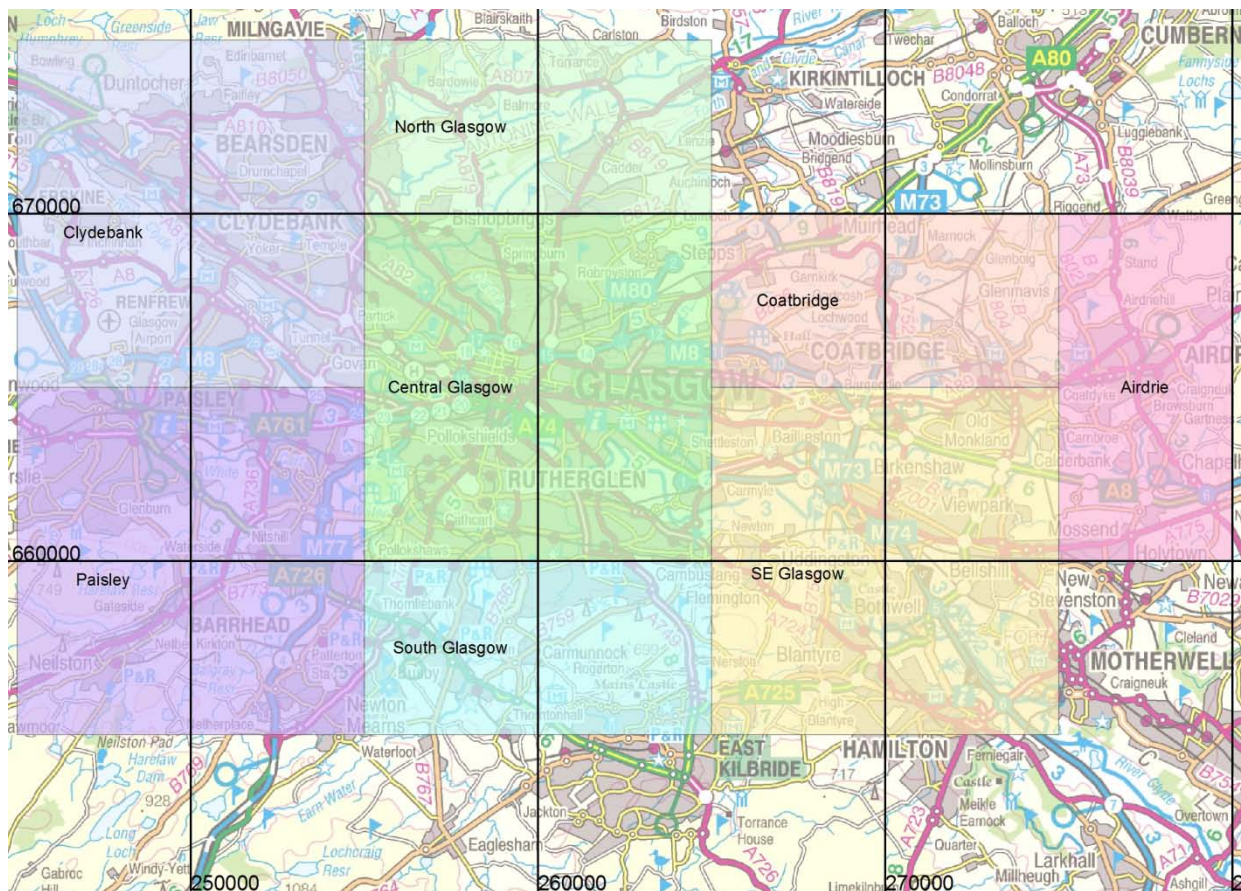


Figure 1 Superficial deposits model areas.

The superficial deposits models of the Glasgow area were constructed in tiled areas (Figure 1) for reasons of computing power and software capability. Currently released to the ASK network is the Central Glasgow tile which has the corner coordinates SW 255000, 660000 to NE 265000, 670000. Later in 2013 an approximately 5 km buffer to cover the Glasgow City Council areal extent will be released to the ASK Network. It will cover an area around the Central Glasgow tile and will have the corner coordinates SW 250000, 655000 to NE 270000, 680000. The model Z values or depths are all given in metres relative to Ordnance Datum (OD) and range from approximately +200 to -150 m.

The Central Glasgow Superficial Deposits Model is most suitable for use at scales between 1:10,000 to 1: 50,000 but is also useful for providing guidance at other scales.

The purpose of the superficial deposits model is to give a broad indication of the likely subsurface sequence from Devensian glacial till to recent made ground across Glasgow for use during the desk study stage, in for example the assessment of potential infrastructure alignments

and locations, and relatively broad scales of planning. However, the model is also well suited to act as a *guide* for studies at more detailed scales, and especially as a basis for planning and improving the economic efficiency of, site investigations. However, users are advised against basing interpretations on extreme exaggeration of vertical or horizontal scales.

The Central Glasgow Superficial Deposits Model was improved in 2012/13 to be fully consistent with surrounding models and to include some new cross-sections in the south-west part of the model, compared to earlier versions supplied by the British Geological Survey to Clyde Gateway Developments Ltd, Glasgow City Council, South Lanarkshire Council and Scottish Enterprise (see section 4.1 below). The new version supplied to the ASK network is version 3 of this model.

1.2 BEDROCK

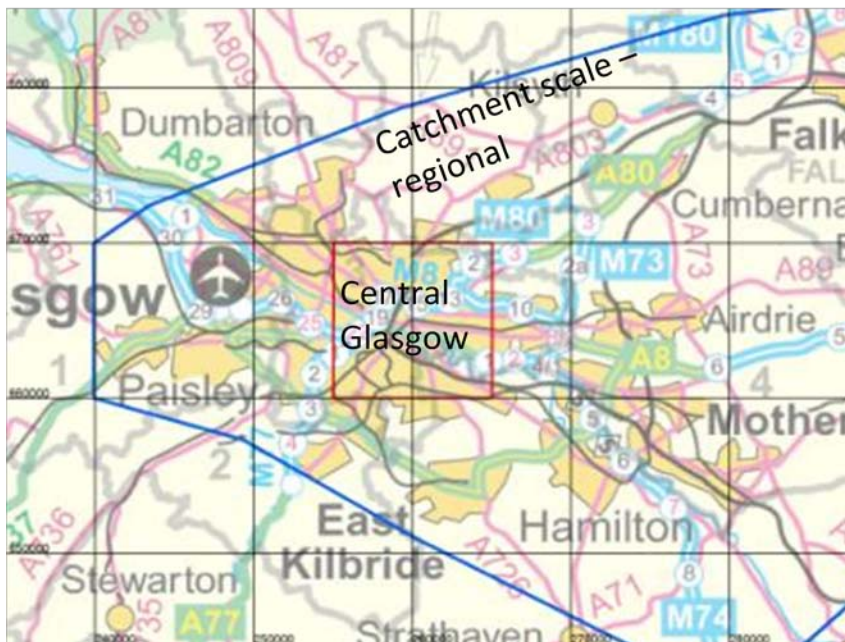


Figure 2 Image showing the extent of the Central Glasgow Bedrock Model and the lower resolution catchment scale model to be provided as the 5 km buffer later in 2013.

Bedrock modelling has been undertaken at a variety of scales by the British Geological Survey in the Clyde Catchment. The model scale is dependent on the data available and used, the number of coal seams, stratigraphic surfaces and faults included in the model.

The Central Glasgow Bedrock Model provided to the ASK network is the same as the Version 2 model supplied to Clyde Gateway Developments Ltd, Glasgow City Council, South Lanarkshire Council and Scottish Enterprise in early 2012. Thus currently released is the Central Glasgow bedrock model which has the corner coordinates SW 255000, 660000 to NE 265000, 670000. Later in 2013 a 5 km buffer bedrock model will be released around the Central Glasgow area which will be of lower resolution and have the corner coordinates SW 250000, 655000 to NE 270000, 680000. The model Z values or depths are all given in metres relative to Ordnance Datum (OD) with the modelled surfaces extend to c.-650 m depth, with the largest faults projecting down to c. -1km

The Central Glasgow Bedrock Model is most suitable for use at scales between 1:10,000 to 1:50,000 but is also useful for providing guidance at other scales.

The purpose of the bedrock model was to specify in 3D the geometry and faulting of some key, mined coal seams for use in outline planning and development, and key stratigraphic surfaces for regional geological correlation.

1.3 DATASETS PROVIDED TO USERS

The Central Glasgow Superficial Deposits Model was created in the GSI3D modelling software and has been converted to ASCII grids and ESRI (ArcGIS) raster grids for supply to users, and encrypted within Lithoframe Viewer software for 3D visualisation, synthetic sections and boreholes. Outlines of the modelled unit extent are supplied as ESRI shapefiles. The Lithoframe Viewer software and an accompanying manual can be downloaded from the BGS website at <http://www.bgs.ac.uk/services/3Dgeology/lithoframeSamples.html>. Note that in some versions of Windows it is necessary to paste in the .GSIPRe file location and name to the 'Load model' screen, and that creating a horizontal slice in the Lithoframe Viewer takes a considerable time (20 minutes).

The resolution of the superficial deposit grids is 25 m.

The Central Glasgow Bedrock Model was created in the GOCAD® modelling software and has been converted to ASCII grids for supply to users. The surface grids and a combined fault grid were exported from the modelled TIN's using a grid spacing of 25 m.

The codes used in the exported model grids are described in section 2. In addition the qualifiers _t, _b and _th are used to indicate the unit top, base and thickness respectively.

Examples of uncertainty layers are also provided as images or ASC files, these are described further in section 8 of this report.

2 Modelled surfaces/volumes

2.1 SUPERFICIAL DEPOSITS

Code	Geological Unit
Water	Water
MGR-ARTDP	Made Ground (made and worked ground undifferentiated)
HEAD-XCZSV	Head
PEAT-P	Peat
LDE-XCZSP	Lacustrine deposit
LAWSG-XCZSVP	Law Sand and Gravel Member
KELV-XCZSP	Strathkelvin Sand and Silt Member
GOSA-XCZSV	Gourock Sand Member
KARN-XSV	Killarn Sand and Gravel Member
LIWD-XCZS	Linwood Clay Member
PAIS-XCZS	Paisley Clay Member
BRON-XSVZ	Bridgeton Sand Member
RSSA-XSV	Ross Sand Member
RSSA-XSZ	Ross Sand Member (silt, sand)
BHSE-XSV	Broomhouse Sand and Gravel Formation (sand and gravel)
BHSE-S	Broomhouse Sand and Gravel Formation (sand)
BILL1-XZCS	Bellshill Clay Member
WITI-DMTN	Wilderness Till Formation
SUPD-XZC	Clay and silt
CADR-XSV	Cadder Sand and Gravel Formation
BRL1-XCZ	Broomhill Clay Formation
BNTI-DMTN	Baillieston Till Formation

Table 1 Lithostratigraphic units included in the superficial deposits model of Central Glasgow

Table 1 indicates the lithostratigraphic units in the Central Glasgow Superficial Deposits Model. In Table 1, X indicates that each lithology is represented (e.g. XSV is a unit containing sand and gravel as opposed to SV which would be a gravelly sand), where S=sand, C=clay, Z=silt, V=gravel, DMTN= diamicton, ARTDP= artificial deposits, P= Peat. Some lenses are modelled in addition, as described below.

A description of each of the modelled units is given below in approximately ascending order (Central Glasgow and beyond).

2.2 BAILLIESTON TILL FORMATION (BNTI)

The Baillieston Till Formation rests directly on bedrock in the Clydebank and North Glasgow areas and is overlain by the Broomhill Clay Formation, Cadder Sand and Gravel Formation, or an un-named silt unit. It consists of stiff to very stiff, red-brown, silty, sandy clay with gravel, cobbles, boulders and some sand layers. The Baillieston Till Formation is considered to be a pre late-Devensian till.

2.3 BROOMHILL CLAY FORMATION (BRLL)

The Broomhill Clay Formation consists of laminated clayey-silt with sandy partings and rests directly on bedrock or is underlain by the Ballieston Till. It is overlain by the dense sand and gravel of the Cadder Sand and Gravel Formation.

2.4 CADDER SAND AND GRAVEL FORMATION (CADR)

The Cadder Sand and Gravel Formation is found mainly to the north of the River Clyde, with deposits up to 60 m thick associated with the deep bedrock depressions of the Kelvin buried valley-system. In some areas (e.g. SE Glasgow) the unit has been truncated by the emplacement of overlying tills and by modern river erosion, resulting in modification of the original depositional morphology of the unit.

The Cadder Sand and Gravel Formation consists of bedded and trough cross-bedded dense sand or silty sand, with gravel and some cobbles. There are occasional clay lenses. Cross bedding has been described in exposed outcrops in the Kelvin Valley. Significant deformation of the upper parts of the unit has occurred due to overriding of the Late Devensian ice sheet (Wilderness Till). The Cadder Sand and Gravel Formation is thought to have originated as outwash deposits, possibly fluvial or deltaic, formed in front of the advancing late Devensian ice sheet. The sands have yielded bones and teeth of woolly rhinoceros, from which Rolfe (1996) reported a radiocarbon age of 27.5 ¹⁴C ka BP.

2.4.1 Silt Unit (SUPD_XZC)

A small, un-named silt unit consisting of a firm to stiff, red-brown clay or silty clay with a little sand, overlies the Cadder Sand and Gravel Formation in places. It has silt laminations and may contain sand seams and traces of fine gravel. The unit contains a till lense and is overlain by the Wilderness Till Formation. It may represent deposition from pro-glacial lakes developing in association with Cadder Formation sand and gravel as the Late Devensian ice sheet advanced up the Firth of Clyde and into the Clyde Valley.

2.5 WILDERNESS TILL FORMATION (WITI)

The Late Devensian Wilderness Till Formation is the most extensive unit in the Clyde area and is named after temporary sections seen in the Wilderness Plantation area north of Bishopbriggs. It is characterised by a diamicton comprising isolated boulders, gravel and cobbles in a firm to stiff sandy, silty to clayey matrix (Browne and McMillan, 1989). Commonly it rests directly on bedrock, but is underlain in places by the Cadder Sand and Gravel, Broomhill Clay or Ballieston Till formations. Drumlins, which are large mounds formed during emplacement of the till below moving glacier ice, are a characteristic landform associated with the Wilderness Till. Drumlins constrain the hilly terrain of Central Glasgow and the surrounding areas where the Wilderness Till is exposed at the surface. Along the Clyde valley, buried drumlins cause considerable spatial variation in the thickness of overlying sediment deposits.

2.6 BROOMHOUSE SAND AND GRAVEL FORMATION (BHSE)

The Broomhouse Sand and Gravel is named after the Broomhouse area of eastern Glasgow where it generally overlies the Wilderness Till Formation. It comprises glaciofluvial ice-contact deposits, which produce features such as esker ridges, mounds, isolated flat-topped kames and kettleholes. Overall, the most abundant deposit is sand, except in esker ridges where gravel dominates. In some places it may also contain cobbles, clay and silt. Sands are planar and trough cross-bedded, ripple laminated and horizontally laminated, gravels are typically massive or crudely bedded. Deposits are up to 25 m thick and flow directions were towards the east (Browne and MacMillan, 1989). The noted occurrences of Broomhouse Sand and Gravel Formation are coincident with areas of dense borehole data. This implies that there may be more

extensive deposits in the area than have been modelled, but a lack of data means that further deposits remain undetected. The Broomhouse Sand and Gravel Formation includes three members: the Bellshill Clay Member, the Greenoakhill Sand and Gravel Member and the Ross Sand Member. The Bellshill Clay and Ross Sand members have been modelled.

2.6.1 Bellshill Clay Member (BILL1 and BILL2)

The Bellshill Clay Member, a member of the Broomhouse Sand and Gravel Formation, comprises silt-clay with wisps, laminae and bands of silt and sometimes sand, deposited, along with the coarser-grained Ross Sand Member (described below), in lakes forming at the margin of the Clyde glacier. The largest of these lakes, a pro-glacial lake called 'Lake Clydesdale', formed to the south-east of an ice margin marked by the 'Blantyferme terminal moraine' (south-east of Glasgow) as the glacier retreated towards the north-west (Browne and McMillan, 1989). The area covered by 'Lake Clydesdale', located south of Bothwell, thus contains the most extensive development of the Bellshill Clay Member. Away from the Clyde valley, the Bellshill Clay Member is found in hollows in the Wilderness Till Formation drumlins. Its occurrence can be best explained by a level for Lake Clydesdale at about 62 m O.D. although it is found as high as 80 m O.D. in other areas.

Near the margins of the former glacial lake, the Bellshill Clay Member inter-fingers with coarser silty-sand and gravelly-sand of the Ross Sand Member, deposited in deltaic systems by meltwater streams entering Lake Clydesdale. The Bellshill Clay Member is thus found below, adjacent to, and in places above, the Ross Sand Member (e.g. Fig 3 of Browne and McMillan, 1989), and the relationship of the units has been simplified for modelling purposes by splitting the Bellshill Clay Member into two units that lie above (BILL2) and below (BILL1) the Ross Sand Member. The Bellshill Clay Member is overlain by the Bridgeton Sand Member, Paisley Clay Member and later alluvial deposits.

The following rules were used to model the Bellshill Clay Member:

- 1) The Bellshill Clay member is distinguished from the overlying Paisley Clay Member by the presence of intervening sand units and more abundant sandy laminae within the silt-clay strata.
- 2) The Bellshill Clay Member includes some lithologies coded in the BGS borehole database as diamicton. These are described in logs as clay rich gravel layers and are hypothesised to be boulder clay. The decision to include diamicton in this group is supported by descriptions of the Bellshill Clay Member lithologies in Browne and McMillan (1989). These units are interpreted to be thin till bands deposited during minor oscillations in the position of the ice front during the period of existence of Lake Clydesdale.
- 3) Where it outcrops at the surface, the Bellshill Clay Member is equivalent to the Glaciolacustrine Deltaic Devensian (GLDDD), consisting of silty clay with some predominant silt and sand partings, depicted on BGS digital 1:10 000 and 1:50 000 scale superficial geology maps (DiGMapGB).

2.6.2 Ross Sand Member (RSSA)

The main lithologies of the Ross Sand Member, a member of the Broomhouse Sand and Gravel Formation, are sand or sand and silt, with clays at the base and thin local gravel layers. As noted above, the deposits are glacio-lacustrine in origin with deposits found in the south-east of Glasgow interpreted to have formed in deltaic systems at the margins of glacial 'Lake Clydesdale' (Browne and MacMillan, 1989). For modelling purposes, the lithologies have been separated into two units, a main unit of sand with minor gravel (RSSA-XSV), and an underlying finer-grained sand with silt (RSSA-XSZ).

The Ross Sand Member is found draping the Wilderness Till Formation and is overlain by deposits of the Bridgeton Sand Member and Paisley Clay Member. Where it outcrops at the

surface, the Ross Sand Member corresponds to glaciofluvial deltaic (and/or subaqueous fan) deposits (GFDD) on BGS DiGMapGB.

2.7 CLYDE CLAY FORMATION

The Clyde Clay Formation is part of the British Coastal Deposits Group and includes mainly Late Devensian deposits from marine isotope stages 2, 2a-b $\delta^{18}\text{O}$. The formation is modelled as component members including the Bridgeton Sand Member, the Paisley Clay Member, the Linwood Clay Member, the Killearn Sand and Gravel Member and the Innerleven Gravel Member.

2.7.1 Bridgeton Sand Member (BRON)

The Bridgeton Sand Member is characterised by fine to medium, massive dense sand or silty sand. Locally, fine to coarse gravel and boulders occur in a sandy matrix. There is some flat bedding but generally the deposits are massive. The unit is largely confined to the Clyde valley, where it overlies the Wilderness Till and Broomhouse Sand and Gravel formations. In the Clydebank area, The Bridgeton Sand Member is 20–30 m thick occupying a 2.5 km wide ‘strip’ confined by a bedrock depression following the line of the River Clyde. The deepest part of the depression, which contains the greatest thickness of the Bridgeton Sand Member, lies approximately 300–400 m to the south of the modern river.

Browne and MacMillan (1989) suggested that the sands were deposited as submarine outwash fans formed during catastrophic draining of pro-glacial Lake Clydesdale to the north-west along the line of the Clyde valley following breaching of the glacier dam.

2.7.2 Paisley Clay Member (PAIS)

The Paisley Clay Member comprises laminated clay and silt-clay deposited in a glaciomarine setting. In borehole records it is described as a grey and grey-brown, occasionally laminated, clayey silt and silty clay, often with a mottled appearance. The retreating glaciers are believed to have been to the north-west, in the sea lochs of the Southern Highlands. Relative sea level was high when deposition of the Paisley Clay Member commenced and some clays were deposited at elevations up to 40 m above OD (Browne and MacMillan, 1989). The Paisley Clay Member drapes the underlying topography, and appears to thin out over the tops of drumlins (Wilderness Till Formation) and bedrock ‘highs’. To the south-east of Glasgow, all clay units at the ground surface within lowland areas (max. 40–45m O.D.) are assumed to be deposits of the Paisley Clay Member. All DiGMapGB polygons coded as raised marine deposits-Devensian (RMDV) were interpreted as outcrops of the Paisley Clay Member.

2.7.3 Linwood Clay Member (LIWD)

The Linwood Clay member overlies the Paisley Clay Member and consists of up to 35m of clay or silt clay, with some sandy lenses and very sparse gravel. Occasional boulders encountered in some boreholes may be dropstones. The Linwood Clay Member is stratigraphically distinguished from the underlying Paisley Clay largely on the basis of an increase in the diversity of marine fossils contained within the strata (Browne and MacMillan, 1989). However, borehole logs rarely contain palaeontological information and thus the Paisley and Linwood Clay members are difficult to distinguish in many areas. Generally, the Linwood Clay Member is considered to be less laminated than the underlying Paisley Clay Member. In some areas, especially near outcrop, the Paisley Clay Member is orange to red-brown in colour, compared to the more grey-brown of the Linwood Clay. In an embayment of the palaeoshoreline in the vicinity of Paisley-Renfrew (much of sheet NS46NE), the Paisley Clay is significantly thinner than deposits found further east up the Clyde valley. By contrast, the Linwood Clay Member becomes more dominant in this area and is considerably thicker than the underlying Paisley Clay Member .

2.7.4 Killearn Sand and Gravel Member (KARN)

The Killearn Sand and Gravel Member is a patchy deposit comprising of varying proportions of sand and some clay layers. The deposits are commonly found in inter-drumlin areas along the Clyde Valley, and the distribution suggests that it formed as a result of marginal marine processes associated with the Late Devensian marine incursion responsible for the deposition of the Paisley and Linwood clays. The maximum level of the marine incursion thought to be responsible for the Killearn Member is approximately 34–36 m above OD (Browne and McMillan 1989, Hall et al. 1998 and Rose 1975).

The Killearn Member is thought to have been deposited in a range of environments, including beaches, and fluvial and deltaic systems. Rose (1975) describes temporary sections north of Erskine bridge where Killearn Member sands and gravels overlie Broomhouse Sand and Gravel Formation. Here the Killearn is well bedded, with beds displaying shallow dip to the south-west. It is suggested that the sedimentary structures are consistent with shoreline processes, hence that the unit in this area is a beach gravel. Rose (1975) also describes ice wedge casts suggesting that a periglacial environment persisted in the period following the deposition of the Killearn Member. It is suggested that these features may have formed during the cold period associated with the cooling event of the Younger Dryas. During this time, glaciers readvanced in the Loch Lomond area, but the Clyde remained ice free.

2.7.5 Inverleven Gravel Member (INVN)

The Inverleven Formation is a 0.5–2.5 m thick unit found in the lower part of the Clyde valley. It consists of boulders, cobbles and gravel in a sandy-clay matrix, and some shells and traces of barnacles on clasts have been found (Browne and McMillan, 1989). It is thought to be a lag deposit representing a Late Devensian erosive, or non-depositional phase in the Clyde Estuary. The Inverleven Gravel Member overlies eroded Linwood and/or Paisley Clay members and Wilderness Till Formation (Browne and McMillan, 1989).

2.8 CLYDEBANK CLAY FORMATION

2.8.1 Longhaugh Sand and Gravel Member (LUGH)

The Longhaugh Formation is known only from site investigation boreholes and there is no standard section for the unit (Browne and McMillan 1989). It consists of silty-sand, with a little gravel, and some boulders towards the base. Some boreholes record shell debris, and this rarely noted characteristic is the only means of distinguishing it from the sand units of the underlying Bridgeton Sand Member and the overlying Gourock Sand Member.

The Longhaugh Formation appears to fill a pre-existing erosive channel aligned along the modern Clyde river, in places lined with coarse boulders and cobbles of the Inverleven Gravel Member. The channel cuts into older deposits, reaching a maximum depth of ~ -20 m OD. This stratigraphy suggests that the channel was cut, probably by tidal processes concentrated through the narrow Erskine gap during a marine incursion in the Late Devensian. It may have developed during a relative sea level fall (c.f Browne and McMillan 1989 p15-16), either associated with the Loch Lomond Re-advance, or with the transition from marine to estuarine conditions that occurred in Late Devensian to early Flandrian times. This interpretation is consistent with the lack of occurrence of the Paisley and Linwood Clay Members in the lower parts of the Clyde Valley. Strong tidal currents may have prevented the deposition of fine clay and silt sediments in this area throughout the late Devensian, with the channel incision occurring at a low stand in sea level.

2.8.2 Erskine Clay Member (ERSK)

The Erskine Clay Member is the fine grained equivalent of the Flandrian age Gourock Sand Member, and is found in the valleys of the River Clyde and White and Black Cart Waters. It is

distinguished from the underlying Linwood Clay Member because it contains significant organic matter. It is mainly found in association with the Gourock Sand Member, which overlies the Erskine Clay in a number of areas. The Erskine Clay Member overlies the Wilderness Till, Broomhouse Sand and Gravel Formation, Bridgeton Sand Member and the Longhaugh Sand and Gravel Member in the Clyde Valley, and the Linwood Clay Member in the valleys of the Black and White Cart. Small patches of peat may be developed on top of the Erskine Clay Member.

2.8.3 Gourock Sand Member (GOSA)

The Gourock Sand Member forms extensive deposits in the Clyde valley and as deposits along the Black and White Cart Waters. It consists of 0.5– ~5 m of fine to coarse sand with some gravel, silt and clay and organic detritus. It passes laterally in some areas into the Erskine Clay Member. The deposits are likely to have formed in estuarine environments, with a fluvial dominance in the east, becoming progressively more marine westwards with shallow channels linked by tidal flats (Browne and McMillan, 1989).

2.9 CLYDE VALLEY FORMATION

2.9.1 Strathkelvin Clay and Silt Member (KELV)

The Strathkelvin Clay and Silt Member is found largely within the Kelvin valley and as deposits along water courses to the north-east of Glasgow. The unit consists of a silt-clay containing abundant layers of silt and some sandy layers and lenses interbedded with peat. In general the clays are dark brown in colour and bands relatively rich in organic detritus also occur. The deposits are either of floodplain or lacustrine origin.

2.9.2 Law Sand and Gravel Member (LAWSG)

The Law Sand and Gravel Member comprises fine to coarse sand with some silt, fine gravel and organic matter deposited in river channels and associated floodplains. The Law Sand and Gravel Member includes recent (currently accumulating) river deposits in the upper Clyde valley and deposits associated with the rivers Kelvin and the Black and White Cart Waters. Small alluvial fans and deposits associated with minor streams are also included in this unit.

2.10 LACUSTRINE DEPOSITS (LDE)

Recent lacustrine deposits, consisting largely of soft silt and clay with organic bands and lenses of peat, are found in and around lakes or former lakes throughout the Clyde area. These lakes have formed in inter-drumlin hollows, and some have been completely infilled by accumulation of sediment and organic matter.

2.11 PEAT

Small deposits of peat, consisting of dark, organic rich, humic material formed by the accumulation of partially decomposed vegetation, are found in places in the Clyde area. Peat accumulations are generally associated with lacustrine and alluvial sediments, where they may form surface deposits, or form bands or lenses within the strata (discussed below).

2.12 MADE GROUND (MGR)

Made ground in the 3D model represents a combination of made and worked ground including filled and partially back-filled pits and quarries - hence it comprises all anthropogenic deposits. Areas of worked ground were primarily identified using DiGMapGB 1:10 000 polygons. These were subsequently altered to encompass areas where boreholes reported additional areas of artificial ground. Alterations were made using the Ordnance Survey maps to identify the extent of industrial areas, housing developments and other information. It is likely that more extensive

but thinner deposits of made ground occur within the model that are not currently represented. This is particularly the case on sheet NS56NE (Kelvinside-Springburn and surrounds) Made ground has not been subdivided.

An accurate 5 m DTM and good quality borehole descriptions would be essential to the more detailed modelling of the made ground in the Clyde Valley area. Subdivision of the made ground will be considered in future models.

2.13 HEAD

Small areas of gravely clay overlying the Bridgeton Sand Formation, Law Sand and Gravel Member or the Paisley Clay Member, usually found at the base of a drumlin with significant relief, are interpreted as Head deposits formed by the remobilisation of the surface deposits on steep hillslopes.

2.14 LENSES

Many of the units described above contain considerable variation in lithologies within their strata. Where thin deposits of a contrasting lithology to the main unit are recorded within borehole descriptions, these have been depicted by the modelling of a lense contained within the 'parent' unit. In general, only lithological sub-units with thicknesses greater than approximately 2 m have been identified as separate lenses.

The lense types used in the Central Glasgow Superficial Deposits Model are: clay_lense, silt_lense, sand_lense, gravel_lense, peat_lense, till_lense, CADR_lense, CADR2_lense, Witi_lense (Boulder clay without the clay fraction (washed out)).

CLAY AND SILT		SAND AND GRAVEL		DIAMICTON	PEAT	LITHOLOGY
Marine	Lac/Fluv	Marine	Lac/Fluv	Glacial	Organic	ORIGIN
		Gourock	Law		Clippens	FLANDRIAN
		Killearn				
Paisley						DEVENSIAN
		Bridgeton				
	Bellshill					
			Ross			
			Broomhouse			
				Wilderness		
			Cadder			

Table 2. Summary of origin of superficial deposits stratigraphic units in the Central Glasgow area: lithology, origin and age (Lac – lacustrine, Fluv – fluvial).

2.15 BEDROCK

The bedrock geology beneath the Central Glasgow Bedrock Model area comprises Upper Carboniferous Coal Measures and Clackmannan Group strata (Table 2, Figure 3). The majority of the strata represent fluvio-deltaic to shallow marine facies consisting of argillaceous rock, sandstone, coal and limestone. The lithostratigraphy is primarily identified from interpretation of borehole records using the established BGS Carboniferous lithostratigraphic framework, lithostratigraphic and biostratigraphic markers (Browne et al., 1999; Hall et al., 1998).

Rockhead unconformity	Combined base of all the superficial deposits
Base Upper Coal Measures (UCMS)	
Glasgow Upper coal (GU)	Worked coal in Scottish Middle Coal Measures Formation (MCMS)
Glasgow Ell Coal (GE)	Worked coal in Scottish Middle Coal Measures Formation (MCMS)
Kiltongue Coal (KILC)	Worked coal in Scottish Lower Coal Measures Formation (LCMS)
Knightswood Gas Coal (KDG)	Worked coal in Scottish Lower Coal Measures Formation (LCMS)
Base Upper Limestone Formation (ULGS) = Index Limestone (ILS)	

Table 2 Summary of modelled bedrock surfaces

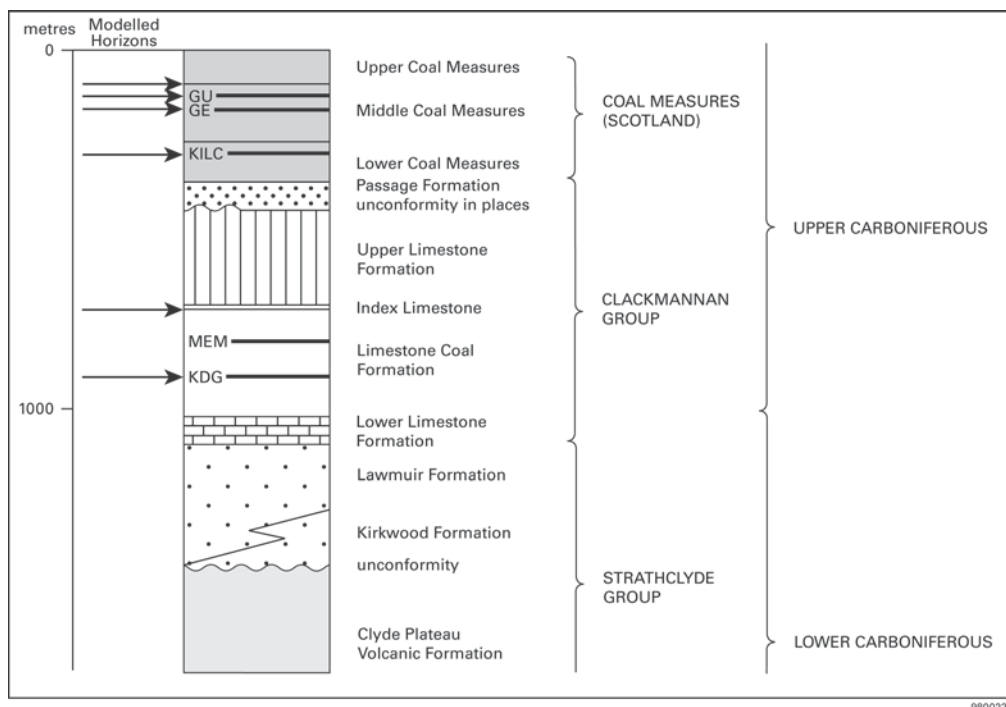


Figure 3 Summary stratigraphy of the modelled Carboniferous strata in the Glasgow area. Coal name codes: Glasgow Upper coal (GU), Glasgow Ell Coal (GE), Kiltongue Coal (KILC), Knightswood Gas Coal (KDG)

GU and GE are the uppermost and most extensively worked coals in the Scottish Middle Coal Measures Formation (MCMS). KILC is quite extensively worked in the Scottish Lower Coal Measures Formation (LCMS). KDG is extensively worked within the Limestone Coal Formation, cropping out on the western side of the area.

3 Modelled faults

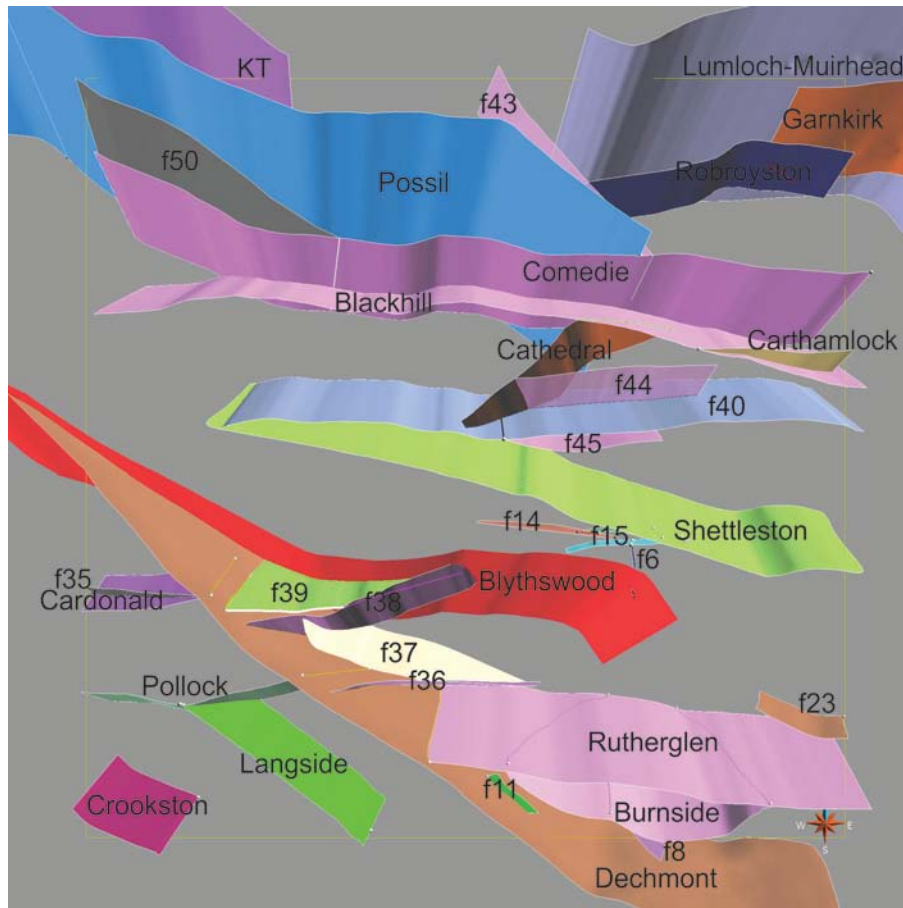


Figure 4 Overview of faults modelled in the bedrock, looking from above.

Dipping faults were included in the bedrock model (Figure 4). Some faults have long-established names, while others are designated informally (e.g. f23). With a few exceptions relating to the inclusion of smaller structures on NS66SW, faults were included in the model if over 30 m throw or 2 km in length.

Fault dips were calculated where data were available using the XYZ positions recorded in subsurface mine abandonment plans and linked to the surface outcrop position. The data were mainly available on NS66SW, showing that the majority of faults have dips c. 60°. One fault, the Burnside Fault (Figure 4), has a c. 45° dip. Occasionally, fault information was encountered in boreholes. However, the amount of fault subsurface position information from mining and boreholes is generally very limited (and not enough to make a fault defining pointset). Thus, faults were created by projection from their mapped outcrop position at 60° (apart from Burnside at 45°).

The availability of mining information on NS66SW meant that faults could be analysed for their continuation at depth. Smaller faults were observed to terminate within 100 m of rockhead (e.g. F6, Figure 4) whereas larger structures (e.g. Rutherglen) extend to the depth of the whole of the model. On NS66NW and NS56SE, mining information was more limited and faults incorporated in the model were generally larger in terms of length (and therefore probably depth also), such that faults on these sheets were extended to a standard depth of 1 km.

4 Model datasets and workflow

General caveats regarding BGS datasets and interpretations can be described:

- Geological observations and interpretations are made according to the prevailing understanding of the subject at the time. The quality of such observations and interpretations may be affected by the availability of new data, by subsequent advances in knowledge, improved methods of interpretation, improved databases and modelling software, and better access to sampling locations.
- Raw data may have been transcribed from analogue to digital format, or may have been acquired by means of automated measuring techniques. Although such processes are subjected to quality control to ensure reliability where possible, some raw data may have been processed without human intervention and may in consequence contain undetected errors.

4.1 SUPERFICIAL DEPOSITS

The Central Glasgow Superficial Deposits Model was constructed in GSI3D using a NEXTMap[®] Digital Elevation Model (©Intermap Technologies, at 50 m resolution), the BGS digital borehole database, BGS 1:10,000 scale digital maps (DiGMapGB, 2009), field slip scans, historic maps and scanned geological cross-sections. Other literature such as BGS regional geological guides and scientific papers influenced the correlation of geological units.

Borehole data were entered to the BGS corporate database BGS Borehole Geology. The spread of borehole data across the area was variable, from extremely closely spaced at site investigation locations to more widely spaced and isolated boreholes.

A standard GSI3D workflow for superficial geological models was followed (Kessler et al., 2008 <http://nora.nerc.ac.uk/3737/1/OR08001.pdf>) for the original component block models. The method principally involves construction of cross-sections between the best quality borehole data followed by envelope construction around the limits of the geological units. GSI3D model calculation then uses envelopes in combination with nodes on the geological surfaces along cross-sections to build geological surfaces by triangulation.

The original component block models (Figure 1) were modelled over a number of years, by a number of geologists and during a period of software development. As such, a number of steps were taken before release of these v2 (or v3 Central Glasgow) models to attempt to resolve inconsistencies between block models. Further details will be documented in a methodology report currently in preparation. The key steps included:

1. Inclusion/exclusion of water bodies to a common standard (see section 5.1.1 below)
2. Creation of a set of merged borehole data for each component block model
3. Creation of a set of master envelopes 'extents' covering the modelled area
4. Inclusion of higher resolution corridor model sections and envelope edits on SW Central Glasgow and Paisley (Monaghan & Whitbread, 2012)
5. Revision to Lexicon and lithology codes describing the modelled units to current best practice and update GVS, legend and GSIPR files
6. Using extended DTM's to cover the edge matching zone for each model

7. Edge matching the block models such that the geological interpretation was consistent across block boundaries. The main issues to be resolved were
 - a. Marine-influenced Gourock Sand Member (GOSA) has been modelled flanking the Clyde in Clydebank and Central Glasgow, but the equivalent unit is modelled as fluvial Law Sand and Gravel Member (LAW) to the south-east of Glasgow, upstream of the tidal limit. This was resolved by taking the GOSA as far as the Uddingston-Bothwell area and then using the LAW upstream of this.
 - b. The Paisley Clay Member was not modelled under Linwood Clay Member in the north-west quarter sheet of Paisley Model meaning that the Paisley Clay in the Clydebank Model ended abruptly at the model edge. Conversely the Linwood Clay was not modelled in the north-east of the Paisley Model. This was resolved by undertaking a revised interpretation of both units on both block models, such that they are now consistent.
 - c. Edge matching issues in envelopes (XY) or sections (Z) viewed as linear jumps. These were resolved by revised interpretations, guided by borehole data in the problem areas.

The following inconsistencies were not attempted to be resolved at this stage:

- Insertion of additional cross-sections to achieve a more uniform cross-section coverage
- Revision of made ground on NS56NE where some boreholes prove made ground that is not modelled.

4.2 BEDROCK

The Central Glasgow Bedrock Model was constructed from all borehole data entered to the BGS corporate database BGS_Borehole_Geology and reaching the modelled horizons, and mine plan, map outcrop and interpreted data.

The spread of borehole data across the area was very variable, from closely-spaced site investigations metres apart, to in extreme cases, boreholes more than a kilometre apart. Data points were concentrated around the outcrop of worked coals and were sparse on stratigraphic surfaces in deeper parts of the basin.

Most borehole data points have a reasonably good level of certainty. Boreholes with very poor quality records or very poorly known sites were not coded into the database. However, there can be uncertainty: in geologically coding short isolated site investigation boreholes; in a drillers record of a borehole (i.e. if not examined by a geologist); or sometimes in the siting of the borehole. However, these should result in errors in location being no greater than about 5–10 m in Z, and perhaps 20–50 m in terms of XY.

Mining data were compiled from all available mine abandonment plans. These consist of spot heights surveyed on the base of worked coal seams underground, and rarely of structure contour elevation data. The distribution of mining data points is variable. These data points have a high confidence level as they were systematically surveyed. Estimates of error on mining data points range from 0–5 m in Z and 0–25 m in XY

The bedrock map represents the outcrop (or subcrop) of stratigraphic horizons at rockhead i.e. very commonly buried beneath superficial deposits. The map data for NS66SW had been revised prior to this study and an updated map published (BGS, 2008). The map data for NS56SE and NS66NW were revised prior to the study but those revisions have not been published. That is the model contains more up-to-date map linework than does the current edition of the 1:10,000-scale

maps (BGS, 1995, 1996). The errors in mapped outcrop line work are extremely variable – from 0–10 m in XYZ where seen at outcrop, to tens of metres where an interpretive outcrop was created from little constraining data.

Expert geological interpretation was added during modelling as two cross-sections, and editing of fault-surface contacts and removal of overlaps during GOCAD® modelling.

A standard BGS GOCAD® modelling workflow using version 2.9.2 P2 structural workflow was employed. GOCAD® (Paradigm, www.pdgm.com) calculates a triangulated mesh based on XYZ data points and then the geologist modeller undertakes various processes to aid geological interpretation in data poor areas.

5 Model assumptions, geological rules used

5.1 SUPERFICIAL DEPOSITS

Where no boreholes were present, basal units were modelled to the BGS rockhead model (see <http://www.bgs.ac.uk/products/onshore/superficialThickness.html>). Where the depth of rockhead in boreholes differed from the rockhead model, the borehole depth was preferred.

5.1.1 Water modelling

A consistent approach to the modelling of water was attempted across all the modelled blocks shown on Figure 1. Firstly, a water envelope (extent) was extracted from the OS 1:10,000 scale Open data digital map. Judgement was then used to extract a subset of the larger water bodies from that map data such that the River Clyde, the lower parts of the White Cart and Black Cart waters and lakes/reservoirs exceeding approximately 150 m in width/length are included in the 3D model.

During modelling, the water bodies were generally fitted to the water envelope and not to the DTM as it is assumed that users will wish to attach the model to an OS map base rather than a derived/simplified NextMap DTM. The result is that in a few places where the DTM and OS map data do not match, water goes uphill. In places (e.g. Balgray reservoir) other artefacts exist within the modelled water (e.g. where there is bridge over the reservoir the DTM/water volume is raised up as the bridge is not modelled as such). The OS Open data envelope (extent) for water which extended to high water mark *was* edited within the intertidal zone in the Clydebank and western Central Glasgow models such that it fitted the NEXTmap® DTM in which ocean surface is levelled to OSGB1936 geodetic datum (i.e. zero metres relative to OD, or low water mark).

The depth of water within the tidal limit of the River Clyde was based approximately on a UK Hydrographic Office Admiralty Chart (International Chart Series c.2000). Where no information is available a sub-rectangular river channel cross-section has been assumed.

The depth of reservoirs was unknown and so they were modelled to water depth of about 10 m.

5.2 BEDROCK

It is assumed that all data types (borehole, mine plan, map) are included with equal weight in the modelled surface calculation, though in reality some data may be more certain than others. Data points were excluded if clearly erroneous, based on the geologist's judgement.

6 Model limitations

6.1 SUPERFICIAL DEPOSITS

- In the absence of borehole information, the model is constrained by map data and the DTM. Modelling highlights areas where the geological map may need revision, but this was not undertaken as part of the modelling exercise. The disparity is minor and commonly relates to anthropogenic activity post-mapping (e.g. made ground).
- Best endeavours (quality checking procedures) were employed to minimise data entry errors but given the diversity and volume of data used, it is anticipated that occasional erroneous entries will still be present (e.g. borehole location).
- The NEXTMap[®] Digital Elevation Model was subsampled from a 5 m resolution to 50 m resolution which means the surface distribution and geometry of a geological unit does not reflect the highest resolution possible. This resolution was chosen based on the size of the area, the resolution of the original modelling and the software capability. Some minor mismatches between geomorphological features and modelled units (including water) and the DTM occur due to the coarse resolution of the DTM.
- The NEXTMap[®] Digital Elevation Model may contain artefacts such as trees or artificial structures such as pylons. The majority of these have been stripped out before modelling. If any of these artefacts were found during the modelling then the effects of these were minimised in the model as much as possible.
- The start heights of boreholes used might differ significantly from the NEXTMap[®] Digital Elevation Model. When modelling, these differences were taken account of by assessing the year the borehole was drilled and assessing the location of the borehole against other data such as historical maps. Therefore the modeller used their own judgment in some areas if the stratigraphy in the borehole did not match the modern day topography and changes in the subsurface (quarrying, landfill etc).
- The thin nature of made ground, and the thin draped form of some areas of Paisley Clay Member deposits, means that these units are poorly shown in visualisations of the 3D model (e.g. in the Lithoframe Viewer 3D window). A substantial number of additional cross-sections ('helper sections') are needed to improve the calculation of thin deposits.
- A known limitation is that for some thin units close to DTM surface and over topographically variable ground (water, MGR, LDE, LAW, KELV, KARN, BILL1, WITI), the superficial deposits ASCII and ESRI grids contain small patches of no modelled surface within the unit envelope (outline, supplied as an ESRI shapefile) where it should exist. The limitation is greatest in the north-east and south-west of the model, where topography is greatest, and the largest area affected is a maximum of 300 by 400 m in made ground in the northeast of the model. This artefact of the modelling software/procedure can be rectified with additional interpreted cross-sections and will be addressed in future work.
- The model does not reflect the full complexity of the superficial deposits geology. In reality, surfaces have been subjected to more glacetectonic deformation than is represented in the model. It is also known that made and artificial ground is more widespread than is shown by the model, particularly on NS56NE as proven by boreholes.
- Smaller rivers, streams and water bodies have not been included in the models (see section 5.1.1)

- The model is attributed with geotechnical and hydrogeological properties. These are simply bulk attributions based on point data in boreholes from Central Glasgow and are provided for general guidance only.
- The cross-section density and therefore model certainty is variable across the model, and is based on complexity and type of geology, borehole density etc
- A subset of the most reliable borehole data has been included to constrain the cross-sections within the model. However there is also a large subset of borehole data that has not been included within the model.
- Whilst effort has been put into making models consistent and edge matched there remain minor inconsistencies between models produced and calculated separately along 5 km grid square boundaries (particularly NS56SE/NS66SW in the vicinity of the River Clyde) and small variations in interpretive style from different modelling geologists. These inconsistencies should not be apparent unless the models are being used at high resolution.

6.2 BEDROCK

General

- The model does not reflect the full complexity of the geology. In reality, geological surfaces will be cut with igneous intrusions/vents and cut by more faults than are modelled.
- The GOCAD® algorithm creates a triangular mesh to try and best fit all data points. For the Central Glasgow model a mesh size of around 120 m was optimal, as there are large areas without any data points. A smaller mesh size was tested but the resultant surfaces did not appear geologically realistic. In areas where there are abundant closely-spaced data (e.g. site investigations, or at a complex part of the outcrop line) a 100–200 m mesh size cannot represent the detail of the data density and complexity. In summary, the meshes are a representation of the geology, using all the data but not fitting all of it exactly. The maximum deviation between a surface and a data point is about 20 m. In the majority of cases the difference between any known data point and the modelled surface is less than 5 m.
- All data points have been checked on data entry and for consistency in the model but there will be some errors that remain – for example boreholes whose site has been incorrectly located in the database, or whose recorded start height is wrong.
- Note that to load to ArcGIS®, the modelled TIN (triangular mesh) files have been converted to grids (rectangular mesh). A grid spacing of 25 m was used so that some detail of fault gaps is preserved. This may give a false impression of the model resolution as the original TIN mesh spacing was 100–200 m. The ArcGIS® grids do not give a clean/fitted together 3D model (e.g. at fault-surface contacts) because of the TIN to grid conversion process.

Faults

- Faults with offsets less than 30 m have not been included in the model because the data were insufficient to constrain them. There will be some areas where adjacent data points have significantly different Z values because they are offset by a fault that is not modelled.
- Fault geometries which terminate against each other at low angles result in numerous thin ‘slivers’. These are difficult to model satisfactorily and are often poorly constrained by

data (e.g. at junction of Blythswood, f38 and f39). There is a small sliver of the Kiltongue Coal surface that penetrates through f38, and should not. Thus thin slivers should be treated as parts of the model with high uncertainty.

- f40 reverses throw along its length, as on the 10,000 scale map, but the fault dip has consistently been modelled at 60° to the north along the whole structure.
- The scale of faults modelled is inconsistent. Some smaller faults have been included (f14,15,6) where they constrain the outcrop of the modelled coals on NS66SW. In other areas of the model, only much larger fault structures have been included. Known, smaller faults have been excluded.

Surfaces

- A more recent and higher resolution model along a linear corridor covering the southwestern part of the Central Glasgow model (Monaghan, 2012b) has highlighted that the KILC, bULGS, KDG modelled surfaces are too high in the Central Glasgow Bedrock Model where they are interpreted in the hangingwall (eastern side) of the Dechmont fault.
- Isopach maps and cross-sections through the stacked surfaces and faults highlight inconsistencies caused by lack of data, particularly on the more deeply buried parts of KDG and Base ULGS/ILS, or by patchy data coverage on one particular surface (e.g. circular inconsistencies caused by data on a particular coal seam from a particular, localised set of boreholes between GU and GE). For example, the variation in general dip between faults 37 and 38 on KILC-ILS-KDG surfaces is inconsistent. Generally the KDG and ILS modelled surfaces are much more uncertain away from their outcrop than the other modelled surfaces, due to lack of data.
- Very small areas of bedrock surfaces lie up to 2 metres higher than the rockhead model causing a local crossover. This happens in areas where there are local low points in the higher resolution mesh of the rockhead surface, and there are no TIN points within that area from the lower resolution bedrock mesh (crossovers at TIN points should have been removed).

7 Model images

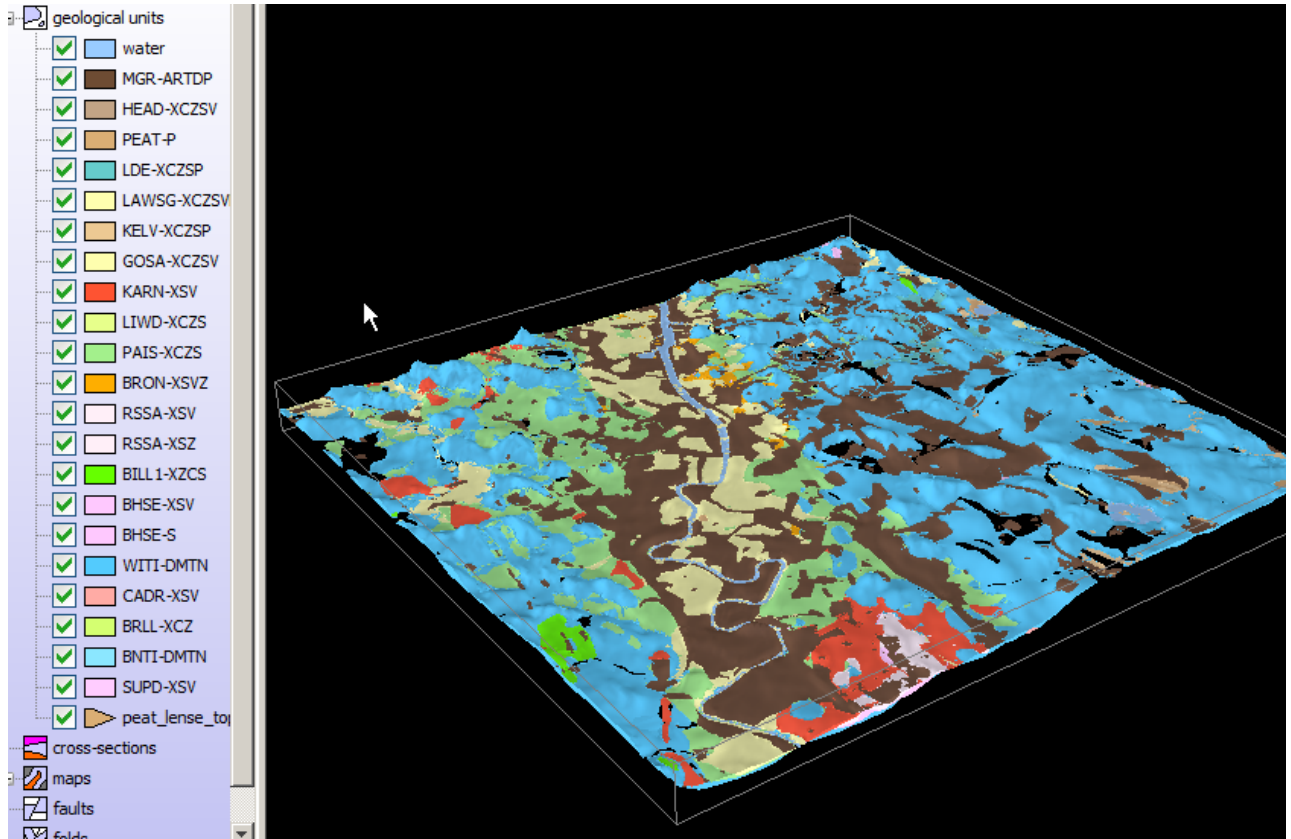


Figure 5 Overview of superficial deposits model of Central Glasgow Superficial Deposits Model, looking NW, ten times vertical exaggeration.

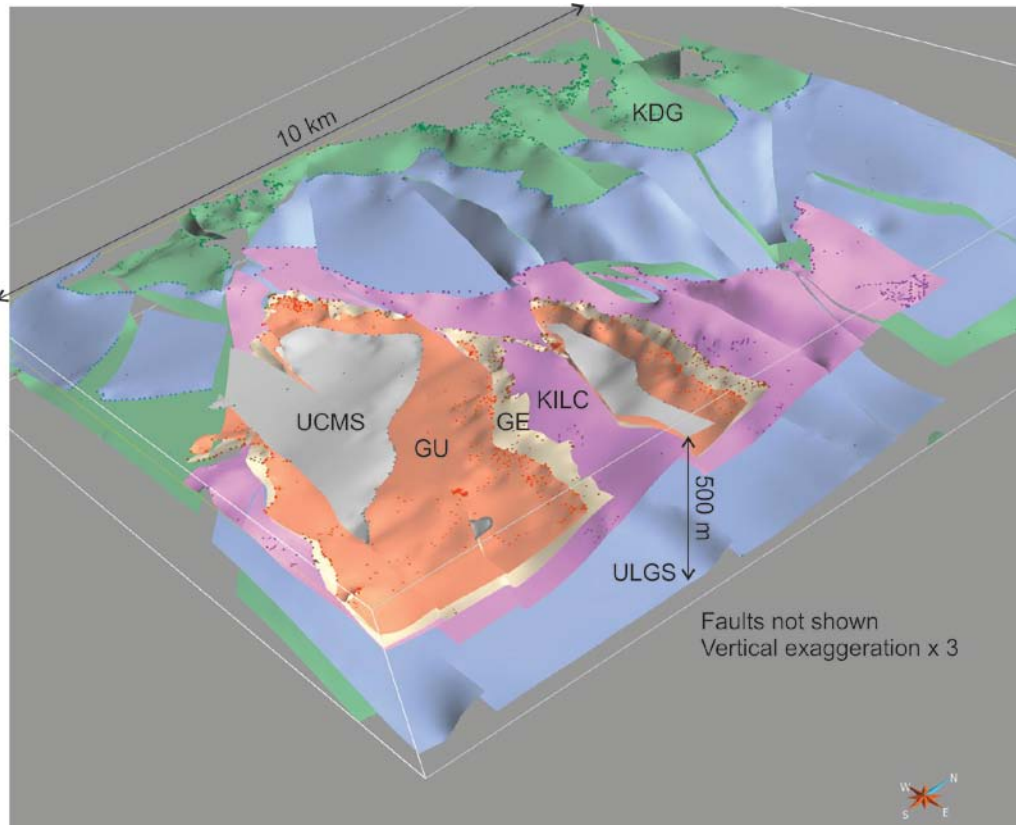


Figure 6 Overview of bedrock surfaces modelled in the Central Glasgow Bedrock Model with data points from boreholes, mine plans and mapped outcrop shown.

8 Model uncertainty

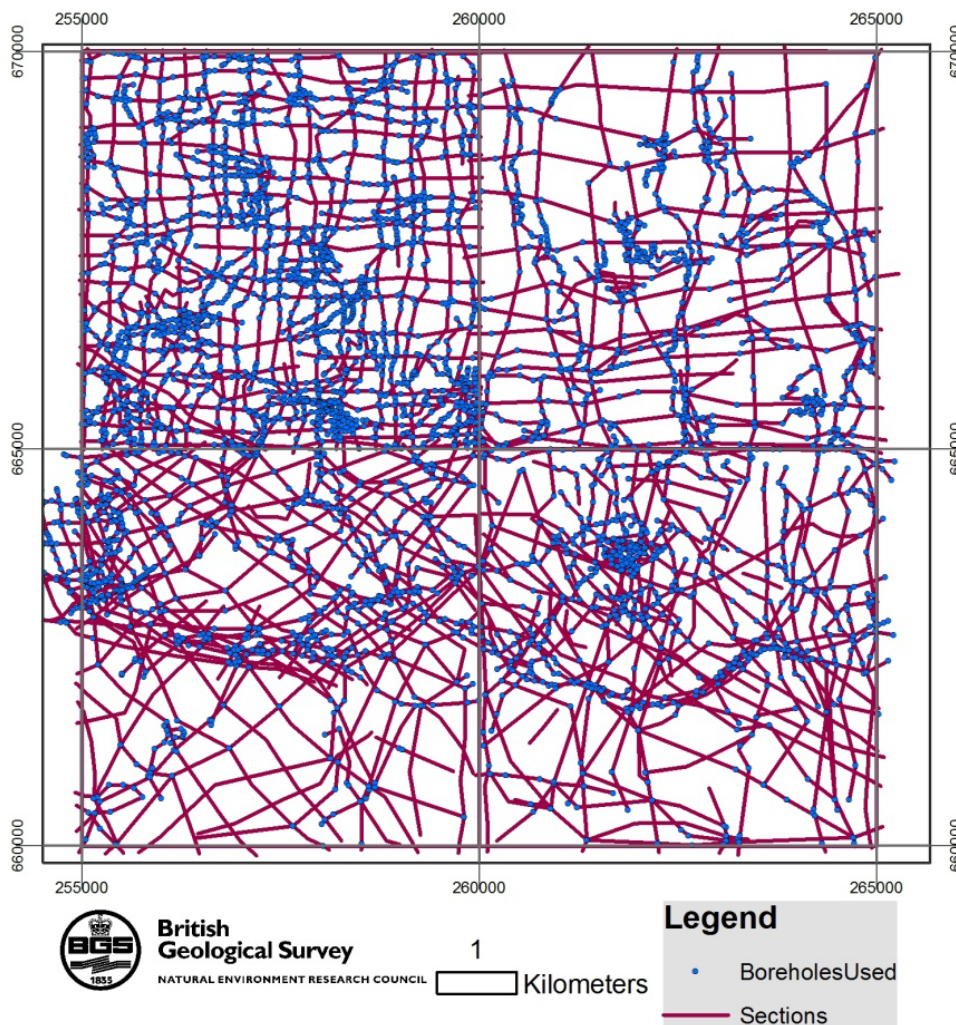


Figure 7 Cross-section array of 383 lines used to constrain the superficial deposits modelled in Central Glasgow, with the 4435 boreholes used shown in blue.

The cross-sections and borehole data points constraining the Central Glasgow Superficial Deposits Model shown in Figure 7 give an indication of the most certain areas of the model containing the most sections and borehole data, and the least certain. A more quantitative approach (described below) is also used in BGS to indicate model uncertainty. Later in 2013 uncertainty layers will be calculated for the whole area of the Clyde 3D geological models and released to the ASK Network.

Currently, ASK Network users are supplied with uncertainty layers from an earlier (2009) version of the 3D geological model for evaluation purposes. The uncertainty layers cover 3 quarter sheets of the Central Glasgow models (NS66SW, NS66NW and NS56SE). Comments on the utility of these uncertainty layers would be welcomed.

The uncertainty layers are supplied in two formats (1) .jpg with .jgw raster image files and (2) ASC files, which can be converted to ArcGIS® grids, for example, using ArcToolBox ('conversion tools' – 'ASCII to raster'). When viewing the converted ASC files, the user needs to colour up the grids based on the 'count' field, which is the relative value of uncertainty for that layer of the model. The largest number represents relative high uncertainty and the smallest number represents relative low uncertainty.

Uncertainty layers were calculated from a combination of data density and geological complexity of the modelled surface. The BGS confidence calculator v1_2 - customised BGS software developed in Matlab - was used. This means the model will be most uncertain where there is little data and where the geological surface dip changes rapidly.

For the superficial deposits model, relative low uncertainty might be considered to have errors of the order of ± 10 m in XYZ, for example, those areas coloured green on Figure 8. In the most uncertain areas where the geology is complex and poorly constrained by borehole data, uncertainty may be approximately ± 70 m in XYZ for example, those areas coloured red on Figure 8.

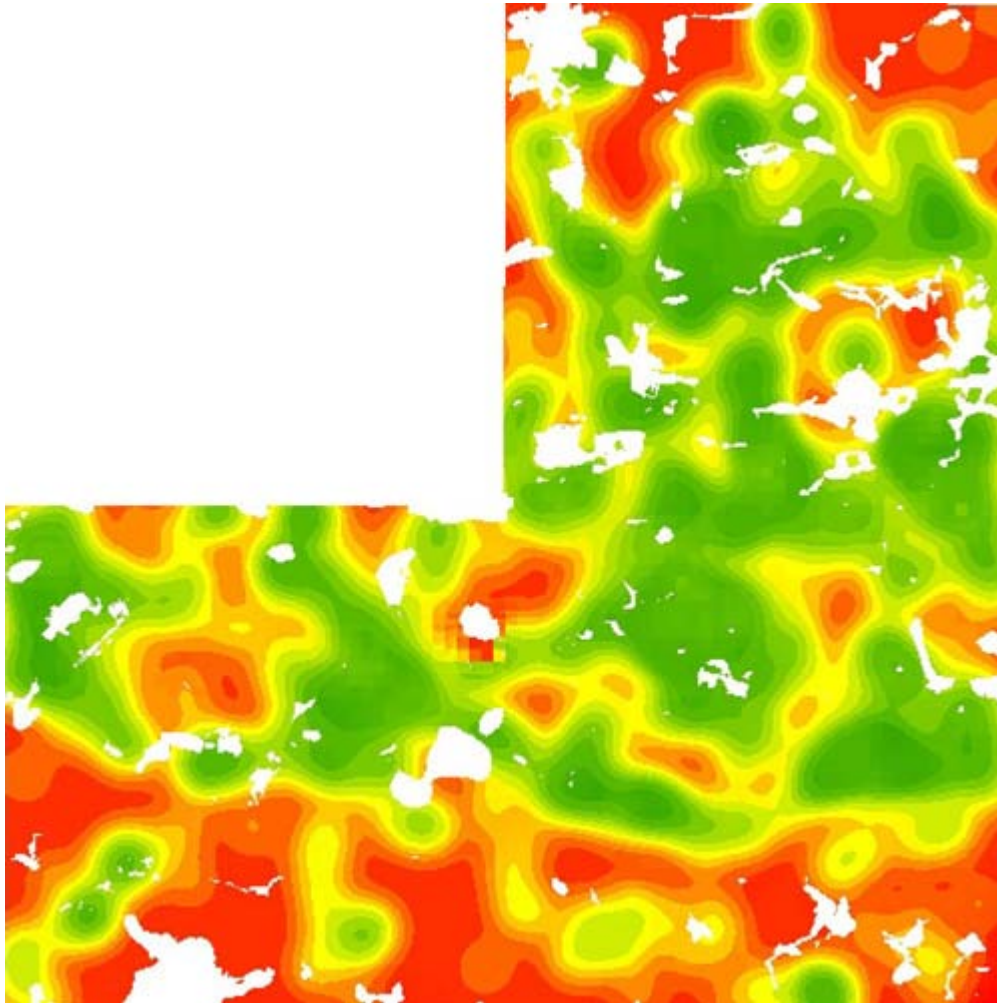


Figure 8 Uncertainty layer for the Wilderness Till coloured up such that the relatively most uncertain areas are red, relatively least uncertain areas are green (covers NS66SW, NS66NW and NS56SE).

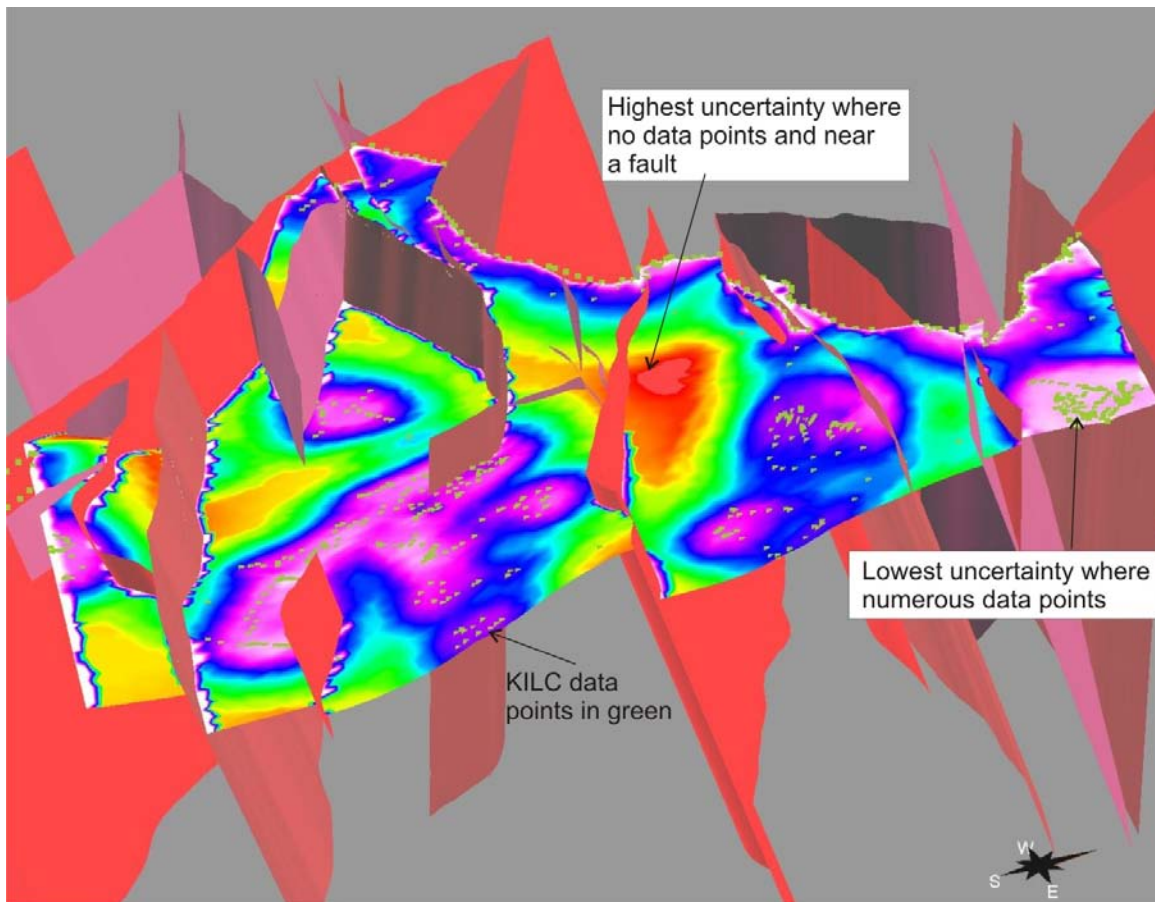


Figure 9 Example of combined uncertainty attribute for KILC draped on the geological surface in GOCAD®

For the bedrock surfaces, lowest uncertainty (highest confidence) areas are those that are well constrained by geological data and where the geology is relatively simple. In these areas, the error on the model might be considered to be of the order of ± 10 m in XYZ; for example, those areas of the KILC uncertainty surface on Figure 9 that are purple or pink.

Highest uncertainty (lowest confidence) areas are areas that are not constrained by any geological data and where the geology is complex i.e. faulted or folded. In these areas, the error on the model might be considered to be of the order of ± 70 m in XYZ; for example those areas of the KILC uncertainty surface on Figure 9 that are orange or red.

9 Conclusions and recommendations for future work

Comments and feedback on the Central Glasgow models are encouraged from ASK network members.

Improvements could be made to the Central Glasgow Superficial Deposits Model by inclusion of more borehole data, increased section density, and improved artificial ground modelling.

Improvements could be made to the bedrock model by addition of higher resolution and more consistent geological interpretation to the deeper surfaces.

Future work aims to address comments made from ASK network users and hopefully to include data provided by them.

These models supercede the published BGS 1: 10,000 scale maps of the Glasgow area. Updates in line with advances made in geological understanding of the area during 3D modelling are yet to be made to the maps.

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British Geological Survey holds most of the references listed below, and copies may be obtained via the library service subject to copyright legislation (contact libuser@bgs.ac.uk for details). The library catalogue is available at: <http://geolib.bgs.ac.uk>.

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