1	INCIDENT RADIATION AND THE ALLOCATION OF NITROGEN WITHIN ARCTIC PLANT
2	CANOPIES; IMPLICATIONS FOR PREDICTING GROSS PRIMARY PRODUCTIVITY.
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4	L. E. Street <sup>*,1</sup> , G. R. Shaver <sup>2</sup> , E.B. Rastetter <sup>2</sup> , M. T. van Wijk <sup>3,4</sup> , Kaye B. A. <sup>2</sup> , M. Williams <sup>1</sup>
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7 8 9 10 11 12 13 14 15 16 17 18 19 20 21 22 23 24	*Corresponding author Email: <u>1.e.street@ed-alumni.net</u> Telephone:+44 (0)131 650 6480 FAX: +44 (0) 131 662 0478 <sup>1</sup> School of GeoSciences University of Edinburgh Edinburgh, EH9 3JN, UK <sup>2</sup> The Ecosystems Center Marine Biological Laboratory Woods Hole, MA 02543 USA <sup>3</sup> Plant Production Systems, Wageningen University, Plant Sciences, Haarweg 333, 6709 RZ Wageningen, Netherlands.
24 25 26 27 28 29	P.O. Box 30709 Nairobi, Kenya
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1 ABSTRACT

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3 Arctic vegetation is characterised by high spatial variability in plant functional type (PFT) 4 composition and gross primary productivity (P). Despite this variability, the two main drivers of 5 P in sub-Arctic tundra are leaf area index  $(L_T)$  and total foliar nitrogen  $(N_T)$ .  $L_T$  and  $N_T$  have been 6 shown to be tightly coupled across PFTs in sub-Arctic tundra vegetation, which simplifies up-7 scaling by allowing quantification of the main drivers of P from remotely sensed  $L_{T}$ . Our 8 objective was to test the  $L_T$ - $N_T$  relationship across multiple Arctic latitudes and to assess  $L_T$  as a 9 predictor of gross primary productivity (P) for the pan-Arctic. Including PFT specific parameters 10 in models of  $L_{\rm T}$ - $N_{\rm T}$  coupling provided only incremental improvements in model fit, but 11 significant improvements were gained from including site-specific parameters. The degree of 12 curvature in the  $L_{\rm T}$ - $N_{\rm T}$  relationship, controlled by a fitted canopy nitrogen extinction co-efficient, 13 was negatively related to average levels of diffuse radiation at a site. This is consistent with 14 theoretical predictions of more uniform vertical canopy N distributions under diffuse light 15 conditions. Higher latitude sites had higher average leaf N content by mass  $(N_{\rm M})$ , and we show 16 for the first time that  $L_{\rm T}$ - $N_{\rm T}$  coupling is achieved across latitudes via *canopy scale* trade-offs between  $N_{\rm M}$  and leaf mass per unit leaf area ( $L_{\rm M}$ ). Site-specific parameters provided small but 17 18 significant improvements in models of P based on  $L_{\rm T}$  and moss cover. Our results suggest that 19 differences in  $L_{\rm T}$ - $N_{\rm T}$  coupling between sites could be used to improve pan-Arctic models of P 20 and we provide unique evidence that prevailing radiation conditions can significantly affect N 21 allocation over regional scales.

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#### 1 INTRODUCTION

2 Photosynthetic  $CO_2$  uptake by vegetation, gross primary productivity (P), is a key component of 3 the carbon cycle and is strongly linked to climate conditions (Beer *et al.*, 2010). The impacts of 4 climate change on P are a major research topic (Piao et al., 2007, Reichstein et al., 2007, Falge et 5 al., 2002, Ziehn et al., 2011). The Arctic climate is warming more quickly than elsewhere on the 6 globe (Bekryaev et al., 2010) and by accelerating nutrient mineralisation in soils, is expected to 7 drive increases in Arctic plant productivity (Hill et al., 2011, Grant et al., 2011, Chapin et al., 8 1995). Arctic terrestrial carbon stocks are large, 98 Pg in North American Arctic soils alone 9 (Ping *et al.*, 2008), so there is a pressing need to understand terrestrial Arctic C balance in order 10 to quantify carbon cycle – climate feedbacks. This requires robust estimation and prediction of P11 over pan-Arctic scales.

12 Quantification and prediction of P is complicated by high spatial and temporal variability 13 (Williams et al., 1999, Street et al., 2007). One approach to extrapolating across space and time 14 uses process-orientated models which can incorporate remotely sensed information on the spatial 15 and temporal drivers of P (Ryu et al., 2011, Sitch et al., 2007). These drivers include light intensity and temperature, as well as plant biomass or leaf area index (m<sup>2</sup> of leaf per m<sup>2</sup> ground). 16 17 Process-based models of P are based on well-understood biochemical processes at the leaf level 18 (Collatz et al., 1992, Farquhar et al., 1980, von Caemmerer et al., 1981) which must be scaled to 19 whole plant canopies. Up-scaling from leaf to canopy is challenging because photosynthesis 20 responds non-linearly to light, and non-linear gradients of light intensity and leaf properties 21 occur within canopies (de Pury et al., 1997, Leuning et al., 1995). Many current models are 22 based on optimization theory (Clark et al., 2011, Ryu et al., 2011) which states that 23 photosynthesis is optimised if leaf-level nitrogen concentration per unit leaf area  $(N_{\rm L})$ , which is

1	directly related to photosynthetic capacity, (Evans, 1989) varies in proportion to vertical light
2	gradients (Field, 1983, Hirose et al., 1987, Hikosaka et al., 1998). However, radiative transfer
3	within canopies in turn depends on structural characteristics such as canopy height, leaf
4	clumping, and leaf angle distribution (Niinemets, 2010, Hikosaka et al., 1997), as well as the
5	amount of diffuse versus direct radiation (Roderick et al., 2001). Canopy structural properties
6	vary between species and plant functional types (Anten et al., 1999, Anten et al., 1995) and the
7	properties of leaves themselves, including the maximum photosynthetic capacity, vary
8	considerably both within and between species (Wright et al., 2004).
9	Previous studies in low Arctic shrub and tussock tundra ecosystems (in N. Sweden and N.
10	Alaska) suggest however, that there are unexpected system level interactions within mixed
11	species plant canopies which can simplify variability in leaf properties and canopy structure at
12	the stand scale. Van Wijk et al. (2005) and Williams & Rastetter (1999) show that per unit
13	ground area, the total leaf nitrogen $(N_T)$ of plant canopies follows a tightly constrained
14	relationship with $L_{\rm T}$ across a range of PFTs, despite large variation in $N_{\rm L}$ at the species level.
15	Average <i>canopy</i> foliar nitrogen concentration per unit leaf area $(N_T/L_T)$ was 1.9 g N m <sup>-2</sup> across a
16	wide range of vegetation types when $L_{\rm T} < 1.0 \text{ m}^2 \text{ m}^{-2}$ . Crucially, they show that the variability in
17	$N_{\rm T}/L_{\rm T}$ was less than would be expected if canopies were randomly constructed from the pool of
18	available species, suggesting convergence of leaf-level properties at the canopy scale, rather than
19	a central tendency towards the mean value of $N_{\rm L}$ .
20	$L_{\rm T}$ and $N_{\rm T}$ have also been shown to be the most important drivers of P in Swedish and
21	Alaskan tundra; a strong correlation between $L_{\rm T}$ (determining light capture) and $N_{\rm T}$ (determining
22	light utilisation by photosynthetic enzymes) optimising P by balancing limitations on

23 photosynthesis (Williams *et al.*, 1999). This has been confirmed in plot scales studies in which

1  $L_{\rm T}$  alone explained 80 % of the variability in P at constant light level (Shaver *et al.*, 2007, Street 2 et al., 2007). The tight coupling between  $L_T$ ,  $N_T$  and P greatly simplifies up-scaling, by reducing species level variability, and allowing remote-sensed estimates of P based on  $L_{\rm T}$  or  $N_{\rm T}$  alone. 3 4 Whether the canopy-scale relationship between  $L_{\rm T}$  and  $N_{\rm T}$  in Sweden and Alaska, and the 5 corresponding linkages to P, are general ecological scaling relationships applicable across the 6 pan-Arctic has not previously been tested. To address this, we ask the following questions: 1) is 7 there a general relationship between  $L_{\rm T}$  and  $N_{\rm T}$  for Arctic vegetation? and 2) is there a general 8 relationship between P and  $L_{\rm T}$  across the Arctic? 9 While soil nutrient availability may dictate patterns of plant abundance and therefore  $L_{\rm T}$ 10 (Shaver et al., 1980, Shaver et al., 1986, 1991), we expect that lower irradiance at higher 11 latitudes will modify the optimal development of leaf area with respect to available nitrogen;  $L_{T}$ -12  $N_{\rm T}$  will be coupled but shifted towards lower  $N_{\rm T}$  per unit leaf area. We expect lower  $N_{\rm T}$  per unit 13 leaf area to result in lower P per unit leaf area, under saturating light conditions. To test these 14 hypotheses, we present destructive measurements of  $N_{\rm T}$  and  $L_{\rm T}$  across five Arctic sites spanning 15 latitudes from 68 - 78 °N. We compare fitted relationships between  $N_{\rm T}$  and  $L_{\rm T}$  to patterns of 16 incident radiation; both total short-wave (SW) radiation and the SW diffuse fraction. We use 1 m 17  $\times$  1 m chamber measurements of P, together with indirect measurements of  $L_{\rm T}$  based on the 18 normalised difference vegetation index (NDVI), to explore the relationship between  $L_{\rm T}$  and P. 19

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#### 1 MATERIALS AND METHODS

A list of symbols and abbreviations is given in Table 1. The present analysis combines new  $N_{\rm T}$ , 2  $L_{\rm T}$ , and P data from near Longyearbyen, Svalbard, from Zackenberg, NE Greenland, and from 3 4 Barrow, Alaska with previously published data from Abisko, northern Sweden and from near 5 Toolik Lake, Alaska (Table 2). We measured  $L_{\rm T}$  and  $N_{\rm T}$  destructively on small (0.03 – 0.09 m<sup>2</sup>) harvest plots (Table 2). We measure CO<sub>2</sub> fluxes on different (1 m<sup>2</sup>) flux plots. We used NDVI to 6 estimate  $L_{\rm T}$  for the 1m<sup>2</sup> flux plots using site-specific calibrations between  $L_{\rm T}$  and NDVI based on 7 8 the smaller harvested plots. To improve the performance of these calibration relationships we 9 also include moss % cover as an explanatory variable in the  $L_{T}$ -NDVI model. We used NDVI to 10 measure  $L_{\rm T}$  for the flux plots because 1) it is non-destructive and 2) it allows for measurement of  $L_{\rm T}$  on larger scales (1m<sup>2</sup> harvests would not have been feasible). The use of NDVI to measure  $L_{\rm T}$ 11 12 also increases the relevance of the analysis for up-scaling, though comparing hand-held NDVI 13 measurements to satellite derived values is beyond the scope of this study.

14

15 *Site descriptions* 

16 LONGYEARBYEN, SVALBARD

We measured CO<sub>2</sub> flux using chamber techniques and sampled vegetation at ten sub-sites, all within approximately 20 km of the town of Longyearbyen, on the island of Spitzbergen (78° 13'N, 15°37'E). Mean annual air temperature (MAT) for Longyearbyen is -5 °C, mean July temperature is 6°C and mean precipitation ~310mm (Forland *et al.*, 2000). Vegetation in the fjord (Adventdalen) in which Longyearbyen is situated ranges from salt marsh on the margins of the estuary, to wet sedge meadow (mostly *Dupontia*, *Carex* and *Eriophorum spp.*) on the flat valley bottom, to dwarf shrub heath communities on well drained slopes. Dwarf shrub vegetation

1 is characterized by locally dominant patches of *Cassiope tetragona*, *Dryas octopetala*, and *Salix* 2 polaris communities (Baddeley et al., 1994). Plots were chosen to sample the range of vegetation types within the area. All field measurements were made 14<sup>th</sup> July - 3<sup>rd</sup> August 2005. 3 4 5 ZACKENBERG, GREENLAND 6 We measured CO<sub>2</sub> flux using chamber techniques and sampled vegetation at sub-sites within 7 approximately 2 km of Zackenberg Research Station (74° 28'N, 20°34'E) and below an 8 elevation of 100 m. Mean annual air temperature (1996 – 2008) for Zackenberg, is –9.0 °C, mean 9 July temperature is 6.1 °C, with a total annual precipitation of 218 mm (www.zackenberg.dk). 10 Vegetation in the area around Zackenberg consists of wet fen and grassland in areas by water 11 tracks, with heath vegetation dominated by *Cassiope tetragona* on better drained level ground. 12 Heath dominated by *Vaccinium uliginosum* or *Dryas* species is more common on exposed slopes. 13 There are also extensive areas of snowbed vegetation dominated by Salix arctica. Plots were 14 chosen to sample the range of vegetation types within the area. All field measurements were 15 made between 8th July – 1st August 2006. 16 17 BARROW, ALASKA 18 Flux measurements, leaf harvests, and reflectance measurements were made in the Barrow Environmental Observatory (BEO) near Barrow, Alaska (71°18'N, 156°40'W) in July 2009. All 19 20 measurements were made within a single, shallow, drained and revegetated thaw lake basin and

21 surrounding shoreline ridges. The area was entirely underlain by permafrost with ice-wedge

22 polygons creating local microrelief. Sites were selected to represent vegetation along a soil

23 moisture gradient, from constantly-flooded, emergent wet-sedge vegetation dominated by

1	rhizomatous sedges and grasses to relatively dry ridges dominated by creeping willow species
2	and grasses. Vegetation, soils, and climate in the Barrow region have been thoroughly described
3	in past research (Brown et al., 1980), including at the BEO adjacent to sites used in this research
4	(Hollister et al., 2005). Annual temperature at Barrow is -11 °C and the long-term average
5	temperature in July is 4 °C although this has increased in recent decades.
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9	Abisko, Sweden
10	We used leaf harvest data from near Abisko in Northern Sweden (68°18'N,18°51'E) collected
11	between the 15 <sup>th</sup> July and 30 <sup>th</sup> July 2002 (van Wijk <i>et al.</i> , 2005). We use CO <sub>2</sub> flux data collected
12	between the 22 <sup>nd</sup> July and 5 <sup>th</sup> August 2004 at two sites nearby the site of vegetation sampling,
13	(the 'Stepps' site and 'Paddus' site), and at another upland site (the 'Latnja' site) (Shaver et al.,
14	2007). MAT at Abisko is -1 °C, mean July temperature is 11 °C and mean annual precipitation is
15	225-475 mm (van Wijk <i>et al.</i> , 2005).
16	
17	Toolik, Alaska
18	We used $L_{\rm T}$ and $N_{\rm T}$ data collected during 1997 as part of the Arctic flux study within the
19	Kuparuk watershed on the northern side of the Brooks Range, AK, USA (Williams et al., 1999).
20	We use $CO_2$ flux data collected between $12^{th}$ July and $4^{th}$ August 2004 in tussock, dry heath and
21	shrub tundra at Toolik Lake (68°38'N,149°36'W) and nearby at Imnavait Creek
22	(68°37'N,149°19'W) (Shaver et al., 2007). Though at the same latitude as Abisko, the climate at

1 Toolik is more continental; MAT at Toolik Lake is -10 °C and mean July temperature is 14 °C.

2 Mean annual precipitation is 200-400 mm (van Wijk *et al.*, 2005).

3

4 Measurements

5 NDVI, LEAF AREA INDEX AND CANOPY N OF HARVEST PLOTS

6 We measured  $L_T$ ,  $N_T$ , and total leaf biomass ( $M_T$ ) of harvested plots at Svalbard (n = 48),

Zackenberg (n = 78), Barrow (n = 23), Toolik (n = 92) and Abisko (n = 94). The size of the
harvested plots was between 0.03 and 0.09 m<sup>2</sup> (Table 2). At Zackenberg, Svalbard, and Barrow
we also measured the NDVI of each plot before harvesting using a Unispec spectral analyser (PP
systems, Haverhill, Massachusetts, USA) following the methods of Street *et al.* (2007). The
Unispec instrument records reflectance spectra from 0.3 mm to 1.0 mm. We held the sensor at a
vertical height (< 1m) such that the field of view equated to the area being harvested (Table 2).</li>
The sensor was positioned over the plot using a 1m ruler held vertically in the centre of the plot;

14 the ruler was removed prior to measurement. We calculated NDVI using the formula:

15 
$$NDVI = (R_{\rm NIR} - R_{\rm VIS})/(R_{\rm NIR} + R_{\rm VIS})$$

16 where  $R_{NIR}$  is reflectance at a wavelength of 0.725–1.0 µm and  $R_{VIS}$  is reflectance at 0.56–0.68 17 µm. Reflected radiation was calculated as a proportion of incident radiation using a standard 18 white (barium sulphate) reference panel. Reference panel readings were taken prior to each 19 measurement. The NDVI of the surface was also measured after the canopy was removed 20 (*NDVI*<sub>post</sub>), within 1 hour of the top of canopy measurement. NDVI measurements were 21 generally made between 10 am and 4 pm.

*L*<sub>T</sub>, *M*<sub>T</sub> and *N*<sub>T</sub> were measured destructively, following the methods of Van Wijk *et al.*(2005). For *Cassiope tetragona* we doubled the one-sided projected leaf area (Campioli *et al.*,

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#### 4 NDVI, LEAF AREA INDEX AND PERCENT COVER OF FLUX PLOTS

5 We measured the NDVI of flux plots in Svalbard, Barrow and Zackenberg using the Unispec 6 analyser in a grid of 9 points, following the methods of Street et. al. (2007). We estimated 7 absolute aerial cover in each flux plot for all vascular plants by species (i.e. total cover can be > 8 100%), and for bryophytes. We did this by placing a  $5 \times 5$  string grid (each square = 0.04 m<sup>2</sup>) 9 over the plot, and visually estimating cover in each square, then calculated an average species 10 cover for the entire plot. Both harvest and flux plots were classified as either deciduous, 11 evergreen, graminoid or mixed vegetation according to the contribution of each functional type 12 to total biomass (for the harvest plots) or cover (for the flux plots). Plots with the abundance of a 13 single PFT > 70 % were classified as that type, otherwise plots were classified as mixed. 14 To estimate  $L_{\rm T}$  for the flux plots from NDVI we used calibration relationships between 15 NDVI and  $L_{\rm T}$  for each plant functional type based on destructive data from the harvested plots. 16 We modelled the relationship between  $L_{\rm T}$  and NDVI using an equation modified from Steltzer & 17 Welker (2006):

2009b). We estimated the percent cover of bryophytes for each plot, and at Svalbard and

Zackenberg re-measured NDVI following removal of the vascular plant canopy.

$$L_{T} = \frac{1}{Kc} \ln \left( \frac{NDVI_{\min} - NDVI_{\max}}{NDVI - NDVI_{\max}} \right)$$
(1)

20 Where  $NDVI_{max}$  is a fitted parameter representing the theoretical maximum possible NDVI for a 21 vegetation canopy (which cannot be measured for open high Arctic canopies) and  $K_c$  is a fitted extinction co-efficient. NDVImin is a parameter representing the minimum possible NDVI of the 22

background surface (i.e. soil or moss). We used the minimum measured value of *NDVI*<sub>post</sub> (0.24)
 as an estimate of *NDVI*<sub>min</sub> when fitting equation 1.

3 To take into account the effect of mosses on NDVI, and therefore potentially improve the accuracy of our L<sub>T</sub> estimates, we also fitted a second model in which we assume that NDVI<sub>min</sub> 4 5 increases as moss cover increases. In other words, the ground surface underneath the vascular 6 canopy on average becomes greener as mosses become more abundant, but at any moss cover 7 there is still a distribution of possible background NDVI values - the minimum of which 8 representing NDVI<sub>min</sub>. We observed a significant correlation between bryophyte cover and  $NDVI_{post}$ , as shown in Figure 1 so we used the slope of this relationship (0.0024 %<sup>-1</sup>) as an 9 10 estimate of the rate of increase in *NDVI*<sub>min</sub> as moss cover increases (*a* in equation 2).

11

$$L_{T} = \frac{1}{Kc} \ln \left( \frac{\left[ aB_{c} + 0.24 \right] - NDVI_{\max}}{NDVI - NDVI_{\max}} \right)$$

13

12

14 Where  $B_c$  is bryophyte cover (%). Note that there are two fitted parameters in both equation 1 15 and 2.

For both Toolik and Alaska we used the previously published values of  $L_{\rm T}$  for each flux measurement plot, based on vegetation specific NDVI -  $L_{\rm T}$  calibrations (Street *et al.*, 2007).

18

#### 19 *Flux measurements*

20 We measured the light response of net ecosystem  $CO_2$  exchange ( $F_c$ ), and ecosystem respiration

21 ( $R_E$ ) in Svalbard, Zackenberg, Barrow, Toolik and Abisko over 1 × 1 m patches of vegetation

22 (Table 2). All flux measurements were made using protocols described in Williams et al. (2006),

23 Street et al. (2007) and Shaver et al. (2007). In Zackenberg we also measured fluxes for 15 0.3 m

(2)

1  $\times$  0.3 m patches using a smaller chamber and destructively sampled  $L_{\rm T}$ . For the Toolik and 2 Abisko CO<sub>2</sub> flux data we used a sub-set of the data presented in Shaver *et al.* (2007) that 3 coincided with the dates of CO<sub>2</sub> flux data collection at other sites (July – early August).

5 Data analysis

6 LAI and total canopy N

7 We aimed to find the parameters for a theoretical model that relates  $N_{\rm T}$  to  $L_{\rm T}$  – and to determine 8 whether the parameters describing the model differ significantly between sites. It has been 9 argued that an exponential decline in foliar N through the canopy is a plant strategy for 10 maximizing canopy photosynthesis with respect to canopy nitrogen (Field, 1983, Hirose *et al.*, 11 1987, Hikosaka *et al.*, 1998).

12

13

$$N_L = N_0 e^{-\gamma L c} \tag{3}$$

14

15 where  $N_{\rm L}$  is the nitrogen concentration of a leaf in the canopy (g N m<sup>-2</sup> leaf),  $N_0$  is the top of the 16 canopy nitrogen concentration (g N m<sup>-2</sup> leaf),  $\gamma$  an extinction coefficient (m<sup>2</sup> ground m<sup>-2</sup> leaf), 17 and  $L_c$  is the cumulative leaf area above the leaf (m<sup>2</sup> leaf m<sup>-2</sup> ground). Total canopy nitrogen is 18 then the integral of equation 3.

19

20  

$$N_{T} = \int_{0}^{L_{T}} (N) dL = \frac{N_{0}}{\gamma} \left( 1 - e^{-\gamma L_{T}} \right)$$
(4)  
21

1 where  $N_{\rm T}$  is the total canopy nitrogen (g N m<sup>-2</sup> ground) and  $L_{\rm T}$  is leaf area index (m<sup>2</sup> leaf m<sup>-2</sup> 2 <sup>2</sup>ground). Equation 4 would be expected to describe the relationship between  $N_{\rm T}$  and  $L_{\rm T}$  if  $N_0$  and 3  $\gamma$  are uniform over the area sampled.

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5 CO_2 flux
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6 To test the relationship between  $L_{\rm T}$  and canopy photosynthesis, we compare the parameters of 7 fitted *P* light response curves to  $L_{\rm T}$  for each  $1 \times 1$  m plot. *P* at each light level was calculated by 8 subtracting  $F_{\rm c}$  from  $R_{\rm E}$ . The light response of photosynthesis was then modelled with a 9 rectangular hyperbola

10

$$P = \frac{P_{\max} \times I}{\left(\frac{P_{\max}}{E_0}\right) + I}$$
(5)

12

11

13 where  $P_{\text{max}}$  is the rate of light saturated canopy level photosynthesis (µmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>), *I* is the 14 incident photosynthetic flux density (µmol photons m<sup>-2</sup> s<sup>-1</sup>),  $E_0$  is the initial slope of the light 15 response curve or canopy-level quantum efficiency at low light levels (µmol CO<sub>2</sub> µmol<sup>-1</sup> 16 photons).

Each light curve was used to predict *P* at 1000 µmol photons m<sup>-2</sup> s<sup>-1</sup> (*P*<sub>1000</sub>). We compared the relationship between *P* near light saturation (*P*<sub>1000</sub>) and leaf area, using only curves where maximum *I* measured during the light curve exceeded 1000 µmol photons m<sup>-2</sup> s<sup>-1</sup>. We assume leaf level photosynthesis at 1000 µmol<sup>-1</sup> photons (*P*<sub>1000</sub>) is approximately linearly related to  $N_L$  (Hirose and Werger 1987) and therefore follows an exponential distribution with canopy depth as  $N_L$ , giving an analogous equation to equation 4 but with a constant term for moss photosynthesis :

$$P_{1000} = \frac{P_0}{\gamma_p} \left( 1 - e^{-\gamma_p \, LAI} \right) + P_m \tag{6}$$

3 where  $P_{1000}$  is canopy-level gross photosynthesis at 1000 µmol photons m<sup>-2</sup> s<sup>-1</sup>,  $P_0$  is top of the 4 canopy *P* per unit leaf area at 1000,  $\gamma_p$  is the extinction of leaf level *P* with canopy depth and  $P_m$ 5 is a constant term for moss photosynthesis.

6 We also compare the relationship between canopy-level quantum efficiency  $E_0$  (the initial 7 slope of the light response curve, equation 5) and  $L_T$ . We assume that  $E_0$  is not strongly related to 8 leaf N concentration at low light levels (Hirose and Werger 1987) and therefore follows a 9 uniform distribution through the canopy and a linear relationship with  $L_T$ . We again include a 10 constant for the photosynthetic activity of mosses:

$$E_0 = \alpha . LAI + E_m \tag{7}$$

12

11

1

2

13 Where  $E_0$  is canopy-level quantum efficiency at low light levels (µmol CO<sub>2</sub> µmol<sup>-1</sup> photons),  $\alpha$  is 14 the increase in  $E_0$  per unit canopy leaf area, and  $E_m$  is a constant representing photosynthetic 15 activity of mosses.

16

#### 17 Statistical analysis

18 We compared alternate statistical models of the relationship between  $N_{\rm T}$  and  $L_{\rm T}$  (equation 4),

19  $P_{1000}$  and  $L_{\rm T}$  (equation 6) and  $E_0$  and  $L_{\rm T}$  (equation 7) using general non-linear least squares fitting

20 procedures in the 'nlme' library for R Version 2.12.1 (R Development Core Team (2008)). We

- 21 initially fit a general model to the whole data set. We then repeated the model fitting including
- site (or PFT) based groupings for the data, both for individual parameters and parameters in
- 23 combination. For example, for the  $N_{\rm T}$  model (equation 4) we compared the general model to

(7)

1	alternate models that included 1) site-specific fitted values of $N_0$ with a general fitted value for $\gamma$ ,
2	2) site-specific fitted values of $\gamma$ with a general fitted value for $N_0$ and 3) site-specific fitted
3	values of both $N_0$ and $\gamma$ . We then repeated the analysis using PFT instead of site to group the
4	data. We do not show results for models with PFT specific values of $N_0$ or $\gamma$ alone, as even with
5	PFT effects for both parameters, we saw little improvement in model fit. For the $P_{1000}$ and $E_0$
6	models we also show results for a restricted set of possible data groupings because improvements
7	in model fit by grouping the data by PFT were small. Alternative models were compared using
8	Akaike's Information Criterion (AIC), Bayes Information Criterion (BIC), loglikelihood
9	(LogLik) ratio tests and the root mean square error (RMSE) of model predictions. We included
10	power variance functions in all $N_{\rm T}$ models, to account for heteroscedasticity in the data. It was
11	necessary to include power variance functions to account for heteroscedasticity in the $E_0$ data,
12	but not in the $P_{1000}$ data.

13

#### 14 Radiation data

15 To investigate the effect of the radiation environment, we compared site-specific fitted 16 parameters for equation 4 to average solar radiation conditions. We used 5 years of data for 17 which we were able to obtain continuous data for every site (1998-2000, 2003 & 2005). We 18 calculated the sum of hourly short-wave (SW) radiation over the growing season (defined as 1<sup>st</sup> 19 June to 31st August) for each year, then calculated the mean and standard deviation across years 20 for each site. We calculated the average diffuse radiation fraction by averaging hourly daytime (defined as the period where incident SW > 20 Wm<sup>-2</sup> or approximately 40  $\mu$ mol m<sup>-2</sup> s<sup>-1</sup> I) diffuse 21 22 fraction over each growing season. For Svalbard we used hourly global SW and diffuse SW data from Ny-Alesund, approximately 100 km NW of Longyearbyen (supplied by the Alfred-Wegner 23

1 Institute, www.awi.de). For Abisko we used hourly global SW data provided by Abisko 2 Scientific Research Station (www.ans.kiruna.se) to model hourly diffuse SW fraction based on 3 the ratio of modelled extraterrestrial to measured global SW according to Erbs et al. (1982). We 4 tested the Erbs model with 2.5 weeks of global and diffuse SW data provided by the ABACUS 5 project (www.geos.ed.ac.uk/abacus). For Toolik we used global SW data from the Toolik Lake 6 Long Term Ecological Research (LTER) database (ecosystems.mbl.edu/ARC) to model daily 7 diffuse SW fraction. We tested the Erbs model results at Toolik with 6 weeks of total and diffuse 8 I from late summer 2008 provided by the Arctic Observing Network (aon.iab.uaf.edu). Global 9 SW was estimated from I using an empirical relationship for that site (see Supplementary 10 Material). For Zackenberg we used hourly global SW data provided by the ClimateBasis 11 programme at Zackenberg research station (available at www.zackenberg.dk/data/). For Barrow we use measured SW and diffuse SW data provided by the US Atmospheric Radiation 12 13 Monitoring program (ARM) (www.arm.gov).

14

15 RESULTS

16 *Leaf area index and total foliar nitrogen* 

17  $L_{\rm T}$  was < 1.0 m<sup>2</sup> m<sup>-2</sup> in 90 % of the 151 harvests carried out at Zackenberg, Svalbard and Barrow 18 reflecting the characteristic short stature, low  $L_{\rm T}$  plant canopies at higher latitude and coastal 19 tundra vegetation compared to lower latitudes. The maximum  $N_{\rm T}$  values recorded (across all 20 sites) were in Svalbard with ~ 4.5 g N m<sup>-2</sup> ground (Fig. 2b). Relationships between  $N_{\rm T}$  and total 21 leaf mass ( $M_{\rm T}$ , g leaf m<sup>-2</sup> ground area), and between  $L_{\rm T}$  and  $M_{\rm T}$  were less well-constrained than 22 the relationship between  $L_{\rm T}$  and  $N_{\rm T}$  both within and across sites (Fig. 2 a b & c, Fig. 3). Canopies 23 with low average N concentration per unit leaf mass, ( $N_{\rm T}/M_{\rm T}$ ) tended to have greater average leaf

1	mass per area $(M_T/L_T)$ (Fig 2d). The highest average canopy $N_T/M_T$ values were in Zackenberg;
2	which in several plots was > 35 mg N g <sup>-1</sup> leaf (Fig 2d). At $L_T < 1$ , average $N_T/M_T$ at each site was
3	greater at Zackenberg (25.8 $\pm$ 1.1 mg N g $^{\text{-1}}$ leaf ), Svalbard (21.4 $\pm$ 0.8 mg N g $^{\text{-1}}$ leaf) and
4	Barrow (24.8 $\pm$ 0.8 mg N g <sup>-1</sup> leaf) than at the lower latitude sites Toolik (18.8 $\pm$ 0.4 mg N g <sup>-1</sup>
5	leaf) and Abisko (12.1 $\pm$ 0.5 mg N g <sup>-1</sup> leaf).
6	Average leaf-level nitrogen concentrations by mass for individual species $(N_{\rm M})$ varied
7	between 16 mg N g <sup>-1</sup> (1.6 % by mass) in <i>Dryas</i> leaves at Zackenberg, and 38 mg N g <sup>-1</sup> (3.8 % by
8	mass) in <i>Polygonum viviparum</i> at Zackenberg. Nitrogen concentrations per unit leaf area $(N_L)$ for
9	individual species varied > 4-fold, from 0.95 g N m <sup>-2</sup> for <i>Saxifraga cernua</i> at Barrow, to > 4.0 g
10	N m <sup>-2</sup> in forbs and graminoids at Zackenberg (Table 3).
11	Including site specific parameters in the $N_{\rm T}$ model (equation 4) resulted in significant
12	improvements in model fit compared to the general model (Table 4). The lowest RMSE
13	(0.18 g N m <sup>-2</sup> ) was achieved by including site specific parameters both for $N_0$ and $\gamma$ , with
14	significantly improvements in fit (assessed with Loglikelihood ratio tests) both over the general
15	model, and models with site specific fitted parameters for either $N_0$ or $\gamma$ alone (Table 4). The
16	overall $R^2$ of modelled versus measured values for a model with site specific $N_0$ and $\gamma$ was 0.93.
17	The average site-specific fitted value for top of canopy nitrogen concentration ( $N_0$ ) was 2.2 g N
18	m <sup>-2</sup> , with a standard deviation of 0.2 g N m <sup>-2</sup> , or coefficient of variation (CV) of 9 %. The
19	average site-specific fitted extinction coefficient, $\gamma$ , was 0.22, with standard deviation across sites
20	of 0.31 or CV of 140 %.
21	Including PFT specific parameters in the $N_{\rm T}$ model also resulted in a significant
22	improvement in model fit, but the reduction in RMSE (from 0.27 to 0.25 g N m <sup>-2</sup> ) and associated

23 increase in LogLiklihood were small compared to the site specific models (Table 4).

1

#### 2 Radiation conditions

Total growing season SW radiation was lowest for Svalbard at  $1310 \pm 100$  MJ m<sup>-2</sup> (1 S.D) and 3 greatest at Zackenberg at  $1707 \pm 111 \text{ MJ m}^{-2}$  (1 S.D) (Fig. 4a.). The Erbs *et al.* (1982) model of 4 5 diffuse radiation fraction performed well when tested with measured data from Abisko (Supp. 6 Material Fig. 2 a & b) though there was a slight bias (daily modelled vs. measured diffuse radiation slope = 1.11, intercept = -0.59,  $R^2 = 0.86$ , RMSE = 0.99 MJ m<sup>-2</sup> day<sup>-1</sup>). This bias was 7 8 corrected for when calculating average growing season diffuse fraction. We also corrected for 9 bias at Toolik Lake (daily modelled vs. measured diffuse radiation slope = 1.02, intercept = 0.69,  $R^2 = 0.72$ , RMSE = 1.38 MJ m<sup>-2</sup> day<sup>-1</sup>). The average diffuse fraction was greatest in Svalbard (80 10 11  $\% \pm 3.5 \%$  (1 S.D.)) and lowest at Zackenberg (62  $\% \pm 3.9 \%$  (1 S.D.)). There was a positive 12 trend in site specific fitted values of  $N_0$  with total growing season radiation, which was not 13 statistically significant (Fig 4a). There was a significant negative trend in site specific fitted 14 values of  $\gamma$  with increasing diffuse radiation fraction (Fig. 4b) (P = 0.018).

15

16 Leaf area index and NDVI

17 We found a significant positive correlation between  $NDVI_{post}$  and  $B_c$  (Spearman  $\rho = 0.73$ ,

18 p < 0.001). NDVI for ground with < 10 % moss cover was  $0.38 \pm 0.0057$  (1 SE), for ground with

19 > 90 % moss cover NDVI was 0.61  $\pm$  0.0028 (1 SE). The slope of the linear relationship (least-

20 squares fit) between  $NDVI_{post}$  and % moss cover was 0.0024 %<sup>-1</sup> (Fig 1). There was a clear

21 relationship between NDVI and observed  $L_{\rm T}$  in Svalbard, for plots with both low and high moss

22 cover (Fig 5a). The relationship between  $L_{\rm T}$  and NDVI was more scattered for Zackenberg, with

23 a clustering of points with > 50 % moss cover at high NDVI values but low  $L_{\rm T}$  values (Fig 5c).

19

1 Including moss cover as an explanatory variable in the  $L_{\rm T}$ -NDVI model (equation 2), 2 parameterised separately for each PFT, we could explain 74 % of the variation in  $L_{\rm T}$  at Svalbard (Fig 5b). If we did not include the effect of mosses on NDVI (i.e. by using equation 1) in the 3 4 PFT specific calibrations, we could explain 70 %. Likewise for Zackenberg, we could explain 5 49 % of the variation in  $L_{\rm T}$  using equation 2 (Fig 5d), but with equation 1 we could explain 37 %. 6 Data from Barrow were limited so we parameterised the  $L_{\rm T}$  -NDVI relationship for graminoids, 7 and combined mixed/forb and deciduous vegetation. We could explain 34 % of the variation in 8 L<sub>T</sub> at Barrow using equation 2 (Fig 5f); NDVI alone (equation 1) explained 23 %. The NDVI 9 calibration relationships used to predict  $L_{\rm T}$  for Abisko and Toolik, published in Street *et. al.* 2007, explained 84 % of the variation in  $L_{\rm T}$ . 10

11

#### 12 Leaf area index and canopy photosynthesis

 $P_{1000}$  increased with  $L_{\rm T}$  for all sites, up to a maximum value of 17 µmol m<sup>-2</sup> s<sup>-1</sup>. The flux plots at 13 Toolik and Abisko had greater  $L_{\rm T}$  than the higher latitude sites, up to 2 m<sup>2</sup> m<sup>-2</sup>, and also had the 14 15 highest values of  $P_{1000}$  (Figure 6a). There was a large degree of overlap in the relationship between  $P_{1000}$  and  $L_{\rm T}$  between sites (Figure 6a). Fitted  $P_0$  for the general  $P_{1000}$  model (equation 6) 16 was 11.8  $\mu$ mol m<sup>-2</sup> s<sup>-1</sup> and  $\gamma_p$  was 0.6 m<sup>-2</sup> ground m<sup>-2</sup> leaf. Including site effects in the model 17 18 resulted in significant but small increases in model fit over the general model (Table 5). The fitting routine was unable to find a solution with site-specific fitted parameters for  $P_0$ ,  $\gamma_P$  and  $P_m$ 19 indicating model over-parameterisation. Fitting equation 6 with site specific parameters for  $\gamma_P$ 20 and  $P_{\rm m}$  gave the best model fit, and resulted in a reduction in RMSE from 1.89 to 1.71  $\mu$ mol m<sup>-2</sup> 21 s<sup>-1</sup>, with a modelled vs. measured  $R^2$  of 0.78. Including PFT specific fitted parameters for  $\gamma_P$  and 22  $P_{\rm m}$ , and  $P_0$  and  $\gamma_{\rm P}$  resulted in small improvements in RMSE, but increased the AIC and resulted 23

in insignificant Loglikelihood ratio tests compared to the general model; indicating that the data
 do not provide support for a PFT specific model.

3	$E_0$ also increased with increasing $L_T$ and had a maximum value of 0.047 µmol CO <sub>2</sub> µmol <sup>-</sup>
4	<sup>1</sup> photons (Figure 6b). There was also a high degree of overlap between sites in Figure 6b, and
5	we found no significant improvements in model fit by including site or PFT specific parameters
6	in the $E_0$ model (equation 7). The general model was associated with the lowest AIC value and
7	had slope of 0.017 $\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup> photons m <sup>-2</sup> and intercept of 0.005 $\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup>
8	photons. $R^2$ of the linear regression between modelled and measured values was 0.67.
9	

10

#### 1 DISCUSSION

2 Is there a pan-Arctic relationship between leaf area index and total foliar nitrogen?

3 The fit of the  $N_{\rm T}$  model was only marginally improved by including PFT specific

4 parameterisations, consistent with the conclusions of Van Wijk *et al.* (2005) and Campioli *et al.* 

5 (2009a) that the relationship between  $L_{\rm T}$  and  $N_{\rm T}$  converges for a wide range of vegetation types.

6 Including site-specific parameters led to much greater improvements in model fit. Our cross-site

7 estimate of 2.2 g N m<sup>-2</sup> for  $N_0$  is close to the average canopy  $N_T/L_T$  of 1.9 g N m<sup>-2</sup> reported by

8 Van Wijk et al (2005) for Abisko and Toolik. Contrary to our original hypothesis we found no

9 evidence that the  $L_T$ - $N_T$  relationship is shifted towards lower N per unit leaf area at higher

10 latitude. Values of  $N_0$ , representing top of canopy N per unit leaf area, were highest at

11 Zackenberg, the second most northerly site. There was also no latitudinal pattern in the fitted

12 parameter for the curvature of the relationship  $(\gamma)$ .

13 The large variability (large CV) in fitted  $\gamma$  compared to  $N_0$  is reflected by the divergence 14 between sites in Fig 2b, which increases as  $L_{\rm T}$  increases; there is a large degree of overlap at low 15  $L_{\rm T}$ . The convergence in the  $N_{\rm T}$ - $L_{\rm T}$  relationship at low  $L_{\rm T}$  occurs despite large variation in average 16 canopy foliar N concentration by mass  $(N_T/M_T)$  and leaf thickness  $(L_T/M_T)$  between as well as 17 within sites. For example, average  $N_T/M_T$  at Zackenberg, Svalbard and Barrow was greater than 18 at Toolik and Abisko. Higher metabolite concentrations in colder climates can be expected as an 19 acclimation response to lower rates of enzyme activity (Chapin et al., 1983), or because of 20 passive lack of dilution resulting from slower growth (Weih et al., 2001). Alternatively, the 21 shorter growing season in Svalbard and Zackenberg may favour short-lived leaves, which tend to 22 be thinner (Shipley *et al.*, 2006) forcing canopies to concentrate N per unit mass. There are 23 therefore site level difference in the way  $L_{\rm T}$ - $N_{\rm T}$  coupling is achieved; in Svalbard and

Zackenberg leaves are thinner with higher N concentration by mass (*Cassiope tetragona*, the only abundant evergreen shrub, being the exception), whereas in Sweden, leaves tend to be thicker, with lower N concentration by mass. This is reflected in the clustering of points along the length of the curve shown in Fig 2d. Trade-offs between leaf properties have been well documented across species at the global scale (Wright *et al.*, 2004, Reich *et al.*, 2004). This is the first time such trade-offs have been documented as an 'emergent' system-level property of mixed species plant canopies.

8

### 9 The influence of radiation

10 Our data suggest that radiation conditions have an impact on N allocation, but that this is 11 not a direct result of changes in latitude and therefore day-length and/or sun angle. We 12 demonstrate a significant correlation between  $L_{\rm T}$ - $N_{\rm T}$  curvature ( $\gamma$ ) and fraction diffuse SW (Fig. 13 4b). While not proof of a causal link, this finding is consistent with theory (Hirose, 2005, 14 Niinemets, 2007) which suggests that canopy N is allocated optimally if  $N_{\rm L}$  declines in 15 proportion to light extinction with canopy depth. Measured gradients in leaf N content in the field are often sub-optimal (Meir et al., 2002, Bond et al., 1999); the degree of N extinction is 16 17 less than would maximise potential carbon gain. Nevertheless, it follows that under diffuse 18 conditions N should be distributed more uniformly due to greater light penetration into the 19 canopy (Roderick et al., 2001, Meir et al., 2002). There are other important influences on canopy 20 light penetration, such as canopy height, leaf angle and geometry (Anten et al., 1995) but the 21 lack of significant improvements in model fit when incorporating PFT specific parameters 22 suggests that the properties of incident radiation may be more important in controlling the 23 distribution of N.

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1 Differences in diffuse radiation between sites can be explained by differences in 2 cloudiness. Zackenberg had the lowest average summer diffuse fraction, and Svalbard the 3 highest. Remotely sensed maps of cloud frequency (Wylie et al., 2005) show high cloud cover 4 (approaching 100%) over the polar oceans and around Svalbard. Cloud frequency over 5 Greenland is lower (around 60 %) than that of terrestrial Arctic at lower latitudes presumably 6 because of the influence of the Greenland ice cap on cloud formation. Low cloud frequency at 7 Zackenberg also explains how the total incident SW radiation can be greater than at lower 8 latitude sites (Fig 4a.).

9 We found no significant relationship between  $N_0$  and total incident SW (or average SW, 10 data not shown), although the trend was positive. Relatively constrained fitted values for  $N_0$  (9 % 11 CV) compared to  $\gamma$  (CV 140%) imply that the rate of extinction of N<sub>L</sub> with canopy depth is a more variable canopy characteristic across sites than top of canopy  $N_{\rm L}$ . The low variability of  $N_0$ 12 13 across Arctic sites is surprising given the host of other factors that affect soil N availability such 14 as soil type, depth, moisture and temperature, and presence and absence of permafrost. We argue 15 however that these soil factors are more likely to affect whole plant, or whole canopy  $N_{\rm T}$  (and  $L_{\rm T}$ ) 16 rather than top of canopy  $N_{\rm L}$ . For example, in a global meta-analysis, Poorter *et. al.* (2009) show 17 that whole plant leaf area is more responsive to nutrient limitation than leaf mass per unit area, 18 which is only affected when growth becomes severely limited. The highest  $N_{\rm T}$  values recorded 19 were in grass dominated communities in Svalbard underneath cliffs, perhaps because of nutrient 20 enrichment by sea birds or run off from snow melt from plateaus above.

21

22 Predicting L<sub>T</sub> from NDVI: the importance of bryophytes

23 Comparing the relationship between P and  $L_T$  required indirect measurements of  $L_T$  based on

1 *NDVI*. The calibration relationships between  $L_{\rm T}$  and *NDVI* suggests that including information on 2 moss abundance in the high arctic can improve the accuracy of  $L_{\rm T}$  estimates. We found a positive 3 relationship between the % cover of mosses and the NDVI of the ground surface after the 4 vascular canopy had been removed. Including the effect of mosses on background NDVI reduced 5 error in  $L_{\rm T}$  prediction in all cases. This improvement is surprising considering the inaccuracies 6 inherent in visually estimating percent cover, and the probable variability in moss NDVI with 7 water content and species. We deal with only peak growing season data in this study, but we 8 expect that the effect of mosses on NDVI will be greater at the start and end of the growing 9 season, when more of the ground surface is visible from above. The large abundances of mosses 10 at Zackenberg (pers. obs) may also act to increase the impact of structural vegetation properties 11 on the  $L_{\rm T}$ -NDVI relationship (Steltzer *et al.*, 2006), because of the effects of canopy structure on 12 the degree to which underlying mosses are visible from above. This could also simply be because 13 the vascular vegetation was most strongly patchy at Zackenberg, where there were more sites 14 with a high proportion of bare ground. Without including information on moss cover at 15 Zackenberg, NDVI was able to explain only 37 % of the variation in  $L_{\rm T}$ .

16

#### 17 The relationship between L<sub>T</sub> and P across Arctic latitudes

We found no significant improvement in model fit by including vegetation type in models of  $P_{1000}$  and  $E_0$ . This is unsurprising as functional convergence of PFTs in Arctic vegetation has previously been reported, and we argue that this is a direct result of the close convergence of the  $L_T-N_T$  relationship within a site (Williams *et al.*, 2006, Street *et al.*, 2007, Shaver *et al.*, 2007). We found a significant improvement in the  $P_{1000}$  model fit by including site level effects on  $\gamma p$ and  $p_m$ , though this improvement was small (RMSE was reduced by < 0.2 µmol m<sup>-2</sup> s<sup>-1</sup>). The effects of soil moisture and vapour pressure deficit will also control *P* through effects on stomata conductance (Rastetter *et al.*, 2010) contributing to noise in the relationship between *P* and  $L_T$ and masking the effect of site in this simple model. We suggest however that an analysis which takes into account changes in stomatal conductance, might reveal clearer site level differences in the relationship between *P* and  $L_T$ .

6

### 7 Implications

8 We show that, at peak growing season, Arctic plant canopies follow a narrowly-9 constrained set of rules that dictate the development of leaf area with respect to canopy nitrogen. 10 The way coupling is achieved however, differs between sites – with a trade-off at the canopy 11 scale between canopy leaf N concentration and leaf thickness. The  $L_{\rm T}$ - $N_{\rm T}$  relationship represents 12 an extremely powerful tool in up-scaling leaf level processes to canopies and ecosystems, by 13 collapsing large amounts of variation in leaf properties within and between species. 14 The mechanism by which  $L_{\rm T}$ - $N_{\rm T}$  coupling is achieved is unknown. A study by Campioli 15 et al. (2009a) showed that the ratio of canopy  $L_T:N_T$  in tundra plant communities is well-16 constrained through time, from 2 weeks after bud burst to 2 weeks before senescence. This 17 suggests either a 'hard-wired' community composition which results in convergence as soon as 18 leaves emerge, or that the reallocation of N within the canopy responds very rapidly to 19 environmental conditions and developing canopy structure. The next steps are to investigate the 20 mechanisms by which  $L_{\rm T}$ - $N_{\rm T}$  coupling is achieved and to test whether similar relationships exist 21 not only for Arctic plant canopies, but for plant canopies globally.

We achieved a small improvement in model fit by including site level parameters in the relationship between P and  $L_{\rm T}$ . After accounting for variation in stomatal conductance this 1 improvement could be more significant and warrants further investigation. There may be 2 potential for improvements in regional *P* models that are based on estimates of  $L_{\rm T}$ , through 3 incorporating differences in the  $L_{\rm T}$ - $N_{\rm T}$  relationship between sites and/or differences in moss 4 cover. Our results suggest that site based differences in the curvature of the  $L_{\rm T}$ - $N_{\rm T}$  relationship 5 are linked to the average fraction of incident radiation that is diffuse. If this is the case, there is 6 potential for improving carbon models without the need for further data input; canopy N 7 extinction parameters could be adjusted based on average incident radiation conditions.

8 Our results also imply that predicted increases in cloudiness as a result of Arctic climate 9 change (Vavrus et al., 2009) may result in redistribution of N within canopies and ecosystems. 10 This reallocation could be significant. If for example, we assume that diffuse fraction directly controls the parameter  $\gamma$ , Fig. 2a indicates that for a canopy with an  $L_{\rm T}$  of 1.5 m<sup>2</sup> m<sup>-2</sup> an increase 11 in average diffuse fraction from roughly 66 % (Toolik) to 80 % (Svalbard) might result in the 12 'optimal' total foliar nitrogen content doubling; from around 2 to 4 g N m<sup>-2</sup>. Re-allocation of N 13 14 on this magnitude would have important implications for other ecosystem processes such as litter 15 decomposition, herbivory, and belowground carbon allocation.

16

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## SUPPORTING INOFORMATION LEGENDS

3	Fig. S1. Global short wave (SW) vs. photosynthetically active radiation (PPFD) data from Toolik
4	Lake Field station for 12 <sup>th</sup> July to 13 <sup>th</sup> September and 24 <sup>th</sup> October to 26 <sup>th</sup> November 2008
5	
6	
7	Fig. S2. a) Modelled and measured total daily diffuse radiation through time at a Abisko, June
8	2008
9	
10	Fig. S3. a) Modelled and measured total daily diffuse radiation through time at Toolik Lake, for

11 a 8 week period in late summer 2008.

# TABLES

Table 1. List of symbols, abbreviations and units

Symbol	Definition	Units
L <sub>T</sub>	Leaf area index	$m^2$ leaf $m^{-2}$ ground
$N_{\mathrm{T}}$	Total foliar nitrogen per unit ground	$g N m^{-2}$ ground
$M_{ m T}$	Total leaf mass per unit ground	g leaf $m^{-2}$ ground
$N_{ m L}$	Nitrogen per unit leaf area	$g N m^{-2}$ leaf area
$N_{ m M}$	Nitrogen per unit leaf mass	$g N g^{-1}$ leaf mass
$L_{\mathrm{M}}$	Specific leaf area	$m^2 leaf g^{-1} leaf$
Ι	Photosynthetic photon flux density	μmol photons m <sup>-2</sup> ground s <sup>-1</sup>
Р	Gross primary productivity	$\mu$ mol CO <sub>2</sub> m <sup>-2</sup> ground s <sup>-1</sup>
NDVI	Normalised difference vegetation index	unitless
<i>NDVI</i> <sub>post</sub>	NDVI measure after vascular canopy removal	unitless
<i>NDVI</i> <sub>min</sub>	Minimum measured value of NDVI <sub>post</sub> (= $0.24$ )	unitless
<i>NDVI</i> <sub>max</sub>	Fitted parameter representing maximum canopy NDVI	unitless
$K_{ m c}$	Fitted extinction coefficient controlling curvature of NDVI-L <sub>T</sub> calibration	$m^2$ ground $m^{-2}$ leaf
a	Intercept of relationship between NDVI <sub>min</sub> and bryophyte cover	unitless
$B_{ m c}$	Slope of relationship between <i>NDVI</i> <sub>min</sub> and bryophyte cover	⁰∕₀ <sup>−1</sup>
$F_{\rm c}$	Net ecosystem exchange	$\mu$ mol CO <sub>2</sub> m <sup>-2</sup> ground s <sup>-1</sup>
$R_{ m E}$	Ecosystem respiration	$\mu$ mol CO <sub>2</sub> m <sup>-2</sup> ground s <sup>-1</sup>
$P_{\rm max}$	Theoretical light saturated photosynthetic rate	$\mu$ mol CO <sub>2</sub> m <sup>-2</sup> ground s <sup>-1</sup>
$E_0$	Initial light use efficiency of photosynthesis	$\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup> photons
$P_{1000}$	Gross primary productivity at 600 $\mu$ mol m <sup>-2</sup> I	$\mu$ mol CO <sub>2</sub> m <sup>-2</sup> ground s <sup>-1</sup>
γ	Nitrogen extinction coefficient	$m^2$ ground $m^{-2}$ leaf
$N_0$	Top of canopy $N_{\rm L}$	g N m <sup>-2</sup> leaf area
γ <sub>p</sub>	P extinction coefficient	$m^2$ ground $m^{-2}$ leaf
$P_0$	Top of canopy $P_{1000}$ per unit leaf area	$\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup> m <sup>-2</sup> leaf s <sup>-1</sup>
$P_{\rm m}$	Constant representing moss photosynthesis	$\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup> m <sup>-2</sup> s <sup>-1</sup>
α	Slope of relationship between $E_0$ and $L_T$	$\mu$ mol CO <sub>2</sub> $\mu$ mol <sup>-1</sup> photons m <sup>-2</sup> leaf
$E_{ m m}$	Constant in relationship between $E_0$ and $L_T$ representing moss $P$	μmol CO <sub>2</sub> μmol <sup>-1</sup> photons

Data type (plot sizes)	Year	Site	n	Source
harvest $(0.03 \text{ m}^{2})$	2009	Barrow, Alaska	23	This study
harvest $(0.03 \text{ m}^2, 0.05 \text{ m}^2, 0.09 \text{ m}^2)$	2006	Zackenberg, NE Greenland	28, 26, 24	This study
harvest $(0.05 \text{ m}^2)$	2005	Longyearbyen, Svalbard	48	This study
harvest $(0.04 \text{ m}^2)$	2001	Abisko, Sweden	92	Van Wijk <i>et al.</i> (2005)
harvest $(0.04 \text{ m}^2)$	1997	Kuparuk watershed, Alaska	94	Williams & Rastetter (1999)
Chamber $P(1 \text{ m}^2)$	2009	Barrow, Alaska	10	This study
Chamber $P (1 \text{ m}^2, 0.09 \text{ m}^2)$	2006	Zackenberg, NE Greenland	26, 15	This study
Chamber $P(1 \text{ m}^2)$	2005	Longyearbyen, Svalbard	36	This study
Chamber $P(1,m^2)$	2004	Toolik Laka Alaska	28	Shaver <i>et al.</i> (2007)
Chamber F (1 m)	2004	TOOTIK Lake, Alaska		& Street et al. (2007)
Chamber $P(1 \text{ m}^2)$	2004	Abisko, Sweden	45	Shaver <i>et al.</i> (2007)

**Table 2.** Summary of data including year, site, number of measurements and literature source.

**Table 3**. Average specific leaf area ( $L_M$ ), leaf nitrogen per unit leaf area ( $N_L$ ) and leaf nitrogen per unit leaf mass ( $N_M$ ) for plant species at Svalbard, Zackenberg and Barrow. Species for which n < 5 are not included. Plant functional types (PFT) are: D = deciduous, E = evergreen, F = forb, G = graminoid.

Site	Species	n	PFT	$N_{\rm L} [{\rm g}{\rm m}^{-2}]$	$N_{\rm M}  [{ m mg g}^{-1}]$	$L_{\rm M}  [{ m m}^2  { m kg}^{-1}]$
Barrow	Salix phlebophylla	7	D	$1.76\pm0.09$	$23.3\pm1.8$	$13.2\pm0.3$
	Stellaria species	9	F	$1.43\pm0.17$	$22.4\pm1.6$	$17.4 \pm 2.2$
	Saxifraga cernua	8	F	$0.95\pm0.07$	$19.4 \pm 1.3$	$20.6\pm1.0$
	Dupontia fisheri	15	G	$2.21\pm0.17$	$23.1\pm1.0$	$11.2\pm0.8$
	Eriophorum scheuchzeri	10	G	$2.51\pm0.23$	$30.3\pm1.9$	$12.5\pm0.6$
	other grasses	9	G	$1.73\pm0.14$	$19.3\pm1.0$	$11.5\pm0.6$
Svalbard	Salix polaris	34	D	$1.75\pm0.05$	$24.2\pm0.8$	$13.9\pm0.4$
	Dryas octopetala	21	Ε	$2.68\pm0.13$	$20.2\pm0.6$	$8.1\pm0.8$
	Cassiope tetragona	6	Ε	$2.16\pm0.20$	$19.8\pm1.2$	$4.6\pm0.2$
	Polygonum viviparum	35	F	$2.06\pm0.07$	$30.8\pm0.3$	$15.5\pm0.5$
	other forbs	20	F	$1.88\pm0.20$	$24.9 \pm 1.1$	$15.1 \pm 1.3$
	Equisetum species	30	Р	$3.14\pm0.18$	$28.4 \pm 1.3$	$9.5\pm0.5$
	Carex species	9	G	$1.90\pm0.17$	$19.3\pm1.3$	$10.6\pm0.9$
	Dupontia fisheri	5	G	$2.09\pm0.20$	$16.1\pm1.3$	$7.8\pm0.2$
	other graminoids	32	G	$2.32\pm0.10$	$21.6\pm0.6$	$9.7\pm0.3$
Zackenberg	Salix arctica	58	D	$2.18\pm0.08$	$30.9\pm1.0$	$14.4\pm0.6$
	Arctostaphylos alpina	14	D	$2.87\pm0.20$	$31.1\pm2.0$	$11.1\pm0.6$
	Vaccinium uliginosum	14	D	$1.44\pm0.05$	$24.6\pm1.3$	$16.2 \pm 1.5$
	Dryas species	23	Ε	$2.12\pm0.09$	$16.0\pm0.7$	$8.0\pm0.5$
	Cassiope tetragona	22	Ε	$2.15\pm0.19$	$23.2\pm0.6$	$5.8\pm0.4$
	Polygonum viviparum	37	F	$4.01 \pm 1.09$	$38.3 \pm 1.0$	$15.9\pm1.3$
	Stellaria species	16	F	$2.15\pm0.57$	$27.0\pm1.5$	$17.6\pm1.9$
	Pedicularis species	8	F	$2.40\pm0.23$	$36.2\pm3.7$	$16.1\pm2.5$
	Equisetum species	12	Р	$2.62\pm0.10$	$27.3 \pm 1.9$	$10.4\pm0.5$
	Carex species	8	G	$2.17\pm0.22$	$24.9 \pm 1.9$	$11.8\pm0.8$
	Dupontia species	6	G	$2.45\pm0.21$	$21.3\pm2.7$	$8.6\pm0.7$
	Eriophorum species	6	G	$2.47\pm0.12$	$25.9 \pm 1.6$	$10.7\pm0.9$
	other graminoids	28	G	$4.47\pm2.03$	$24.9 \pm 1.4$	$11.1 \pm 1.1$
	other grasses	18	G	$5.07 \pm 1.69$	$28.6 \pm 1.6$	$8.2\pm0.9$
	other sedges	5	G	$2.25\pm0.26$	$24.9\pm2.1$	$11.5 \pm 1.2$

**Table 4**. Model description and structure, Akaike's information criteria (AIC), Bayes Information Criteria (BIC), logliklihood (LogLik) and root mean square error (RMSE) for alternative  $N_T$  models (equation 4) describing the data shown in Figure 2a. Parameters are fitted either for the whole data set, or separately for each site or PFT. " - " indicates that grouping factors for site or PFT are not included i.e. there is a single fitted value of the parameter for the whole data set.

$N_{\rm T}$ model description	Grouping structure		Number of parameters	AIC	BIC	LogLik	RMSE [g N m <sup>-2</sup> ground]
	$N_0$	γ	-				-0 0 -
No site or PFT effect	-	-	3	5.9	21.0	$1.06^{*+}$	0.27
Site effects (on $N_0$ and $\gamma$ )	Site	Site	15	-201.3	-140.9	$153.1^{+\& \pounds}$	0.18
Site effects (on $N_0$ only)	Site	-	11	-171.4	-129.1	97.7 <sup>&amp;</sup>	0.23
Site effects (on $\gamma$ only)	-	Site	11	-180.9	-135.5	$102.4^{\text{\pounds}}$	0.19
PFT effects (on $N_0$ and $\gamma$ )	PFT	PFT	15	-86.2	-25.7	59.1 <sup>*</sup>	0.25

£&\$\*+ matching symbols indicates pairs of nested models where the log likelihood ratio test is significant at 0.05 level.

**Table 5**. Model description and structure, Akaike's information criteria (AIC), Bayes Information Criterion (BIC), loglikelihood (LogLik) and root mean square error (RMSE) for alternative  $P_{1000}$  models (equation 6) describing the data displayed in Figure 6a. Parameters are fitted either for the whole data set, or separately for each site/PFT " - " indicates that no grouping factors are included in the model for that parameter i.e. there is a single fitted value of the parameter for the whole data set.

P <sub>1000</sub> model description	Grouping structure		Number of parameters	AIC	BIC	LogLik	RMSE [µmol m <sup>-2</sup> s <sup>-1</sup> ]	
	$P_0$	$\gamma_{\rm p}$	$P_{\rm m}$					·
No site or PFT effect	-	-	-	4	356.9	369.2	-173.4*\$+	1.89
Site effects (on $P_0$ and $\gamma_p$ )	Site	Site	-	12	354.3	386.1	-164.1 <sup>*</sup>	1.73
Site effects (on $\gamma_p$ and $P_m$ )	-	Site	Site	12	353.8	385.8	-163.9 <sup>\$#^</sup>	1.71
Site effects (on $\hat{P}_{m}$ only)	-	-	Site	8	355.2	377.3	$-168.6^{+^{+^{+^{+^{+^{+^{+^{-}}}}}}}$	1.81
Site effects (on $\gamma_p$ only)	-	Site	-	8	358.9	381.0	$-170.5^{\#}$	1.81
Site effects (on $P_0$ only)	Site	-	-	8	357.8	379.9	-169.9	1.81
PFT effects (on $P_0$ and $\gamma_p$ )	PFT	PFT	-	10	360.5	387.5	-169.2	1.77
PFT effects (on $\gamma_p$ and $P_m$ )	-	PFT	PFT	10	358.1	385.1	-168.0	1.78

Pairs of matching symbols (\* + #) indicates nested models where the log likelihood ratio test is significant at 0.05 level. For all light curves maximum *I* exceeded 1000 µmol m<sup>-2</sup> s<sup>-1</sup>

**Table 6**. Model structure, Akaike's information criteria (AIC), Bayes information criterion (BIC), loglikelihood (LogLik) and the root mean square error (RMSE) for alternative  $E_0$  models (equation 7) describing the data displayed in Figure 6b. Parameters are fitted either for the whole data set, or separately for each site/PFT. " - " indicates that no grouping factors are included in the model for that parameter. i.e. there is a single fitted value of the parameter for the whole data set

$E_0$ model description	Grouping structure		Number of parameters	AIC	BIC	LogLik	RMSE [µmol m <sup>-2</sup> s <sup>-1</sup> ]
	α	$E_{ m m}$					·
No site or PFT effect	-	-	3	-638.2	-628.4	323.1	0.0060
Site effects (on $\alpha$ and $E_{\rm m}$ )	Site	Site	11	-628.6	-599.2	326.3	0.0057
Site effects (on $\alpha$ only )	-	Site	7	-631.7	-612.1	323.9	0.0059
Site effects (on $E_{\rm m}$ only )	Site	-	7	-632.6	-612.9	324.3	0.0060
PFT effects (on $\alpha$ and $E_{\rm m}$ )	PFT	PFT	9	-630.9	-606.4	325.5	0.0058

For all light curves maximum I exceeded 600  $\mu$ mol m<sup>-2</sup> s<sup>-1</sup>



**Figure 1.** Post harvest NDVI ( $NDVI_{post}$ ) and percent moss cover (Bc) for Zackenberg and Svalbard.



**Figure. 2.** Inter-comparison of relationships between a)  $N_T$  and  $L_T$  b)  $N_T$  and  $M_T$  c)  $L_T$  and  $M_T$  and d)  $M_T/L_T$  and  $N_T/M_T$  for vegetation harvests at Svalbard 2005 (n = 49), Abisko 2002 (n = 92) and Toolik 1997 (n = 94), Zackenberg (n = 78) and Barrow (n = 23).



**Figure 3.** Relationships between  $N_{\rm T}$  and  $L_{\rm T}$  (left panels) and  $N_{\rm T}$  and  $M_{\rm T}$  (right panels) for vegetation harvests in **a,b**) Svalbard 2005 (n = 49), **c,d**) Zackenberg 2006 (n = 78) and **e,f**) Barrow 2009 (n = 23). Data points classified by vegetation type.



**Figure 4** Site-specific fitted values of **a**)  $N_0$  vs. total growing season short wave radiation and **b**)  $\gamma$  versus diffuse radiation fraction for Abisko, Barrow, Toolik, Svalbard and Zackenberg. Horizontal error bars are standard deviation for 5 years of radiation data. Vertical error bars are 90 % confidence interval for fitted parameters.



**Figure 5** The relationship between NDVI and  $L_{\rm T}$  for the destructively harvested plots (left panels) and between modelled and measured  $L_{\rm T}$  for the same plots, based on equation 2 fitted separately for each plant functional type (right panels) **a,b**) Svalbard, **c,d**) Zackenberg and **e,f**) Barrow. Open symbols show plots where moss cover is > 50 %, filled symbols show plots where moss cover < 50 %.



Figure 6 a) The relationship between  $P_{1000}$  and  $L_T$  and b)  $E_0$  and  $L_T$  for Svalbard, Abisko Toolik, Zackenberg and Barrow.

#### SUPPORTING INFORMATION



**Fig. S1.** Global short wave (SW) vs. photosynthetically active radiation (PPFD) data from Toolik Lake Field station for 12<sup>th</sup> July to 13<sup>th</sup> September and 24<sup>th</sup> October to 26<sup>th</sup> November 2008



Fig. S2. a) Modelled and measured total daily diffuse radiation through time at a Abisko, June 2008



Fig. S3. a) Modelled and measured total daily diffuse radiation through time at Toolik Lake, for a 8 week period in late summer 2008.