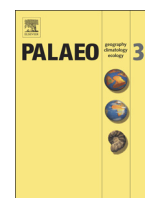


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## Middle to late Wisconsinan climate and ecological changes in northern Alaska: Evidences from the Itkillik River Yedoma

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### ABSTRACT

Continuous paleoenvironmental records covering the period prior to the Last Glacial Maximum in northeastern Beringia are sparse. This study presents a multi-proxy analysis of a 35-m-high yedoma exposure located on the right bank of the Itkillik River in Alaska. The exposure accumulated over 39 thousand years (kyr) during the Middle Wisconsinan Interstadial and the Late Wisconsinan glacial advance. We identified five stratigraphic units based on pollen, carbon and ice content, and isotopic composition ( $\delta^{18}\text{O}$ ) of the sediments. Middle Wisconsinan climate in northern Alaska promoted peat accumulation prior to 33.6 cal kyr BP (calibrated kyr before present). Reconstructed July air temperatures were 1–2 °C lower than current at 34.8 cal kyr BP, consistent with the timing of the interstadial climatic optimum in interior Alaska and Yukon. Colder (by up to 4 °C) and drier conditions characterize the transition from interstadial to glacial conditions between 33.6 and 29.8 cal kyr BP. Late Wisconsinan (29.8–17.9 cal kyr BP) July air temperatures were 2–3 °C lower than today, with grassland vegetation dominated by Poaceae, *Artemisia* and forbs, in contrast to the modern Cyperaceae dominance. Moister and warmer environmental conditions after 17.9 cal kyr BP correspond to the Late Glacial to Early Holocene interval.

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### 1. Introduction

Beringia is the high-latitude region extending from the Taymyr Peninsula in eastern Russia to the Mackenzie River in Canada that remained largely unglaciated throughout the Late Pleistocene glaciations (130 to 11.7 thousand years before present (kyr BP)) (Elias and Brigham-Grette, 2006).

During the Wisconsinan glaciation (71–11.7 kyr BP), Beringia was occupied by a now extinct biome formerly known as steppe-tundra (Yurtsev, 1982) or mammoth steppe (Guthrie, 1990). The plant assemblages and the amplitude of climate changes are thought to have been very heterogeneous across Beringia over the Wisconsinan (Anderson and Lozhkin, 2001), but graminoid and forb tundra growing under colder and drier climate than today seem to have dominated the Late

Glacial Maximum (LGM; 26.5–19.0 kyr BP) landscape (Bigelow et al., 2003).

Most of the eastern Beringia vegetation reconstructions based on pollen analysis were undertaken south of the Brooks Range from the end of the 1980s to the mid-1990s (e.g. Anderson, 1985, 1988; Anderson et al., 1988; Anderson and Brubaker, 1994). Overall, the pollen assemblages corresponding to the cold and dry late Pleistocene conditions in the lowlands of Beringia are dominated by Poaceae (Swanson, 2006). Recent macrofossil and DNA analyses have indicated that there might have been a higher proportion of forbs than was previously thought based on pollen analysis, which tends to underestimate actual vegetation diversity by over-representing high pollen-producers such as the herbs from the Poaceae, Cyperaceae and *Artemisia* genera (Willerslev et al., 2014).

A lot of uncertainty remains about climate and ecological changes over the Middle Wisconsinan Interstadial and the following Late Wisconsinan glacial advance. During the Middle Wisconsinan interstadial (57–29 kyr BP), part of western Beringia experienced temperatures that rose to nearly modern levels, thus allowing the treeline to migrate northward, and into mountain ranges, to approximately its modern limit. In contrast, this pattern is not observed in eastern Beringia

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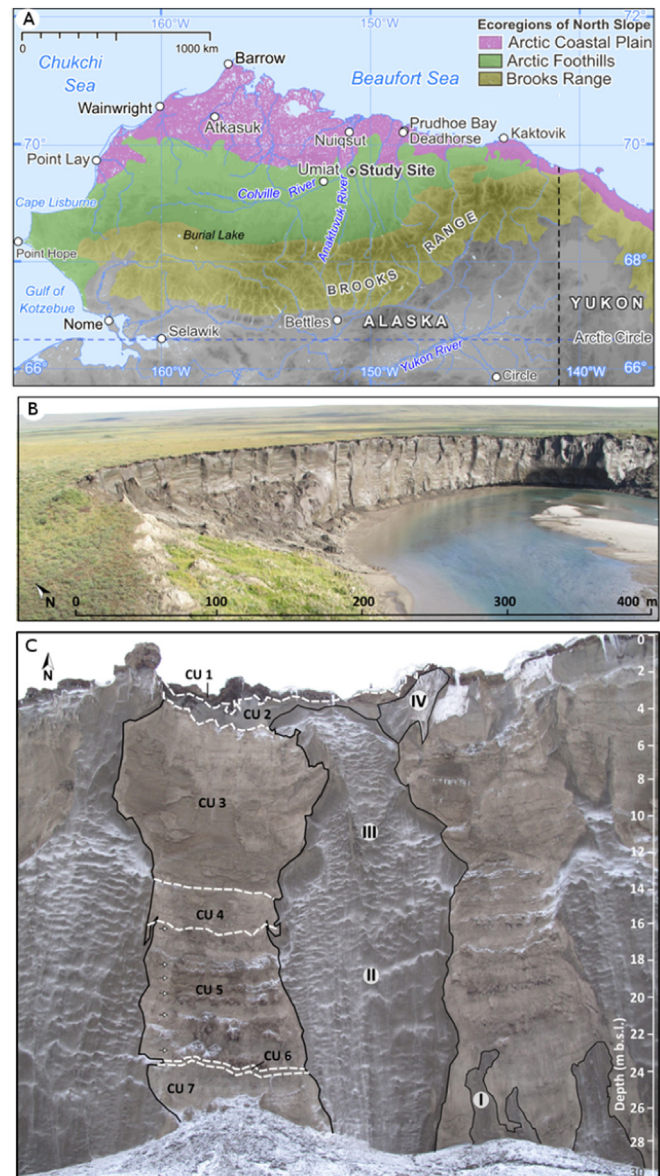
(Anderson, 1985; Anderson and Lozhkin, 2001). The transition from Middle Wisconsinan to Late Wisconsinan is imprecisely dated in eastern Beringia, but floral and faunal data indicate that conditions changed from interstadial to glacial by about 29 kyr BP (Hamilton and Fulton, 1994; Anderson and Lozhkin, 2001). July air temperatures decrease during the LGM seems to have been more substantial in eastern Beringia compared to western Beringia, as fossil beetle evidence from Siberia indicates that regional LGM summer temperatures have been close to modern levels (Alfimov and Berman, 2001).

Recently, an increasing number of multiproxy studies of yedoma sediments have contributed to a better understanding of the paleoecology of Beringia (Wetterich et al., 2008; Kienast et al., 2011; Wetterich et al., 2014). Yedoma is defined as ice- and organic rich syngenetic permafrost that is composed of an active silt deposition from a regional source and accumulation of ground ice under cold and arid conditions during the Wisconsinan (Kanevskiy et al., 2011; Strauss et al., 2012a; Schirmermeister et al., 2013; Murton et al., 2015; Strauss et al., in press). Yedoma contains well-preserved biotic remains and ground-ice features that can be used to reconstruct ecological changes and climate dynamics during the glacial, interglacial and interstadial episodes of northern high latitudes (e.g. Kaplina, 1981, 1986; Guthrie, 1990; Muhs and Bettis, 2003; Shapiro and Cooper, 2003; Sher et al., 2005; Wetterich et al., 2008; Froese et al., 2009; Strauss et al., in press). Most yedoma-based reconstructions have been undertaken at sites located in northeastern Siberia (e.g. Schirmermeister et al., 2002, 2013; Andreev et al., 2002, 2004; Murton et al., 2015), Interior Alaska (e.g. Matthews, 1974; Hamilton et al., 1988; Begét, 1990; McDowell and Edwards, 2001; Schirmermeister et al., 2016), and in the Klondike region in Yukon (e.g. Sanborn et al., 2006; Froese et al., 2009).

To date, few paleoecological studies have been conducted north of the Brooks Range in Alaska and the timing of climate changes and the ecological response of the former tundra-steppe biome to the climate fluctuations during Middle to Late Wisconsinan are poorly known. In this study, we present paleoenvironmental records from the Itkillik yedoma exposure located in northern Alaska. The primary goal is to define environmental responses related to climate changes from the Middle Wisconsinan Interstadial to the Holocene. Our objectives are to (1) reconstruct the regional vegetation using pollen analysis; (2) infer mean July air temperature ( $^{\circ}\text{C}$ ), seasonal air temperature contrast ( $T^{\circ}\text{C July} - T^{\circ}\text{C January}$ ), and July sunshine values (expressed in % of hours of bright sunshine between sunrise and sunset) using a pollen-based modern analog technique (MAT); and (3) establish links between the climate information derived from sediment and ground-ice records (geochemistry, isotopes, volumetric ice content) and the pollen-based reconstructed climate.

## 2. Study area

The Itkillik yedoma exposure ( $69^{\circ}34'0.72''\text{N}$ ,  $150^{\circ}52'7.73''\text{W}$ ) is located in the continuous permafrost zone, at the boundary between the Arctic Coastal Plain and the Arctic Foothills in northern Alaska (Fig. 1a). According to the Anaktuvuk River meteorological station (altitude: 81 m;  $69^{\circ}27'\text{N}$ ,  $151^{\circ}10'\text{W}$ ), the climate is characterized by severe winters with January air temperatures of  $-29.1^{\circ}\text{C}$  and short summers with July temperatures of  $13.0^{\circ}\text{C}$ . The mean annual air temperature is  $-10.2^{\circ}\text{C}$  and the annual total precipitation is 122 mm, with 82 to 107 mm (67 to 88%) falling as snow (Kane et al., 2014). The ground surface is slightly undulating and underlain by a poorly drained, silty mineral sediment with a shallow organic surface layer (Vioreck et al., 1992). Cyperaceae species such as *Eriophorum vaginatum* and *Carex bigelowii* are the primary tussock formers, while shrubs are sparse (Boggs et al., 2012). Previous studies in this area have been undertaken to assess the extent of loess-like sediments in the region (Carter, 1988), permafrost characteristics (Carter, 1988; Kanevskiy et al., 2011; Strauss et al., 2012a; Murton et al., 2015), and the rates and patterns of riverbank erosion (Kanevskiy et al., 2016).



**Fig. 1.** Study site. A) Location map of the study site at the Arctic Coastal plain and Arctic Foothills border (USGS, 2001); B) Photograph of the exposed bluff (photograph by Mikhail Kanevskiy, July 2007); C) Photograph of the vertical cross section studied with cryostratigraphic units (CU1–7) and stages of aggradation of ice-wedges (I–IV) Kanevskiy et al. (2011), and buried peat layers identified with arrows (photograph by Jens Strauss, May 2012).

The studied sequence is a vertical bluff exposed since the mid-1990s by thermo-erosional processes caused by the meandering of the Itkillik River (Kanevskiy et al., 2016). The surface elevation ( $\sim 85$  m above sea level) ranges from 30 to 35 m above the local river level and the total length of the exposed face exceeds 680 m (Fig. 1b). The cliff face reveals undisturbed sections of late Pleistocene syngenetic permafrost that formed during the Middle and Late Wisconsinan. Based on data from field work conducted in 2007, seven cryostratigraphic units (CU1–CU7 in Fig. 1c, Supplementary Table 1) are present within the exposed yedoma face (Kanevskiy et al., 2011). Three stages of ice-wedge aggradation are recognized at the site (Strauss et al., 2012b). The silty sediments contain sparse organic material including rootlets and wood and six peat layers were visually identified between 26 and 16 m below surface level (b.s.l.), based on their colour, continuity and the presence of visible organic remnants (Strauss et al., 2012b) (Fig. 1c). The cliff is currently eroding at a rate of  $\sim 20$  m year $^{-1}$  (Kanevskiy et al., 2016).

### 3. Methods

#### 3.1. Sampling

A sediment column in the central part of the exposure was sampled by rappelling from the top of the cliff in May 2012. An overhang between 20.5 and 9.7 m b.s.l. was impossible to reach. Samples with the same stratigraphic position were extracted from adjacent sediment columns to fill the gap. Surfaces were first cleaned with a hammer, a pic or an axe, and then a hand-held core-drill with a 4.7 cm diameter hole saw was used to extract 45 sediment samples and 48 samples from an adjacent ice wedge. The drilling head was cleaned between each sample extraction. Samples were cleaned of the drilling mud and kept frozen during fieldwork and transport to laboratory. Subsamples from the same depths were used for biogeochemistry, ice content analysis, geochronology, and pollen extraction (Strauss et al., 2012b). The biogeochemistry data and 10 radiocarbon dates from Kanevskiy et al. (2011) obtained from boreholes drilled in 2007 were also used to increase the sampling resolution. Although the locations of the 2012 sampling sites were located approximately 100 m away from the 2007 sites due to riverbank erosion, no differences in cryostratigraphy were detected between the sampling locations of these two field campaigns. The general location of the peat layers was similar between the two field campaigns, as well as the organic C patterns, confirming that the general stratigraphy was similar between the two sampling locations.

#### 3.2. Biogeochemistry

To quantify and determine the nature of carbon in the sediments, total carbon (TC) and total organic carbon (TOC) were measured. 45 intact samples of ~50 cm<sup>3</sup> (including coarse organic material present in the sediment matrix) were freeze-dried (sublimator 3–4–5, Zirbus Technology) and homogenized by grinding (Pulverisette 5 planetary mill, Fritsch) before being measured by a carbon analyzer (Vario EL III, Elementar). Two 5 mg sub-samples per sample were taken and encapsulated in two zinc capsules. To ensure complete oxidation of the sample during measurements ~10 mg of tungsten-(VI)-oxide were added. Background signals were detected by measuring a blank capsule at the beginning and calibration standards after each 20 samples. The quantified measurement accuracy is <0.1 wt%. Total organic carbon content was measured with a TOC analyzer (Vario Max C, Elementar). Four calibration standards were measured after each 14 samples. The measurement accuracy is <0.1 wt%.

As isotopic values can be used as a paleoclimate proxy,  $\delta^{18}\text{O}$  from 45 samples obtained from the ice wedge (Fig. 1c; central ice wedge) were analyzed based on mass spectrometric detection of isotopes (Hoefs, 2008). For this purpose a Finnigan MAT Delta-S mass spectrometer was combined with two equilibration units (MS Analysetechnik, Berlin). Before measuring a sequence of water samples, the laboratory reference standard NGT1 (snow from the North Greenland Traverse) was used as reference material. Additionally, for quality control, standards of different isotopic composition were employed for each equilibration unit. Standard deviations were generally better than 0.1‰. The detected values of  $\delta^{18}\text{O}$  were further recalculated to the Vienna Standard Mean Ocean Water (VSMOW) standard with the software ISODAT®.

#### 3.3. Ice content

The characteristics of the ice are indicative of the climatic conditions prevailing during the aggradation. To calculate the volumetric ice content of sediments, we weighed 48 frozen samples using a balance (0.1 g accuracy, Kern FCB 8 K0.1) and determined their volume by water displacement. To verify these volumetric measurements, we calculated the volume of the dry sample and the volume of the wet sample by assuming an ice density of 0.9167 g cm<sup>-3</sup> and a solids density of 2.65 g cm<sup>-3</sup> (adapted from Strauss et al., 2012a).

#### 3.4. Age-depth model

The plant material or bulk peat from 20 samples was isolated for radiocarbon dating (Table 1). Whenever available, we chose plant aerial macro remains over bulk peat or unidentifiable remains to avoid introducing variability in the dating (e.g. Nelson et al., 1988). When this was not possible, unidentified plant remains likely composed of rootlets and aerial components were dated. Radiocarbon dates were calibrated using IntCal13 for terrestrial northern hemisphere material (Reimer et al., 2013). Six infinite ages were excluded from the age-depth model produced by the R Bacon package (Blaauw and Christen, 2011). The Bacon model divides a core into many thin vertical sections using a Bayesian approach and then, through millions of Markov Chain Monte Carlo (MCMC) iterations, an age-depth model is built based on the accumulation rate for each of these sections (Blaauw and Christen, 2011). Ages are reported as the mean of each iteration for a 1 cm increment with the minimum and maximum of  $1\sigma$ . Period chronology is presented as rounded cal. kyr BP, omitting the uncertainty estimate for clarity.

#### 3.5. Pollen

Pollen extraction using heavy-liquid separation was performed on 19 samples obtained from the main sampling profile and 7 samples from nearby sub-profiles. The treatment involved acid digestion of carbonate with hydrochloric acid (HCl), organic matter deflocculating with potassium hydroxide (KOH), sieving (15 and 250  $\mu\text{m}$ ), heavy liquid density separation using sodium polytungstate (SPT) and acetolysis. *Eucalyptus globulus* exotic markers suspensions (Greer laboratory) were added in order to quantify pollen concentrations. Pollen samples were mounted in glycerin and analyzed under light microscope (Leitz, Wetzlar, D-69301) at 400 $\times$  magnification. When concentrations were over 1500 grains cc<sup>-1</sup> at least 300 terrestrial pollen grains were counted, otherwise 150 pollen grains were counted. Pollen diagram was graphed using the C2 software (Juggins, 2007). All identified pollen grains of trees, shrubs and herbs were used to calculate the pollen sum. Determination of pollen units was done using a temporally constrained incremental sum of squares cluster analysis (CONISS; Grimm, 1987). For CONISS, the relative frequencies (in percent) of the pollen taxa were square-root transformed. Other palynomorphs such as algae and spores (undifferentiated) were not included in the calculation of pollen concentrations or for the CONISS dendrogram.

#### 3.6. Climate reconstruction

Climate was reconstructed using the modern analogue technique (MAT, e.g. Guiot, 1990) and the North American and Greenland Modern Pollen Database (Whitmore et al., 2005). The database contains 57 taxa of tree, shrubs and grass. Each of the identified taxa at the Itkillik site was present in the database.

Modern pollen samples from the western boreal subarctic and arctic biomes north of 50°N and west of 100°W were considered as potential modern analogues candidates ( $n = 981$  modern pollen assemblages). The five closest analogs were used to compute the reconstructed climate variables. The accuracy of reconstruction can be assessed from the slope and the coefficient of multiple determination ( $R^2$ ), as well as from the standard deviation of the difference between inferred and observed values that corresponds to the root mean square error (RMSE) (e.g. Sawada et al., 2004). July air temperatures were reconstructed ( $R^2 = 0.84$ ; RSME = 1.4 °C) to allow a comparison with available beetle-based regional paleo-climate reconstructions (Elias, 2001; Kurek et al., 2009). Seasonal air temperature contrast ( $R^2 = 0.82$ ; RSME = 3.1 °C) and July sunshine values ( $R^2 = 0.85$ ; RSME = 3.0%) were selected for the reconstruction as they also have good  $R^2$  and RSME values (See Supplementary Fig. 1). Current climate estimates were obtained by calculating the average values from the nearby modern analogs and the

**Table 1**  
Radiocarbon ( $^{14}\text{C}$ ) age of the Itkillik Yedoma with depth. When the macrofossil remains were identified, the dated plant part is mentioned. Otherwise, it is noted as "Plant macro remains".

Depth (m)	Age $^{14}\text{C}$ yr BP	Age cal kyr BP	Laboratory number	Material	Source
0.7	5320 ± 35	6100 ± 107	OS-66781	in situ peat	Kanevskiy et al., 2011
1.3	8610 ± 35	9594 ± 68	OS-66782	in situ peat	Kanevskiy et al., 2011
2.3	14,370 ± 100	17,510 ± 312	Poz-51,789	Plant macro remains	This study
3.0	14,300 ± 50	17,408 ± 196	OS-66783	Twigs	Kanevskiy et al., 2011
4.0	16,550 ± 75	19,949 ± 245	OS-66784	Twigs	Kanevskiy et al., 2011
5.4	19,950 ± 180	23,999 ± 429	Poz-51,791	Plant macro remains	This study
9.3	24,290 ± 190	28,312 ± 408	Poz-51,795	Plant macro remains	This study
11.0	29,300 ± 200	33,466 ± 437	OS-66785	Twigs	Kanevskiy et al., 2011
13.7	25,450 ± 280	29,642 ± 744	Poz-51,797	Plant macro remains	This study
15.0	26,300 ± 130	30,609 ± 309	OS-66786	Twigs	Kanevskiy et al., 2011
16.0	23,900 ± 110	27,966 ± 262	OS-66780	Twigs	Kanevskiy et al., 2011
19.5	28,330 ± 390	32,357 ± 965	Poz-51,796	Plant macro remains	This study
20.6	43,500 ± 1700	47,403 ± 2506	Poz-51,788	in situ peat	This study
23.0	41,700 ± 460	45,137 ± 777	OS-66787	Fine-grained organic matter	Kanevskiy et al., 2011
23.8	>45,000	Infinite	Poz-51,793	Plant macro remains	This study
25.9	>45,000	Infinite	Poz-51,794	Plant macro remains	This study
27.5	>46,000	Infinite	Poz-51,792	Plant macro remains	This study
28.9	>48,000	Infinite	Poz-51,798	Plant macro remains	This study
30.9	>48,000	Infinite	OS-66777	in situ peat	Kanevskiy et al., 2011
31.8	>47,500	Infinite	OS-66778	Fine-grained organic matter	Kanevskiy et al., 2011

Anaktuvuk River station (July air temperature: 9 °C; Seasonal air temperature contrast: 35.5 °C; July sunshine: 39.6%).

Similarity between fossil and modern pollen assemblages was assessed using the squared chord distance (SCD) (Overpeck et al., 1985). Values of SCD vary between zero and two, with higher values indicating greater dissimilarity. The adopted threshold for the analog is 0.21, which is in agreement with suggested values for MAT-based reconstructions from fossil pollen assemblages in the Arctic biome in Fréchet et al. (2008). SCD threshold values over 0.20 for the Arctic biome reduces the probability of generating false-negative errors, such as the omission of contrasting pollen assemblages originating from sites experiencing similar climate and therefore representing viable analogues (Fréchet et al., 2008). Climate reconstructions were performed with the Bioindic package (Guiot and Gally, 2014) built on the R platform. More details on the MAT approach used here are given in Fréchet et al. (2008).

## 4. Results

### 4.1. Sediment and ground-ice records

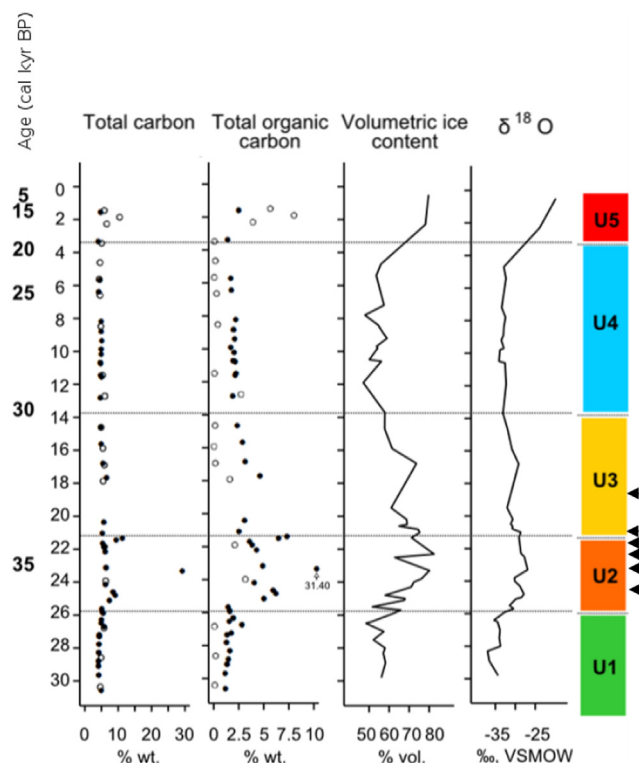
Sedimentary and biogeochemical characteristics of the sediments were grouped into five units based on pollen and climate reconstructions for consistency (U1-5) (Fig. 2, see Section 4.3 and 4.4). Overall, TC shows an average of 4.9 wt% and is mostly inorganic (generally over 75% of TC), while the volumetric ice content including segregated and pore ice averages 62%.

Sediments in the first unit (U1: 30.0–25.3 m b.s.l.) are composed of yellowish-grey silt. The U1 TC values are as low as 2.7–4.7 wt% with an average of 0.4 wt% as organic compounds. The volumetric ice content of sediments between ice-wedges varies from 48.5 to 65.9 vol%. Stable isotope analysis shows an average  $\delta^{18}\text{O}$  value of  $-33.6\text{‰}$ . Isotopic anomaly decreases slightly towards the top of the unit.

Sediments in the second unit (U2: 25.3–20.9-m) are composed of organic-rich silt with medium grained sand that grades upwards into silty peat with graminoid remains. TC values increase in U2, varying between 4.0 and 29.4 wt% with a TOC average of 4.5 wt%. Systematic differences in TOC values between the two sampling campaigns may be due to the differences in TOC measurement methods. However, in both cases, TOC enrichment matches with the position of peat layers located in this unit and U3 (23.4–24.2, 21.5–22.7, 19–19.8, 17–17.6, ~16.0 m b.s.l.). At 22.5 m b.s.l., TC reaches 29.4 wt%, while the TOC reaches 31.4 wt% making U2 the most organic rich unit. Volumetric ice content of segregated ice increases upwards and reaches values as high as 82.5 vol%. The peak

in ice content at 23.3 m b.s.l. corresponds to an ancient position of the ice-rich upper horizon of permafrost, known as the intermediate layer (Shur, 1988; Shur et al., 2011).  $\delta^{18}\text{O}$  decreased to average values of  $-28.7$  with lower values (until U5) concordant to 23.3 m b.s.l.

Sediments in the third unit (U3: 20.9–13.3 m b.s.l.) are composed of yellow-grey silt with inclusions of organic material, rootlets and twigs. The TC value of U3 varies from 10.4 to 3.6 wt% (average TOC of 1.9 wt%) and tends to slightly decrease upwards. Highest TOC values



**Fig. 2.** Biogeochemical, ice content, and  $\delta^{18}\text{O}$  data from the Itkillik Yedoma exposure over ~39 kyr. Total carbon and total organic carbon are represented in weight percent (wt%); volumetric ice content values are represented in volume percent (vol%);  $\delta^{18}\text{O}$  anomaly values are represented in parts per thousand compared to the Vienna Standard Mean Ocean Water (‰, VSMOW). Note that data shown as empty circles are from the 2007 sampling campaign (Kanevskiy et al., 2011). Data from 2012 (this study) are illustrated as black filled circles. Units (U1 to U5) are given on the right-hand side. Peat layers are identified with the black arrows.

correspond to the location of the uppermost peat layer (~16 m b.s.l.). As volumetric ice content decreases by about 17.4 vol% (from 73.8 to 56.5 vol%) over the unit, stable isotope values for wedge ice become more stable with an average value of  $-30.9\%$  for  $\delta^{18}\text{O}$ .

Sediments in the fourth unit (U4: 13.3–3.3 m b.s.l.) are yellow-grey silts with small in situ inclusions of poorly decomposed rootlets and twigs. The U4 TC values vary between 3.0 and 4.9 wt%, while TOC is low (average of 0.6 wt%). Volumetric ice contents are lower, with values varying between 47.1 and 59.0 vol%. The  $\delta^{18}\text{O}$  values are very different from modern values, with an average of  $-33.0\%$ .

In the last unit (U5: 3.3–0 m), organic-rich silt with some clay and very fine sand with peat and moss become dominant at the surface. This unit shows a net increase in TOC reaching 8.18 wt% (average of 2.45 wt%). TC remains elevated at the bottom of this unit and reaches about 17.2 wt% near the surface. Volumetric ice content exceeds 75 vol% (78.9% average) indicating the position of the modern intermediate layer (Shur, 1988; Shur et al., 2011; French and Shur, 2010). Stable isotope analysis shows rapidly decreasing values going upward the unit with  $\delta^{18}\text{O}$  that reaches nearly modern values.

#### 4.2. Geochronology

Except for two overlapping radiocarbon dates (11 and 16 m b.s.l.), ages generally increase with depth (Fig. 3). Sediment accumulation varies between 0.01 and 0.33  $\text{cm yr}^{-1}$  with an average rate of 0.16  $\text{cm yr}^{-1}$ . The fastest vertical sediment accumulation occurred prior to 28.1 cal kyr BP with average rates over 0.17  $\text{cm yr}^{-1}$ . From 31.7 to 30.7 cal kyr BP and from 29.9 to 28.4 cal kyr BP average accumulation rates reached about 0.25  $\text{cm yr}^{-1}$ . Slower accumulation rates were recorded from 28.4 to 17.8 cal kyr BP (average rates are below 0.10  $\text{cm yr}^{-1}$ ) and especially after 17.1 cal kyr BP (average rates of 0.01  $\text{cm yr}^{-1}$ ). It should be noted that the landscape-scale sedimentation rates were much slower due to the presence of ice wedges. Landscape-scale rates of accumulation cannot be calculated based on our data.

#### 4.3. Pollen

Based on the CONISS cluster analysis, the pollen diagram was subdivided into five units (U1–5) corresponding to changes in taxa

composition and abundance (Fig. 4). The profile is dominated by herbaceous taxa, notably Poaceae and Cyperaceae. Pollen grains of boreal trees such as *Picea* and *Pinus* are from distal sources as they are not currently present at the study site, and were very likely not present during the Middle and Late Wisconsinan period.

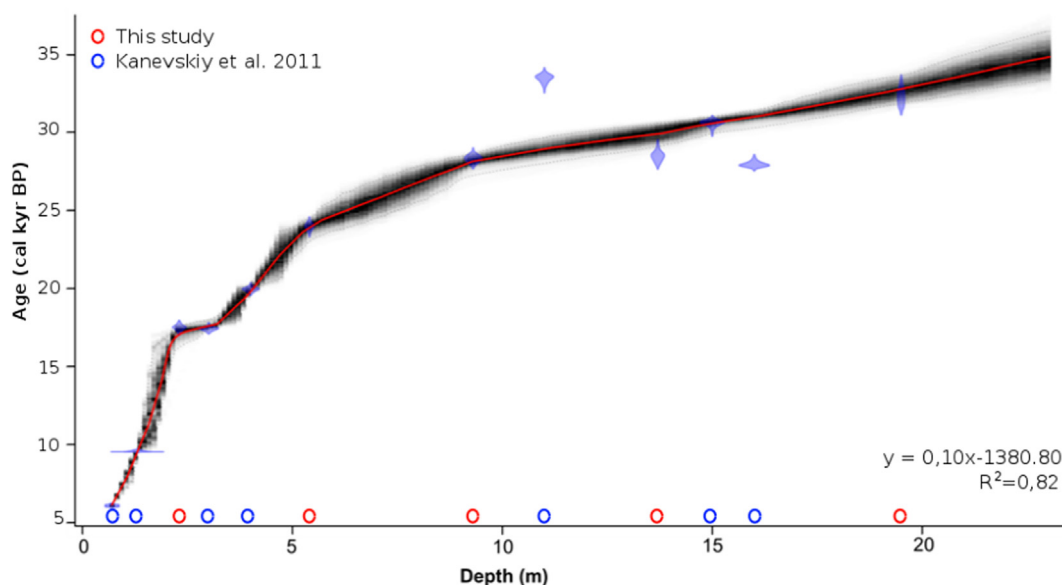
U1 (30.0–25.3 m b.s.l., >34.8 cal kyr BP) is characterized by a low pollen concentration (<2500 pollen grains  $\text{cc}^{-1}$ ). We observed pollen grains of Poaceae (>50%), Cyperaceae (~9–23%), and to a lesser extent, of *Artemisia*, Tubuliflorae (sub-family of Asteraceae), and Brassicaceae. *Picea* and *Pinus* reach their highest values of the entire sequence (~10%).

U2 (25.3–20.9-m b.s.l., >34.8–33.6 cal kyr BP) shows a 44% increase in Cyperaceae pollen grains concentration (from 28 to 72%), while Poaceae decreases by 47% (from 65 to 18%). Other minor herbs and forbs are less represented than in the first unit. *Zygnema* algae occurrence is high and reaches an abundance peak of 20% at ~35 cal kyr BP. The percentage of boreal tree taxa (*Picea*, *Pinus*) substantially decreases to <2% from the previous unit. An upward increase of total pollen concentration was observed over this unit, with a maximum value reaching ~5000 grains  $\text{cc}^{-1}$ .

U3 (20.9–13.3 m b.s.l., 33.6–29.8 cal kyr BP) is notable for having high Cyperaceae abundance at the bottom of the unit. It decreases upward by 42% (from 94 to 52%). Inversely, Poaceae increased by 10% (from 2 to 12%) concurrently with arboreal pollen (from 1 to 9%) over this unit. Minor herbs and forbs abundance and diversity are lower than before, but tend to increase at the end of the unit. Total pollen concentration slightly decreases compared to U2 to about 3400 grains  $\text{cc}^{-1}$ .

U4 (13.3–3.3 m b.s.l., 29.8–17.9 cal kyr BP) shows a clear dominance of Poaceae over Cyperaceae. This last taxon tends to increase upwards (31% on average, going from 64 to 33% for Poaceae and 35%, going from 4 to 39% for Cyperaceae). Minor herbs and forbs abundance and diversity reach their highest values, with peaks of *Artemisia*, Asteraceae (sub-family of Tubuliflorae) and Brassicaceae. Boreal tree taxa stay abundant (~10%). The lowest pollen concentrations of about 2000 grains  $\text{cc}^{-1}$  were measured in this unit.

The pollen assemblage of U5 (3.3–0 m b.s.l., 17.9 cal kyr BP–modern) is dominated by Cyperaceae (67–96%), in agreement with current surface vegetation. Forbs are poorly represented and Poaceae have decreased compared to the previous unit (<10%). Pollen concentration exceeds 2500 grains  $\text{cc}^{-1}$ .



**Fig. 3.** Age-depth model showing calibrated  $^{14}\text{C}$  dates (blue). The grey stippled lines show 95% confidence intervals and the red curve shows the best model based on the weighted mean age for each depth. The blue envelopes represent the probability distribution of the calibrated ages as violin plots. The circles at the bottom of the plot indicate the sampling campaign from which the radiocarbon dates were obtained. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.) (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)



anomaly tends to decrease with time from  $-2$  to  $-1$  °C. Average seasonal air temperature contrast anomaly is  $-1$  °C. July sunshine anomaly decreases to nearly modern value (from 3 to  $<1\%$ ).

For U3, the MAT found five valid analogs with an average SCD for the closest analogues of 0.13. July air temperature is close to modern values between 33.6 and 31.4 cal kyr BP and decreases drastically to reach an anomaly of  $-2.5$  °C by 30.4 cal kyr BP. Seasonal air temperature contrast is similar to current values. July sunshine is approximately equivalent to the modern value for most of the unit. Climate reconstructions for U3 should be interpreted cautiously as they are based on only three pollen assemblages.

Six out of seven pollen assemblages for U4 have good analogs with an average SCD value of 0.12 for the closest (first) analog. However, the bottom sample of the unit (11.9 m b.s.l.; 29.3 cal kyr BP) has only one analog, indicating that its pollen assemblage is very different from modern ones. July air temperature increases slightly from 3 to 2 °C (lower value) through the unit. Seasonal air temperature contrast varies between 1 and 2 °C lower than current. July sunshine varies from 1 to 3% higher than modern values.

Paleoclimate reconstructions for U5 are based on only two pollen assemblages with a high SCD average value of 0.17 for the closest analog. MAT found five valid analogues for the two surface samples. July air temperature is close to the modern value. Seasonal air temperature contrast and July sunshine anomalies vary respectively by about  $-2$  °C and  $-2\%$  compared to current value over the unit.

## 5. Discussion

### 5.1. Chronology of paleoenvironmental records

Lack of chronological control in U1 and early U2 prevents us from assessing a precise age for the sediments deeper than 23.3 m b.s.l. (prior to  $\sim 35$  cal kyr BP). Extrapolation from average accumulation rates indicates that if there was no interruption in sediment deposition, U1 is probably no older than 39 cal kyr BP and is likely encompassed into the Middle Wisconsinan interstadial. Thus, we interpret proxy data changes over U1–3 ( $\sim 39$  to 29.8 cal kyr BP) as climatic fluctuations occurring during the interstadial. However, as the assignment of U1–3 is

based on an extrapolation of finite  $^{14}\text{C}$  ages and there is no evidence that sedimentation is continuous between U1 and U2, we remain very cautious about the interpretation of U1 and early U2. Radiocarbon dating on Middle Wisconsinan permafrost samples is complicated due to detection limits and cryoturbation that can relocate material (e.g. Wetterich et al., 2014). For instance, the reversed ages at 11 and 16 m b.s.l. (Table 1) could be explained by erosion and redistribution of younger sediments or cryogenic relocation processes. U4 age is related to the following Late Wisconsinan glacial advance in the surrounding areas (29.8 to 17.9 cal kyr BP), while U5 corresponds to the Late Glacial transition into the Holocene ( $<17.9$  cal kyr BP) (Fig. 6). Considering the limited age control, we also remain cautious about the chronology of U5.

The chronology of sediment accumulation rate fluctuations coincides roughly with sediment unit boundaries. Therefore, we interpret rapid accumulation rates as active Pleistocene sedimentation processes and the subsequent slow-down observed about 17.9 cal kyr BP as the transition to Holocene, characterized by more active organic accumulation and less active deposition of loess.

### 5.2. Beringian palaeoenvironments

As yedoma form by long distance deposition of eolian sediments, the pollen found at the Itkillik site is interpreted as being representative of regional vegetation and climate. Compared to many Beringian diagrams, pollen assemblages from the Itkillik yedoma are poorly diversified with only 14 minor taxa (Fig. 4). The most important patterns observed across the units are the shifts between Cyperaceae and Poaceae dominance in pollen assemblages. Thus, it is possible that shifts in grass and sedges are almost exclusively driving the reconstructed changes in regional climate. When Cyperaceae dominated the pollen assemblages (U2, U3 and U5), ice and organic carbon contents are higher locally at the site. Isotopic values were also more similar to modern values of snow, implying warmer and moister conditions that were favorable to peat aggradation. Inversely, when Poaceae and other minor taxa percentage are higher, other proxies indicate drier, windier, and colder conditions that are promoting silt accumulation.

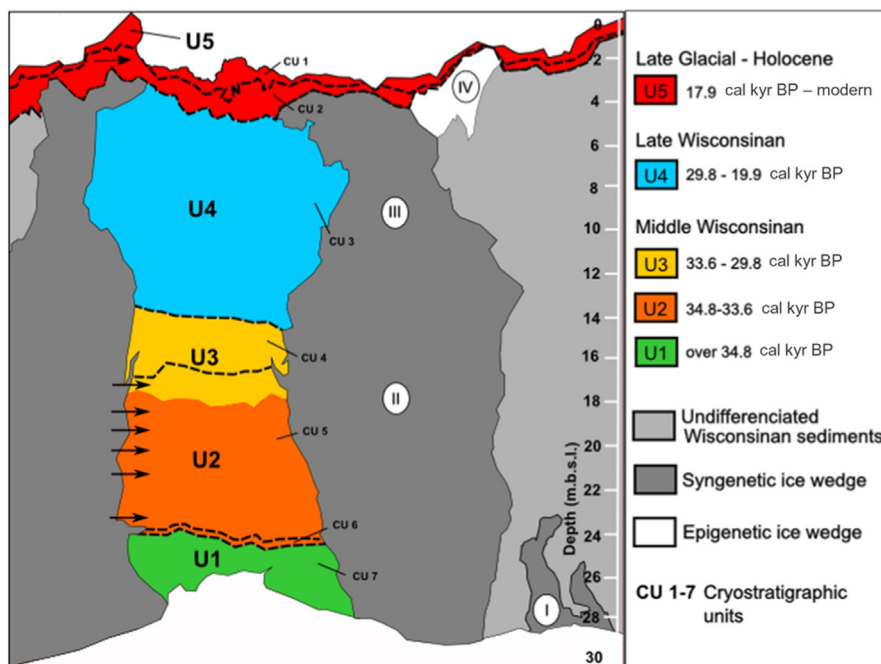


Fig. 6. Sketch of sediment units exposed with cryostratigraphic units (CU 1–7) and stages of aggradations of ice-wedges (I–IV) (based on Kanevskiy et al., 2011). Note that the dashed lines represent the boundary between CU and that the black arrows indicate the locations of the peat layers.

### 5.3. Unit 1

Presence of *Poaceae* and *Artemisia*, and a high diversity of forbs taxa in U1, which we interpret as approximately covering the ~39–36 cal kyr BP interval, suggest that dry conditions prevailed in a grass environment (Zazula, 2002). Reconstructed July air temperatures indicate colder than modern climate conditions during that time interval. The poor analog quality found for U1 is interpreted as conflicting pollen rain inputs from contrasted vegetation across eastern Beringia. Considering the chronological uncertainties, we must remain cautious about the interpretation of reconstructed climate for U1. The low organic carbon content is indicative of cold conditions that do not favor carbon accumulation. The lower  $\delta^{18}\text{O}$  values reported in U1 compared to modern snow values of  $-22.3\%$  reported for northern Alaska (Sloat, 2013) and to modern ice-wedges ( $-25.0\%$ ,  $n = 3$ ) sampled at the Itkillik River site is also indicative of colder than modern conditions. The lower ice content and micro-lenticular cryostratigraphic units (Kanevskiy et al., 2011) suggest syngenetic permafrost aggradation under severe periglacial conditions.

In eastern Beringia, a cold phase approaching glacial conditions was reported in interior Alaska from 39 to 34–33 cal kyr BP, and in northwestern Alaska, from 40 to 35 cal kyr BP (Anderson and Lozhkin, 2001) (Fig. 7). July air temperatures ranging around modern values are reported in western Beringia during most part of the Karginsky interval (the western Beringian equivalent of the Middle Wisconsinan, see Supplementary Table 2) (Anderson and Lozhkin, 2001; Elias and Brigham-Grette, 2006). In the central and western areas of western Beringia, *Larix* forests were relatively common through much of the Middle Wisconsinan (Anderson and Lozhkin, 2001). As for Western Beringia, warmer and moister climatic conditions, and favorable edaphic conditions have led to the expansion of *Picea* populations in the interior of Yukon (Anderson and Lozhkin, 2001) and could explain the high pollen representation of this taxon prior to ~36 kyr BP at the Itkillik site. Under a strong low-pressure atmospheric system, Pinaceae pollen can be transported over thousands of km and create noise in pollen records (Campbell et al., 1999). However, as chronological control is lacking and as proposed for the lower unit of Ahaliork Lake located near the study site (Eisner and Colinvaux, 1990), high *Picea* presence could be related to the Sangamonian interglacial (130–71 kyr BP).

### 5.4. Unit 2

Pollen assemblages in U2 are dominated by Cyperaceae and a higher representation of *Sphagnum* spores, implying more humid conditions (Ager, 2003). July air temperatures and seasonal air temperature contrast increased, as July sunshine declined. Concomitant decreases in  $\delta^{18}\text{O}$  values also support the assertion of a warmer and moister climate during this period. The vegetation is interpreted as a moist Cyperaceae

meadow in a low centered ice-wedge polygons landscape. Organic carbon and ice content enrichment across U2 indicate, respectively, environmental conditions favorable for organic matter accumulation and greater moisture availability around 35 kyr BP.

Despite the chronological uncertainties in unit U2, it seems chronologically related to the Middle Wisconsinan interstadial optimum reported across eastern Beringia (Anderson and Lozhkin, 2001). The interstadial was characterized by higher summer insolation than the preceding and following stages. The extent of alpine glaciation was reduced. With higher sea level (60–80 m lower than present), there was more efficient transport of moist air masses from the North Pacific Ocean towards Alaska (Bender, 2013). The significant increase in volumetric ice content (reaching over 75 to 80% in U2) at the Itkillik site is related to the formation of a buried intermediate layer (~24 m b.s.l.), suggesting surface stabilization, vegetation accumulation, and increased ground-ice accumulation (cryostratigraphic unit 6 from Kanevskiy et al., 2011). The climatic reconstruction suggests Itkillik site experienced a warmer interval from 36 to 33.6 cal kyr BP. In agreement with our observations, higher lake levels are reported across interior and northern Alaska between 37.2 and 33 cal kyr BP (Finkenbinder et al., 2015). This is occurring 2–3 kyr prior to the warm interval reconstructed from the Fox permafrost tunnel record in interior Alaska (from 33 to 34 to 30 kyr BP) (Hamilton et al., 1988).

Formation of U2 was likely triggered by peat accumulation which occurred from 34.8 to 33.6 cal kyr BP (cryostratigraphic unit 5 in Kanevskiy et al., 2011). The peat layers resulted in the significant increase in organic carbon content, representing 40 to 45% of total carbon. Higher OC contents relate to changes in primary productivity and plant assemblages that are characteristic of Middle Wisconsinan yedoma deposits (Schirmermeister et al., 2013). Peat layers of similar age are described in several yedoma exposures in western Beringia (Tomirdiario, 1980; Kaplina, 1981; Hopkins, 1982; Tomirdiario and Chyornen'kiy, 1987) and eastern Beringia (Anderson and Lozhkin, 2001). Peat layers in yedoma sequences are related to buried paleosols indicating interstadial accumulation (Ruhe, 1983) and often can be observed above buried intermediate layers (Gubin, 2002; Kanevskiy et al., 2011).

Similar to climate reconstructions developed for sites located in Interior Alaska and Yukon (Anderson and Lozhkin, 2001), the climate reconstruction of the interstadial optimum from Itkillik suggests a climate that was warmer than during glacial advance episodes but drier than during the Holocene.

### 5.5. Unit 3

The silt accumulation above the peat layer of U3 suggests a return of regional wind erosion, eolian transport and sedimentation which resulted in higher sedimentation accumulation rates (see unit 4 in Kanevskiy et al., 2011). Concurrently, the climate reconstruction and

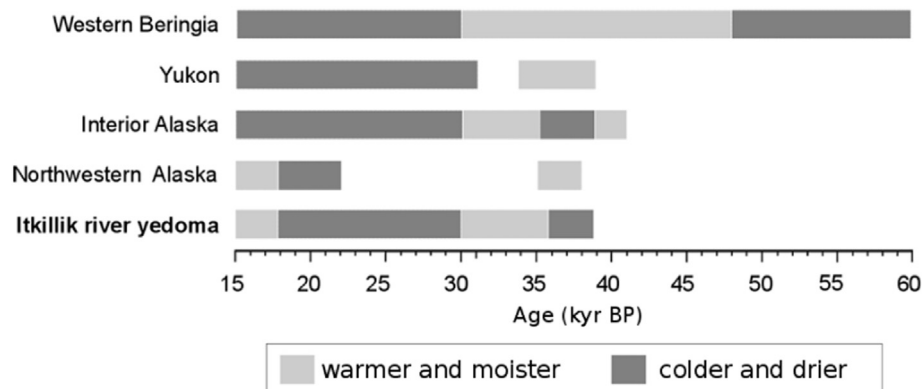


Fig. 7. Simplified synthesis of paleoclimate fluctuations of Beringia. Itkillik River yedoma (this study); Northwestern Alaska (Burial Lake: Kurek et al., 2009); Interior Alaska (Fox permafrost tunnel: Begét, 1990; Anderson and Lozhkin, 2001); Yukon (Anderson and Lozhkin, 2001); Western Beringia (Wetterich et al., 2014).



isotopic composition during the 33.6–29.8 cal kyr BP interval at the Itkillik site (U3) support the idea of a transition stage towards drier climate conditions. The climate during this period is associated with higher July sunshine value (%) and greater variability than modern (Fig. 5). Increasing sunshine reconstructed from MAT could be associated with reduced cyclone frequency and therefore strong anticyclonic atmospheric circulation pattern (Serreze et al., 2000). However, the uppermost buried peat layer at the exposure is aged about 30.4 cal kyr BP and the climate reconstruction suggests that until that time, climate conditions were not as harsh as during the LGM.

#### 5.6. Unit 4

The climate reconstructions for the 29.8–17.9 cal kyr BP interval at the Itkillik site show that July air temperatures were on average 2–3 °C colder than modern, with Poaceae and forbs dominating the vegetation, suggesting drier conditions than episodes dominated with Cyperaceae. The  $\delta^{18}\text{O}$  values reaching  $-30\%$  are coherent with the pollen-based climate reconstruction pointing that both summer and winter air temperatures were much lower. The dry tundra-dominated environment of Itkillik during the Late Wisconsinan is interpreted as a steppe. Tundra and woody communities are reported in Interior Yukon and southern Alaska over the same period (Anderson and Lozhkin, 2001). Seasonal air temperature contrast is low compared to the previous units. Ice wedges are thinner (<5 m) than in units 1 to 3, suggesting conditions favoring faster eolian sedimentation. The sediment column is characterized by high silt and low ice content (cryostatigraphic unit 3 from Kanevskiy et al., 2011). The low seasonal temperature contrast and OC content is indicative of low productivity condition and shorter growing seasons (Anderson and Lozhkin, 2001). High values of July sunshine between 29.8 and 17.9 cal kyr BP suggests increased continentality and thus drier conditions potentially due to lower sea levels during the LGM. The harshest conditions near 30 cal kyr BP are related to a rapid lowering of the sea level and glacial advance. Also, perennial sea ice cover (Bartlein et al., 1991) likely prevented the Arctic Ocean from providing moisture to the North Slope of Alaska. Today, winter air temperature and precipitation have the greatest controls on  $\delta^{18}\text{O}$  values in Alaska and may explain the large  $\delta^{18}\text{O}$  anomalies observed over that interval. However, warmer and moister conditions are observed from 22.2 cal kyr BP and could be related to ice retreat following the regional early Itkillik II glaciation advance reported in the Brooks Range (Hamilton et al., 1988).

A cooling episode following the Middle Wisconsinan optimum occurred in Interior Alaska 30–26 kyr BP (Anderson and Lozhkin, 2001). Starting around 30 kyr BP, the increasing extent and thickness of continental ice sheets during the LGM led to a eustatic drop from about 60 m to about 120 m lower than present-day sea levels (Clark et al., 2009). Several beetle assemblages from eastern Beringia indicate that LGM mean July air temperatures in Alaska were lower ( $-1$  to  $-4$  °C) than modern while western Beringia beetle assemblages suggest July air temperatures as warm as modern (Elias, 2001). A high degree of regional variability among climate reconstructions of eastern Beringia sites make the interpretation of climate difficult; not all fossil assemblages that date to the LGM interval reflect cooling compared to modern temperature averages at the study sites (Elias, 2001).

From  $\sim 25.0$  to 17.0 cal kyr B.P., mean July temperatures from Zagoskin Lake and Burial Lake average 2.3 and 3.8 °C colder than today, respectively (Kurek et al., 2009). Greater continentality across Beringia may explain the main differences between LGM and Holocene climates (Guthrie, 2001). Reconstruction from Viau et al. (2008) based on 47 fossil pollen sites from the North American Pollen Database show that July air temperature was around 4 °C lower during full glacial conditions. The coldest interval in northern Alaska is related to the early Itkillik II glaciation advance in the Brooks Range from 27 to 21 cal kyr BP (Hamilton, 2003). Burial Lake (355 km south-west of the study site) proxy data indicate xeric and windy conditions in northern Alaska

during the LGM (Abbott et al., 2010). Based on Chironomidae assemblages, Kurek et al. (2009) report a 3.5 °C July air temperature cooling during the LGM in Northern Alaska that matches our July air temperature reconstruction.

#### 5.7. Unit 5

Vegetation changed over this period from Poaceae to Cyperaceae-dominated and moister conditions comparable to modern ones were taking place. These changes are not accompanied by an increase in woody taxa as observed in other regions of eastern Beringia. Increases in relative abundances of *Betula*, *Salix*, *Populus*, and Ericaceae occurred across eastern Beringia during deglaciation and are related to higher temperatures and precipitation (Ager and Brubaker, 1985). However, presence of peat and Cyperaceae dominance in the upper part of the pollen assemblage is concordant with the modern landscape, essentially colonized by sedges and sparse shrubs.

An abrupt decrease in silt accumulation from 17.1 cal kyr BP (Fig. 2) indicates the end of the Itkillik River yedoma buildup. Similarly, for the Burial Lake site, higher lake level and termination of eolian silt input from 17 cal kyr BP were reported (Abbott et al., 2010). Mann et al. (2002) reported that peat deposition began and shrub vegetation became widespread north of the Brooks Range about 12.5 kyr BP, probably in response to the advent of a warmer and wetter climate. Today, the presence of peat in the Arctic Foothills is widespread. The  $\delta^{18}\text{O}$  anomaly values decrease to  $>-25\%$  is concordant with the modern snow ratio (22.3%; Sloat, 2013). Climate reconstruction and isotopic composition of adjacent ice wedges support the presence of a more seasonally contrasted, warmer and moister climate. Less sunshine could be related to higher cyclonic atmospheric circulation compared to the LGM because of higher sea level that promoted the presence of warmer and moister air mass from the North Pacific. However, we cannot rule out the possibility that lower sunshine may also be due to higher fog coverage. Increased organic carbon content may indicate higher primary productivity. High volumetric ice content at 0.6–1.5 m b.s.l. is related to the suspended cryostructure typical of the intermediate layer (cryostatigraphic unit 2 from Kanevskiy et al., 2011). The main difference in U2–3 and U5 climates is not the temperature but the aridity.

With only two post-LGM sediment samples, it is not possible to infer from our data climate and ecological changes related to short events such as the last glacial re-advance during the Itkillik II glaciation and the Dryas intervals. The occurrence of modern epigenetic ice wedges (generation IV, Fig. 6) indicates that these ice wedges formed after the deposition of the host sediments (French and Shur, 2010).

## 6. Conclusions

The Itkillik yedoma formed by long distance deposition of eolian sediments. It preserves an important palaeoenvironmental record for northern Alaska that reflects climate and ecological changes over 35 kyr during the Middle Wisconsinan Interstadial and the following Late Wisconsinan glacial advance. Over that period, shifts between Cyperaceae and Poaceae dominance were observed in the pollen spectra. When Cyperaceae were dominant, ice and organic carbon contents were higher locally at the site. When  $\delta^{18}\text{O}$  values were similar to modern values, implying warmer and moister conditions, peat aggradation was promoted. Inversely, when Poaceae and other minor taxa were dominant, sediment and ground-ice records indicated drier and colder climate conditions, which promoted silt-rich accumulation. The climate fluctuations timing are consistent with interior Alaska and Yukon records. Similar to interior Alaska, the study site experienced cold conditions before 35 cal kyr BP and a following warmer interval until 30 cal kyr BP. The 35 cal kyr BP Middle-Wisconsinan optimum record at Itkillik is in agreement with reports from both interior Alaska and Yukon.

Our study demonstrates the usefulness of combining the use of pollens, carbon and ground ice records yedoma to study Beringia

paleoenvironments. As for most studies covering the Middle Wisconsinan - Late Wisconsinan time period, uncertainty remains about the chronology of major vegetation and climate shifts. A better understanding of regional atmospheric circulation patterns and edaphic conditions could enhance the quality of interpretations of the ecological heterogeneity across Beringia during the climate fluctuations related to the Wisconsinan glaciation.

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