

This is a repository copy of Distributions of salt-marsh foraminifera along the coast of SW California, USA: Implications for sea-level reconstructions.

White Rose Research Online URL for this paper: https://eprints.whiterose.ac.uk/115666/

Version: Accepted Version

Article:

Avnaim-Katav, Simona, Gehrels, Willem Roland, Brown, Lauren et al. (2 more authors) (2017) Distributions of salt-marsh foraminifera along the coast of SW California, USA: Implications for sea-level reconstructions. Marine Micropaleontology. pp. 25-43. ISSN 0377-8398

https://doi.org/10.1016/j.marmicro.2017.02.001

Reuse

This article is distributed under the terms of the Creative Commons Attribution-NonCommercial-NoDerivs (CC BY-NC-ND) licence. This licence only allows you to download this work and share it with others as long as you credit the authors, but you can't change the article in any way or use it commercially. More information and the full terms of the licence here: https://creativecommons.org/licenses/

Takedown

If you consider content in White Rose Research Online to be in breach of UK law, please notify us by emailing eprints@whiterose.ac.uk including the URL of the record and the reason for the withdrawal request.



1 Distributions of salt-marsh foraminifera along the coast of SW California, USA:

implications for sea-level reconstructions

2

Simona Avnaim-Katav^{a,b*}, W. Roland Gehrels^c, Lauren N. Brown^b, Elizabeth Fard^b,
 Glen M. MacDonald ^{a,b}

6 7

8

^a University of California, Los Angeles, Institute of the Environment and

Sustainability, La Kretz Hall, Suite 300, Box 951496, Los Angeles, CA 90095-1496,
 USA.

b University of California, Los Angeles, Department of Geography, 1255 Bunche

12 Hall, Box 951524, Los Angeles, CA 90095, USA.

^c Environment Department, University of York, Heslington, York, YO10 5NG,

14 United Kingdom.

15 16

*Corresponding author: Telephone: +909-734-9661, Email: Simona100@ucla.edu

17 18 19

20

21

Reference: Avnaim-Katav, S., Gehrels, W.R., Brown, L.N., Fard, E., MacDonald, G.M., 2017. Distributions of salt-marsh foraminifera along the coast of SW California, USA: implications for sea-level reconstructions. Marine Micropaleontology 131, 25-43, doi:10.1016/j.marmicro.2017.02.001.

222324

Abstract

25 26

27

28

29 30

31

Salt-marsh foraminifera are frequently used around the world as proxies in paleoenvironmental studies of sea-level change. Quantitative reconstructions of sea-level change use transfer functions which are based on the vertical zonation of salt-marsh foraminifera with respect to the tidal frame. This paper explores for the first time the environmental factors that control the foraminiferal assemblages in Southern California marshes using modern surface samples (1 cm thick) from two marshes Seal

Beach and Tijuana Estuary. The dead foraminiferal assemblages demonstrate distinct zonation across the salt-marsh surfaces which is primarily related to elevation. Other

variables less important than elevation such as O₂, temperature, salinity and pH

35 additionally control the distribution pattern of these assemblages.

The tidal flat and low marshes are characterized by high abundances of *Miliammina*

37 fusca and calcareous species. The middle marsh is dominated by Jadammina

38 macrescens and Trochammina inflata, while the high marsh zone is dominated

39 by $Trochamminita\ irregularis$, $Miliammina\ petila$, $J.\ macrescens\ and\ T.$

40 *inflata*. Regression modelling was used for the development of a sea-level transfer

function based on a combined training set of surface samples from the two study sites.

The performance of the Weighted Average – Partial Least Squares (WA-PLS) transfer function suggests a robust relationship between the observed and estimated elevations

44 $(r^2_{\text{Jack}} = 0.72)$, and is capable of predicting former sea levels to a precision of ± 0.09

m. Our results can be used for future paleoenvironmental reconstructions along

the Southern California coast, an area that has experienced changes in sea level in the

past and will be affected by future sea-level rise coupled with climate and

anthropogenic changes, resulting in wide impacts on the natural coastal habitats in this region.

Keywords: foraminiferal assemblages; California; salt marshes; environmental relations; redundancy analysis; transfer function; relative sea level;

1 Introduction

Salt-marsh sediment records are used to quantitatively reconstruct late Holocene relative sea-level changes (e.g., Kemp et al., 2011), with fairly high precision (up to ±0.05 m, cf. Gehrels and Woodworth, 2013). In combination with long-term tide gauge records they have provided sea-level reconstructions of the transition from relatively low rates of change during the late Holocene in the order of tenths of mm yr⁻¹ to accelerated modern rates (early 20th century) in the order of mm yr⁻¹, both in the Northern and Southern Hemispheres (Gehrels and Woodworth, 2013). Quantifying regional patterns of sea-level change using proxies from salt-marsh sediments provides information on drivers of local to regional sea-level change and ultimately on global climate change (e.g., mass-balance changes of ice sheets and glaciers) (e.g., Milne et al., 2002; Milne et al., 2009; Kopp et al., 2016). Benthic foraminifera preserved in salt-marsh sediments are frequently used for reconstructing past sea-level changes due to their vertical zonation in the modern intertidal zone (e.g., Scott and Medioli, 1980; Scott et al., 1984; Gehrels, 1994; Horton et al., 1999a; Kemp et al., 2009a; Horton and Edwards, 2005), especially in micro- to mesotidal settings (e.g., Scott et al., 2001; Barlow et al., 2013 and references therein; Kemp and Telford, 2015 and references therein). Elevation relative to the tidal frame (or frequency of tidal submergence) is the controlling environmental parameter that is of interest to sea-level studies (Gehrels, 2000). Studies exploring the ecology of modern salt-marsh foraminifera show that their spatial distribution can also be controlled by other environmental factors, such as salinity (e.g., Murray, 1971; Patterson, 1990; de Rijk and Troelstra, 1997), pH (e.g., Woodroffe et al., 2005; Barnett et al., 2016), grain-size (e.g., Matera and Lee, 1972; de Rijk and Troelstra, 1997; Scott et al., 1998; Barnett et al., 2016) and organic carbon concentration (e.g., de Rijk and Troelstra, 1997; Milker et al., 2015a). However, these environmental parameters and others (e.g., vegetation cover, geochemical properties) are significantly correlated with tidal elevation in many salt marshes (e.g., de Rijk and Troelstra, 1997; Horton, 1999a; Horton et al., 2003; Horton and Edwards, 2005; Milker et al., 2015a).

Many quantitative sea-level studies use regression methods to develop predictive transfer functions, capable of inferring past relative sea levels from fossil foraminifera in subsurface salt-marsh deposits (e.g., Gehrels, 1999; Horton et al., 1999b; Edwards and Horton, 2000; Gehrels, 2000; Edwards et al., 2004b; Gehrels et al., 2006; Horton and Edwards, 2006; Kemp et al., 2009b; Gehrels et al., 2012). In marshes where salt-marsh foraminifera are unimodally distributed along the environmental gradient, Weighted Averaging (WA), Locally Weighted Weighted-Averaging (LW-WA) and even to a higher degree, Weighted Averaging Partial Least-Squares (WA-PLS), are considered as the most robust models for reconstructing sea level (ter Braak and Juggins, 1993; ter Braak et al., 1993; Birks, 1995, 2010; Juggins and Birks, 2012; Kemp and Telford, 2015). Where linear species—environment responses are demonstrated, Partial Least Squares (PLS) is the most widely used technique (e.g., Stone and Brooks, 1990; Rossi et al., 2011). These methods are reviewed in detail by

100 Birks (1995, 2003, 2010), Barlow et al. (2013) and Kemp and Telford (2015), and

101 have been applied to reconstruct past sea-level changes in a wide range of geographical areas, including the Atlantic coast of the USA (e.g., Gehrels, 2000; 102 Edwards et al., 2004a,b; Gehrels et al., 2002, 2004, 2005; Horton et al., 2006; Kemp 103 et al., 2009a, 2009b, 2011, 2013, 2015; Wright et al., 2011), the Atlantic coast of 104 Europe (e.g., Gehrels et al., 2001; Horton and Edwards, 2005, 2006; Leorri et al., 105 2010; Long et al. 2014; Barlow et al., 2014; Barnett et al., 2015), eastern Canada (e.g., 106 107 Patterson et al., 2004; Gehrels et al., 2005; Barnett et al., 2016), west coast of Canada (Guilbault et al., 1996), South Africa (Franceschini et al., 2005; Strachan et al., 2014), 108 Iceland (Gehrels et al., 2006; Saher et al., 2015), Australia (Woodroffe, 2009; Gehrels 109 et al., 2012), New Zealand (Gehrels et al., 2008; Grenfell et al., 2012) and Malaysia 110 (Culver et al., 2015). Despite this vast literature, few studies have been conducted 111 along the Pacific coast of the USA. Exceptions are the studies of intertidal 112 foraminifera from Oregon, USA, to reconstruct relative sea-level changes caused by 113 co-seismic subsidence (e.g., Hawkes et al., 2010, 2011; Engelhart et al., 2013) and to 114 investigate the influence of taphonomic processes, small-scale variability and infaunal 115 distribution on the accuracy of sea-level reconstructions (Milker et al., 2015a). 116 117 However, neither a comprehensive modern training set, nor transfer function based on recent developments and methodological recommendations (Barlow et al., 2013; 118 Kemp and Telford, 2015) have been developed for salt marshes in California. With 119 that being said, previous studies, during the mid-late 1970's, from Tijuana salt-marsh, 120 Southern California, qualitatively divided the marsh into foraminiferal zones (Scott, 121 1976; Scott and Medioli, 1978) which were applied in Holocene paleoenvironmental 122 123 reconstructions (Scott et al., 2011). 124

Salt-marsh ecosystems are of high economic value and societal and ecological significance. They provide ecosystem services for tourism and commercial fisheries, as nurseries and refuge areas for a variety of organisms, for carbon sequestration, and for protecting water quality by filtering runoff and reducing the pollutant load entering estuaries. Moreover, salt marshes are a natural protection against coastal flooding, erosion and sea-level rise (e.g., Gedan et al., 2010). For the Southern California coast, with its dense population and ongoing disappearance of natural coastal habitats due to anthropogenic modification, variable tectonic regimes and climate change, the need to investigate the coastal marshes is pressing. In order to create a reference baseline for future regional paleogeographic, paleoecological and paleo sea-level study this study aims to: (1) identify and describe the distribution patterns of the living and dead foraminiferal assemblages in two coastal salt marshes: Tijuana and Seal Beach; (2) constrain by quantitative ordination methods the relationships between foraminiferal assemblages and the most important environmental drivers influencing their distribution; (3) examine whether the recent foraminiferal assemblages in these salt marshes show a distinct zonation relative to the tidal frame; (4) develop an ecological response function using the modern training set (i.e., a foraminifera based transfer function) suitable for precise relative sea-level reconstructions in an area that is susceptible to both sea-level change and tectonic land movements.

144145146

125126

127

128

129

130

131

132

133

134

135136

137

138139

140

141

142

143

2 Study area

150 The salt marsh at Tijuana is part of the Tijuana River Estuary located north of the US-

- Mexico border (32° 33' N, 117° 07' W; Fig. 1). The salt marsh is part of the National
- 152 Estuarine Research Reserve which has over 1000 ha of coastal marsh environments
- 153 (Zedler et al., 1986). Just over 150 km to the northwest of Tijuana, Sea Beach consists
- of approximately 390 ha of salt marshes incorporated in a National Wildlife Refuge
- that lies within the US Naval Weapons Station at Seal Beach (33° 44' N, 118° 05' W).
- 156 The Seal Beach and Tijuana sites represent salt marshes with generally similar tidal
- characteristics and overall climatic and oceanographic settings representative of
- Southern California. Tides are semidiurnal and have an observed mean range close to
- the upper micro-tidal range, <2 m (Table 1), as measured from 1983 to 2001 at the
- National Oceanic and Atmospheric Administration (NOAA) tide-gauge stations in
- Los Angeles, ~16 km north of Seal Beach, and Imperial Beach, ~2 km north of
- 162 Tijuana.

163164

- Historically, salt marshes on the Southern California coast have developed in response
- to the drowning of river valleys when sea levels reached their approximate current
- position about 5000 years before present (Zedler, 1982). Although salt marshes in
- 167 California have seen a great reduction in area due to human influence, Tijuana salt
- marsh is considered a pristine ecosystem that preserves many of its natural qualities.
- Moreover the northern arm of the marsh, which was chosen for this study, is known
- for its simple elevation gradient and its relative physical stability (Zedler, 1977). The
- marsh at Seal Beach is a similar habitat that has been least affected by human
- disturbance thanks to its designation as a wildlife reserve. Moreover, this marsh is the
- only remaining undeveloped part of the Anaheim Bay estuary. Prior to development,
- this estuary had large fringing freshwater wetlands, salt flats, and alkali meadows, and
- freshwater creeks may have flowed perennially into it (Grossinger et al., 2011).
- 176 Reclaimed areas of the estuary adjacent to the Seal Beach salt marsh include military,
- municipal and industrial infrastructure.

178

- Natural hydrologic disturbances at Tijuana include extreme flood events which can
- deposit up to 8 or 9 cm of sediment in low marsh areas near channels (Cahoon et al.,
- 181 1996; Wallace et al., 2005). In the past century, some reduction in streamflow has
- been observed due to damming upstream in the watershed. Even so, hydrologic
- extremes of flood and drought at Tijuana have been known to cause periods of tidal
- closure and hypersalinity leading to vegetation die-off (Zedler, 2010).

185

- The most significant natural hydrologic change at the salt marsh at Seal Beach
- occurred about 300 years before present with the isolation of the marsh due to
- changes in the flow of the Santa Ana River (Leeper III, 2015). Historically
- sedimentation at Seal Beach appears to have kept pace with sea-level rise in the recent
- past, but the lack of terrestrial sediment input makes the marsh at Seal Beach all the
- more vulnerable to accelerated sea-level rise (Kirwan and Megonigal, 2013). This
- danger is compounded by the fact that Seal Beach lies directly along the San Andreas
- 193 Fault and is suspected to have suffered elevation loss due to tectonic subsidence
- 194 (Leeper III, 2015). Additionally, the region surrounding Seal Beach has observed a
- 4.13 mm/yr average lowering in surface elevation from 1994 to 2012 (Takekawa et
- al., 2013) due to oil extraction.

- Marsh vegetation in Southern California is often classified into zones based on
- elevation (Peinado et al., 1994; Zedler et al., 1986; Zedler, 1977). Zonation is due to a

complex combination of biotic tolerances and interspecific competition (Engels et al., 2011). The low elevations, high salinity habitats of Southern California salt marshes are dominated by Spartina alterniflora. Mid-marsh regions are frequently co-dominated by a number of species including Sarcocornia pacifica, Batis maritima, and Jaumea carnosa. Species such as Distichlis spicata, Frankenia grandifolia and Limonium californicum are found in the mid-to-high marsh areas. However, the boundary between mid and high marsh vegetation is less definite, and plants from these zones also colonize the marsh-upland transition area. Arthrocnemum subterminale and Monanthochloe littoralis are species which show a preference for the highest elevations in the marsh. Marsh-upland transition can most clearly be seen by the appearance of shrub-type plants, often those of the native chaparral vegetation such as *Artemisia californica*, *Rhus lauriana*, and *Baccharis pilularis*.

The climate of Southern California is Mediterranean and experiences hot, dry summers and warm, wet winters. At Tijuana Estuary, the average annual high temperature is 21°C with an average low of 12.9°C and an average annual precipitation of 26.6 cm. At Seal Beach, average annual high temperatures are 23.4°C and lows are 13°C with an annual average of 31.1 cm of precipitation (usclimatedata.com). Drought and extreme flooding, as previously mentioned, are not uncommon and pose the most significant climatic impacts on marsh vegetation (Zedler et al., 1986; Zedler, 2010).

3 Materials and methods

3.1. Field sampling

We established twenty stations in Seal Beach salt marsh and seventeen stations in Tijuana salt marsh (Fig. 1, Table 2) along transects perpendicular to the primary direction of tidal inundation. Stations covered the sub-environments within the intertidal zone from high marsh to the tidal flat, where possible. Sampling coincided with approximately spring tide in the fall in order to standardize the timing of collection and to maintain perennial persistency in foraminiferal distribution (i.e., we avoided sampling during bloom periods) (e.g., Buzas, 1965; Jones and Ross, 1979; Horton and Edwards, 2003; Horton and Murray, 2006, 2007; Schönfeld et al., 2012). The stations at Seal Beach were sampled in mid-October 2015 and resampled (only for faunal analysis) three days after storm activity generated by Hurricane Patricia on 23 October, 2016 (www.nhc.noaa.gov). However, our results showed that the study area was not impacted by the storm and we used the replicate dead foraminiferal data at Seal Beach to enhance our interpretations. The stations at Tijuana were sampled in early December, 2015.

Station locations were determined with Differential Global Positioning System (DGPS) and post processed with the Online Positioning User Service (OPUS) generated by NOAA in order to provide simplified access to high-accuracy National Spatial Reference System (NSRS) coordinates i.e., to standardize and correct for spatial changes in MSL. The elevations, with a precision of ≤4cm (3cm in average), were referenced to the North American vertical datum (NAVD88) computed using GEOID12B.

248 At each station vegetation cover was described and pore-water salinity, temperature, O₂ concentration and pH were measured using hand-held EXTECH DO700 meter 249 following de Rijk (1995).

250 251

Fifty seven surface marsh sediment samples with a standardized volume of 10 cm³ 252 from the uppermost centimeter (10 cm² by 1 cm thick) were sampled for foraminiferal 253 analysis. This sampling strategy follows the assumption that the intertidal 254 foraminifera are primarily shallow infaunal (classification follows Buzas et al. 255 (1993)), thus our surface samples sufficiently represent the modern intertidal 256 257 environment. This hypothesis is supported by research in other salt marshes, for example in Oregon (Hawkes et al., 2010; Milker et al., 2015a), Nova Scotia (Scott and 258 Medioli, 1980), Maine (Gehrels, 1994), Massachusetts (de Rijk, 1995), North Carolina 259 (Culver and Horton, 2005), New Brunswick, Canada (Patterson et al., 2004) and the 260 UK (Horton, 1997; Horton and Edwards, 2006). Nevertheless, there are a considerable 261 number of studies that reported living infaunal salt-marsh foraminifera which might 262 alter the composition of fossil assemblages (e.g., Hippensteel et al., 2002; Culver and 263 Horton, 2005 and references therein) due to seasonal and local environmental 264 conditions and bioturbation (Buzas et al., 1993; de Stigter et al., 1998; Culver and 265 Horton, 2005; Tobin et al., 2005; Horton and Edwards, 2006). We tested the influence 266 267 of infaunal foraminifera in the marshes of Southern California by analyzing foraminifera at 1 cm resolution down 10cm long mini cores from the mid-marsh (one 268 core along each transect).

269 270 271

272 273

274

275

276

277

278

279

280

All foraminifera surface and core samples were preserved on the day of sampling in rose Bengal solution (2 g rose Bengal /1 95%- ethanol) for two weeks to distinguish living from dead specimens at the time of collection (Walton, 1952; Murray and Bowser, 2000; Schönfeld et al., 2012). Specimens were considered living at the time of collection when all chambers were stained completely bright red, except for the last chamber (e.g., Horton and Edwards, 2006; Milker et al., 2015a). Additionally the solution was buffered by calcium carbonate powder to prevent dissolution of calcareous tests. At each sampling station a second set of sediment samples (one per station) was taken and stored in sealed plastic bags for grain-size distribution and loss on ignition (LOI) analyses. All samples were stored in the refrigerator at 4°C prior to further laboratory analyses.

281 282 283

284

3.2. Laboratory techniques **Quantitative Sedimentological Analyses**

285 286 287

288

289

290

297

Grain-size distributions of the surface samples were determined using a Malvern Mastersizer 2000 Laser Diffraction Particle Size Analyzer coupled to a Hydro 2000G large-volume sample dispersion unit at the Paleoclimatology and Paleotsunami Laboratory, California State University, Fullerton. Prior to analysis, samples were treated with 25-50 ml of 30% hydrogen peroxide (H₂O₂) to remove organic matter.

291 292 Standard laboratory procedures for grain-size distributions are detailed in Kirby et al.

293 (2015).

Loss-on-Ignition (LOI; a proxy for percent organic matter- %TOM) at 550 °C and 294 295 950 °C (% total carbonate - %TC) were measured according to the method of Heiri et

al. (2001). For LOI measurements, samples were wet sieved to remove the material 296

larger than 2 mm, dehydrated at 105°C for 24 hours, and burned in a muffle furnace at

550°C for 4 hours. LOI was calculated by the difference in sample dry weight at 105°C and 550°C. Samples were then burned in a furnace at 950°C for 1 hour to calculate carbonate content.

Foraminiferal analysis

A total of fifty seven surface samples and 20 core samples were analyzed for foraminifera from the two marshes. These samples were wet-sieved through 500 μ m and 63 μ m sieves. The >500 μ m fraction was examined for larger foraminifera before being discarded. The fraction between 63 and 500 μ m was subdivided into eight equal aliquots following the most reliable method for splitting a foraminiferal sample described in Horton and Edwards (2006) using a wet splitter (Scott and Hermelin, 1993). Where possible, at least 300 tests (dead and live) were counted in water which enabled easy detection of rose Bengal stained foraminifera and prevented drying of the organic residue (e.g., de Rijk, 1995; Horton and Edwards, 2006). In the short core samples, living specimens from below the surface (1-10 cm) sediment layer were considered as infaunal.

Taxonomic identifications follow Gehrels and van de Plassche (1999), Horton and Edwards (2006), Hawkes et al. (2010), Wright et al. (2011), and Milker at al. (2015a, b). The genus Jadammina is now considered a junior synonym of Entzia (Kaminski and Filipescu, 2011). However, we still employ the genus *Jadammina* in this study in order to avoid confusion since the majority of salt-marsh foraminiferal researchers use it. Juvenile specimens of Trochammina inflata and Jadammina macrescens that were difficult to distinguish from each other because of their small sizes were lumped into a single group and assigned as juvenile Trochamminids. Specimens of the genus Ammobaculites were combined into a single group, because these species were often broken making it difficult to identify them to the species level (Kemp et al., 2009a; Milker et al., 2015a). All counts were expressed as numerical abundance of living and dead foraminifera numbers per 10 cm³ bulk sediment and as a relative abundance (%) of species out of the total group. All foraminiferal data can be found in the supplementary material (Supplementary data 1, 2 and 3). Scanning electron microscope photographs of key species were taken at the Department of Earth, Planetary and Space Sciences, UCLA.

3.3. Statistical analysis

Dead surface foraminiferal assemblages were statistically analyzed to minimize the seasonal fluctuations commonly recorded in live populations and because they most accurately reflect the subsurface assemblages indicating that taphonomic processes are minimal (e.g., Culver and Horton, 2005, Milker et al., 2015a and references therein). Two samples with low counts (<4 specimens) were removed from the statistical analyses while the remaining 55 surface samples (average of 417 individuals per sample) with 84 as the lowest specimens number were included.

Multivariate statistical methods (unconstrained Cluster Analysis, CA, and detrended correspondence analysis, DCA) are used to identify and classify the distribution of groups and subgroups in the foraminiferal community into homogeneous faunal zones (clusters) along each transect (e.g., Frezza and Carboni, 2009; Phipps et al., 2010).

347 A Q-mode CA was processed by PRIMER version 6 software (Plymouth Routines In Multivariate Ecological Research, UK). The data of the common benthic species 348 (>1% relative abundance) were double-root transformed in order to down-weight the 349 350 relative contribution of highly abundant species. These transformed abundances were used to build a similarity matrix calculated between every pair of samples comprising 351 a Bray-Curtis similarity (Bray and Curtis, 1957) as the distance metric, and used to 352 353 reconstruct a dendrogram for hierarchical clustering (group average linkage). 'Similarity profile' (SIMPROF) permutation tests were used to identify significant 354 groupings, such that the group being subdivided had a significant (p < 0.05) internal 355 356 structure. A 'similarity percentages' (SIMPER) routine was used in order to identify species that contributed most to the similarity within each sample cluster, as well as to 357 dissimilarities with other clusters. 358

359 360

361

362 363

364

365366

367 368

369

370

371372

373

374375

376

377

378

A DCA was carried out in order to provide further information about the patterns of variation in the faunal data and to determine the type of response displayed by the species distribution to one or more environmental gradients, a unimodal or linear response (e.g., Leps and Smilauer, 2003). Choosing an appropriate ordination technique to quantify the direct species—environment relationships relies on determining the extent of species turnover (the beta diversity in community composition) measured in standard deviation (SD) units along an environmental gradient, as calculated by DCA (e.g., Leps and Smilauer, 2003; Barlow et al., 2013). Gradient lengths of SD <3 point toward linear faunal response along the environmental gradient, whereas those of SD >4 indicate unimodal species environment relationships. DCA showed a linear species response and therefore RDA was used. RDA was applied to quantify the direct relationship between the distribution of benthic foraminifera to the elevation and abiotic ecological variables including pore-water and sediment properties (Table 2). This ordination technique is based on a linear species—environment relationship, where the axes are linear combinations of the environmental variables. It integrates variance as well as correlation in its search for relationships between the two sets of biotic and abiotic variables (Leyer and Wesche, 2007). In order to further test the correlation between the species distribution and elevation we used the parametric correlation coefficient Pearson's r.

379 380 381

382

383

384 385

386

387

388

389

390

Both DCA and RDA were applied using Canoco, version 4.55 software (Leps and Smilauer, 2003; Ter Braak and Smilauer, 2002). For DCA, detrending by segments was selected and for RDA environmental parameters were standardized and Monte Carlo permutation tests (2000 permutations) were performed. In both statistical methods species data were log transformed because they don't show normal distribution based on the Shapiro-Wilk test calculated using the PAST software package vs. 3.11 (Hammer et al., 2001) and also due to the rather uneven vertical distribution of the samples (Gehrels, 2000; Telford and Birks, 2011). We calculated inter-correlations between environmental parameters and statistical significance as well as the Pearson's correlation coefficient (r value) with the PAST software package vs. 2.15 (Hammer et al., 2001).

391392393

394

395

396

DCA results were also used as an initial step for choosing an appropriate unimodal or linear-based transfer function. In order to develop transfer functions ('ecological response functions') for predicting marsh surface elevations, the relationship between water depth and the relative abundances of foraminiferal taxa within the combined

397 training data set was empirically modeled in the C2 program (version 1.7.4; Juggins, 2011). Although the difference in tidal range between the two sites is small, sample 398 elevations were normalized using the Standardised Water Level Index (SWLI). 399 commonly applied in salt-marsh sea-level studies (Gehrels, 1999; Horton et al., 1999; 400 Wright et al., 2011; Barlow et al., 2013; Kemp and Telford, 2015). The SWLI value is 401 calculated as: (100 × (sample elevation – local mean tidal level) / (high water 402 403 reference level – local mean tidal level)) + 100. We used the mean higher high water as the high water reference level (Kemp and Telford, 2015). Although Wright et al. 404 (2011) recommend to use the highest occurrence of foraminifera as the high water 405 406 reference level in SWLI calculations, we could not establish this datum because our highest samples still contained foraminifera. All elevations are expressed relative to 407 the same datum and a SWLI of 100 is MTL and 200 is MHHW. 408

409 410

411

412413

414

415 416

417

418 419

420

421 422

423

424

Model performance was evaluated based on the cross-validated (leave-one-out/ jackknifed) correlation between model prediction elevation and measured elevation (r^2 _{iack}), the root mean squared error of prediction (RMSEP) and the tendency of the model to overestimate (underestimate) parts of the training set gradient, as mean and maximum bias (Birks, 1998; Juggins and Birks, 2012; Kemp and Telford, 2015). The observed residual scatter and observed and predicted values after model prediction in the produced scatterplots were also taken in consideration as complementing the evaluated model performance (Telford et al., 2004). Sampling of the modern analogue along transects (e.g., evenly stratified by elevation and or within one site/marsh) may lead to spatial autocorrelation (increased resemblance of samples) and thus might negatively impact the transfer function model performance (Legendre and Fortin, 1989; Telford and Birks, 2009). The issue of spatial autocorrelation is reduced by sampling two different salt marshes and combining the results into one modern training set. Consequently, we present a robust foraminifera-based transfer function from a modern analogue potentially suitable for RSL reconstructions in Southern California.

425 426 427

4 Results

4.1. The environmental properties of surface sediment

429 430

428

Water temperature (WT) varies between 24 °C and 32 °C and between 15 °C and 20 431 °C in Seal Beach and Tijuana salt marshes, respectively (Fig. 2; Table 2). Differences 432 in WT between the two marshes are related to differences in sampling months 433 (October vs. December) and differences within each marsh are related to sampling 434 hours during the day, in each transect. The pH values of the pore-water show no 435 observable spatial trend and range commonly from 6.1 to 7.8 in Seal Beach excluding 436 one sample with the lowest pH (4.3) and from 6.4 to 7.3 in Tijuana. Low O₂ 437 concentration is measured in the most elevated stations of Seal Beach and thereafter a 438 decreasing trend is observed ranging from 8 to 0.3 mg/l. O₂ concentration were rather 439 low, <1.6 3 mg/l, at all stations in Tijuana excluding two stations with values of 6.6 440 and 3.3 mg/l. Salinity values in Seal Beach show a decreasing trend from the high 441 marsh samples (~40 ‰) to the low marsh samples (~34 ‰) whereas in Tijuana the 442 443 values varies between 27 ‰ and 38 ‰ in all stations apart of one high marsh sample with 20 %. The organic matter content is fluctuating in both transects and ranges from 444 67 to 1% and from 12 to 75% in Seal Beach and Tijuana salt marshes, respectively. 445 The organic matter content in both transects was the highest in the mid marsh 446

447 sediments (apart from two samples in Seal Beach in which sparse vegetation occurs) and it decreases somewhat towards the low marsh edge due to the decrease in 448 vegetation cover. The carbonate content varied between 1% and 22% and between 1% 449 and 11% in Seal Beach and Tijuana salt marshes, respectively, with the higher values 450 occurring in the mid marsh samples in both transects (apart from two samples in Seal 451 Beach which had also low organic matter content). There was no clear trend in the 452 453 grain-size distribution along the intertidal zone of both transects. Tijuana salt marsh is dominated generally by silty sediment (62-77%), except in two sandy high marsh 454 stations (TJE1, TJE2; Fig. 2; Table 2). Seal Beach stations are characterized by silty 455 456 to sandy sediments with low concentrations of clay (<14%). Increasing sand concentrations in the lower salt-marsh stations of Seal Beach (SB8-SB17) are 457 probably related to a tidal velocity gradient. 458

459 460

4.2. Distribution of live (rose Bengal stained) foraminifera in the surface samples

461 462

463

464

465 466

467

468 469

470

471

The distribution of the live dominant taxa identified along transects from Seal Beach and Tijuana salt marshes are summarized in Figs. 3 and 4, respectively. Nine different agglutinated taxa were found in the live (rose Bengal stained) populations from both sites which were dominated by Jadammina macrescens, Trochammina inflata, Miliammina fusca, and Ammobaculites spp. Other species such as Trochamminita irregularis, Miliammina petila and Scherochorella moniliformis were observed in rather low numbers (Plate 1). Calcareous species were found mainly in the tidal flat and low marsh sediments. High abundances of juvenile calcareous miliolids were encountered in several high marsh samples. These were probably transported by tides via nearby tidal creeks (Plate 2) as suggested by the high sand content in the same samples.

472 473 474

475

476

477

478 479

480

481 482

483

484

The live species exhibit observable zonation relative to elevation. At both sites, the relative abundance of *M. fusca* is larger at lowest elevations. The tidal flat sediments at Seal Beach (St. 16) are dominated by this species with general abundances between 92 and 496 specimens/10 cm³ sediment volume and S. moniliformis with 108 and 216/10 cm³ (Fig. 3). In the low marsh sediments at Seal Beach (St. SB8-SB17) and at Tijuana (St. TJE10-TJE 15) these two species occur in decreasing numbers (Figs. 3 and 4). In the middle marsh these species are replaced by high concentrations of J. macrescens and T. inflata (St. TJE7-TJE 9 and TJE 16), while at Seal Beach (St. SB3-SB7) Ammobaculites spp. also occur. The highest stations (St. SB1 and SB18-SB20 and St. TJE1-TJE 3 and TJE17) are characterized by higher numbers of J. macrescens and T. inflata and occurrence of T. irregularis and M. petila (Figs. 3 and 4, Supplementary data 1 and 2).

485 486 487

4.3. Distribution patterns of dead foraminiferal assemblages in the surface samples

488 489 490

491

492

The dead foraminiferal assemblages are dominated by the same taxa as the live (rose Bengal stained) populations at the same stations with occasional higher concentrations in both marshes (Figs. 3 and 4).

The dead foraminiferal zonation defined by the Q mode cluster analyses mimics the 493 494

vascular plant zonation in Seal Beach saltmarsh whereas in Tijuana it differs slightly

from the zonation of vascular plants. Three main sample groups (clusters 1, 2 and 3) 495 were distinguished in each salt marsh (Figs. 5 and 6). 496

At Seal Beach, cluster 1 contains samples from the low marsh (St. SB17 and SB8-498 SB15 and their replicates) with elevations between 1.37-1.67 m NAVD88 designated 499 as cluster 1b and samples from tidal flat (St. SB16, 2SB16) with only one elevation of 500 0.98 m NAVD88 labeled as cluster 1a. Both sub-clusters are characterized by J. 501 macrescens (7-68%), T. inflata (18-57%), calcareous species (1-25%) and M. fusca 502 503 (1-29%). The latter species is more abundant in cluster 3a where it is also accompanied by S. moniliformis (5-11%). Cluster 2 consists entirely of mid marsh 504 stations (St. SB2-SB7 and their replicates) with elevations between 1.48 and 1.58 m 505 506 NAVD88 (Fig. 5, Table 2, Supplementary data 1). This cluster is divided into two sub-clusters 2a and 2b. Sub cluster 2a is characterized by *J. macrescens* (21-75%), 507 Ammobaculites spp. (12-62%) and T. inflata (<20%). Sub cluster 2b is dominated by 508 J. macrescens (57-89%) and accompanied by T. inflata (<29%) and calcareous 509 species (<5%), Cluster 3 includes stations from the high marsh (St. SB1, SB18-SB20 510 and their replicates) with elevations between 1.56 and 1.79 m NAVD88. The most 511 significant species contributing to this cluster are J. macrescens (>46%), T. inflata 512 513 (<38%) and *T. irregularis* (<7%).

514

At Tijuana, cluster 1 consists in part of the low marsh stations (St. TJE12-TJE15) with 515 the lowest elevations between 1.39 and 1.48 m NAVD88 (Fig. 6, Table 2, 516 Supplementary data 2). This cluster is dominated by *J. macrescens* (21-52%), 517 Trochammina inflata (18-48%), calcareous species (11-38%), M. fusca (4-9%) and by 518 519 S. moniliformis (1-3%). Cluster 2 contains samples from the mid to low marsh (St. TJE4 – TJE11and TJE16) with elevations between 1.42-1.59 m NAVD88. Samples in 520 this cluster are dominated by J. macrescens (35-72%), T. inflata (20-51%) and 521 522 calcareous species (1-13%). Cluster 3 encompasses the high marsh samples (St. TJE1-3 and TJE17) with elevations between 1.57-1.93 m NAVD88. The dominant species 523 in this cluster are similar to those occurring in cluster 3 at Seal Beach including J. 524 macrescens (>40%), T. inflata (12-44%), T. irregularis (4-12%) and M. petila (<7%) 525 (Fig. 6). 526

527 528

4.4. Surface and infaunal foraminiferal distribution in the short cores

529530531

532

533

534535

536

537

538

539

540 541

542

543544

545

546

The infaunal foraminiferal distributions in the short cores are shown in Fig. 7. The down core distribution of dead foraminifera is given in Supplementary data 1 and 2. Similar foraminiferal species are currently living (rose Bengal stained) at the surface (0-1 cm depth) of the transition from mid to low marsh (St. SB5) of Seal Beach and of the mid marsh (St. TJE8) of Tijuana with general abundances of 872/10 cm³ and 340/10 cm³, respectively. The living assemblage at the surface of Seal Beach is dominated by J. macrescens (624/10 cm³), calcareous species (192/10 cm³) and T. inflata (56/10 cm³). The living assemblage at the surface of Tijuana is dominated by T. inflata (188/10 cm³), J. macrescens (108/10 cm³), and calcareous species (24/10 cm³). Living infaunal specimens of the same surface living species are observed between 1 cm and maximum down to 6 cm, showing general decreasing numbers with core depth. In the shallowest depths (1-2 cm), the total highest numbers of infaunal specimens representing 16% and 10% out of the total assemblages, in Seal Beach and Tijuana, respectively, are observed. The most abundant infaunal species is T. inflata (672/10 cm³ or 80/10 cm³ at 1-2 cm in Seal Beach and Tijuana, respectively) which also found deeper (down to 6 cm) compared to other infaunal species.

4.5. The relationship between foraminiferal composition and environmental variables

The short lengths of the first DCA axis in turnover (less than 2 SD) units for each transect and for the combined surface assemblage of both marshes (Table 3a-c) indicate that the foraminifera demonstrate a linear response to one or more environmental gradients and thus a linear ordination method (i.e., RDA) is expected to perform well. The applied Monte Carlo tests suggest a significant influence (p < 0.05; Table 4a) of several environmental parameters on the species distribution in the modern data set of Seal Beach. O₂ explains 24.1% of the variance in the data set, elevation, salinity, carbonate content and temperature, each justifying values between 16.3% and 19.9% of the variance in the data set. These significant environmental variables in Seal Beach are correlated to the first two axes, calculated with the RDA, which explain 51% of the cumulative variance of the species data and 94% of the species-environment relationship (Table 4a, Fig. 8A). However, in Tijuana, elevation and TOM were found to be the significant environmental parameters, explaining 43.2% and 8.3% of the variance in the data set, respectively (Table 4b, Fig. 8B). Axes one and two explain 56% of the total variance in the foraminiferal data (Table 4b).

The combined RDA results for both salt marshes show that elevation is the most significant environmental parameter explaining 17.1% of the variance in the data set. Other parameters explaining the variance in the data set include O_2 (12.2%), temperature (7.6%), salinity (10.9%), pH (2.1%) with the remainder (50%) consisting of inter-correlations between variables (Table 4c). The first two axes calculated with the RDA explain 38% of the cumulative variance of the species data and 93% of the species-environment relationship (Table 4c, Fig. 9).

 In general, the results of the RDA (Figs. 8 and 9) support the Q-mode CA for both sites (Figs. 5 and 6). The distance between the samples with relatively good separation of the three assemblages within each transect is distinguished in the three dendrograms (1, 2, 3; Figs. 8 and 9) and shows the dissimilarities between the sample clusters and the similarities within each sample cluster. In the RDA ordination diagrams, cluster 3 samples representing high marsh sediments in both salt marshes are plotting around the elevation arrow while cluster 1 samples, representing mid to low marsh and low to tidal flat sediments, with lower elevations plotting mostly in the opposite direction.

 Several samples of cluster 2 plot close to the carbonate content, temperature and O_2 arrows, all pointing in the same direction (Figs. 8A and 9). The arrow of salinity is in the same direction as the arrow of elevation in Seal Beach indicating an intercorrelation between the two and reflecting a similar general trend from higher values in the high marsh to lower values towards the low marsh (Figs. 8 and 2).

The relationships between elevation, the most significant environmental parameter (Table 4c) and abundance of the six strongly correlative foraminiferal species are shown in Fig. 9. *Jadammina macrescens* and *T. irregularis* are positively correlated with elevation, although rather weakly (r = 0.2). *Miliammina petila* occurs in low relative abundances at high marsh sites (i.e., in Tijuana) does not seem to have any correlation with elevation (r = -0.03). Conversely, *M. fusca*, calcareous spp. and *S.*

moniliformis occurring in the low marsh to tidal flat stations, are negatively correlated with elevation (Figs. 5, 6, 8 and 9) (r values range from -0.3 to -0.5). Trochammina inflata occurs in high relative abundances at low marsh as well as at high marsh sites, notably in Seal Beach, and does not seem to have any correlation with elevation (r = -0.1) (Fig. 8A). Ammobaculites spp. that have higher abundances in the mid marsh sediments, cluster 2 of Seal Beach, are positively correlated with O₂ (Figs. 8A and 9).

4.6. Development of a foraminifera-based transfer function

DCA of the Southern California training set, combining Seal Beach and Tijuana salt marshes with the SWLI as the environmental variable, produced a gradient length of 1.98 (Table 3c), allowing us to apply linear regression models, i.e., PLS (Birks, 1995). Nonetheless, WA-PLS also performs well for a linear species distribution along such an environmental gradient and while providing 'the minimal adequate model' (Birks, 1998 and reference therein), occasionally it outperforms PLS (Ter Braak et al., 1993). For this reason we tested also the WA-PLS model predictions.

Prior to the development of the foraminifera-based transfer function we performed a screening exercise. Three samples that contain high abundances of *Ammobaculites* species (>26%) and belonging to cluster 2a (Fig. 5) were excluded from the original 55 salt-marsh samples. The reasoning behind this decision is that these species are usually found broken in surface marsh sediments (Kemp et al., 2009a; Milker et al., 2015a) and to a greater extent they are hardly preserved within the subsurface

fossilized record and thus, are not considered as useful proxy species for sea-level 620 studies.

621

597

598

599

600

601

602 603

604 605 606

607

608 609

610

611

612 613

614

615 616

617

618 619

622

623

624 625

626

627

628

629

630

631 632

634 635

636

637

638

639

640 641

642

643 644

645

646

In order to improve the predictive ability of the transfer function we removed samples that have a difference between the predicted and observed values larger than 25% of the total observed elevation range, between 120.7 and 222.8 m NAVD88 (following Edwards et al., 2004). Consequently, PLS was developed from a screened training set of 43 samples and 8 species whereas a WAPLS was developed from 44 samples and 8 species. The final component in each transfer function was chosen according to the highest r^2 and the lowest RMSEP and maximum bias values if the reduction in prediction error exceeds 5% for this component compared to the next lowest component (Ter Braak and Juggins, 1993). This decision path was merely applied to the first three components in order not to add statistical complexity (Wright et al.,

2011; Barlow et al., 2013). 633

> The performance of both regression methods is presented in Table 5. The statistical output of the PLS foraminiferal transfer function shows that the second component of our model performed better than the first component providing the highest correlation between the observed and predicted values with $r^2_{\text{Jack}} = 0.49$, and the lowest rootmean squared-error of prediction (RMSEP Jack) of 14.7 SWLI. Whereas the WAPLSbased- transfer-function shows that the third component performed significantly better than the other components and also compared to the PLS model, with the lowest RMSEP_{Jack} of 11.9 SWLI, higher $r^2_{Jack} = 0.72$ and lower maximum bias values (22.7). The transfer function shows a stronger performance between observed and predicted SWLIs as also reflected in the scatter plots for the WAPLS (Fig. 10). A trend was found in the residuals of the WAPLS component 2, but this trend was weaker with the use of component 3. Nevertheless the residual scatter is suggesting that some

elevation predictions are underestimated especially from the landward upper edge of the gradient (high marsh).

649 650

647

648

5 Discussion

651 652 653

654 655

656

657

658

659

660

661

662

663

5.1. Environmental controls on modern foraminiferal distribution

Among the environmental variables controlling the spatial foraminiferal distribution across salt marshes (e.g., salinity, organic carbon concentration, grain-size, pH) (e.g., Patterson, 1990; de Rijk and Troelstra, 1997; Scott et al., 1998; Woodroffe et al., 2005; Milker et al., 2015a; Barnett et al., 2016), elevation which is associated with tidal inundation (or subaerial exposure) is frequently identified as the primary control, at least indirectly, responsible for their zonation (Gehrels, 2000; Horton et al., 2003; Edwards et al., 2004; Horton and Edwards, 2005; Barlow et al., 2013; Milker et al., 2015a). Consequently, these organisms have been extensively used as proxies for sealevel reconstruction (e.g., Scott and Medioli,1980; Patterson et al., 2004; Barlow et al., 2013).

664 665

The results of the Q-mode CA and the RDA from Southern California further support 666 the hypothesis that foraminiferal assemblages are strongly related to elevation and 667 hence tidal submergence. At Tijuana elevation explains 43% of the total variance in 668 the foraminiferal assemblages whereas at Seal Beach, 16% of the variance of the 669 670 species data is explained by elevation (Fig. 8, Table 4a-b). Despite the difference in 671 the elevational control between marshes, the combined RDA results of both data sets still show that elevation is the primary factor influencing the foraminiferal zonation, 672 explaining 17.1% of the variance in the data set (Fig. 9, Table 4c) similar to studies in 673 the Magdalen Islands in the Gulf of St. Lawrence, Canada (Barnett et al., 2016), 674 Oregon, USA (Hawkes et al., 2010) and the UK (Horton and Edwards, 2006). 675 Elevation is followed by four additional environmental variables explaining lower 676 variance in the data set including O₂ (12.2%), temperature (7.6), salinity (10.9), pH 677 (2.1) and with the remainder (50%) consisting of inter-correlations between variables 678 (Table 4c). Inter-correlations between environmental variables and foraminifera are 679 known from other intertidal studies since varied factors also may be governed by the 680 frequency and duration of tidal exposure as controlled by elevation (Horton et al., 681 2003; Horton and Edwards, 2005). For example in Seal Beach a significant (p < 0.05, 682 r = 0.36) correlation between elevation and salinity was observed (Fig. 8b). 683 Pore-water salinity is an additional key factor (after elevation) known to control 684 marsh foraminiferal distributions (Murray, 2006), for example in tidal marshes with 685 extremely low gradients having complex topography such as of the Great Marshes of 686 eastern Massachusetts (e.g., de Rijk and Troelstra, 1997). Salinity of the sediment 687 pore-water generally varies considerably due to the combined influence of the tidal 688 cycle and infiltration of seawater, desiccation, precipitation and seepage of fresh 689 ground water. The Southern California marshes experience frequent droughts under 690 dry climate conditions. Pore-water salinity decreases from the high marsh (~40 %) to 691 the low marsh (\sim 34 ‰) at Seal Beach marsh (Fig. 2). This trend is also measured by 692 693 the RDA results where salinity explains 16.8% of the variance in the species data and it has a positive correlation with elevation (Fig. 8b, Table 4b). On the other hand, 694 salinity values at Tijuana were more variable and thus had no significant influence on 695 the distribution pattern of the foraminifera. It seems that this difference between both 696

marshes is associated with their topographic characteristics with Tijuana transect having a simple gradient with no topographic complications whereas Seal Beach has more complex and irregular topography. Other studies also imply that when the topography is irregular, there is no very clear vertical foraminiferal zonation with respect to the tidal frame (de Rijk, 1995; de Rijk and Troelstra, 1997; Müller-Navarra et al., 2016). Nonetheless, since salinity varies considerably through the tidal cycle and our measurements were done only at the time of the sample collection during summer when evaporation is playing an important role, not much significance is attached to our measurements of this environmental parameter.

At Tijuana marsh, elevation is followed by TOM explaining 8.3% of the variance in the data set (Fig. 8b, Table 4b). Mainly mid to low samples dominated by *J. macrescens* and *T. inflata* were positively related to the TOM (Figs. 6, 8b). These species are known to prefer higher amounts of organic matter (Armynot du Châtelet et al., 2009). When the data are combined, the clear control of TOM on the distribution of foraminifers in Tijuana becomes masked.

Further environmental parameters probably affect the distributions patterns of foraminifera in Seal Beach, for example O₂ and temperature (Fig. 8). Substrate oxygenation is an important factor determining the degree of infaunal foraminiferal dispersion into the substrate (Walker and Goldstein, 1999). O₂ and temperature also influence organic matter degradation and thus the preservation of agglutinant species organic cement (Berkeley et al., 2007 and references therein). However, the potential of preservation of agglutinant species is also inherent selectively by different species (Berkeley et al., 2007 and references therein). At Seal Beach, O₂ and temperature are positively related to mid marsh sediment that are the only samples characterized by higher abundances of Ammobaculites spp. (Fig. 8A). It may be that their preservation specifically at the mid marsh is associated with more stable environmental conditions compared to the salt marsh edges. Similar fragility was related to the coarseness of the grains making up the tests of M. fusca explained their poor preservation (de Rijk and Troelstra, 1999) compared to other more packed grains of relatively robust species such as T. inflata (Berkeley et al., 2007 and references therein). It is worth mentioning that this study was not designed to determine the temporal variability of intertidal foraminifera in the surface sediments from a time series survey (e.g., by monthly sampling). Consequently, we cannot attribute more significance to these factors which fluctuate daily, through the tidal cycle and seasonally.

5.1.1. Vertical zonation of dead for aminifera

The comparison between the dead foraminiferal assemblages across the salt-marsh surfaces of Seal Beach and Tijuana show that their distributions are very similar to each other and are also comparable to foraminiferal distributions in other parts of the world. This indicates that despite the recent hydrological changes, the foraminiferal distributions are robust to these changes and their prominent elevational zonation (with site-specific differences) is not affected.

The dead foraminiferal assemblages of the tidal flat and low marsh of the studied sites are characterized by the occurrence of *M. fusca* and calcareous species. The agglutinant *M. fusca* is known as an indicative species of lower elevations, for example along the North America's Pacific coast (e.g., Nelson et al., 2008; Hawkes et

747 al., 2010; Engelhart et al., 2013, Milker et al., 2015a) and the North America's Atlantic coast (e.g., Edwards et al., 2004). In the middle marsh the dominant dead 748 species include J. macrescens, T. inflata and Ammobaculites spp. Trochammina 749 750 inflata and more prominently J. macrescens are also dominating higher elevations of the high marshes at both sites where they are accompanied by T. irregularis and M. 751 petila. Jadammina macrescens and T. inflata have frequently been reported from 752 753 middle marsh (Guilbault et al., 1996; Jennings and Nelson, 1992; Nelson et al., 2008; Hawkes et al., 2010; Engelhart et al., 2013, Milker et al., 2015a, b) and the highest 754 marsh environments (e.g., Patterson, 1990; de Rijk and Troelstra, 1997; Horton et al., 755 756 1999; Wang and Chappell, 2001; Hippensteel et al., 2002; Horton et al., 2003; 757 Edwards et al., 2004a; Franceschini et al., 2005; Woodroffe et al., 2005). Trochamminita irregularis (often grouped with T. salsa) has commonly been detected 758 as a dominant species in the high and highest marsh in North American Pacific 759 marshes (Hawkes et al., 2010; Engelhart et al., 2013; Milker et al., 2015a, b) and in 760 New Zealand (e.g., Hayward et al., 2004; Southall et al., 2006) and Tasmania (Callard 761 et al., 2011). Similarly M. petila has been also observed in the middle and high 762 763 marshes from Oregon (Engelhart et al., 2013). The agglutinant Balticammina pseudomacrescens is an important sea-level indicator (Gehrels and van de Plassche, 764 1999) dominating higher elevation of mid to highest marsh zones in North American 765 766 Pacific marshes, for example in Oregon salt marshes (Hawkes et al., 2010, 2011; 767 Engelhart et al., 2013; Milker et al., 2015a, b) but is absent in the studied salt marshes in California. 768 769

5.2. The influence of taphonomic processes

770771

772 773 Processes as sub-surface foraminiferal production (infauna) and taphonomic changes vary temporally and spatially, from one intertidal system to another, and may 774 influence the fossil assemblages (e.g., Martin, 1999; Berkeley et al., 2007). 775 Accordingly there is a clear need to explore these processes in each depositional 776 environment, as a base for employing an adequate surface sampling strategy aiming at 777 accurate paleoenvironmental interpretations based on the microfossil record. At Seal 778 Beach and Tijuana, test production mirrors rather proportionally the standing crop 779 numbers, a principle discussed in Berkeley et al. (2007) (Figs. 3 and 4). Samples with 780 high numbers of standing crops yielded high total numbers (per 10 cm³ sediment 781 782 volume) of dead foraminifera though exceeding the living ones, often by over an order of magnitude. This observation applies to most species excluding S. 783 moniliformis and calcareous species in both marshes which showed opposite trends 784 785 with higher numbers of live specimens compare to the dead numbers. This is probably 786 due to taphonomic processes and low preservation potential (Figs. 3 and 4). Dead agglutinant specimens of S. moniliformis were found only in lower elevations. These 787 are very fragile and tend to break easily due to the loss of organic cement by oxidation 788 (Berkeley et al., 2007). On the low marsh and tidal flat lower numbers of dead 789 790 calcareous species were observed compared to the live foraminifera (Figs. 3 and 4, Supplementary data 1 and 2). This difference along with apparent dissolution (Plate 2, 791 Fig. 6b) is probably associated with reduced mean pH values of 6.5 and 6.9 in Seal 792 Beach and in Tijuana, respectively, causing carbonate dissolution and lower 793 794 preservation of the calcareous tests. Early diagenetic dissolution and loss of calcareous tests related to a pH <~7, which results in a carbonate undersaturation is a 795 known process from intertidal environments along the North-West American coastal 796

marshes (e.g., Hawkes et al., 2010; Milker et al., 2015a), salt marshes in the
Mediterranean region (Cundy et al., 2000; Shaw et al 2016) as well as from other
temperate and tropical coasts (Wang and Chappell, 2001; Edwards and Horton, 2000;
Horton and Murray, 2006; Berkeley et al., 2009). Moreover, a general spatial trend of
better preservation at higher elevations was observed at Tijuana (Fig. 4), possibly
associated with diagenetic controls on the preservation potential of calcareous and
agglutinated tests.

5.3. Surficial vs. subsurface foraminifera

 Most paleoenvironmental studies of intertidal marshes use the foraminiferal assemblages in surface sediment (upper 1 to 2 cm) as a modern counterpart assuming that they are characterized primarily by epifaunal production (e.g., Scott and Medioli, 1980; Gehrels, 1994; de Rijk, 1995; Horton, 1999; Patterson et al., 2004; Woodroffe et al., 2005; Horton and Edwards 2006; Kemp et al., 2009a) or shallow infaunal (Buzas et al., 1993). However, several studies have shown that intertidal foraminifera can live infaunally as deep as several decimeters in some marsh settings (e.g., Goldstein et al., 1995; Ozarko et al., 1997; Goldstein and Watkins, 1998; Saffert and Thomas, 1998; Goldstein and Watkins, 1999; Patterson et al., 1999; Hippensteel et al., 2002; Culver and Horton, 2005; Duchemin et al., 2005; Tobin et al., 2005; Berkeley et al., 2007; Leorri and Martin, 2009; Milker et al., 2015a). In such cases sampling the uppermost surface sediment (0-1 cm) would produce poor modern analogues (Duchemin et al., 2005). For example Ozarko et al. (1997) and Patterson et al. (1999) relied on a deeper surface sampling (to 10 cm), covering the entire distribution of the living infauna. Relying on such thick intervals may be misleading as they might represent a longer time span, in some cases a decade or longer, in some salt marsh settings (Milker et al., 2015a).

 At Seal Beach and Tijuana the highest numbers of living specimens at the transition to middle marsh (St. SB5) and middle marsh (St. TJE8) are found in the surface samples (Fig. 7). The numbers of infaunal specimens, restricted to the first six cm of the cores, are low compared to the total foraminiferal assemblage and show a decreasing trend from 16% and 10% between 1 and 2 cm at Seal Beach and Tijuana, respectively, to nil (Fig. 7). These results are in accordance with other salt-marsh studies reporting living species predominantly in the top few centimeters of salt marsh in North Carolina (Culver and Horton, 2005), New Brunswick, Canada (Patterson et al., 2004) and the UK (Horton, 1997; Horton and Edwards, 2006).

Furthermore, the same infaunal species are presently living in the surface samples indicating that none of them are exclusively infaunal. Additionally the dominant infaunal species *T. inflata*, *J. macrescens* and calcareous miliolids are also dominating the dead surface and fossil assemblages in both marshes. Consequently, our results show that the influence of infaunal species on the dead assemblage is minor since most reproduction is concentrated near the sediment surface and that the same post-depositional processes influence fossil assemblages as the dead surface assemblages. These observations enable us to use the dead assemblages in the upper surface sediments as an adequate modern analogue for accurate quantitative paleoenvironmental interpretations.

The dead fossil assemblages in the mini cores analyzed also reveal interesting results. At Tijuana, similar species that dominate the surface dead assemblages occur along the entire 10 cm core with similar ratios and abundances. However at Seal Beach the down core shows a fossilized dead assemblage different from the surface dead assemblage. *Miliammina fusca* dominates the entire foraminiferal assemblage in the lower part of the core (Supplementary data 1 and 2) and numbers gradually decrease towards the top. This faunal change is indicating a change from a low marsh setting to a mid-marsh setting (Figs. 3-6).

5.4. Robustness of the transfer functions for relative sea-level estimates

This study demonstrates that elevation related to tidal inundation is a primary control on the foraminiferal distributions across the two salt marshes in Southern California and our data are thus suitable for developing a transfer function for reconstructing past sea-level changes (described in chapter 5.1). The WAPLS transfer function (component 3) is the favored model with the best statistical performance compared to the PLS model (described in chapter 4.6) (Table 5, Fig. 10, Supplementary data 4). The performance of the transfer function is also evaluated by the positive correlation between transfer function model prediction error (RMSEP Jack) and tidal and elevation ranges at the studied sites (e.g., Callard et al., 2011; Barlow et al., 2013; Mills et al., 2013). Earlier studies show that salt-marsh foraminifera-based transfer function models produce RMSEPs values between ~3–23% of the tidal range in a micro-tidal setting characterized by a <2 m tidal range (summaries in Callard et al. (2011) and Barlow et al. (2013)). The elevation range of the modern analogue is additionally influencing the error terms of RSL reconstruction, with a common RMSEPs between 5 and 15% of the elevation range of samples produced in most studies. In the current study the RMSEP value (11.9 SWLI or 0.09 m NAVD88) is 5.7% of the tidal range (1.6 m; Table 1) and 10% of the elevation range of the samples (0.95 m; Table 2). These values are within the common ranges presented by Barlow et al. (2013).

6 Conclusions

 This study, for the first time, quantifies the environmental drivers controlling the distribution of foraminiferal assemblages in two coastal salt marshes along the Southern California coast, Seal Beach and Tijuana. The Q-mode CA and the RDA results show that elevation is the primary factor influencing the foraminiferal zonation. Seal Beach has more topographical complexity than Tijuana salt marsh, and thus more environmental parameters (e.g., O₂, salinity), apart from elevation, affect the foraminiferal distribution. The dead assemblages in the upper surface sediments are an adequate modern analogue for quantitative paleoenvironmental interpretations due to i) concentrated presence in the upper cm of sediment enabling statistical analyses, ii) presence of the same dominant foraminifera in the live and dead assemblages in both marshes indicating similar influence of post-depositional processes, iii) exceedance of the abundances of the surface dead foraminifera compared to the live abundances, reflecting preservation of several generations of foraminifera.

The dead surface assemblages in both marshes show a distinct zonation with respect to elevation, similar to earlier studies, however with some notable site-specific variability. The tidal flat and low marsh are characterized by higher concentrations of *M. fusca* and calcareous species, the middle marsh is dominated by *J. macrescens*, *T. inflata* and *Ammobaculites* spp (primarily at Seal Beach) and the high marsh zone is dominated by *J.macrescens*, *T. inflata* accompanied by *T. irregularis* and *M. petila*.

The development of WA-PLS foraminiferal transfer functions based on the training set combining data from two salt marshes increases the likelihood that fossil assemblages are adequately represented in the modern training set. The regional training set also reduces problems of spatial autocorrelation. The performance of the transfer functions is supported statistically by a robust relationship between the observed and the predicted elevations ($r^2_{\text{Jack}} = 0.72$), the lower maximum bias values in the training set and by a precision of the reconstructions of 0.09 m NAVD88 (RMSEP Jack) which is 5.7% of the tidal range and 10% of the elevation range of the samples. Our results agree with other foraminiferal studies from micro-tidal saltmarsh environments and can be used for future paleoenvironmental reconstructions in the study area.

Acknowledgments

This study was supported by the US Department of Interior establishing grant #Y561461:03 for the Southwest Climate Science. Dr. Avnaim-Katav also expresses her gratitude to the Leon H. Charney School of Marine Sciences (University of Haifa, Israel) for the support via the Helmsley Charitable Trust. Thanks are due to Matthew E. Kirby (Cal-State Fullerton) for the granulometric analysis and to Andrew Fricker and Scott Lydon for their help with the field work. We gratefully acknowledge the Seal Beach Wildlife Reserve and the Tijuana River National Estuarine Research Reserve for site access. Matt Zebrowski of the Cartography Laboratory of the Department of Geography and Ming-Chang Liu of the Department of Earth, Planetary and Space Sciences, at UCLA (University of California Los Angeles), are thanked for the graphics and the SEM, respectively. Editor Ric Jordan and the two anonymous reviewers are deeply appreciated for their constructive comments and suggestions

References

- 935 Barlow, N. L., Shennan, I., Long, A. J., Gehrels, W. R., Saher, M. H., Woodroffe, S. A., &
- 936 Hillier, C. (2013). Salt marshes as late Holocene tide gauges. *Global and Planetary*
- *Change*, 106, 90–110.
- 938 Barlow, N. L., Long, A. J., Saher, M. H., Gehrels, W. R., Garnett, M. H., & Scaife, R. G.
- 939 (2014). Salt-marsh reconstructions of relative sea-level change in the North Atlantic during
- 940 the last 2000 years. *Quaternary Science Reviews*, 99, 1–16.

resulting in a significantly improved manuscript.

- 941 Barnett, R.L., Gehrels, W.R., Charman, D.J., Saher, M.H., & Marshall, W.A. (2015). Late
- Holocene sea-level change in Arctic Norway. *Quaternary Science Reviews*, 107, 214–230.

- 943 Barnett, R. L., Garneau, M., & Bernatchez, P. (2016). Salt-marsh sea-level indicators and
- transfer function development for the Magdalen Islands in the Gulf of St. Lawrence,
- 945 Canada. *Marine Micropaleontology*, 122, 13–26.
- 946 Berkeley, A., Perry, C. T., Smithers, S. G., Horton, B. P., & Taylor, K. G. (2007). A review of
- 947 the ecological and taphonomic controls on foraminiferal assemblage development in intertidal
- 948 environments. Earth-Science Reviews, 83(3), 205–230.
- Berkeley, A., Perry, C. T., & Smithers, S. G. (2009). Taphonomic signatures and patterns of
- 950 test degradation on tropical, intertidal benthic foraminifera. *Marine Micropaleontology*, 73(3),
- 951 148–163.
- 952 Birks, H. J. B. (1995). Quantitative palaeoenvironmental reconstructions. Statistical modelling
- 953 *of quaternary science data. Technical guide*, *5*, 161–254.
- Birks, H. J. B. (1998). DG Frey and ES Deevey Review 1: Numerical tools in
- palaeolimnology–Progress, potentialities, and problems. *Journal of Paleolimnology*, 20(4),
- 956 307–332.
- 957 Birks, H.J.B. (2003). Quantitative paleoenvironmental reconstructions from Holocene
- biological data. In: Mackay, A.W., Battarbee, R.W., Birks, H.J.B., Oldfield, F. (Eds.), Global
- 959 Change in the Holocene. Arnold, London, pp. 342–357.
- 960 Birks, H.J.B. (2010). Numerical methods for the analysis of diatom assemblage data, In:
- 961 Smol, J.P., Stoermer, E.F. (Eds.), The Diatoms: Applications for the Environmental and Earth
- 962 Sciences, Second ed. University Press, Cambridge, UK, pp. 23–54.
- 963 Bray, J. R., & Curtis, J. T. (1957). An ordination of the upland forest communities of southern
- 964 Wisconsin. Ecological monographs, 27(4), 325–349.
- 965 Buzas, M. A. (1965). The distribution and abundance of foraminifera in Long Island Sound:
- 966 Smithsonian Institution Miscellaneous Collection, v. 149, 94 p.
- 967 Buzas, M. A., Culver, S. J., & Jorissen, F. J. (1993). A statistical evaluation of the
- 968 microhabitats of living (stained) infaunal benthic foraminifera. *Marine*
- 969 *Micropaleontology*, 20(3-4), 311–320.
- 970 Cahoon, D. R., Lynch, J. C., & Powell, A. N. (1996). Marsh vertical accretion in a southern
- 971 California estuary, USA. Estuarine, Coastal and Shelf Science, 43(1), 19–32.
- 972 Callard, S. L., Gehrels, W. R., Morrison, B. V., & Grenfell, H. R. (2011). Suitability of salt-
- 973 marsh foraminifera as proxy indicators of sea level in Tasmania. *Marine*
- 974 *Micropaleontology*, 79(3), 121–131.
- 975 Culver, S. J., & Horton, B. P. (2005). Infaunal marsh foraminifera from the outer banks,
- 976 North Carolina, USA. *The Journal of Foraminiferal Research*, 35(2), 148–170.
- 977 Culver, S.J., Leorri, E., Mallinson, D.J., Corbett, D.R., & Shazili, N.A.M. (2015). Recent
- 978 coastal evolution and sea-level rise; Setiu Wetland, Peninsular Malaysia. Palaeogeography,
- 979 Palaeoclimatology, Palaeoecology, 417, 406–421.

- 980 Cundy, A.B., Kortekaas, S., Dewez, T., Stewart, I.S., Collins, P.E.F., Croudace, I.W.,
- 981 Maroukian, H., Papanastassiou, D., Gaki-Papanastassiou, P., Pavlopoulos, K., & Dawson, A.
- 982 (2000). Coastal wetlands as recorders of earthquake subsidence in the Aegean: a case study of
- 983 the 1894 Gulf of Atalanti earthquakes, central Greece. *Marine Geology*, 170 3–26.
- De Rijk, S. (1995). Salinity control on the distribution of salt marsh foraminifera (Great
- 985 Marshes, Massachusetts). *The Journal of Foraminiferal Research*, 25(2), 156–166.
- 986 De Rijk, S., & Troelstra, S. R. (1997). Salt marsh foraminifera from the Great Marshes,
- 987 Massachusetts: environmental controls. *Palaeogeography*, *Palaeoclimatology*,
- 988 *Palaeoecology*, *130*(1), 81–112.
- 989 De Rijk, S., & Troelstra, S. (1999). The application of a foraminiferal actuo-facies model to
- 990 salt-marsh cores. Palaeogeography, Palaeoclimatology, Palaeoecology, 149(1), 59–66.
- 991 De Stigter, H. C., Jorissen, F. J., & Van der Zwaan, G. J. (1998). Bathymetric distribution and
- 992 microhabitat partitioning of live (Rose Bengal stained) benthic foraminifera along a shelf to
- bathyal transect in the southern Adriatic Sea. The Journal of Foraminiferal Research, 28(1),
- 994 40–65.
- 995 Du Châtelet, É. A., Bout-Roumazeilles, V., Riboulleau, A., & Trentesaux, A. (2009).
- 996 Sediment (grain size and clay mineralogy) and organic matter quality control on living
- 997 benthic foraminifera. Revue de micropaléontologie, 52(1), 75–84.
- 998 Duchemin, G., Jorissen, F. J., Redois, F., & Debenay, J. P. (2005). Foraminiferal
- 999 microhabitats in a high marsh: consequences for reconstructing past sea
- 1000 levels. Palaeogeography, Palaeoclimatology, Palaeoecology, 226 (1), 167–185.
- 1001 Edwards, R. J., & Horton, B. P. (2000). Reconstructing relative sea-level change using UK
- salt-marsh foraminifera. *Marine Geology*, 169(1), 41–56.
- 1003 Edwards, R. J., Wright, A. J., & Van de Plassche, O. (2004a). Surface distributions of salt-
- marsh foraminifera from Connecticut, USA: modern analogues for high-resolution sea level
- studies. *Marine Micropaleontology*, *51*(1), 1–21.
- 1006 Edwards, R. J., Van De Plassche, O., Gehrels, W. R., & Wright, A. J. (2004b). Assessing sea-
- 1007 level data from Connecticut, USA, using a foraminiferal transfer function for tide
- level. Marine Micropaleontology, 51(3), 239–255.
- Engelhart, S. E., Horton, B. P., & Kemp, A. C. (2011). Holocene sea-level changes along the
- 1010 United States' Atlantic coast. Oceanography, 24(2), 70–79,
- Engelhart, S. E., Horton, B. P., Vane, C. H., Nelson, A. R., Witter, R. C., Brody, S. R., &
- 1012 Hawkes, A. D. (2013). Modern foraminifera, δ 13 C, and bulk geochemistry of central
- 1013 Oregon tidal marshes and their application in paleoseismology. *Palaeogeography*,
- 1014 Palaeoclimatology, Palaeoecology, 377, 13–27.
- 1015 Engels, J. G., Rink, F., & Jensen, K. (2011). Stress tolerance and biotic interactions determine
- plant zonation patterns in estuarine marshes during seedling emergence and early
- establishment. *Journal of Ecology*, 99(1), 277–287.

- 1018 Franceschini, G., Mc Millan, I. K., & Compton, J. S. (2005). Foraminifera of Langebaan
- Lagoon salt marsh and their application to the interpretation of late Pleistocene depositional
- environments at Monwabisi, False Bay coast, South Africa. South African Journal of
- 1021 Geology, 108(2), 285–296.
- Frezza, V., & Carboni, M. G. (2009). Distribution of recent foraminiferal assemblages near
- the Ombrone River mouth (Northern Tyrrhenian Sea, Italy). Revue de
- 1024 *micropaléontologie*, *52*(1), 43–66.
- 1025 Gedan, K. B., Kirwan, M. L., Wolanski, E., Barbier, E. B., & Silliman, B. R. (2010). The
- present and future role of coastal wetland vegetation in protecting shorelines: Answering
- recent challenges to the paradigm. *Climatic Change 106*, 7–29.
- 1028 Gehrels, W. R. (1994). Determining relative sea-level change from salt-marsh foraminifera
- and plant zones on the coast of Maine, USA. *Journal of Coastal Research*, 10, 990–1009.
- Gehrels, W. R. (1999). Middle and late Holocene sea-level changes in eastern Maine
- reconstructed from foraminiferal saltmarsh stratigraphy and AMS 14 C dates on basal
- 1032 peat. *Quaternary Research*, 52(3), 350–359.
- Gehrels, W. R., & van de Plassche, O. (1999). The use of *Jadammina macrescens* (Brady)
- and Balticammina pseudomacrescens Brönnimann, Lutze and Whittaker (Protozoa:
- 1035 Foraminiferida) as sea-level indicators. *Palaeogeography*, *Palaeoclimatology*,
- 1036 *Palaeoecology*, 149(1), 89–101.
- 1037 Gehrels, W. R. (2000). Using foraminiferal transfer functions to produce high-resolution sea-
- level records from salt-marsh deposits, Maine, USA. *The Holocene*, 10(3), 367–376.
- 1039 Gehrels, W. R., Roe, H. M., & Charman, D. J. (2001). Foraminifera, testate amoebae and
- 1040 diatoms as sea-level indicators in UK saltmarshes: a quantitative multiproxy
- approach. Journal of Quaternary Science, 16(3), 201–220.
- 1042 Gehrels, W.R., Belknap, D.F., Black, S., & Newnham, R.M. (2002). Rapid sea-level rise in
- the Gulf of Maine, USA, since AD 1800. *The Holocene*, *12*, 383–389.
- Gehrels, W.R., Milne, G.A., Kirby, J.R., Patterson, R.T., & Belknap, D.F. (2004). Late
- Holocene sea-level changes and isostatic crustal movements in Atlantic Canada. Quaternary
- 1046 International, 120, 79–89.
- Gehrels, W. R., Kirby, J. R., Prokoph, A., Newnham, R. M., Achterberg, E. P., Evans, H.,
- 1048 Black, S., & Scott, D. B. (2005). Onset of recent rapid sea-level rise in the western Atlantic
- 1049 Ocean. *Quaternary Science Reviews*, 24(18), 2083–2100.
- 1050 Gehrels, W. R., Marshall, W. A., Gehrels, M. J., Larsen, G., Kirby, J. R., Eiríksson, J.,
- Heinemeier, J., & Shimmield, T. (2006). Rapid sea-level rise in the North Atlantic Ocean
- since the first half of the nineteenth century. *The Holocene*, 16(7), 949–965.
- 1053 Gehrels, W. R., Hayward, B. W., Newnham, R. M., & Southall, K. E. (2008). A 20th century
- acceleration of sea-level rise in New Zealand. Geophysical Research Letters, 35(2) 1–5,
- **1055** L02717.

- Gehrels, W. R., Callard, S. L., Moss, P. T., Marshall, W. A., Blaauw, M., Hunter, J., Milton,
- J.A., & Garnett, M. H. (2012). Nineteenth and twentieth century sea-level changes in
- 1058 Tasmania and New Zealand. Earth and Planetary Science Letters, 315, 94–102.
- 1059 Gehrels, W. R., & Woodworth, P. L. (2013). When did modern rates of sea-level rise
- start? Global and Planetary Change, 100, 263–277.
- Goldstein, S. T., Watkins, G. T., & Kuhn, R. M. (1995). Microhabitats of salt marsh
- foraminifera: St. Catherines Island, Georgia, USA. *Marine Micropaleontology*, 26(1), 17–29.
- 1063 Goldstein, S. T., & Watkins, G. T. (1998). Elevation and the distribution of salt-marsh
- Foraminifera, St. Catherines Island, Georgia; a taphonomic approach. *Palaios*, 13(6), 570–
- 1065 580.
- Goldstein, S. T., & Watkins, G. T. (1999). Taphonomy of salt marsh foraminifera: an example
- from coastal Georgia. Palaeogeography, Palaeoclimatology, Palaeoecology, 149(1), 103–
- 1068 114.
- Grenfell, H.R., Hayward, B.W., Nomura, R., & Sabaa, A.T. (2012). A foraminiferal proxy
- record of 20th century sea-level rise in the Manukau Harbour, New Zealand. *Marine and*
- 1071 Freshwater Research, 63, 370-384.
- 1072 Grossinger, R. Stein, E.D. Cayce, K. Askevold, R. Dark, S., & Whipple, A. (2011). Historical
- 1073 Wetlands of the Southern California Coast: An Atlas of US Coast Survey T-sheets, 1851-
- 1074 1889. Technical Report 589. Southern California Coastal Water Research Project, Costa
- 1075 Mesa, CA and San Francisco Estuary Institute, Oakland, CA.
- 1076 Guilbault, J. P., Clague, J. J., & Lapointe, M. (1996). Foraminiferal evidence for the amount
- of coseismic subsidence during a late Holocene earthquake on Vancouver Island, west coast
- of Canada. Quaternary Science Reviews, 15(8), 913–937.
- Hammer, O., Harper, D.A.T., & Ryan, P.D. (2001). PAST: Palaeontological statistics package
- for education and data analysis. *Palaeontol. Electron.* 4, 1–9 (online).
- Hawkes, A. D., Horton, B. P., Nelson, A. R., & Hill, D. F. (2010). The application of
- 1082 intertidal foraminifera to reconstruct coastal subsidence during the giant Cascadia earthquake
- 1083 of AD 1700 in Oregon, USA. *Quaternary International*, 221(1), 116–140.
- 1084 Hawkes, A. D., Horton, B. P., Nelson, A. R., Vane, C. H., & Sawai, Y. (2011). Coastal
- subsidence in Oregon, USA, during the giant Cascadia earthquake of AD 1700. Quaternary
- 1086 Science Reviews, 30, 364–376.
- Hayward, B. W., Grenfell, H. R., Nicholson, K., Parker, R., Wilmhurst, J., Horrocks, M.,
- 1088 Swales, A., & Sabaa, A. T. (2004). Foraminiferal record of human impact on intertidal
- estuarine environments in New Zealand's largest city. Marine Micropaleontology, 53(1), 37–
- 1090 66.
- Hippensteel, S. P., Martin, R. E., Nikitina, D., & Pizzuto, J. E. (2002). Interannual variation of
- marsh foraminiferal assemblages (Bombay Hook National Wildlife Refuge, Smyrna, DE): Do
- 1093 foraminiferal assemblages have a memory? The Journal of Foraminiferal Research, 32(2),
- 1094 97–109.

- Horton, B. P. (1997). Quantification of the indicative meaning of a range of Holocene sea-
- level index points from the western North Sea (Doctoral dissertation, University of Durham).
- Horton, B. P. (1999). The distribution of contemporary intertidal foraminifera at Cowpen
- Marsh, Tees Estuary, UK: implications for studies of Holocene sea-level
- changes. Palaeogeography, Palaeoclimatology, Palaeoecology, 149(1), 127–149.
- Horton, B. P., Edwards, R. J., & Lloyd, J. M. (1999). A foraminiferal-based transfer function:
- implications for sea-level studies. *The Journal of Foraminiferal Research*, 29(2), 117–129.
- Horton, B. P., & Edwards, R. J. (2003). Seasonal distributions of foraminifera and their
- implications for sea-level studies: SEPM (Society for Sedimentary Geology) Special
- 1104 Publication no. 75, p. 21–30.
- Horton, B. P., Larcombe, P., Woodroffe, S. A., Whittaker, J. E., Wright, M. R., & Wynn, C.
- 1106 (2003). Contemporary foraminiferal distributions of a mangrove environment, Great Barrier
- 1107 Reef coastline, Australia: implications for sea-level reconstructions. *Marine Geology*, 198(3),
- 1108 225–243.
- Horton, B. P., & Edwards, R. J. (2005). The application of local and regional transfer
- functions to the reconstruction of Holocene sea levels, north Norfolk, England. *The*
- 1111 *Holocene*, 15(2), 216–228.
- Horton, B. P., & Edwards, R. J. (2006). Quantifying Holocene sea level change using
- intertidal foraminifera: lessons from the British Isles. Cushman Foundation for Foraminiferal
- 1114 Research Special Publication 40.
- Horton, B. P., & Murray, J. W. (2006). Patterns in cumulative increase in live and dead
- species from foraminiferal time series of Cowpen Marsh, Tees Estuary, UK: Implications for
- sea-level studies. *Marine Micropaleontology*, 58(4), 287–315.
- Horton, B. P., & Murray, J. W. (2007). The roles of elevation and salinity as primary controls
- on living foraminiferal distributions: Cowpen Marsh, Tees Estuary, UK. Marine
- 1120 *Micropaleontology*, *63*(3), 169–186.
- Horton, B. P., Corbett, R., Culver, S. J., Edwards, R. J., & Hillier, C. (2006). Modern
- saltmarsh diatom distributions of the Outer Banks, North Carolina, and the development of a
- 1123 transfer function for high resolution reconstructions of sea level. Estuarine, Coastal and Shelf
- 1124 Science, 69(3), 381–394.
- 1125 Jennings, A. E., & Nelson, A. R. (1992). Foraminiferal assemblage zones in Oregon tidal
- 1126 marshes; relation to marsh floral zones and sea level. The Journal of Foraminiferal
- 1127 Research, 22(1), 13–29.
- Jones, G. D., & Ross, C. A. (1979). Seasonal distribution of foraminifera in Samish Bay,
- 1129 Washington. *Journal of Paleontology*, *53*, 245–257.
- 1130 Juggins, S. (2011). C2 Data Analysis Version 1.7. 2. Newcastle upon Tyne: University of
- 1131 Newcastle.

- 1132 Juggins, S., & Birks, H. J. B. (2012). Quantitative environmental reconstructions from
- biological data. In *Tracking environmental change using lake sediments* (pp. 431–494).
- 1134 Springer Netherlands.
- Kaminski, M.A., & Filipescu, S., (eds), (2011). Proceedings of the Eighth International
- 1136 Workshop on Agglutinated Foraminifera. Grzybowski Foundation Special Publication, 16,
- **1137** 29–35.
- 1138 Kemp, A. C., Horton, B. P., & Culver, S. J. (2009a). Distribution of modern salt-marsh
- 1139 foraminifera in the Albemarle–Pamlico estuarine system of North Carolina, USA:
- implications for sea-level research. *Marine Micropaleontology*, 72(3), 222–238.
- 1141 Kemp, A. C., Horton, B. P., Corbett, D. R., Culver, S. J., Edwards, R. J., & van de Plassche,
- 1142 O. (2009b). The relative utility of foraminifera and diatoms for reconstructing late Holocene
- sea-level change in North Carolina, USA. *Quaternary Research*, 71(1), 9–21.
- 1144 Kemp, A. C., Horton, B. P., Donnelly, J. P., Mann, M. E., Vermeer, M., & Rahmstorf, S.
- 1145 (2011). Climate related sea-level variations over the past two millennia. *Proceedings of the*
- 1146 National Academy of Sciences, 108(27), 11017–11022.
- 1147 Kemp, A.C., Horton, B.P., Vane, C.H., Bernhardt, C.E., Corbett, D.R., Engelhart, S.E.,
- Anisfeld, S.C., Parnell, A.C., & Cahill, N. (2013). Sea-level change during the last 2500 years
- in New Jersey, USA. Quaternary Science Reviews, 81, 90–104.
- 1150 Kemp, A. C., & Telford, R. J. (2015). Transfer functions. *Handbook of Sea-Level Research:*
- 1151 *John Wiley and Sons, Chichester*, 470–499.
- Kirby, M.E., Knell, E.J., Anderson, W.T., Lachniet, M.S., Palermo, J., Eeg, H., Lucero, R.,
- Murrieta, R., Arevalo, A., Silveira, E., & Hiner, C.A. (2015). Evidence for insolation and
- Pacific forcing of late glacial through Holocene climate in the Central Mojave Desert (Silver
- 1155 Lake, CA). Quaternary Research, 84, 174–186.
- 1156 Kirwan, M. L., & Megonigal, J. P. (2013). Tidal wetland stability in the face of human
- impacts and sea-level rise. *Nature*, *504*(7478), 53–60.
- 1158 Kopp, R. E., Simons, F. J., Mitrovica, J. X., Maloof, A. C., & Oppenheimer, M. (2013). A
- probabilistic assessment of sea level variations within the last interglacial stage. *Geophysical*
- 1160 *Journal International*, 193(2), 711–716.
- 1161 Leeper, R. J. (2015). Abrupt subsidence in the Seal Beach Wetlands, southern
- 1162 California (Doctoral dissertation, CALIFORNIA STATE UNIVERSITY, FULLERTON).
- Legendre, P., & Fortin, M. J. (1989). Spatial pattern and ecological analysis. *Vegetatio*, 80(2),
- 1164 107–138.
- Leorri, E., & Martin, R. E. (2009). The input of foraminiferal infaunal populations to sub-
- fossil assemblages along an elevational gradient in a salt marsh: application to sea-level
- studies in the mid-Atlantic coast of North America. Hydrobiologia, 625(1), 69–81.
- Leorri, E., Gehrels, W. R., Horton, B. P., Fatela, F., & Cearreta, A. (2010). Distribution of
- 1169 foraminifera in salt marshes along the Atlantic coast of SW Europe: Tools to reconstruct past
- sea-level variations. *Quaternary International*, 221(1), 104–115.

- 1171 Lepš, J., & Šmilauer, P. (2003). Multivariate analysis of ecological data using CANOCO.
- 1172 Cambridge university press.
- 1173 Leyer, I., & Wesche, K. (2007). Multivariate Statistik in der Ökologie: Eine Einführung.
- 1174 Springer-Verlag.
- 1175 Long, A. J., Barlow, N. L. M., Gehrels, W. R., Saher, M. H., Woodworth, P. L., Scaife, R. G.,
- Brain, M.J., & Cahill, N. (2014). Contrasting records of sea-level change in the eastern and
- western North Atlantic during the last 300 years. Earth and Planetary Science Letters, 388,
- **1178** 110–122.
- 1179 Martin, R.E., 1999. Taphonomy and temporal resolution of foraminiferal assemblages. In:
- 1180 Sen Gupta, B.K. (Ed.), Modern Foraminifera. Kluwer Academic Publishers, Dordrecht, pp.
- **1181** 281–298.
- 1182 Matera, N. J., & Lee, J. J. (1972). Environmental factors affecting the standing crop of
- foraminifera in sublittoral and psammolittoral communities of a Long Island salt
- 1184 marsh. *Marine Biology*, 14(2), 89–103.
- Milker, Y., Horton, B. P., Nelson, A. R., Engelhart, S. E., & Witter, R. C. (2015a). Variability
- of intertidal foraminiferal assemblages in a salt marsh, Oregon, USA. Marine
- **1187** *Micropaleontology*, *118*, 1–16.
- Milker, Y., Horton, B. P., Vane, C. H., Engelhart, S. E., Nelson, A. R., Witter, R. C., Khan,
- 1189 N. S., & Bridgeland, W. T. (2015b). Annual and seasonal distribution of intertidal
- foraminifera and stable carbon isotope geochemistry, Bandon Marsh, Oregon, USA. *The*
- 1191 *Journal of Foraminiferal Research*, 45(2), 146–155.
- Milne, G. A., Gehrels, W. R., Hughes, C. W., & Tamisiea, M. E. (2009). Identifying the
- causes of sea-level change. *Nature Geoscience*, 2(7), 471–478.
- Mills, H., Kirby, J., Holgate, S., & Plater, A. (2013). The Distribution Of Contemporary
- 1196 Saltmarsh Foraminifera In A Macrotidal Estuary: An Assessment Of Their Viability For Sea-
- 1197 Level Studies. Journal of Ecosystems and Ecography, 3(3), 1–16 doi:10.4172/2157-
- 1198 7625.1000131

- 1199 Müller-Navarra, K., Milker, Y., & Schmiedl, G. (2016). Natural and anthropogenic influence
- on the distribution of salt marsh foraminifera in the Bay of Tümlau. *Journal of Foraminiferal*
- 1201 *Research*, 46(1), 61–74.
- Murray, J.W. (1971). Living foraminiferids of tidal marshes: a review. *Journal of*
- 1203 Foraminiferal Research 1, 156–161.
- 1204 Murray, J. W., & Alve, E. (1999). Natural dissolution of modern shallow water benthic
- foraminifera: taphonomic effects on the palaeoecological record. *Palaeogeography*,
- 1206 Palaeoclimatology, Palaeoecology, 146(1), 195–209.
- Murray, J. W., & Bowser, S. S. (2000). Mortality, protoplasm decay rate, and reliability of
- 1208 staining techniques to recognize 'living' foraminifera: a review. The Journal of Foraminiferal
- **1209** *Research*, *30*(1), 66–70.

- 1210 Murray, J. W. (2006). Ecology and applications of benthic foraminifera. Cambridge
- 1211 University Press.
- 1212 Nelson, A. R., Sawai, Y., Jennings, A. E., Bradley, L. A., Gerson, L., Sherrod, B. L., Sabean,
- 1213 J., & Horton, B. P. (2008). Great-earthquake paleogeodesy and tsunamis of the past 2000
- years at Alsea Bay, central Oregon coast, USA. *Quaternary Science Reviews*, 27(7), 747–768.
- 1215 Ozarko, D.L., Patterson, R.T., & Williams, H.F.L. (1997). Marsh foraminifera from Nanaimo,
- 1216 British Columbia: infaunal habitat and taphonomic implications. *Journal of Foraminiferal*
- 1217 Research, 27, 51–68.
- 1218 Patterson, R. T. (1990). Intertidal benthic foraminiferal biofacies on the Fraser River Delta,
- 1219 British Columbia: modern distribution and paleoecological importance. *Micropaleontology*,
- **1220** *36*, 183–199...
- Patterson, R. T., Guilbault, J. P., & Clague, J. J. (1999). Taphonomy of tidal marsh
- foraminifera: implications of surface sample thickness for high-resolution sea-level
- studies. *Palaeogeography, Palaeoclimatology, Palaeoecology, 149*(1), 199–211.
- Patterson, R. T., Gehrels, W. R., Belknap, D. F., & Dalby, A. P. (2004). The distribution of
- salt marsh foraminifera at Little Dipper Harbour New Brunswick, Canada: implications for
- development of widely applicable transfer functions in sea-level research. Quaternary
- 1227 International, 120(1), 185–194.
- Peinado, M., Alcaraz, F., Delgadillo, J., De La Cruz, M., Alvarez, J., & Aguirre, J. L. (1994).
- 1229 The coastal salt marshes of California and Baja California. *Vegetatio*, 110(1), 55–66.
- 1230 Phipps, M. D., Kaminski, M. A., & Aksu, A. E. (2010). Calcareous benthic foraminiferal
- biofacies along a depth transect on the southwestern Marmara shelf
- **1232** (Turkey). *Micropaleontology*, *56*, 377–392.
- Rossi, V., Horton, B. P., Corbett, D. R., Leorri, E., Perez-Belmonte, L., & Douglas, B. C.
- 1234 (2011). The application of foraminifera to reconstruct the rate of 20th century sea level rise,
- 1235 Morbihan Golfe, Brittany, France. *Quaternary Research*, 75(1), 24–35.
- 1236 Saffert, H., & Thomas, E. (1998). Living foraminifera and total populations in salt marsh peat
- 1237 cores: Kelsey Marsh (Clinton, CT) and the Great Marshes (Barnstable, MA). Marine
- **1238** *Micropaleontology*, *33*(3), 175–202.
- 1239 Saher, M.H., Gehrels, W.R., Barlow, N.L.M., Long, A.J., Haigh, I.D., & Blaauw, M. (2015).
- 1240 A 600 year multiproxy record of sea-level change and the influence of the North Atlantic
- 1241 Oscillation. Quaternary Science Reviews, 108, 23–36.
- 1242 Schönfeld, J., Alve, E., Geslin, E., Jorissen, F., Korsun, S., & Spezzaferri, S. (2012). The
- 1243 FOBIMO (FOraminiferal BIo-MOnitoring) initiative—Towards a standardised protocol for
- soft-bottom benthic foraminiferal monitoring studies. *Marine Micropaleontology*, 94, 1–13.
- 1245 Scott, D. B. (1976). Brackish-water foraminifera from southern California and description of
- 1246 Polysaccammina ipohalina n. gen., n. sp. The Journal of Foraminiferal Research, 6(4), 312–
- **1247** 321.

- 1248 Scott, D. S., & Medioli, F. S. (1978). Vertical zonations of marsh foraminifera as accurate
- indicators of former sea-levels. *Nature* 272, 528–531.
- 1250 Scott, D. B., & Medioli, F. S. (1980). Quantitative studies of marsh foraminiferal distributions
- 1251 in Nova Scotia; implications for sea level studies. Special Publications-Cushman Foundation
- 1252 for Foraminiferal Research.
- 1253 Scott, D. B., Medioli, F. S., & Duffett, T. E. (1984). Holocene rise of relative sea level at
- Sable Island, Nova Scotia, Canada. *Geology*, 12(3), 173–176.
- 1255 Scott, D. B., & Hermelin, J. O. R. (1993). A device for precision splitting of
- micropaleontological samples in liquid suspension. *Journal of Paleontology*, 67(01), 151–
- **1257** 154.
- 1258 Scott, G., Thompson, L., Hitchin, R., & Scourse, J. (1998). Observations on selected salt-
- marsh and shallow-marine species of agglutinated foraminifera: grain size and mineralogical
- selectivity. Journal of *Foraminiferal Research*, 28, 261–267.
- 1261 Scott, D. B., Medioli, F. S., & Schafer, C. T. (2001). Monitoring in coastal environments
- *using foraminifera and thecamoebian indicators.* Cambridge University Press.
- 1263 Scott, D. B., Mudie, P.J., & Bradshaw, J.S. (2011). Coastal evolution of Southern
- 1264 California as interpreted from benthic foraminifera, ostracodes, and pollen. *Journal of*
- 1265 Foraminiferal Research, 41, 285–307.
- 1266
- Shaw, T.A., Kirby, J.R., Holgate, S., Tutman, P., & Plater, A.J. (2016). Contemporary salt-
- 1268 marsh foraminifera distribution from the Adriatic Coast of Croatia and its potential for Sea-
- 1269 Levels studies. *Journal of Foraminiferal Research*, 46 (3) 314–332.
- 1270 Southall, K. E., Gehrels, W. R., & Hayward, B. W. (2006). Foraminifera in a New Zealand
- salt marsh and their suitability as sea-level indicators. Marine Micropaleontology, 60(2), 167–
- **1272** 179.
- 1273 Stone, M., & Brooks, R. J. (1990). Continuum regression: cross-validated sequentially
- 1274 constructed prediction embracing ordinary least squares, partial least squares and principal
- 1275 components regression. *Journal of the Royal Statistical Society*, 52, 237–269.
- 1276 Strachan, K.L., Finch, J.M., Hill, T., & Barnett, R.L. (2014). A late Holocene sea-level curve
- for the east coast of South Africa. South African Journal of Science, 110 (1/2), art. # 2013—
- 1278 0198, 9 pp.
- 1279 Takekawa, J. Y., Thorne, K. M., Buffington, K. J., Freeman, C. M., & Block, G. (2013).
- 1280 Evaluation of subterranean subsidence at Seal Beach National Wildlife Refuge. U.S.
- 1281 Geological Survey, Western Ecological Research Center, Vallejo, CA.
- Telford, R. J., Heegaard, E., & Birks, H. J. B. (2004). All age-depth models are wrong: but
- 1283 how badly? *Quaternary Science Reviews*, 23(1), 1–5.
- 1284 Telford, R. J., & Birks, H. J. B. (2009). Evaluation of transfer functions in spatially structured
- environments. Quaternary Science Reviews, 28(13), 1309–1316.

- Telford, R.J., & Birks, H.J.B. (2011). A novel method for assessing the statistical significance
- of quantitative reconstructions inferred from biotic assemblages. Quaternary Science
- 1288 Reviews, 30, 1272–1278.
- 1289 Ter Braak, C. J., & Juggins, S. (1993). Weighted averaging partial least squares regression
- 1290 (WA-PLS): an improved method for reconstructing environmental variables from species
- **1291** assemblages. *Hydrobiologia*, 269(1), 485–502.
- 1292 Ter Braak, C. J., & Smilauer, P. (2002). CANOCO reference manual and CanoDraw for
- 1293 Windows user's guide: software for canonical community ordination (version 4.5).
- ter Braak, C.J.F., Juggins, S., Birks, H.J.B., & van de Voet, H. (1993). Weighted averaging
- 1295 partial least squared regression (WA-PLS): definition and comparison with other methods for
- 1296 species-environmental calibration. In: Patil, G.P., Rac, C.R. (Eds.), Multivariate
- Environmental Statistics. Elsevier Science Publishers, Amsterdam, pp. 525–560.
- Tobin, R., Scott, D.B., Collins, E.S., & Medioli, F.S. (2005). Infaunal benthic foraminifera in
- some North American marshes and their influence on fossil assemblages. *Journal of*
- 1300 Foraminiferal Research, 35, 130–147.
- Wallace, K., Callaway, J. C., & Zedler, J. (2005). Evolution of tidal creek networks in a high
- 1302 sedimentation environment: a 5-year experiment at Tijuana Estuary, California. Estuaries and
- 1303 Coasts, 28(6), 795–811. Retrieved from
- http://www.springerlink.com/index/7H2P240120801675.pdf
- 1305 Walker, S. E., & Goldstein, S. T. (1999). Taphonomic tiering: experimental field taphonomy
- of molluses and foraminifera above and below the sediment–water
- interface. Palaeogeography, Palaeoclimatology, Palaeoecology, 149(1), 227–244.
- 1308 Walton, W. R. (1952). Techniques for recognition of living foraminifera. Cushman
- 1309 Foundation for Foraminiferal Research, 3, 56–60.
- 1310 Wang, P., & Chappell, J. (2001). Foraminifera as Holocene environmental indicators in the
- 1311 South Alligator River, northern Australia. *Quaternary International*, 83, 47–62.
- Woodroffe, S.A. (2009). Testing models of mid to late Holocene sea-level change, North
- 1313 Queensland, Australia. *Quaternary Science Reviews* 28, 2474–2488.
- 1314 Woodroffe, S. A., & Horton, B. P. (2005). Holocene sea-level changes in the Indo-
- Pacific. *Journal of Asian Earth Sciences*, 25(1), 29–43.
- Wright, A. J., Edwards, R. J., & van de Plassche, O. (2011). Reassessing transfer-function
- performance in sea-level reconstruction based on benthic salt-marsh foraminifera from the
- 1318 Atlantic coast of NE North America. *Marine Micropaleontology*, 81(1), 43–62.
- 1319 Zedler, J. B. (1977). Salt marsh community structure in the Tijuana Estuary, California.
- 1320 Estuarine and Coastal Marine Science, 5(1), 39–53. doi:10.1016/0302-3524(77)90072-X
- 1321 Zedler, J. B. (1982). The ecology of southern Californiia coastal salt marshes: a community
- profile. Washington, D.C. FWS/OBS-81/54.

- 1323 Zedler, J. B. (2010). How frequent storms affect wetland vegetation: A preview of climate-
- change impacts. Frontiers in Ecology and the Environment, 8(10), 540–547.
- 1325 doi:10.1890/090109
- Zedler, J. B., Covin, J., Nordby, C., Williams, P., & Boland, J. (1986). Catastrophic events
- reveal the dynamic nature of salt-marsh vegetation in Southern California. *Estuaries*, 9(1),
- 1328 75–80. doi:10.1007/BF02689746
- Zedler, J. B., Nordby, C. S., & Kus, B. E. (1986). The Ecology of Tijuana Estuary, California.
- 1330 USGS, Estuarine Profile.

Table captions

13321333

- Table 1. The nearest tide gauge station for each site with observations of Mean Higher
- High Water (MHHW), Mean Tide Level (MTL) and Mean Sea-Level (MSL). Tidal
- water level heights (m), above North American vertical datum (NAVD88), are
- measured over the period from 1983 to 2001 at the National Oceanic and Atmospheric
- 1338 Administration (NOAA).

1339

- Table 2. Sample locations, elevation, pore-water and sediment properties of the
- surface marsh samples collected at Seal Beach (denoted with the initials SB) and
- 1342 Tijuana (denoted with the initials TJE) during spring tide on mid-October, and early
- December 2015, respectively. TOM (total organic matter) and carbonate content were
- determined by Loss-on-Ignition at 550 and 950 °C, respectively. Detailed
- granulometric data detailed in Supplementary data No. 1 and 2. (nd = no data; asterisk
- denote stations in which a mini core was collected in addition to the surface sample).

1347

- Table 3. Statistical results of Detrended Correspondence Analysis (DCA) for Seal
- Beach (a), Tijuana (b) and the combined surface assemblage of both marshes (c).

1350

- Table 4. Statistical results of Redundancy Analysis (RDA) for Seal Beach (a), Tijuana
- 1352 (b) and the combined surface assemblage of both marshes (c).

1353

- Table 5. Results of the transfer functions performance criteria for foraminifera
- training set using two regression models: Partial Least Squares and Weighted
- 1356 Averaging-Partial Least Squares (see text for details). Given are the cross-validated
- (jack-knifed) correlation (r^2) between observed and estimated elevation in the modern
- data, the mean and maximum bias, the Root Mean Squared Error of Prediction
- 1359 (RMSEP) and the change of the RMSEP (in %) from one component to the next.

1360

1361

Figure captions

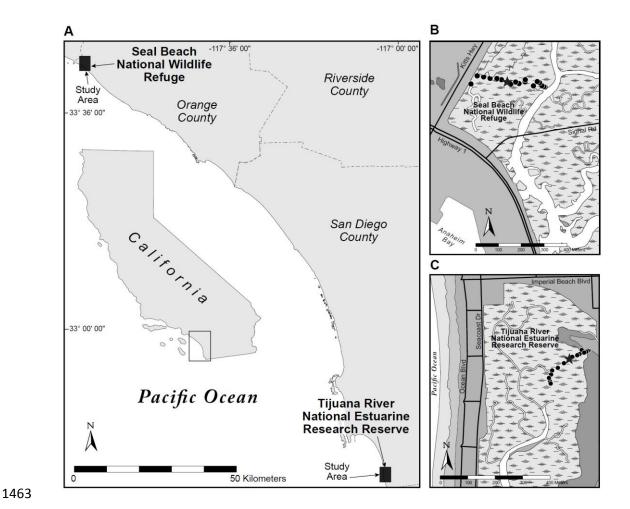
- Fig. 1. Location of the study area on the North American Pacific coast, southern
- California (A), and sites mentioned in the text including Seal Beach salt-marsh (B)
- and Tijuana salt-marsh (C) with transects highlighted showing the surface station

- locations (asterisk denote stations in which a mini core was collected in addition to
- the surface sample). See also Tables 1 and 2.
- Fig. 2. Environmental variables from sampled transects in Seal Beach and Tijuana salt
- marshes including: temperature, pH, O₂, salinity, organic matter, carbonate content,
- sand, silt and clay content and vertical profile relative to the North American vertical
- datum (NAVD88). Sample elevation profiles with the tidal datum for MHHW (mean
- highest high water). See also Tables 1 and 2.
- Fig. 3. Absolute abundances of the most abundant live (rose Bengal stained) (A) and
- dead foraminifera B) and their counts in the Seal Beach salt-marsh surface
- samples collected during mid-and late October, 2015. Sample elevation profiles are
- 1377 shown.
- Fig. 4. Absolute abundances of the most abundant live (rose Bengal stained) (A) and
- dead foraminifera (B) and their counts in the Tijuana salt- marsh surface samples
- collected during December, 2015. Sample elevation profiles are shown
- Fig. 5. Dendrogram of Q-mode cluster analysis of the dead foraminifera from Seal
- Beach salt-marsh. The most significant species contributing to each cluster, based on
- the 'similarity percentages' (SIMPER) routine, are shown on top of each cluster (J.m.
- 1384 J. macrescens; A.spp. Ammobaculites spp.; T.in. T. inflata; Cal. Calcareous
- species; T.ir. T. irregularis; M.f. M. fusca; S.m S. moniliformis. The percentages
- of the most common dead species are given below the dendrogram.
- Fig. 6. Dendrogram of Q-mode cluster analysis of the dead foraminifera from Tijuana
- salt-marsh. Taxa that make significant contributions to the similarity within each
- cluster, based on SIMPER routine, are shown on top of each cluster (abbreviations are
- given in Fig. 5) and the percentages of the most common dead species are given
- below the dendrogram.
- Fig. 7. Total and individual live (rose Bengal stained) foraminiferal numbers (per
- 1393 10cm³ sediment volume), and total percentages of live (rose Bengal stained)
- specimens relative to the total populations in each mini core (10cm deep) taken from
- the middle marshes of Seal Beach and Tijuana, St. 5 and St. 8, respectively.
- Fig. 8. Redundancy Analysis (RDA) ordination triplots with forward selected
- significant environmental variables (Table 4) showing retrospective projection of the
- surface samples—species—environmental variables for Seal Beach (A) and Tijuana
- 1399 (B). Clusters sample distinguished in the Q-mode CA (Figs. 5 and 6) were
- incorporated into the RDA results.
- 1401 Fig. 9. Redundancy Analysis (RDA) ordination triplot with forward selected
- significant environmental variables (Table 4) showing retrospective projection of the
- surface samples—species—environmental variables for the combined data sets of
- Seal Beach and Tijuana. Clusters sample distinguished in the Q-mode CA (Figs. 5 and
- 1405 6) were incorporated into the RDA results.

- Fig. 10. Scatterplots showing the relationship between observed standardized water
- level index (SWLI) against model predicted SWLI and residuals versus observed

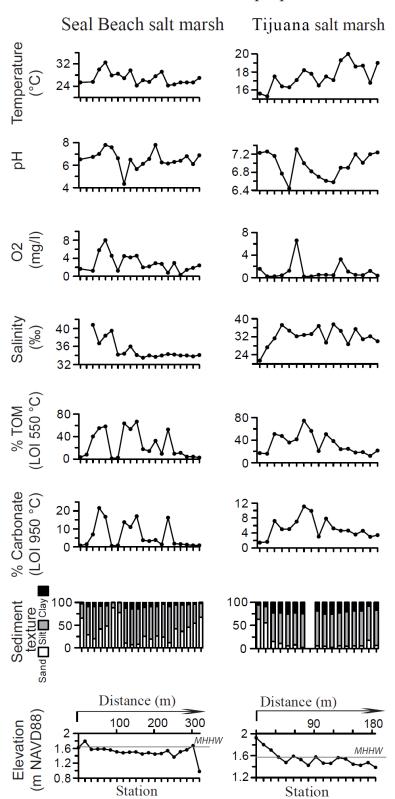
1409 1410 1411	SWLI using WA-PLS (component 3) transfer function in the modern data set derived from the southern California salt marshes.
1412	
1413	Plate captions
1414	
1415	Plate 1. Agglutinant foraminifera from surface sediments collected in Seal Beach and
1416	Tijuana salt marches, along the coast of southern California.
1417	The scale bars of Figs. 2c, 3b equal 20 μm, of Figs. 2a-b equal 50 μm, of Figs. 1a-c,
1418	4a-b, 5b, 6a-c, 7c, 8, 9c, 10b-e equal 100 μm, of Figs. 3a, 5a, 7a-b, 9a-b, 10a equal
1419	200 μm.
1420	
1421	(1a-b) Miliammina fusca (Brady, 1870), side view. (1c) M. fusca, aperture view. (2a-
1422	b) Miliammina petila Saunders, 1958, side view. (2c) M. petila, aperture view. (3a)
1423	Scherochorella moniliformis (Siddall, 1886), side view. (3b) S. moniliformis, aperture
1424	view. (4a) Trochamminita irregularis Cushman & Brönnimann, 1948, side view. (4b)
1425	T. irregularis, side and aperture view. (5a) Ammobaculites dilatatus Cushman &
1426	Brönnimann, 1948, side view. (5b) A. dilatatus, aperture view. (6a-b) Ammobaculites
1427	sp., side view. (6c) Ammobaculites sp., aperture view. (7a-b) Ammobaculites spp.,
1428	side view. (7c) Ammobaculites spp., aperture view. (8) Ammobaculites spp., side
1429	view. (9a) Trochammina inflata (Montagu, 1808), spiral view. (9b) T. inflata,
1430	umbilical view. (9c) T. inflata, aperture view. (10a-b) Jadammina macrescens (Brady,
1431	1870), spiral view. (10c) J. macrescens, umbilical view. (10d) J. macrescens,
1432	umbilical view showing collapsed chambers occurred soon after the specimen was
1433	dried. (10e) J. macrescens, aperture view with secondary apertures.
1434	
1435	
1436	Plate 2. Calcareous intertidal taxa from surface sediments collected in Seal Beach and
1437	Tijuana salt marches, along the coast of southern California.
1438	The scale bars of Figs. 1, 3 and 5c equal 50 µm, of Figs. 2, 4a-b, 5a-b and 6a-b equal
1439	100 μm.
1440	·
1441	(1) Cornuspira sp., side view. (2) Quinqueloculina sp. side view. (3) miliolid, side
1442	view. (4a) <i>Trichohyalus aguayoi</i> (Bermúdez), spiral view. (4b) <i>T. aguayoi</i> , umbilical
1443	view. (5a) Ammonia sp. spiral view. (5b) Ammonia sp. umbilical view. (5c) Ammonia
1444	sp. aperture view. (6a) <i>Elphidium</i> sp side view. (6b) <i>Elphidium</i> sp. aperture view.
1445	$\mathbf{r}_{\mathbf{r}}$ $\mathbf{r}_{\mathbf{r}}$ $\mathbf{r}_{\mathbf{r}}$ $\mathbf{r}_{\mathbf{r}}$ $\mathbf{r}_{\mathbf{r}}$ $\mathbf{r}_{\mathbf{r}}$
1446	
1447	
1448	Supplementary data captions
1449	
1450	Supplementary data 1. Seal Beach salt-marsh: general data and census foraminiferal
1451	data.
1452	
1453	Supplementary data 2. Tijuana salt-marsh: general data and census foraminiferal data.
1454	<u> </u>

1455	Supplementary data 3. Taxonomic reference list of species presented in text and in
1456	Supplementary data 1 and 2.
1457	
1457	
1458	Supplementary data 4. Scatterplots showing the relationship between observed
1459	standardized water level index (SWLI) against model predicted SWLI and residuals
1460	versus observed SWLI using PLS (component 2) transfer function in the modern data
1461	set derived from the southern California salt marshes.
1462	



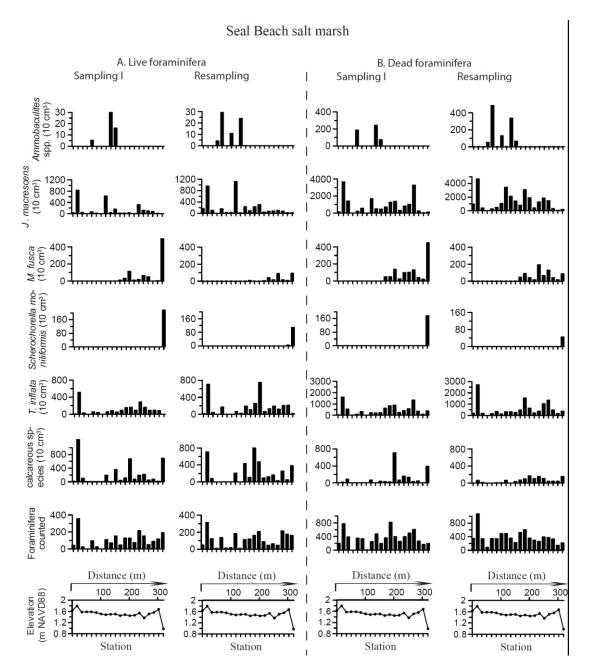
1464 Figure 1

Environmental properties

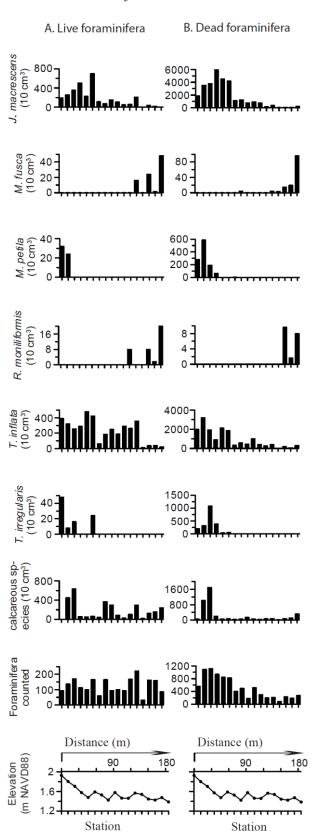


1466

1467 Figure 2

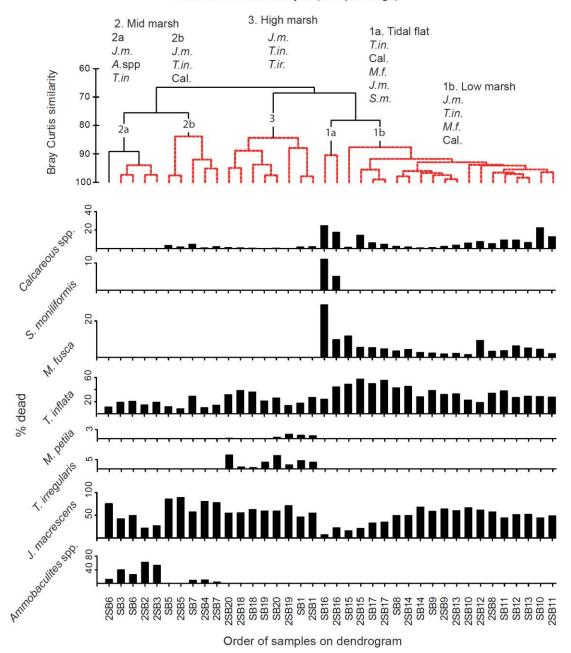


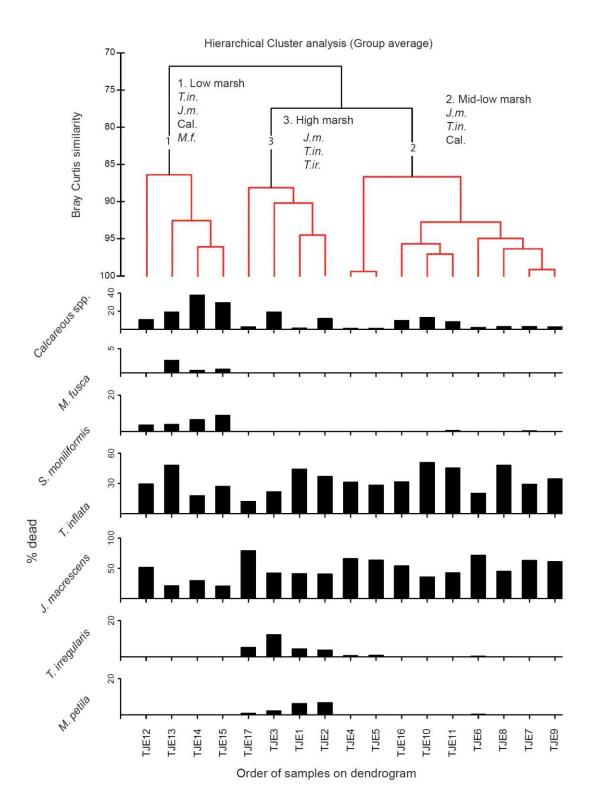
Tijuana salt marsh



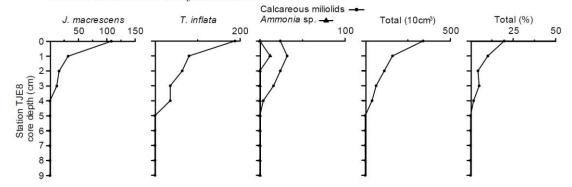
1470

1471 Figure 4

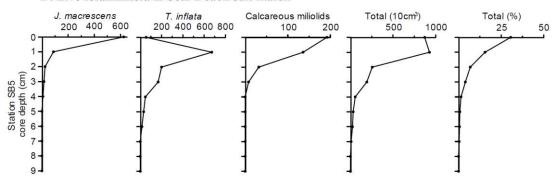




A. Live foraminifera in Tijuana salt marsh

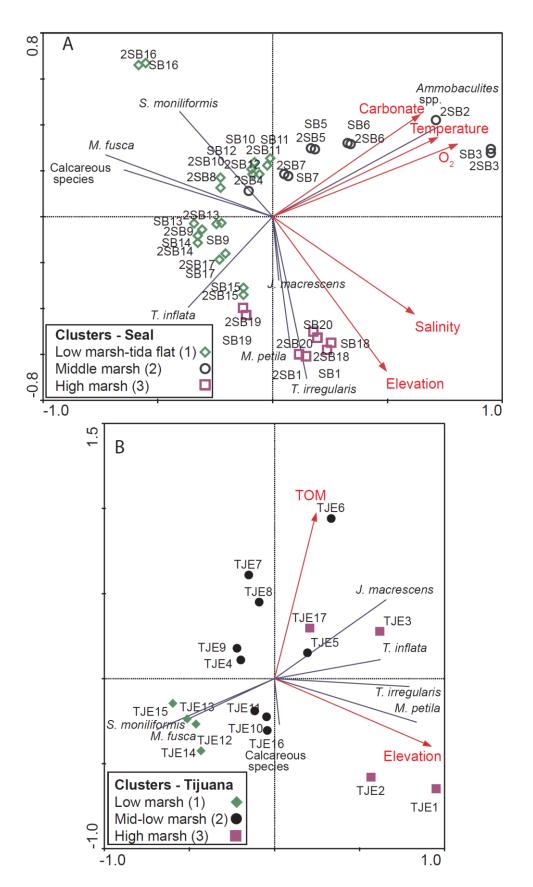


B. Live foraminifera in Seal Beach salt marsh

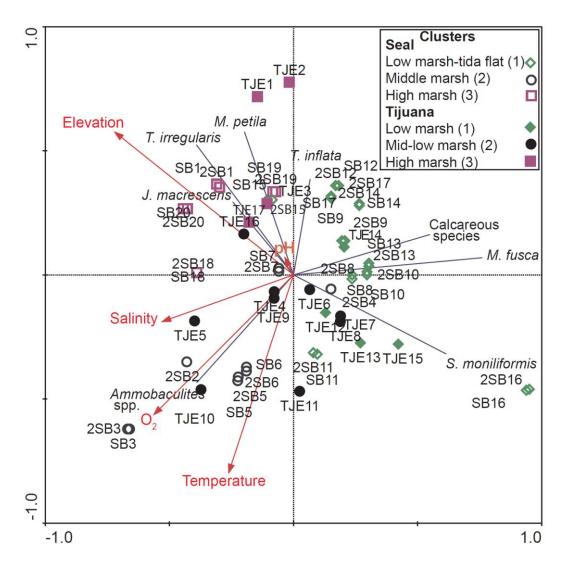


1476

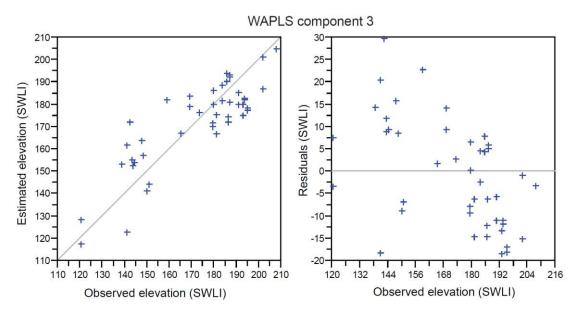
1477 Figure 7



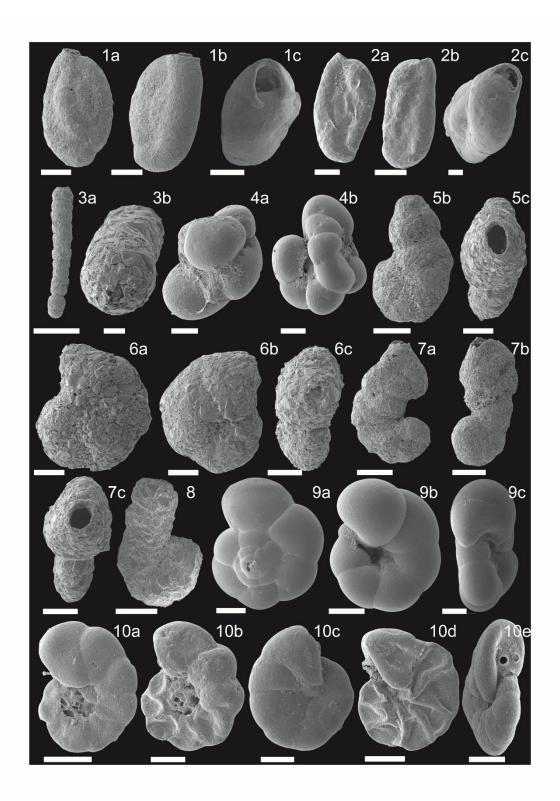
1479 Figure 8

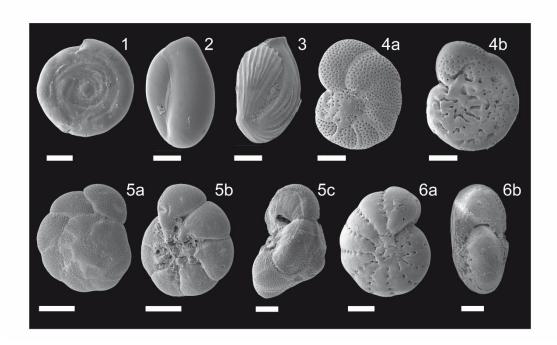


1480



1483 Figure 10





1487 Plate 2

Site	Nearest tidal	MHHW	MTL	MSL	MLLW	Mean
	station (ID)					diurnal
						range
Seal Beach	Los Angeles (9410660)	1.61	0.81	0.8	-0.06	1.67
Tijuana River Estuary	Imperial Beach (9410120)	1.56	0.77	0.77	-0.07	1.64

1490 Table 1

Sample name	Coordinates		Elevation (m NAVD88)	Por	e-water	prope	erties	Sedin	nent proper	rties (%)	
	Latitude (N)	Longitude (E)		p H	Temp.		Salinit y (‰)	ТОМ	Carbonate	Sand	Silt	Clay
SB19	33 44 25.3116	-118 5 17.7163	1.63±0.03	6.5	25.4	1.7	nd	3.4	0.7	65.4	32.1	2.5
SB20	33 44 25.3334	-118 5 17.8292	1.79±0.04	nd	nd	nd	nd	8.1	1.5	29.1	62.0	8.8
SB1	33 44 26.4765	-118 5 16.7351	1.57±0.04	6.7	25.6	1.2	40.8	40.2	6.9	20.7	70.7	8.5
SB2	33 44 26.3093	-118 5 15.6391	1.58±0.04	7.0	30.0	5.8	36.7	55.1	21.6	41.7	50.1	8.1
SB3	33 44 26.0878	-118 5 14.5482	1.58±0.04	7.8	32.5	8.0	38.4	58.1	16.7	48.3	44.1	7.5
SB18	33 44 26.0348	-118 5 13.3831	1.56±0.03	_	27.9	4.6	39.5	1.2	0.6	88.0	12.0	0.0
SB4		-118 5 12.6228		6.6	28.5	1.2	34.2	2.6	0.9	77.6	21.7	0.6
SB5*	33 44 25.5088	-118 5 11.7917	1.48±0.03	4.3	26.9	4.5	34.4	63.5	13.7	10.9	76.9	12.1
SB6	33 44 25.1860	-118 5 11.0653	1.50±0.03	6.5	29.7	4.2	36.0	53.4	11.0	7.5	78.6	13.8
SB7		-118 5 10.2254		5.7	24.3	4.6	34.1	66.6	17.0	8.3	77.5	14.1
SB8	33 44 25.2602	-118 5 9.64825	1.45±0.03	6.1	26.2	2.0	33.5	17.7	3.7	25.2	64.3	10.4
SB9		-118 5 8.60514		_	25.6	2.2	34.0	14.4	3.3	19.4	71.6	8.9
SB10		-118 5 8.57033		_	27.6	2.9	33.7	32.6	3.8	26.4	65.0	8.5
SB11		-118 5 7.2019			29.2	2.8	34.0	9.5	1.4	24.0	71.8	4.1
SB12		-118 5 6.5347		_	24.3	0.8	34.3	52.7	16.2	11.4	79.7	8.8
SB13		-118 5 6.16139		_	24.7	3.0	34.2	9.6	1.8	42.1	52.0	5.7
SB14		-118 5 5.76425		_	25.4	0.3	34.0	11.3	1.6	36.4	58.0	5.5
SB17		-118 5 5.61920			25.4	1.4	34.0	4.4	1.2	45.2	50.2	4.5
SB15		-118 5 5.44184		_	25.4	1.9	33.8	4.4	0.8	54.0	42.2	3.8
SB16		-118 5 5.23403		_	27.0	2.4	34.1	2.7	0.8	67.2	31.1	1.6
TJE1		-117 7 39.37079		_	15.6	1.6	21.4	17.0	1.5	63.2	30.9	5.9
TJE2		-117 7 39.59188		_	15.3	0.2	27.3	15.8	1.7	55.3	35.7	9.0
TJE3		-117 7 39.8134		_	17.5	0.2	31.3	50.8	7.2	15.0	61.6	23.2
TJE17		-117 7 40.4346		_	16.4	0.4	37.2	47.4	5.0	11.3	65.6	22.9
TJE4		-117 7 40.10398		_	16.3	1.2	34.7	35.8	5.1	6.0	67.9	25.9
TJE5		-117 7 40.4217		_	17.1	6.6	32.2	41.4	7.1	8.0	69.3	22.6
TJE6		-117 7 40.76285		_	18.2	0.2	32.8	74.5	11.1	1.9	73.3	24.6
TJE7		-117 7 41.07820		_	17.8	0.2	33.2	56.3	9.9	nd	nd	nd
TJE16		-117 7 41.78665		6.7	16.5	0.5	36.8	20.7	3.1	5.7	75.6	18.6
TJE8*		-117 7 42.14508		_	17.5	0.5	29.4	50.5	7.9	4.9	69.8	25.1
TJE9		-117 7 42.96428		_	17.1	0.4	37.5	38.3	5.2	2.0	72.4	25.4
TJE10		-117 7 43.82999		_	19.3	3.3	34.6	23.9	4.6	5.9	70.8	23.1
TJE11		-117 7 44.63298		_	20.0	1.1	28.7	24.6	4.7	5.0	74.3	20.6
TJE12		-117 7 44.79428			18.6	0.5	35.4	18.0	3.6	6.1	74.8	19.0
TJE13		-117 7 45.10929		_	18.7	0.4	30.9	18.7	4.6	5.6	75.9	18.3
TJE14		-117 7 44.9764:		_	16.8	1.2	32.2	12.0	3.0	17.7	74.1	8.1
TJE15		-117 7 44.8901			19.0	0.4	30.0	21.5	3.5	6.7	76.6	16.6

a. Seal Beach	Axis 1	Axis 2	Axis 3	Axis 4
Eigenvalues	0.318	0.165	0.024	0.012
Lengths of gradient	2.017	1.519	1.19	1.214
Cumulative				
percentage variance of				
species data	40.6	61.7	64.7	66.3
b. Tijuana				
Eigenvalues	0.252	0.012	0.008	0.002
Lengths of gradient	1.468	0.447	0.407	0.438
Cumulative				
percentage variance of				
species data	62.1	65.1	67.2	67.7
c. Combined sites				
Eigenvalues	0.318	0.214	0.022	0.009
Lengths of gradient	1.981	1.847	0.907	1.004
Cumulative				
percentage variance of				
species data	40.9	68.3	71.1	72.3

1494 Table 3

a. Seal Beach	Axis 1	Axis 2	Axis 3	Axis 4	F-value	P value	Captured
							variance
Eigenvalues	0.344	0.164	0.019	0.008			
Species-environment correlation	0.826	0.813	0.517	0.427			
Cumulative percentage variance							
Of species data	34.4	50.8	52.7	53.5			
Of species-environment relatio	63.8	94.2	97.7	99.1			
Correlations							
O ₂ (mg/L)					11.41	0.0005	24.1
Elevation					8.92	0.0005	16.3
Salinity (‰)					3.66	0.0105	16.7
Carbonate (wt.%)					2.79	0.0255	17.7
Temperature (°C)					3.02	0.0215	19.9
b. Tijuana							
Eigenvalues	0.501	0.06	0.235	0.102			
Species-environment correlation			0	0			
Cumulative percentage variance							
Of species data	50.1	56.1	79.6	89.8			
Of species-environment relation	89.4	100	0	0			
Correlations							
Elevation					11.39	0.0005	43.2
TOM (wt.%)					4.11	0.007	8.6
c. Combined sites							
Eigenvalues	0.238	0.141	0.023	0.006			
Species-environment correlation	0.805	0.66	0.535	0.319			
Cumulative percentage variance							
Of species data	23.8	37.9	40.2	40.7			
Of species-environment relation	58.2	92.7	98.3	99.7			
Correlations							
Elevation					10.95	0.0005	17.1
O ₂ (mg/L)					8.41	0.0005	12.2
Salinity (‰)					3.66	0.0075	7.6
Temperature (°C)					2.88	0.027	10.9
pH					3.15	0.0175	2.1

PLS	R ² _{Jack}	Ave.Bias _{Jack}	Max.Bias _{Jack}	RMSEP _{Jack}	%Change
Component 1	0.242	-0.030	32.224	18.033	
Component 2	0.492	-0.696	21.553	14.737	18.281
Component 3	0.507	-0.628	20.451	14.527	1.420
Component 4	0.543	-0.481	26.127	14.218	2.132
Component 5	0.578	-0.253	28.639	13.557	4.643
WAPLS					
Component 1	0.338	0.148	27.778	18.576	
Component 2	0.511	-0.180	20.976	16.002	13.860
Component 3	0.724	0.621	22.757	11.964	25.232
Component 4	0.680	-0.016	26.186	12.995	-8.620
Component 5	0.683	-0.289	27.531	12.964	0.240

1498 Table 5

(C)	S	(A)	S	S	S	S	S	S	S	Γ .
B59	3B58	B5 7-8	B5 6	B5 5-6	B5 4-5	B5 3-4	SB5 2-3	3B5 1-2	SB5 0-1	Sample Name
9-10	8-9		6-7							
9 1	8 1	7 1	6 1	5 1	4 1	3 1	2 1	1 1	0 1	/-
0	ļ -	0		0.		0.		0		Aliquot/8
.125	0.125	.125	0.125	.125	0.125	.125	0.125	.125	0.125	Aliquot sample size
										Census live
					1	2	3	11	78	Jadammina macrescens
						1	4	17	24	Calcareous miliolids
			2	4	6	21	25	84	7	Trochammina inflata
		1	1					4		juvenile Trochamminids
0	0	1	3	4	7	24	32	116	109	Total live
										Census dead
1										Ammobaculites dilatatus
9	4									Ammobaculites sp.
184	232	168	346	373	254	320	207	386	211	Jadammina macrescens
339	178	39	4	3	11	25	7	1		Miliammina fusca
	3	1	2	3	2	1	4	9	∞	Calcareous miliolids
1	1									Scherochorella moniliformis (forma R
				1	1	1				Textulariid sp.
32	26	49	148	118	151	192	184	204	28	Trochammina inflata
						2				Trochammina sp.
4	9	15	41	27	57	38	31	31		juvenile Trochamminids
570	453	272	541	525	476	579	433	631	247	Total dead
570	453	273	544	529	483	603	465	747	356	Total live+dead
										Proccessed data
										Live BF/10 cm ³
0	0	0	0	0	∞	16	24	88	624	Jadammina macrescens
0	0	0	0	0	0	8	32	1:	1:	Calaanaana mailialida
								136	192	Calcareous miliolids
0	0	0	16	32	48	168	200	672	56	Trochammina inflata
0	0	∞	24	32	56	192	256	928	872	total live populations 10cm 3
0	0	0	1	1	1	4	7	16	31	%live relative to the total populations
										Dead BF/10 cm ³
1472	1856	1344	2768	2984	2032	2560	1656	3088	1688	Jadammina macrescens
2712	1424	312	32	24	88	200	56	8	0	Miliammina fusca
0	24	8	16	24	16	8	32	72	64	Calcareous miliolids
256	208	392	1184	944	1208	1536	1472	1632	224	Trochammina inflata
			4		Φ	6	2	2		
	_	!	_	I	I	I	_	I		l

Sample Name		Hour	HGPS Latitude	HGPS Longitude	Name	DGPS Latitude	DGPS Longitude	Elevation	Error	Distrance between srations 01 as ref.	рН Temp. C	_	Salinity ppt (560) % Organic Content	% Carbonate Content	% sand % silt	% clay	Top cm Sediment Bulk		Remarks during	Remarks during analysis	`	Ammobaculites spp. Ammonia sp.		Cornuspira involvens	Elphidium spp. Jadammina macrescens	Miliammina fusca	Miliammina petila Calcareous miliolids	Scherochorella moniliformis (forma R	Reophax spp. Textulariid sp.	Trichohyalus aguayoi Trochammina inflata	Trochamminita ir.	Unidentified very coarse agglutinated
TJE	_	_		117° 07.656			-117 7 39.37079		0.026		7.2 16		_	_	63 31	6	х х	Frankenia salina ; Distichli			0.125				24		4			49		
TJE	_	_				32 34 27.01151	-117 7 39.59188	1.803		5.09902			_	_	55 36		х х	Frankenia salina ; Distichli	ŭ		0.125			\sqcup	32		3 56	\perp		40		
_	3 3			117° 07.665	-		-117 7 39.81347	1.705		11.7047			_		15 62	_	х х	Frankenia salina ; Sarcoco			0.125				44		72	\perp		7 32		
TJE	17 4	16:30	32° 34.451	117° 07.677	TJE17	32 34 27.11920	-117 7 40.43469	1.576			6.8 16		_	_	11 66	23	х х	Frankenia salina ; Sarcoco	High Man	sh 1	0.125				63		7			36		
	4 5			117° 07.670			-117 7 40.10398	1.476		20.24846					6 68	26	х х	Sarcocornia californica ; so			0.125				28		6			60		
_	5 6	_		117° 07.675	_		-117 7 40.42177	1.591		28.4605			_	7.05	8 69	23	х х		Mid mars		0.125	4			87		4	\perp		53		_
	6 7			117° 07.680			-117 7 40.76285			37.94733		0.2 3			2 73	25	х х	Transition from Sarcocorn	·		0.25	5		1	28		5	\perp		15		_
_	7 8	_	32° 34.442	117° 07.684	TJE7		-117 7 41.07826	1.424	0.029	46.8188			_	_			х х	Frankenia salina ; Spartina	Midmars	h 2	0.25	1		1	18		89			46		_
_	16 9	15:40					-117 7 41.78665	1.579		66.7308						19		Jaumea carnosa; Sarcocor			0.125			1	19		36	\perp		31		4
TJE		_		117° 07.704	_	32 34 26.13974	-117 7 42.14508	1.464		77.25283			_	_		25		Jaumea carnosa ; Sarcoco			0.25				27		21			47		
_	9 11			117° 07.716		32 34 25.63580	-117 7 42.96428			102.8834			_	_			х х	Jaumea carnosa ; Batis ma	Mid-to-lo	ow (2	0.25				12					6 72		
_	10 12		-	117° 07.732		32 34 25.10969	-117 7 43.82999	1.567		130.5986			_			23	х х	Sarcocornia californica; S	Low mars	sh 3	0.375	32			23		6	3		98		
TJE	11 13	_		117° 07.743		32 34 24.81891	-117 7 44.63298	1.541		153.3949			_	4.66		21	х х	Pure Spartina foliosa	Low mars	sh 2	0.25	41		2	52	4	28		1	2 89		
_	12 14	_		117° 07.747		32 34 24.34597	-117 7 44.79428	_		163.6154					6 75		х х	Spartina foliosa on mudfla	Low mars	sh 8	1	8		3	3 4		5			12		
TJE	13 15	14:38	32° 34.400	117° 07.753	TJE13	32 34 23.94699	-117 7 45.10929	1.425	0.025	177.2484			_	_	6 76	18	х х	Pure Spartina foliosa	Low mars	sh 5	0.625	1 52		2	25	15	21	5	2	2 23		1
TJE	14 16	14:54	32° 34.392	117° 07.750	TJE14	32 34 23.53768	-117 7 44.97645		0.025	181.604				-	18 74	8	х х	Pure Spartina foliosa	Low mars	sh 5	0.625	59	1	7	7 12	1	30	1 3	2	1 25	17	
TJE	15 17	15:07	32° 34.388	117° 07.750	TJE15	32 34 23.27933	-117 7 44.89011	1.387	0.025	184.9108	7.2 19	0.4 3	22	3.45	7 77	17	х х	Pure Spartina foliosa	Low mars	sh 2	0.25	47		1	l	12	12	5		5	3	

85	159	161	32	221	167	92	100	93	164	59	165	98	112	169	136	92	Total live
					Ì										0.		Census dead
																	Ammobaculites spp.
61	38	23	2	7	18	1	1			2	3	1					Ammonia sp.
	1				1												Bolivina sp.
1																	Cornuspira involvens
4	4	6	3	2	5	5			1	2	2						Elphidium spp.
56	55	49	44	92	72	184	232	96	312	287	522	560	743	470	440	232	Jadammina macrescens
24	12	9	3	1		4	2		2 1	7	2	Ö	ώ	Ŏ	Ö	2	Miliammina fusca
_	-									1			8	24	73	35	Miliammina petila
10	24	13	4	8			14	17	14	4	3	6	24	208	123	7	Calcareous miliolids
) 2	1	6					_							8	ü		Scherochorella moniliformis (forma R
7										1							Reophax spp.
7	3	ω															Textulariid sp.
4	4	2		1	3	2								4	5		Trichohyalus aguayoi
74	33	112	25	98	103	104	248	56	144	81	232	264	11	240	400	248	Trochammina inflata
_	3.	2	<u> </u>	3	ß	4	8	0,	4	1	32 7	54 5	112 48	Ю 135	0 40	18 24	
23	11	11	4	8	1	6	21	9	24	22	54	14	8 1	35 32	0 6	4 16	Trochamminita irregularis
3	1	1					1		4	2	4	4		2		6 1	juvenile Trochamminids
273	18	23	85	2:	2(3	5.	1	49	4(∞,	82	93	1:	1(563	Unidentified agglutinated miliolid Total dead
73	186	234	Oi	217	203	302	516	78	496	401	823	850	936	1113	1087	53	Total dead
																	Proccessed data
0	0	2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	Live BF/10 cm ³
			8			0) 0	0) 4) 0	0	0	0	0	Ammobaculites spp.
188	94	83		164	85					20	32						Ammonia sp.
0 0	2 0	0 0	0 0	0 0	0 0	0	0 0	0 8	0 0	0 0	0 0	0 0	0 0	0 0	0 0	0 0	Bolivina sp.
) 3		0												Cornuspira involvens
4 0	11 1	3 2	3 4	8		0 2	0 1	0 1	4 7	4 1	0 6	0 2	0 5	0 3	0 2	0 1	Elphidium spp.
	19	40		208	61	48	108	152	72	112	696	224	504	352	256	192	Jadammina macrescens
48	2	24	0	16	0	0	0	0	0	0	0	0	0	0	0	0	Miliammina fusca
0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	24	32	Miliammina petila
48	48	34	5	112	16	0	8	288	356	20	32	48	56	576	448	0	Calcareous miliolids
20	2	8	0	0	8	0	0	0	0	0	0	0	0	0	0	0	Scherochorella moniliformis (forma R
0	5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	Reophax spp.
0	3	3	0	4	0	0	0	0	0	0	0	0	0	0	0	0	Textulariid sp.
0	2	3	0	8	0	24	0	0	0	0	0	0	0	56	0	0	Trichohyalus aguayoi
20	40	37	12	356	261	288	188	248	184	60	424	480	288	256	320	392	Trochammina inflata
0	0	0	0	0	0	0	0	0	0	0	24	0	0	16	∞	48	Trochamminita irregularis
12	27	19	0	8	13	∞	20	48	36	20	112	32	48	96	32	72	juvenile Trochamminids
0	0	2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	Unidentified very coarse agglutinated
240	155	123	16	292	101	24	84	296	364	44	2	48	56	632	448	0	Live general Calcareous /10 cm3
340	5 254.4	3 257.6	32	2 884	1 445	368	400	5 744	4 656	236	1320	784	896	2 1352	3 1088	736	Live absolute abundances /10 cm3
_	1.4	7.6		_	J1	~	_	_	31	31	ŏ	1	J1	2	88	J,	
																	Dead BF/10 cm ³

	0 0 0 0 8 24 8 0 0 4	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0	. dds ani dd ani	1856 3520 3760 5944 4480 4176 1148 1248 768 928 736	0 58 0 19 0 64 0 0	56 1 984	0 4 0 64 0 9 0 0 0 0 0	Simple familia	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 40 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	210 1984 3200 1920 896 2112 1856 324 576 448 992 416	320 1080 384 40	Spirimendochaminide 128 48 256 8 112 432 88 96 72 84 24	0 0 0 0 0 0 0 0	56 1024 1696 192 56 64 32 60 32	Em 201/ save pend appointed appointe		Ammopac	0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 2 1 0 0	State	4.3.2.2.2.0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	O Calcareous n	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 Textulariid sp.		53 6 29 0 19 2 25 0 28 0 47 0 78 0		o o o o o o o O Unidentified very coarse agglutinated	0 41 47 6 6 5 19 55 40 21 7		Ammondanies spp.		O O O O O O O O O O O O O O O O O O O	0 4 0 3 0 0 0 0 0 0 0 0 0 0 0 0 2 0	41 C 40 C 42 C 79 C 66 C 63 C 72 C 63 C 64 C 65 C	200	5 1 7 1:	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 4 0 3 0 2 0 1 0 3 0 2 0 2 0 2 0 3 0 4 1 3	100	1 12 3 3 5 0 0 5 5 0 0 5 0 0 0 0 0 0 0 0 0 0	7-1-11-11-11-11-11-11-11-11-11-11-11-11-	Dildentified aggintinated miliolid
0	0	0	0	0	768	0 0	136	5 0	0	0	0	448	0	72	0	136	1424	(0 0	0	1.1	0	20 0	0		0 0	0	0	33 (0 6	0	40	0	0	0	0	_) 0) 1	0 0	0	0 (0 3	31 0) 5	5 0	1
0	4	0	0	0		0 0	56	0	0	0	0	_	0	-				(0 0	0	0	_		0	21	0 0	0	0	_	5	0	21	0	0	0	0) 0	3	0	0	0 (-	_) 4	0	3
0	48	3	0	13	192	0	0	0	0	0	8	275	0	3		69	541	() 1	9 0	0	_	14 0	0	4	180	0	0.5	59 () 2	0	23	0	9	0	0		35 () () ()	0	0	0 .		54 U	0 0	. O	1
0	28	0	0	8	368	4 0	32	0	0	0	4	392	0	32		72	868	(_	9 0	0	1	24 1.8	3 0	13	0 0	0	0.9	40 () 1	0	33	0	3	0	0	1 4	42 () (4	0	0	0		45 0) 4	4 0	8
0	2	0	0	3	44	3 0	4	0	0	0	0	25	0	4	0	9	85	() 2	5 0	0	9	13 0	0	16	0 0	0	0	38 (0 0	0	50	0	2	0	0	4	52 4	1 0) 5	0	0	0	0 2	29 0	5 5	5 0	1
0	37	0	0	10	78	14 0	21	10	0	5	3	179	0	18	0	70	374	1	1 3	2 0	0	1	16 9.3	3 0	13	3.1 0	1	1.2	14 (7	1	48	0	10	0	0	3	21 4	4 0) 6	3	0	1	1 4	48 0) 5	5 0	1
0	61	2	0	6	88	19 0	38	2	0	5	6	53	0	18	0	112	298	() 3	7 1	0	4	7.5 0.6	6 0	19	0.6 2	1	0.6	16 () 1	1 0	61	0	20	1	0	2	30 (6	5 0) 13	3 1	0	2	2 1	18 0	ა 6	5 0	3
0	244	0	4	16	224	96 0	40	8	28	28	16	296	0	92	0	320	1092	() 5	5 0	0	1	0 14	. 0	14	5.9 0	0	0	5.9) 4	0	71	0	22	0	0	1	21 9) () 4	1	3	3	1 7	27 0) 8	3 0	2

Taxonomic reference list of species presented in text and in Supplementary data 1 and
 2
 1511

1512

1513 Ammobaculites spp. Wright et al., 2011, p. 59, Fig. A2/4

1514

1515 *Ammonia* spp. Murray, 1979, p. 57, Figs. 18a–l; Horton and Edwards, 2006; p. 73, P3,

1516 Figs. 10a–c, Figs. 11a–c, Fig. 12a–c.

1517

1518 *Cornuspira* spp. Loeblich & Tappan, 1988, p. 322, Pl. 51, Figs. 7–8

1519

1520 *Trichohyalus aguayoi* (Bermúdez, 1935). Bermudez, 1935, p. 204, pl. 15, Figs. 10–

1521 14

1522

- 1523 Elphidium spp. Murray, 1979, p. 53, Figs. 16a–d; Hayward et al., 1999, p. 219, P17,
- 1524 Figs. 1–28; Murray, 2006, p. 65, Fig. 4.2, 11; Horton and Edwards, 2006; p. 75, P4,
- 1525 figs. 15–20.

1526

- 1527 Jadammina macrescens (Brady, 1870). Murray, 1979, p. 27, Figs. 6k-m; Gehrels and
- van de Plassche, 1999, p.98, P1, Figs. 1–5; Hayward et al., 1999, p. 217, P1, Figs. 27–
- 29; Horton and Edwards, 2006; p. 67, P1, Figs. 4a–d; Hawkes et al., 2010, p. 133, P1,
- 1530 Figs. 7a–d; Wright et al., 2011, p. 58, Fig. A1/5.

1531

- 1532 Miliammina fusca (Brady, 1870). Murray, 1979, p. 24, Figs. 5d–f; Hayward et al.,
- 1533 1999, p. 217, P1, Figs. 25, 6; Edwards et al., 2004; p. 16, P1, Fig. 7; Horton and
- Edwards, 2006; p. 67, P1, Figs. 5a, b. Wright et al., 2011, p. 59, Fig. A2/2.

1535

- 1536 Miliammina petila Saunders, 1958. Saunders, 1958, p. 87, pl. 1, figs. 10, 11; Milker et
- al., 2015, p. 5, Pl. 1, Figs. 3–4.

1538

- 1539 *Quinqueloculina* spp. Murray, 1979, p. 35, Figs. 9a–i; Hayward et al., 1999, p. 223;
- 1540 P4, Figs. 26–28; p. 225, P5, Figs. 9, 10; Horton and Edwards, 2006; p. 71, P2, Figs.
- 1541 9a, b.

1542

- 1543 Scherochorella moniliformis Siddall, 1886. Murray, 1979, p. 24, Fig. 5b; Horton and
- Edwards, 2006; p. 67, P1, Figs. 6a–c; Wright et al., 2011, p. 58, Fig. A1/6.

1545 1546

- 1547 *Trochammina inflata* (Montagu, 1808). Hayward et al., 1999, p. 219, P2, Figs. 6–8;
- 1548 Edwards et al., 2004; p. 16, P1, Figs. 14, 15; Horton and Edwards, 2006; p. 71, P2,
- 1549 Figs. 8a-d; Wright et al., 2011, p. 58, Fig. A1/9; Kemp et al., 2012; p. 29, P1, Figs. 7-
- 1550 8. Appendix A, Appendix B

1551

- 1552 Trochamminita irregularis Cushman & Bro"nnimann, 1948. Loeblich & Tappan,
- 1553 1988, p. 67, Pl. 51, Figs. 1–5; Hawkes et al. 2010, p. 18, Pl. 1, Figs. 3a–b; Milker et
- 1554 al., 2015, p. 5, Pl. 1, Fig. 11.

15551556

1557

1558 Reference

1560

- 1561 Bermudez, P.J. (1935). Foraminiferos de la costa norte de Cuba. Memorias de la
- 1562 Sociedad
- 1563 Cubana de Historia Natural, 9, 129–224.

1564

- Edwards, R. J., Wright, A. J., & van de Plassche, O. (2004). Surface distribution of
- salt-marsh foraminifera from Connecticut, USA: Modern analogues for high-
- resolution sea level studies. *Marine Micropaleontology*, *51*, 1–21.

1568

- Gehrels, W. R., & van de Plassche, O. (1999). The use of Jadammina macrescens
- 1570 (Brady) and Balticammina pseudomacrescens Brönnimann, Lutze and Whittaker
- 1571 (Protozoa: Foraminiferida) as sea-level indicators. *Palaeogeography*,
- 1572 Palaeoclimatology, Palaeoecology, 149(1), 89–101.

1573

- Hawkes, A. D., Horton, B. P., Nelson, A. R., & Hill, D. F. (2010). The application of
- intertidal foraminifera to reconstruct coastal subsidence during the giant Cascadia
- earthquake of AD 1700 in Oregon, USA. *Quaternary International*, 221(1), 116–140.

1577

- Hayward, B. W., Grenfell, H. R., Reid, C. M., & Hayward, K. A. (1999). Recent New
- 1579 Zealand Shallow-Water Benthic Foraminifera: Taxonomy, Ecologic Distribution,
- 1580 Biogeography, and use in
- 1581 Paleoenvironmental Assessments: Institute of Geological and Nuclear Sciences
- 1582 Monograph 21, Lower Hutt, New Zealand, 264 p.

1583 1584

- Horton, B. P., & Edwards, R. J. (2006). Quantifying Holocene sea level change using
- intertidal foraminifera: lessons from the British Isles. Cushman Foundation for
- 1587 Foraminiferal Research Special Publication 40.

1588 1589

- 1590 Kemp, A. C., Horton, B. P., Vann, D. R., Engelhart, S. E., Grand Pre, C. A., Vane, C.
- H., Nikitina, D., & Anisfeld, S. C. (2012). Quantitative vertical zonation of salt-marsh
- 1592 foraminifera
- 1593 for reconstructing former sea level; an example from New Jersey, USA. *Quaternary*
- 1594 *Science Reviews*, *54*, 26–39.

1595

- Loeblich, A. R., & Tappan, H. (1988). Foraminiferal genera and their classification:
- 1597 Van Nostrand Reinhold Company, New York, 1694p.

1598

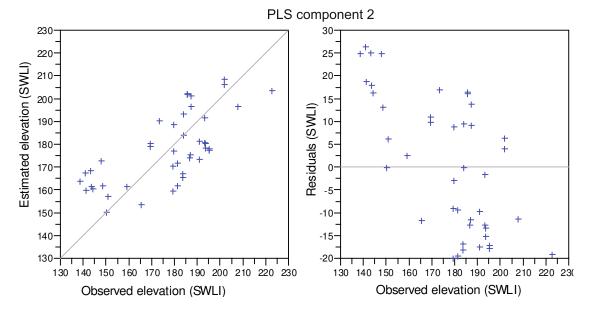
- 1599 Milker, Y., Horton, B. P., Nelson, A. R., Engelhart, S. E., & Witter, R. C. (2015).
- Variability of intertidal foraminiferal assemblages in a salt marsh, Oregon,
- 1601 USA. Marine Micropaleontology, 118, 1–16.

1602

- Murray, J. W. (1979). British nearshore foraminiferids, in Kermack, D. M., and
- Barners, R. S. K., (eds), Synopsis of the British Fauna (New Series) No 16: Academic
- 1605 Press, London, 62 p.

- 1607 Murray, J. W. (2006). Ecology and Applications of Benthic Foraminifera: Cambridge
- 1608 University Press, Cambridge, 426 p.

1609	
1610	Saunders, J.B. (1958). Recent foraminifera of mangrove swamps and river estuaries
1611	and their fossil counterparts in Trinidad. <i>Micropaleontology</i> , 4, 79–92.
1612	
1613	Wright, A. J., Edwards, R. J., & van de Plassche, O. (2011). Reassessing transfer-
1614	function performance in sea-level reconstruction based on benthic salt-marsh
1615	foraminifera from the Atlantic coast of NE North America. Marine
1616	Micropaleontology, 81(1), 43–62.
1617	
1618	



1620 Supplementary data 4