

# Optimal mapping of terrestrial gamma dose rates using geological parent material and aerogeophysical survey data

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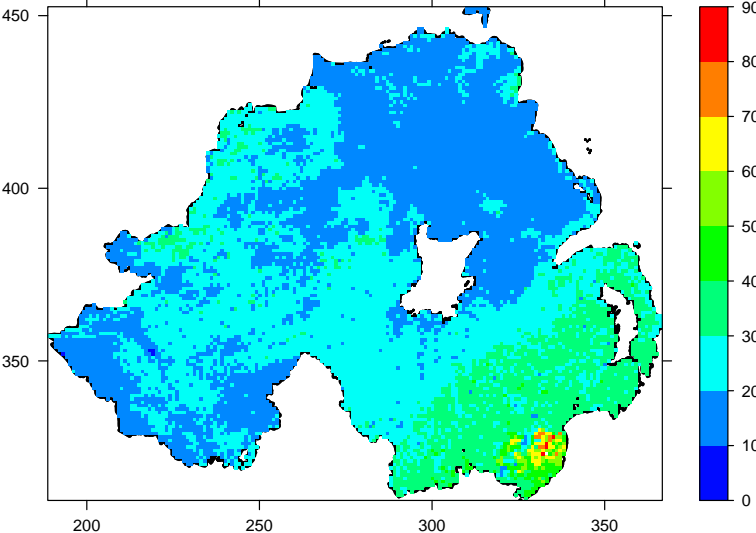
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# Illustrated contents entry



This optimal map of outdoor terrestrial gamma dose rate (nGy h<sup>-1</sup>) across Northern Ireland combines *in situ* measurements and dose estimates from a national scale, airborne radiometric survey.

# Abstract

Regulatory authorities need ways to estimate natural terrestrial gamma radiation dose rates ( $\text{nGy h}^{-1}$ ) across the landscape accurately, to assess its potential deleterious health effects. The primary method for estimating outdoor dose rate is to use an *in situ* detector supported 1 m above the ground, but such measurements are costly and cannot capture the landscape-scale variation in dose rates which are associated with changes in soil and parent material mineralogy. We investigate the potential for improving estimates of terrestrial gamma dose rates across Northern Ireland (13 542  $\text{km}^2$ ) using measurements from 168 sites and two sources of ancillary data: i) a map based on a simplified classification of soil parent material, and ii) dose estimates from a national-scale, airborne radiometric survey. We used the linear mixed modelling framework in which the two ancillary variables were included in separate models as fixed effects, plus a correlation structure which captures the spatially correlated variance component. We used a cross-validation procedure to determine the magnitude of the prediction errors for the different models. We removed a random subset of 10 terrestrial measurements and formed the model from the remainder ( $n=158$ ), and then used the model to predict values at the other 10 sites. We repeated this procedure 50 times. The measurements of terrestrial dose vary between 1 and 103 ( $\text{nGy h}^{-1}$ ). The median absolute model prediction errors ( $\text{nGy h}^{-1}$ ) for the three models declined in the following order: no ancillary data (10.8) > simple geological classification (8.3) > airborne radiometric dose (5.4) as a single fixed effect. Estimates of airborne radiometric gamma dose rate can significantly improve the spatial prediction of terrestrial dose rate.

# 1 Introduction

2 Exposure to natural, gamma radiation has deleterious effects on health by causing dam-  
3 age to nuclear DNA.<sup>1</sup> Regulatory authorities need effective ways to improve estimates  
4 of exposure to natural radiation, which accounts for more than 80% of the total dose  
5 to which humans are typically exposed. Typically half of this natural radiation is re-  
6 lated to radon gas and its decay products. Terrestrial gamma radiation contributes on  
7 average 13% of the average annual dose to the UK population.<sup>2</sup> Exposure to terrestrial  
8 gamma radiation varies widely across the landscape, largely due to geological variation  
9 of naturally occurring radioactive materials in rocks, soils and in building materials.  
10 The mineralogical composition of the upper 1 m of the land surface is largely controlled  
11 by the soil parent material (PM) – the combination of bedrock or overlying superficial  
12 (Quaternary) deposits where present. Groups of PM account for a large proportion of  
13 the variation in the concentrations of the three dominant gamma emitting radionuclides  
14 (potassium (K), thorium (eTh) and Uranium (eU)), particularly in recently glaciated  
15 landscapes where the geochemistry of the soil and PM are closely associated.<sup>3</sup> At large  
16 scales (e.g. >1000 km<sup>2</sup>), the spatial variation in absorbed dose rate will be greater  
17 than its temporal variation because dose rate is source-dominated. In other words, it  
18 largely reflects the gamma source rather than the temporal variations in attenuation  
19 of the source due to variations in soil and atmospheric conditions.

20 Measurements of absorbed dose rates in air (nGy h<sup>-1</sup>), which express gamma  
21 ray intensity from radioactive materials in the earth and atmosphere, are typically  
22 undertaken using an *in situ* detector supported 1 m above the ground. Due to con-  
23 straints of both time and cost, only a limited number of ground-based measurements  
24 can be undertaken across the landscape which cannot capture the spatial variation in  
25 absorbed dose rates. For example, in a survey across the UK a total of 128 measure-  
26 ments were undertaken,<sup>4</sup> a sampling density of around one site per 2000 km<sup>2</sup>, which is  
27 insufficient to account for the variation in PM groups across this landscape. Estimates  
28 of absorbed dose based on interpolation for locations where measurements have not

29 been undertaken will be subject to a substantial level of uncertainty.

30 It may be possible to significantly reduce these uncertainties using additional  
31 sources of landscape-scale data. Such approaches have been applied widely in the dig-  
32 ital soil mapping community<sup>5</sup> where covariates are often used within a geostatistical  
33 framework to account for the spatial variation in soil properties. One such covariate  
34 is measurement of natural emissions of gamma radiation from aerogeophysical sur-  
35 veys flown at low altitudes which cover large areas. Although terrestrial and airborne  
36 measurements are likely to be correlated, they will not estimate the same dose rate  
37 because: i) measurements will be undertaken at different times with differing atmo-  
38 spheric pressures and so gas fluxes from the ground will vary, ii) they have different  
39 measurement supports (the size and shape – in three dimensions – of the volume of ma-  
40 terial constituting the source of the gamma radiation for which a single measurement  
41 is determined) and, iii) corrections for cosmic sources may be somewhat different.

42 To our knowledge, low spatial resolution terrestrial measurements of gamma ra-  
43 diation dose ( $\text{nGy h}^{-1}$ ) have not previously been compared to high-resolution, airborne  
44 estimates of gamma radiation to investigate whether the latter could be used as an-  
45 cillary variables to map the former with smaller prediction errors. Such an approach  
46 requires geostatistical techniques in which optimal spatial predictions of a primary  
47 property – terrestrial dose rate – are made using both primary measurements and sec-  
48 ondary sources. The linear mixed modelling framework has been successfully applied  
49 to explain variations in airborne radiometric estimates of K concentration in the soil  
50 across part of central England based on variations in soil geochemistry and PM.<sup>6</sup>and to  
51 improve spatially explicit estimates of soil organic carbon concentration in soils across  
52 northern Ireland. <sup>7</sup>

53 In this paper we compare data from two sets of terrestrial dose measurements  
54 <sup>8,4</sup> with measurements of airborne gamma radiation across all of Northern Ireland. <sup>9</sup>  
55 We explore the relative importance of soil PM in determining gamma emissions from  
56 K, eTh and eU. We use geostatistical models to compare the magnitude of errors

57 in estimating terrestrial gamma dose rate at unsampled locations by utilising: i) the  
58 original terrestrial survey data, ii) a simple geological classification and, iii) the airborne  
59 survey data. We undertake this analysis using a cross-validation procedure so that  
60 prediction errors could be compared. We comment on the implications of our findings  
61 for mapping adsorbed dose rate at landscape scales.

## 62 **Methods**

### 63 **Study region**

64 There is a wide range of bedrock types and Quaternary deposits across Northern Ireland  
65 which give rise to strong contrasts in near surface mineralogy, and thus concentrations  
66 of gamma emitting radionuclides (Figure 1). The north-eastern quadrant is dominated  
67 by Palaeogene basalt lava and lacustrine sedimentary rocks (Lough Neagh) at the sur-  
68 face, whilst the north-west is dominated largely by Dalradian psammite and semipelite.  
69 There are sedimentary outcrops of mudstone, sandstone and limestone across central  
70 to south-west Northern Ireland, which are mainly Carboniferous in age (with a De-  
71 vonian component). The south-east comprises Ordovician and Silurian marine sedi-  
72 mentary rocks of the Southern-Uplands-Down-Longford Terrane with younger igneous  
73 complexes. Extensive Palaeogene granite bedrock forms the Mourne mountains to the  
74 south-east, close to the other felsic igneous complexes: the contemporaneous Sileve  
75 Gullion complex and the older Newry complex.

76 Extensive Quaternary deposits cover much of Northern Ireland including peat  
77 in upland areas, with large regions of superficial glacial till, glacial sands and gravels  
78 and river alluvium. We created a PM map of all of Northern Ireland by merging the  
79 bedrock and Quaternary geological maps in a GIS. This created a set of polygons, each  
80 having one of 65 unique parent material codes associated with it.

### 81 **Airborne radiometric survey**

82 The Tellus airborne geophysical survey of Northern Ireland was flown in 2005 and 2006

83 and comprised more than 82 000 km of flight lines. High-resolution radiometric data  
84 were acquired with an Exploranium GR820 256-channel NaI(Tl) gamma spectrome-  
85 ter system of four downward (32 litres) and one upward looking (8 litres) crystal sets.  
86 Data from the sensor were recorded every second corresponding to approximately 60-m  
87 intervals in the direction of flight. Survey lines were spaced 200 m apart with flight  
88 lines oriented at 345° and 165°. The aircraft flew 56 m above ground in rural areas,  
89 but higher (244 m) above elevated structures and urban areas. The ground area or  
90 ‘footprint’, from which most of the radiation recorded comes, has the form of an ellipse  
91 elongated in the direction of flight. For example, at 56 m above the ground, 75% of  
92 the measured radiation comes from an ellipse with a major diameter of 220 m in the  
93 flight direction and minor diameter of 150 m.<sup>13</sup>A complete description of the airborne  
94 geophysical systems deployed and the processing methods used is given by Hautaniemi  
95 et al.<sup>14</sup> Procedures for processing the airborne radiometric data were based on those  
96 described in the reference manuals of the International Atomic Energy Authority<sup>15</sup> and  
97 the Australian Geological Survey Organisation.<sup>16</sup> The processing included corrections  
98 for aircraft and cosmic background radiation, altitude of the aircraft and spectral inter-  
99 actions. Gamma radiation measured by the airborne system comes from a thin surface  
100 layer of about 30 cm in rock, rather more in less dense, unconsolidated material such  
101 as mineral soil. Estimated activities of <sup>137</sup>Cs were also determined, based on a full  
102 spectrum processing.<sup>9</sup> The corrected count rates were used to estimate the concentra-  
103 tion of three natural radioelements within the conventional energy ranges: potassium  
104 (K, 137-157 MeV), equivalent uranium (eU, derived from <sup>214</sup>Bi, 166-186 MeV) and  
105 equivalent thorium (eTh, derived from <sup>208</sup>Tl, 241-281 MeV).<sup>15,17</sup>The survey yielded *ca.*  
106 1.2 million values for K(%), eTh (mg kg<sup>-1</sup>) and eU (mg kg<sup>-1</sup>) and activities of <sup>137</sup>Cs  
107 (kBq m<sup>-2</sup>). Total gamma dose rate (nGy h<sup>-1</sup> was then computed from the total count  
108 rate, which included the activities of of the three naturally occurring elements (K, eTh,  
109 and eU) and also <sup>137</sup>Cs.

110 To assess the relative importance of PM in determining the concentration of each

111 of the three gamma-emitting elements (K, eTh and eU) plus dose rate, we assigned a  
112 code to every airborne survey site based on the 65 main types of PM (described above).  
113 We explored these data by constructing box and whisker plots for each of the three  
114 elements (K, eTh and eU) and also dose rate using the PM codes as the classes. In each  
115 case we calculated the proportion of variance accounted for by the PM classification  
116 using one-way ANOVA. For the purpose of producing a map of terrestrial gamma  
117 dose rate incorporating the airborne estimates we used the linear mixed modelling  
118 framework (see below). For each node on a one kilometre grid across all of the land  
119 surface of Northern Ireland we calculated the average of the nearest eight airborne dose  
120 estimates and used these as a fixed effect for prediction at the grid nodes.

#### 121 **Terrestrial dose rate survey data**

122 In 1989, a nation-wide survey of terrestrial gamma outdoor dose rates was conducted  
123 across Northern Ireland<sup>8</sup> at 158 sites. Measurements were taken from each 10 km square  
124 of the Ordnance Survey grid throughout Northern Ireland using a Mini-Instruments  
125 Environmental Monitor Type 6-80 with an energy compensated Geiger-Muller tube  
126 MC-71 at a height of 1 m above the ground. For use in this present study, measurements  
127 of corrected terrestrial gamma dose rate ( $\text{nGy h}^{-1}$ ) were taken from this published  
128 report and georeferenced using the associated Irish National Grid coordinates.<sup>8</sup> for use  
129 in our study. A further ten measurements of terrestrial gamma dose rate for sites across  
130 Northern Ireland were available for sites that were part of the UK Soil and Herbage  
131 Survey.<sup>4</sup> Both sets of terrestrial measurements were pre-corrected for cosmic radiation.  
132 The methods from these two surveys were sufficiently similar for the results to be  
133 combined into a single dataset of measurements at 168 sites (Figure 2).

134 First we assigned each terrestrial survey location a PM code using the 65-fold  
135 classification described previously. We then simplified these PM codes using expert  
136 knowledge to form a set of aggregated classes. For example, peat deposits typically  
137 contain small quantities of mineral material and so have low total gamma dose rates.



138 Where peat deposits have been mapped they are more than 1 m deep, so such sites  
139 were assigned to the single class ‘peat’. Similar rules were established to aggregate the  
140 classes, forming an initial set of 11. These were given codes which partly related to  
141 their origin; ‘BASA’=basalt, ‘TILL’=glacial till, ‘SDST’= sandstone, ‘ALV’=alluvium,  
142 etc. Any remaining PM groups more difficult to aggregate were grouped in a mixed  
143 class with the code ‘MIX’.

#### 144 **Comparison of terrestrial and airborne dose measurements**

145 The sample support – the length, area or volume across which each sample measure-  
146 ment is made – were different for the terrestrial and airborne gamma surveys. For the  
147 former it was an area of a few metres and much larger for the latter as described above.  
148 In addition, their locations were not coincident. To estimate the airborne gamma dose  
149 rate at the terrestrial survey sites (Figure 2) we calculated the average of the nearest  
150 airborne survey sites. To identify the neighboring airborne survey locations we used  
151 the *ann* function in the *yaImpute* package<sup>18</sup> in the R environment.<sup>19</sup> We calculated  
152 the average of the nearest 5, 8 and 10 neighbours. The average of 8 neighbours had  
153 the strongest linear (Pearson) correlation ( $r=0.8$ ) with the terrestrial survey dose rates  
154 (see Figure 3) and so we used these data in our subsequent analyses.

#### 155 **Geostatistical analysis using the linear mixed model**

156 We wished to explore to what extent geological class or airborne estimates of dose  
157 rate could account for the spatial variation in terrestrial gamma dose rates. For this  
158 purpose we used the linear mixed model, which we can write as

$$\mathbf{z} = \mathbf{X}\boldsymbol{\tau} + \mathbf{Y}\boldsymbol{\eta} + \boldsymbol{\varepsilon} . \quad (1)$$

159 Here the vector  $\mathbf{z}$  contains our  $n$  observations of terrestrial gamma dose rate. Matrix  
160  $\mathbf{X}$  is an  $n \times p$  design matrix that associates each of the  $n$  observations with a value  
161 of each of the  $p$  fixed effects, here a set of dummy variables that identify the parent

162 material class at each site or airborne gamma dose rates or both. The vector  $\boldsymbol{\tau}$  contains  
 163 the  $p$  fixed-effect coefficients. The vector  $\boldsymbol{\eta}$  contains  $q$  random effects, realizations of  
 164 a variable  $\eta$ , that are associated with the  $n$  observations by the  $n \times p$  design matrix  
 165 ( $\mathbf{Y}$ ), which here is the identity matrix and  $n = q$ . We assume that  $\eta$  is a spatially  
 166 correlated random variable and  $\boldsymbol{\varepsilon}$  is a vector of independent random errors. These  
 167 terms are independent of each other, and so we may write

$$\begin{bmatrix} \boldsymbol{\eta} \\ \boldsymbol{\varepsilon} \end{bmatrix} \sim \mathcal{N} \left( \begin{bmatrix} \mathbf{0} \\ \mathbf{0} \end{bmatrix}, \begin{bmatrix} \sigma^2 \xi \mathbf{R} & \mathbf{0} \\ \mathbf{0} & \sigma^2 \mathbf{I} \end{bmatrix} \right), \quad (2)$$

168 where  $\sigma^2$  is the variance of the independent error,  $\xi$  is the ratio of the variance of  $\eta$   
 169 to  $\sigma^2$  and  $\mathbf{R}$  is the correlation matrix of  $\boldsymbol{\eta}$ . Note that we make an explicit assump-  
 170 tion that the random terms are jointly Gaussian. The term  $\boldsymbol{\varepsilon}$ , the nugget, represents  
 171 both independent measurement errors and variation that arises from processes that are  
 172 spatially dependent over shorter distances than those that separate the closest pairs of  
 173 sampling points. Under the assumption that  $\boldsymbol{\eta}$  is drawn from a second-order stationary  
 174 random process, the correlation matrix  $\mathbf{R}$  will depend only on the relative locations  
 175 of our observations given some specified correlation function  $C(\cdot)$  with one or more  
 176 parameters that characterize the spatial dependence; so

$$\mathbf{R}_{i,j} = \text{Corr} [\boldsymbol{\eta}(\mathbf{s}_i), \boldsymbol{\eta}(\mathbf{s}_j)] = C(\mathbf{s}_i - \mathbf{s}_j). \quad (3)$$

177 The correlation function may be one of several authorized functions.<sup>20</sup> The correlation  
 178 function could be more complex with parameters that describe spatial anisotropy, but  
 179 our exploratory analysis suggested that such elaboration was unnecessary. We esti-  
 180 mated the parameters of the exponential function, which we represent by the vector  
 181  $\boldsymbol{\theta}$ , along with  $\sigma^2$  and  $\xi$  by residual maximum likelihood (REML). The REML solu-  
 182 tion removes dependence of the estimates of the parameters in  $\boldsymbol{\theta}$  on the fixed effects  $\boldsymbol{\tau}$   
 183 which are nuisance parameters in this problem and which would increase the bias of  
 184 estimates based on maximum likelihood or method-of-moments.<sup>21</sup> Once the variance  
 185 parameters are estimated we can obtain estimates of the fixed effects ( $\boldsymbol{\tau}$ ) by generalized  
 186 least squares. More details are given elsewhere.<sup>22,23</sup>

187 Before fitting the linear mixed models, we undertook exploratory statistical analy-  
188 sis of the terrestrial and airborne survey gamma dose data. If data are strongly skewed  
189 (e.g. absolute skewness coefficients  $>1$ ) this can present problems for geostatistical  
190 analysis because a variogram calculated from such data may be strongly biased. The  
191 skewness coefficient for the untransformed, terrestrial dose data showed some positive  
192 skewness (skewness coefficient=1.49). In deciding whether to transform these data  
193 prior to our geostatistical analyses we had to consider the other models we planned to  
194 fit. Specifically, when using the airborne survey data as a fixed effect, the correlation  
195 function is calculated from the set of terrestrial dose rate residuals which may be close  
196 to normally distributed. If this were the case, it may be preferable not to transform the  
197 data because we wish to make comparisons between the different models using data  
198 on the same scale. To investigate this further, we fit a simple, least squares model  
199 between the paired (n=168) airborne dose (predictor) and terrestrial dose rate (pre-  
200 dictand) data; the residuals were close to normally distributed (skewness coefficient =  
201 0.56). We therefore chose to undertake all our analyses on the original, untransformed  
202 data to ensure that the results were consistent and comparable.

203 We used the *lme* function in the R package *nlme*<sup>24</sup> which fits linear mixed models  
204 which has an option to include a spatial correlation structure. We fitted both spatial  
205 and exponential covariance models and selected one of these using the log likelihood  
206 statistic. The residual likelihood statistic is only comparable between models with  
207 common fixed effects. We found that the exponential correlation function was optimal  
208 for estimating the spatial structure of the terrestrial gamma dose rate data with a  
209 range of fixed effects. The following fixed effects models were tested for the prediction  
210 of the terrestrial gamma dose rate:

- 211 1. The mean value only.
- 212 2. The mean value plus parent material class.
- 213 3. The mean value plus the airborne gamma dose rate ( $\text{nGy h}^{-1}$ ).

214 4. The mean value plus airborne gamma dose rate ( $\text{nGy h}^{-1}$ ) plus parent material  
215 class.

216 The decision on whether to consider both airborne dose rate and parent material as  
217 fixed effects was made by comparing the model fits (in this case by maximum likelihood  
218 because likelihood estimates for two models by REML are not directly comparable) by  
219 reference to their log-likelihood ratio using the ANOVA command in R. The result ( $P$ -  
220 value=0.18) suggested that there was no basis for inclusion of both fixed effects so the  
221 fourth model in the list above is not considered further.

222 We compared the performance of the three different models using cross validation.  
223 We selected a random subset of 158 from the total 168 sites to build each of the  
224 models, and then estimated the dose rates at the remaining 10 sites. We repeated  
225 this procedure 50 times which gave 500 predictions that we compared to the measured  
226 values by calculating the absolute error. We then calculated the mean absolute error  
227 (MAE) and median absolute error (MedAE) and the bias for the estimates ( $n=500$ ) for  
228 each model. Using the airborne data as a fixed effect (model 3 above), we created an  
229 optimal map of terrestrial dose rate across all of Northern Ireland by estimating values  
230 on a 1 km grid using this model; the empirical, best least squares unbiased predictor  
231 or E-BLUP.<sup>23</sup>

## 232 Results and discussion

233 The summary statistics for terrestrial and airborne dose rates (Table 1) show that there  
234 is a positive bias to the distribution of the latter compared to the former; both the  
235 median and the maximum are larger. This may be in part be explained by differences  
236 in the procedures for estimating dose used in the terrestrial and airborne systems, and  
237 perhaps also the more complete coverage of the airborne survey; remote areas were not  
238 accessed by the ground based survey.

239 Based on a one-way ANOVA, classes of the original parent material map ac-  
240 count for large proportions of the variance for each of the concentrations of the three

241 radiogenic elements and dose rate, estimated from the airborne survey: 47% (K), 44%  
242 (eTh), 22% (eU) and 52% (dose). The relationship between each of these variables and  
243 15 PM classes – selected to represent a broad range of mineralogy – is shown in Figure 4  
244 as a box and whisker plot. There are clear differences between the distributions for  
245 many of these classes. For example, areas with sandstone and granitic bedrock have  
246 amongst the largest median K and eTh concentrations which accounts for the larger  
247 dose estimates over these lithologies. By contrast, basalt bedrock has low concentra-  
248 tions of all three gamma emitting elements – particularly where this is covered by peat  
249 deposits (Figure 4).

## 250 **Models and cross-validation**

251 The parameters for the fixed effects in each of the three models are shown in Table 2;  
252 the parameters of the spatially correlated components are shown in Table 3. Where  
253 airborne survey data are not included in the model, the PM classes with the most  
254 significance for predicting terrestrial dose rate are basalt bedrock and superficial peat;  
255 these have the smallest  $P$ -values (Table 2). The airborne survey data are substantially  
256 more significant ( $P$ -values  $<0.0001$ ) than the PM codes for predicting terrestrial dose  
257 rate. In each case, a substantial component of variance is captured by the spatial  
258 correlation structure. The fixed effects capture a substantial proportion of the variance  
259 in terrestrial dose rate.

260 The results of the cross-validation ( $n=500$ ) in predicting terrestrial dose rate are  
261 summarised in Table 4. The mean and median absolute predictions errors, according  
262 to inclusion of fixed effects in the model, decline in the following order: mean value >  
263 mean value + PM > mean value + airborne >. The simple geological classification  
264 results in a modest reduction in the MedAE for dose prediction; declining from 10.8 to  
265 8.29 nGy h<sup>-1</sup>. A greater reduction in prediction error results on including the airborne  
266 dose estimates (MedAE=5.39 nGy h<sup>-1</sup>). The bias in the prediction errors also declines  
267 in the same order as for the prediction errors reported above; the smallest prediction

268 bias is for the model in which airborne dose data are included as a fixed effect.

### 269 **Optimal map of terrestrial gamma dose rate**

270 The map of terrestrial gamma dose rate (Figure 5) provides a far greater level of detail  
271 than the interpolation of the data values presented in the original report on terrestrial  
272 gamma dose rates.<sup>8</sup> The largest dose rates are associated with the granite bedrock  
273 of the Mourne mountains and the arenites across the south-eastern part of Northern  
274 Ireland (Figure 1). The smallest dose rates occur across the area of basalt bedrock,  
275 where the common occurrence of superficial, organic peat deposits dilutes the gamma  
276 signal from the minerals derived from the underlying rock.

277 Our study demonstrates that using a simplified map delineating twelve classes of  
278 PM can significantly improve the estimation of terrestrial gamma dose. The relation-  
279 ships between PM and terrestrial gamma dose may be quite different in landscapes  
280 which have been subject to longer (geological) periods of weathering such as tropical  
281 regions, and in such environments the distribution of geological parent materials may  
282 be less effective in accounting for variations in gamma dose. In some areas, human  
283 exposure to naturally occurring radioactive materials may be increased as a result  
284 of mining and transportation of raw materials. This technological enhancement of  
285 naturally occurring radioactive materials (TENORM) may result as a by-product of  
286 a variety of industrial and other activities including mining, extracting, concentrat-  
287 ing, processing or combusting raw materials containing naturally occurring radioactive  
288 materials.<sup>25</sup> Including the distribution of PM to improve mapping of terrestrial dose  
289 cannot account for the occurrence and distribution of TENORM.

290 Our findings show that terrestrial gamma dose rates across much of Northern  
291 Ireland are generally small ( $< 30 \text{ nGy h}^{-1}$ ). High-resolution maps of gamma dose rate  
292 are likely to be of most use for regulatory authorities in areas where dose rates are  
293 greatest ( $>60 \text{ nGy h}^{-1}$ ). Our study shows that airborne radiometric is likely to be the  
294 most effective covariate for enhancing maps of terrestrial dose rate for a smaller set of

295 ground-based measurements.

296 Most of the primary legacy data used in our study were a set of 158 terrestrial  
297 gamma dose measurements undertaken in 1989 and published as a paper report which  
298 included the grid coordinates of each site.<sup>8</sup> Our study highlights the importance of  
299 ensuring that such legacy data are available so that new datasets may be combined  
300 with them to generate enhanced outputs – such as our optimal map of terrestrial dose  
301 rate.

## 302 **Conclusions**

303 The main findings from our study are:

- 304 1. A map of geological parent material classes across all of Northern Ireland accounted  
305 for 52% of the variation in gamma dose rate estimated from an aerogeophysical  
306 survey. Those areas with the largest estimated dose rates had geological parent  
307 materials with greater concentration of gamma emitting elements. For example,  
308 some of the largest dose rates were observed over areas of granite, the soils over  
309 which have large concentrations of both K and eTh.
- 310 2. Using a series of terrestrial gamma dose rate measurements ( $\text{nGy h}^{-1}$ :  $n=168$ )  
311 across all of Northern Ireland, a simplified PM map ( $n=12$  classes) significantly  
312 improved their estimation using a statistical model and a cross-validation proce-  
313 dure. The median absolute error in estimated dose rate declined from 10.8 (no  
314 PM classification) to  $8.29 \text{ nGy h}^{-1}$  (with a PM map).
- 315 3. Incorporating the estimated dose rates from the airborne survey led to a greater  
316 reduction in the error of estimating terrestrial gamma dose rate. The median  
317 absolute error from the cross-validation analysis after applying the linear mixed  
318 model was  $5.39 \text{ nGy h}^{-1}$ . The map of terrestrial gamma dose rate across all of  
319 Northern Ireland incorporating the airborne data shows a resolution which would  
320 not be possible based on ground-based measurements because the latter are too

321           costly to collect at the resolution of the airborne estimates.

322   4. Terrestrial gamma dose rates across much of Northern Ireland are generally small  
323       ( $< 30 \text{ nGy h}^{-1}$ ).

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396 **List of Figures and Captions**

397 **Figure 1** Simplified bedrock map of across Northern Ireland.

398 **Figure 2** The spatial distribution of terrestrial gamma dose rate measurements from  
399 the Northern Ireland survey (n=158; Caulfield and Ledgerwood, 1989) and the  
400 UK soil and herbage survey (n=10; Tyler and Copplestone, 2007).

401 **Figure 3** Scatterplot of airborne and terrestrial gamma dose rate measurements for  
402 168 terrestrial measurement sites across Northern Ireland. The airborne dose  
403 rates are the average of the eight airborne sites closest to each terrestrial site.

404 **Figure 4** Box and whisker plot showing the estimated concentrations of the three  
405 gamma emitting elements (K, eTh and eU) and estimated dose rate in the  
406 upper part of the solum for 15 selected combinations of PM for all the air-  
407 borne survey sites. The width of each box is proportional to the number of  
408 sites in each class. The PM codes refer to combinations of the Quaternary  
409 (TILL=till, ALV=alluvium, GSG=glacial sands and gravels, PEAT=peat) and  
410 specific bedrock (SDST=Quarry sandstone, PSSP=psammite and semi pelite,  
411 LMST=limestone, BASA=basalt, GRAN=Granite, COSD=conglomerate, LATU=Cromarty  
412 Sandstone, ROCK=a range of smaller bedrock types with no Quaternary deposit  
413 present) deposits.

414 **Figure 5** Map of terrestrial gamma dose rate ( $\text{nGy } h^{-1}$ ) across Northern Ireland on  
415 a 1-kilometre grid. The estimates were generated using a linear mixed model  
416 in which the fixed effects were the mean value (measured terrestrial dose) and  
417 measured airborne dose rate plus a spatially correlated component. Coordinates  
418 are kilometres on the Irish National Grid.

419 **Table 1** Summary statistics for surveys of terrestrial and airborne gamma dose rate  
 420 across Northern Ireland.

421

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	Terrestrial gamma dose (nGy h <sup>-1</sup> )	Airborne gamma dose (nGy h <sup>-1</sup> )
Minimum	1.0	<sup>a</sup> -9.0
Mean	23.2	33.3
422 Median	23.0	32.6
Maximum	103	320
Standard deviation	13.3	22.5
Skewness	1.49	1.19
Number of sites	<sup>b</sup> 168	1 230 440

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423 <sup>a</sup> the greater uncertainties in the airborne measurements and the corrections applied  
 424 to account for cosmic radiation and aircraft background can lead to negative estimates  
 425 of dose rate

426 <sup>b</sup> includes data from the NI survey (Caulfield and Ledgerwood, 1989) and the UK soil  
 427 and herbage survey (Tyler and Copplestone, 2007)

428 **Table 2** Parameter estimates of the fixed effects for four linear mixed models (see text)  
 429 used for the estimation of terrestrial dose rate. The values shown are those based on a  
 430 model fit to data from all 168 terrestrial dose sites.

Model parameters	Value	Std.Error	<i>t</i> -value	<i>P</i> -value
a) mean only				
Intercept	24.7	6.43	3.85	$2 \times 10^{-5}$
b) mean + PM class				
<sup>a</sup> Intercept	26.1	9.59	2.72	0.0075
BASA	-11.6	4.12	-2.81	0.006
GRAN	-2.04	5.24	-0.39	0.70
GSG	1.11	4.37	0.25	0.80
LAT	-5.70	5.11	-1.12	0.27
PEAT	-10.2	4.38	-2.32	0.02
431 PSSP	-8.06	4.61	-1.75	0.08
ROCK	-2.20	4.06	-0.54	0.59
SDSM	2.03	5.28	0.38	0.70
SDST	-1.36	3.63	-0.37	0.71
TILL	-4.19	3.97	-1.05	0.29
MIX	-3.26	4.40	-0.74	0.46
c) mean + airborne dose				
Intercept	9.36	4.43	2.11	0.0363
<sup>b</sup> Air	0.395	0.045	8.68	<0.0001

432 <sup>a</sup> the coefficient for the PM class ALV (alluvium) is not shown because all the other  
 433 PM class intercept values are differences from the ALV coefficient.

434 <sup>b</sup> refers to inclusion of the airborne radiometric data as a fixed effect.

435 **Table 3** Parameter estimates of the spatially correlated variance components for four  
 436 linear mixed models (see text) used for the estimation of terrestrial dose rate. The  
 437 values shown are those based on a model fit to data from all 168 sites. In all cases an  
 438 exponential spatial correlation model was selected.

Model	Mean	Mean+PM	Mean+airborne
range (metres)	66416	40372	53266
nugget variance	52.4	38.9	43.0
sill variance	166	86.6	67.9

439

440 **Table 4** Cross-validation statistics for estimation of terrestrial gamma dose rate (n=500)  
 441 using linear mixed models with different fixed effects. Ten sites were selected randomly  
 442 from the full dataset (n=168) and models were formed from data at the other 158 sites.  
 443 Estimates of terrestrial gamma dose rate were then computed for the ten independent  
 444 sites. The procedure was repeated 50 times to give a total of 500 estimates which were  
 445 compared with the actual values to calculate mean absolute error (MAE), median ab-  
 446 solute error (MedAE) and bias. PM refers to parent material class and airborne refers  
 447 to the estimate of airborne radiometric dose rate.

Model	MAE (nGy h <sup>-1</sup> )	MedAE (nGy h <sup>-1</sup> )	bias (nGy h <sup>-1</sup> )
Mean value only	12.1	10.8	-10.2
Mean + PM	10.1	8.29	-3.26
Mean value + airborne	7.14	5.39	-0.72



Figure 1:

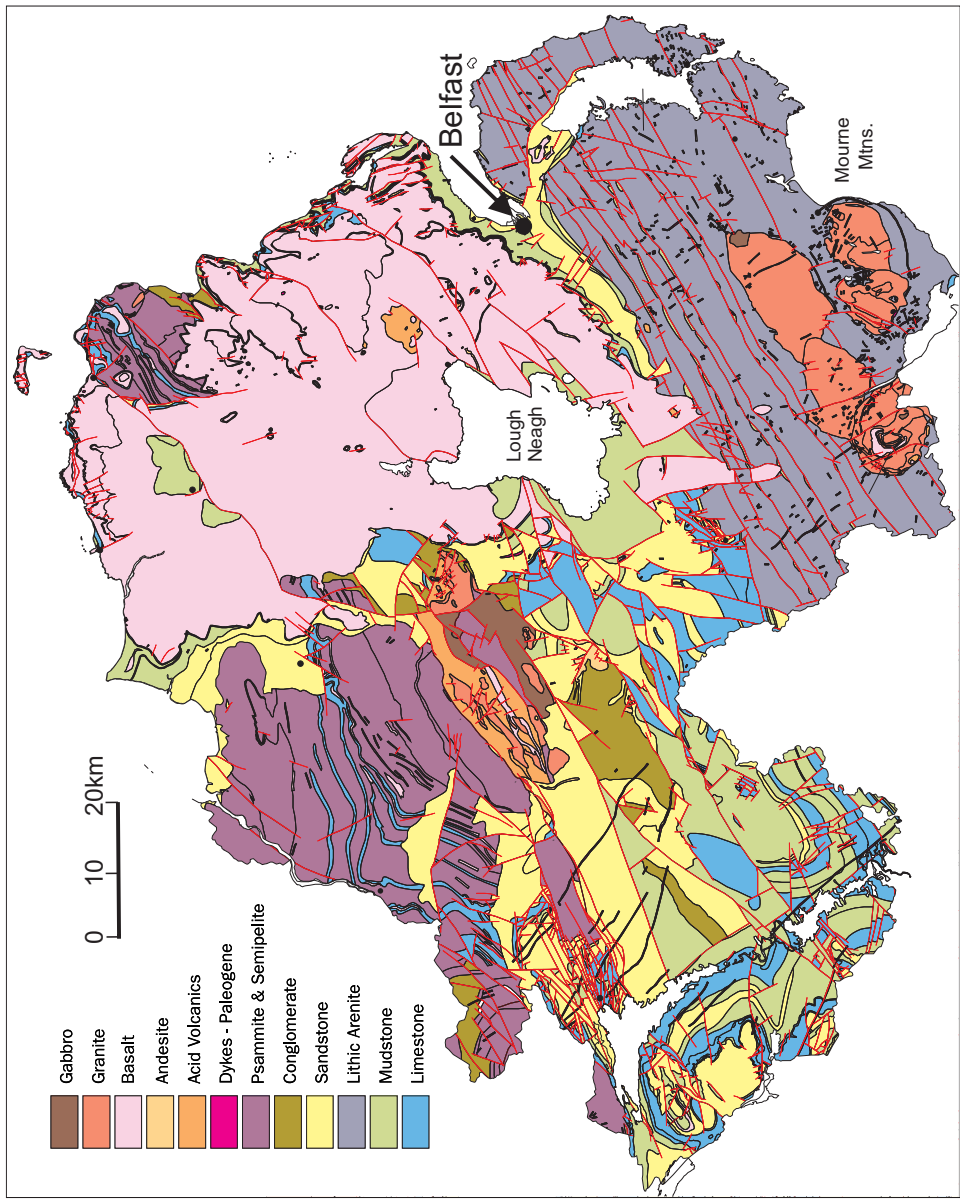


Figure 2:

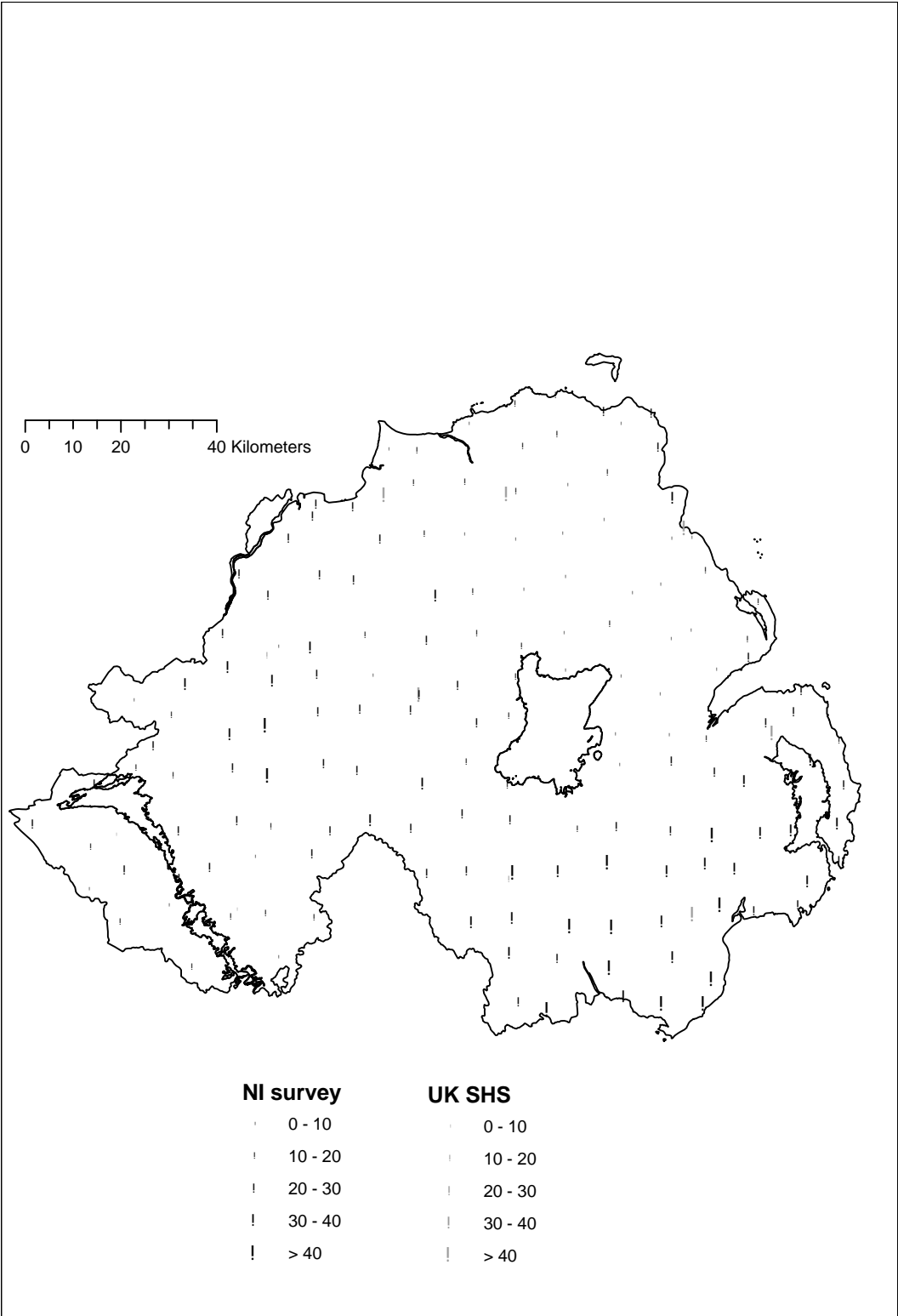


Figure 3:

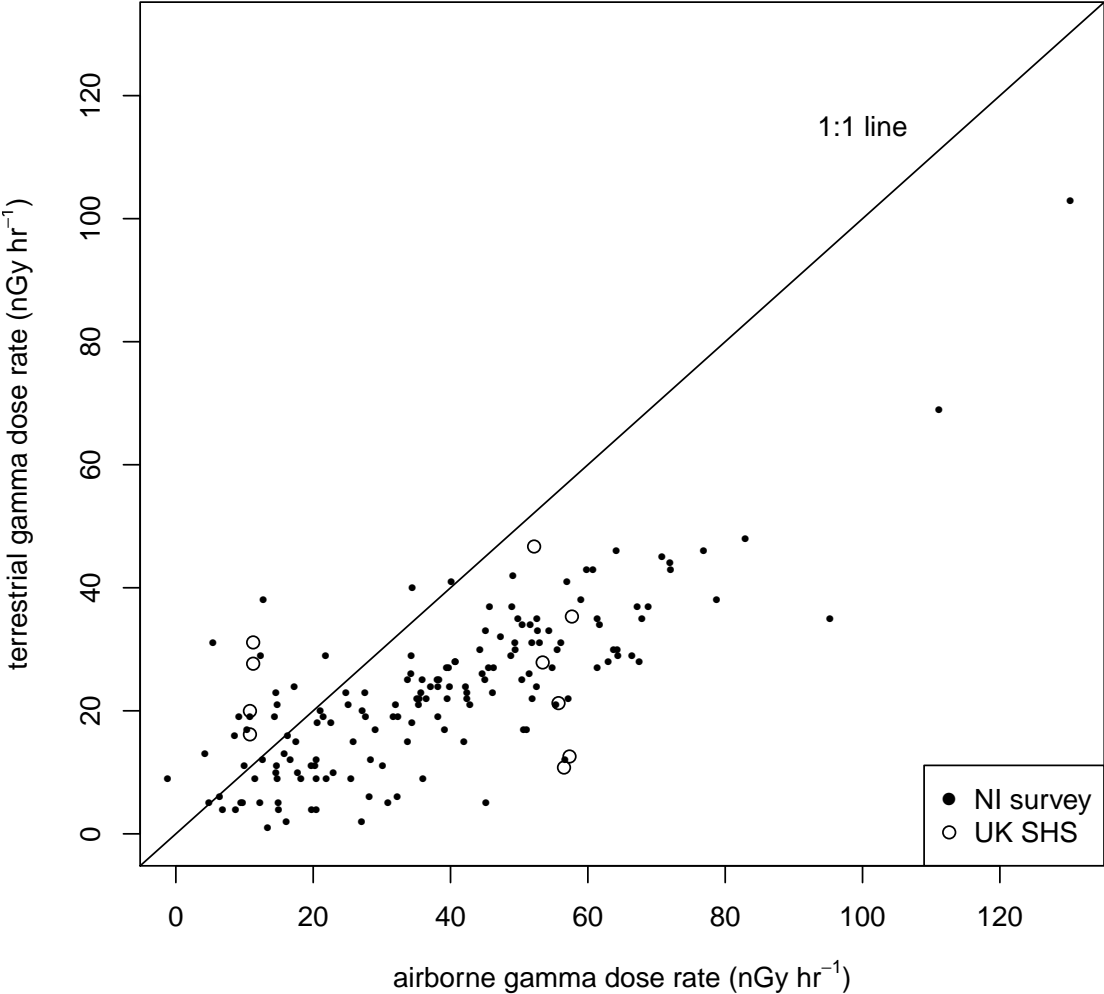


Figure 4:

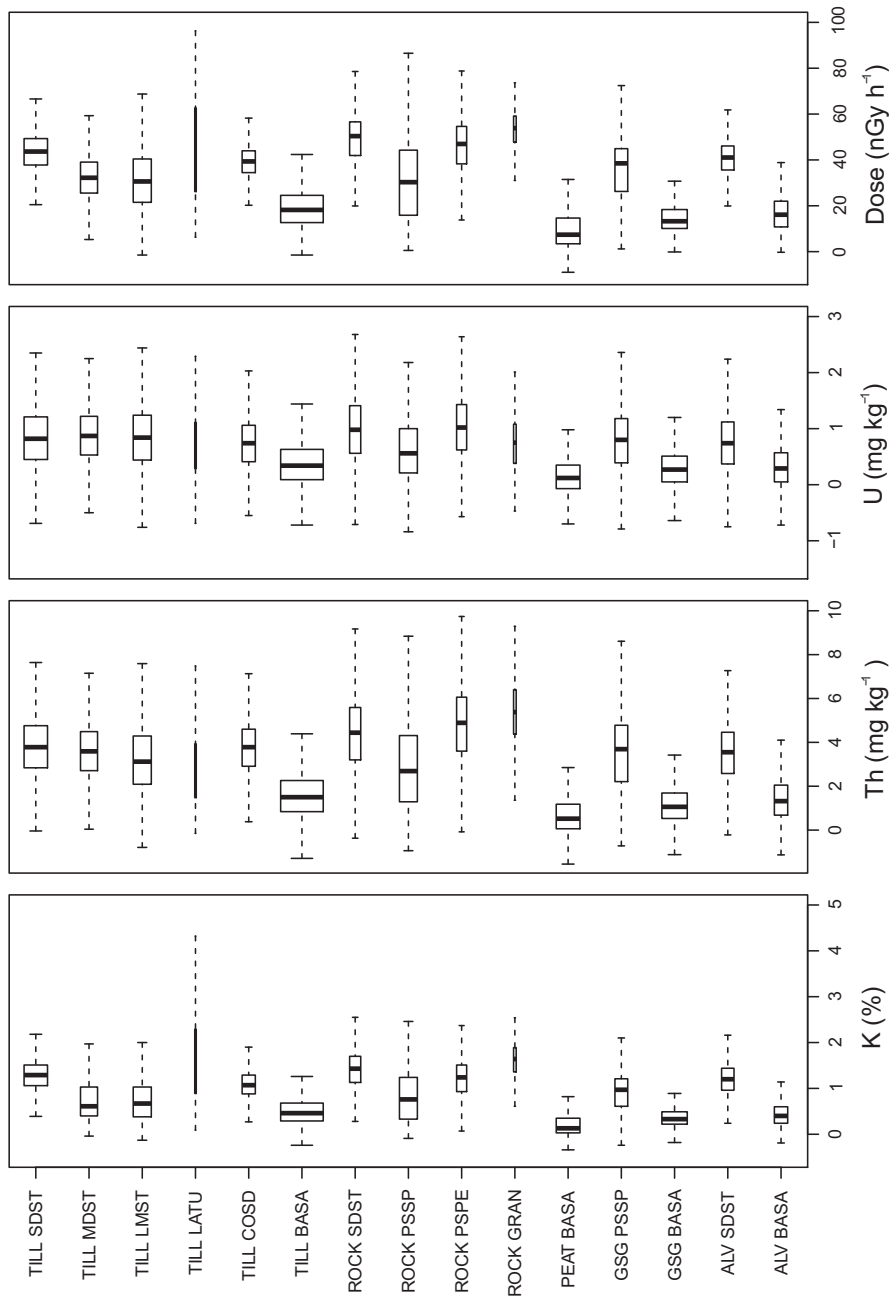


Figure 5:

