

# A novel synthesis of the N-13 labeled atmospheric trace gas peroxyntic acid

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(Received November 29, 2010; accepted in revised form January 14, 2011)

*Synthesis / N-13 / Nitrogen oxides / Snow-air interface / Atmosphere / Cryosphere / PNA / HNO<sub>4</sub> / HO<sub>2</sub>NO<sub>2</sub>*

**Summary.** Radioactively labeled trace gases have been successfully used to study heterogeneous chemistry of atmospheric relevance. Here we present a new synthesis of gas-phase peroxyntic acid labeled with <sup>13</sup>N (H<sup>13</sup>NO<sub>4</sub>) to study the interaction of HNO<sub>4</sub> with ice and snow surfaces. A yield of about 30% for HNO<sub>4</sub> was determined. The main by-products were HNO<sub>3</sub> and HNO<sub>2</sub>. Exposure of an ice packed bed flow tube to these species revealed that the interaction with the surface scale in the order HNO<sub>3</sub> > HNO<sub>4</sub> = HNO<sub>2</sub> > NO<sub>2</sub>.

## 1. Introduction

The interaction of atmospheric trace gases with the Earth's snow cover or ice surfaces is of high environmental relevance [1, 2]. The atmospheric trace gas that we focus on in this work is the nitrogen oxide peroxyntic acid (HNO<sub>4</sub>). Nitrogen oxides have been of particular interest in atmospheric science, because their gas-phase concentration directly influences the ozone levels and the oxidative capacity of the atmosphere.

For example, in the lower atmosphere, HNO<sub>3</sub> adsorbs to and is thus scavenged by ice clouds [3]. Similarly, on the ground, trace gas uptake by snow may significantly alter the concentration both in the air above snow-covers and in the snow itself [4, 5]. Nitrous acid (HNO<sub>2</sub>) partitions less to ice or snow, but its interactions with the ice are still strong enough to significantly slow down its diffusion through surface snow [6, 7]. This longer residence time in the snow pack, as compared to non-interacting species such as NO<sub>2</sub>, makes its photolytic dissociation and thus its role as source of the strong oxidant OH more probable. Recently, Slusher *et al.* concluded that the uptake of HNO<sub>4</sub> by snow is of similar magnitude than observed for HNO<sub>3</sub> [8]. The finding that deposition to snow is a major loss process for gas-phase HNO<sub>4</sub> above Antarctic surface snow was based on steady-state calculations and comparison to field measurements. Taken its significant fraction of total nitrogen oxides in cold

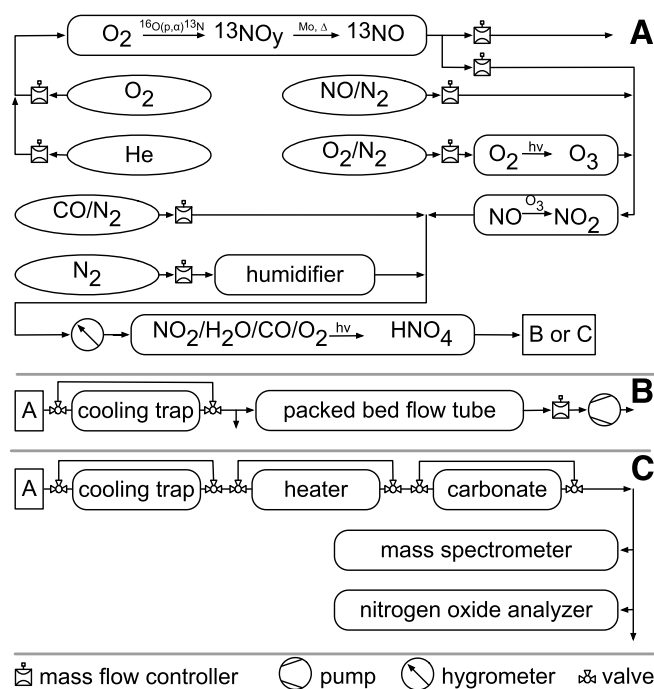
regions of the atmosphere such as in Antarctica and the upper troposphere [9, 10], its uptake would significantly impact the budget of gas-phase nitrogen oxides. Also, as HNO<sub>4</sub> has a significant photo-dissociation cross section [11], its deposition should be taken into account when discussing the vivid photochemistry of nitrogen oxides in snow-packs [4]. Sound conclusions are however hampered by missing data from well-controlled laboratory experiments on the interaction of HNO<sub>4</sub> with ice or any other environmental surface.

<sup>13</sup>N with a half-life of 10 min has been used in the past to label nitrogen oxides to be used as tracer in chemical experiments with relevance to Earth's atmosphere [12]. The main aim of this work was to develop a new synthesis route for <sup>13</sup>N labeled HNO<sub>4</sub> in the gas phase by an association reaction of <sup>13</sup>NO<sub>2</sub> with HO<sub>2</sub>. <sup>13</sup>N isotopes are obtained from the PROTRAC facility at the Paul Scherrer Institute. In the first part of the work, we describe the yield of the product (HNO<sub>4</sub>) and of other nitrogen oxide by-products by means of a chemical ionization mass spectrometer and of a chemiluminescence detector. This work was done with an excess of H<sup>14</sup>NO<sub>4</sub>. As neither method can differentiate <sup>13</sup>N from other nitrogen isotopes, the production of <sup>13</sup>N labeled nitrogen oxides was verified using a packed bed flow tube for separation and radioactive decays for detection in a second step. This method has been called thermochromatography before [13]. In this work, ice was chosen as stationary phase, because the separation of HNO<sub>3</sub>, HNO<sub>2</sub>, NO<sub>2</sub>, and of NO in an ice packed bed flow tube has been shown before [14] and additionally, first information on the partitioning of HNO<sub>4</sub> between ice and air can be gained.

## 2. Experimental

Fig. 1 shows a scheme of the experimental set-up consisting of the production of <sup>13</sup>N in a gas target and of the synthesis of HNO<sub>4</sub> (Fig. 1a). Included is also a scheme of the packed bed flow tube that was used to verify the production of H<sup>13</sup>NO<sub>4</sub> (Fig. 1b) and of the analytical set-up to characterize the HNO<sub>4</sub> synthesis (Fig. 1c). The tubing of the flow system consisted of perfluoro-alkoxy copolymer (PFA) 4 mm i.d. Gas flows were controlled by mass flow controllers (Brooks Instruments) or by mass flow regulators (Vögtlin Instruments) both of which have a 1% full-scale accuracy.

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**Fig. 1.** Scheme of the experimental set-up. (a) shows the production of  $^{13}\text{N}$  isotopes and the synthesis of  $\text{HNO}_4$  molecules. The synthesis was coupled to either a packed bed flow tube for detection of radioactively labeled nitrogen oxides (b), or to analytical devices for detection and characterization of also non-labeled molecules (c).

## 2.1 Production of $^{13}\text{N}$

The production of  $^{13}\text{N}$  isotopes *via* the reaction  $^{16}\text{O}(p, \alpha)^{13}\text{N}$  has been described in detail before [12]. In brief, a flow of 10%  $\text{O}_2$  (99.9995%, AirLiquid) in He (99.9999%, Messer) passed through a gas target at  $1 \text{ L min}^{-1}$  flow velocity and at 2 bar pressure that was continuously irradiated with an 11 MeV proton beam – provided by the accelerator facilities at Paul Scherrer Institute, Switzerland. The primary  $^{13}\text{N}$  molecules and radicals were reduced to  $\text{NO}$  over a molybdenum catalyst at 653 K, immediately after the target cell. The resulting gas was continuously transported to the laboratory through a 580 m long capillary. The radiation chemistry in the target cell also led to the production of non-labeled nitrogen oxides at around  $2 \times 10^{11} \text{ molecules cm}^{-3}$  from nitrogen impurities in the carrier gas supplies.

## 2.2 Synthesis of $\text{HNO}_4$

$^{13}\text{N}$  labeled and/or non-labeled  $\text{HNO}_4$  was continuously synthesised in the gas phase by reaction of  $\text{NO}_2$  with  $\text{HO}_2$ . For this, a  $\text{H}_2\text{O}/\text{O}_2/\text{CO}/\text{NO}_2$  mixture in a  $\text{N}_2$  (Carbagas, 99.999%) gas flow passing an 8 mm i.d. quartz tube was irradiated by an excimer UV lamp (Heraeus) emitting light at a wavelength of 172 nm. Typical mixing ratios of the gas phase species at atmospheric pressure were  $3.5 \times 10^{18} \text{ molecules cm}^{-3}$   $\text{H}_2\text{O}$ ,  $2.3 \times 10^{16} \text{ molecules cm}^{-3}$   $\text{CO}$ ,  $1.3 \times 10^{16} \text{ molecules cm}^{-3}$   $\text{O}_2$ , and  $9.4 \times 10^{12} \text{ molecules cm}^{-3}$   $\text{NO}_2$ . Ultrapure water (0.054  $\mu\text{S}$ , Millipore) was dosed by passing the gas flow through a home-built, temperature regulated humidifier consisting of a heated Teflon tube immersed in water. The gases were mixed from certified gas bottles of synthetic air (Carbagas, 20%  $\text{O}_2$  in  $\text{N}_2$

(99.999%)), of  $\text{CO}$  (Carbagas, 10%  $\text{CO}$  in  $\text{N}_2$  (99.999%)), and of  $\text{NO}$  (Carbagas, 10 ppm  $\text{NO}$  in  $\text{N}_2$  (99.999%)) as source for  $\text{NO}_2$ , see below.

$\text{NO}_2$  was quantitatively synthesized by mixing a gas flow containing  $\text{NO}$ , from the PROTRAC target and/or from the certified gas bottle, and  $\text{O}_3$  in a reactor of 2 L volume.  $\text{O}_3$  was produced by irradiation of a flow of dry synthetic air with a Hg pen-ray lamp at 185 nm. The irradiation time and the flow of  $\text{O}_2$  were adjusted to achieve a small excess of  $\text{O}_3$  for full conversion of  $\text{NO}$  to  $\text{NO}_2$ , but not more than  $2 \times 10^{12} \text{ molecules cm}^{-3}$  excess to prevent further oxidation of the  $\text{NO}_2$ .

## 2.3 Detection of $\text{HNO}_4$ and by-products

A chemical ionization mass spectrometer was used to monitor  $\text{HNO}_4$  in the gas phase. The mass spectrometer has been described elsewhere [15], the detection scheme was adapted from Slusher [10]. The strength of this mass spectrometer method is that – without further use of specific chemical traps – several nitrogen oxides can be detected simultaneously based on specific cluster ions [16]. These clusters were produced by reaction with  $\text{SF}_6^-$  in a home-made ionization chamber situated in front of the mass spectrometer entrance orifice. In detail, a flow of  $600 \text{ mL min}^{-1}$  of the gas flow exiting the  $\text{HNO}_4$  synthesis was mixed with a flow of  $1205 \text{ mL min}^{-1}$   $\text{N}_2$  and  $\text{SF}_6^-$  at 11 mbar total pressure. The  $\text{SF}_6^-$  ions were produced by passing  $\text{SF}_6$  (Carbagas, 1%  $\text{SF}_6$  in Ar (99.999%)) in  $\text{N}_2$  through a  $^{210}\text{Po}$ -ionizer (NRD, p-2031). To enhance the formation of ions, a negative voltage of  $-136 \text{ V}$  was applied to the ionizer and the inner walls of the ionisation chamber. Charged clusters entered the mass spectrometer from the ionization chamber *via* an orifice at  $-10 \text{ V}$ .

The following clusters have been described and were also observed in this work:  $(\text{HF})^- \text{NO}_4$  with mass 98 from reaction of  $\text{HNO}_4$  with  $\text{SF}_6^-$  [10],  $(\text{HF})^- \text{NO}_3$  with mass 82 from  $\text{HNO}_3$  [16],  $(\text{HF})^- \text{NO}_2$  with mass 66 for  $\text{HNO}_2$  [17],  $\text{NO}_2^-$  with mass 46 for  $\text{NO}_2$  [16]. A complication of mass spectrometry is that several different species might produce identical fragments.  $\text{HNO}_4$ , for example, has been described to break apart upon reaction with  $\text{SF}_6^-$  leading also to the formation of  $(\text{HF})^- \text{NO}_2$  clusters that are typically used to monitor  $\text{HNO}_2$  [10]. To quantify this effect the gas flow containing the nitrogen oxides was heated to 373 K for quantitative destruction of  $\text{HNO}_4$  to form  $\text{NO}_2$ .  $\text{HNO}_4$  is thermally unstable [18] and the exposure to 373 K at our experimental flow conditions led to a quantitative decomposition of  $\text{HNO}_4$  as verified by observation of the mass spectrometer's signal at a mass to charge ratio ( $m/z$ ) of 98. To heat the gas flow 2 m of the PFA tube were wrapped around an aluminium support and covered by temperature-regulated heating wire (Wisag AG, Switzerland). It was found that the  $(\text{HF})^- \text{NO}_2$  signal decreased by about 10%. From this we conclude that the  $(\text{HF})^- \text{NO}_2$  fragment at  $m/z = 66$  has also a non-negligible contribution from  $\text{HNO}_4$  at our detection conditions. We further found that  $m/z = 66$  also rose in absence of nitrogen oxides, as soon as  $\text{O}_3$  was added. This correlation of  $\text{O}_3$  and the signal at  $m/z = 66$  was observed earlier [19], and is assigned to  $\text{O}_3^- (\text{H}_2\text{O})$  clusters of mass 66. To correct for this, and derive the fraction of the signal that is not

caused by O<sub>3</sub>, O<sub>3</sub> was also monitored at mass 48 (O<sub>3</sub><sup>-</sup>). The ratio of  $m/z = 48$  to  $m/z = 66$  was determined in the absence of nitrogen oxides and used to compute the intensity of  $m/z = 66$  due to O<sub>3</sub> during the synthesis. This later signal was then subtracted from the raw signal at  $m/z = 66$  to give an estimate of the  $m/z = 66$  traces originating from HNO<sub>2</sub>. No interferences were observed for  $m/z = 82$  or  $m/z = 46$ . Overall, with this measurement method, HNO<sub>4</sub>, HNO<sub>3</sub>, and NO<sub>2</sub> can be monitored with high selectivity, while HNO<sub>2</sub> measurements are less reliable. The mass spectrometer data allow direct analysis of relative trends in the individual nitrogen oxide's abundance with changing synthesis settings. For a quantitative analysis the mass spectrometer needs to be calibrated.

Quantification of the nitrogen oxides in the reactor and calibration of the mass spectrometer was done by means of a chemiluminescence NO monitor equipped with a molybdenum converter (Monitor Labs 9841). This converter reduces nitrogen oxides to NO and its use thus allows detecting the sum of all nitrogen oxides present in a sample (NO<sub>x</sub>), by-passing it selectively quantifies NO. Please note that the presence of CO interferes with the NO measurements, so that NO cannot be quantified in this study. NO<sub>x</sub> measurements *via* the converter were not affected by the presence of CO, presumably because the molybdenum converter eliminates CO. To further differentiate individual nitrogen oxides selective chemical traps were used. The performance of the traps was verified by observing the individual traces in the mass spectrometer. HNO<sub>3</sub> is not detected by this NO monitor, because it is removed from the gas flow prior to entering the molybdenum converter by steel components of the instrument. Its concentration can be given by drop in the instrument's signal intensity when the synthesis is started. Scrubbing HNO<sub>4</sub> and HNO<sub>2</sub> from the gas phase in a carbonate trap and measuring the remaining nitrogen oxide content in the gas phase quantifies NO<sub>2</sub>. The carbonate trap was made from firebrick granulate that was soaked with 1.5% aqueous Na<sub>2</sub>CO<sub>3</sub> (Fluka, p.a.) solution, dried, and placed in a 100 × 6 mm glass tube. The ends of the glass tube were filled with glass wool to keep the covered firebrick in its position. Due to their acidity, carbonate traps HNO<sub>4</sub>, and HNO<sub>2</sub>; but neither NO<sub>2</sub> nor NO. HNO<sub>4</sub> was quantified by heating the gas flow to 373 K, by which HNO<sub>4</sub> is converted to NO<sub>2</sub>, before it enters the carbonate trap. The measured gas-phase concentration corresponds then to HNO<sub>4</sub> and NO<sub>2</sub>. From this measurement, also HNO<sub>2</sub> can be derived as fraction of nitrogen oxides that is removed by the carbonate trap. All this requires careful calibration of the nitrogen oxide analyzer that was done with a certified bottle of NO (Carbagas, 10 ppm NO in N<sub>2</sub> (99.999%)).

## 2.4 Packed bed flow tube

The main feature of the packed bed flow tube is a negative temperature gradient along a bed of packed ice spheres – 500 μm in diameter each – as described previously [14]. One end, where the gas flow enters the packed bed flow tube was cooled with a circulating cooling liquid regulated at 250 K. The other end where the gas flow exits the apparatus, was immersed in liquid nitrogen. The temperature inside the flow

tube was measured with a Pt-100 thermo element (MTS, Switzerland) prior to the measurements.

To start an experiment, a packed bed flow tube was placed in the apparatus and exposed to the temperature gradient for 30 min to allow the temperature equilibrium to be reached at any place in the ice flow tube. Then, the carrier gas containing the <sup>13</sup>N- and <sup>14</sup>N-nitrogen oxides was fed into the packed bed flow tube. After 30 min the experiment was stopped, and the flow tube was removed, sealed and immersed in an open bath of liquid nitrogen to stop any further migration of species. In some experiments, the gas flow passed a cooling trap prior to entering the packed bed flow tube to freeze out components from the carrier gas flow with a high partitioning tendency to surfaces. Different geometries and thus surface to volume ratios were used depending on the demands on the capacity. Typically, a quartz tube, 50 mm i.d. and 200 mm in length was filled with quartz spheres, to enhance the surface area, and cooled to 268 K by ethanol cooling liquid circulating around its wall. The cooling trap was operated at least 1 h prior to experiments.

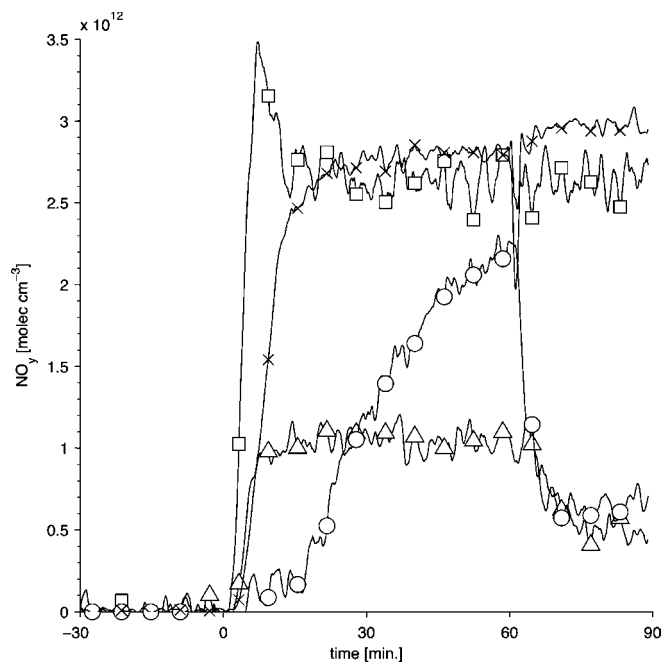
The distribution of the <sup>13</sup>N-nitrogen oxides on the ice surface along the flow tube is the primary observable of the experiment and was measured by means of a coincident γ-counter. The coincident γ-counter consisted of two bismuth-germanate-detectors; 3 cm in diameter, mounted face to face with a gap of 35 mm. Lead shields were used to reduce the detection width to a slit of 5 mm distance. Coincident γ-counting leads to optimum counting efficiency and low background counting rates (1 cps or less), because annihilation of positrons following the β<sup>+</sup>-decay of <sup>13</sup>N results in two γ-rays in opposite direction to each other.

For each experiment, a new packed bed flow tube was prepared as described earlier [7, 14]: Spraying of ultra pure water into liquid nitrogen rapidly froze small ice droplets. After a minimum of 2 d of annealing at 253 K, the ice spheres were sieved with calibrated sieves (Retsch, Germany) to grain sizes between 400 and 600 μm, filled into a PFA tube (8 mm inner diameter, 360 mm length), and stored again for at least 12 h at –20 °C. This preparation was done in a walk-in cold room, during transport to the laboratory, the ice columns were cooled with cooling elements in an insulated box. Based on the mass of the ice filling and the mean diameter of the spheres a mean ice surface area of the packed bed of 20 cm per cm length and a packing density of 70% can be calculated.

## 3. Results and discussion

Irradiation at 172 nm of a H<sub>2</sub>O/NO<sub>2</sub>/O<sub>2</sub> mixture leads to the production of OH and HO<sub>2</sub> radicals by photolysis of water (Eqs. 1 and 2), both of which react with NO<sub>2</sub> to form HNO<sub>3</sub> (Eq. 3) and HNO<sub>4</sub> (Eq. 4), respectively. Central to the synthesis of HNO<sub>4</sub> with high yields is to shift the OH to HO<sub>2</sub> ratio towards an excess of HO<sub>2</sub> and thus amplify HNO<sub>4</sub> yields. Fig. 2 shows that this can be achieved in the presence of CO as scavenger. CO reacts fast with OH but not with HO<sub>2</sub> (Eqs. 2 and 5) [20].





**Fig. 2.** Products of the  $\text{HNO}_4$  synthesis with time at  $8.7 \times 10^{12}$  molecules  $\text{cm}^{-3}$  initial  $\text{NO}_2$ . The line plot shows the calibrated traces as measured with the mass spectrometer of  $\text{HNO}_4$  (line with crosses),  $\text{HNO}_3$  (circles),  $\text{HNO}_2$  (triangles), and  $\text{NO}_2$  (squares). On the relative time axis 0 denotes the time when the  $\text{NO}_2$  flow was switched on to start the synthesis of  $\text{HNO}_4$ . 60 min later the gas flow was passed over a cooling trap at 268 K situated just after the reactor.



Fig. 2 displays the gas phase concentrations of  $\text{HNO}_4$ ,  $\text{HNO}_3$ ,  $\text{HNO}_2$ , and  $\text{NO}_2$  in the reactor with time as measured with the mass spectrometer during irradiation in absence (Fig. 2, -30 to 0 min) and in the presence of  $\text{NO}_2$  (Fig. 2, 0–90 min). It can be clearly seen that, at a high concentration of  $2 \times 10^{16}$  molecules  $\text{cm}^{-3}$  CO,  $\text{HNO}_4$  is the main product with a concentration in the reactor of  $3 \times 10^{12}$  molecules  $\text{cm}^{-3}$ .  $\text{HNO}_3$  and  $\text{HNO}_2$  are also detected but at lower concentrations of  $2 \times 10^{12}$  molecules  $\text{cm}^{-3}$  and  $1 \times 10^{12}$  molecules  $\text{cm}^{-3}$ , respectively.  $\text{HNO}_3$  is most likely formed *via* Eq. (3).  $\text{HNO}_2$  can potentially be formed by reaction of NO with OH, where NO comes from reaction of  $\text{NO}_2$  with O- or with H-radicals. H-radicals are an intermediate product of the reaction 1 and 5, O-radicals originate from photolysis of  $\text{O}_3$ .

Also shown is a significant decrease of 76%  $\text{HNO}_3$  in gas-phase reaching the mass spectrometer when the flow is feed through a cooling trap at 268 K (Fig. 2, 60–90 min). This loss can be explained by the high tendency of  $\text{HNO}_3$  to stick to surfaces even at high temperatures. The mass spectrometer traces of  $\text{HNO}_4$  and of  $\text{HNO}_2$  did also show a response to the cooling trap. But these are caused by variations in humidity in the chemical ionization mass spectrometer by adsorption of water in the cooling trap and not by changes in the concentration of  $\text{HNO}_4$  or  $\text{HNO}_2$ , respectively. This conclusion is supported by independent measurements of the total nitrogen oxide concentration in the gas

flow with the chemiluminescence detector the performance of which is unaffected by gas-phase humidity. The signal from this nitrogen oxide analyzer that is sensitive to NO,  $\text{NO}_2$ ,  $\text{HNO}_2$ , and  $\text{HNO}_4$ , but not to  $\text{HNO}_3$ , did not show any reduction when the gas passed the cooling trap. Additionally, it is well known that the concentration of water in the chemical ionization chamber can significantly affect the formation and yield of ion production [16]. For example, water might exchange with HF in  $(\text{HF})^- \text{NO}_4$  clusters of  $m/z = 98$  leading to  $(\text{H}_2\text{O})^- \text{NO}_4$  clusters of  $m/z = 96$ . A slight decrease in relative humidity, when the gas flow is fed over the cooling trap, might thus explain the higher abundance of  $m/z = 98$  and the decrease in observed intensity of the  $m/z = 96$  cluster (data not shown). Similarly, the cluster  $(\text{HF})^- \text{NO}_2$  with  $m/z = 66$  might be sensitive to humidity. Additionally, and likely more important, to derive the concentration of  $\text{HNO}_2$  based on the measured trace with  $m/z = 66$ , the portion of the mass spectrometer signal originating from the  $\text{O}_3^- \text{H}_2\text{O}$  cluster (also at  $m/z = 66$ ) was subtracted (see Experimental). As the later cluster heavily depends on the water content, this analysis becomes questionable when humidity is not stable during an experiment. As the intensity of the cluster  $(\text{H}_2\text{O})^- \text{NO}_3$  did not change during passage through the cooling trap, we conclude that the cluster  $(\text{HF})^- \text{NO}_3$  does not respond to changes in humidity under our experimental conditions and consequently, the drop in signal intensity observed at  $m/z = 82$  can be fully attributed to loss of  $\text{HNO}_3$  from the gas phase. Note that the temperature of the cooling trap was always above the dew point of water in the carrier gas to prevent riming.

In summary, in presence of CO in large excess,  $\text{HNO}_4$ ,  $\text{HNO}_3$ ,  $\text{HNO}_2$ , and  $\text{NO}_2$  have been identified as products by mass spectrometer measurements. Calibration of the mass spectrometer traces by and comparison to measurements with a chemiluminescence analyzer showed that  $\text{HNO}_4$  is the main product, but that substantial amounts of  $\text{HNO}_3$  and  $\text{HNO}_2$  are present.  $\text{HNO}_3$  can be significantly reduced by use of a cooling trap at 268 K, whereas  $\text{HNO}_4$  and  $\text{HNO}_2$  pass this trap unhindered.

More information on the effect of CO on the product distribution of the photolysis can be seen in Fig. 3 showing the concentration of nitrogen oxides with increasing CO content in the gas phase. In the absence of CO, about 80% of the initial  $\text{NO}_2$  is oxidized, forming primarily  $\text{HNO}_3$ . This synthesis route to  $\text{HNO}_3$  is well established and  $\text{HNO}_3$  yields of up to 90% can be reached at higher humidity [12]. This approach has been used as  $\text{HNO}_3$  source in a number of earlier experiments [21, 22]. The level of  $\text{HNO}_3$  drops rapidly to below 20% with increasing concentration of CO in the reactor and, simultaneously, the  $\text{HNO}_4$  share increases from below 5 to over 30%. The major leap in the  $\text{HNO}_4$  yield is observed at CO concentrations of  $1.0 \times 10^{16}$  molecules  $\text{cm}^{-3}$ , further increase in CO seems to lead to no significant increases in  $\text{HNO}_4$  yields, presumably because the self-reaction of  $\text{HO}_2$  to form  $\text{H}_2\text{O}_2$  becomes more important.  $\text{H}_2\text{O}_2$  has not been quantified in this study due to a lack of a method to calibrate the mass spectrometer signal. A high count rate of  $m/z = 140$  was observed and can be attributed to  $\text{H}_2\text{O}_2$  at our experimental conditions. Bardwell *et al.* have observed and explained this mass as product of a multi-step reaction

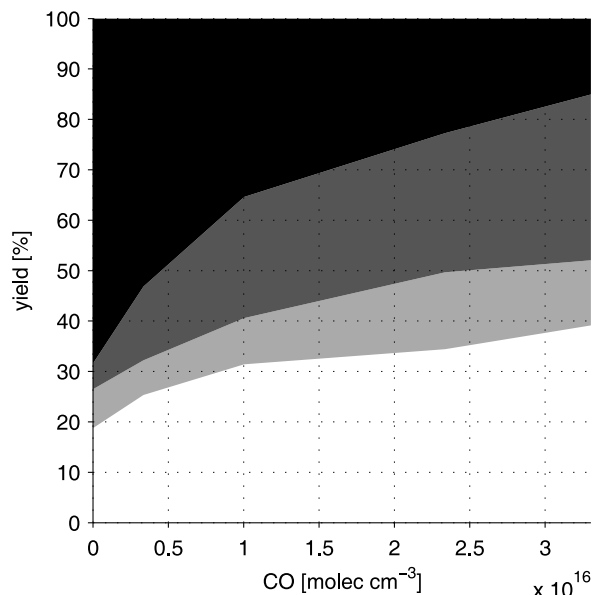
from HO<sub>2</sub> and SF<sub>6</sub><sup>-</sup> previously [23]. We have excluded the presence of HO<sub>2</sub> in the sample gas immediately before the sample inlet by adding NO<sub>2</sub> to the gas flow after the reactor. In the presence of HO<sub>2</sub>, HNO<sub>4</sub> should have been formed, which was not observed. We thus argue that the fragment of  $m/z = 140$  also originates from H<sub>2</sub>O<sub>2</sub>. HNO<sub>2</sub> shows no response to varying CO concentrations with a constant yield of about 10%. NO<sub>2</sub> concentration increases slightly from about 20–40% with increasing CO.

The observed trends for HNO<sub>4</sub> and HNO<sub>3</sub> with increasing CO as determined with the specific traps and the nitrogen oxide analyzer (Fig. 3) are supported by the mass spectrometer measurements. The very selective, but relative, raw mass spectrometer signals of both species at various CO concentrations show the same trend. In contrast, the mass spectrometer data of NO<sub>2</sub> indicate the highest yields in the absence of CO followed by a plateau with increasing CO levels, which is quite different from the increasing NO<sub>2</sub> levels with rising CO concentration as measured with the nitrogen oxide analyzer. An explanation might be that NO turns up as further by-product at high CO concentrations. NO does not interfere with the NO<sub>2</sub> signal of the mass spectrometer, but contributes to the NO<sub>2</sub> budget as determined with the carbonate trap and chemiluminescence analyzer. The same might hold for some organic nitrates that can potentially also be formed in the reactor during the photolysis in the presence of CO.

To increase readability, Figs. 2 and 3 do not show any error bars. Uncertainties of the mass spectrometer measurements, based on the standard deviation of 3 repeated experiments at  $2.3 \times 10^{16}$  molecules cm<sup>-3</sup> CO, are 42% of the mean signal, 18, 14, and 4% for NO<sub>2</sub>, HNO<sub>2</sub>, HNO<sub>3</sub>, and HNO<sub>4</sub>. NO<sub>2</sub> is generally difficult to measure by chemical ionization mass spectrometry, because its reaction with SF<sub>6</sub><sup>-</sup> is slow which makes this measurement method for NO<sub>2</sub> insensitive and uncertain [24]. The reproducibility of HNO<sub>2</sub> measurements is quite low, because of the intense data handling necessary to correct the raw signals for interferences. Measurements of HNO<sub>3</sub> by the chemical ionisation mass spectrometer are very sensitive and accurate, but HNO<sub>3</sub> might be lost at tubing walls during transport to the mass spectrometer resulting in a lower reproducibility of the amount that actually enters the mass spectrometer. HNO<sub>4</sub> measurements show an excellent reproducibility.

The main uncertainty of the chemiluminescence measurements originates most likely from fluctuations of the chemical trap performance. Standard deviations of 16, 19, 6, and 22% for NO<sub>2</sub>, HNO<sub>2</sub>, HNO<sub>3</sub>, and HNO<sub>4</sub> were determined.

To summarize, addition of CO to the gas phase leads to a significant increase of HNO<sub>4</sub> as product with a maximum yield of about 30% of the initial NO<sub>2</sub>. The fraction of HNO<sub>3</sub> decreases significantly with higher CO levels, from the main product in the absence of CO to a share smaller than the HNO<sub>4</sub> content at CO levels higher than  $1.0 \times 10^{16}$  molecules cm<sup>-3</sup>. HNO<sub>2</sub>, on the other hand, shows a rather invariant yield with higher CO levels and accounts for about 1/3 of the HNO<sub>4</sub> concentration. The use of a cold-trap at 268 K further reduces the HNO<sub>3</sub> mixing ratio in the gas flow.



**Fig. 3.** Products of the HNO<sub>4</sub> synthesis with increasing CO concentration. The area graph displays the proportion of each nitrogen oxide as quantified by chemical traps and by the chemiluminescence detector relative to the amount of total nitrogen oxides: NO<sub>2</sub> (white), HNO<sub>2</sub> (light gray), HNO<sub>4</sub> (dark gray), HNO<sub>3</sub> (black).

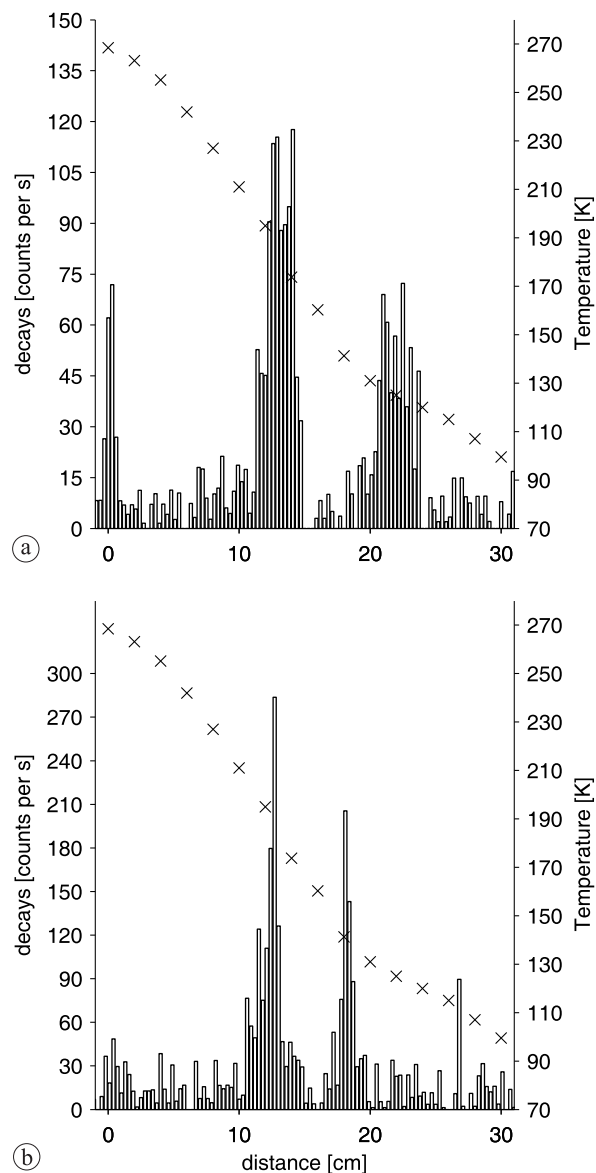
#### 4. Chromatography of <sup>13</sup>N-nitrogen oxides

Based on the above, the following settings were chosen for the production of <sup>13</sup>N labeled HNO<sub>4</sub> during the packed bed flow tube experiments:  $2.3 \times 10^{16}$  molecules cm<sup>-3</sup> CO,  $1.3 \times 10^{16}$  molecules cm<sup>-3</sup> O<sub>2</sub>,  $9.4 \times 10^{12}$  molecules cm<sup>-3</sup> NO<sub>2</sub>,  $3.5 \times 10^{18}$  molecules cm<sup>-3</sup> H<sub>2</sub>O. With these settings about  $2.6 \times 10^{12}$  molecules cm<sup>-3</sup> HNO<sub>4</sub>,  $2.2 \times 10^{12}$  molecules cm<sup>-3</sup> HNO<sub>3</sub>,  $1.4 \times 10^{12}$  molecules cm<sup>-3</sup> HNO<sub>2</sub>, and  $3.2 \times 10^{12}$  molecules cm<sup>-3</sup> NO<sub>2</sub> exited the reactor.

The resulting distribution of this mixture along the packed bed flow tube, as shown in Fig. 4a, reveals three distinct zones of increased radioactive decays centred around temperatures of 267 K with a 95% confidence interval of 252–282 K, 184 K (170–198 K), and of 126 K (116–136 K). The confidence interval reflects uncertainties in determining the centre position of the zone by a Gaussian fit to the data and uncertainties in temperature determination at that position by linear regressions.

The evolution of these peaks is a direct consequence of the partitioning of each species between the gas phase passing through the flow tube and the surface of the packed bed [13]. As this interaction is both strongly temperature dependent and species-specific, mixtures of trace gases are separated during their migration along the packed bed process reflecting their individual partitioning behaviour [14]. Thus the results presented in Fig. 4 indicate that the gas flow contains 3 types of nitrogen oxides species, with weak, intermediate and strong partitioning to the ice phase, respectively.

When the gas flow is passed *via* a cooling trap at 268 K prior to entering the packed bed flow tube, the radioactive decays at the beginning of the packed bed do not exceed the background level and only two peaks are evident at 189 K (175–204 K) and at 135 K (126–144 K). The first peak appears at an identical position and temperature as the cor-



**Fig. 4.** The distribution of radioactive decays from  $^{13}\text{N}$ -labeled nitrogen oxides along packed bed flow tubes. Zero denotes the beginning of the ice bed. Also shown are the temperature profiles inside the flow tube (crosses). (a) The gas flow from the synthesis was directly passed over the packed bed flow tube for 30 min. (b) The gas flow passed a cooling trap at 268 K prior to the packed bed flow tube.

responding peak in Fig. 4a. The second peak is shifted to a shorter distance and slightly higher temperature, but as the confidence intervals overlap this difference is statistically not significant.

The vanished peak at 267 K can clearly be assigned to  $\text{HNO}_3$  based on measurements with the mass spectrometer that showed that  $\text{HNO}_3$  is the only species scrubbed by this cold trap. The strong interaction with ice, or dominant partitioning to the surface at relatively high temperatures, is also supported by recent IUPAC recommendations on the partitioning of  $\text{HNO}_3$  between ice and air, which predict a high surface concentration of  $2 \times 10^3$  molecules  $\text{cm}^{-2}$  in equilibrium for each gas-phase molecule  $\text{cm}^{-3}$  at 267 K [25]. The partitioning of  $\text{NO}_2$  and  $\text{HNO}_2$  to ice is much weaker compared to  $\text{HNO}_3$ : Significant adsorption of  $\text{NO}_2$  to ice has not been observed at temperatures above 195 K [26, 27].  $\text{HNO}_2$  partitioning to ice surfaces occurs at higher tem-

perature (above 170 K) [7, 28, 29]. Consequently, we can assign  $\text{HNO}_2$  deposition to the second zone at 184 K and  $\text{NO}_2$  to the third zone at 126 K. Again, the current IUPAC recommendations for  $\text{HNO}_2$  support this conclusion with  $1 \times 10^4$  molecules  $\text{cm}^{-2}$   $\text{HNO}_2$  on the surface per molecule  $\text{cm}^{-3}$  in the gas-phase at 184 K.

This assessment of  $\text{HNO}_3$ ,  $\text{HNO}_2$ , and  $\text{NO}_2$  is supported by our earlier experiments where we have used selective traps to identify the deposition zones of nitrogen oxides along a negative temperature gradient in a packed bed flow tube [14]. During those earlier experiments settings such as flow velocity (75 and  $360 \text{ cm}^3 \text{ min}^{-1}$ ) and ice surface area in the packed bed flow tube (4 and  $10.9 \text{ cm}^2 \text{ cm}^{-1}$ ) were similar but not identical to those in our current investigation. However, systematic variation of these settings during our earlier work was found not to result in modified retention behaviour. Direct comparison of the deposition properties seems thus well suited.  $\text{NO}_2$  was found to migrate to a position with a temperature of 132 K with a 95% confidence interval of 18 single experiments of  $\pm 14$  K, which is in excellent agreement with 5 recent experiments, for which a mean of 130 K ( $\pm 9$  K) can be determined.  $\text{HNO}_2$  was found to migrate to temperatures at 186 K ( $\pm 20$  K for 8 experiments), which overlaps with the intermediate peak as presented in Fig. Fig. 4 at 191 K ( $\pm 17$  K for 5 experiments).

Li *et al.* have observed that  $\text{HNO}_4$  is released from ice surfaces at temperatures above 210 K, whereas  $\text{HNO}_3$  was only released at higher temperatures of 246 K or above [30]. These results imply for our experimental set-up that  $\text{HNO}_4$  deposits at lower temperatures than  $\text{HNO}_3$ , *i.e.* at the same position as  $\text{HNO}_2$  in the second peak (Fig. 4a). This conclusion is supported by comparing the peak areas as measure for the amount of nitrogen oxides deposited to the concentration of nitrogen oxides in the carrier gas as determined by chemiluminescence measurements. For this, the Gauss fits to the distribution of radioactive decays along the packed bed flow tube were integrated. The areas of peak 1 : peak 2 : peak 3 scale with the ratio 0.5 : 1.7 : 1. Taking that pure  $\text{NO}_2$  deposits in peak 3 and pure  $\text{HNO}_3$  in peak 1 we can compare this to the ratio of  $\text{NO}_2$  :  $\text{HNO}_3$  of 0.7 as determined by quantification of the nitrogen oxide budget by the chemiluminescence analyzer (Fig. 3). This agreement is excellent, taking into account that a fraction of  $\text{HNO}_3$  is lost during its way to the packed bed flow tube by surface adsorption on the tubing walls. The mass balance analysis of nitrogen oxides further reveals that neither  $\text{HNO}_2$ , nor  $\text{HNO}_4$  alone, can explain the observed high intensity of peak 2. The ratio of  $\text{HNO}_2$  to  $\text{NO}_2$  is only 0.4 and the ratio of  $\text{HNO}_4$  to  $\text{NO}_2$  is 0.8; both ratios are less than the ratio of the area of peak 2 to peak 3 of 1.7. Thus it is most likely that both species together deposit in peak 2.

The uptake to ice surfaces heavily depends on the amount of total uptake to the surface, as at high surface concentrations, adsorbate–adsorbate interactions evolve. For our experiments, a surface concentration of  $\text{HNO}_4$  on the ice of  $1 \times 10^{14}$ – $4 \times 10^{14}$  molecules  $\text{cm}^{-2}$  can be calculated. For this calculation, a typical length of the deposition zone of 5 cm,  $\text{HNO}_4$  gas phase concentration entering the packed bed flow tube of  $1 \times 10^{12}$ – $5 \times 10^{12}$  molecules  $\text{cm}^{-3}$  with a flow of  $300 \text{ mL min}^{-1}$  for 30 min, were used. The resulting surface concentration is significantly lower than during the early

study by Li *et al.* in which a formal monolayer, here defined as  $1 \times 10^{15}$  molecules cm<sup>-2</sup>, was exceeded [30].

We conclude that the interaction of HNO<sub>4</sub> with ice surfaces is significantly weaker than observed for HNO<sub>3</sub>, it rather resembles the interaction of HNO<sub>2</sub> with ice. One might thus expect that the partitioning coefficient at the temperature of peak 2 for HNO<sub>4</sub> is similar to the recommendations for HNO<sub>2</sub>, and we propose the same partitioning coefficient  $K_{\text{LinC}}$  of  $7.6 \times 10^{-5}$  cm<sup>-1</sup> also for HNO<sub>4</sub> at 189 K. The partitioning coefficient  $K_{\text{LinC}}$  is defined as concentration of adsorbed species per surface area divided by concentration of gas-phase species per volume in equilibrium and is frequently used for parameterization of surface uptake in atmospheric chemistry models.

## 5. Conclusion and outlook

A new synthesis route to H<sup>13</sup>NO<sub>4</sub> was developed based on the gas-phase reaction of <sup>13</sup>NO<sub>2</sub> + HO<sub>2</sub>. At high concentrations of CO, HNO<sub>4</sub> is the main product with a yield of 30%. HNO<sub>3</sub> and HNO<sub>2</sub> are the most important nitrogen oxide by-products. HNO<sub>3</sub> can be scrubbed from the gas flow in a cold trap at 268 K. Exposure of this mixture to a temperature gradient along a packed ice bed leads to separation of HNO<sub>3</sub> from HNO<sub>2</sub>, HNO<sub>4</sub>, and from NO<sub>2</sub>. This migration behaviour reveals that the interaction of HNO<sub>4</sub> with ice surfaces is similar to that of HNO<sub>2</sub>-ice but much weaker than the HNO<sub>3</sub>-ice interaction. Thus, the surface partitioning coefficient ( $K_{\text{LinC}}$ ) that quantitatively describes the uptake to the ice surface lies in the order of  $7.6 \times 10^{-5}$  cm<sup>-1</sup> also for HNO<sub>4</sub> at 189 K. This implies that the uptake of HNO<sub>4</sub> by surface snow in Antarctica or ice crystals in the upper troposphere, where temperatures of 190 K can be reached, is a very potential loss process.

Precise measurement of the surface partitioning and especially its temperature dependence is urgently recommended. For this isothermal chromatographic methods are well suited [7, 29]. The advantage of isothermal methods is that the highly temperature-dependend partitioning coefficient is the observable of such experiments, and can be directly measured at the temperatures of interest. The strength of the gradient method is rather that it directly reflects the relative intensity of trace-gas-ice interactions for mixtures exposed to the ice surface.

Scrubbing of HNO<sub>2</sub> and HNO<sub>3</sub> from the gas flow is required as the concentration of HNO<sub>2</sub> in the synthesis presented here is too high for isothermal experiments.

*Acknowledgment.* We gratefully thank M. Birrer for the excellent technical support. We acknowledge the staff of the PSI accelerator facilities for supplying stable proton beams. We thank I. Zimmermann and J. Graell for their work on this project during their internships with us. We thank R. Eichler for discussion on the thermochromatography experiments. This project was supported by the Swiss National Science Foundation, project no. 200021\_121857 and the EU FP6 SCOUT-O3 project (GOCE-CT-2004- 505390) funded through the Swiss Federal Office of Education and Science.

## References

- Huthwelker, T., Ammann, M., Peter, T.: [The uptake of acidic gases on ice](#). Chem. Rev. **106**, 1375–1444 (2006).
- Abbatt, J.: [Interactions of atmospheric trace gases with ice surfaces: Adsorption and reaction](#). Chem. Rev. **103**, 4783–4800 (2003).
- Popp, P. J., Marcy, T. P., Watts, L. A., Gao, R. S., Fahey, D. W., Weinstock, E. M., Smith, J. B., Herman, R. L., Troy, R. F., Webster, C. R., Christensen, L. E., Baumgardner, D. G., Voigt, C., Kaercher, B., Wilson, J. C., Mahoney, M. J., Jensen, E. J., Bui, T. P.: Condensed-phase nitric acid in a tropical subvisible cirrus cloud. Geophys. Res. Lett. **34**, L24812 (2007).
- Grannas, A. M., Jones, A. E., Dibb, J., Ammann, M., Anastasio, C., Beine, H., Bergin, M., Bottenheim, J., Boxe, C. S., Carver, G., Chen, G., Crawford, J. H., Dominé, F., Frey, M. M., Guzman, M. I., Heard, D. E., Helmig, D., Hoffmann, M. R., Honrath, R., Huey, L. G., Hutterli, M., Jacobi, H.-W., Klán, P., Lefer, B., McConnell, J., Plane, J., Sander, R., Savarino, J., Shepson, P. B., Simpson, W. R., Sodeau, J. R., Von Glasow, R., Weller, R., Wolff, E. W., Zhu, T.: An overview of snow photochemistry: Evidence, mechanisms and impacts. Atmos. Chem. Phys. **7**, 4329–4373 (2007).
- Dominé, F., Shepson, P.: [Air-snow interactions and atmospheric chemistry](#). Science **297**, 1506–1510 (2002).
- Pinzer, B., Kerbrat, M., Huthwelker, T., Gäggeler, H. W., Schneebeli, M., Ammann, M.: Diffusion of NO<sub>x</sub> and HONO in snow: A laboratory study. J. Geophys. Res. **115**, D03304 (2010).
- Kerbrat, M., Huthwelker, T., Gäggeler, H. W., Ammann, M.: [Interaction of nitrous acid with polycrystalline ice: Adsorption on the surface and diffusion into the bulk](#). J. Phys. Chem. C **114**, 2208–2219 (2010).
- Slusher, D., Huey, L., Tanner, D., Chen, G., Davis, D., Buhr, M., Nowak, J., Eisele, F., Kosciuch, E., Mauldin, R., Lefer, B., Shetter, R., Dibb, J.: Measurements of pernitric acid at the South Pole during ISCAT 2000. Geophys. Res. Lett. **29** (2002).
- Jaeglé, L., Jacob, D., Brune, W., Wennberg, P.: [Chemistry of HO<sub>2</sub> radicals in the upper troposphere](#). Atmos. Environ. **35**, 469–489 (2001).
- Slusher, D., Pitteri, S., Haman, B., Tanner, D., Huey, L.: [A chemical ionization technique for measurement of pernitric acid in the upper troposphere and the polar boundary layer](#). Geophys. Res. Lett. **28**, 3875–3878 (2001).
- Staikova, M., Donaldson, A., Francisco, J.: [Overtone-induced reactions on the HO<sub>2</sub>/NO<sub>2</sub> potential surface](#). J. Phys. Chem. A **106**, 3023–3028 (2002).
- Ammann, M.: Using <sup>13</sup>N as tracer in heterogeneous atmospheric chemistry experiments. Radiochim. Acta **89**, 831–838 (2001).
- Eichler, B., Zvara, I.: Evaluation of the enthalpy of adsorption from thermochromatographical data. Radiochim. Acta **30**, 233–238 (1982).
- Bartels-Rausch, T., Eichler, B., Zimmermann, P., Gäggeler, H. W., Ammann, M.: [The adsorption enthalpy of nitrogen oxides on crystalline ice](#). Atmos. Chem. Phys. **2**, 235–247 (2002).
- Guimbaud, C., Bartels-Rausch, T., Ammann, M.: An atmospheric pressure chemical ionization mass spectrometer (APCI-MS) combined with a chromatographic technique to measure the adsorption enthalpy of acetone on ice. Int. J. Mass. Spectrom. **226**, 279–290 (2003).
- Huey, L. G.: [Measurement of trace atmospheric species by chemical ionization mass spectrometry: Speciation of reactive nitrogen and future directions](#). Mass. Spectrom. Rev. **26**, 166–184 (2007).
- Longfellow, C., Imamura, T., Ravishankara, A., Hanson, D.: [HONO solubility and heterogeneous reactivity on sulfuric acid surfaces](#). J. Phys. Chem. A **102**, 3323–3332 (1998).
- Zabel, F.: [Unimolecular decomposition of peroxy nitrates](#). Z. Phys. Chem. **188**, 119–142 (1995).
- Abbatt, J., Oldridge, N., Symington, A., Chukalovskiy, V., McWhinney, R. D., Sjostedt, S., Cox, R. A.: [Release of gas-phase halogens by photolytic generation of OH in frozen halide-nitrate solutions: An active halogen formation mechanism?](#) J. Phys. Chem. A **114**, 6527–6533 (2010).
- Schultz, M., Heitlinger, M., Mihelcic, D., Volzthomas, A.: Calibration source for peroxy-radicals with built-in actinometry using H<sub>2</sub>O and O<sub>2</sub> photolysis at 185 nm. J. Geophys. Res.-Atmos. **100**, 18811–18816 (1995).
- Vlasenko, A., Huthwelker, T., Gäggeler, H. W., Ammann, M.: Kinetics of the heterogeneous reaction of nitric acid with mineral

- dust particles: An aerosol flowtube study. *Phys. Chem. Chem. Phys.* **11**, 7921–7930 (2009).
22. Guimbaud, C., Arens, F., Gutzwiller, L., Gäggeler, H. W., Ammann, M.: Uptake of  $\text{HNO}_3$  to deliquescent sea-salt particles: A study using the short-lived radioactive isotope tracer N-13. *Atmos. Chem. Phys.* **2**, 249–257 (2002).
  23. Bardwell, M., Bacak, A., Raventos, M., Percival, C., Sanchez-Reyna, G., Shallcross, D.: Kinetics of the  $\text{HO}_2 + \text{NO}$  reaction: A temperature and pressure dependence study using chemical ionisation mass spectrometry. *Phys. Chem. Chem. Phys.* **5**, 2381–2385 (2003).
  24. Huey, L., Hanson, D., Howard, C.: Reactions of  $\text{SF}_6^-$  and  $\text{I}^-$  with atmospheric trace gases. *J. Phys. Chem.* **99**, 5001–5008 (1995).
  25. Ammann, M., Atkinson, R., Cox, R., Crowley, J., Hynes, R., Jenkin, M., Rossi, M., Troe, J., Wallington, T., IUPAC Subcommittee: Evaluated kinetic and photochemical data for atmospheric chemistry: Heterogeneous reactions on ice. <http://www.iupac-kinetic.ch.cam.ac.uk/index.html> (2008).
  26. Saastad, O., Ellermann, T., Nielsen, C.: On the adsorption of NO and  $\text{NO}_2$  on cold  $\text{H}_2\text{O}/\text{H}_2\text{SO}_4$  surfaces. *Geophys. Res. Lett.* **20**, 1191–1193 (1993).
  27. Leu, M.: Heterogeneous reactions of  $\text{N}_2\text{O}_5$  with  $\text{H}_2\text{O}$  and HCl on ice surfaces – implications for antarctic ozone depletion. *Geophys. Res. Lett.* **15**, 851–854 (1988).
  28. Fenter, F., Rossi, M.: Heterogeneous kinetics of HONO on  $\text{H}_2\text{SO}_4$  solutions and on ice: Activation of HCl. *J. Phys. Chem.* **100**, 13765–13775 (1996).
  29. Chu, L., Diao, G., Chu, L.: Heterogeneous interaction and reaction of HONO on ice films between 173 and 230 K. *J. Phys. Chem. A* **104**, 3150–3158 (2000).
  30. Li, Z., Friedl, R., Moore, S., Sander, S.: Interaction of peroxyntitric acid with solid  $\text{H}_2\text{O}$  ice. *J. Geophys. Res.* **101**, 6795–6802 (1996).