

1 **Effects of elevated CO₂ on phytoplankton community biomass and**
2 **species composition during a spring *Phaeocystis* spp. bloom in the**
3 **western English Channel**

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15 Running head: Effects of elevated CO₂ during a spring *Phaeocystis* bloom at coastal station L4 in the
16 western English Channel.

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21 **Abstract**

22 A 21-year time series of phytoplankton community structure was analysed in relation to
23 *Phaeocystis* spp. to elucidate its contribution to the annual carbon budget at station L4 in the
24 western English Channel (WEC).

25 Between 1993-2014 *Phaeocystis* spp. contributed ~4.6% of the annual phytoplankton carbon
26 budget at station L4. During the March – May spring bloom period, the mean *Phaeocystis* spp.
27 biomass constituted 17% with a maximal contribution of 47% in 2001. Upper maximal weekly
28 values above the time series mean ranged from 63 – 82% of the total phytoplankton carbon
29 (~42 – 137 mg carbon (C) m³) with significant inter-annual variability in *Phaeocystis* spp..
30 Maximal biomass usually occurred by the end of April, although in some cases as early as mid-
31 April (2007) and as late as late May (2013).

32 The effects of elevated pCO₂ on the *Phaeocystis* spp. spring bloom were investigated during a
33 fifteen-day semi-continuous microcosm experiment. The phytoplankton community biomass
34 was estimated at ~160 mg carbon C m³ and was dominated by nanophytoplankton (40%,
35 excluding *Phaeocystis* spp.), *Phaeocystis* spp. (30%) and cryptophytes (12%). The smaller
36 fraction of the community comprised picophytoplankton (9%), coccolithophores (3%),
37 *Synechococcus* (3%), dinoflagellates (1.5%), ciliates (1%) and diatoms (0.5%). Over the
38 experimental period, total biomass increased significantly by 90% to ~305 mg C m³ in the high
39 CO₂ treatment while the ambient pCO₂ control showed no net gains. *Phaeocystis* spp. exhibited
40 the greatest response to the high CO₂ treatment, increasing by 330%, from ~50 mg C m³ to over
41 200 mg C m³ and contributing ~70% of the total biomass.

42 Taken together, the results of our microcosm experiment and analysis of the time series suggest
43 that a future high CO₂ scenario may favour dominance of *Phaeocystis* spp. during the spring
44 bloom. This has significant implications for the formation of hypoxic zones and the alteration of
45 food web structure including inhibitory feeding effects and lowered fecundity in many copepod
46 species.

47 KEY WORDS: Ocean acidification, *Phaeocystis* spp., natural phytoplankton populations, western
48 English Channel.

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50 1. Introduction

51 While coastal zones account for just 7% of the global ocean surface, their role in the global
52 carbon cycle is crucial (Wollast, 1998). Supporting an estimated 10-15% of global ocean net
53 annual primary production, coastal regions are responsible for more than 40% of oceanic
54 carbon sequestration (Muller-Karger, 2005). Atmospheric CO₂ concentration has increased by
55 around 33% over pre-industrial levels, with an on-going annual increase of ~0.4% (Wolf-
56 Gladrow et al. 1999, Raven et al. 2005, Alley et al. 2007). The dissolution of this excess CO₂ into
57 the surface ocean directly affects the carbonate system which has lowered pH by ~0.1 units,
58 from 8.21 to 8.10 over the last ~250 years. Further decreases are predicted by 0.3-0.4 pH units
59 by the end of this century (Doney et al., 2009; Orr et al., 2005), a phenomenon commonly
60 referred to as ocean acidification (OA). The physiological and ecological aspects of the
61 phytoplankton response to this changing environmental factor holds the potential to alter
62 phytoplankton community composition, community biomass and to feedback to biogeochemical
63 cycles (Boyd and Doney, 2002).

64 Marine phytoplankton have been shown to exhibit sensitivity to elevated partial pressure of
65 CO₂ in seawater (pCO₂) in growth and photosynthetic rates, in both laboratory studies using
66 model species in culture and on natural populations in the field (e.g. Endo et al., 2013; Eggers et
67 al., 2013; Feng et al., 2009; Hare et al., 2007; Schulz et al., 2008; Tortell et al., 2002). Since a
68 wide variety of processes are affected, high variability in responses has been reported across
69 and within taxa. For example, the response of diatoms under elevated pCO₂ is not straight
70 forward. Diatom dominated natural communities exhibited no increase in growth under pCO₂
71 elevated to 800 µatm during shipboard incubations (Tortell et al., 2000). The diatom
72 *Skeletonema costatum* also showed no increase in growth during laboratory studies at 800 µatm

73 pCO₂ (Burkhardt and Riebesell, 1997), but increased growth rates at 750 µatm pCO₂ during a
74 mesocosm experiment (Kim et al., 2006). Feng et al., (2010) observed dominance of the large
75 centric *Chaetoceros* spp. relative to the smaller pennate *Cylindrotheca closterium* regardless of
76 other experimental factors during shipboard incubations under pCO₂ elevated to 760 ppm.
77 Conversely, Coello-Camba et al., (2014) observed significant increases in growth of smaller
78 centric diatoms ($\leq 7 \mu\text{m}$) and a decline in growth rates of larger centric diatoms ($\geq 11 \mu\text{m}$)
79 under pCO₂ elevated to 1000 ppm during bottle incubations of an arctic phytoplankton
80 community. Coccolithophores exhibit no change in growth rates but increased particulate
81 organic carbon content and decreased inorganic carbon content (calcification) (Barcelos e
82 Ramos et al., 2010; Feng et al., 2008), whereas for the lesser-studied *Phaeocystis* spp. a decrease
83 or no change in growth rates have been observed (Chen and Gao, 2011; Thoisen et al., 2015).
84 The few studies on natural populations suggest that elevated pCO₂ may lead to a shift in
85 community composition with consequences for overall rates of primary production through the
86 pCO₂ influence on photosynthesis, elemental composition and calcification of marine
87 phytoplankton (Riebesell, 2004).

88 *Phaeocystis* spp. are ubiquitous with a unique polymorphic life-cycle, alternating free-living
89 solitary ($\sim 6 \mu\text{m}$ in size) and colonial ($\sim 2 \text{mm}$ in diameter) cells, a process that changes organism
90 bio-volume by 6 to 9 orders of magnitude (Verity et al., 2007). As such *Phaeocystis* spp. can
91 outcompete other phytoplankton and form massive blooms (up to 10 mg C m^{-3}) with impacts on
92 food webs, global biogeochemical cycles and climate regulation (Schoemann et al., 2005). Since
93 *Phaeocystis* spp. produce dimethylsulfoniopropionate (DMSP) their blooms also provide an
94 important source of dimethylsulphide (DMS) (Stefels et al., 1995) playing a key role in the
95 transfer of carbon and sulphur between ocean and atmosphere and vice versa (Liss et al., 1994).
96 While not a highly toxic algal species, *Phaeocystis* spp. are considered a harmful algal species
97 (HAB) when biomass reaches sufficient concentrations to cause anoxia and mucus foam which
98 can clog the feeding apparatus of zooplankton and fish (Eilertsen & Raa, 1995). Along the
99 European coasts of the North Sea dense blooms of *Phaeocystis globosa* known to impact

ecosystem function have been well documented as localised events for many years (Lancelot and Mathot, 1987). More recently, other continents with nutrient enriched waters have reported massive blooms such as the Arabian Gulf and southeast coastal waters of China (Lancelot et al., 2002; Schoemann et al., 2005). In these ecosystems, *P. globosa* bloom formation is predominantly a result of anthropogenic factors, i.e. nutrient inputs via riverine and land runoff routes (Cadée and Hegeman, 2002). Consequently, *Phaeocystis* spp. have been identified as key water disturbance indicator species (Tett et al., 2007) and recommendations for decreasing its abundance to that of non-problem areas and good ecological status have been made in the scope of the OSPAR strategy (Ospar, 2005) and the Water Framework Directive of the European Union (2000/60/EC) (Lancelot et al., 2009). Dense blooms of *Phaeocystis antarctica* colonies are also observed in naturally nutrient rich waters, however, such as the Ross Sea (DiTullio et al., 2000), Greenland Sea and Barents Sea (Eilertsen et al., 1989; Wassmann et al., 1990). Given the ecological, global biogeochemical and climate regulation relevance, *Phaeocystis* spp. are a highly suitable and a significant model phytoplankton species to study in the context of OA.

Phytoplankton species composition, abundance and biomass has been measured at the time-series station L4 in the western English Channel (WEC) since 1992, to evaluate how global changes could drive future shifts in phytoplankton community structure and carbon biogeochemistry. The goals of the present study were to investigate: 1) the effects of elevated pCO₂ on phytoplankton community structure and the relative species contribution to community biomass during the spring bloom succession to *Phaeocystis* spp. and 2) assess the natural variability in phytoplankton community structure and the carbon biomass of *Phaeocystis* spp. at station L4 over two decades (1993-2014). We hypothesized that community biomass will increase under a high CO₂ regime and that succession to *Phaeocystis* spp. may be reduced in magnitude.

2. Materials and Methods

125

126 **2.1. Time series – Phytoplankton community composition**

127 Station L4 (50° 15'N, 4° 13'W) is located 13km SSW of Plymouth in a water depth of ~54m
128 (Harris, 2010) and is regarded as one of Europe's principal coastal time series sites. Sampling is
129 conducted on a weekly basis (weather permitting) and has been on-going since 1988
130 (<http://www.westernchannelobservatory.org>). Phytoplankton taxonomic composition was
131 enumerated from seawater samples collected from 10m depth, fixed with 2% (final
132 concentration) Lugol's iodine solution and analysed by inverted light microscopy using the
133 Utermöhl counting technique (Utermöhl, 1958; Widdicombe *et al.*, 2010). For phytoplankton
134 carbon biomass values; taxa-specific mean cell bio-volumes were calculated following Kovala &
135 Larrance, (1966) and converted to carbon using the equations of Menden-Deuer & Lessard,
136 (2000). Additionally, samples for the current study were taken from 10 m depth for the analysis
137 of the macro-nutrients; nitrite, nitrate, silicate, phosphate and ammonium and analysed using
138 the methods described by Woodward & Rees, (2001).

139 **2.2. Elevated pCO₂ perturbation experiment**

140 Experimental seawater containing a natural phytoplankton community was sampled at station
141 L4 (50° 15'N, 4° 13'W) (**Fig 1.**) on 13th April 2015 via a high volume, wide aperture trace-metal
142 clean manual diaphragm pump system from 10 m depth (40 L volume). The experimental
143 seawater was pre-filtered through a 200 µm Nitex mesh to remove zooplankton grazers, into
144 two 20 L acid-cleaned carboys. In addition, 320 L of seawater was collected into sixteen 20 L
145 acid-cleaned carboys from the same depth for use as experimental media. Immediately upon
146 return to the laboratory the media seawater was filtered through an in-line 0.2 and 0.1 µm filter
147 (Acropak™, Pall Life Sciences) then stored in the dark at 11°C until use. The experimental
148 seawater was gently and thoroughly mixed and transferred in equal parts from each carboy (to
149 ensure homogeneity) to sixteen 2.5 L borosilicate incubation bottles (2 sets of 8 replicates). The
150 remaining experimental seawater was sampled for initial (T₀) concentrations of nutrients,
151 chlorophyll *a*, total alkalinity and dissolved inorganic carbon and was also used to characterise

152 the starting experimental phytoplankton community. A semi-continuous closed incubation
153 culture system linked the replicate incubation bottles to two 22 L media reservoirs filled with
154 the filtered seawater media which was aerated with CO₂ free air and 5% CO₂ in air, precisely
155 mixed using a mass flow controller (Bronkhorst UK Limited). This CO₂ enriched seawater media
156 was then used for the microcosm dilutions, adjusted as per the following treatments: 1)
157 Ambient pCO₂ (control at ~340 μatm, matching station L4 in situ values) and 2) Elevated pCO₂
158 (high CO₂ at ~800 μatm, predicted for the end of this century assuming the IPCC 'business as
159 usual' scenario' (Alley et al., 2007)).

160 Initial nutrient concentrations (measured at 1.4 μM nitrate + nitrite and 0.05 μM phosphate on
161 13th April 2015) were amended to 8μM nitrate+nitrite and 0.5 μM phosphate replicating mean
162 pre-spring bloom values at station L4. As the phytoplankton community was in the transitional
163 phase from diatoms to nanophytoplankton, the in situ silicate concentration was maintained to
164 reproduce the silicate concentrations typical of this time of year (Smyth et al., 2010). Media
165 transfer and sample acquisition was facilitated by peristaltic pumps and daily dilution rates
166 were set between 10-13% of the incubation bottle volume following 24 hrs acclimation. Thus,
167 CO₂ enriched seawater was added to the high CO₂ treatment replicates every 24 hrs, acclimating
168 the natural phytoplankton population to increments of elevated pCO₂ from ambient to ~800
169 μatm over 8 days followed by maintenance at ~800 μatm as per the method described by
170 Schulz *et al.*(2009). This protocol was preferred since some phytoplankton species are inhibited
171 by the mechanical effects of direct bubbling (Riebesell et al., 2010; Shi et al., 2009) which can
172 cause a reduction in growth rates and the formation of aggregates (Love et al., 2016).

173 All glassware was autoclaved and all sampling equipment, PTFE and Marprene™ tubing was
174 rigorously washed with 10% HCl, rinsed thoroughly with milli-Q ultrapure™ water (Millipore
175 Corporation, USA), followed by a thorough rinse with FSW. The incubation bottles were
176 maintained at 11 °C in a flow-through seawater bath (temperature was monitored twice daily)
177 to replicate in situ temperature on the day of sampling. Light was supplied by a cool white LED

178 light bank at irradiance of $\sim 200 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$ on a 16 : 8 hour light : dark cycle.
179 Incubation bottles were inverted and gently agitated twice each day to maintain phytoplankton
180 cells in suspension and also prior to sampling to ensure homogeneity.

181 **2.3. Analytical methods – experimental seawater**

182

183 **2.3.1. Chlorophyll *a***

184 Chlorophyll *a* (chl *a*) was measured every three days in each incubation bottle. 100 mL triplicate
185 samples from each replicate were filtered onto 25 mm GF/F filters, extracted in 90% acetone
186 overnight at -20 °C and chl *a* was estimated on a Turner Trilogy™ fluorometer using the non-
187 acidified method of Welschmeyer (1994). The fluorometer was calibrated against a stock chl *a*
188 standard (*Anacystis nidulans*, Sigma Aldrich, UK), the concentration of which was determined
189 with a Perkin Elmer™ spectrophotometer at wavelengths 663.89 and 750.11 nm.

190 **2.3.2. Carbonate system**

191 70 mL samples for total alkalinity (TA) and dissolved inorganic carbon (DIC) analysis were
192 collected every three days from each experimental replicate, stored in amber borosilicate
193 bottles with no head space and fixed with 40 μL of super-saturated Hg_2Cl_2 solution for later
194 determination (Apollo SciTech™ Alkalinity Titrator AS-ALK2; LiCOR™ 7000 DIC analyser).
195 Duplicate measurements were made for TA and triplicate measurements for DIC. Carbonate
196 system parameter values for media and treatment samples were calculated from TA and DIC
197 measurements using the programme CO₂sys (Pierrot et al., 2006) with dissociation constants of
198 carbonic acid of Mehrbach *et al.*, (1973) refitted by Dickson and Millero (Dickson and Millero,
199 1987).

200 **2.3.3. Phytoplankton community analysis**

201 Phytoplankton community analysis was performed by flow cytometry (Becton Dickinson Accuri
202™ C6) for the 0.2 to 18 μm size fraction following Tarran *et al.*, (2006). FlowCAM™ (Fluid

203 Imaging Technologies™) flow-through analysis was used for the 18-100 µm size fraction
 204 following the method described by Poulton & Martin, (2010) and light microscopy was used to
 205 enumerate cells > 100 µm (BS EN 15204,2006). For flow cytometry, 2 mL samples fixed with
 206 glutaraldehyde to a final concentration of 2% were taken every five days, flash frozen in liquid
 207 nitrogen and stored at -80 °C for later analysis. For FlowCAM, 250 mL samples fixed with acid
 208 Lugol's iodine to a final concentration of 1% were also taken every five days and stored in cool,
 209 dark conditions. Concentrated aliquots were filtered through a 100 µm Nitex mesh prior to
 210 analysis to avoid larger phytoplankton cells blocking the FlowCAM flow cell. The Nitex mesh
 211 was back-washed by gently pipetting 3mL of sample media on the reverse side to remove
 212 residual large cells. This was carefully decanted into Hydro-Bios™ settling chambers prior to
 213 analysis using an Olympus™ IMT-2 inverted light microscope. All cells in the settling chambers
 214 were identified and enumerated and the calculated cell concentrations combined with the
 215 FlowCAM data.

216 **2.3.4. Phytoplankton community biomass**

217 The smaller size fraction identified and enumerated through flow cytometry;
 218 picophytoplankton, nanophytoplankton, *Synechococcus*, coccolithophores and cryptophytes
 219 were converted to carbon biomass (mg C m³) using a spherical model to calculate mean cell
 220 volume:

$$221 \left(\frac{4}{3} * \pi * r^3\right)$$

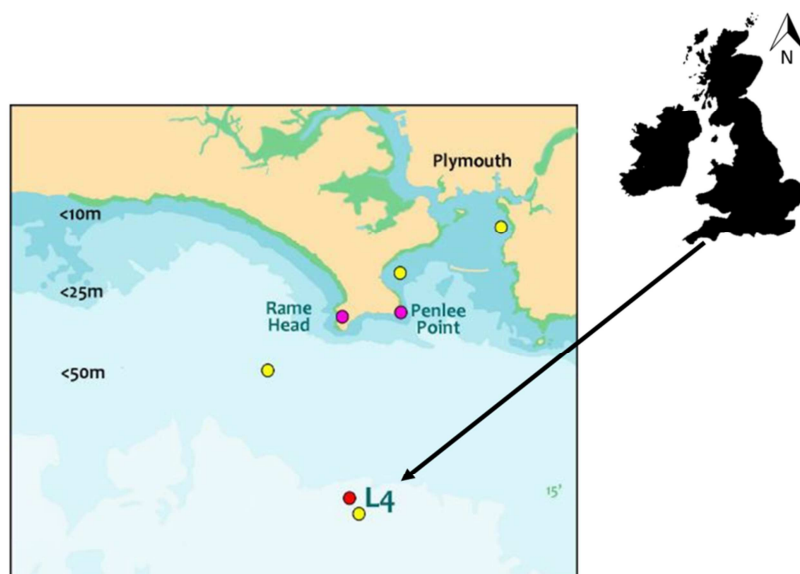
222 and a conversion factor of 0.22 pg C µm⁻³ (Booth, 1988). A conversion factor of 0.285 pg C µm⁻³
 223 was used for coccolithophores (Tarran et al., 2006) and a cell volume of 113 µm³ and carbon
 224 cell⁻¹ value of 18 pg applied for *Phaeocystis* spp. (Widdicombe *et al.*, 2010) . *Phaeocystis* spp.
 225 were identified and enumerated separately to the nanophytoplankton class due to high
 226 observed abundance. Mean cell measurements of individual species/taxa were used to calculate

227 cell bio-volume for the 18 μm + size fraction according to Kovala and Larrance (1966) and
228 converted to biomass according to the equations of Menden-Deuer & Lessard, (2000).

229

230 **2.4. Statistical analysis**

231 Weekly biomass values from the L4 time-series were averaged over years to elucidate the
232 variability and seasonal cycles of *Phaeocystis* spp. and total phytoplankton carbon biomass
233 (excluding *Phaeocystis* spp.). Trends in *Phaeocystis* spp. biomass over time were investigated
234 using linear regression performed with a Generalised Least Squares model (gls) incorporating
235 an auto-regressive correlation structure of the order (1) to account for auto correlation. A
236 generalised linear model (glm) was used to test for differences in biomass between years. In
237 order to test for effects of high CO_2 and to account for possible time dependence of the
238 measured response variables (Chl *a*, C:chl *a*, total community biomass and biomass of individual
239 species), glm models with the factors pCO_2 and time were applied to the data following target
240 pCO_2 equilibration between days (T)9 and (T)14. Where main effects were established, pairwise
241 comparisons were performed using the method of Herberich *et al.*, (2010) for data with non-
242 normality and/or heteroscedasticity. Analyses were conducted using the R statistical package (R
243 Core Team (2014). R: A language and environment for statistical computing. R Foundation for
244 Statistical Computing, Vienna, Austria).



245

246 **Fig 1.** Location of coastal station L4, western English Channel247 **3. Results**

248

249 **3.1. Station L4 time-series, *Phaeocystis* spp. biomass in the WEC**

250 Over the time series from 1993 to 2014, the annual mean total phytoplankton biomass sampled
 251 at L4 was 1646 (\pm 521 sd) mg C m³, with annual mean *Phaeocystis* spp. biomass of 72 (\pm 69 sd)
 252 mg C m³. Maximum total annual phytoplankton biomass occurred in 1997 (3206 mg C m³) and
 253 minimum values were in 2007 (998 mg C m³) when the associated annual *Phaeocystis* spp.

254 biomass were 121 and 75 mg C m³ respectively (**Fig 2. A.**). *Phaeocystis* spp. contributed 4.6% of
 255 the total phytoplankton annual carbon budget, which peaked at ~16% in 2001 (~270 mg C m³).

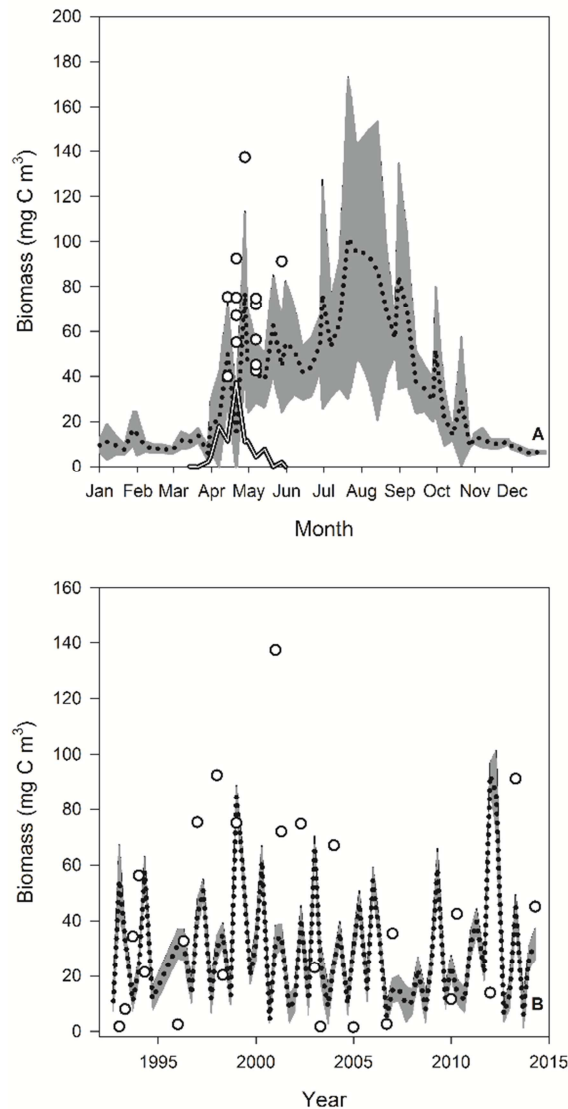
256 **Fig 2. B.** shows the biomass trends over the March – May seasonal bloom period. *P. globosa* and
 257 *P. pouchetii* were both recorded at the L4 time-series site, but were grouped as *Phaeocystis* spp.
 258 due to the inherent difficulties in distinguishing single cells using microscopy.

259 Weekly *Phaeocystis* spp. biomass recorded from January to December throughout the time
 260 series ranged from below the limit of detection to 137 mg C m³ (**Fig 3. A.**). Bloom initiation

261 occurred as early as mid to late March (2007 and 2011). The bloom peak (taken as an increase

262 in biomass $> 0.5 \text{ mg C m}^3$) usually occurred by the end of April, though in 2007 it was mid-April
263 and in 2013 it was late May. The mean yearly maximal biomass peak was $41.6 (\pm 39.3 \text{ sd}) \text{ mg C}$
264 m^3 (**Fig 3. B.**). The generalised linear model highlighted significant inter-annual variability in
265 *Phaeocystis* spp. biomass between 1993 and 2014 with biomass in 2001 significantly greater
266 than any other year throughout the time series period, when maximal biomass reached 137 mg
267 C m^3 ($z = 3.355, p < 0.001$). Half of the 21 years analysed showed a maximal peak *Phaeocystis*
268 spp. biomass range between $42 - 137 \text{ mg C m}^3$ (above the time series mean maxima peak),
269 significantly higher than all other years over the time-series ($z = -6.695, p < 0.0001$). Biomass as
270 low as between $0.09 - 1.5 \text{ mg C m}^3$ however, was observed during the seasonal bloom in 6 out of
271 the 21 years (**Table 3.**).

272



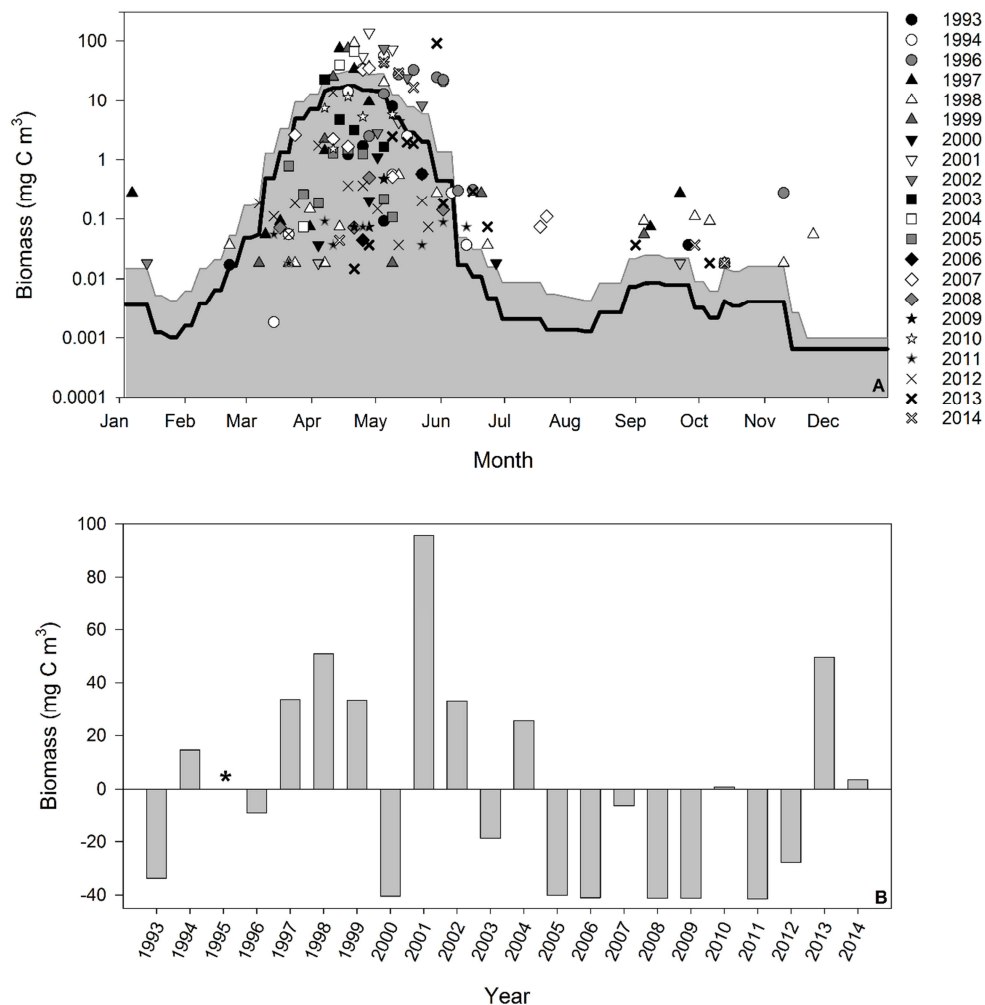
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274 **Fig 2. A.** Temporal monthly profile of total phytoplankton carbon biomass at station L4 between
 275 1993 – 2014. Dotted line is weekly mean total phytoplankton excluding *Phaeocystis* spp., grey
 276 area is standard deviation, inverse white line is mean *Phaeocystis* spp. carbon biomass and open
 277 circles are maximal weekly *Phaeocystis* spp. biomass values above the time series mean maxima
 278 value of 41.6 mg C m³. **B.** Temporal seasonal profile of total phytoplankton carbon biomass
 279 during the spring bloom period (March – May) between 1993 – 2014. Dotted line is mean total
 280 phytoplankton excluding *Phaeocystis* spp., grey area is standard deviation and open circles are
 281 monthly maximal *Phaeocystis* spp. biomass values.

282

283 Total phytoplankton biomass peaked during the spring bloom period, typically between March
 284 and May and was highest (215 mg C m³) when *Phaeocystis* spp. biomass increased above the
 285 time series mean peak maxima (to 56 mg C m³). Throughout the time-series *Phaeocystis* spp.
 286 maximal biomass (above the mean maximal biomass) contributed between 26% (1994) – 82%

287 (1998 and 2013) of phytoplankton carbon, while total *Phaeocystis* spp. biomass over the March
 288 – May bloom period contributed a mean value of 17% to total phytoplankton carbon, between
 289 20%-47% during years of high biomass above the mean maxima (i.e. 2001 and 2010) and
 290 between 0.1% to 13.9% during years of low biomass (i.e. 2006, 2001 and 1999). Mean bloom
 291 duration over the time-series was 22 days with a range between 6 days (2000) and 44 days
 292 (1994). Biomass showed a positive slope from 1993-2014 indicating an overall increase in
 293 *Phaeocystis* spp. carbon over the time-series, however this relationship was not significant ($T =$
 294 $1.86, p = 0.07$).



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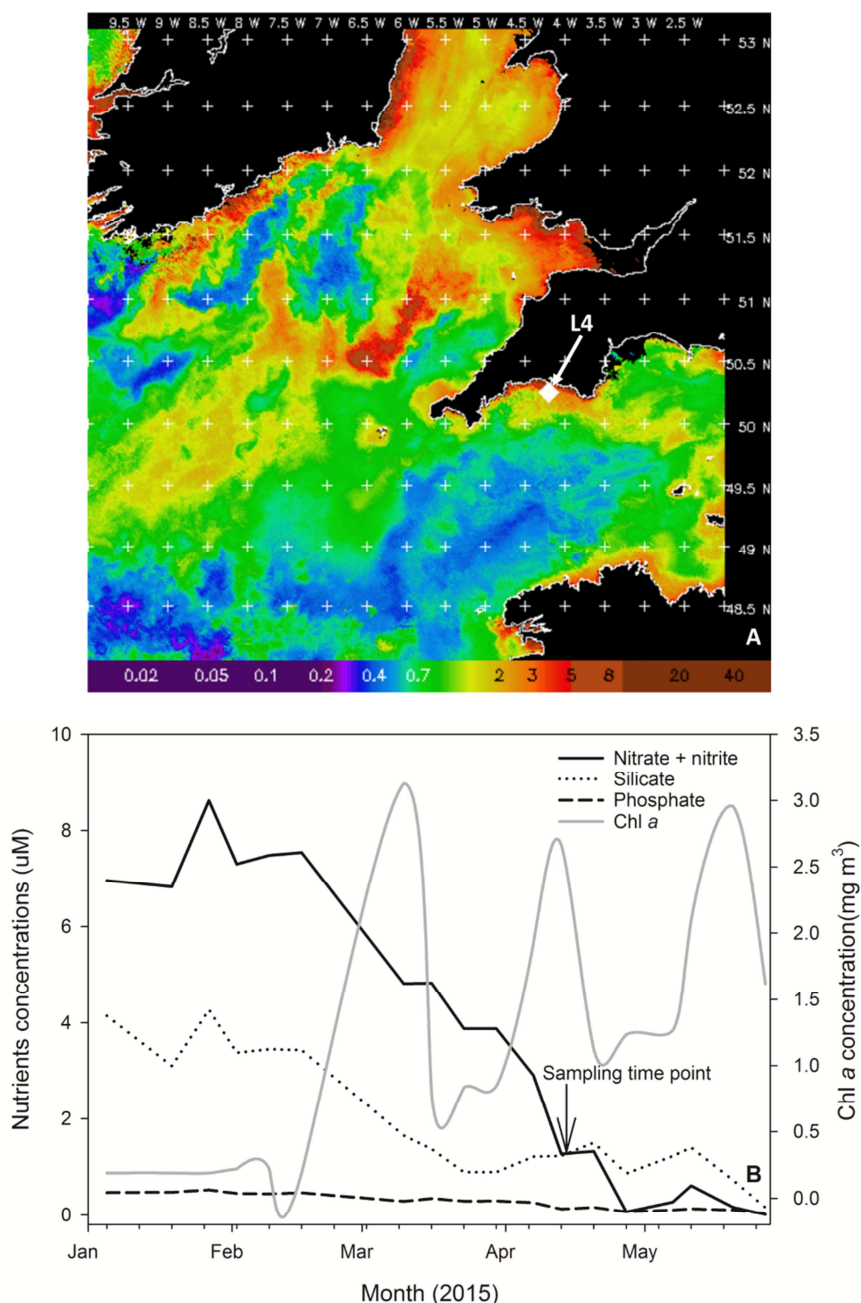
296 **Fig 3. A.** Seasonal profiles of *Phaeocystis* spp. carbon biomass (common log scale) between 1993
 297 – 2014. Black line is smoothed running average over the time-series, grey area is the standard
 298 deviation and all symbols are observed data values by year. **B.** Annual anomalies of maximal
 299 *Phaeocystis* spp. carbon biomass above and below the time series mean maxima of 41.6 mg C m⁻³.
 300 * Insufficient data 1995.

301 **3.2. Elevated pCO₂ perturbation experiment**

302

303 Chl *a* concentration in the WEC and the Celtic Sea ranged between 0.4 – 6.0 mg m³ from 8th –
304 14th April (**Fig 4. A.**). Declining nitrate and silicate concentrations at station L4 from February
305 coincided with a chl *a* peak in early March, indicating the presence of an early diatom bloom. A
306 second chl *a* peak was evident during April, indicating that our sample timing coincided with the
307 successional phase from diatoms to the nanophytoplankton functional group / *Phaeocystis* spp.
308 (**Fig 4. B.**). A diverse diatom community dominated by *Cocinodiscus wailesii* was observed in
309 200 µm net trawl samples from station L4 from late February into March and *Phaeocystis* spp.
310 colonies were observed throughout April (data not shown).

311



312

313 **Fig 4. A.** MODIS weekly composite chl *a* image of the western English Channel covering the
 314 period 8th – 14th April (coincident with the week of phytoplankton community sampling for the
 315 present study), processing courtesy of NEODAAS. The position of coastal station L4 is marked
 316 with a white triangle. **B.** Profiles of weekly nutrient and Chl *a* measurements from station L4 at a
 317 depth of 10 m over the first half of 2015 in the months prior to experimental phytoplankton
 318 community sampling (indicated by black arrow and text).

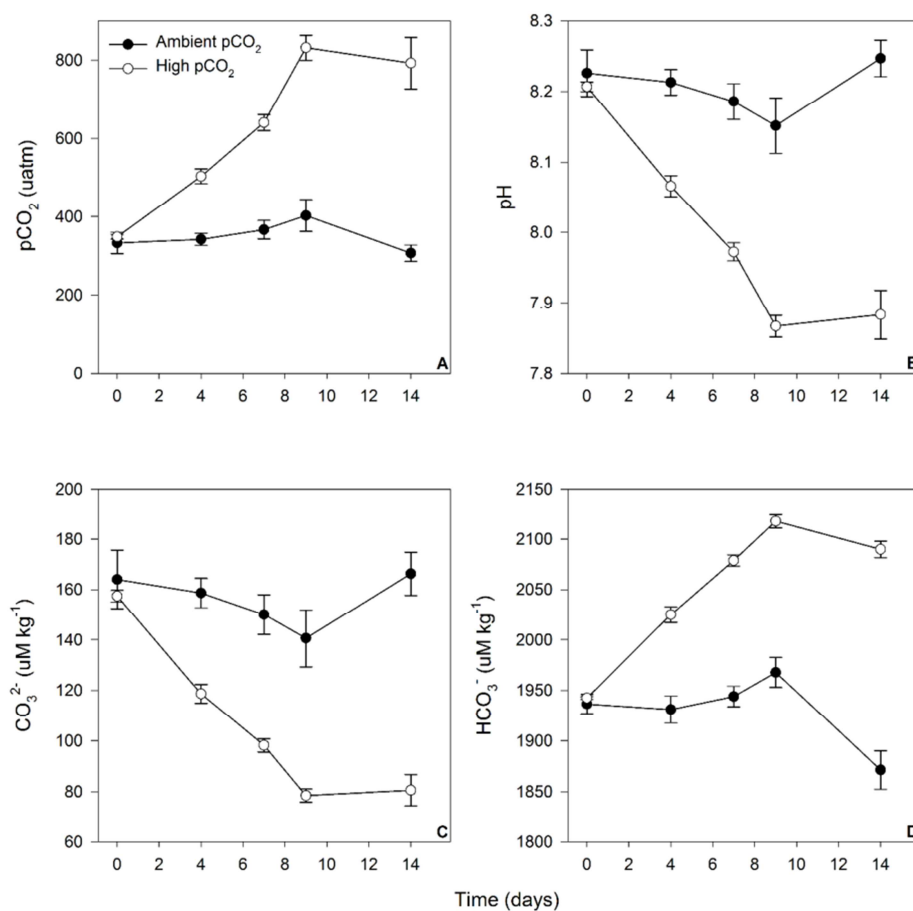
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320 3.2.1. Carbonate system

321 Mean pCO_2 values of media for the control and high CO_2 treatments were $361 (\pm 56 \text{ sd})$ and

322 $1006 (\pm 61 \text{ sd}) \mu\text{atm}$ respectively. Due to the dilution volume regime of the phytoplankton

323 community incubations, full equilibration to the target pCO₂ value (800 μatm) within the high
 324 CO₂ treatment incubations was achieved at T8. The high pCO₂ treatment incubations were
 325 slowly acclimated to rising pCO₂ over 8 days while the ambient control pCO₂ incubations were
 326 acclimated at the same ambient carbonate system values as that from station L4 on the day of
 327 sampling. Following equilibration, the mean pCO₂ values within the control and high CO₂
 328 incubations were 350 (± 95 sd) and 812 (± 39 sd) μatm respectively (**Fig 5. A-D**).



329

330 **Fig 5.** Carbonate system values of the experimental phytoplankton incubations. **A.** partial
 331 pressure of CO₂ in seawater (pCO₂), **B.** pH on the NBS scale, **C.** carbonate concentration (CO₃²⁻)
 332 and **D.** bicarbonate concentration (HCO₃⁻) were estimated from direct measurements of total
 333 alkalinity and dissolved inorganic carbon.

334 3.2.2. Chlorophyll *a*

335 Mean chl *a* concentration in the experimental seawater at T0 was 4.15 (± 0.38 sd) mg m³. The
 336 concentration dropped between T0 and T6 which in the control was 1.2 (± 0.27 sd) mg m³ at T6

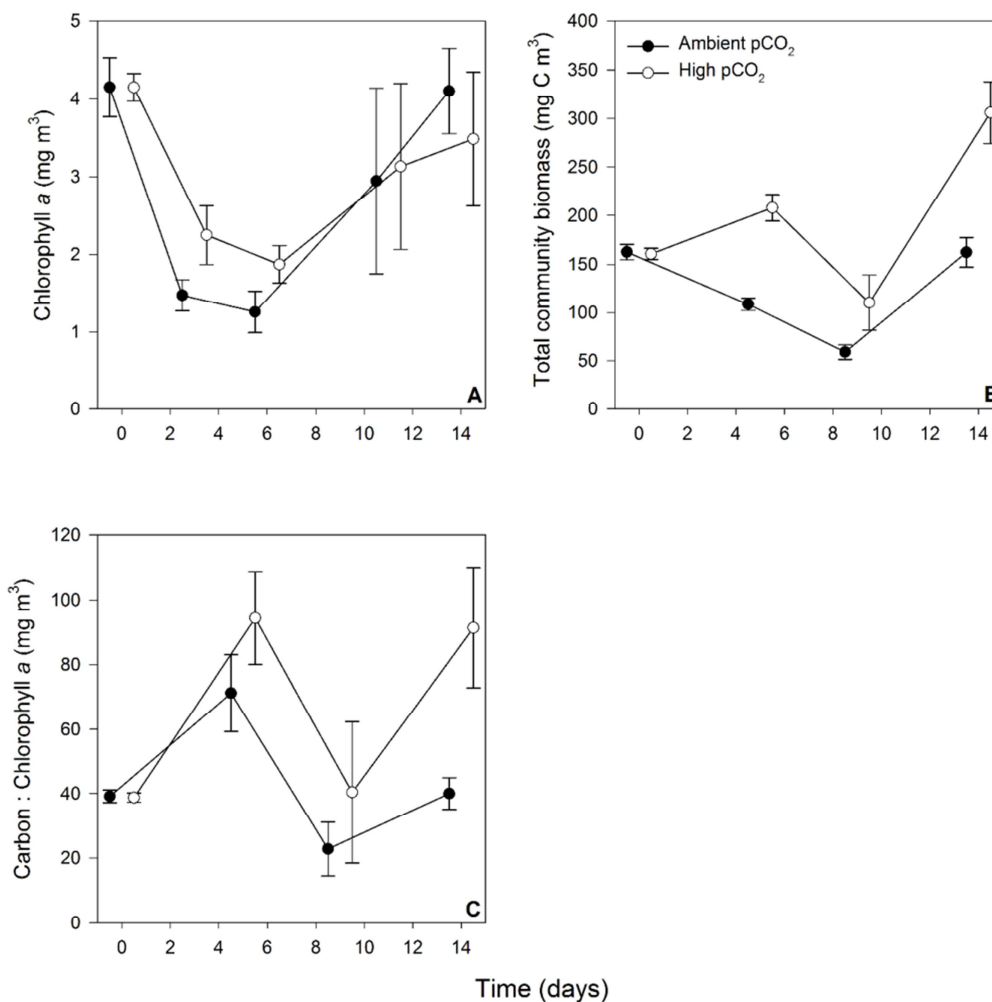
337 and in the high CO₂ treatment was 1.87 (\pm 0.24 sd) mg m³. Both control and high CO₂ treatments
338 showed a positive response to media dilutions from T6 onwards with no significant difference
339 in increased chl *a* concentration following target pCO₂ equilibration at T8. Final concentrations
340 were close to starting values, indicating that community net production was in balance with the
341 dilution rate. There was a significant increase in chl *a* concentration in both treatments over
342 time ($z = 2.437, p < 0.05$) (**Table 1.**). Mean chl *a* values on the final day of the experiment (T14)
343 were 4.1 (\pm 0.55 sd) and 3.5 (\pm 0.86 sd) mg m³ for the control and high CO₂ treatment
344 respectively (**Fig 6. A.**).

345 **3.2.3. Phytoplankton biomass**

346 The starting biomass was estimated at ~160 mg C m³ in both treatment groups. The community
347 was dominated by nanophytoplankton (excluding *Phaeocystis* spp. ~40%), *Phaeocystis* spp.
348 (~30%) and cryptophytes (~12%). Picophytoplankton contributed ~9% of total biomass while
349 the remaining 10% comprised diatoms, phytoflagellates, *Synechococcus*, ciliates,
350 coccolithophores and dinoflagellates in low abundance.

351 While total community biomass in both treatments declined to T9, the biomass in the high CO₂
352 treatment increased significantly from 110 to 305 mg C m³ between T9 and T14 ($z = 12.89, p <$
353 0.0001) (**Table 1., Fig 6. B.**) showing a 90% increase. The control community also increased
354 significantly between T9 and T14 and was restored to the initial starting value of 160 mg C m³
355 showing no overall net gain over the experimental period. Pairwise comparisons between the
356 treatments showed the high CO₂ treatment total biomass to be significantly greater than the
357 control at T14 ($t = 10.787, p < 0.001$). In both treatments there was a significant difference in
358 C:chl *a* (mg C m³:mg chl *a* m³) over time ($z = 6.684, p < 0.0001$). For the control, the ratio was
359 22.75 at T9 and 39.97 at T14, whereas for the high CO₂ treatment the ratio significantly
360 increased to 40.38 at T9 and 90.4 at T14, ($z = 6.778, p < 0.0001, \text{Table 1., Fig 6. C.}$).

361 *Phaeocystis* spp. decreased in the control community from T0 to T5 followed by a sharp increase
 362 at T9 to T14, from 17 to 91 mg C m⁻³. It dominated the control community, contributing 56% of
 363 overall biomass, more than any other group at T14 and increased by almost 90% compared to



364

365 **Fig 6. A.** Time series of chl *a*, **B.** total phytoplankton community biomass and **C.** carbon:chl *a*
 366 ratio. Following equilibration to experimental target pCO₂ (800 μatm), no significant response
 367 to elevated pCO₂ was observed in chl *a* between the ambient control and high CO₂ treatments,
 368 however both treatments showed a significant increase in chl *a* over time. A highly significant
 369 increase in total community biomass was observed in the elevated CO₂ treatment compared to
 370 that of the ambient control. Note: time points have been displaced to display standard deviation.

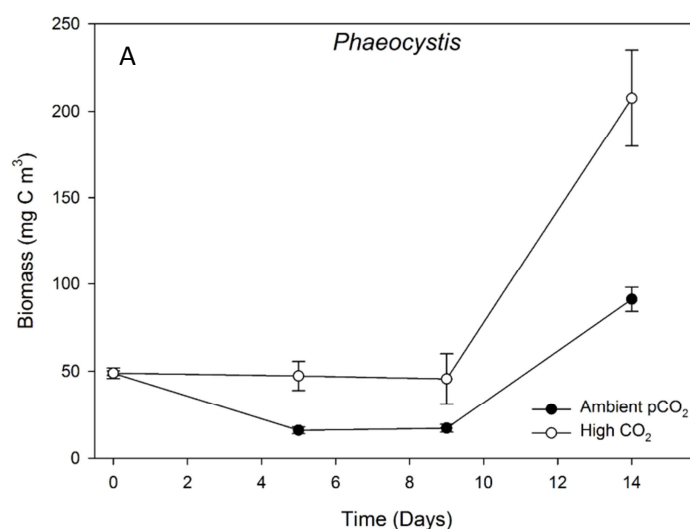
371

372 its initial T0 value. In the high CO₂ treatment however, there was a significant increase in

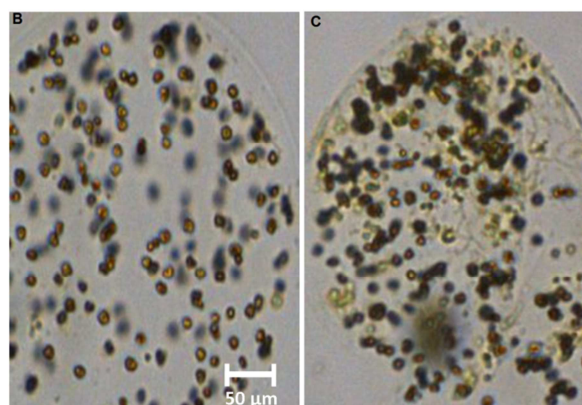
373 *Phaeocystis* spp. relative to the control community at T9 to T14, from 45 to 207 mg C m⁻³ (**Fig 7.**

374 **A.**), and it dominated the community within this treatment contributing ~70% of total

375 community biomass, increasing by 330% compared to the T0 starting biomass ($z = 3.219$, $p =$
 376 < 0.001) (**Table 2.**). Pairwise comparisons showed *Phaeocystis* spp. biomass in the high CO_2
 377 treatment to be significantly higher than the control at T5 ($t = 9.632$, $p < 0.001$), T9 ($t = 5.139$, p
 378 < 0.001) and T14 ($t = 10.811$, $p < 0.001$). Between T9 and T14, colonies of *Phaeocystis* spp. were
 379 observed in both treatments in the FlowCAM images, which provided a qualitative assessment
 380 of colony presence (**Fig 7. B. & C.**).



381



382

383 **Fig 7. A.** The highly significant response of *Phaeocystis* spp. to elevated pCO₂ in comparison to
 384 the ambient pCO₂ control. Note, biomass is the sum of solitary and colonial cells. **B. & C.** Image
 385 capture from FlowCAM analysis at T9 and T14 established the presence of *Phaeocystis* spp.
 386 colonies bound into a gelatinous matrix.

387 Nanophytoplankton biomass (excluding *Phaeocystis* spp.) declined in both treatment groups
 388 from T0 to T5, which was greater in the control community, from ~65 to 15 mg C m³ compared
 389 to ~65 to 23 mg C m³ in the high CO₂ treatment. Biomass increased significantly in both

390 treatments between T9 and T14 ($z = 4.141$, $p < 0.001$) (**Table 2., Fig 8. A.**).

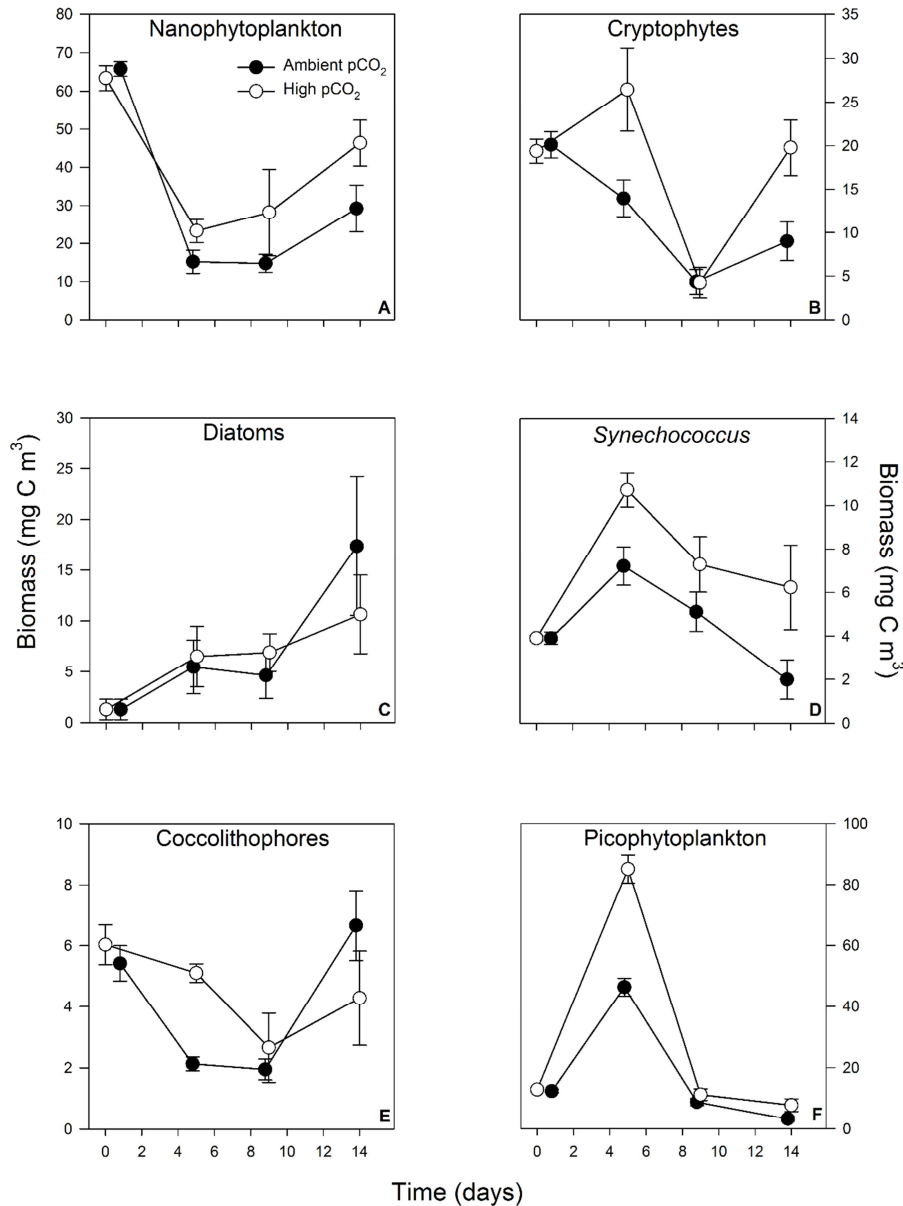
391 Nanophytoplankton showed an overall net loss in biomass at T14 in the high CO₂ treatment
392 (~46 mg C m³ at T14 compared to a starting biomass of ~63 mg m³ at T0, a decrease of 27%). A
393 pairwise comparison however, showed nanophytoplankton biomass to be significantly greater
394 in the high CO₂ treatment compared to the control at T14 ($t = 5.297$, $p < 0.001$).

395 Following an initial (acclimation) response of increased biomass between T0 and T5,
396 picophytoplankton and *Synechococcus* both declined over time. The high CO₂ treatment
397 maintained significantly higher biomass of picophytoplankton relative to the control following
398 target pCO₂ equilibration (T9 to T14, $t = 5.470$, $p < 0.001$) and significantly greater biomass of
399 *Synechococcus* which showed a net gain at T14 compared to its starting biomass value, an
400 increase of ~60% (3.9 to 6.2 mg C m³, **Table 2., Fig 8. D. & F.**, pairwise comparison, $t = 5.239$, p
401 < 0.001). An initial short term response of increased biomass was also observed with
402 cryptophytes in the high CO₂ treatment (T0 to T5), followed by a decrease in both control and
403 high CO₂ treatments (T5 to T9). Between T9 and T14 however, cryptophyte biomass increased
404 significantly in both treatments although to a greater extent in the high CO₂ treatment (pairwise
405 comparison, $t = 7.332$, $p < 0.001$) where it was restored to the starting value of ~20 mg C m³
406 compared to 9 mg C m³ in the control (**Table 2., Fig 8. B.**).

407 Dinoflagellate biomass was greater in the high CO₂ treatment compared to the control at T9,
408 however by T14 there was no significant difference between treatments (**Table 2.**). Flagellate
409 biomass (not including *Phaeocystis* spp.) remained low and exhibited a decline in both the
410 control and high CO₂ treatments, though the variability was high. With a mean of ~ 0.3 mg C m³
411 throughout the experiment, flagellates were the lowest biomass contributor and showed no
412 significant difference between treatments.

413 Ciliate biomass declined between T9 and T14 relative to the control. Both coccolithophore and
414 diatom biomass increased between T9 and T14 in both treatments, however the increase in the
415 high CO₂ treatment was lower compared to the control although not significantly for diatoms

416 (Table 2., Fig 8. C. & E.; pairwise comparisons for coccolithophores $t = -3.272, p < 0.02$; and
 417 diatoms $t = -2.266, p < 0.276$). Diatom community biomass was dominated by the chain forming
 418 *Chaetoceros curvisetus* and pennates *Proboscia alata* and *P. truncata* species. Smaller biomass



419

420 **Fig 8.** Response of individual components of the experimental phytoplankton community to
 421 elevated pCO₂. Cells were enumerated and converted to carbon biomass.
 422 contributions were made by *Chaetoceros socialis*, *C. decipiens*, *C. eibonii*, *Leptocylindrus danicus*,
 423 *Pseudonitzschia* spp. and *Thalassiosira* spp. At T9 *Proboscia* spp. contributed 66% and 62%
 424 (~1.5 and 2.2 mg C m⁻³) of total diatom biomass in the control and high CO₂ treatments while *C.*

425 *curvisetus* contributed 32% and 30% (~ 0.5 and 0.7 mg C m^{-3}) respectively (mean values). At T14
426 *Proboscia* spp. contributed 32% and 30% (~ 2.85 and 1.6 mg C m^{-3}) of total diatom biomass
427 respectively in the control and high CO_2 treatments, while *C. curvisetus* contributed 52% and
428 34% respectively (~ 4.6 and 1.8 mg C m^{-3}).

429

430 **4. Discussion**

431

432 **4.1. Trends in *Phaeocystis* spp. biomass from time-series analysis**

433 Previous analysis of the L4 phytoplankton time-series (1992-2007) elucidated distinct seasonal
434 and inter-annual changes in functional type composition as well as significant long term trends
435 in abundance. Over the study period, diatom abundance decreased while coccolithophorids, the
436 dinoflagellate *Prorocentrum cordatum* and some heterotrophic dinoflagellates and ciliates
437 increased in abundance (Widdicombe et al., 2010b). Analysis of 3 years of pCO_2 observations at
438 station L4 show an inter-annual trend of low pCO_2 in spring to high pCO_2 in autumn with a
439 concentration range of $\sim 250 \mu\text{atm}$ to $\sim 440 \mu\text{atm}$ (2005, 2007 and 2008). Metabolic processes
440 (i.e. photosynthesis during phytoplankton blooms), solubility and advection have been shown to
441 control seawater pCO_2 at station L4, with spring and summer showing the greatest atmospheric
442 CO_2 drawdown during stratified conditions, with CO_2 outgassing during the breakdown of
443 stratification during autumn (Litt et al., 2010).

444 Analysis of the *Phaeocystis* spp. biomass time-series at station L4 highlighted: 1) significant
445 inter-annual variability in biomass and 2) the occurrence of spring peak biomass between mid-
446 April to late May which contributed on average 17% of phytoplankton biomass between March
447 – May.

448 *Phaeocystis* spp. exhibit high temporal variability in the North Atlantic and North Sea which is
449 controlled by both meteorological and nutrient regimes (Gieskes et al., 2007).

450 On a global level, using 5057 observations from 1955-2009 Vogt *et al.*, (2013) showed that 64%
451 of *Phaeocystis* spp. biomass, was recorded during spring (northern hemisphere) with more
452 observations in the month of April compared to March and May, which we also observed in our
453 time-series analysis. Vogt *et al.*, (2013) also showed that the minimum and maximum
454 *Phaeocystis* spp. biomass was between 2.9×10^{-5} mg C m³ and 5.4×10^3 mg C m³, with a global
455 mean of 45.7 mg C m³ for both northern and southern hemispheres, which is similar to the
456 mean at station L4 (41.6 mg C m³). Inclusion of the colony mucus matrix carbon caused a
457 significant increase in the global mean to 183.8 mg C m³, highlighting the effect of colony mucus
458 on the carbon budget.

459 In the South-eastern North Sea there was a gradual decrease in *Phaeocystis* spp. abundance
460 from 1948 – 1970 based on Continuous Plankton Recorder survey (CPR) data (Gieskes and
461 Kraay, 1977) . Over a 12 year period (1973 – 1985), the spring maxima in *Phaeocystis* spp. and
462 bloom duration increased in the Marsdiep tidal inlet of the western Wadden Sea as a
463 consequence of eutrophication (Cadée and Hegeman, 1986). In the North East Atlantic, CPR data
464 also showed a decline in *Phaeocystis* spp. abundance from 1946 – 1987, whereas in the southern
465 North Sea between 1980 – 1987 *Phaeocystis* spp. abundance increased (Owens *et al.*, 1989).
466 More recently (1988 – 2001) in the North Sea Southern Bight Belgian Coastal Zone, variations
467 in *Phaeocystis* spp. abundance are reported to be regulated by winter nitrate enrichment
468 supplied by riverine pulses, which are controlled by local meteorological conditions associated
469 with the Winter North Atlantic Oscillation Index (NAO_w) (Breton *et al.*, 2006).

470

471 **4.2. Elevated pCO₂ Perturbation experiment**

472 From the microcosm experiments we found that elevated pCO₂ to ~800 µatm caused 1) a
473 significant increase in total community biomass by 90%, from ~160 to 305 mg C m³, 2)
474 significant changes in community structure from a nanophytoplankton (40%) and *Phaeocystis*
475 spp. (30%) dominated community to a *Phaeocystis* spp. (~70%) dominated community and 3)

476 either positive or negative responses in the rest of the phytoplankton community. Both diatoms
477 and coccolithophores were the only other phytoplankton groups to show a constant increase in
478 biomass in the control treatment which was greater than that of the high CO₂ treatment, but
479 only significant for coccolithophores.

480 The overall increase in total community biomass followed the same trend as previous studies
481 conducted on natural phytoplankton community CO₂ enrichments (Feng et al., 2009; Hare et al.,
482 2007b; Riebesell et al., 2007; Tortell et al., 2008b). The only groups/species to show an overall
483 net gain in biomass in the high CO₂ treatment, irrespective of significant increases (or
484 decreases) relative to the control community following target pCO₂ equilibration were diatoms,
485 *Phaeocystis* spp. and *Synechococcus*. A number of other studies have shown that the growth
486 rates of specific diatoms can increase by 5% to 33% following 20 generations acclimated at
487 elevated pCO₂ between 750 – 1000 µatm (*Phaeodactylum tricornutum*, *Thalassiosira*
488 *pseudonana*, *T. guillardii*, *T. weissflogii*, *T. punctigera* and *Cocinodiscus wailesii*) (Wu et al., 2014,
489 2010) and that the highest growth occurred in diatoms > 40 µm in diameter (*T. punctigera* and
490 *C. wailesii*). Similarly in some natural phytoplankton communities exposed to elevated pCO₂
491 (750 ppmv) diatoms and prymnesiophytes become dominant, making up 60% and 30% of the
492 total biomass (Tortell *et al.*, 2002). Tortell *et al.*, (2008a) also observed in Ross Sea
493 phytoplankton, a shift in dominance from *Phaeocystis antarctica* (contributing > 90%
494 community biomass) to large chain-forming diatoms (*Chaetoceros* spp.) within high CO₂
495 treatments (800 ppmv).

496 Solitary and colonial cells of *P. globosa* exposed to high CO₂ (750 ppm) have been shown to
497 exhibit a differential response. During 14 day incubations, solitary cell biomass decreased by
498 46%, while the number of colonies increased by 26%. Maximum growth rates of colonies
499 significantly increased by 30%, but no change in growth rate was observed in solitary cells.

500 Increased particulate organic carbon (POC), nitrogen (PON) and cellular C:N ratios were also

501 observed under elevated CO₂ (Wang et al., 2010). This suggests that elevated CO₂ may enhance
502 carbon export.

503 In a monoculture study using a *P. globosa* isolate from the South China Sea, Chen *et al.*, (2014)
504 recently showed that high CO₂ (1000 ppmv) caused a decrease in non-photochemical energy
505 loss. Short term exposure to elevated pCO₂ in combination with low, medium and high light
506 levels (25, 200 and 800 $\mu\text{mol photons m}^2 \text{s}^{-1}$) resulted in reduced growth rates under high light
507 conditions, however little effect was observed under low light conditions in short term
508 incubations. This is in agreement with a similar study using a North Sea *P. globosa* isolate where
509 Hoogstraten *et al.*, (2012) demonstrated decreased growth rates and photosynthetic efficiency
510 over a 6 day incubation period. However, following acclimation to experimental conditions after
511 9 and 14 generations, Chen *et al.*, (2014) observed enhanced growth rates, increased cellular chl
512 *a* and photosynthetic activity that had recovered to values equivalent of the control, which
513 contradicts the findings of Hoogstraten *et al.*, (2012). The authors concluded that effects of
514 elevated CO₂ on *P. globosa* are strongly influenced not just by irradiance, but also the stage of
515 acclimation to acidification. In our study, no decline in *Phaeocystis* spp. biomass was observed
516 over the first 8 days under high CO₂ exposure, but biomass remained constant, followed by a
517 significant increase to T14. Our irradiance was equal to that of the high light level applied by
518 Chen *et al.*, (2014). This highlights the importance of distinguishing between short-term
519 'cellular stress' related responses and acclimated responses, as well as the time period over
520 which experimental incubations are performed.

521 For *P. antarctica* in the Southern Ocean exposed to high CO₂ (current ambient, 600 and 800
522 ppmv), +2°C temperature increments (i.e. 2, 4 and 6°C respectively) and +50 $\mu\text{mol photons m}^2 \text{s}^{-1}$
523 irradiance increments (i.e. 50, 100 and 150 $\mu\text{mol photons m}^2 \text{s}^{-1}$) under both Fe replete and
524 limited conditions, there was a 64% decrease in growth rates at 800 ppmv (Fe replete) and a
525 46% decrease under Fe limiting conditions (Xu *et al.*, 2014). The Fe replete treatment increased
526 the percentage of solitary cells by 136% compared to 258% in the Fe limited treatment. Cellular

527 chl *a* decreased in the same treatments, but no influence was observed on cellular POC. The
528 experiment also assessed the competition between *P. antarctica* and the diatom *Fragilariopsis*
529 *cylindrus* and showed that the diatom dominated the population after day 8 at 800 ppmv CO₂.
530 We did not examine the effects of macro or micro nutrients, temperature or irradiance on our
531 *Phaeocystis* spp. community and all macro nutrients were replete. The findings of Xu *et al.*,
532 (2014) contrast what we observed and probably reflect differences between monoclonal and
533 two-species competition incubations and our experiment on a natural phytoplankton
534 community. Macro nutrients in our study were equal to maximal *in situ* values (L4 mean winter
535 values - 8µM nitrate+nitrite and 0.5 µM phosphate) providing favourable growth conditions for
536 *Phaeocystis* spp. in both the control and elevated CO₂ treatments.

537 Monoclonal incubations of *Phaeocystis* spp. under different CO₂ treatments can produce variable
538 responses related to species, strain or ecotype as well as due to differences in experimental
539 approaches. Incubations of a natural Arctic phytoplankton community under four pH
540 treatments (pH 8.0, 7.7, 7.4 and 7.1) showed that growth rates of *P. pouchetii* were unaffected by
541 pH from 8.0 – 7.4. There was however, a 50% decrease in growth rates at pH 7.1 (Thoisen *et al.*,
542 2015).

543 Accurate identification of *Phaeocystis* spp. relies on composite independent investigations
544 combining light microscopy, transmission and scanning electron microscopy as well as flow
545 cytometry for a complete identification of the morphotype (Rousseau *et al.*, 2007). Such an
546 investigation was beyond the scope and resources available for this study. Therefore, on the
547 basis of the FlowCam image capture, current knowledge of global geographical distributions of
548 the different *Phaeocystis* species (Verity *et al.*, 2007) and records from station L4
549 phytoplankton community time-series (Widdicombe *et al.*, 2010), it is likely the *Phaeocystis*
550 species observed in the microcosm exp. were a combination of *P. globosa* and *P. pouchetii*. The
551 FlowCam image capture was not used to enumerate *Phaeocystis* spp. colonies since the samples
552 were preserved with lugol's iodine which is known to cause colony disaggregation (Rutten *et al.*,

553 2005) which can cause an underestimation in group biomass. Flow cytometric analysis causes
554 cleavage of cell aggregations through the sheer force of sheath fluid (Dubelaar and van der
555 Reijden, 1995), and thus provides more accurate enumeration of single cells disaggregated from
556 colonies. This technique however, does not enable us to distinguish between colonial
557 cells/colonies and free living solitary cells.

558 In our experiment, the biomass range of *Phaeocystis* spp. in the control was within the range
559 measured throughout the L4 time series (~48 - 91 mg C m³ compared to in situ values of
560 between 33 - 137 mg C m³). The response of *Phaeocystis* spp. to the high CO₂ treatment
561 (increase of 330% from initial starting value to ~207 mg C m³) is above the maxima measured
562 at L4. Schoemann *et al.*, (2005) illustrated the difficulties in estimating *Phaeocystis* spp. biomass
563 due to the high carbon content of the polysaccharide matrix. The difference in carbon content
564 between a solitary (without the mucus matrix carbon contribution) and colonial cell has been
565 estimated to be 42.85 - 107.85 pg C cell⁻¹ based on empirical methods (Jahnke and Baumann,
566 1987). Since we fixed samples in Lugol's iodine, the C biomass could be underestimated. The
567 steady state incubation conditions in our exp., (irradiance, nutrients, temperature, dilution and
568 mixing regime) may have preferentially selected for *Phaeocystis* spp. compared to the
569 fluctuating conditions of the natural environment. Sommer, (1985) demonstrated that
570 variability in resource supply controls the number and relative proportion of coexisting species
571 and using monoclonal cultures subjected to carbonate system manipulation can result in
572 significantly different growth rates of the same species (e.g. Shi *et al.*, 2009).

573

574 **4.3. Implications**

575 Dense blooms of *Phaeocystis* spp. in some ecosystems can be responsible for fish and shell-fish
576 mortality and alteration of fish taste (Levasseur *et al.*, 1994; Peperzak & Poelman, 2008).
577 *Phaeocystis* spp. colony mucous matrix can inhibit copepod grazing, and therefore affect food
578 web structure through predator-prey size mis-match (Nejstgaard *et al.*, 2007). Several studies

579 have found consumption rates of *Phaeocystis* spp. by copepods to be significantly lower than
580 consumption of co-occurring diatoms and heterotrophic protists during *Phaeocystis* spp.
581 blooms, showing preferential feeding strategies towards more palatable and nutritious prey
582 sources (e.g. Gasparini *et al.*, 2000; Rousseau *et al.*, 2000; Verity, 2000). Additionally,
583 carbohydrates excreted by *Phaeocystis* spp. that coagulate to form transparent exopolymer
584 particles (TEP) have strong inhibitory feeding effects on both nauplii and adult copepods (Dutz
585 *et al.*, 2005). *Phaeocystis* spp. can also be inadequate as a food source for some copepods (e.g.
586 *Calanus helgolandicus*, *Temora stylifera* and *Acartia tonsa*), which can lead to negative effects on
587 fecundity and egg production (Tang *et al.*, 2001; Turner *et al.*, 2002). Stabell *et al.*, (1999)
588 extracted toxins from both *P. pouchetii* in culture and seawater samples collected during a
589 *Phaeocystis* spp. bloom, which have anaesthetic properties and can be toxic to fish larvae.
590 Exotoxins produced by *Phaeocystis* spp. during the spring bloom in the northern Norwegian
591 coast can also induce stress in cod larvae (*Gadus morhua*) (Eilertsen and Raa, 1995). Mass fish
592 mortalities have been linked to *Phaeocystis* spp. blooms in the Irish Sea (Rogers and Lockwood,
593 2009) and south-eastern Vietnamese coastal waters (Tang *et al.*, 2004). The mass transport and
594 sedimentation of a *P. globosa* bloom in 2001 in the Oosterschelde estuary (North Sea) caused
595 anoxic conditions that led to the mass mortality of 10 million kg of *Mytilus edulis* with a market
596 value of 15 to 20 million euro (Peperzak and Poelman, 2008). *Phaeocystis* spp. are also known
597 to produce and release the cytotoxic α , β , γ , δ -unsaturated aldehyde 2-trans-4-decadienal (DD),
598 which can inhibit mitotic cell divisions. An increase in DD concentration can have a negative
599 effect on the growth rates of the diatoms *Skeletonema costatum*, *Chaetoceros socialis* and
600 *Thalassiosira antarctica* (Hansen and Eilertsen, 2007). In addition, the odorous foam produced
601 by *Phaeocystis* spp. blooms can wash up on beaches and create anoxic conditions in the surface
602 sediment which can lead to mortality of the intertidal benthic community (Desroy and Denis,
603 2004; Spilmont *et al.*, 2009). These foam deposits also have a deleterious effect on coastal
604 tourism (Lancelot and Mathot, 1987). Our microcosm experiments suggest that future high CO₂

605 scenarios could increase *Phaeocystis* spp. blooms at station L4 in the WEC which could
606 adversely affect ecosystem functioning, food web structure, fisheries and tourism.

607 **5. Conclusion**

608 Microcosm experiments showed that *Phaeocystis* spp. carbon biomass increased by 330% at
609 elevated pCO₂ (~800 µatm) over a 15-day period. This study suggests that future high pCO₂
610 concentrations in the WEC may favour the dominance of *Phaeocystis* spp. biomass during the
611 spring bloom, with associated negative impacts on ecosystem function and food web structure.

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Table 1. Results of generalized linear models testing for the effects of pCO₂ and time on measured phytoplankton community parameters. Significance results are given as:
* p < 0.05 and *** P < 0.0001.

<u>Response variable</u>	<u>n</u>	<u>df</u>	<u>z-value</u>	<u>p</u>	<u>sig</u>
Chl <i>a</i> (mg m³)					
pCO ₂	32	29	1.032	0.52061	
Time	32	29	2.437	< 0.05	*
Carbon:Chl <i>a</i> (mg m³)					
pCO ₂	32	29	6.778	<0.0001	***
Time	32	29	6.684	<0.0001	***
Total community biomass (mg C m³)					
pCO ₂	32	29	12.890	<0.0001	***
Time	32	29	20.48	<0.0001	***

Table 2. Results of generalized linear models testing for the effects of pCO₂ and time on

individual phytoplankton species biomass, (n = 32). Significance results are given as:
 * p < 0.05, and *** P < 0.0001.

<u>Response variable</u>	<u>Parameter</u>	<u>df</u>	<u>z-value</u>	<u>p</u>	<u>sig</u>
Ciliate biomass (mg C m ³)	pCO ₂	29	2.532	0.17	
	Time	29	0.622	0.539	
Coccolithophore biomass (mg C m ³)	pCO ₂	29	-1.723	0.095	
	Time	29	5.763	<0.0001	***
Cryptophyte biomass (mg C m ³)	pCO ₂	29	-2.060	0.039	*
	Time	29	3.513	<0.0001	***
Diatom biomass (mg C m ³)	pCO ₂	29	-1.068	0.286	
	Time	29	5.648	<0.001	***
Dinoflagellate biomass (mg C m ³)	pCO ₂	29	1.567	0.117	
	Time	29	1.205	0.228	
Flagellate biomass (mg C m ³)	pCO ₂	29	-0.342	0.732	
	Time	29	0.345	0.73	
Nanophytoplankton biomass (mg C m ³)	pCO ₂	29	2.36	0.018	*
	Time	29	3.697	<0.0001	***
<i>Phaeocystis</i> biomass (mg C m ³)	pCO ₂	29	3.707	<0.0001	***
	Time	29	15.636	<0.0001	***
Picophytoplankton biomass (mg C m ³)	pCO ₂	29	-1.448	0.148	
	Time	29	-4.331	<0.0001	***
<i>Synechococcus</i> biomass (mg C m ³)	pCO ₂	29	-2.334	0.027	*
	Time	29	-5.407	<0.0001	***

627

628

Table 3. Inter-annual differences in *Phaeocystis* spp. carbon biomass at station L4 between 1993 – 2014 tested with a generalised linear model. Significance results are given as:
* $p < 0.05$, ** $p < 0.01$ and *** $p < 0.001$ (n = 21).

<u>Response variable (Year)</u>	<u>df</u>	<u>z value</u>	<u>p</u>	<u>sig</u>
1993	20	1.458	0.144737	
1994	20	2.055	0.039883	*
1996	20	2.1	0.035769	*
1997	20	2.567	0.010258	*
1998	20	2.343	0.01915	*
1999	20	2.522	0.011669	*
2000	20	-1.009	0.313129	
2001	20	3.355	0.000794	***
2002	20	2.473	0.013411	*
2003	20	1.153	0.249038	
2004	20	2.649	0.008075	**
2005	20	-1.48	0.138741	
2006	20	-0.806	0.420182	
2007	20	1.764	0.077781	
2008	20	-0.847	0.39682	
2009	20	-0.86	0.389549	
2010	20	1.74	0.081802	
2011	20	-2.298	0.021548	*
2012	20	0.2	0.841177	
2013	20	2.34	0.019302	*
2014	20	2.405	0.016167	*

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