

Climate change may drive cave spiders to extinction

Stefano Mammola ^{1,2}, Sara L. Goodacre ³ and Marco Isaia ^{1,2*}

1. Department of Life Sciences and Systems Biology, University of Torino, Torino, Italy

2. IUCN SSC Spider & Scorpion Specialist Group, Torino, Italy

3. School of Biology, University of Nottingham, Nottingham, UK

* Corresponding author: marco.isaia@unito.it (**Tel:** 011 6704544; **Fax:** 011 6704508)

Keywords: Subterranean biology, Ecological Niche Modeling, Local extinction, Global Warming, Pleistocene glaciation, Araneae, *Troglohyphantes*, MaxEnt

Running title: Predicting cave spiders' extinction

Authors contribution statement

MI and SM conceived the study, collected and analyzed the data. SLG provided fundamental arguments for the discussion. All authors contributed to the writing of the manuscript.

20 **Abstract**

21 Subterranean ecosystems present ideal opportunities to study mechanisms underlying responses to
22 changes in climate because species within them are often adapted to a largely constant temperature. We
23 have characterized the thermal conditions of caves in the Western Alps, and relate these hypogean climate
24 data to the occurrence of *Troglohyphantes* spiders (Araneae, Linyphiidae). Our data indicate that present
25 distributions reflect Pleistocene glaciation events and also point to specific responses as a consequence of
26 changes in temperature. Constant temperatures recorded inside caves provide an approximation of the
27 mean annual temperature outside, thus we extend the results to a regional scale. We used ecological niche
28 modeling to predict habitat suitability both in the Pleistocene and under future global warming scenarios.
29 These analyses point toward a future decline in habitat suitability for subterranean spiders and the potential
30 extinction of the most restricted endemic species. When compared with other species that live in confined
31 habitats such as islands and mountains, we expect cave species to be as much, if not more, vulnerable to
32 climate change.

33 **Introduction**

34 The potential impact of global climate change on a broad range of organisms occurring in different
35 ecosystems and showing a very diverse pattern of distributions is well documented (Walther et al. 2002,
36 Parmesan and Yohe 2003, Parmesan 2006, Walther 2010, Chen et al. 2011, Dawson et al. 2011, Bellard et
37 al. 2012). However, the effects on biological communities in hypogean ecosystems have not been well
38 explored (but see Brandmayr et al. 2013, Wynne et al. 2014, Sánchez-Fernández et al. 2016).

39 From an ecological standpoint, hypogean ecosystems are considered extreme, lacking solar radiation,
40 receiving poor energy inputs, and through being characterized by low abundance and diversity of organisms
41 (Poulson and White 1969, Culver and Pipan 2009, Romero 2009). Most caves have almost constant
42 temperature over the year, with an annual variability decreasing progressively with increasing distance from
43 the entrance until a remarkable temperature steadiness is reached (Moore and Sullivan 1964, Smithson
44 1991), usually varying by only a few tenths of a degree over the year (Badino 2010).

45 According to direct field observations (Fejér and Moldovan 2013, Dominguez-Villar et al. 2015) and
46 theoretical models (Badino 2004, Covington and Perne 2015), anthropogenic global warming may
47 significantly influence and modify the underground climate in different ways to that which happens outside.
48 The most important difference refers to the thermal inertia of caves, which explains time lags of up to several
49 decades in the warming of the system (Badino 2004). This increased flux of energy from the atmosphere to
50 the subterranean environment is expected to be deposited primarily in the superficial sections of caves
51 (Badino 2004) and in other superficial subterranean habitats (Mammola et al. 2016). Energy fluxes
52 accumulating underground may have dramatic cascade effects on both physical and biological components
53 of hypogean ecosystems. By affecting air density gradients, for instance, small changes in cave temperature
54 modify air circulation, with potential fallouts on speleothem growth, seasonal ventilation rates, and processes
55 of condensation and corrosion (Dominguez-Villar et al. 2015).

56 Given the reduced physiological tolerance to temperature fluctuations in subterranean species (e.g. Bernabò
57 et al. 2011, Mermillod-Blondin et al. 2013, Novak et al. 2014, Rizzo et al. 2015) as well as their low dispersal
58 capacity (Sánchez-Fernández et al. 2016), ongoing human-driven climatic shifts are also expected to affect
59 subterranean biocoenosis and ecosystems (Wynne et al. 2014). From a conservation perspective, the
60 ecological and biogeographical peculiarity of the subterranean fauna enhances the risk of local extinction
61 (Cardoso 2011a, b, 2012) and strengthens the importance of considering subterranean organisms as model
62 species for ecological studies dealing with climatic changes (Chevaldonné and Lejeune 2003, Wynne et al.

63 2014, Mammola et al. 2015b).

64 We investigated the response of hypogean specialized spiders to increases in temperature in the Western
65 sector of the Alps. Among the range of subterranean taxa occurring in the Western Alps, we chose the cave-
66 dwelling spiders *Troglohyphantes* (Araneae, Linyphiidae) as model organisms. The reliability of this model
67 comes from empirical observations pointing out their preference for narrow, stable and cool microclimatic
68 conditions (Deeleman-Reinhold 1978, Isaia and Pantini 2010, Isaia et al. 2011, 2016, Mammola and Isaia
69 2016). Accordingly, *Troglohyphantes* spiders are characterized by moderate to high levels of
70 *troglobiomorphy* (*sensu* Juberthie and Decu 1994) and show restricted or point-like ranges of distribution
71 (e.g. Brignoli 1971, Deeleman-Reinhold 1978, Isaia and Pantini 2010, Isaia et al. 2011). On top of that,
72 populations of *Troglohyphantes* are usually extremely structured from a genetic point of view, lacking shared
73 haplotypes between caves and showing highly reduced dispersal ability (Mammola et al. 2015a).

74 We modelled presence/absence of *Troglohyphantes* spiders in relation to cave microclimate, past glacial
75 dynamics and other cave features in order to assess their sensitivity to potential subterranean climatic
76 variation induced by climate change.

77 Because constant temperatures recorded inside caves provide an approximation of the mean annual
78 temperature outside (Moore and Sullivan 1964, Smithson 1991, Badino 2004, 2010), in a second step we
79 extend the results to a bigger dataset. We used Ecological Niche Modeling (ENM) to extend to a wider scale
80 the response of our model organism to climatic changes, estimating the impact in terms of loss of habitat
81 suitability. In particular, we hypothesized i) a significant relationship between the occurrence of
82 *Troglohyphantes* spiders and different climatic conditions as a function of their adaptation to narrow and
83 specific ranges of temperature; and ii) that climate affected the distribution of our model species in the past
84 and will affect it in the future, through a decline in habitat suitability.

85

86 **Materials and methods**

87 **Study area**

88 We studied caves in the administrative regions of Piemonte and Valle d'Aosta (Western Alps, Italy), a
89 coherent biogeographic area of the Alps in which the distribution of *Troglohyphantes* spiders has been well
90 documented (e.g. Isaia and Pantini 2010, Isaia et al. 2011, 2016). We used a regional scale for analyses
91 instead of smaller or larger scales, as the former has been shown to be appropriate for species-level
92 responses to global warming (Walther et al. 2002). The Western Alps contain more than 2,500 caves
93 developed in carbonate and non-carbonate substrates (AGSP, 2016), plus a variety of artificial hypogean
94 habitats such as bunkers and abandoned mines.

95 We selected 33 hypogean localities over a north-south range in the Western Alps (Fig. 1), from the Tanaro
96 Valley (South; Province of Cuneo) to the Strona Valley (North; Province of Verbania). These included karst
97 and tectonic caves as well as artificial subterranean sites, with an altitudinal range of 415–2,357 m a.s.l, a
98 planimetric development range of 7– 3,440 m and an ascending/descending structure (i.e. drop) range from
99 –143 to +184 m. We deliberately choose localities with an established, recently and well-documented spider
100 fauna (see Isaia et al. 2011). A description of each site is provided in Supplementary Material Appendix 1.

101

102 **Sampling design**

103 Temperature (T) was monitored from 2012 to 2014 using Hygrochron™ devices (accuracy: ± 0.5 °C) placed
104 inside the cave (T_{int}) and at cave entrances (T_{ext}). Measurements were taken every 3 hours and the devices
105 were replaced after 12 months to allow download of data part way through the study.

106 The presence of *Troglohyphantes* at each site was verified during all surveys. The final presence/absence
107 dataset was verified and integrated with available literature data (Isaia and Pantini 2010, Isaia et al. 2011,
108 2016), wherever possible.

109

110 **Fine-scale modeling**

111 Data from the Hygrochron™ devices were used to generate values of mean annual temperature, daily and
112 annual range at the cave entrance (T_{ext}) and at the deeper zone (T_{int}). Sites were sorted according to
113 previous glacial history (ICE), as identified through an in-field geological survey by Motta (2014) or by using
114 paleo-glacial reconstruction for sites not included in the Motta study (Ehlers et al. 2011)—see Supplementary

115 Material Appendix 1. The relationship between temperature variables was explored using linear regression
116 models—*lm* command in the "stats" R package (R Development Team 2015). Presence/absence of spiders
117 was modelled as a function of temperature and the glacial history (ICE) via generalized linear models (*glm*
118 command in the "stats" R package) assuming a Bernoulli error structure (0–1 discrete). Prior to model fitting,
119 we tested for multi-collinearity among covariates and excluded collinear variables (Zuur et al. 2010). We
120 tested for spatial autocorrelation among environmental variables and in species occurrence via Moran's I test
121 in the "ape" R package (Paradis et al. 2004), using the Gittleman and Kot (1990) method. After fitting the
122 initial model, we carried out model reduction by sequentially deleting terms—backward elimination—
123 according to AICc values (Hurvich and Tsai 1989). The process was repeated until a minimum adequate
124 model of fixed effects remained (Zuur et al. 2009).

125

126 **Wide scale modeling, past and future predictions**

127 Once the model had been fitted using the fine-scale dataset, we extended the detected trends to a wider
128 scale (regional) using ecological niche modeling. We modelled the current distribution of *Troglohyphantes* in
129 the study area, transferring the results into ancestral (Last Glacial Maximum; ~22,000 years ago; hereinafter
130 LGM) and future (2050 and 2070) climatic conditions in order to estimate variations in the distribution ranges
131 relative to past and future climate changes.

132

133 *Occurrence points*

134 Geo-referenced occurrence localities of *Troglohyphantes* spp. were taken from Isaia et al. (2011). This
135 dataset covers 361 caves and virtually all possible thermal ranges experienced by *Troglohyphantes* spiders
136 in the Western Alps. *T. lucifuga* and *T. lucifer*, which are not exclusively restricted to subterranean habitats
137 (Isaia and Pantini 2010, Isaia et al. 2016), were excluded from our analysis. Prior to model fitting, we
138 designated a sampling bias grid (Phillips et al. 2009, Syfert et al. 2013). Within each cell in the grid—which
139 had a width equal to the grain of the bioclimatic rasters—all the points apart from one were randomly
140 removed (see Newbold 2010). In this way, we minimised spatial autocorrelation, given that proximate
141 localities likely comprise similar environmental variables.

142

143 *Explanatory variables*

144 We represented present-day climate with 19 bioclimatic variables (1950–2000; Hijmans et al. 2005;

145 resolution: 30"; available at: www.worldclim.org; Table 1), and analogous data layers for past and future
146 climate. For both past and future climate, multiple GMCs were used to take into account variation and
147 uncertainty among simulations (Kageyama et al. 2001, Beaumont et al. 2008, Diniz-Filho et al. 2009). Paleo-
148 bioclimatic variables for the LGM were derived from simulations available from three General Circulation
149 Models (GCMs) from the Coupled Model Intercomparison Project Phase 5 (CMIP5; Taylor et al. 2012).
150 These were CCSM4, MIROC-ESM and MPI-ESM-P (resolution of 2.5'). Future data represent downscaled
151 and calibrated climate projections from three GCMs (CCSM4, BCC-CSM1-1, MIROC5) under two
152 Representative Concentration Pathway scenarios (rcp 2.6 and rcp 8.5 respectively) for both 2050 and 2070,
153 at a resolution of 30". Ice cover during Pleistocene was derived from paleo-reconstruction, by converting the
154 shape of glacial masses provided by Ehlers et al. (2011) into a binary raster. This categorical variable was
155 further included in the models. Prior to model computing, we investigated multi-collinearity among covariates
156 via Pearson r correlation (Zuur et al. 2009, 2010), selecting a final set of uncorrelated bioclimatic variables
157 ($r < \pm 0.7$) (Table 1).

158

159 *Algorithm choice and projections*

160 A wide suite of algorithms exists in the realm of bioclimatic envelope modeling (Peterson et al. 2011).
161 Considering the conceptual ground of the "no free lunch" theorem (Ho and Pepyne 2002), Qiao et al. (2015)
162 recently suggested that no single best algorithm should be chosen *a priori* for modeling the ecological
163 niche/species distribution of a certain organism. We thus began the analysis of this particular dataset with an
164 assessment of performance of potential competing algorithms (see also Elith et al. 2006). In particular, we
165 tested the performance of three well-acknowledged statistical techniques, belonging to the three existing
166 categories of algorithms for computing ENM/SDM: regression [Generalized linear models (GLM); McCullagh
167 and Nelder 1989], regression trees [Boosted Regression Trees (BRT); Friedman, 2001] and machine-
168 learning [Maximum entropy model (MaxEnt); Philips et al. 2006]. Details of the modeling procedure and
169 settings for each technique are provided in Supplementary Material Appendix 2.

170 As is widely recommended in the literature (e.g. Peterson et al. 2011, Saupe et al. 2012, Merow et al. 2013),
171 models were calibrated within the accessible area—often referred to as area **M** (Barve et al. 2011; details in
172 Supplementary Material Appendix 2). **M** is the geographic extent hypothesized to fall within the long-term
173 dispersal and colonization potential for a particular species over its evolutionary history. Considering the
174 reduced dispersal potential of *Troglohyphantes* spiders, in our case area **M** was approximated by masking
175 the bioclimatic layers with a rectangle delimiting the Western alpine chain.

176 We computed the models using the three modelling techniques in order to relate the occurrence points to the
177 set of non-collinear variables. For each technique, we ran twenty replicate of models via a loop in R, keeping
178 a random partition of 20% of the points for each run to assess the predictive ability of each algorithm.
179 Predictive abilities of the three models were obtained by calculating two different metrics (average of the
180 twenty runs): the area under the Receiver Operating Characteristics (ROC) curve (AUC; Fielding and Bell
181 1997) and the True Skill Statistics (TSS; Allouche et al. 2006), whereby an AUC < 0.60 and/or a TSS < 0.1
182 denotes a predictive ability no better than random. We employed two metrics because it is not recommended
183 to rely on a single measure of model fit for comparing different techniques (Elith and Graham 2009).

184 We finally retained the modeling technique showing the best performance. A final model was generated
185 using the same parameters and calibrated with the complete occurrence dataset.

186 Subsequently, we transferred the model into each LGM and future GCMs, and we calculated the median
187 value across all projection for each combination of GCM. We estimated environmental suitability for each
188 projection by rescaling raster cell values above the threshold of 0.4 (see discussion in Liu et al. 2005)
189 between 0 and 1. Areas below a threshold value of 0.4 were set to 0.

190

191

192

193 **Results**

194 **Fine-scale modelling**

195 As expected, we observed a strong linear relationship between mean annual T_{int} and mean annual T_{ext}
196 (Estimated $\beta \pm se$: 0.901 ± 0.041 , $p < 0.001$), with only 3 out of 33 caves showing T_{int} values 1 °C higher (Grotta
197 di Bergovei and Fessura di Verrogne) or lower (Grotta Testa di Napoleone) than T_{ext} values (Supplementary
198 material Appendix 2, Fig. A1).

199 No variables considered in the GLM analysis were found to be spatially autocorrelated (Moran's I test, all
200 $p > 0.05$) apart from the mean annual T_{int} range ($p < 0.01$) which therefore was not considered further. Moran's I
201 test results are reported in the Supplementary Material Appendix 2, Table A1. Daily range of T_{int} was highly
202 collinear with past ice cover (ICE) and thus excluded from the analysis (Supplementary material Appendix 2,
203 Fig. A2). All variables related to T_{ext} (mean and range) were highly collinear with T_{int} (Pearson $r > \pm 0.7$), and
204 thus not included in the model (Supplementary material Appendix 2, Table A2). Backward elimination

205 revealed that variables related to cave features—type of cave, total development, drop—were not influential
206 in determining the presence/absence of *Troglohyphantes* and these were therefore not considered in further
207 analytical steps. Accordingly, the structure of the minimum adequate model included mean annual T_{int} and
208 the past glacial cover (ICE) as covariates. Outcomes of the GLM pointed to a significant decrease in the
209 probability of occurrence of *Troglohyphantes* with increasing mean annual T_{int} values (Estimated $\beta \pm se$: -
210 0.543 ± 0.249 , $p=0.02$). We also found a significant negative effect of glacial history, namely a lower
211 probability of occurrence of the model organisms in caves covered by the ice during past glaciations (ICE,
212 test relative to the level "uncovered"; Estimated $\beta \pm se$: -2.584 ± 1.052 , $p=0.01$; Fig. 2).

213

214 **Model on the regional dataset, past and future predictions**

215 MaxEnt performed better than other algorithms tested in terms of predictive ability (MaxEnt: mean $AUC \pm sd$ =
216 0.873 ± 0.09 ; mean $TSS \pm sd$ = 0.51 ± 0.08 ; GLM: $AUC \pm sd$ = 0.676 ± 0.378 ; mean $TSS \pm sd$ = 0.32 ± 0.221 ; BRT:
217 $AUC \pm sd$ = 0.78 ± 0.02 ; $TSS \pm sd$ = 0.43 ± 0.12). We therefore retained the former technique for further analysis
218 and projections. Uncorrelated variables introduced in the final model are reported in Table 1. The final model
219 included mean annual temperature (Bio01), mean diurnal range (Bio02), and ice cover (ICE)—response
220 plots are reported in Supplementary Material Appendix 2, Fig. A3. Suitable areas predicted by the model for
221 the current climate (Fig. 3) roughly overlay the known distribution of *Troglohyphantes* in the study area, with
222 higher suitability predicted in the low-medium mountain belt, ranging from the districts of Maritime to Graian
223 Alps. The most suitable and unfragmented habitat is found at the southern border of the Alps. Based on the
224 model projections, the average probability of occurrence is 0.68 across all known populations of
225 *Troglohyphantes*.

226 With regard to LGM projections (Fig. 3), the district of Maritime Alps (South-western Alps) and the
227 surrounding plains appear to contain the most suitable areas of habitat. Areas covered by ice were predicted
228 to be largely unsuitable, with areas bordering the glacial masses possibly acting as microrefugia (*sensu* Rull
229 2009).

230 Future forecasts based on different emission scenarios indicate a general decline of suitability throughout the
231 distribution range (Fig. 4), with major effects at higher emissions. For 2050 we predict a general decrease in
232 suitable areas (threshold 0.4) of ~ 45% for the low emission (rcp 2.6) and of ~ 55% for the high emission
233 scenario (rcp 8.5) in respect to the present-day projection. Habitat loss is further intensified in the 2070 high-
234 emission scenario, with an overall reduction of ~ 70%. The predicted loss of suitable habitat appears to be

235 more intense at lower altitudes. In particular, for the high-emission scenario in 2070 approximately half of the
236 current localities are predicted to be unsuitable (Fig. 5).

237 Across all known populations, the average decrease in the probability of occurrence of *Troglohyphantes*
238 drops from 0.68 to 0.51 in the low emission scenario (both 2050 and 2070). For the high emission scenario,
239 the probability of occurrence drops to 0.40 for 2050 and 0.17 for 2070.

240 **DISCUSSION**

241 **Caves as laboratories for the study of climate change**

242 In contrast with fluctuating surface temperatures, temperature stability in the subterranean domain allows the
243 detection of temperature changes over longer time periods. Indeed, this working hypothesis provides a
244 stimulating starting point for the ecological modeling of hypogean ecosystems. In spite of that, correlative
245 distribution modelling techniques have been rarely—and only recently—applied to these particular biological
246 systems (e.g., Bryson et al. 2014, Naranjo et al. 2014, Fišer et al. 2015, Mammola et al. 2015a, Soto-
247 Centeno et al. 2015, Christman et al. 2016, Copper-Bohannon et al. 2016, Ferreira et al. 2016, Sánchez-
248 Fernández et al. 2016).

249 However, the link between the external climate and the subterranean habitat has been well documented in
250 speleological literature (Moore and Sullivan 1964, Smithson 1991, Badino 2004, 2010, Covington and Perne
251 2015). The general explanation for this relationship is found in the physics of the infiltrating meteoric fluids.
252 Over a geological timeframe, a given rock layer will be forced to the thermal equilibrium of the water flowing
253 through it. As a consequence, the temperature of the rock surrounding the cave equilibrates to the mean
254 temperature of the infiltrating waters, which, in turn, is equal to the mean temperature outside (Badino 2010).
255 Depending on the thickness of the rock layer (i.e. the cave depth), the amplitude of the sinusoidal input
256 signal (i.e. the external temperature range) can be shifted in time—i.e. there is a time lag between input and
257 output maxima (Badino 2004), but this lag is likely to have a duration of decades. As an example, for the
258 Postojna cave (Slovenia) Domínguez-Villar et al. (2015) estimated a warming time lag of ~ 20 years, which is
259 orders of magnitude less time than the glacial time frames that are the focus of the study. In practical terms,
260 this implies that our future projections are mostly relevant for the superficial subterranean habitats—including
261 the outermost sectors of caves—whereas they should be shifted in time of at least 20 years to obtain a more
262 precise time scale for future habitat modifications in deep hypogean layers.

263 In an evolutionary sense, such a time lag is not large, but in an anthropogenic sense it might be significant.
264 As a consequence, spiders that we see today may still be in the lag phase i.e. not yet fully adapted to the
265 current temperature.

266 We note that, in general, correlative distribution models prioritize the use of extreme values—maxima and
267 minima—instead of average values. Whilst extremes are often believed to reflect a realistic relationship with
268 species-specific physiological constraints (Hijmans et al., 2005), in the case of caves with constant
269 temperatures and reduced temperature excursions, average values are likely to be more appropriate. Our

270 use of average values thus likely increases the robustness of our results because GCMs models are better
271 at predicting average than extreme values (but see Zimmerman et al. 2009).

272

273 **Interpreting the effects of past climate change**

274 Two contrasting theories attempt to explain the origin and the distribution of the subterranean fauna (Culver
275 and Pipan 2009, 2010). The theory of the active colonization (Rouch and Danielopol 1987) or adaptive shift
276 hypothesis (Howarth, 1980) puts great emphasis on the process of active colonization of the hypogean
277 domain, with species being driven by the opportunity to occupy new, unexploited ecological niches. On the
278 other hand, the theory of relicts and refuges (Botosaneanu and Holsinger 1991) invokes long-term climatic
279 changes, such as glaciation cycles and other large-scale climate upheavals, as the main factors that
280 prompted the colonization of the subterranean habitat and causing the obliteration of surface-dwelling
281 populations (Holsinger 1988, Botosaneanu and Holsinger 1991).

282 Although this dichotomous vision exists, the case of our model species possibly supports the latter theory of
283 relict and refuges, given that the occurrence of *Troglohyphantes* appears to be significantly linked to the
284 glacial history. This pattern can be explained by taking into account the self-evident fact that, most likely,
285 natural populations do not survive in caves filled with ice (Culver and Pipan 2010). Remarkably, in our case
286 only two sister species (*T. lucifuga* and *T. lucifer*) were found to occur in caves that were covered by ice
287 during the Pleistocene. *T. lucifuga* is a common and widespread species in the North Western Alps showing
288 a low degree of troglobiomorphism, whereas the recently described *T. lucifer* represents its ecological
289 vicariant in the nearby alpine districts of Cottian and Southern Graian Alps (Isaia et al. 2016). Both species
290 occur in epigean as well as hypogean habitats (Isaia and Pantini 2010, Isaia et al. 2016), an occurrence that
291 could be explained by *T. lucifuga* and *T. lucifer* having colonized the area after the glaciers retreated.

292

293 **Global change in hypogean ecosystems**

294 Little attention has been given to the effects of global climate change on hypogean organisms and
295 ecosystems relative to their epigean counterparts. In this study, we show the extent to which cave-dwelling
296 biological communities are likely to be affected by changes in temperature. Specialized hypogean arthropods
297 are known to have a fine-tuned thermal tolerance to the constant and narrow temperature ranges of the
298 subterranean habitat (Delay 1978, Huevy and Kingsolver 1989, Rizzo et al. 2015). This extends even to a
299 greater reduction in thermal tolerance observed in species living in the buffered portion of caves compared

300 with those living in the vicinity of the cave entrance (Lencioni et al. 2010, Bernabò et al. 2011, Novak et al.
301 2014).

302 However, from a global change perspective, such an adaptation turns out to be a strong limitation. Indeed,
303 while most invertebrates living close to the surface—troglaphiles—have retained their ability to withstand
304 temperature variations, many specialized hypogean organisms—troglbionts—have lost such
305 thermoregulatory mechanisms (Novak et al. 2014).

306 Modelling at both local and regional spatial scales indicates that subterranean *Troglohyphantes* spiders are
307 restricted to particular climatic environments, occurring only in caves with annual mean temperatures below
308 10–11°C. This finding implies that in most mid-altitude areas of the W-Alps, where annual mean
309 temperatures are around 10–12°C, populations of *Troglohyphantes* are close to their physiological limit. Due
310 to predicted temperature rises (IPCC 2014), it is therefore expected that large areas of currently suitable
311 habitat will become progressively unsuitable, therefore falling outside the climatic niche of our model
312 species.

313 Our results, indeed, point to a likely dramatic decline in habitat suitability (Fig. 5) for most of the species
314 included in the analysis—*T. bornensis*, *T. iulianae*, *T. konradi*, *T. lanai*, *T. nigraerosae*, *T. pedemontanus*,
315 and *T. pluto*. On the other hand, a higher chance of survival is expected in *T. vignai* because it shows a
316 wider distribution and is also recorded at higher altitudes (above 2,000 m; Isaia et al. 2011, Mammola et al.
317 2015a). In addition to troglbiomorphic species being restricted to areas of temperature stability, caves are
318 often isolated from each other and species within them, including spiders, often exhibit reduced vagility (e.g.,
319 Hedin 1997, Snowman et al. 2010, Yao et al. 2016). In turn, this latter factor significantly increases the threat
320 of global extinction because the rate of recolonisation is low (Cardoso et al. 2011a, b). This is particularly
321 evident for *Troglohyphantes* spiders, for which we previously demonstrated a very low dispersal ability within
322 caves (Mammola and Isaia 2016) and extreme levels of population differentiation (Mammola et al. 2015a).

323 It is interesting to note that in the case of other troglbiomorphic species, the assessment of global warming
324 impacts on subterranean species based exclusively on correlative niche models is suggested to be
325 unreliable if thermal tolerance is wider than that reflected by current distributions (Sánchez-Fernández et al.
326 2016). Physiological mechanisms that might account for such an observation include the ability to synthesize
327 stress proteins (e.g. heat shock proteins; Bernabò et al. 2011). Whilst the data for these particular species
328 (Leiodidae beetles) are compelling, the finding seems unlikely to apply more generally to the glacial relict,
329 cold-adapted spiders in our study. These have a rather different evolutionary history to that of the beetles
330 and are not shown to be able to over-express similar, protective proteins. It is clear, however, that an intrinsic

331 degree of uncertainty remains associated with interpreting the results of any statistical model of this kind.
332 Accordingly, as a follow up of this study, we aim to combine multiple lines of evidence (see Peterson 2009)—
333 such as thermal tolerance experiments and population genetic studies—in order to explore our predicted
334 patterns further (see Krehenwinkel et al. 2015 for an example on spiders). Testing the physiological
335 response of subterranean model species to altered climatic condition could help to identify their specific
336 upper thermal limits, and thus provide experimental evidence for our findings. In parallel, molecular analysis
337 could establish contemporary population structure and dispersal modes, thereby increasing our
338 understanding of the ability of subterranean species to escape and survive climate changes.

339

340 **Conclusions**

341 Together with other species living in confined habitats such as islands (e.g., Bellard et al. 2014, Yalor and
342 Kumar 2016) and mountain summits (e.g., Beniston 2003, Xu et al. 2009, Yoccoz et al. 2010), we expect
343 cave adapted species to be as vulnerable, if not more so, to climate change. However, the principal
344 mechanisms underlying the response of cave-dwelling species to global climate change are as yet poorly
345 described. Caves are excellent natural laboratories for the study of global changes in environment because
346 of their remarkable ecological and thermal stability. Our findings emphasise the importance of considering
347 subterranean organisms as model species for ecological studies dealing with climatic changes, and to
348 extend such investigations to other subterranean systems worldwide.

349

350 **Acknowledgments**

351 This work is part of the work package 6 (WP6) of the CAVELAB project “From microclimate to climate
352 change: caves as laboratories for the study of the effects of temperature on ecosystems and biodiversity”,
353 funded by Compagnia di San Paolo and University of Turin—Progetti di Ricerca di Ateneo 2011, cod.
354 ORTO11T92F. We are indebted to Giovanni Badino for providing physical argumentations about
355 subterranean meteorology and to all the people involved in CAVELAB. A special thank goes to Mauro
356 Paschetta and Enrico Lana, for helping us place and retrieve several I-button devices. We warmly thank the
357 graphic designer Irene Frigo, for helping us to create the graphical abstract. We are grateful to Lisa Comte,
358 Boris Leroy and an anonymous reviewer for their help in improving the quality of the manuscript through their
359 constructive comments during the review process.

360

361 **Supplementary materials**

362 **Appendix 1** - Raw data of the fine-scale dataset, with details about sampling sites.

363 **Appendix 2** - Details on modeling procedure. Additional figures and tables.

364

365 **Literature cited**

366

367 Allouche, O. et al. 2006. Assessing the accuracy of species distribution models: prevalence, kappa and the
368 true skill statistic (TSS). – J. Appl. Ecol. 43: 1223–1232.

369

370 AGSP 2016. Catasto speleologico del Piemonte e della Valle D'Aosta. – Associazione Gruppi Speleologici
371 Piemontesi, Italy, < <http://sellarenato.interfree.it> >.

372

373 Badino, G. 2004. Cave temperature and Global Climatic Change. – Int. J. Speleol. 33: 103–114.

374

375 Badino, G. 2010. Underground meteorology. What's the weather underground? – Acta Carsol. 39: 427– 448.

376

377 Barve, N. et al. 2011. The crucial role of the accessible area in ecological niche modeling and species
378 distribution modeling. – Ecol. Model. 222: 1810–1819.

379

380 Beaumont, L. J. et al. 2008. Why is the choice of future climate scenarios for species distribution modelling
381 important? – Ecol. Lett. 11: 1135–1146.

382

383 Bellard, C. et al. 2012. Impacts of climate change on the future of biodiversity. – Ecol. Lett. 15: 365–377.

384

385 Bellard, C. et al. 2014. Impact of sea level rise on the 10 insular biodiversity hotspots. – Global Ecol.
386 Biogeogr. 23: 203–212.

387

388 Beniston, M. 2003. Climatic change in mountain regions: a review of possible impacts. –Springer. [page
389 numbers]

390

391 Bernabò, P. et al. 2011. Cold stenothermal cave-dwelling beetles do have an HSP70 heat shock response. –
392 J. Therm. Biol. 36: 206–208.

393

- 394 Botosaneanu, L. and Holsinger, J. 1991. Some aspects concerning colonization of the subterranean realm -
395 especially subterranean waters: a response to Rouch and Danielopol, 1987. – *Stygologia* 6: 11–39.
396
- 397 Brandmayr, P. et al. 2013. Hypogean carabid beetles as indicators of global warming? – *Env. Res. Lett.* 8:
398 1–11.
399
- 400 Brignoli, P. M. 1971. Note su ragni cavernicoli italiani (Araneae). – *Fragm. Entomol.* 7: 129–229.
401
- 402 Bryson, Jr R. W. et al. 2014. Caves as microrefugia: pleistocene phylogeography of the troglomorphic North
403 American scorpion *Pseudouroctonus reddelli*. – *BMC Evol. Biol.* 14: 9.
404
- 405 Cardoso, P. 2012. Diversity and community assembly patterns of epigean vs. troglobiont spiders in the
406 Iberian Peninsula. – *Int. J. Speleol.* 41(1): 83–94.
407
- 408 Cardoso, P. et al. 2011a. Adapting the IUCN red listing criteria for invertebrates. – *Biol. Conserv.* 144: 2432–
409 2440.
410
- 411 Cardoso, P. et al. 2011b. The seven impediments in invertebrate conservation and how to overcome them. –
412 *Biol. Conserv.* 144: 2647–2655.
413
- 414 Chen, I. C. et al. 2011. Rapid range shifts of species associated with high levels of climate warming. –
415 *Science* 333(6045): 1024–1026.
416
- 417 Chevaldonné, P. and Lejeune, C. 2003. Regional warming-induced species shift in northwest Mediterranean
418 marine caves. – *Ecol. Lett.* 6: 371–379.
419
- 420 Christman, M. C. et al. 2016. Predicting the occurrence of cave-inhabiting fauna based on features of the
421 earth surface environment. – *PLoS ONE* 11: e0160408.
422
- 423 Cooper-Bohannon, R. et al. 2016. Predicting bat distributions and diversity hotspots in southern Africa. –

424 Hystrix 27.

425

426 Covington, M. D. and Perne, M. 2015. Consider a cylindrical cave: a physicist's view of cave and karst
427 science. - Acta Carsol. 44: 363–380.

428

429 Culver, D.C. and Pipan, T. 2009. The Biology of Caves and other Subterranean Habitats. – Oxford Univ.
430 Press.

431

432 Culver, D.C. and Pipan, T. 2010. Climate, abiotic factors, and the evolution of subterranean life. Acta Carsol.
433 - 39: 539–577.

434

435 Dawson, T. P. et al. 2011. Beyond predictions: biodiversity conservation in a changing climate. – Science
436 332: 53–58.

437

438 Deeleman-Reinhold, C. L. 1978. Revision of the cave-dwelling and related spiders of the genus
439 *Troglohyphantes* Joseph (Linyphiidae), with special reference to the Yugoslav species. Opera
440 Academia Scientiarum et Artium Slovenica.

441

442 Delay, B. 1978. Milieu souterrain et écophysiological de la reproduction et du développement des
443 Coléoptères Bathysciinae hypogés. – Mém. Biospéol. 5: 1–349.

444

445 Diniz-Filho, J. A. F. et al. 2009. Partitioning and mapping uncertainties in ensembles of forecasts of species
446 turnover under climate change. – Ecography 32: 897–906.

447

448 Domínguez-Villar, D. et al. 2015. Is global warming affecting cave temperatures? Experimental and model
449 data from a paradigmatic case study. – Clim. Dynam. 45: 569–581.

450

451 Ehlers, J. et al. 2011. Quaternary glaciations - extent and chronology. A closer look. Vol. 15. – Elsevier.
452 [chapter & page numbers?]

453

454 Elith, J. and Graham, C. H. 2009. Do they? How do they? WHY do they differ? On finding reasons for

455 differing performances of species distribution models. – *Ecography* 32: 66–77.

456

457 Elith, J. et al. 2006. Novel methods improve prediction of species' distributions from occurrence data. –
458 *Ecography* 29: 129–151.

459

460 Ferreira, M. I. et al. 2016. The genus *Zelurus* Hahn, 1826, in Brazilian caves: description of new species and
461 comments on the potential distribution of the genus in South America. – *Zootaxa* 4170: 250–270.

462

463 Fejér, A. and Moldovan, O. T. 2013 Population size and dispersal patterns for a *Drimeotus* (Coleoptera,
464 Leiodidae, Leptodirini) cave population. – *Subterr. Biol.* 11: 31–44.

465

466 Fielding, A. H. and Bell, J. F. 1997. A review of methods for the assessment of prediction errors in
467 conservation presence/absence models. – *Environ. Cons.* 24: 38–49.

468

469 Fišer, Ž. et al. 2015. Morphologically cryptic amphipod species are “ecological clones” at regional but not at
470 local scale: a case study of four *Niphargus* species. – *PloS ONE* 10(7): e0134384.

471

472 Friedman, J. H. 2001. Greedy function approximation: a gradient boosting machine. – *Ann. Stat.* 29: 1189–
473 1232.

474

475 Gittleman, J. L. and Kot, M. 1990. Adaptation: statistics and a null model for estimating phylogenetic effects.
476 – *Syst. Zool.* 39: 227–241.

477

478 Hedin, M. C. 1997. Molecular phylogenetics at the population/species interface in cave spiders of the
479 southern Appalachians (Araneae: Nesticidae: *Nesticus*). – *Mol. Biol. Evol.* 14: 309–324.

480

481 Hijmans, R. J. et al. 2005. Very high resolution interpolated climate surfaces for global land areas. – *Int. J.*
482 *Climatol.* 25: 1965–1978.

483

484 Ho, Y. C. and Pepyne, D. L. 2002. Simple explanation of the no-free-lunch theorem and its implications. – *J.*

485 Optim. Theory Appl. 115: 549–570.

486

487 Holsinger, J. R. 1988. Troglobites: the evolution of cave-dwelling organisms. – Am. Sci. 76: 147–153.

488

489 Howarth, F. G. 1980. The Zoogeography of specialized cave animals: a bioclimatic model. - Evolution 34:
490 394–406.

491

492 Huevy, R. B. and Kingsolver, J. G. 1989. Evolution of thermal sensitivity of ectotherm performance. – Trends
493 Ecol. Evol. 4: 131–135.

494

495 Hurvich, C. M. and Tsai C. L. 1989. Regression and time series model selection in small samples. –
496 Biometrika 76: 297–307.

497

498 IPCC 2014. Climate Change 2014: Synthesis Report. Contribution of Working Groups I, II and III to the Fifth
499 Assessment Report of the Intergovernmental Panel on Climate Change. – IPCC.

500

501 Isaia, M. and Pantini, P. 2010. New data on the spider genus *Troglohyphantes* (Araneae, Linyphiidae) in the
502 Italian Alps, with the description of a new species and a new synonymy. – Zootaxa 2690: 1–18.

503

504 Isaia, M. et al. 2011. Subterranean arachnids of the Western Italian Alps (Arachnida: Araneae, Opiliones,
505 Palpigradi, Pseudoscorpiones). – Mon. Mus. Reg. Sc. Nat.

506

507 Isaia, M. et al. 2016. Advances in the systematics of the spider genus *Troglohyphantes* (Araneae,
508 Linyphiidae) - Syst. Biodivers.[page numbers]

509

510 Juberthie, C. and Decou, V. 1994. Structure et diversité du domaine souterrain; particularités des habitats et
511 adaptations des espèces. – Encycl. Biospeol. 1: 5–22.

512

513 Kageyama, M. et al. 2001. The Last Glacial Maximum climate over Europe and western Siberia: a PMIP
514 comparison between models and data. – *Clim. Dynam.* 17: 23–43.

515

516 Krehenwinkel, H. et al. 2015. Eco-Genomic analysis of the poleward range expansion of the wasp spider
517 *Argiope bruennichi* shows rapid adaptation and genomic admixture. – *Glob. Change Biol.* 21: 4320–
518 4332.

519

520 Lencioni, V. et al. 2010. Cold resistance in two species of cave-dwelling beetles (Coleoptera: Cholevidae). –
521 *J. Therm. Biol.* 35: 354–359.

522

523 Liu, C. et al. 2005. Selecting thresholds of occurrence in the prediction of species distributions. – *Ecography*
524 28: 385–393.

525

526 Mammola, S. and Isaia, M. 2016. The ecological niche of a specialized subterranean spider. – *Invertebr.*
527 *Biol.* 135: 20–30.

528

529 Mammola, S. et al. 2015a. Alpine endemic spiders shed light on the origin and evolution of subterranean
530 species. – *PeerJ* 3: e1384.

531

532 Mammola, S. et al. 2015b. Seasonal dynamics and micro-climatic preference of two Alpine endemic
533 hypogean beetles. – *Int. J. Speleol.* 44: 239–249.

534

535 Mammola, S. et al. 2016. Ecology and sampling techniques of an understudied subterranean habitat: the
536 Milieu Souterrain Superficiel (MSS). – *The Science of Nature* 103(11-12): 88.

537

538 McCullagh, P. and Nelder, J. A. 1989. *Generalized Linear Models* (2nd ed.). – Chapman and Hall.

539

540 Mermillod-Blondin, F. et al. 2013. Thermal tolerance breadths among groundwater crustaceans living in a

541 thermally constant environment. – J. Exp. Biol. 216: 1683–1694.

542

543 Merow, C. et al. 2013. A practical guide to MaxEnt for modeling species' distributions: what it does, and why
544 inputs and settings matter. – *Ecography* 36: 1058–1069.

545

546 Moore, G. W. and Sullivan, G. N. 1964. Out of phase seasonal temperature fluctuations in Cathedral Cave,
547 Kentucky. – *Geol. Soc. Am.* 76: 313.

548

549 Motta, M. 2014. The definition of the extension of quaternary glaciers within alpine valleys, and his
550 application to study of troglobites. – *EDIS* 1: 439–444.

551

552 Naranjo, M. et al. 2014. ¿Dónde buscar troglobiontes? Ensayo de una cartografía predictiva con MaxEnt en
553 Gran Canaria (islas Canarias). – *Arxius de Miscel·lània Zoològica* 12: 83–92.

554

555 Newbold, T. 2010. Applications and limitations of museum data for conservation and ecology, with particular
556 attention to species distribution models. – *Prog. Phys. Geog.* 34: 3–22.

557

558 Novak, T. et al. 2014. Cold tolerance in terrestrial invertebrates inhabiting subterranean habitats. – *Int. J.*
559 *Speleol.* 43: 265–272.

560

561 Paradis, E. et al. 2004. APE: analyses of phylogenetics and evolution in R language. – *Bioinformatics* 20:
562 289–290.

563

564 Parmesa, C. and Yohe, G. 2003. A globally coherent fingerprint of climate change impacts across natural
565 systems. – *Nature* 421: 37–42.

566

567 Parmesan, C. 2006. Ecological and evolutionary responses to recent climate change. – *Annu. Rev. Ecol.*
568 *Evol. Syst.*: 637–669.

569

570 Peterson, A. T. 2009. Phylogeography is not enough: the need for multiple lines of evidence. – Front.
571 Biogeogr. 1: 19–25.

572

573 Peterson, A. T. et al. (2011) Ecological niches and geographical distributions: a modeling perspective. –
574 Princeton Univ. Press.

575

576 Phillips, S. J. et al. 2006. Maximum entropy modeling of species geographic distributions. – Ecol. Model.
577 190: 231–259.

578

579 Phillips, S. J. 2009. Sample selection bias and presence-only distribution models: implications for
580 background and pseudo-absence data. – Ecol. Appl. 19: 181–197.

581

582 Phillips, S. J. 2011. A brief tutorial on Maxent. – AT&T Research, Princeton.

583

584 Poulson, T. L. and White, W. B. 1969. The cave environment. – Science 165: 971–981.

585

586 Qiao, H. et al. 2015. No silver bullets in correlative ecological niche modelling: insights from testing among
587 many potential algorithms for niche estimation. – Methods Ecol. Evol. 6: 1126–1136.

588

589 R Development Core Team 2015. R: A language and environment for statistical computing. – R Foundation
590 for Statistical Computing, Austria, < <http://www.R-project.org/> >.

591

592 Rizzo, V. et al. 2015. Lack of evolutionary adjustment to ambient temperature in highly specialized cave
593 beetles. – BMC Evol. Biol. 15: 10.

594

595 Romero, A. 2009. Cave biology. – Cambridge Univ. Press.

596

597 Rouch, R. and Danielpol, D. 1987. L'origine de la faune aquatique souterraine, entre le paradigme du refuge
598 et le modèle de la colonisation active. – Stygologia 3: 345–372.

599

600 Rull, V. 2009. Microrefugia. – J. Biogeogr. 36: 481–484.

601

602 Sánchez-Fernández, D. et al. 2016. Thermal niche estimators and the capability of poor dispersal species to
603 cope with climate change. – Sc. Rep. 6: 23381.

604

605 Saupe, E. E. et al. 2012. Variation in niche and distribution model performance: the need for a priori
606 assessment of key causal factors. – Ecol. Model. 237: 11–22.

607

608 Smithson, A. 1991. Inter-relationships between cave and outside air temperatures. – Theor. Appl. Climatol.
609 44: 65–73.

610

611 Snowman, C. V. et al. 2010. Caves as islands: mitochondrial phylogeography of the cave-obligate spider
612 species *Nesticus barri* (araneae: Nesticidae). – J. Arachnol. 38: 49–56.

613

614 Soto-Centeno, J. A. et al. 2015. The importance of late Quaternary climate change and karst on distributions
615 of Caribbean mormoopid bats. – Am. Mus. Novit. 3847: 1–32.

616

617 Syfert, M. M. et al. 2013. The effects of sampling bias and model complexity on the predictive performance
618 of MaxEnt Species Distribution Models. – PLoS ONE 8: e55158.

619

620 Taylor, S. and Kumar, L. 2016. Global climate change impacts on pacific islands terrestrial biodiversity: a
621 review. – Trop. Conserv. Sci. 9: 203–223.

622

623 Taylor, K. E. et al. 2012. An overview of CMIP5 and the experiment design. – B. Am. Meteorol. Soc. 93:
624 485–49.

625

626 Walther, G. R. 2010. Community and ecosystem responses to recent climate change. – Phil. Trans. R. Soc.
627 Lond., B, Biol. Sci. 365: 2019–2024.

628

629 Walther, G. R. et al. 2002. Ecological responses to recent climate change. – Nature 416: 389–395.

630

631 Wynne, J.J. et al. 2014. Disturbance relicts in a rapidly changing world: The Rapa Nui (Easter Island) factor.

632 – Bioscience 64: 711–718.

633

634 Xu, J. et al. 2009. The melting Himalayas: cascading effects of climate change on water, biodiversity, and

635 livelihoods. – Conserv. Biol. 23: 520–530.

636

637 Yao, Z. et al. 2016. High endemism at cave entrances: a case study of spiders of the genus *Uthina*. – Sc.

638 Rep., 6: 35757.

639

640 Yoccoz, N. G. et al. 2010. Impact of climatic change on alpine ecosystems: inference and prediction. – J.

641 Alp. Res.: 98–104.

642

643 Zimmermann, N. E. et al. 2009. Climatic extremes improve predictions of spatial patterns of tree species. –

644 Proc. Natl. Acad. Sc. USA 106: 19723–19728.

645

646 Zuur, A. F. et al. 2010. A protocol for data exploration to avoid common statistical problem. – Methods Ecol.

647 Evol. 1: 3–14.

648

649 Zuur, A. F. et al. 2009. Mixed effect models and extensions in ecology with R. – Springer.

650

651

652

653

654

655

656

657

658

659

660

661

662

663 **Tables and Figures**

664 **Table 1.** List of the variables selected to construct the regional scale model. The uncorrelated variables
665 (Pearson $r < \pm 0.7$) used in the model are highlighted. Percent contribution and permutation importance
666 (*sensu* Phillips 2011) of each variable for building the final MaxEnt model is reported.

Variable	Description	Unit	Percent contribution	Permutation importance
Bio01	Annual Mean Temperature	°C	29.6 %	27.4
Bio02	Mean Diurnal Range	°C	32.7 %	32.5
Bio03	Isothermality (Bio02/Bio07) (*100)	°C	-	
Bio04	Temperature Seasonality (standard deviation*100)	°C	-	
Bio05	Max Temperature of Warmest Month	°C	-	
Bio06	Min Temperature of Coldest Month	°C	-	
Bio07	Temperature Annual Range (Bio05– Bio06)	°C		
Bio08	Mean Temperature of Wettest Quarter	°C	-	
Bio09	Mean Temperature of Driest Quarter	°C	Discarded (<5%)	Discarded
Bio10	Mean Temperature of Warmest	°C	-	

Variable	Description	Unit	Percent contribution	Permutation importance
	Quarter			
Bio11	Mean Temperature of Coldest Quarter	°C	-	
Bio12	Annual Precipitation	mm	-	
Bio13	Precipitation of Wettest Month	mm	-	
Bio14	Precipitation of Driest Month	mm	-	
Bio15	Precipitation Seasonality (Coefficient of Variation)	mm	-	
Bio16	Precipitation of Wettest Quarter	mm	-	
Bio17	Precipitation of Driest Quarter	mm	-	
Bio18	Precipitation of Warmest Quarter	mm	Discarded (<5%)	Discarded
Bio19	Precipitation of Coldest Quarter	mm	-	
Alt	Altitude a.s.l	m	-	
Ice	Ice cover during Last Glacial Maximum	Categorical	37.7 %	40.1

667

668

669 **Figure legends**

670 **Figure 1.** Hypogean localities selected in the present work. Ice cover during Last Glacial Maximum is
671 reported in the top-right inset map. A code in brackets referring to the Italian speleological cadaster is
672 provided for each locality—more details in Supplementary Material Appendix 1.

673

674 **Figure 2.** Predicted values (filled lines) and 95% confidence intervals (dashed lines) of the effect of mean
675 annual cave temperature (T_{int}) on the probability of presence of *Troglohyphantes* spiders according to glacial
676 history, derived from GLM analysis.

677

678 **Figure 3.** Projected environmental suitability of Western Alps for troglobiomorphic species of
679 *Troglohyphantes* spiders according to current and Last Glacial Maximum (LGM) climatic conditions. LGM
680 projection represents the average of three General Circulation Models (GCMs). Only suitable areas over the
681 threshold of 0.4 are shown.

682

683 **Figure 4.** Projected future environmental suitability of the Western Alps for troglobiomorphic species of
684 *Troglohyphantes* spiders in 2050 and 2070 according to the two Representative Concentration Pathways
685 (rcp) considered in the work. Each map represents the average of the projection into three General
686 Circulation Models (GCMs). Only suitable areas over the threshold of 0.4 are shown.

687

688 **Figure 5.** Current and projected future environmental suitability (high emissions scenario for 2070) of the
689 Western Alps for *Troglohyphantes* spiders. Orange dots represent current localities of *Troglohyphantes*.
690 Localities of *T. lucifuga* and *T. lucifer* are not shown in the map. Only suitable areas over the threshold of 0.4
691 are shown.

692