

# Use of data assimilation methods in coastal sediment models

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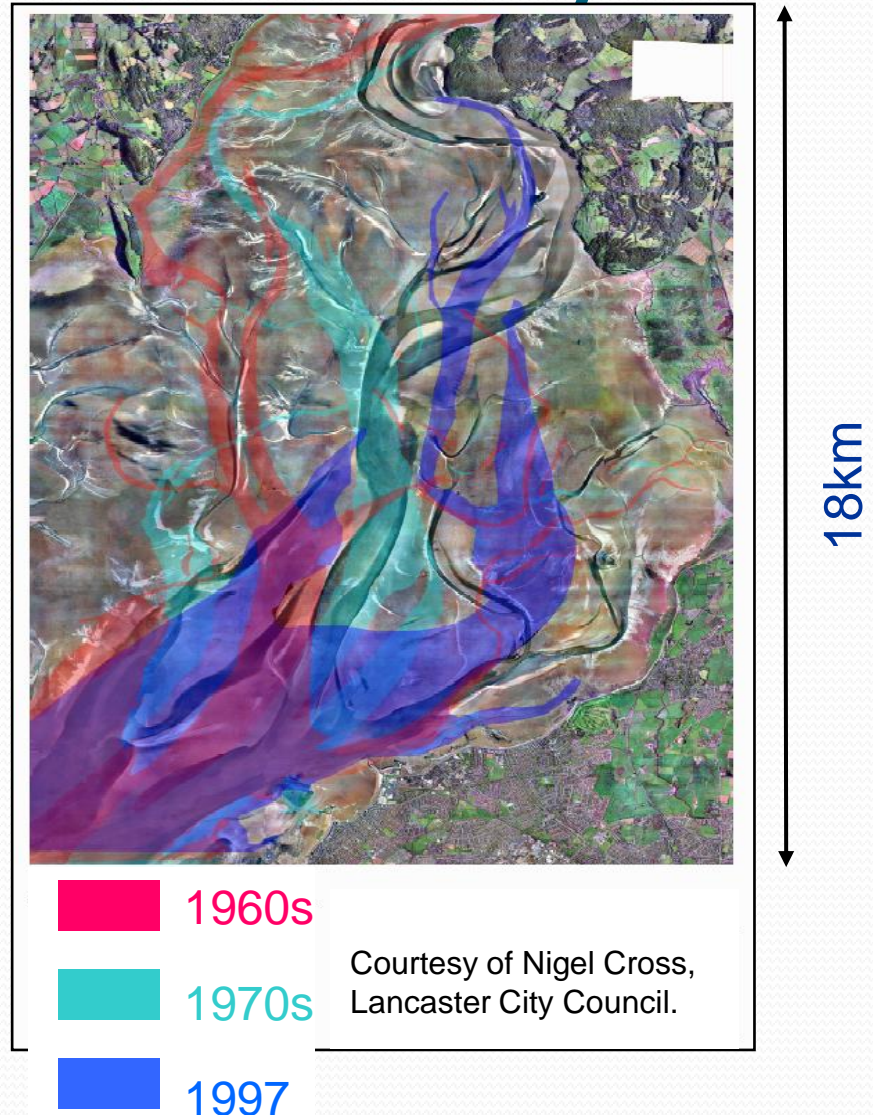
With thanks to Kevin Horsburgh, POL and the Environment Agency

# Outline

- Motivation: morphodynamics and flooding
- Observational data and morphodynamic model
- Data assimilation
- Parameter estimation
- Conclusions

# Case study: Morecambe Bay

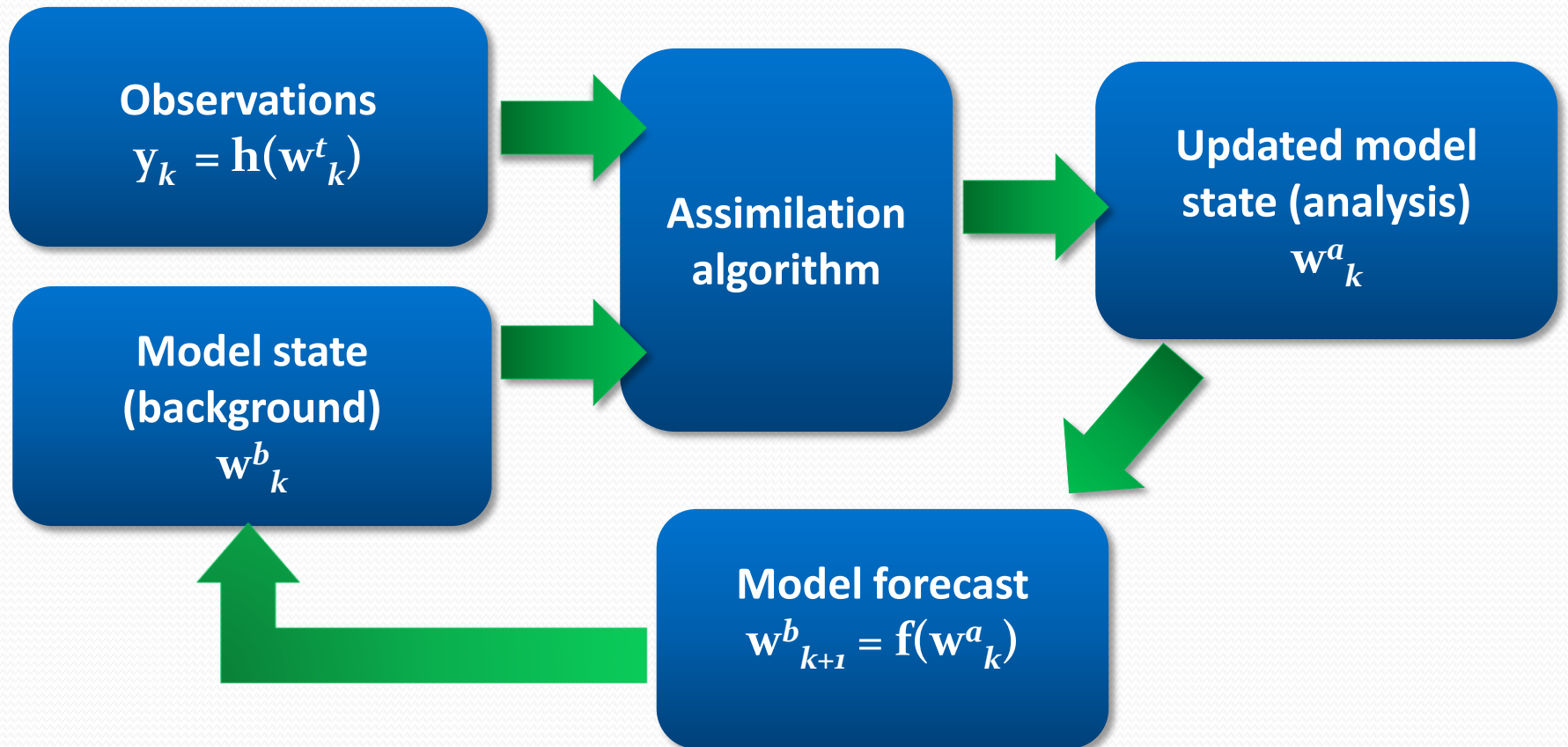
- Channels in Morecambe Bay can move by several kilometres in just a few years
- Channel movement
  - impacts on habitats in the bay
  - affects access to ports
  - has implications for flooding during storm events
- For example, Morecambe can be flooded by storm waves propagating up the deep-water channels



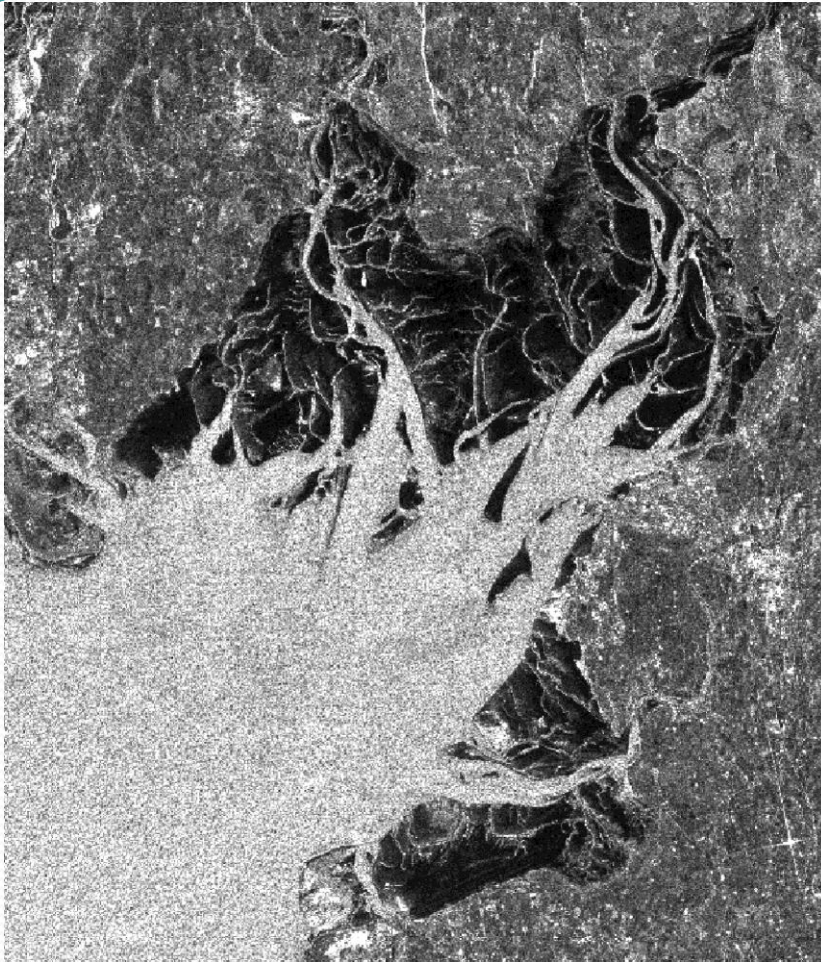
# The general idea

- Coastal bathymetry is dynamic and evolves with time
  - water action erodes, transports, and deposits sediment, which changes the bathymetry, which alters the water action, and so on
- Accurate knowledge of coastal bathymetry at the time of a storm event would allow improved flood forecasting using coastal flood inundation models
  - but it is impractical to continually monitor large coastal areas in anticipation of a storm
- A solution may be to run an operational coastal area morphodynamic model
  - and keep the model on track using data assimilation
- As observations become available they can be used to nudge model bathymetry back towards true bathymetry
  - these observations may be infrequent and only sample a small part of the model domain
- Data assimilation can also be used for parameter estimation

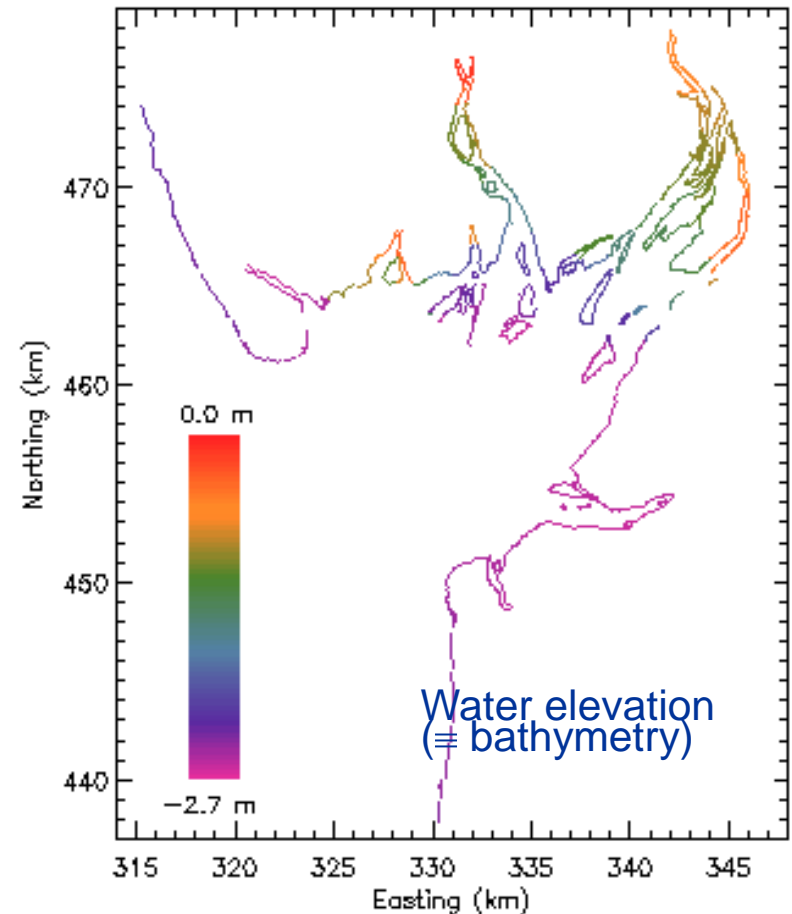
# Sequential data assimilation



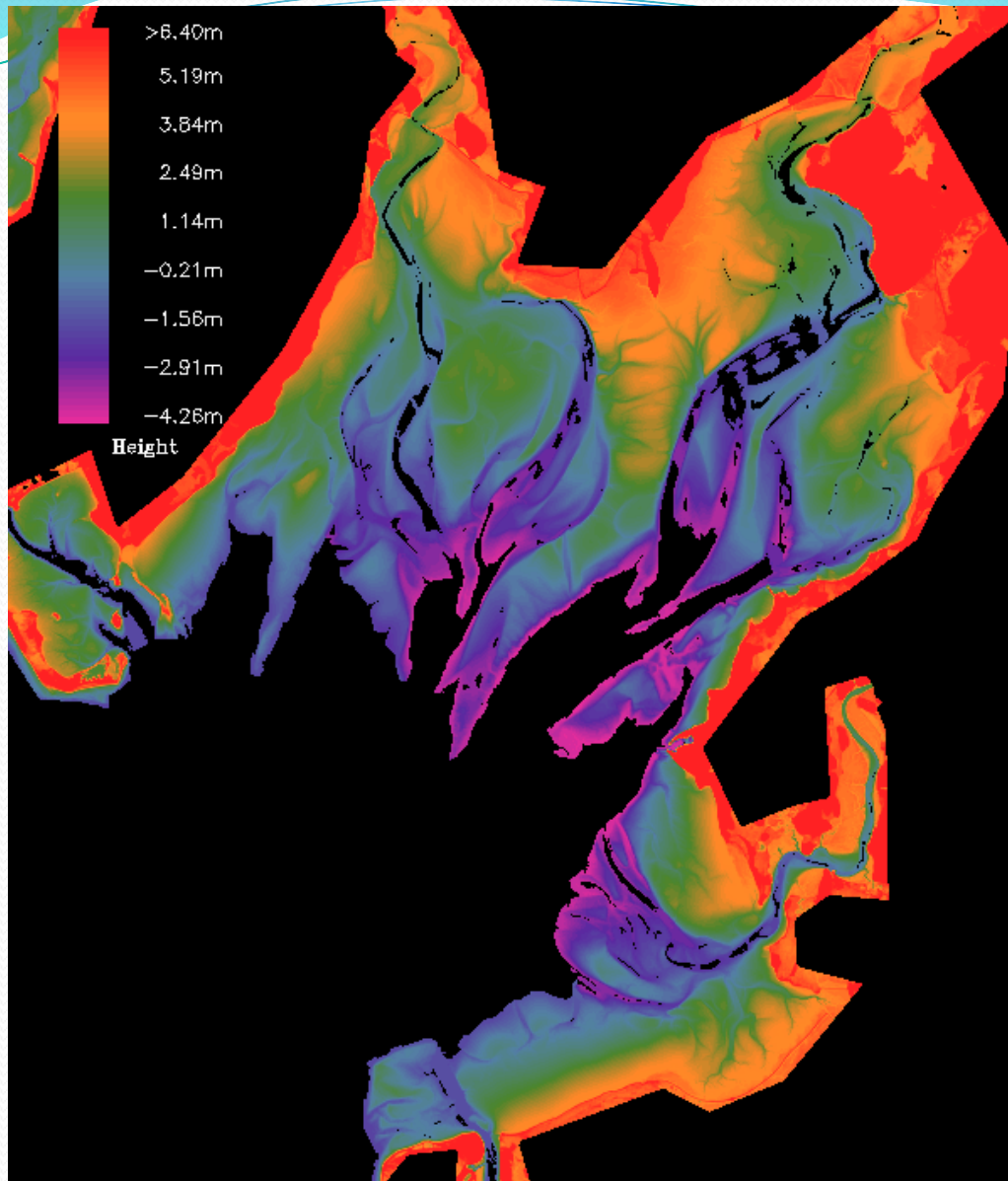
# Waterline data for Morecambe Bay



An image of the land-sea boundary  
(satellite SAR image © ESA)

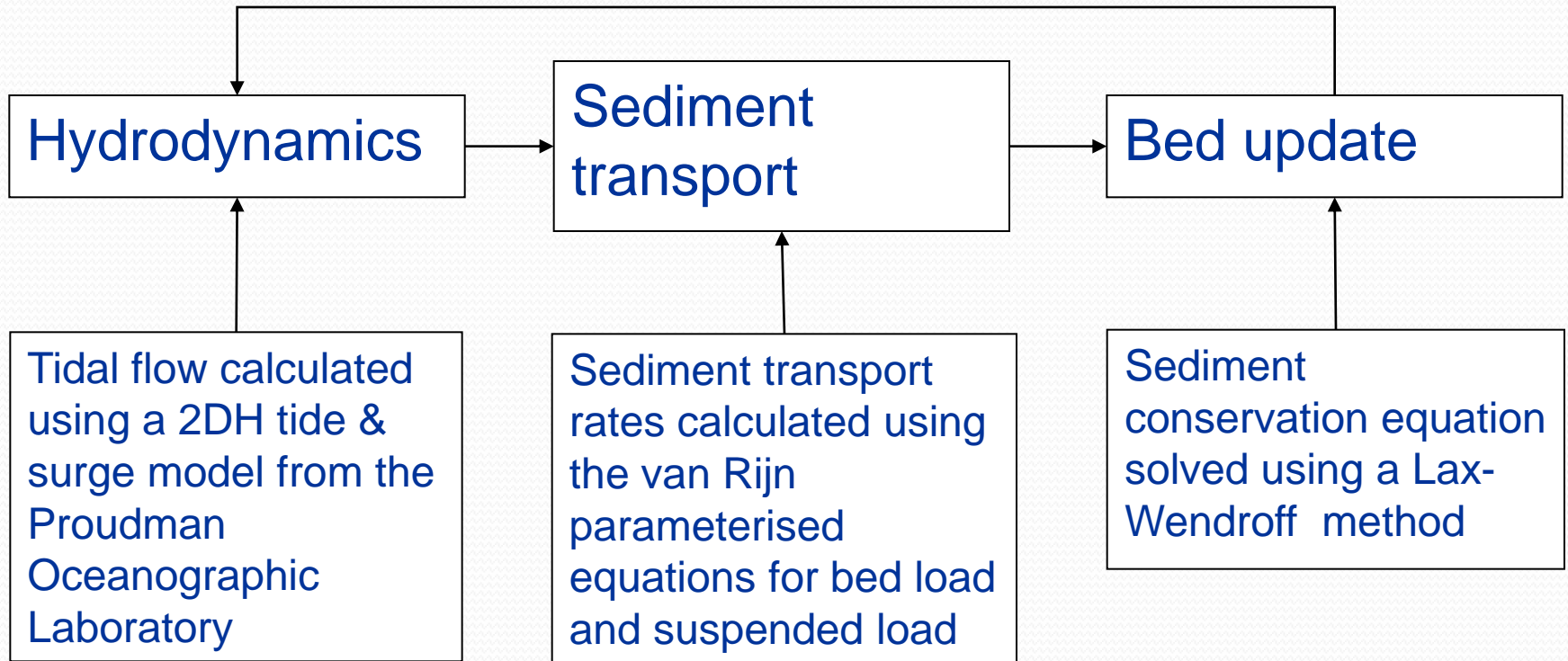


Water elevation predicted along the boundary by a hydrodynamic model with surge component



*LiDAR data from  
15/11/05 to be  
used for  
validation.*

# Morphodynamic model

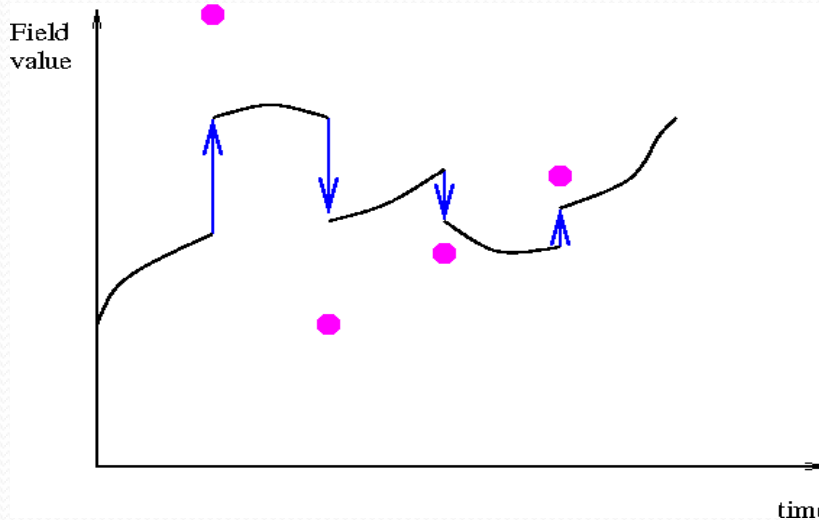


Spatial resolution ~ 240m

A simple model, relative to state-of-the-art engineering models, but adequate for assessing the benefits of data assimilation



# 3D Var data assimilation

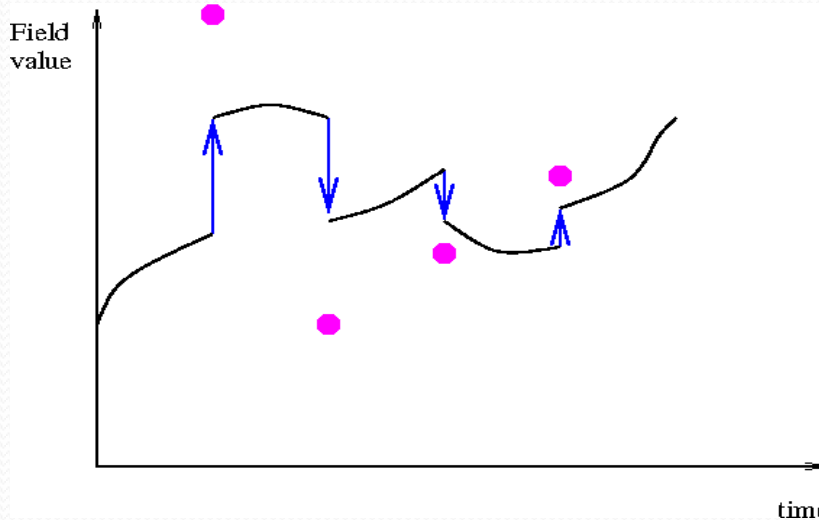


- Analysis state is found by minimising a cost function

$$J(\mathbf{w}_k) = \underbrace{(\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b)}_{\text{background term}} + \underbrace{(\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))}_{\text{observation term}}$$

$\mathbf{B}$  and  $\mathbf{R}$  are the covariance matrices of the background and observation errors

# 3D Var data assimilation



- Analysis state is found by minimising a cost function

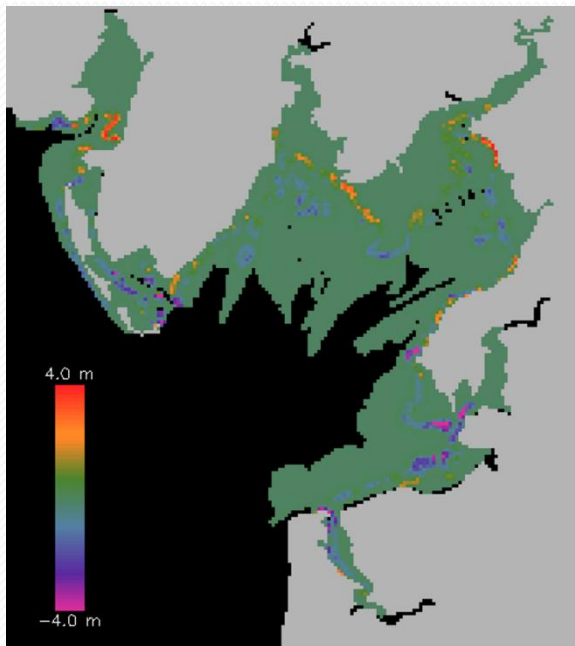
$$J(\mathbf{w}_k) = (\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b) + (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))$$

background term observation term

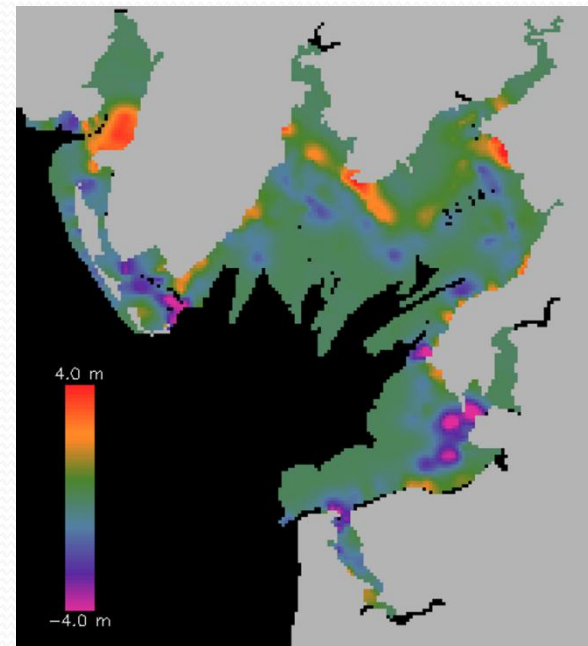
**B** and **R** are the covariance matrices of the background and observation errors  
Choice of matrix is crucial to success of assimilation scheme

# Choosing the lengthscale for B

Increment (Analysis-Background) for assimilation of 1 waterline

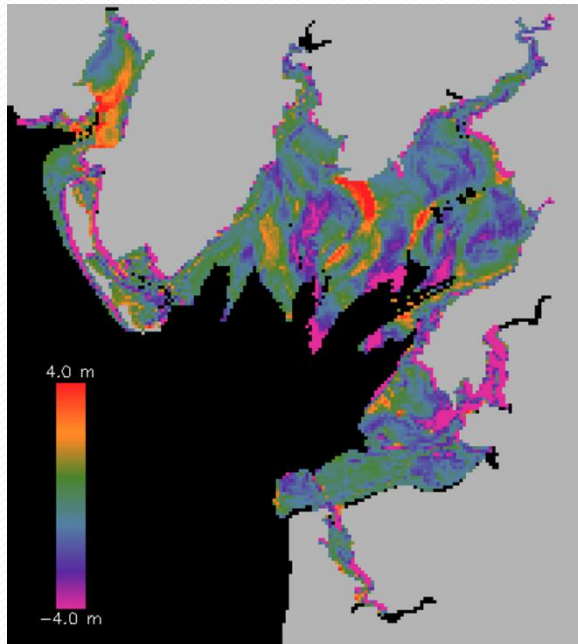


$L = 0.5$

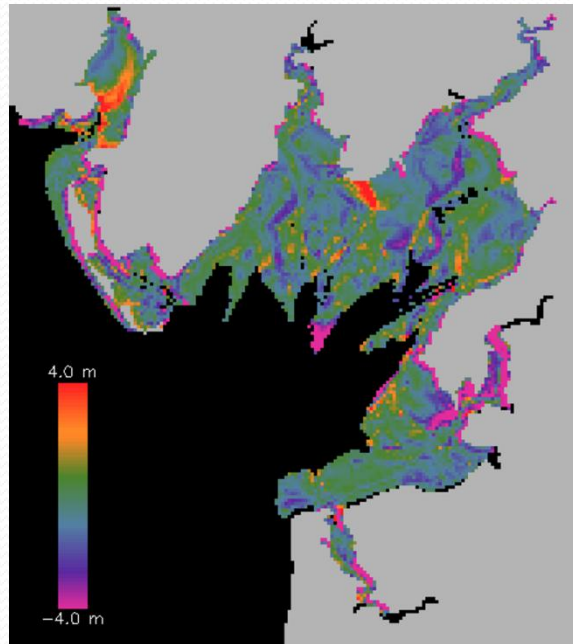


$L = 3.0$

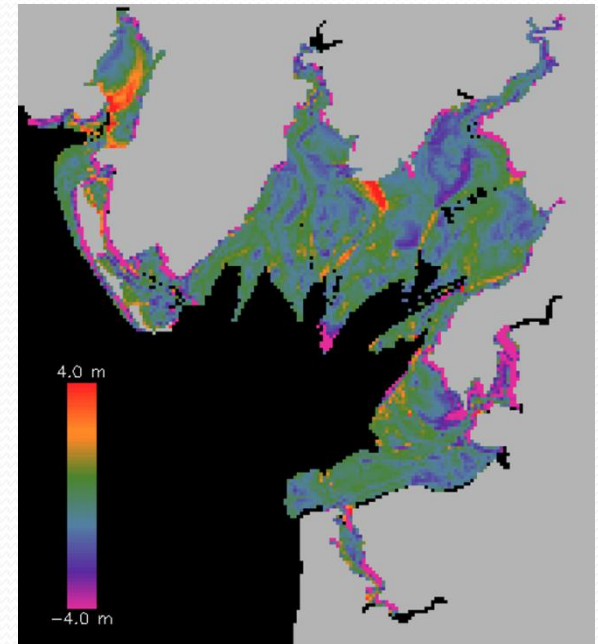
# Comparison with lidar validation data



No assimilation



With assimilation  
Calibrated params



With assimilation  
Weighted ensemble mean  
Parameter ensemble

# Parameter estimation

- Numerical models suffer from errors in their initial conditions and parameters.
- Initial conditions are often estimated by data assimilation; combining model predictions with observational data to produce an updated model state (the analysis) whilst keeping the model parameters fixed.
- Even with perfect initial data, inaccurate model parameters will lead to the growth of prediction errors.
- Using *state augmentation* in data assimilation we can estimate model parameters concurrently with the state.

# Joint state-parameter estimation

- Model state & parameter evolution

$$\mathbf{z}_{k+1} = \mathbf{f}(\mathbf{z}_k, \mathbf{p})$$

$$\mathbf{p}_{k+1} = \mathbf{p}_k.$$

- Augmented system model

$$\mathbf{w}_{k+1} = \begin{pmatrix} \mathbf{z}_{k+1} \\ \mathbf{p}_{k+1} \end{pmatrix} = \begin{pmatrix} \mathbf{f}(\mathbf{z}_k, \mathbf{p}_k) \\ \mathbf{p}_k \end{pmatrix} = \tilde{\mathbf{f}}(\mathbf{w}_k)$$

- Observations

$$\mathbf{y}_k = \tilde{\mathbf{h}}(\mathbf{w}_k) + \boldsymbol{\delta}_k = \tilde{\mathbf{h}} \begin{pmatrix} \mathbf{z} \\ \mathbf{p} \end{pmatrix} + \boldsymbol{\delta}_k = \mathbf{h}(\mathbf{z}_k) + \boldsymbol{\delta}_k$$

# Background error covariance

$$J(\mathbf{w}_k) = (\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b) + (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))$$

background term      observation term

State background error covariance

$$\mathbf{B}_k = \begin{pmatrix} \mathbf{B}_{zz_k} & \mathbf{B}_{zp_k} \\ (\mathbf{B}_{zp_k})^T & \mathbf{B}_{pp_k} \end{pmatrix}$$

state-parameter cross covariance

parameter background error covariance

For joint state-parameter estimation, it is important that the a priori cross-covariances between the parameters and the state are well specified.

# A hybrid approach

Combines ideas from 3D-Var and the extended Kalman filter (EKF)

- assumes  $\mathbf{B}_{zz}$  and  $\mathbf{B}_{pp}$  fixed
- uses a flow dependent state-parameter cross covariance  $\mathbf{B}_{zp_k}$

$$\mathbf{B}_{k+1} = \begin{pmatrix} \mathbf{B}_{zz} & \mathbf{N}_k \mathbf{B}_{pp} \\ \mathbf{B}_{pp} \mathbf{N}_k^T & \mathbf{B}_{pp} \end{pmatrix}$$

where  $\mathbf{N}_k = \left. \frac{\partial \mathbf{f}(\mathbf{z}, \mathbf{p})}{\partial \mathbf{p}} \right|_{\mathbf{z}_k^a, \mathbf{p}_k^a}$



# A simple sediment transport model

Based on the sediment conservation equation

$$\frac{\partial z}{\partial t} = - \left( \frac{1}{1 - \varepsilon} \right) \frac{\partial q}{\partial x}$$

with

$$q = Au^n$$

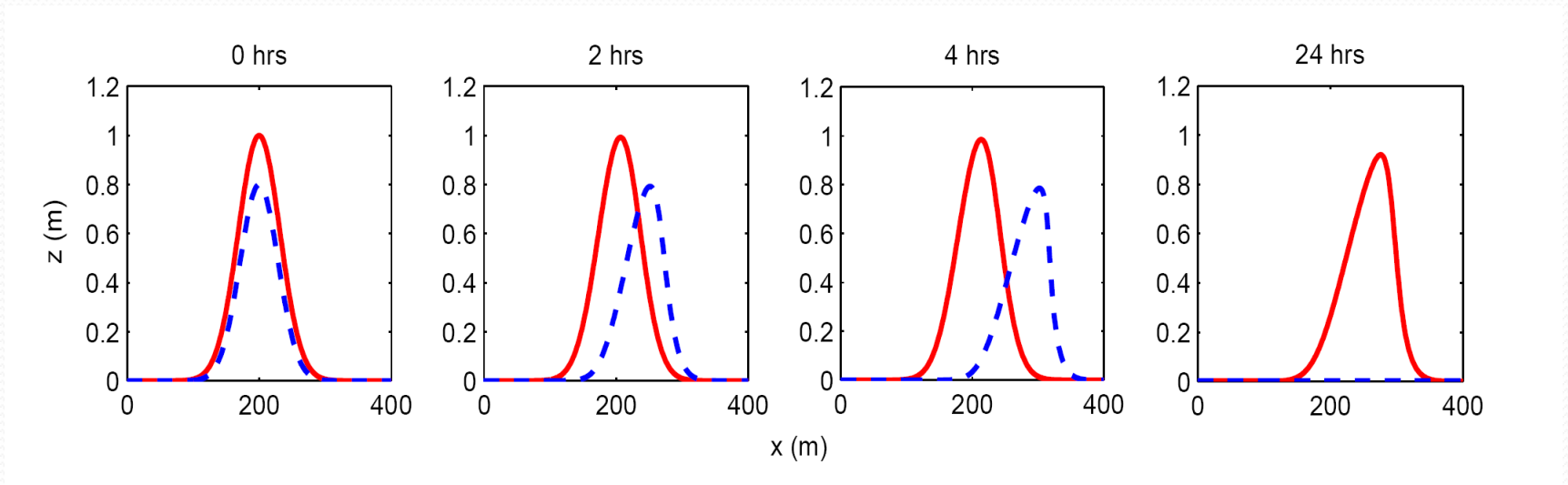
where  $z(x,t)$  is the bathymetry,  $t$  is time,  $\varepsilon$  is the sediment porosity,  $q$  is the sediment transport rate,  $u(x,t)$  is the depth averaged current and  $A$  and  $n$  are constant parameters.

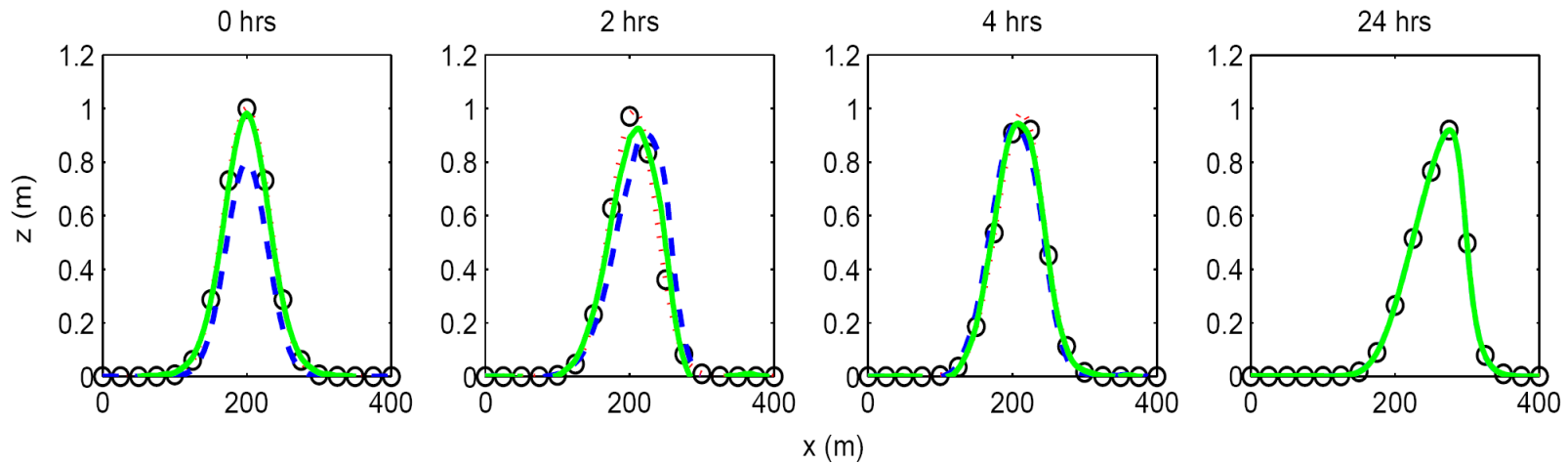
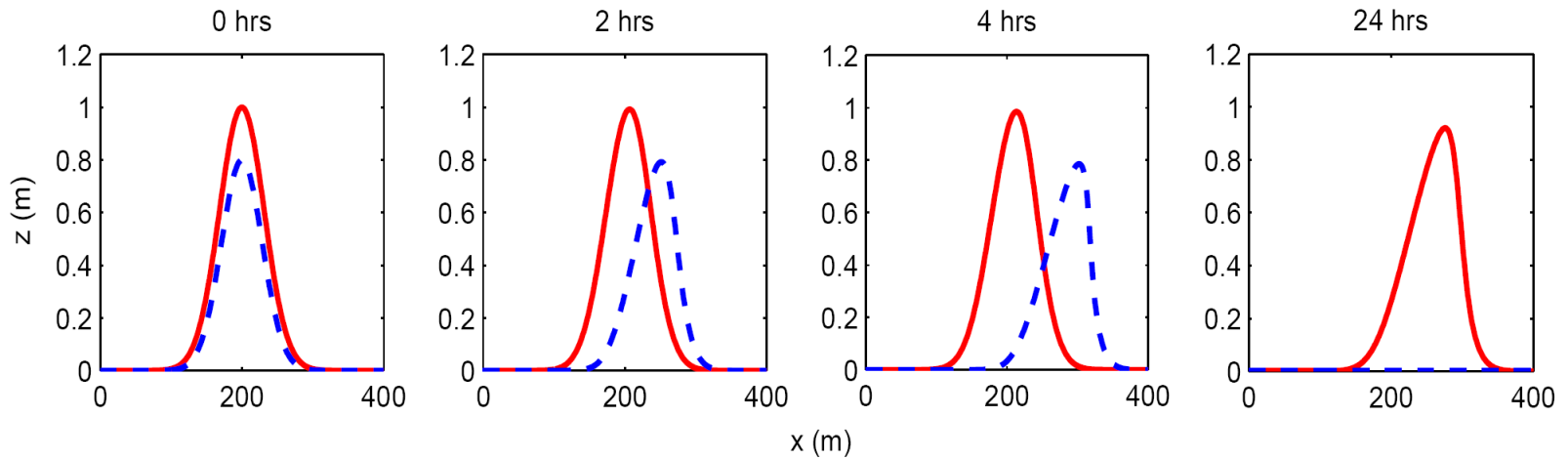
Can we use data assimilation to estimate the parameters  $A$  and  $n$ ?

# Experiments

- Identical twin:
  - reference solution generated using 'true' parameter values  $A = 0.002 \text{ ms}^{-1}$  and  $n = 3.4$
  - model then re-run with incorrect initial bathymetry and parameter values
- Observations assimilated sequentially at regular time intervals
  - taken from reference solution & assumed perfect
  - the 3D-Var cost function is minimized iteratively using a quasi-Newton descent algorithm
- Background error covariances
  - $\mathbf{B}_{zz}$  fixed
  - $\mathbf{B}_{zpk}$  time varying

# without data assimilation ...

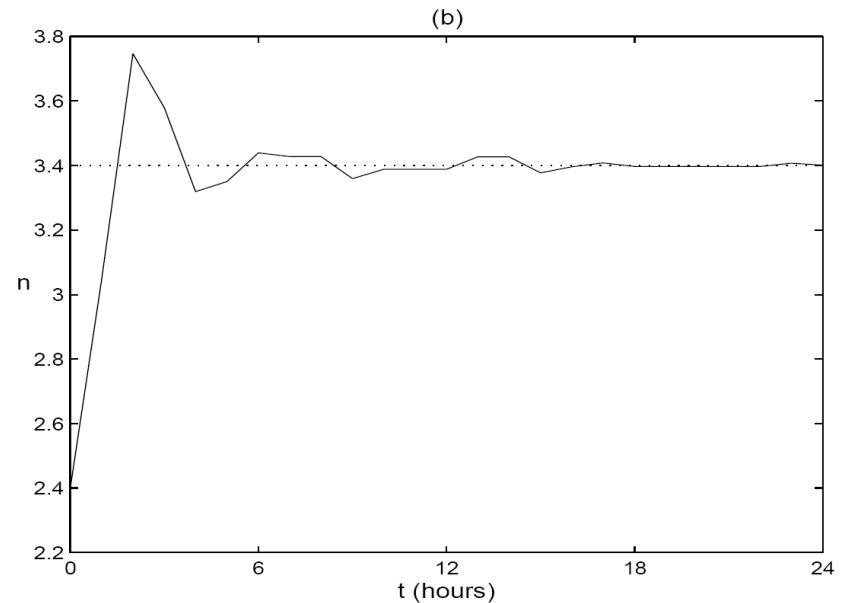
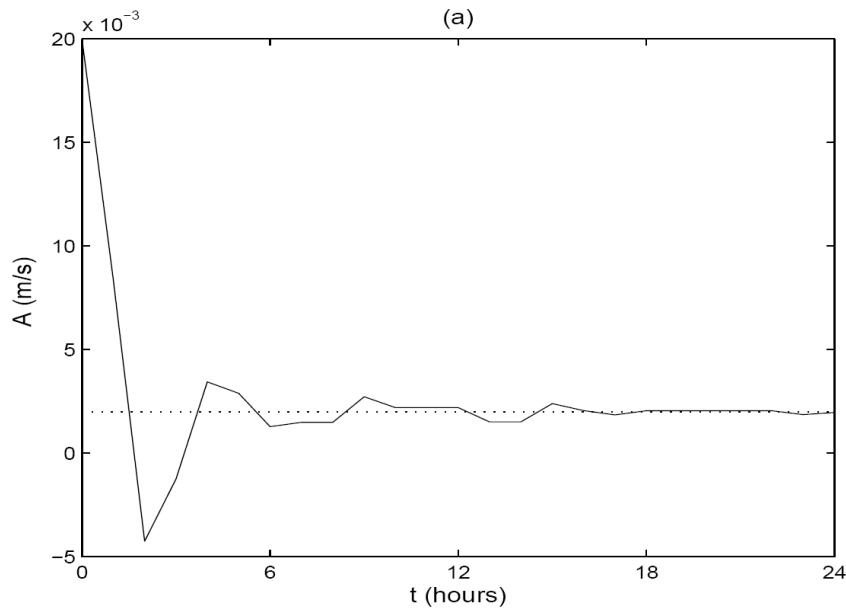




with data assimilation

# Parameter estimates

Initial estimates (a)  $A_o = 0.02 \text{ ms}^{-1}$  (b)  $n_o = 2.4$



$(A_{\text{true}} = 0.002 \text{ ms}^{-1}, n_{\text{true}} = 3.4)$

# Summary

- Up-to-date knowledge of near-shore coastal bathymetry is important in flood prediction and risk management
- Assimilated SAR waterline data into a model of Morecambe Bay to keep the model on track
- Best results are obtained using an ensemble of parameters
- Developed a new hybrid data assimilation scheme for joint estimation of model parameters and state.
- Recovers the true parameter values to a good level of accuracy, even when observations are noisy.
- Relatively simple to implement and computationally inexpensive to run
- Method also successfully applied to a range of simple dynamical system models.
- Expect this new technique to be easily transferable to more complex models.