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High-frequency Observations of the Isotopic Composition of Soil, Stem and Root Respiration in a Danish Beech Forest

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Introduction

Recent advances in laser technology have allowed for real-time

Methods and Objectives

The study had the following objective:

measurements of the isotopic composition of C and O in CO₂ thereby providing new ways for partitioning ecosystem CO₂ fluxes into contributions of single processes. In this study, we present isotopic concentration data (i.e. δ^{13} C) and CO₂ fluxes measured at high temporal resolution from intact soil, trenched soil, tree stems, tree roots and intact beech leaves of a Danish beech forest ecosystem by an Aerodyne Single CW-Quantum Cascade Laser Trace Gas Analyzer for CO₂ Isotopes.



Figure 1 (left): One of eight automated soil respiration chambers (8100-104 Long-Term Chamber, LI-COR Biosciences). Figure 2 (right): The Aerodyne Single CW-Quantum Cascade 🚺 Laser Trace Gas Analyzer for CO₂ Isotopes (Aerodyne Research Inc.).



•Combine the Aerodyne Single CW-Quantum Cascade Laser Trace Gas Analyzer for CO₂ Isotopes (Figure 1) with the LI-COR LI-8100/8150 flux chamber system to yield an isotopic signal of the CO₂ fluxes from various compartments of a forest ecosystem at a high temporal scale.

The Aerodyne laser was combined in parallel with the LI-COR LI-8100/8150 system that controlled 13 automated, closed chambers, thus providing δ^{13} C of the measured CO₂ fluxes. The chamber fluxes were each measured automatically every 2 hours. Eight of these chambers were LI-COR long term soil flux chambers (Figure 1), with four measuring fluxes from intact soil plots and 4 measuring fluxes from trenched soil plots. The remaining were custom made chambers with two measuring fluxes from tree stems, two measuring fluxes from coarse roots and one measuring fluxes from a tree branch with leaves. By applying the Keeling plot methodology to the change in $\delta^{13}C$ of CO_2 during the 5 minutes of chamber closure, the $\delta^{13}C$ of the CO_2 fluxes was determined.

Results

The system has been running since January 2016. During the cold months of winter and spring, where fluxes were low, the change in CO₂ concentration during chamber closure time was found to be too small for proper determination of δ^{13} C, resulting in a low R² value of the keeling plot fit. However for the high fluxes during June and July a sufficient change in δ^{13} C during chamber closure time was measured resulting in a high R² value of the determined δ^{13} C. An example of the bihourly δ^{13} C of soil CO₂ efflux can be seen in figure 3.

Example of bihourly δ^{13} C of soil CO₂ efflux



The average δ^{13} C of the CO₂ fluxes were calculated for each chamber for June and July (Figure 4). δ^{13} C from soil fluxes was found to be more enriched than δ^{13} C from root respiration, which has also been found in previous studies (e.g. Millard et al. 2010). No clear difference was however found between intact soils and trenched soils.

Figure 3: Example of 10 days of soil CO₂ effluxes for one trenched soil plot measured by a LI-COR soil respiration chamber (In blue), and the δ^{13} C of the respired CO₂ (In red) determined by a Keeling plot. A measurement was performed every two hours. $\delta^{13}C$ with R² values less than 0.90 from the Keeling plot have been removed. Error bars show standard deviation.

δ^{13} C of ecosystem CO₂ fluxes



Conclusions and Outlook

The Aerodyne laser and the LI-COR LI-8100/8150 system was successfully combined to yield the δ^{13} C of various ecosystem CO₂ fluxes. However, the determined δ^{13} C of the respired CO₂ showed a high variation.

To increase the precision of the determined $\delta^{13}C$ we will:

•Increase chamber closure time, to increase the change in $\delta^{13}C$ during a chamber measurement.

•Try to decrease drift in the δ^{13} C signal measured by the Aerodyne laser.

Acknowledgements

Figure 4: The determined δ^{13} C of the respired CO₂ for each individual flux chamber. Each data point is based on all measurements during June and July (only 10 days in June for the leaf chamber during daytime) for one flux chamber, after removal of δ^{13} C values with an R² value less than 0.90 from the Keeling plot. Error bars show standard deviation.

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Reference: Millard, P., Midwood, A.J., Hunt, J.E., Barbour, M.M., Whitehead, D., 2010. Quantifying the contribution of soil organic matter turnover to forest soil respiration, using natural abundance δ¹³C. Soil Biol. Biochem. 42, 935–943. doi:10.1016/j.soilbio.2010.02.010

